



Element-based Galerkin Methods in Geophysical Fluid Dynamics Modeling

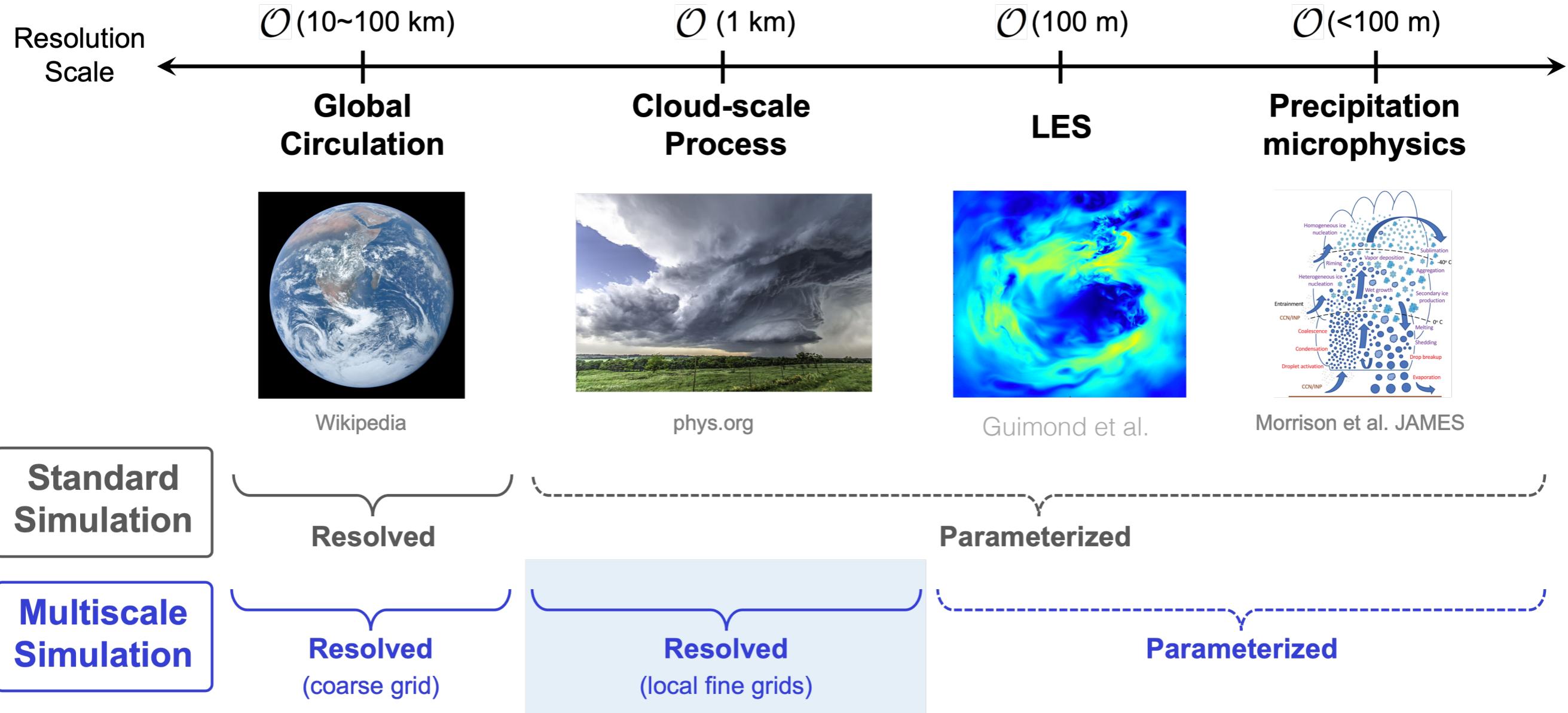
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Dartmouth College

Frank Giraldo, Distinguished Professor and Chair
Department of Applied Mathematics
Naval Postgraduate School
Monterey CA 93943

fxgirald@nps.edu

<http://frankgiraldo.wix.com/mysite>

Motivation



- Global Circulation Model (GCM) resolution approaching scales $\Delta h = \mathcal{O}(1 \text{ km})$ and $\Delta v = \mathcal{O}(100 \text{ m})$
- Vertical direction is special in that most “physics” occurs in this direction (i.e., moist processes such as the formation of clouds and precipitation of rain)
- Will show results for GCM, CRM, and LES

Acknowledgements



- **NUMA Model(s)**
 - Funded by ONR, NSF, DARPA, AFOSR
 - Collaborators: Soonpil Kang, Kiran Jadhav, FXG (NPS); Jim Kelly (NRL-DC); Felipe Alves (Un Reddy (LLNL), Yassine Tissaoui, Simone Marras (NJIT); Steve Guimond (Hampton Univ)
 - U.S. Navy's NEPTUNE dynamics based on NUMA. Both NUMA and NEPTUNE are space-weather capable.
 - NUMA is part of the Nvidia SDK Testing Analysis suite that is run nightly across various hardware and compilers.
 - <https://frankgiraldo.wixsite.com/mysite/numa>
 - <https://frankgiraldo.wixsite.com/mysite/xnuma>
- **NUMO Model(s)**
 - Funded by DOE, ONR
 - Collaborators: Yao Gahounzo, Michal Kopera (Boise State); Wieslaw. Maslowski, FXG (NPS);
 - NUMO is an incompressible NS solver used for studying the effects of ice-ocean interactions at, e.g., fjords.
 - h-NUMO is a multi-layer SWE equation developed by Yao Gahounzo (PhD thesis at Boise State Univ.).
 - <https://frankgiraldo.wixsite.com/mysite/numo>
 - <https://github.com/ygahounzo/h-NUMO>

Governing Equations

Atmospheric Equations:

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0$$

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} + \frac{1}{\rho} \nabla P + \nabla \phi + 2\Omega \times \mathbf{u} = 0$$

$$\frac{\partial E}{\partial t} + \mathbf{u} \cdot \nabla E + \delta(\gamma - 1)E \nabla \cdot \mathbf{u} = 0$$

$$P = (1 - \delta)P_A \left(\frac{\rho R \theta}{P_A} \right)^\gamma + \delta(\gamma - 1)\rho e$$

Ocean Equations:

$$\nabla \cdot \mathbf{u} = 0$$

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} + \frac{1}{\rho_0} \nabla P + \frac{\rho}{\rho_0} g \mathbf{k} + 2\Omega \times \mathbf{u} = 0$$

$$\frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = \nabla \cdot (k_T \nabla T)$$

$$\frac{\partial S}{\partial t} + \mathbf{u} \cdot \nabla S = \nabla \cdot (k_S \nabla S)$$

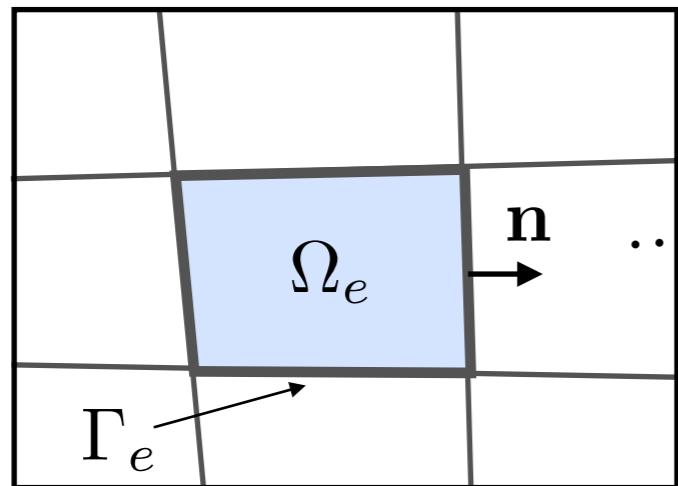
$$\rho = f(T, S)$$

- NUMA carries 4 different forms of compressible Euler/Navier-Stokes (non-hydrostatic) including conservation/balance law forms.
- NUMO carries Boussinesq incompressible Navier-Stokes (non-hydrostatic) and h-NUMO is a hydrostatic version.

Element-based Galerkin Methods [1]

Domain decomposition

$$\Omega = \sum_{e=1}^{N_e} \Omega_e$$

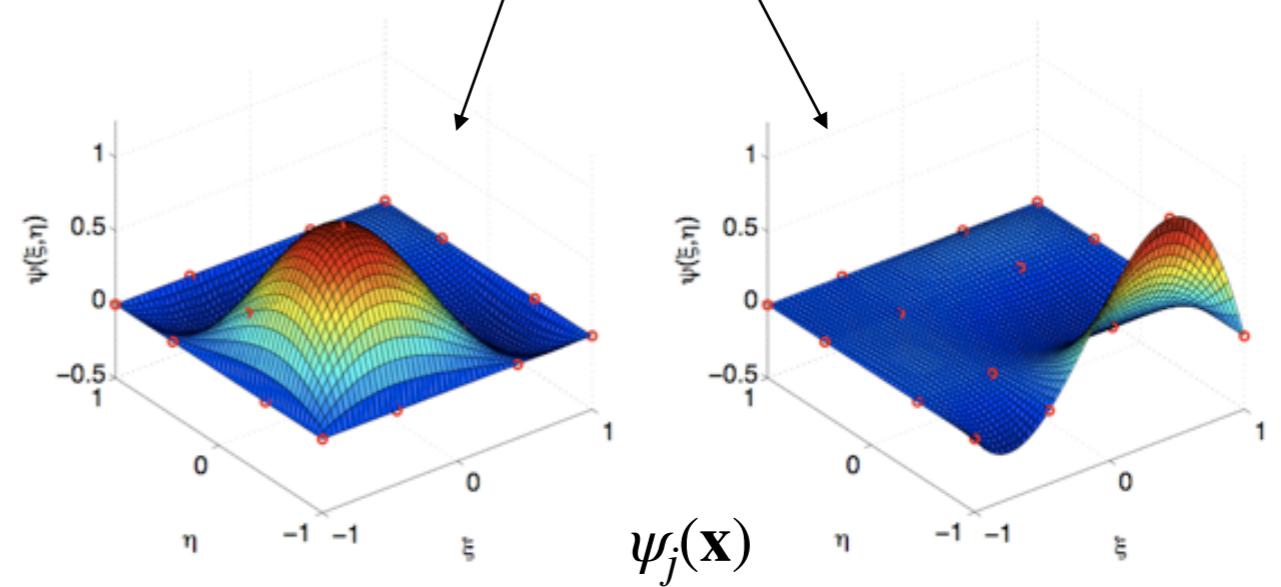
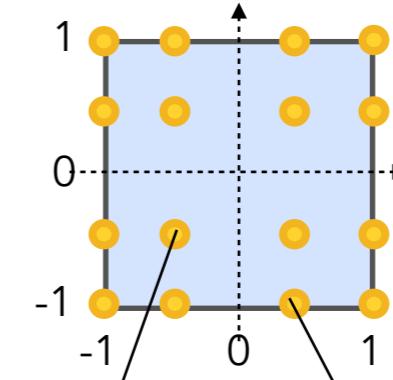


Approximate local solution as:

$$q_N^{(e)}(\mathbf{x}, t) = \sum_{j=1}^{M_N} \psi_j(\mathbf{x}) q_j^{(e)}(t)$$

Reference element

- Legendre-Gauss-Lobatto points



Basis functions - Lagrange polynomials

Element-based Galerkin Methods

Governing equation:

$$\frac{\partial \mathbf{q}}{\partial t} + \nabla \cdot \mathbf{F}(\mathbf{q}) = \mathbf{S}(\mathbf{q})$$

Approximate the **global** solution as:

$$q_N^{(e)}(\mathbf{x}, t) = \sum_{j=1}^{M_N} \psi_j(\mathbf{x}) q_j^{(e)}(t) \quad \begin{aligned} \mathbf{F}_N &= \mathbf{F}(\mathbf{q}_N) \\ \mathbf{S}_N &= \mathbf{S}(\mathbf{q}_N) \end{aligned}$$

Define residual:

$$R\left(q_N^{(e)}\right) \equiv \frac{\partial \mathbf{q}_N^{(e)}}{\partial t} + \nabla \cdot \mathbf{F}_N^{(e)} - \mathbf{S}_N^{(e)} = \epsilon$$

Problem statement:

Find $\mathbf{q}_N \in \mathcal{S}$ $\forall \psi \in \mathcal{S}$ $\left\{ \begin{array}{l} \mathcal{S}_{CG} = \{\psi \in H^1(\Omega) : \psi \in P_N(\Omega_e) \ \forall \Omega_e\} \\ \mathcal{S}_{DG} = \{\psi \in L^2(\Omega) : \psi \in P_N(\Omega_e) \ \forall \Omega_e\} \end{array} \right.$

such that

$$\int_{\Omega_e} \psi_i R(\mathbf{q}_N) d\Omega_e = 0$$

Element-based Galerkin Methods

Integral form:

$$\int_{\Omega_e} \psi_i R(\mathbf{q}_N^{(e)}) d\Omega_e = 0$$

$$\int_{\Omega_e} \psi_i \frac{\partial \mathbf{q}_N^{(e)}}{\partial t} d\Omega_e + \int_{\Omega_e} \psi_i \nabla \cdot \mathbf{F}_N^{(e)} d\Omega_e - \int_{\Omega_e} \psi_i S_N^{(e)} d\Omega_e = 0$$

Integration by parts: $\int_{\Omega_e} \psi_i \nabla \cdot \mathbf{F}_N^{(e)} d\Omega_e = \int_{\Omega_e} \nabla \cdot (\psi_i \mathbf{F}_N^{(e)}) d\Gamma_e - \int_{\Omega_e} \nabla \psi_i \cdot \mathbf{F}_N^{(e)} d\Omega_e = \int_{\Gamma_e} \mathbf{n} \cdot \psi_i \mathbf{F}_N^{(e)} d\Gamma_e - \int_{\Omega_e} \nabla \psi_i \cdot \mathbf{F}_N^{(e)} d\Omega_e$

$$\int_{\Omega_e} \psi_i \frac{\partial \mathbf{q}_N^{(e)}}{\partial t} d\Omega_e + \int_{\Gamma_e} \mathbf{n} \cdot \psi_i \mathbf{F}_N^{(e)} d\Gamma_e - \int_{\Omega_e} \nabla \psi_i \cdot \mathbf{F}_N^{(e)} d\Omega_e - \int_{\Omega_e} \psi_i S_N^{(e)} d\Omega_e = 0$$

face integral volume integrals

Matrix form:

$$M_{ij}^{(e)} \frac{d\mathbf{q}_j^{(e)}}{dt} + \sum_{f=1}^{N_{faces}} \left(\mathbf{M}_{ij}^{(e,f)} \right)^T \mathbf{F}_j^{(e,f,*)} - \left(\tilde{\mathbf{D}}_{ij}^{(e)} \right)^T \mathbf{F}_{ij}^{(e)} - S_i^{(e)} = 0$$

$$M_{ij}^{(e)} = \int_{\Omega_e} \psi_i \psi_j d\Omega_e$$

mass matrix

$$\mathbf{M}_{ij}^{(e,f)} = \int_{\Gamma_e} \psi_i \psi_j \mathbf{n}^{(e,f)} d\Gamma_e$$

face mass matrix

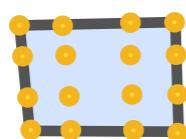
$$\tilde{\mathbf{D}}_{ij}^{(e)} = \int_{\Omega_e} \nabla \psi_i \psi_j d\Omega_e$$

differentiation matrix

Element-based Galerkin Methods

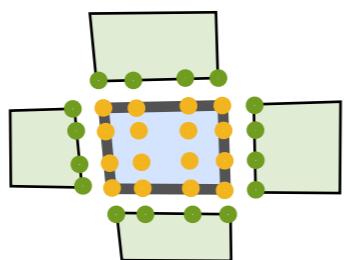
$$\int_{\Omega_e} \psi_i \frac{\partial \mathbf{q}_N^{(e)}}{\partial t} d\Omega_e + \int_{\Gamma_e} \psi_i \mathbf{n} \cdot \mathbf{F}_N^{(*)} d\Gamma_e - \int_{\Omega_e} \nabla \psi_i \cdot \mathbf{F}_N^{(e)} d\Omega_e - \int_{\Omega_e} \psi_i \mathbf{S}_N^{(e)} d\Omega_e = \mathbf{0}$$

- Evaluate “volume” integrals on element interiors



$$R^{(e)} := \int_{\Omega_e} \nabla \psi_i \cdot F_N^{(e)} d\Omega_e + \int_{\Omega_e} \psi_i S_N^{(e)} d\Omega_e$$

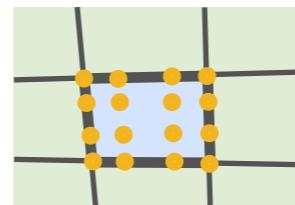
- Evaluate flux integrals



$$R^{(e)} := R^{(e)} - \int_{\Gamma_e} \psi_i n \cdot F_N^{(*)} d\Gamma_e$$

CG: cancels at interior element edges

- Direct Stiffness Summation



$$R = \bigwedge_{e=1}^{N_e} R^{(e)}$$

DG: matrices are block diagonal except for the flux matrix

- Multiply by inverse global mass matrix and evolve time step

$$\frac{dq_i}{dt} = M_i^{-1} R_i$$

We rely on inexact integration so M is diagonal

- NUMA/NUMO carry CG and DG, xNUMA CG, and ATUM/h-NUMO DG

SBP methods

- An interesting area of research in EBG concerns SBP methods [**see MS01, MS03, MS05 for more recent developments in this area**]

- For conservation laws discretized as: $M_i \frac{d\mathbf{q}_i^{(e)}}{dt} + D_{ij} f_j^{(e)} = \mathbf{0}$ we rewrite as

$$M_i \frac{d\mathbf{q}_i^{(e)}}{dt} + \sum_{i=1}^{M_N} D_{ij} \circ \hat{f}_{ij}^{(e)} = \mathbf{0}$$

- with, e.g., the kinetic-energy-preserving flux [2] being $\hat{f}_{ij}^{(e)}(q, \mathbf{u}) = 2\{\{\mathbf{u}_i\}\}\{\{q_j\}\}$ and other options for entropy-stable flux [3]
- Will show the benefits of these methods in the Results section

EBG GFD Models

Atmospheric Models

- NUMA* (NPS, USA)
- NEPTUNE (U.S. Navy's global NWP model)
- ATUM* (NPS, USA)
- KIM (South Korea)
- LFRic* (UK Met Office)
- CAM-SE (in CESM at NCAR, USA)
- EAM (in E3SM at DOE, USA)

Ocean Models

- NUMO* (BSU/NPS, USA)
- h-NUMO* (BSU, USA)
- SLIM3D* (UCL, Belgium)
- Fluidity (Imperial College, UK)
- RiCOM (NIWAR, NZ)
- SCHISM (VIMS, USA)

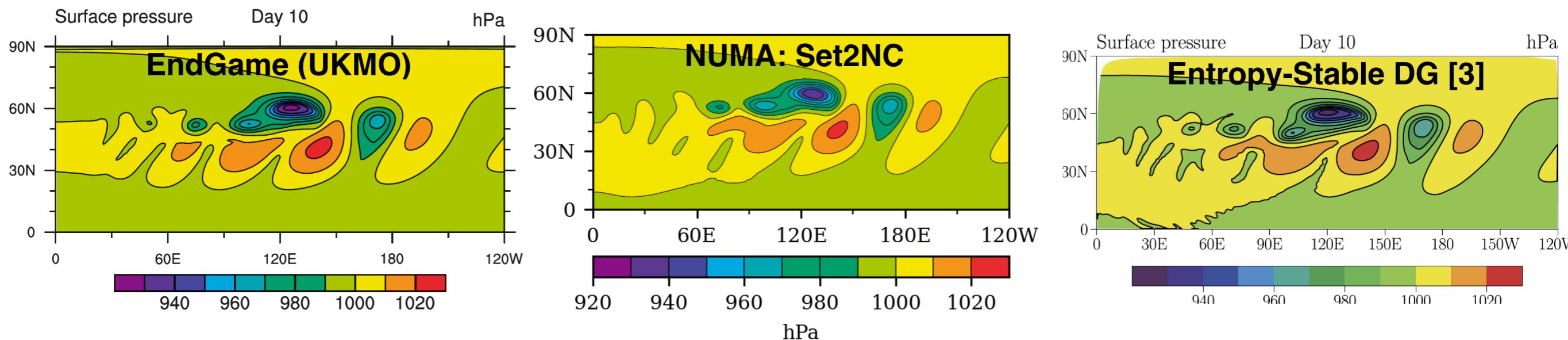
*use EBG in all 3 directions. *use high-order EBG ($P>1$)

Results

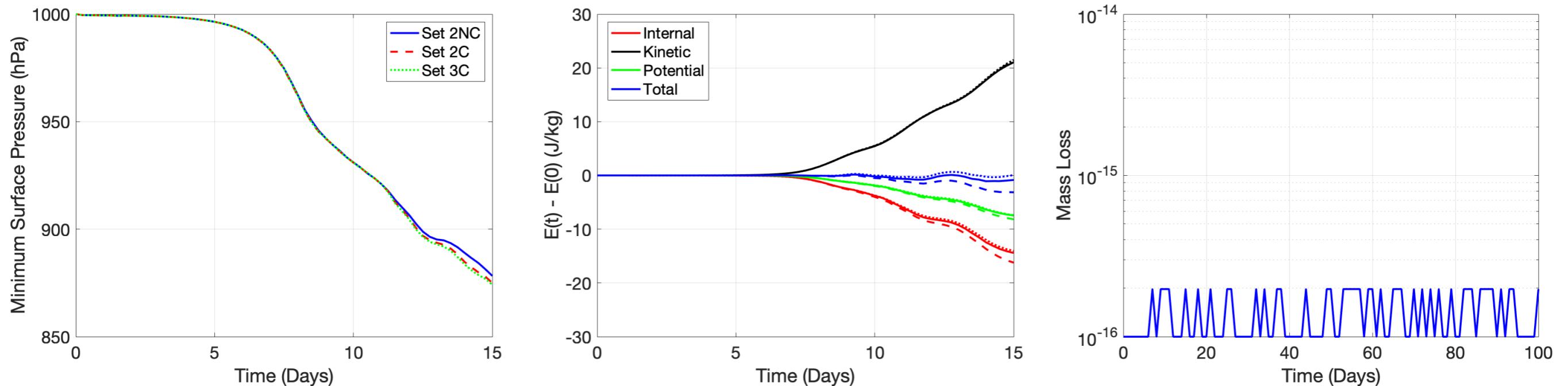
1. Baroclinic Instability on the Sphere (NUMA, GCM scale)
2. Tropical Cyclone Simulations to better understand rapid intensification (NUMA, CRM/LES scales)
3. Density Current with Salinity (NUMO, LES scale)
4. Double Gyre (h-NUMO, GCM scale)

Result 1: Baroclinic Instability (GCM scale) [4]

Baroclinic Instability on the sphere for 10-day forecast simulation with Grid Resolution: $\Delta h = 100 \text{ km}$ and $\Delta v = 1 \text{ km}$. ES-DG uses no viscosity (conserves total energy exactly) and remains stable for all time

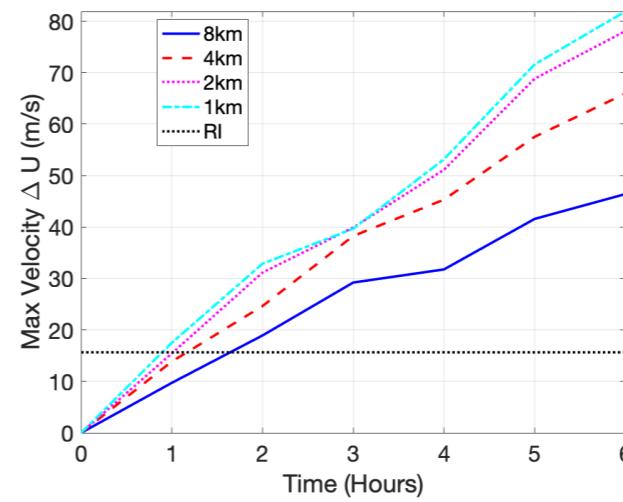
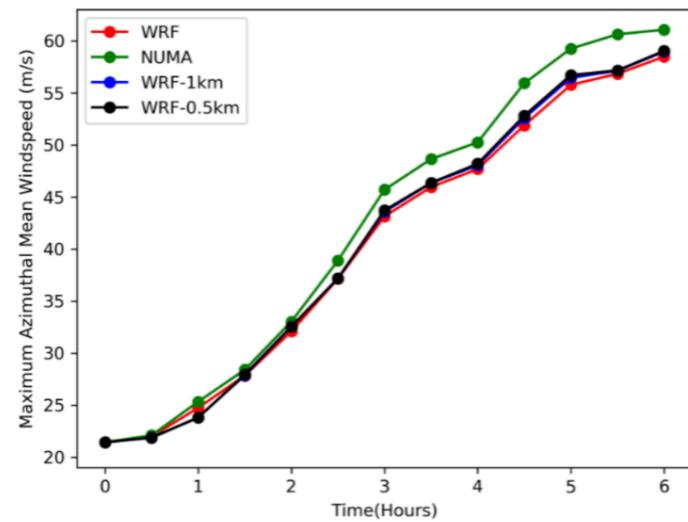


All NUMA equation sets give similar results. Mass conservation is imperative.



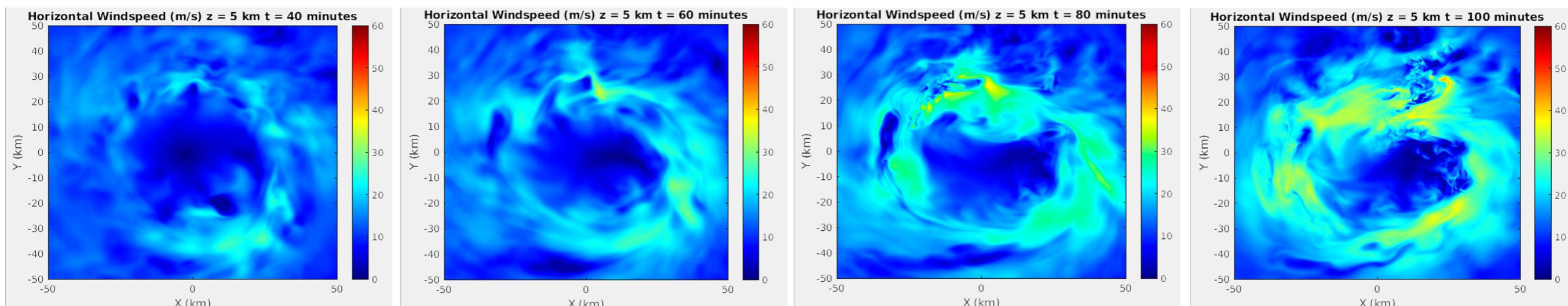
Result 2: Tropical Cyclone (CRM scale)

Tropical Cyclone Rapid Intensification (RI) (where max wind velocities exceed 30 mph within 24-hour period) is important to understand extreme weather. Here is a sample plot from [5] where we show our model vs the U.S. National Hurricane Center model. RI requires running with minimal dissipation and high resolution.



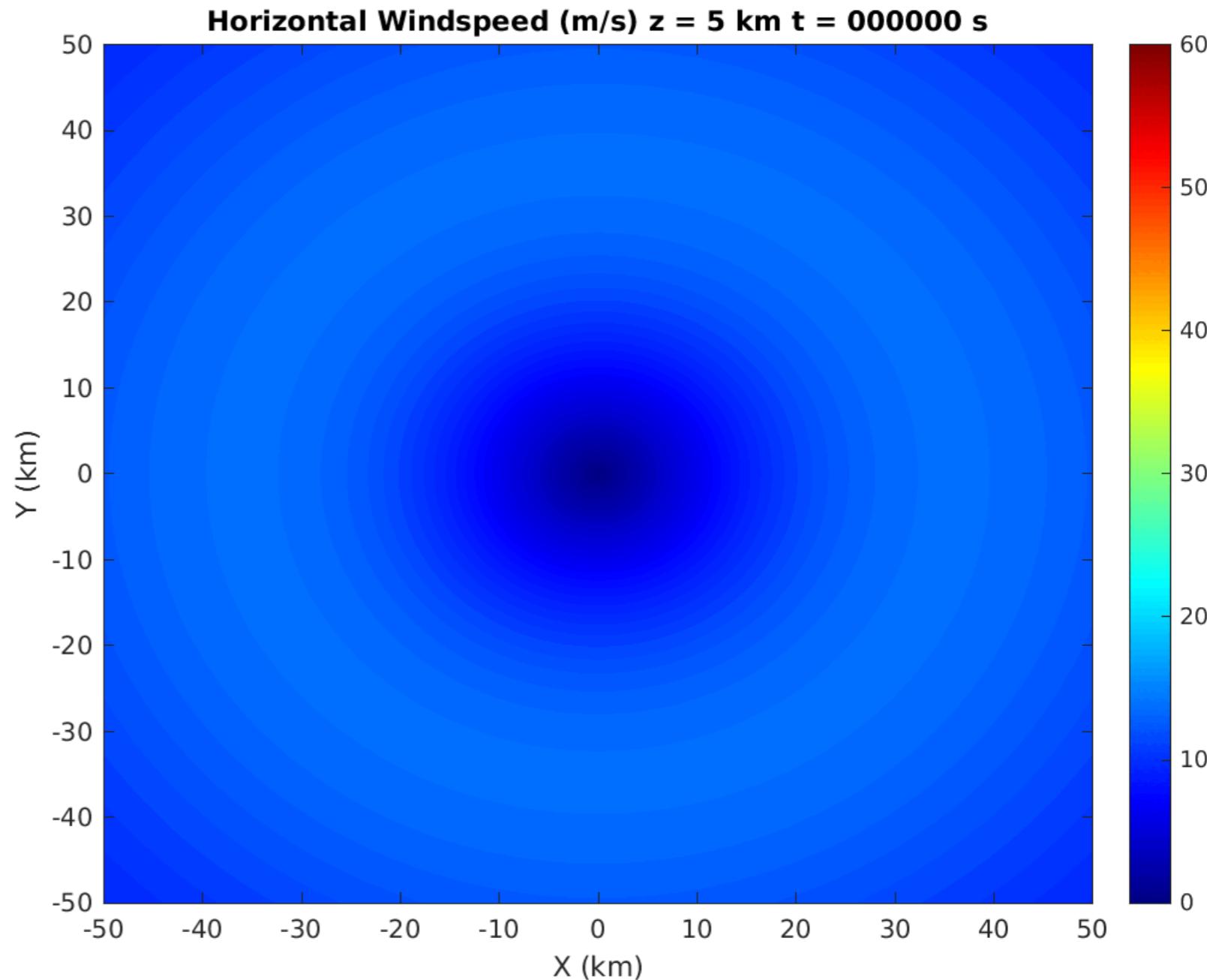
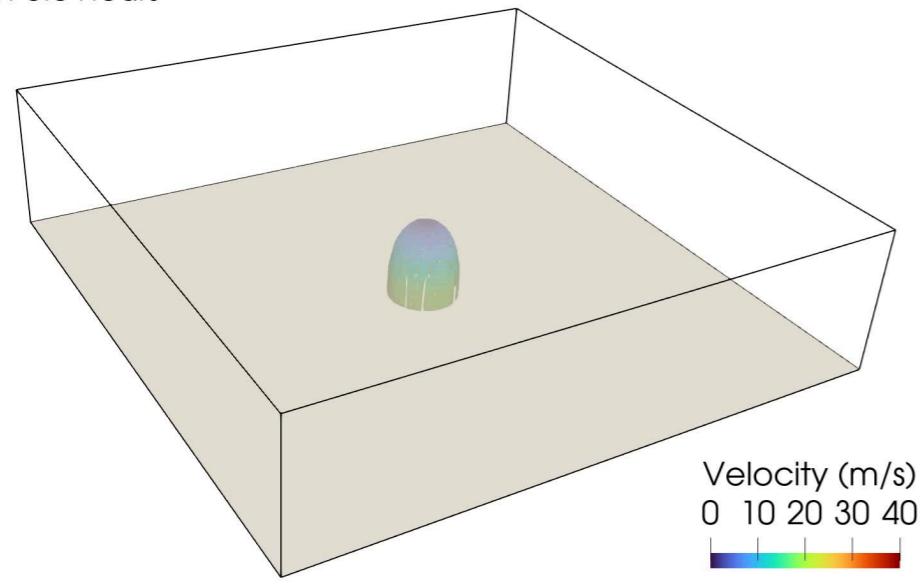
To properly capture RI requires LES models at $\Delta x = \mathcal{O}(100\text{ m})$. Running a CRM at this resolution is still too expensive (cannot be done on a regular basis). Our simulations in [5] were run at $\Delta x = \mathcal{O}(2\text{ km})$ (with 80 million DoF, cost is 1 simulation hour/wallclock hour on 5000 AMD Epyc cores). LES requires 32 billion DoF with Δt 20x smaller ($\mathcal{O}(10^4)$ more expensive). (The simulation shown below used 10 billion DoF (reduce the domain size); courtesy of Steve Guimond UMBC)

To make such simulations possible require leveraging: 1) multi-scale modeling frameworks (MMF), 2) dynamic adaptive mesh refinement, 3) new time-integration and preconditioning strategies, and 4) HPC.



Result 2: NUMA Tropical Cyclone (LES scale) [6]

Time: 0.0 hours



The eye-wall of Hurricane Guillermo undergoing RI. Real data consists of 4D observational latent heat forcing derived from airborne radar. Model resolution is $\Delta x = \Delta y = \mathcal{O}(60 \text{ m})$ and $\Delta z = \mathcal{O}(100 \text{ m})$, on 9000 AMD Epyc cores); courtesy of Steve Guimond, Hampton University)

Result 3: Density Current (LES-scale) [7]

$$Ra = 5 \times 10^6$$

$$Sc = 1$$

$$v_h = 1.17 \text{ m}^2/\text{s}$$

$$\Delta x, \Delta y = 6 - 25 \text{ m}$$

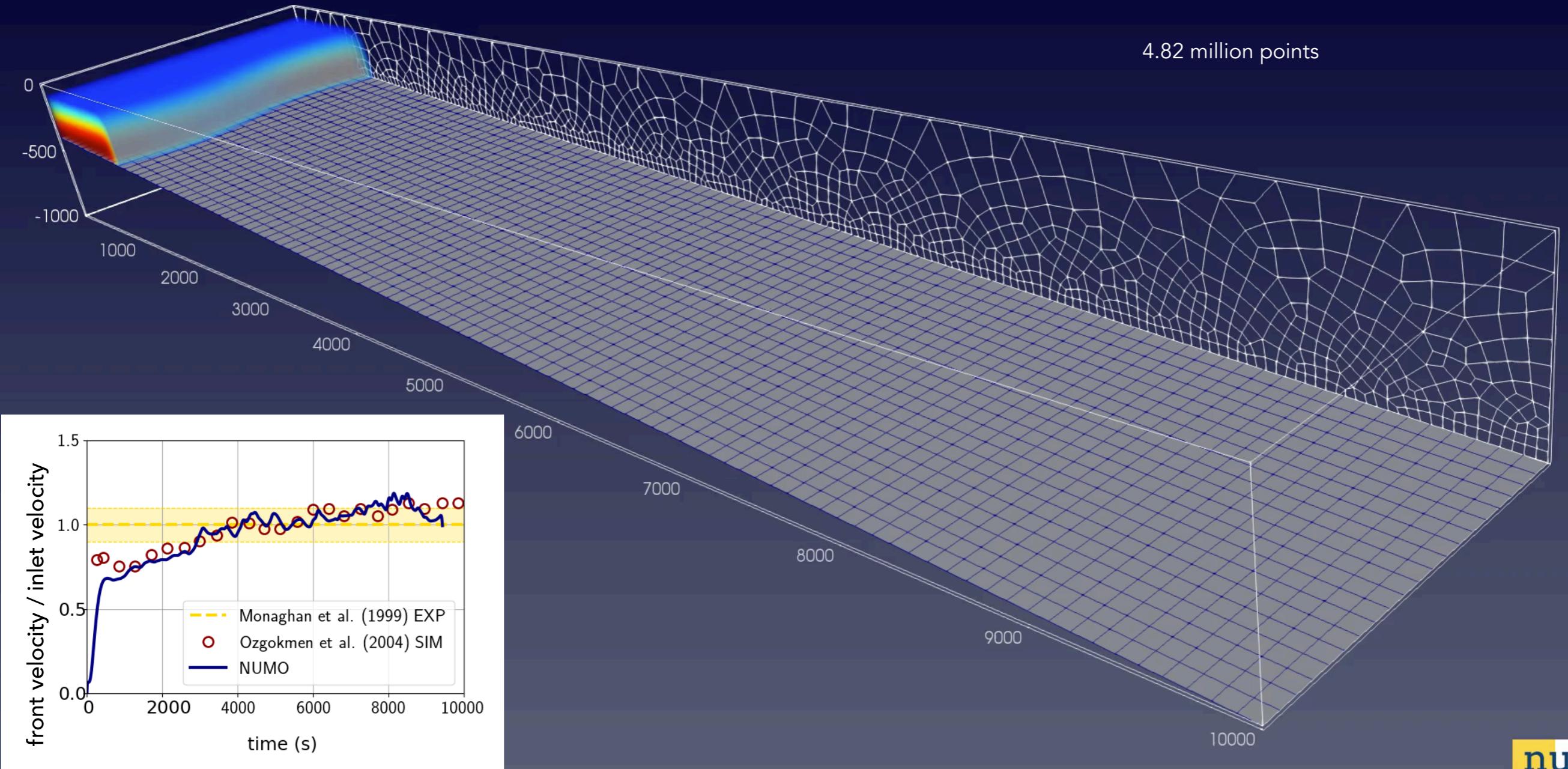
$$t = 10.0\text{s}$$

$$v_v = 2.34 \times 10^{-2} \text{ m}^2/\text{s}$$

$$\Delta z = 25\text{m}$$

38,560 elements

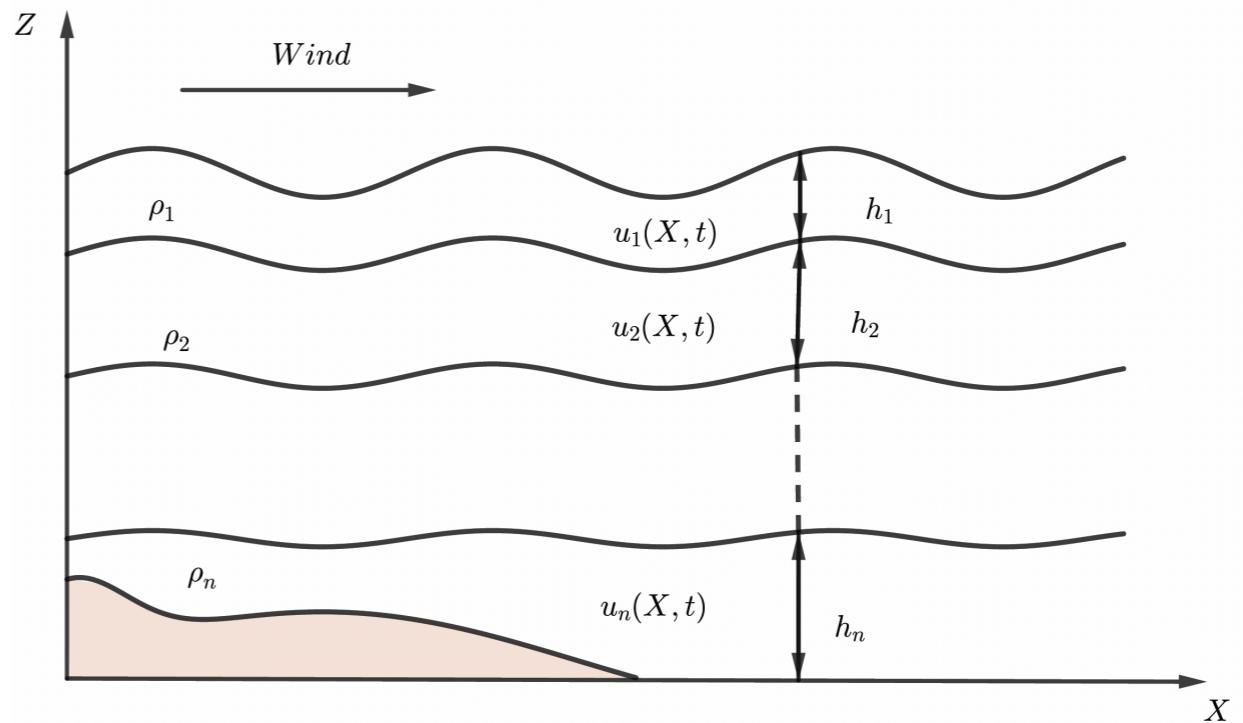
4.82 million points



numo

Result 4: Double-Gyre (GCM scale) [8]

- Multi-layered shallow water equations
- Hydrostatic
- Isopycnal coordinates (layers of constant density)
- Equations are solved for Δp_k and \mathbf{U}_k



$$\frac{\partial \Delta p_k}{\partial t} + \nabla \cdot \mathbf{U}_k = 0,$$

$$\begin{aligned} \frac{\partial \mathbf{U}_k}{\partial t} + \nabla \cdot \left(\frac{\mathbf{U}_k \otimes \mathbf{U}_k}{\Delta p_k} + H_k \mathbf{I} \right) + f \mathbf{k} \times \mathbf{U}_k &= g (p_{k-1} \nabla z_{k-1} - p_k \nabla z_k) \\ &\quad + g \Delta \tau_k + \nabla \cdot (\nu \Delta p_k \nabla \mathbf{u}_k), \end{aligned}$$

$\Delta p_k = p_k - p_{k-1}$ is the layer pressure

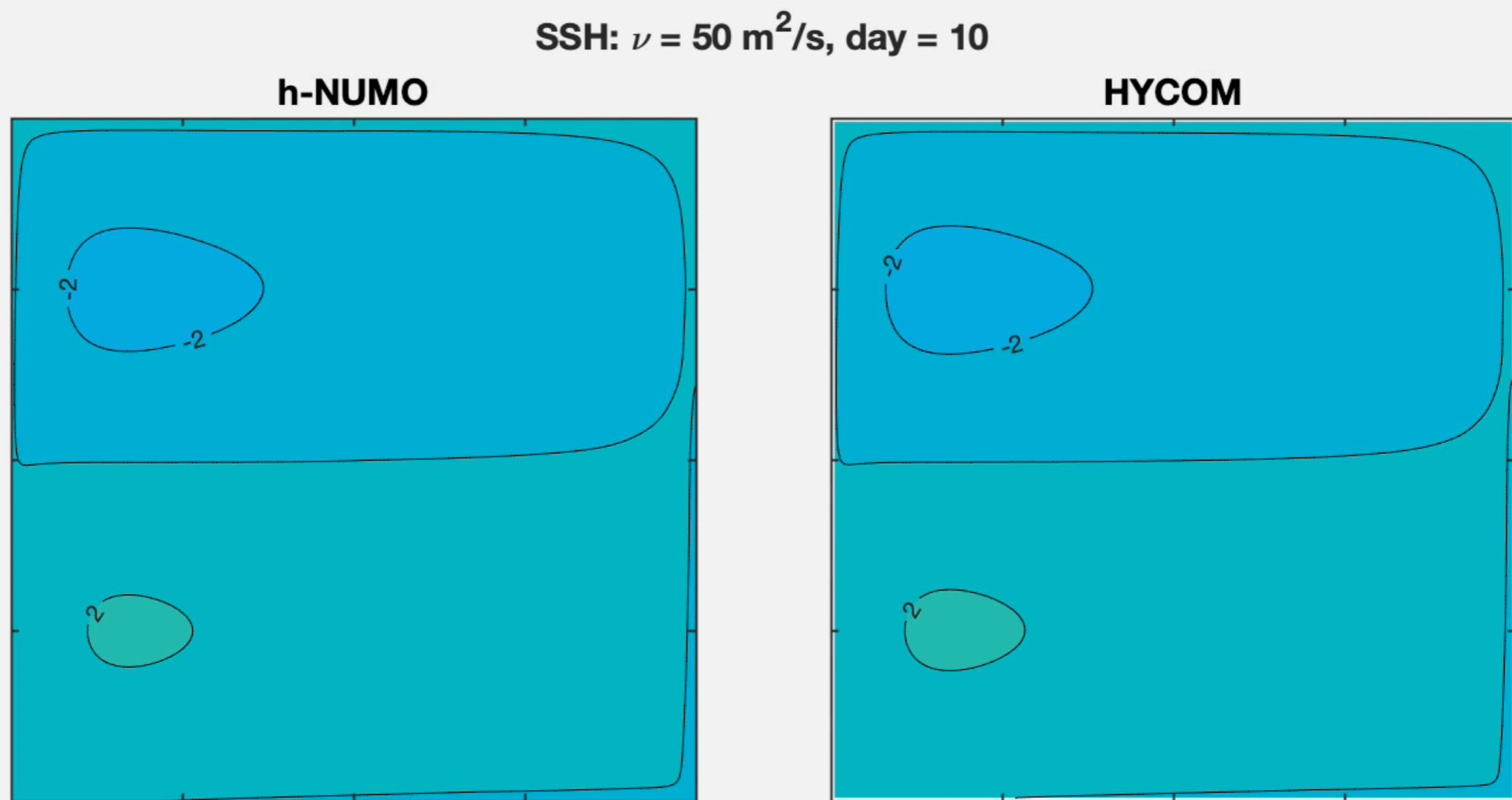
$\mathbf{U}_k = \mathbf{u}_k \Delta p_k$ is the layer momentum

$$H_k(x, y, t) = g \int_{z_k}^{z_{k-1}} P(x, y, z, t) dz$$

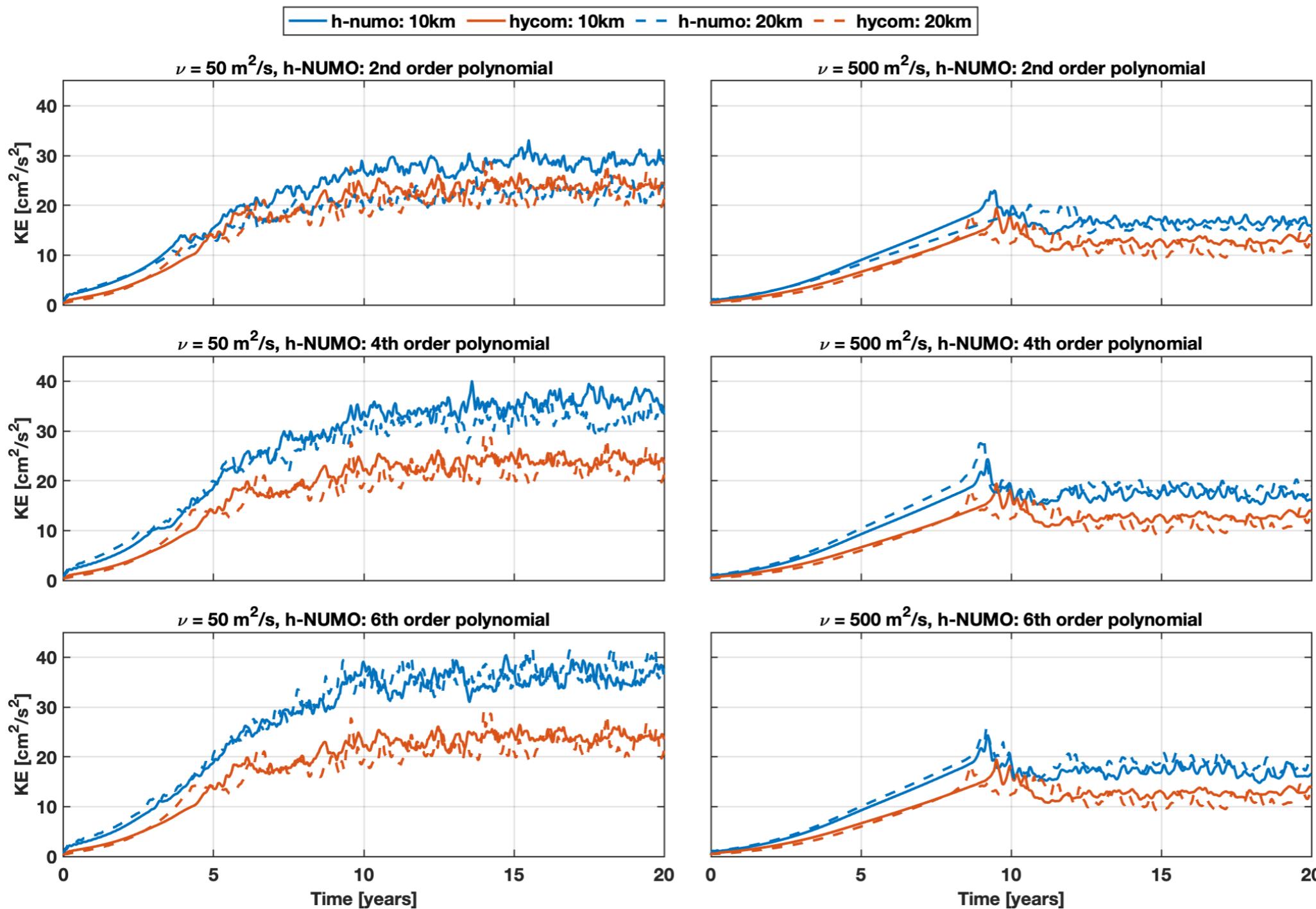
$\Delta \tau_k$ is the shear stress

Result 4: Double-Gyre (GCM scale)

- Comparison between h-NUMO and HYCOM; HYCOM is the U.S. Navy's operational global ocean model
- 2 layers using at $\Delta x = 10 \text{ km}$ resolution integrated for 20 years with $\nu = 50 \text{ m}^2/\text{s}$
- xy-slice showing sea surface height



Result 4: Double-Gyre (GCM scale)



h-NUMO
shows higher
KE for the
same
resolution
and viscosity

Closing Remarks

- **Element-based Galerkin Methods**
 - The minimal inherent dissipation of EBG/high-order methods offers the possibility of making advances in better understanding of important weather/ocean dynamics (e.g., TC RI in atmosphere and more accurate kinetic energy in the ocean).
 - Some (i.e., continuous Galerkin) are entirely non-dissipative and need help to maintain stability. Others (i.e., discontinuous Galerkin) have more mechanisms to introduce dissipation but still need some help. Entropy-Stable (ES) methods allow for mathematically rigorous construction of stable methods but have high complexity (5-10x more FLOPS) [**see MS01,MS03,MS05 today**]. Can we extend these to CG?
 - Efficiency continues to be the Achilles heel of these methods.

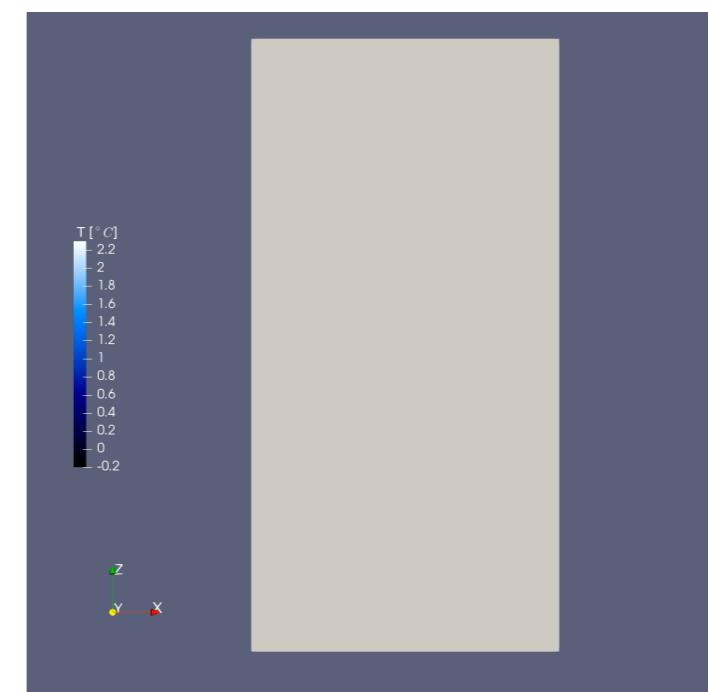
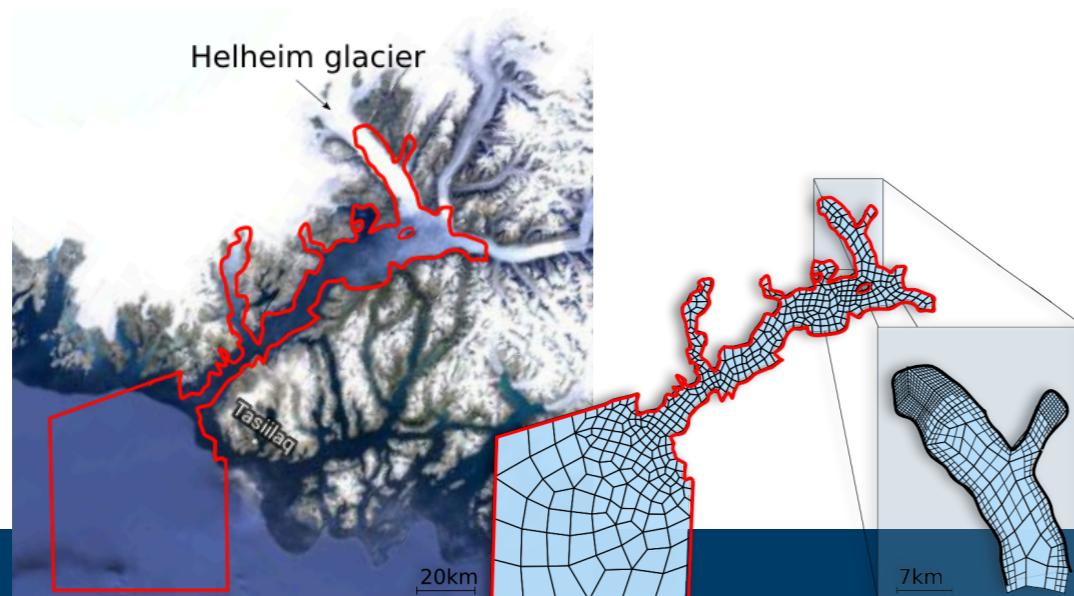
Closing Remarks

- **Future of Atmospheric Models**
 - As resolutions of GCMs approach cloud-resolving scales $\mathcal{O}(1 \text{ km})$, many of the important processes are resolved explicitly (less reliance on parameterizations).
 - However, the LES scale $\mathcal{O}(100 \text{ m})$ is quite out of reach at the moment.
 - To curtail the cost of LES-scale simulations requires using: (1) MMF (e.g., see [7]); (2) HPC particularly on GPUs [**see MS09 and Noel Chalmers' plenary Tuesday morning**]; and (3) Scientific Machine Learning/ROM [**see MS02, MS04, MS06, MS13**].

Closing Remarks

- **Future of Ocean Models**

- Recent MOU on Ocean Sciences between NPS and Stanford Doerr School of Sustainability provides an exciting opportunity to develop a new EBG ocean model to couple with NUMA/NEPTUNE.
- Currently, all operational ocean models are hydrostatic (few exceptions). Is hydrostatic sufficient or do we need to switch to nonhydrostatic as the atmospheric community recently did?
- SBP methods should play a key role in developing stable high-order ocean models.
- In some ways, EBG may be more optimal for ocean models due to the need for unstructured grids (e.g., Greenland fjords).



Melting Ice at Ice/Ocean Interface [5]

Closing Remarks

- **Final Thoughts**
 - Did not discuss time-integration (TI) but much work to be done. A brief list includes: (1) Multi-rate methods [**see Carol Woodward's plenary on Wednesday**], (2) implicit methods for SBP [**see talks on SBP minis on this topic today**]. TI for GFD needs to consider the disparate spatial scales between the horizontal and vertical directions.
 - High-order methods are in very good shape and are competing with other methods for operational/industrial-type applications in terms of both stability and performance (e.g., 2024 Gordon Bell award to E3SM team)
 - NAHOMCon is a great venue for us to continue to share our experiences to improve HO methods - thanks to the organizers and participants and let's keep it going!