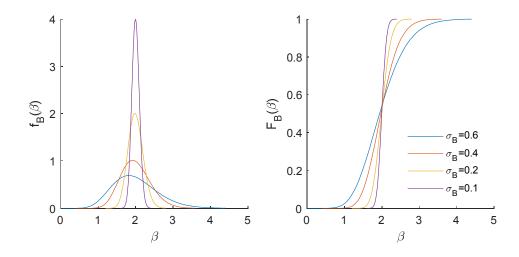


A single point source generates earthquakes of magnitude with a truncated exponential distribution (NM<sub>min</sub>=2,  $M_{min}$  = 4,  $M_{max}$  = 8,  $\bar{\beta}$  = 2,  $\sigma_{\beta}$  = {0.1, 0.2, 0.4, 0.6}). Use the Sadigh et al. 1997 GMM (strike-slip, rock) with untruncated sigma to compute the seismic hazard curve for PGA for site located 20 km from the hypocenter.

To account for the uncertainty in  $\beta$ , use a Gamma probability density function (Keller, 2014) with mean  $\bar{\beta} = 2$  and standard deviation  $\sigma_{\beta}$ , truncated at  $\pm 2\sigma_{\beta}$ . The functional for the truncated Gamma pdf is

$$f_B(\beta) = \frac{\beta^{k-1} e^{-\beta/\theta}}{\Gamma(k\theta)\theta^k} \cdot \frac{1}{F_B(\bar{\beta} + 2\sigma_b) - F_B(\bar{\beta} - 2\sigma_b)}$$

with  $\bar{\beta} - 2\sigma_b \leq \beta \leq \bar{\beta} + 2\sigma_b$ . In this expression,  $F_B(\beta) = \frac{1}{\Gamma(k)} \gamma(k, \beta/\theta)$  is cumulative distribution function for the Gamma distribution, the  $k = \frac{\bar{\beta}^2}{\sigma_{\beta}^2}$  is the shape parameter,  $\theta = \frac{\sigma_{\beta}^2}{\bar{\beta}}$  is the scale parameter, and  $\gamma(s,x) = \int_0^x t^{s-1} e^{-t} dt$  is the incomplete lower Gamma function. For the different  $\sigma_{\beta}$  values considered, the pdf and cdf of  $\beta$  are shown herein.



Evaluating Sadigh et al. 1997 (strike-slip, rock) for PGA and r = 20 km leads to

$$\ln PGA = \begin{cases} -0.624 + 1.1m - 2.1\ln(20 + \exp(1.29649 + 0.250 \cdot m)) & \text{if } m \le 6.5\\ -1.274 + 1.1m - 2.1\ln(20 + \exp(-0.48451 + 0.524 \cdot m)) & \text{if } m > 6.5 \end{cases}$$

$$\sigma_{lnPGA} = \begin{cases} 1.39 + 0.14m & if \ m < 7.21 \\ 0.38 & if \ m \ge 7.21 \end{cases}$$

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Then, the probability that PGA exceeds a test value y is  $P(Sa > y|m) = 1 - \Phi\left(\frac{\ln y - \ln PGA}{\sigma_{lnPGA}}\right)$ , where  $\Phi$  is the cumulative function for the standard normal distribution. The probability density function for a Truncated Exponential magnitude is

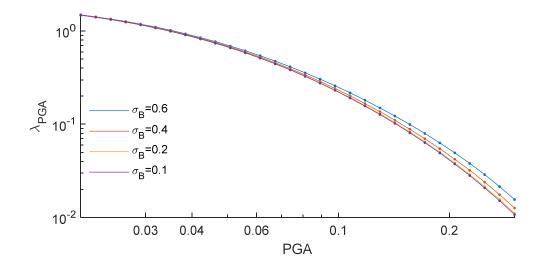
$$f_M(m|\beta) = \frac{\beta e^{-\beta(m-M_{min})}}{1 - e^{-\beta(M_{max} - M_{min})}}$$

Finally, based on total probabilities, the annual rate of exceedance  $\lambda_{PGA}$  is

$$\lambda_{PGA} = NM_{min} \int_{M_{min}}^{M_{max}} \int_{0}^{\infty} P(Sa > y|m) f_{M}(m|\beta) f_{B}(\beta) d\beta dm$$
 
$$\lambda_{PGA}(y) = NM_{min} \int_{M_{min}}^{M_{max}} \int_{0}^{\infty} H(PGA - y) \frac{\beta e^{-\beta(m - M_{min})}}{1 - e^{-\beta(M_{max} - M_{min})}} \frac{\beta^{k-1} e^{-\beta/\theta}}{\Gamma(k\theta) \theta^{k}} d\beta dm$$

The platform SeismicHazard performs this double integral using two-way Gauss quadrature, leading to

$$\Lambda_{PGA}(y) = NM_{min} \sum_{i=1}^{n_{gp}} \sum_{j=1}^{n_{M}} H(PGA(m_{j}) - y) f_{m}(m_{j}, \beta_{i}) f_{\beta}(\beta_{i}) \omega_{i} \omega_{j}$$



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## Independent MATLAB calculation

```
NMmin = 2;
Mmin = 4;
Mmax = 8;
Beta = 1i
sigmaB = [0.6 \ 0.4 \ 0.2 \ 0.1];
Rrup = 20i
nsigB = 2;
     = linspace(Mmin,Mmax,100);
      = logsp(0.02, 0.3, 30);
У
lny = log(y);
hold on
for ii=1:length(sigmaB)
    bint = [max(Beta-nsigB*sigmaB(ii),0.01),Beta+nsigB*sigmaB(ii)];
         = linspace(bint(1),bint(2),100);
         = makedist('gamma',Beta^2/sigmaB(ii)^2,sigmaB(ii)^2/Beta);
    pd
    pd
         = truncate(pd,bint(1),bint(2));
    [m,beta] = meshgrid(M,b);
    rrup = ones(size(m))*Rrup;
    [lnpga,sig] = Sadigh1997(0,m,rrup,'rock','strike-slip');
    PGA = exp(lnpga);
    lambda = zeros(size(y));
    for i=1:length(y)
            = 1-normcdf((lny(i)-lnpga)./(sig));
        fm = beta.*exp(-beta.*(m-Mmin))./(1-exp(-beta*(Mmax-Mmin)));
             = pdf(pd,beta);
             = P.*fm.*fB;
        lambda(i) = NMmin*trapz(M,trapz(b,f));
    plot(y,lambda,'.-
', 'displayname', ['\sigma_B=', sprintf('%g', sigmaB(ii))]),
end
xlabel('PGA')
ylabel('\lambda_{PGA}')
set(gca,'yscale','log','xscale','log','xtick',[0.03 0.04 0.06 0.1 0.2
0.3])
L=legend;
L.Box='off';
```

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## Independent MATLAB implementation

```
NMmin = 2;
Mmin
       = 5;
Mmax = 6.5;
      = 1;
beta = b*log(10);
rrup = 20;
     = linspace(Mmin,Mmax,100000);
       = [-0.624 \ 1.0 \ 0.000 \ -2.100 \ 1.29649 \ 0.250 \ 0.0];
C
lny = C(1)+C(2)*M+C(4)*log(rrup+exp(C(5)+C(6)*M));
       = [1.39 \ 0.14 \ 0.38 \ 7.21];
sigma = C(1)-C(2)*M;
    = logsp(0.01,1,30);
lambda = zeros(size(y));
for i=1:length(y)
    xhat = (log(y(i)) - lny)./sigma;
         = 1-normcdf(xhat);
         = beta*exp(-beta*(M-Mmin))./(1-exp(-beta*(Mmax-Mmin)));
    lambda(i) = NMmin*trapz(M,P.*fm);
end
loglog(y,lambda,'.-')
```

## References

- Keller, M., A. Pasanisi, M. Marcilhac, T. Yalamas, R. Secanell, and G. Senfaute (2014). A Bayesian methodology applied to the estimation of earthquake recurrence parameters for seismic hazard assessment, Qual. Reliab. Eng. Int. 30, no. 7, 921–933
- Poulos, A., Monsalve, M., Zamora, N., & de la Llera, J. C. (2018). An Updated Recurrence Model for Chilean Subduction Seismicity and Statistical Validation of Its Poisson Nature. Bulletin of the Seismological Society of America, 109(1), 66-74.