# The Influence of Temperature on Ozone Production under

varying NO<sub>x</sub> Conditions – a modelling study

J. Coates<sup>1</sup> and T. Butler<sup>1</sup>

<sup>1</sup>Institute for Advanced Sustainability Studies, Potsdam, Germany

January 5, 2016

Include Katie and Noelia as co-authors for providing ERA and WRF-Chem data or does an acknowledgement suffice?

Abstract

Ground-level ozone is a secondary air pollutant produced during the degradation of emitted volatile organic compounds (VOCs) and nitrogen oxides ( $NO_x$ ) in the presence of sunlight. As ozone production is dependent on photochemical processes, meteorological factors such as temperature influence the ozone levels. Temperature directly influences ozone production through speeding up the rates of the chemical processes producing ozone and increasing the emissions of VOCs, such as isoprene, from vegetation. In this study, we used a box model to reproduce the non-linear relationship of ozone on  $NO_x$  and temperature from previous observational studes. Faster chemistry was responsible for an increase in ozone of up to 20 ppbv while increased isoprene emissions added a further 11 ppbv of ozone under high- $NO_x$  conditions. The decrease in the lifetime of peroxy nitrates with increased temperature was the main contributor to the increased production of ozone with temperature. At 40 °C, the thermal decomposition of peroxy nitrates was responsible for up to 45 % of the normalised  $O_x$  production. The rate of increase in ozone with temperature from our box model simulations was about half of both the observed and modelled, using the 3D WRF-Chem model, rate of the observed increase in ozone with temperature over central Europe.

## $_{\scriptscriptstyle 3}$ 1 Introduction

Surface-level ozone  $(O_3)$  is a secondary air pollutant formed during the photochemical degradation

of volatile organic compounds (VOCs) in the presence of nitrogen oxides (NO  $_{\rm x} \equiv {\rm NO} + {\rm NO}_2).$  Due

to the photochemical nature of ozone production, meteorological variables such as temperature strongly influence ozone production (Jacob and Winner, 2009). by Otero et al. (2016) showed that temperature was a major meteorological driver for summertime ozone in many areas of central Europe.

Temperature primarily influences ozone production in two ways: speeding up the reaction rates of many chemical reactions leading to ozone production and increasing emissions of VOCs from biogenic sources (BVOCs). In general, emissions of anthropogenic VOCs (AVOCs) are not typically dependent on temperature, however evaporative emissions of AVOCs tend to increase with temperature (Rubin et al., 2006). The review of Pusede et al. (2015) provides further details of the temperature-dependent processes impacting ozone production.

Studies over the US (Sillman and Samson, 1995; Dawson et al., 2007; Pusede et al., 2014)
noted that increased temperatures tend to lead to higher ozone levels, often exceeding local
air quality guidelines. Some studies (Sillman and Samson, 1995; Dawson et al., 2007) included
regional modelling to simulate the observed increases in ozone with temperature. In these studies,
the increase of ozone with temperature was attributed to the decrease in the lifetime of PAN
(peroxy acetyl nitrate) at higher temperatures and increased emissions of BVOCs, in particular
isoprene, from vegetation.

Pusede et al. (2014) used an analytical model constrained by observations over San Joaquin Valley, California to infer a non-linear relationship of ozone production with temperature and NO<sub>x</sub>, similar to the well-known non-linear relationship of ozone production on NO<sub>x</sub> and VOC levels (Sillman, 1999). Morever, Pusede et al. (2014) showed that temperature can be used as a surrogate for VOC levels when considering the relationship of ozone across NO<sub>x</sub> gradients.

Environmental chamber studies have also been used to analyse the relationship of ozone with temperature for a particular mixture of VOCs. The chamber experiments of Carter et al. (1979) and Hatakeyama et al. (1991) showed increases in ozone with temperature linked to increased PAN decomposition at temperatures greater than 303 K.

Despite many studies considering the effects of temperature on ozone production from an observational and chamber study perspective, there are no modelling studies (to our knowledge) focusing on the detailed chemical processes of the influence of temperature on ozone production across  $NO_x$  gradients. Regional modelling studies have concentrated on reproducing ozone levels over regions with known meteorology and  $NO_x$  conditions and then varying the temperature and not considering the relationship of ozone with  $NO_x$  with temperature. The review of Pusede

et al. (2015) also highlights a lack of modelling studies looking at the non-linear relationship of ozone on temperature across  $NO_x$  gradients.

In this study, we use an idealised box model to determine how ozone levels vary with temperature and across  $NO_x$  gradients. We determine whether faster chemistry or increased BVOC emissions have a greater influence on instantaneous ozone production with higher temperature at different  $NO_x$  conditions. Rasmussen et al. (2013) indicated that changing the chemical mechanism used by a model may also change the simulated ozone-temperature relationship to investigate this, we repeated all simulations using various chemical mechanisms.

# 66 2 Methodology

### 67 2.1 Model Setup

We used the MECCA box model to determine the important chemical processes for ozone production under different temperatures and  $NO_x$  conditions. The MECCA box model was set up as described in Coates and Butler (2015) and updated to include vertical mixing with the free troposphere and a diurnal cycle for the PBL height. The supplementary material includes further details of these updates.

Simulations were performed to broadly simulate urban conditions of central Europe with equinoctical conditions. The simulations started at 06:00 with a total run time of two days.

Methane was fixed at 1.7 ppmv throughout the model run, carbon monoxide (CO) and ozone were initialised at 200 ppbv and 40 ppbv and then allowed to evolve freely throughout the simulation. All VOC emissions were held constant until noon of first day simulating a plume of freshly-emitted VOC.

Model runs were repeated using a temperature-dependent and temperature-independent source
of BVOC emissions to determine whether increased emissions of BVOC or faster chemistry is more
important for the increase of ozone with temperature. MEGAN2.1 (Guenther et al., 2012) specified
the temperature-dependent BVOC emissions of isoprene and Sect. 2.3 provides further details. We
considered only isoprene emissions from vegetation as isoprene emissions are the most important
on the global scale (Guenther et al., 2006). Only temperature-dependent emissions of isoprene
were considered, all other emissions were constant in all simulations. In reality, many other BVOC
are emitted from varying vegetation types (Guenther et al., 2006) and increased temperature
can also increase AVOC emissions through increased evaporation (Rubin et al., 2006).

Table 1: Total AVOC emissions in 2011 in tonnes from each SNAP category assigned from TNO-MACC\_III emission inventory and temperature-independent BVOC emissions in tonnes from Benelux region assigned from EMEP. The allocation of these emissions to MCMv3.2, CRIv2, CB05, MOZART-4 and RADM2 species is found in the supplementary material.

	SNAP1	SNAP2	SNAP34	SNAP5	SNAP6	SNAP71
Belgium	4494	9034	22152	5448	42809	6592
Netherlands	9140	12173	29177	8723	53535	16589
Luxembourg	121	44	208	1371	4482	1740
Total	13755	21251	62648	15542	100826	24921
	SNAP72	SNAP73	SNAP74	SNAP8	SNAP9	BVOC
Belgium	2446	144	210	6448	821	7042
Netherlands	3230	1283	1793	10067	521	1462
Luxembourg	1051	6	324	643	0	2198
Total	6727	1433	2327	17158	1342	10702

All simulations were repeated using different chemical mechanisms to investigate how well the relationship of ozone with temperature across NO<sub>x</sub> gradients is represented by different 89 representations of ozone production chemistry. The reference chemical mechanism was the 90 near-explicit Master Chemical Mechanism, MCMv3.2, (Jenkin et al., 1997, 2003; Saunders et al., 91 2003; Rickard et al., 2015). The reduced chemical mechanisms in our study were Common 92 Representative Intermediates, CRIv2 (Jenkin et al., 2008), Model for ozone and related chemical tracers, MOZART-4 (Emmons et al., 2010), Regional Acid Deposition Model, RADM2 (Stockwell et al., 1990) and the Carbon Bond Mechanism, CB05 (Yarwood et al., 2005). Coates and Butler (2015) described these chemical mechanisms and the implementation of these chemical 96 mechanisms in MECCA. These reduced chemical mechanisms were chosen as they are commonly 97 used by modelling groups in 3D regional and global models (Baklanov et al., 2014).

Box model simulations were performed by systematically varying the temperature between 288 and 313 K (15 – 40 °C). The only source of  $NO_x$  emissions in the box model was a constant source of NO emissions. Box model runs were performed with the NO emissions systematically varied from 5.0  $\times$  10<sup>9</sup> to 1.5  $\times$  10<sup>12</sup> molecules (NO) cm<sup>-2</sup> s<sup>-1</sup> at each temperature used in this study. At 20 °C, these NO emissions corresponded to peak  $NO_x$  mixing ratios of 0.02 ppbv and 10 ppbv respectively.

#### 2.2 VOC Emissions

105

Emissions of urban AVOC over central Europe were taken from TNO-MACC\_III emission inventory for the Benelux (Belgium, Netherlands and Luxembourg) region for the year 2011.

TNO-MACC\_III is the updated version of the TNO-MACC\_II emission inventory created using
the same methodology as Kuenen et al. (2014) and based upon improvements to the existing
emission inventory during AQMEII-2 (Pouliot et al., 2015).

Temperature-independent emissions of isoprene and monoterpenes from biogenic sources were calculated as a fraction of the total AVOC emissions from each country in the Benelux region.
This data was obtained from the supplementary data available from the EMEP (European Monitoring and Evaluation Programme) model (Simpson et al., 2012). Temperature-dependent emissions of isoprene are detailed in Sect. 2.3.

AVOC emissions were allocated to SNAP (Selected Nomenclature for Air Pollution) source 116 categories. Table 1 shows the tonnes of VOC emissions from each SNAP category and the 117 temperature-independent BVOC emissions. These categorised AVOC emissions were assigned to 118 chemical species and groups based on the country specific profiles for Belgium, the Netherlands 119 and Luxembourg provided by TNO. Most individual chemical species are represented by the 120 MCMv3.2, otherwise the individual contributions of a group of VOC were further split into individual components using the detailed speciation of Passant (2002). For example, 'xylenes' 122 are one of the component chemical groups in many SNAP categories but the MCMv3.2 treats 123 xylenes as the individual isomers (m-, o-, p-xylene) and the contributions of the individual 124 isomers to a SNAP category was provided by Passant (2002). This approach was also used in von 125 Schneidemesser et al. (2016) to allocate AVOC emissions from different solvent sector speciations 126 to MCMv3.2 species. 127

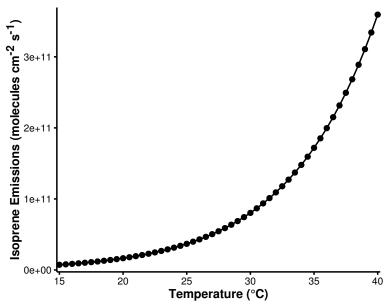
The VOC emissions represented by the MCMv3.2 were mapped to the mechanism species representing VOC emissions in each reduced chemical mechanism based on the recommendations of the source literature. The VOC emissions in the reduced chemical mechanisms were weighted by the carbon numbers of the MCMv3.2 species and the emitted mechanism species. The supplementary data outlines the primary VOC and calculated emissions with each chemical mechanism.

## 2.3 Temperature Dependent Isoprene Emissions

134

Temperature-dependent emissions of isoprene were estimated using the MEGAN2.1 algorithm for calculating the emissions of VOC from vegetation (Guenther et al., 2012). Emissions from plants are dependent on variables including temperature, radiation and age but for the purpose of our study all variables except temperature were held constant.

Figure 1: The estimated isoprene emissions (molecules isoprene  $cm^{-2} s^{-1}$ ) using MEGAN2.1 at each temperature used in the study.



The MEGAN2.1 parameters were chosen to give similar isoprene mixing ratios at 20 °C to the temperature-independent emissions of isoprene in order to compare the effects of increased isoprene emissions with temperature. The estimated emissions of isoprene with MEGAN2.1 using these assumptions, are illustrated in Fig. 1 and show the expected exponential increase in isoprene emissions with temperature (Guenther et al., 2006).

The estimated emissions of isoprene at 20 °C lead to 0.07 ppbv of isoprene in our simulations while at 30 °C, the increased emissions of isoprene using MEGAN2.1 estimations lead to 0.35 ppbv of isoprene in the model. A measurement campaign over Essen, Germany (Wagner and Kuttler, 2014) measured 0.1 ppbv of isoprene at temperature 20 °C and 0.3 ppbv of isoprene were measured at 30 °C. The similarity of the simulated and observed isoprene mixing ratios indicates that the MEGAN2.1 variables chosen for calculating the temperature-dependent emissions of isoprene were suitable for simulating urban conditions over central Europe.

## 3 Results and Discussion

# 3.1 Ozone as a Function of $NO_x$ and Temperature

Figure 2 depicts the maximum mixing ratio of ozone as a function of the total  $NO_x$  emissions on the first day of simulations and temperature when using a temperature-independent and temperature-dependent source of isoprene emissions for each chemical mechanism. A non-linear

Figure 2: Contours of maximum ozone mixing ratios as a function of the total  $\mathrm{NO_x}$  emissions on the first day and temperature for each chemical mechanism using a temperature-dependent and temperature-independent source of isoprene emissions.

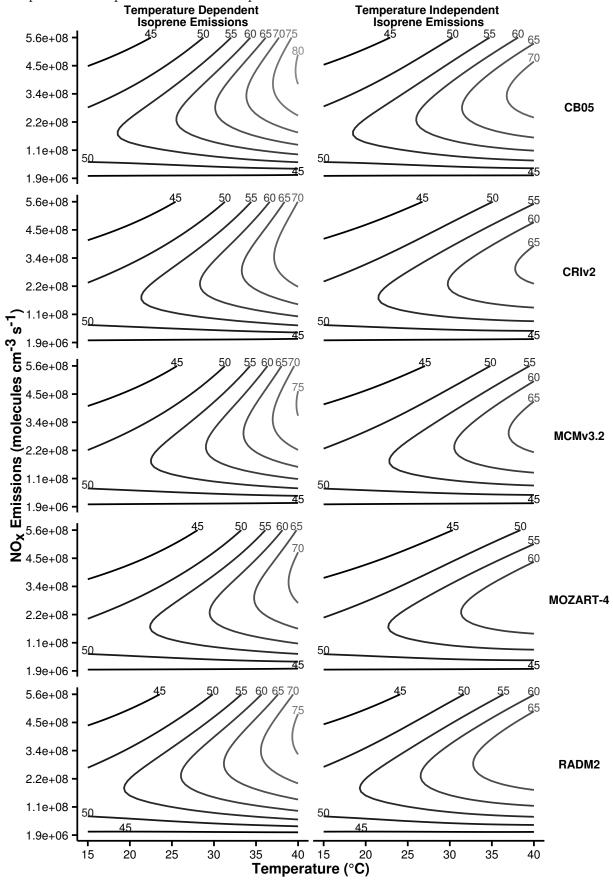
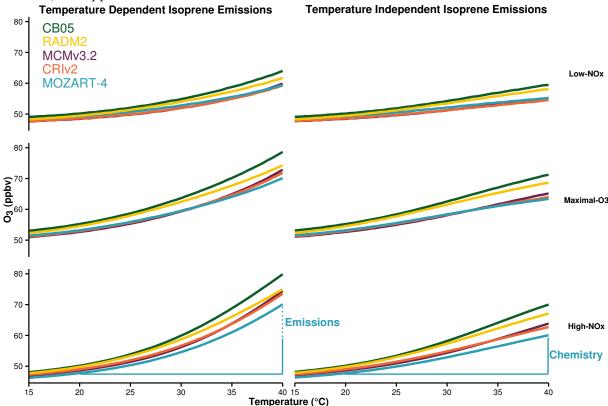


Figure 3: Ozone mixing ratios at each temperature are allocated to different  $NO_x$ -regimes of Fig. 2. The differences in ozone mixing ratios due to chemistry (solid line) and emissions (dotted line) are represented graphically for MOZART-4 with High- $NO_x$  conditions and summarised in Table 2, the approach was used to calculate the differences with each chemical mechanism.



relationship of ozone mixing ratios with  $\mathrm{NO_x}$  and temperature is reproduced by each chemical mechanism. This non-linear relationship is similar to that determined by Pusede et al. (2014) using an analytical model constrained to observational measurements over the San Joaquin Valley in California.

Higher ozone mixing ratios are produced when using a temperature-dependent source of isoprene emissions in Fig. 2. The highest mixing ratios of ozone are produced at high temperatures and high emissions of  $\mathrm{NO}_{\mathrm{x}}$  regardless of the source of isoprene emissions. Conversely, the least amount of ozone is produced with low emissions of  $\mathrm{NO}_{\mathrm{x}}$  over the whole temperature range (15 – 40 °C) when using both a temperature-independent and temperature-dependent source of isoprene emissions.

The contours of ozone mixing ratios as a function of  $NO_x$  and temperature can be split into three  $NO_x$  regimes (Low- $NO_x$ , Maximal- $O_3$  and High- $NO_x$ ), similar to the  $NO_x$  regimes defined for the non-linear relationship of ozone with VOC and  $NO_x$ . The Low- $NO_x$  regime corresponds to regions with little increase in ozone with temperature, also called  $NO_x$ -sensitive regime. The High- $NO_x$  (or  $NO_x$ -saturated) regime is when ozone levels increase rapidly with temperature

Table 2: Increase in ozone mixing ratio (ppbv) due to chemistry and emissions at 40  $^{\circ}$ C from reference temperature (20  $^{\circ}$ C) in the NO<sub>x</sub>-regimes of Fig. 3.

Chemical	Source of	Increase in Ozone at 40 °C from 20 °C (ppbv)			
Mechanism Difference		$Low-NO_x$	${ m Maximal-O}_3$	$\operatorname{High-NO}_{\operatorname{x}}$	
MCMv3.2	Chemistry	6.8	12.5	15.2	
	Emissions	4.6	7.7	10.6	
CRIv2	Chemistry	6.0	11.1	13.7	
	Emissions	4.8	7.9	10.8	
MOZART-4	Chemistry	6.0	10.2	12.3	
	Emissions	4.1	6.7	10.0	
CB05	Chemistry	9.3	16.0	19.9	
	Emissions	4.6	7.4	9.8	
RADM2	Chemistry	8.6	14.1	17.3	
	Emissions	3.8	5.7	7.8	

and the contour ridges correspond to regions of maximal ozone production and we call this the Maximal- $O_3$  regime. Pusede et al. (2014) showed that temperature can be used as a proxy for VOC, thus we assigned the ozone mixing ratios from each box model simulation to a  $NO_x$  regimes based on the ratio of  $HNO_3$  to  $H_2O_2$  used by Sillman (1995) to designate ozone to  $NO_x$  regimes based on  $NO_x$  and VOC levels.

Fig. 3 illustrates the mean ozone mixing ratio at each temperature in the  $NO_x$  regimes for each chemical mechanism and each source of isoprene emissions. The absolute increase in ozone at 40 °C from 20 °C due to faster chemistry is the difference between ozone mixing ratios at 40 °C and 20 °C when using a temperature-independent source of isoprene emissions. When using a temperature-dependent source of isoprene emissions, the difference in ozone mixing ratios at 40 °C from 20 °C minus the increase due to faster chemistry, gives the absolute increase in ozone due to increased isoprene emissions. These differences are represented graphically in Fig. 3 and summarised in Table 2.

Table 2 shows that the absolute increase in ozone with temperature due to faster chemistry is larger than the absolute increase in ozone due to increased isoprene emissions for each chemical mechanism and each  $NO_x$  regime. The absolute increase in ozone is largest with High- $NO_x$  conditions and lowest with Low- $NO_x$  conditions (Fig. 3 and Table 2). The increase in ozone mixing ratio at 40 °C from 20 °C due to faster chemistry with High- $NO_x$  conditions is almost double that with Low- $NO_x$  conditions. We determine the chemical processes responsible for the increases in ozone mixing ratios with temperature by analysing  $O_x$  production budgets in Sect. 3.2.

The largest differences in ozone mixing ratios using the reduced chemical mechanisms (CRIv2,

 $^{193}$  MOZART-4, CB05 and RADM2) from the MCMv3.2 occur in the Maximal- $\mathrm{O_3}$  and High- $\mathrm{NO_x}$  regimes. Table 2 also indicates that all reduced chemical mechanisms except RADM2 have similar  $^{195}$  increases in ozone due to increased isoprene emissions to MCMv3.2. RADM2 produces 3 ppbv  $^{196}$  less ozone than the MCMv3.2 due to increased isoprene emissions in each  $\mathrm{NO_x}$  regime, indicating  $^{197}$  that this difference is due to how isoprene degradation chemistry is treated in RADM2.

The Tagged Ozone Production Potential (TOPP) defined in Butler et al. (2011) is a measure 198 of the number of molecules of ozone produced per molecule of VOC emitted. Coates and 199 Butler (2015) compared ozone production in different chemical mechanisms to the MCMv3.2 200 using TOPPs and showed that less ozone is produced per molecule of isoprene emitted using 201 RADM2 than with MCMv3.2. The degradation of isoprene has been extensively studied and 202 it is well-known that methyl vinyl ketone (MVK) and methacrolein are signatures of isoprene 203 degradation. All chemical mechanisms in our study except RADM2 explicitly represent MVK 204 and methacrolein (or in the case of CB05, a lumped species representing both these secondary 205 degradation products). RADM2 does not represent methacrolein and the mechanism species 206 representing ketones (KET) represents a mixture of acetone and methyl ethyl ketone (MEK). Thus 207 the secondary degradation of isoprene in RADM2 is unable to represent the ozone production 208 from the further degradation of the signature secondary degradation products of isoprene, MVK 209 and methacrolein. Updated versions of RADM2, RACM (Stockwell et al., 1997) and RACM2 210 (Goliff et al., 2013), sequentially included methacrolein and MVK and with these updates the 211 TOPP value of isoprene approached that of the MCMv3.2 (Coates and Butler, 2015). 212

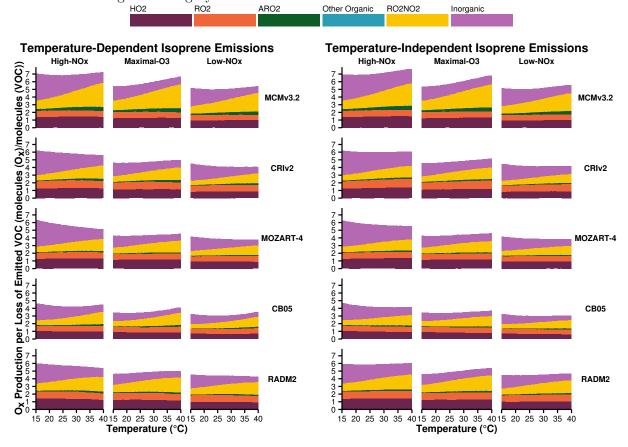
### 3.2 Ozone Production Budgets

213

The total day-time production budgets of  $O_x$  ( $\equiv O_3 + NO_2 + O$ ) normalised by the total rate of oxidation of the emitted VOC are displayed in Fig. 4. The  $O_x$  production budgets are assigned to each  $NO_x$  regime for each chemical mechanism and source of isoprene emissions. The budgets are allocated to the major sources, where 'HO2', 'RO2', 'ARO2' represent the reactions of NO with HO2, alkyl peroxy radicals and acyl peroxy radicals respectively. 'RO2NO2' represents the thermal decomposition of peroxy nitrates, 'Inorganic' represents all inorganic contributions to  $O_x$  production (primarily the de-excitation of  $O^1D$  to O) and any other remaining organic reactions producing  $O_x$  are included in the 'Other Organic' category.

In Fig. 4, a similar number of molecules of  $O_x$  per molecule of emitted VOC oxidised are produced in High-NO<sub>x</sub> conditions when using either temperature-dependent or

Figure 4: Day-time  $O_x$  production budgets normalised by the total oxidation rate of emitted VOC in the  $NO_x$ -regimes of Fig. 3. The budgets are allocated to categories of inorganic reactions, peroxy nitrate (RO2NO2) decomposition, reactions of NO with HO2, alkyl peroxy radicals (RO2) and acyl peroxy radicals (ARO2). All other reactions contributing to  $O_x$  budgets are allocated to the 'Other Organic' category.



temperature-independent isoprene emissions for each chemical mechanism, the same occurs for the Maximal- $O_3$  and Low- $NO_x$  regimes. Thus the increases in ozone production due to increased emissions of isoprene with temperature are balanced by the faster oxidation rates at higher temperatures. The largest amount of  $O_x$  is produced in the High- $NO_x$  regime and the lowest amount of  $O_x$  is produced in the Low- $NO_x$  regime, mirroring the  $O_3$  mixing ratios in the different  $NO_x$  regimes in Fig. 3. For example, MCMv3.2 produces seven molecules of  $O_x$  per molecule of emitted VOC oxidised in High- $NO_x$  conditions decreasing to about six and five molecules of  $O_x$  per molecule of emitted VOC oxidised in the Maximal- $O_3$  and Low- $NO_x$  regimes.

Thermal decomposition of RO2NO2 contributes the most to the normalised  $O_x$  production at higher temperatures in Fig. 4, this contribution shows a strong dependency on temperature and is analysed further in Sect. 3.2.1. The contributions of the reaction of NO with peroxy radicals (HO2, RO2 and ARO2 in Fig. 4) to the normalised production of  $O_x$  do not increase strongly with temperature indicating that the faster oxidation of emitted VOC at higher temperatures

 $_{237}$   $\,$  produces more peroxy radicals which when reacting NO fuels  $\mathrm{O}_{\mathrm{x}}$  production.

The reduced chemical mechanisms produce up to two molecules of  $O_x$  per molecule of emitted VOC oxidised less than the MCMv3.2 in each  $NO_x$  regime despite the reduced chemical mechanisms producing similar absolute amounts of ozone to the MCMv3.2 (Fig. 2 and Fig. 3). At high temperatures, up to 86 % of the total difference in the normalised  $O_x$  production using the reduced chemical mechanisms from the MCMv3.2 is due to differences in the contribution from peroxy nitrate (RO2NO2) decomposition.

#### 244 3.2.1 Peroxy Nitrates

Peroxy nitrates are formed from the reactions of peroxy radicals  $(RO_2)$  with  $NO_2$  (R1) and are an important reservoir species for both peroxy radicals and  $NO_x$ .

$$\mathrm{RO}_2 + \mathrm{NO}_2 + \mathrm{M} \rightleftharpoons \mathrm{RO}_2 \mathrm{NO}_2 + \mathrm{M} \tag{R1}$$

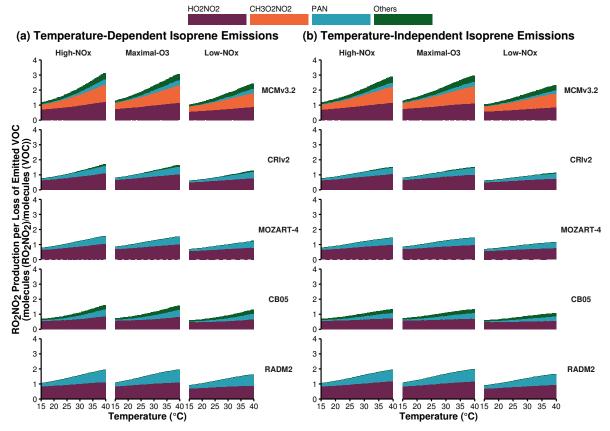
The chemical bond of  $RO_2NO_2$  is relatively weak and thermal decomposition is the most important chemical process with the rate of thermal decomposition increasing strongly with temperature. At low temperatures,  $RO_2NO_2$  can accumulate and be transported downwind of emissions of the sources of its precursors (VOC and  $NO_x$ ) and after thermal decomposition the release of  $NO_2$  and peroxy radicals can promote production of  $O_3$  (Moxim et al., 1996).

Peroxy nitrates are formed from both alkyl and acyl peroxy radicals produced during the

Peroxy nitrates are formed from both alkyl and acyl peroxy radicals produced during the secondary degradation of emitted VOC. The most important alkyl peroxy nitrates are pernitric acid  $(HO_2NO_2)$  and methylperoxy nitrate  $(CH_3O_2NO_2)$ , while peroxy acetyl nitrate  $(PAN, CH_3C(O)O_2NO_2)$  and peroxy propionyl nitrate  $(PPN, C_2H_5C(O)O_2NO_2)$  are important acyl peroxy nitrates. The alkyl peroxy nitrates have a weaker  $RO_2-NO_2$  bond than acyl peroxy nitrates hence alkyl peroxy nitrates have a shorter lifetime than acyl peroxy nitrates. At 298 K, the lifetime of  $CH_3O_2NO_2$  is 0.5 seconds while PAN has a lifetime of 51 minutes (Orlando and Tyndall, 2012).

Each chemical mechanism used in our study represents  $HO_2NO_2$  and PAN, although in many reduced chemical mechanisms the PAN mechanism species represents  $CH_3C(O)O_2NO_2$ and other acyl peroxy nitrates. This representation of PAN in reduced chemical mechanisms can overestimate PAN levels compared to more detailed chemical mechanisms (Luecken et al., 1999). The near-explicit MCMv3.2 represents a range of peroxy nitrates including  $CH_3O_2NO_2$ 

Figure 5: Day-time  $\mathrm{RO_2NO_2}$  production budgets normalised by the total oxidation rate of emitted VOC in the  $\mathrm{NO_x}$ -regimes of Fig. 3. The total budgets are allocated to the most important peroxy nitrates and all other contributions included as 'Others'.



265 and about 280 acyl peroxy nitrates.

Figure 5 displays the normalised production budgets of  $RO_2NO_2$  by the total oxidation rate of the emitted VOC for each chemical mechanism in each  $NO_x$  regime with a temperature-independent and temperature-dependent source of isoprene emissions. The contribution of  $CH_3O_2NO_2$  to normalised  $RO_2NO_2$  production in MCMv3.2 is missing from the budgets of each reduced chemical mechanism as  $CH_3O_2NO_2$  is not represented in any of the reduced chemical mechanisms. In fact when removing the contribution of  $CH_3O_2NO_2$  to normalised  $RO_2NO_2$  production in MCMv3.2, the normalised  $RO_2NO_2$  production of the reduced chemical mechanisms is similar to that in the MCMv3.2 for each  $NO_x$  regime and regardless of isoprene source. Including  $CH_3O_2NO_2$  chemistry in reduced chemical mechanisms would improve the representation of the total  $RO_2NO_2$  production having the added benefit of improving the representation of  $O_x$  production budgets in Fig. 4.

Figure 6: The maximum 8 h mean ozone from the box model simulations allocated to the different  $NO_x$  regimes for each chemical mechanisms (solid lines). The box model ozone-temperature correlation is compared to the summer 2007 ERA-Interim data (black circles) and WRF-Chem output using MOZART-4 (purple boxes) and RADM2 (orange triangles).

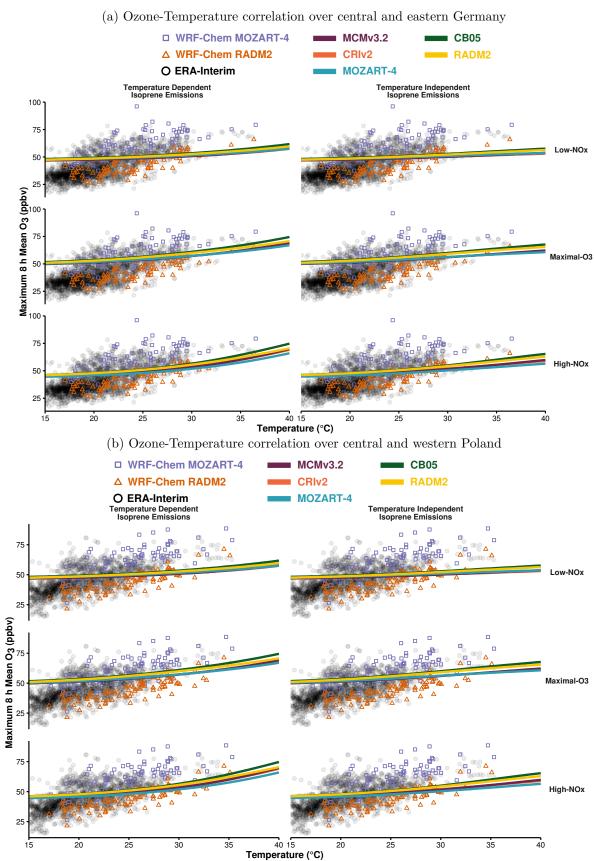


Table 3: Slopes ( $m_{O3-T}$  in ppbv per  $^{\circ}C$ ) of the linear fit to the ozone-temperature correlations in Fig. 6

(a) Slope of linear fit of the ERA-Interim observational data and WRF-Chem model output using MOZART-4 and RADM2 chemistry over central and eastern Germany and western and central Poland.

	Germany	Poland
ERA-Interim	2.15	1.94
WRF-Chem with MOZART-4	2.05	2.00
WRF-Chem with RADM2	1.78	1.77

(b) Slope of linear fit of box model experiments for each chemical mechanism, source of isoprene emissions allocated to the three  $NO_x$ -regimes.

Mechanism	Isoprene Emissions	Low-NO <sub>x</sub>	${\bf Maximal \hbox{-} O_3}$	$\operatorname{High-NO}_{\mathrm{x}}$
MCMv3.2	Temperature Dependent	0.42	0.74	0.93
	Temperature Independent	0.28	0.51	0.59
CRIv2	Temperature Dependent	0.40	0.71	0.90
	Temperature Independent	0.25	0.47	0.55
MOZART-4	Temperature Dependent	0.38	0.65	0.81
	Temperature Independent	0.25	0.44	0.49
CB05	Temperature Dependent	0.52	0.89	1.12
	Temperature Independent	0.39	0.67	0.79
RADM2	Temperature Dependent	0.48	0.79	0.97
	Temperature Independent	0.37	0.61	0.70

## 3.3 Comparison to Observations and 3D Model Simulations

277

288

289

290

This section compares the results from our idealised box model simulations to real-world 278 observations and model output from a 3D model verifying the applicability of our results 279 to more realistic atmospheric conditions. Otero et al. (2016) showed that over the summer (JJA) 280 months, temperature is the main meteorological driver of ozone production over many regions of central Europe using the observational data set of the ERA-Interim re-analysis of Schnell et al. (2015). This data set includes the daily maximum temperature and daily maximum 8 h mean of 283 ozone for the years 1998–2012 over Europe. Model output from the 3D WRF-Chem regional 284 model set-up over the European domain for simulations of the year 2007 using MOZART-4 and 285 RADM2 chemistry from Mar (2015) was used to further compare the box model simulations to a 286 model including more meteorological processes than the box model. 287

Figure 6 compares the summer data of 2007 from observations (ERA-Interim), WRF-Chem simulations and the maximum 8 h mean ozone from the box model simulations using a temperature-independent and temperature-dependent source of isoprene emissions for each

chemical mechanism and allocated to the different  $NO_x$ -regimes. In Fig. 6, only days where the 291 daily maximum temperature corresponded to the temperature range in our study (15–40 °C) 292 were considered from the ERA-Interim observational data set. We selected two regions of the 293 gridded domains of the observations and WRF-Chem output, central and eastern Germany (Fig. 6a) and central and western Poland (Fig. 6b), as the summertime ozone values in these 295 regions are correlated with temperature (Otero et al., 2016). Table 3 summarises the slopes 296  $(m_{O3-T})$  of the linear fits of all the ozone-temperature correlations displayed in Fig. 6 in ppbv of 297 ozone per °C determining the rate of change of ozone with temperature. 298

299

300

301

304

305

306

307

308

309

310

311

312

313

314

The spread of the ERA-Interim ozone-temperature values over both Germany and Poland are generally captured by the combined WRF-Chem simulations with both MOZART-4 and RADM2 chemistry. However, the ozone-temperature data from WRF-Chem using MOZART-4 chemistry reproduces the higher ozone values with temperature from ERA-Interim but not the 302 lower values. On the other hand, WRF-Chem with RADM2 only reproduces the lower ozone 303 values of the ERA-Interim data. The differences between the ozone produced using WRF-Chem with MOZART-4 and RADM2 shall be addressed in Mar (2015). The rate of change of ozone with temperature from the WRF-Chem simulations using MOZART-4 is closer to the ERA-Interim data than the WRF-Chem simulations using RADM2 (Table 3a).

The box model simulations using a temperature-independent source of isoprene emissions do not reproduce the range of observed ozone-temperature values, also indicated by the lower m<sub>O3-T</sub> values of the box model simulations than ERA-Interim in Table 3. Also, the differences in ozone production from the different chemical mechanisms with the box model are insignificant compared to the spread of the ERA-Interim and WRF-Chem data. When using a temperature-dependent source of isoprene emissions in the box model, the rate of change of ozone with temperature in the box model approaches that of the observed data.

 $High-NO_x$  conditions with a temperature-dependent source of isoprene produce the most 315 similar ozone-temperature slope to the observational data but this is still lower by a factor of two 316 than the observed ozone-temperature slope. In particular, the box model simulations over-predict 317 the ozone values at lower temperatures and under-predict the ozone values at higher temperatures 318 compared to the ERA-Interim data, regardless of the chemical mechanism. Similarly, the rate 319 of change of ozone with temperature in the box model is less-sensitive than WRF-Chem using 320 MOZART-4 or RADM2 chemistry. 321

The main reason for the box model simulations being less sensitive to temperature than the

observations is related to the set-up of the experiments compared to observations. Observations consider the total effect of temperature on ozone, while models represent the individual effects of temperature on ozone. In other words, observational studies look at the total derivative of ozone with temperature while models look at the partial derivatives of the temperature-dependent processes influencing ozone (Rasmussen et al., 2013).

$$\frac{d[\mathcal{O}_3]}{dT} = \frac{\partial[\mathcal{O}_3]}{\partial[\mathcal{B}\mathcal{V}\mathcal{O}C]} \frac{\partial[\mathcal{B}\mathcal{V}\mathcal{O}C]}{\partial T} + \frac{\partial[\mathcal{O}_3]}{\partial\mathcal{C}hemistry} \frac{\partial\mathcal{C}hemistry}{\partial T} + \frac{\partial[\mathcal{O}_3]}{\partial\mathcal{S}tagnation} \frac{\partial\mathcal{S}tagnation}{\partial T} + \dots$$

In our simulations, we focused on instantaneous production of ozone from a freshly-emitted source of VOC not considering stagnant atmospheric conditions. In these atmospheric conditions, the ozone built-up from the previous day is not transported away from the region and can lead to increased ozone levels with the production of fresh ozone from new emissions. Otero et al. (2016) showed that the previous day's ozone was also an important driver for observed ozone production over Europe. 3D models such as WRF-Chem that simulate more realistic atmospheric conditions would play a valuable role for future work that would also include stagnant conditions.

In this study, we determined the effects of temperature on ozone production using a box

## 29 4 Conclusions

330

model over a range of temperatures and  $NO_x$  conditions using a temperature-independent and 331 temperature-dependent source of isoprene emissions. These simulations were repeated using 332 reduced chemical mechanism schemes (CRIv2, MOZART-4, CB05 and RADM2) typically used 333 in 3D models and compared to the near-explicit MCMv3.2 chemical mechanism. Each chemical mechanism produced a non-linear relationship of ozone with temperature and  $\mathrm{NO}_{\mathrm{x}}$  with the 335 most ozone produced at high temperatures and high emissions of NO<sub>x</sub>. Conversely, lower NO<sub>x</sub> 336 levels led to a minimal increase of ozone with temperature. Thus air quality in a future with 337 higher temperatures predicted with climate change would benefit from dramatically reducing 338  $NO_x$  emissions. 339 Faster chemistry at higher temperatures was responsible for a greater absolute increase in 340 ozone than increased isoprene emissions. Faster thermal decomposition of peroxy nitrates at 341 higher temperatures contributed the most to ozone production with each chemical mechanism and 342 all  $NO_x$  conditions. The contribution of peroxy nitrates using reduced chemical mechanisms was 343 larger than in MCMv3.2 due to the inclusion of methylperoxy nitrate (CH<sub>3</sub>O<sub>2</sub>NO<sub>2</sub>) chemistry that is not included in any other chemical mechanism used in this study. Including methylperoxy nitrate chemistry in reduced chemical mechanisms would minimise the differences in the production of ozone from reduced chemical mechanisms to the MCMv3.2 at higher temperatures.

The rate of change of ozone with temperature using observational data (ERA-Interim) over Europe was twice as high as when using the box model. This was due to the box model not representing stagnant atmospheric conditions that are inherently included in observational data and models including meteorology, such as WRF-Chem. Future work looking at the influence of temperature on ozone should include stagnant conditions to represent more realistic atmospheric conditions. Any modelling work addressing this should also consider a range of  $NO_x$  conditions as this strongly influenced the amount of ozone produced in our study.

# References

- A. Baklanov, K. Schlünzen, P. Suppan, J. Baldasano, D. Brunner, S. Aksoyoglu, G. Carmichael,
- J. Douros, J. Flemming, R. Forkel, S. Galmarini, M. Gauss, G. Grell, M. Hirtl, S. Joffre, O. Jorba,
- E. Kaas, M. Kaasik, G. Kallos, X. Kong, U. Korsholm, A. Kurganskiy, J. Kushta, U. Lohmann,
- A. Mahura, A. Manders-Groot, A. Maurizi, N. Moussiopoulos, S. T. Rao, N. Savage, C. Seigneur,
- R. S. Sokhi, E. Solazzo, S. Solomos, B. Sørensen, G. Tsegas, E. Vignati, B. Vogel, and Y. Zhang.
- online coupled regional meteorology chemistry models in Europe: current status and prospects.
- Atmospheric Chemistry and Physics, 14(1):317–398, 2014.
- 363 T.M. Butler, M.G. Lawrence, D. Taraborrelli, and J. Lelieveld. Multi-day ozone production
- potential of volatile organic compounds calculated with a tagging approach. Atmospheric
- 365 Environment, 45(24):4082-4090, 2011.
- William P. L. Carter, Arthur M. Winer, Karen R. Darnall, and James N. Pitts Jr. Smog chamber
- studies of temperature effects in photochemical smog. Environmental Science & Technology, 13
- 368 (9):1094–1100, 1979.
- J. Coates and T. M. Butler. A comparison of chemical mechanisms using tagged ozone production
- potential (TOPP) analysis. Atmospheric Chemistry and Physics, 15(15):8795–8808, 2015.
- John P. Dawson, Peter J. Adams, and Spyros N. Pandis. Sensitivity of ozone to summertime
- climate in the eastern USA: A modeling case study. Atmospheric Environment, 41(7):1494 –
- зтз 1511, 2007.

- L. K. Emmons, S. Walters, P. G. Hess, J.-F. Lamarque, G. G. Pfister, D. Fillmore, C. Granier,
- A. Guenther, D. Kinnison, T. Laepple, J. Orlando, X. Tie, G. Tyndall, C. Wiedinmyer, S. L.
- 376 Baughcum, and S. Kloster. Description and evaluation of the Model for Ozone and Related
- chemical Tracers, version 4 (MOZART-4). Geoscientific Model Development, 3(1):43-67, 2010.
- Wendy S. Goliff, William R. Stockwell, and Charlene V. Lawson. The regional atmospheric
- chemistry mechanism, version 2. Atmospheric Environment, 68:174 185, 2013.
- A. Guenther, T. Karl, P. Harley, C. Wiedinmyer, P. I. Palmer, and C. Geron. Estimates of global
- terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from
- Nature). Atmospheric Chemistry and Physics, 6(11):3181–3210, 2006.
- A. B. Guenther, X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons, and
- 384 X. Wang. The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1):
- an extended and updated framework for modeling biogenic emissions. Geoscientific Model
- 386 Development, 5(6):1471-1492, 2012.
- Shiro Hatakeyama, Hajime Akimoto, and Nobuaki Washida. Effect of temperature on the
- formation of photochemical ozone in a propene-nitrogen oxide (NOx)-air-irradiation system.
- 389 Environmental Science & Technology, 25(11):1884–1890, 1991.
- Daniel J. Jacob and Darrell A. Winner. Effect of climate change on air quality. Atmospheric
- Environment, 43(1):51-63, 2009. Atmospheric Environment Fifty Years of Endeavour.
- M. E. Jenkin, S. M. Saunders, V. Wagner, and M. J. Pilling. Protocol for the development of the
- Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile
- organic compounds. Atmospheric Chemistry and Physics, 3(1):181–193, 2003.
- M.E. Jenkin, L.A. Watson, S.R. Utembe, and D.E. Shallcross. A Common Representative
- <sup>396</sup> Intermediates (CRI) mechanism for VOC degradation. Part 1: Gas phase mechanism development.
- 397 Atmospheric Environment, 42(31):7185 7195, 2008.
- Michael E. Jenkin, Sandra M. Saunders, and Michael J. Pilling. The tropospheric degradation of
- volatile organic compounds: a protocol for mechanism development. Atmospheric Environment,
- 31(1):81 104, 1997.
- 401 J. J. P. Kuenen, A. J. H. Visschedijk, M. Jozwicka, and H. A. C. Denier van der Gon.
- <sup>402</sup> TNO-MACC II emission inventory; a multi-year (2003–2009) consistent high-resolution european

- emission inventory for air quality modelling. Atmospheric Chemistry and Physics, 14(20):
- 404 10963–10976, 2014.
- 405 D.J. Luecken, G.S. Tonnesen, J.E. Sickles, and II. Differences in noy speciation predicted by
- three photochemical mechanisms. Atmospheric Environment, 33(7):1073 1084, 1999.
- Kathleen A. Mar. WRF-Chem Simulations over Europe: Model Evaluation and Chemical
- 408 Mechanism Comparison, 2015.
- 409 W. J. Moxim, H. Levy, and P. S. Kasibhatla. Simulated global tropospheric PAN: Its transport
- and impact on NO<sub>x</sub>. Journal of Geophysical Research: Atmospheres, 101(D7):12621–12638, 1996.
- John J. Orlando and Geoffrey S. Tyndall. Laboratory studies of organic peroxy radical chemistry:
- an overview with emphasis on recent issues of atmospheric significance. Chem. Soc. Rev., 41:
- 413 6294–6317, 2012.
- N. Otero, J. Sillmann, J. L. Schnell, H. Rust, and T. M. Butler. Synoptic and meteorological
- drivers of extreme ozone concentrations over europe. Environmental Research Letters, page In
- Preparation, 2016.
- N. Passant. Speciation of UK emissions of non-methane volatile organic compounds. Technical
- report, DEFRA, Oxon, UK., 2002.
- 419 George Pouliot, Hugo A.C. Denier van der Gon, Jeroen Kuenen, Junhua Zhang, Michael D. Moran,
- and Paul A. Makar. Analysis of the emission inventories and model-ready emission datasets of
- Europe and North America for phase 2 of the AQMEII project. Atmospheric Environment, 115:
- 422 345–360, 2015.
- S. E. Pusede, D. R. Gentner, P. J. Wooldridge, E. C. Browne, A. W. Rollins, K.-E. Min, A. R.
- Russell, J. Thomas, L. Zhang, W. H. Brune, S. B. Henry, J. P. DiGangi, F. N. Keutsch, S. A.
- 425 Harrold, J. A. Thornton, M. R. Beaver, J. M. St. Clair, P. O. Wennberg, J. Sanders, X. Ren,
- T. C. VandenBoer, M. Z. Markovic, A. Guha, R. Weber, A. H. Goldstein, and R. C. Cohen.
- On the temperature dependence of organic reactivity, nitrogen oxides, ozone production, and
- the impact of emission controls in San Joaquin Valley, California. Atmospheric Chemistry and
- Physics, 14(7):3373-3395, 2014.
- Sally E. Pusede, Allison L. Steiner, and Ronald C. Cohen. Temperature and Recent Trends in
- the Chemistry of Continental Surface Ozone. Chemical Reviews, 115(10):3898–3918, 2015.

- 432 D. J. Rasmussen, Jianlin Hu, Abdullah Mahmud, and Michael J. Kleeman. The ozone-climate
- penalty: Past, present, and future. Environmental Science & Technology, 47(24):14258–14266,
- 434 2013. PMID: 24187951.
- Andrew Rickard, Jenny Young, M. J. Pilling, M. E. Jenkin, Stephen Pascoe, and S. M. Saunders.
- The Master Chemical Mechanism Version MCM v3.2. http://mcm.leeds.ac.uk/MCMv3.2/,
- 437 2015. [Online; accessed 25-March-2015].
- <sup>438</sup> Juli I. Rubin, Andrew J. Kean, Robert A. Harley, Dylan B. Millet, and Allen H. Goldstein.
- 439 Temperature dependence of volatile organic compound evaporative emissions from motor vehicles.
- Journal of Geophysical Research: Atmospheres, 111(D3), 2006. D03305.
- 441 S. M. Saunders, M. E. Jenkin, R. G. Derwent, and M. J. Pilling. Protocol for the development of
- the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic
- volatile organic compounds. Atmospheric Chemistry and Physics, 3(1):161–180, 2003.
- J. L. Schnell, M. J. Prather, B. Josse, V. Naik, L. W. Horowitz, P. Cameron-Smith, D. Bergmann,
- 445 G. Zeng, D. A. Plummer, K. Sudo, T. Nagashima, D. T. Shindell, G. Faluvegi, and S. A. Strode.
- 446 Use of north american and european air quality networks to evaluate global chemistry–climate
- 447 modeling of surface ozone. Atmospheric Chemistry and Physics, 15(18):10581–10596, 2015.
- 448 Sanford Sillman. The use of NOy, H2O2, and HNO3 as indicators for ozone-NOx-hydrocarbon
- 449 sensitivity in urban locations. Journal of Geophysical Research: Atmospheres, 100(D7):
- 450 14175–14188, 1995.
- Sanford Sillman. The relation between ozone, NOx and hydrocarbons in urban and polluted
- rural environments. Atmospheric Environment, 33(12):1821 1845, 1999.
- 453 Sanford Sillman and Perry J. Samson. Impact of temperature on oxidant photochemistry in
- urban, polluted rural and remote environments. Journal of Geophysical Research: Atmospheres,
- 455 100(D6):11497–11508, 1995.
- D. Simpson, A. Benedictow, H. Berge, R. Bergström, L. D. Emberson, H. Fagerli, C. R. Flechard,
- 457 G. D. Hayman, M. Gauss, J. E. Jonson, M. E. Jenkin, A. Nyíri, C. Richter, V. S. Semeena,
- S. Tsyro, J.-P. Tuovinen, Á. Valdebenito, and P. Wind. The EMEP MSC-W chemical transport
- model technical description. Atmospheric Chemistry and Physics, 12(16):7825–7865, 2012.

- William R. Stockwell, Paulette Middleton, Julius S. Chang, and Xiaoyan Tang. The second
- 461 generation regional acid deposition model chemical mechanism for regional air quality modeling.
- Journal of Geophysical Research: Atmospheres, 95(D10):16343–16367, 1990.
- William R. Stockwell, Frank Kirchner, Michael Kuhn, and Stephan Seefeld. A new mechanism
- for regional atmospheric chemistry modeling. Journal of Geophysical Research: Atmospheres,
- 102(D22):25847-25879, 1997.
- 466 E. von Schneidemesser, J. Coates, A. J. H. Visschedijk, H. A. C. Denier van der Gon, and T. M.
- Butler. Variation of the NMVOC speciation in the solvent sector and the sensitivity of modelled
- tropospheric ozone. Atmospheric Environment, page In preparation, 2016.
- 469 Patrick Wagner and Wilhelm Kuttler. Biogenic and anthropogenic isoprene in the near-surface
- urban atmosphere A case study in Essen, Germany. Science of The Total Environment, 475:
- 104 115, 2014.
- Greg Yarwood, Sunja Rao, Mark Yocke, and Gary Z. Whitten. Updates to the Carbon Bond
- <sup>473</sup> Chemical Mechanism: CB05. Technical report, U. S Environmental Protection Agency, 2005.