

The Influence of Temperature on Ozone Production under varying NO_x Conditions – a modelling study

J. Coates¹, K. Mar¹ and T. Butler¹

¹Institute for Advanced Sustainability Studies, Potsdam, Germany

March 5, 2016

Abstract

Ground-level ozone is a secondary air pollutant produced during the degradation of emitted volatile organic compounds (VOCs) in the presence of sunlight and nitrogen oxides (NO_x). Temperature directly influences ozone production through speeding up the rates of the chemical processes producing ozone and increasing the emissions of VOCs, such as isoprene, from vegetation. In this study, we used a box model to reproduce the non-linear relationship of ozone on NO_x and temperature from previous observational studies. An increase in ozone of up to 20 ppbv was due to faster reaction rates while increased isoprene emissions added a further 11 ppbv of ozone under high- NO_x conditions. The increase of ozone with temperature is coupled with an increase in OH with temperature resulting in increased VOC reactivity with temperature that causes an increase in net O_x production of ~ 1 molecule of O_x per loss of VOC. The rate of increase in ozone with temperature from our box model simulations was about half the rate of increase in ozone with temperature over central Europe compared to both observed values and WRF-Chem output. The missing sensitivity in our simulations compared to observations and 3D model output is related to not including stagnant atmospheric conditions (coupling of high temperature and low wind speeds) in our experiment.

1 Introduction

Surface-level ozone (O_3) is a secondary air pollutant formed during the photochemical degradation of volatile organic compounds (VOCs) in the presence of nitrogen oxides ($\text{NO}_x \equiv \text{NO} + \text{NO}_2$). Due to the photochemical nature of ozone production, meteorological variables such as temperature

strongly influence ozone production (Jacob and Winner, 2009). Otero et al. (2016) showed that temperature was a major meteorological driver for summertime ozone in many areas of central Europe.

Temperature primarily influences ozone production in two ways: speeding up the reaction rates of many chemical reactions leading to ozone production, and increasing emissions of VOCs from biogenic sources (BVOCs). While emissions of anthropogenic VOCs (AVOCs) are generally not dependent on temperature, evaporative emissions of AVOCs increase with temperature (Rubin et al., 2006). The review of Pusede et al. (2015) provides further details of the temperature-dependent processes impacting ozone production.

Studies over the US (Sillman and Samson, 1995; Dawson et al., 2007; Pusede et al., 2014) noted that increased temperatures tend to lead to higher ozone levels, often exceeding local air quality guidelines. Some studies (Sillman and Samson, 1995; Dawson et al., 2007) included regional modelling to simulate the observed increases in ozone with temperature. In these studies, the increase of ozone with temperature was attributed to the shorter lifetime of PAN (peroxy acetyl nitrate) at higher temperatures and increased emissions of BVOCs, in particular isoprene, from vegetation.

Pusede et al. (2014) used an analytical model constrained by observations over San Joaquin Valley, California to infer a non-linear relationship of ozone production with temperature and NO_x , similar to the well-known non-linear relationship of ozone production on NO_x and VOC levels (Sillman, 1999). Moreover, Pusede et al. (2014) showed that temperature can be used as a surrogate for VOC levels when considering the relationship of ozone across NO_x gradients.

Environmental chamber studies have also been used to analyse the relationship of ozone with temperature. The chamber experiments of Carter et al. (1979) and Hatakeyama et al. (1991) showed increases in ozone from a VOC mix with temperature linked to increased PAN decomposition at temperatures greater than 303 K.

Despite many studies considering the effects of temperature on ozone production from an observational and chamber study perspective, there are no modelling studies (to our knowledge) focusing on the detailed chemical processes of the influence of temperature on ozone production under different NO_x conditions. Regional modelling studies (using a single chemical mechanism) have concentrated on reproducing ozone levels over regions with known meteorology and NO_x conditions then only varying the temperature. These modelling studies did not consider the relationship of ozone with NO_x and temperature. The review of Pusede et al. (2015) also highlights

a lack of modelling studies looking at the non-linear relationship of ozone on temperature under different NO_x conditions.

Comparisons of different chemical mechanisms, such as Emmerson and Evans (2009) and Coates and Butler (2015), showed that different representations of tropospheric chemistry influenced ozone production. These studies did not consider whether the ozone-temperature relationship differed between chemical mechanisms. The study of Rasmussen et al. (2013) also noted that changing the chemical mechanism used by a model may also change the simulated ozone-temperature relationship. Comparing the ozone-temperature relationship predicted by different chemical mechanisms is important for modelling of future air quality due to the expected increase in heatwaves.

In this study, we use an idealised box model to determine how ozone levels vary with temperature under different NO_x conditions. We determine whether faster chemical reaction rates or increased BVOC emissions have a greater influence on instantaneous ozone production with higher temperature under different NO_x conditions. Furthermore, we compare the ozone-temperature relationship produced by different chemical mechanisms by repeating all simulations with various chemical mechanisms.

2 Methodology

2.1 Model Setup

We used the MECCA box model to determine the important gas-phase chemical processes for ozone production under different temperatures and NO_x conditions. The MECCA box model was set up as described in Coates and Butler (2015) and updated to include vertical mixing with the free troposphere using a diurnal cycle for the PBL height. The supplementary material includes further details of these updates.

Simulations were performed to broadly simulate urban conditions representative of central Europe with equinoctical conditions. The simulations started at 06:00 with a total run time of two days. Methane was fixed at 1.7 ppmv throughout the model run, carbon monoxide (CO) and ozone were initialised at 200 ppbv and 40 ppbv and then allowed to evolve freely throughout the simulation. All VOC emissions were held constant until noon of first day simulating a plume of freshly-emitted VOC.

Separate box model simulations were performed by systematically varying the temperature

between 288 and 313 K (15 – 40 °C). The only source of NO_x emissions in the box model was a constant source of NO emissions. Box model runs were performed with the NO emissions systematically varied between 5.0×10^9 and 1.5×10^{12} molecules (NO) cm⁻² s⁻¹ at each temperature used in this study. At 20 °C, these NO emissions corresponded to peak NO_x mixing ratios of 0.02 ppbv and 10 ppbv respectively, this range of NO_x mixing ratios covers the NO_x conditions in pristine and urban conditions (von Schneidemesser et al., 2015).

All simulations were repeated using different chemical mechanisms to investigate whether the relationship of ozone with temperature across NO_x gradients changes using different representations of ozone production chemistry. The reference chemical mechanism was the near-explicit Master Chemical Mechanism, MCMv3.2, (Jenkin et al., 1997, 2003; Saunders et al., 2003; Rickard et al., 2015). The reduced chemical mechanisms in our study were Common Representative Intermediates, CRIV2 (Jenkin et al., 2008), Model for ozone and related chemical tracers, MOZART-4 (Emmons et al., 2010), Regional Acid Deposition Model, RADM2 (Stockwell et al., 1990) and the Carbon Bond Mechanism, CB05 (Yarwood et al., 2005). Coates and Butler (2015) described these chemical mechanisms and the implementation of these chemical mechanisms in MECCA. These reduced chemical mechanisms were chosen as they are commonly used by modelling groups in 3D regional and global models (Baklanov et al., 2014).

Model runs were repeated using a temperature-dependent and temperature-independent source of BVOC emissions to determine whether increased emissions of BVOC or faster reaction rates of chemical processes is more important for the increase of ozone with temperature. MEGAN2.1 (Guenther et al., 2012) specified the temperature-dependent BVOC emissions of isoprene, Sect. 2.3 provides further details. As isoprene emissions are the most important on the global scale, we considered only isoprene emissions from vegetation (Guenther et al., 2006). Only isoprene emissions were dependent on temperature, all other emissions were constant in all simulations. In reality, many other BVOC are emitted from varying vegetation types (Guenther et al., 2006) and increased temperature can also increase BVOC emissions through increased evaporation (Rubin et al., 2006).

2.2 VOC Emissions

Emissions of urban BVOC over central Europe were taken from TNO-MACC_III emission inventory for the Benelux (Belgium, Netherlands and Luxembourg) region for the year 2011. TNO-MACC_III is the updated TNO-MACC_II emission inventory created using the same

Table 1: Total AVOC emissions in 2011 in tonnes from each anthropogenic source category assigned from TNO-MACC_III emission inventory and temperature-independent BVOC emissions in tonnes from Benelux region assigned from EMEP. The allocation of these emissions to MCMv3.2, CRIV2, CB05, MOZART-4 and RADM2 species are found in the supplementary material.

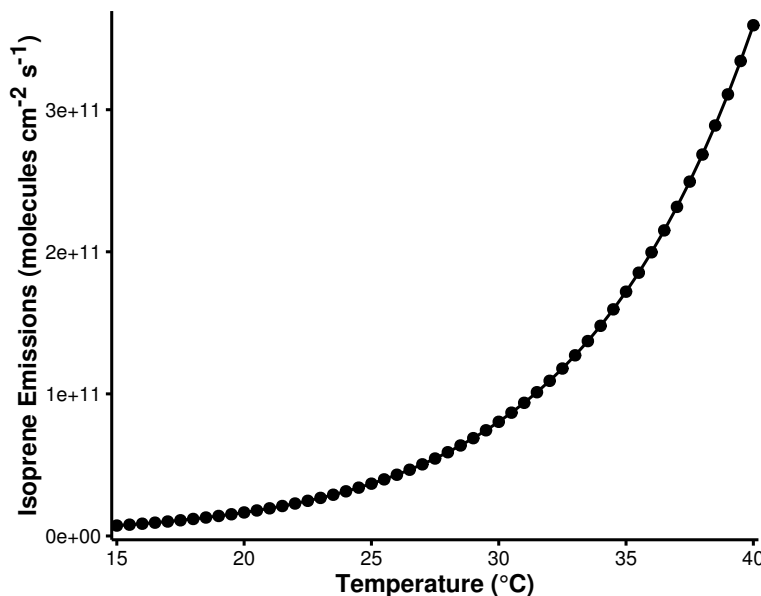
	Public Power	Residential Combustion	Industry	Fossil Fuel
Belgium	4494	9034	22152	5448
Netherlands	9140	12173	29177	8723
Luxembourg	121	44	208	1371
Total	13755	21251	62648	15542
	Solvent Use	Road Transport: Gasoline	Road Transport: Diesel	Road Transport: Others
Belgium	42809	6592	2446	144
Netherlands	53535	16589	3230	1283
Luxembourg	4482	1740	1051	6
Total	100826	24921	6727	1433
	Road Transport: Evaporation	Non-road Transport	Waste	BVOC
Belgium	210	6448	821	7042
Netherlands	1793	10067	521	1462
Luxembourg	324	643	0	2198
Total	2327	17158	1342	10702

methodology as Kuenen et al. (2014) and based upon improvements to the existing emission inventory during AQMEII-2 (Pouliot et al., 2015).

Temperature-independent emissions of isoprene and monoterpenes from biogenic sources were calculated as a fraction of the total AVOC emissions from each country in the Benelux region. This data was obtained from the supplementary data available from the EMEP (European Monitoring and Evaluation Programme) model (Simpson et al., 2012). Temperature-dependent emissions of isoprene are detailed in Sect. 2.3.

Table 1 shows the quantity of VOC emissions from each source category and the temperature-independent BVOC emissions. These categorised AVOC emissions were assigned to chemical species and groups based on the country specific profiles for Belgium, the Netherlands and Luxembourg provided by TNO. Most individual chemical species are represented by the MCMv3.2, otherwise the individual contributions of a group of VOC were further split into individual components using the detailed speciation of Passant (2002). For example, ‘xylenes’ are one of the component chemical groups in many source categories but the MCMv3.2 treats xylenes as the individual isomers (m-, o-, p-xylene) and the contributions of the individual

Figure 1: The estimated isoprene emissions (molecules isoprene $\text{cm}^{-2} \text{s}^{-1}$) using MEGAN2.1 at each temperature used in the study.



isomers to a source category was provided by Passant (2002). This approach was also used in von Schneidmesser et al. (2016) to allocate AVOC emissions from different solvent sector speciations to MCMv3.2 species.

For simulations done with other chemical mechanisms, the VOC emissions represented by the MCMv3.2 were mapped to the mechanism species representing VOC emissions in each reduced chemical mechanism based on the recommendations of the source literature and Carter (2015). The VOC emissions in the reduced chemical mechanisms were weighted by the carbon numbers of the MCMv3.2 species and the emitted mechanism species, thus keeping the amount of emitted carbon constant between simulations. The supplementary data outlines the primary VOC and calculated emissions with each chemical mechanism.

2.3 Temperature Dependent Isoprene Emissions

Temperature-dependent emissions of isoprene were estimated using the MEGAN2.1 algorithm for calculating the emissions of VOC from vegetation (Guenther et al., 2012). Emissions from nature are dependent on many variables including temperature, radiation and age but for the purpose of our study all variables except temperature were held constant.

The MEGAN2.1 parameters were chosen to give similar isoprene mixing ratios at 20 °C to the temperature-independent emissions of isoprene in order to compare the effects of increased isoprene emissions with temperature. The estimated emissions of isoprene with MEGAN2.1

Table 2: Increase in ozone mixing ratio (ppbv) due to chemistry and emissions at 40 °C from reference temperature (20 °C) in the NO_x-regimes of Fig. 3.

Chemical Mechanism	Source of Difference	Increase in Ozone at 40 °C from 20 °C (ppbv)		
		Low-NO _x	Maximal-O ₃	High-NO _x
MCMv3.2	Emissions	4.6	7.7	10.6
	Chemistry	6.8	12.5	15.2
CRIV2	Emissions	4.8	7.9	10.8
	Chemistry	6.0	11.1	13.7
MOZART-4	Emissions	4.1	6.7	10.0
	Chemistry	6.0	10.2	12.3
CB05	Emissions	4.6	7.4	9.8
	Chemistry	9.3	16.0	19.9
RADM2	Emissions	3.8	5.7	7.8
	Chemistry	8.6	14.1	17.3

using these assumptions are illustrated in Fig. 1 and show the expected exponential increase in isoprene emissions with temperature (Guenther et al., 2006).

The estimated emissions of isoprene at 20 °C lead to 0.07 ppbv of isoprene in our simulations while at 30 °C, the increased emissions of isoprene using MEGAN2.1 estimations lead to 0.35 ppbv of isoprene in the model. A measurement campaign over Essen, Germany (Wagner and Kuttler, 2014) measured 0.1 ppbv of isoprene at temperature 20 °C and 0.3 ppbv of isoprene were measured at 30 °C. The similarity of the simulated and observed isoprene mixing ratios indicates that the MEGAN2.1 variables chosen for calculating the temperature-dependent emissions of isoprene were suitable for simulating urban conditions over central Europe.

3 Results and Discussion

3.1 Ozone as a Function of NO_x and Temperature

Figure 2 depicts the maximum mixing ratio of ozone as a function of the total NO_x emissions on the first day of simulations and temperature when using a temperature-independent and temperature-dependent source of isoprene emissions for each chemical mechanism. A non-linear relationship of ozone mixing ratios with NO_x and temperature is reproduced by each chemical mechanism. This non-linear relationship is similar to that determined by Pusede et al. (2014) using an analytical model constrained to observational measurements over the San Joaquin Valley in California.

Higher ozone mixing ratios are produced when using a temperature-dependent source of isoprene emissions (Fig. 2). The highest mixing ratios of ozone are produced at high temperatures

Figure 2: Contours of maximum ozone mixing ratios (ppbv) as a function of the total NO_x emissions on the first day and temperature for each chemical mechanism using a temperature-dependent and temperature-independent source of isoprene emissions.

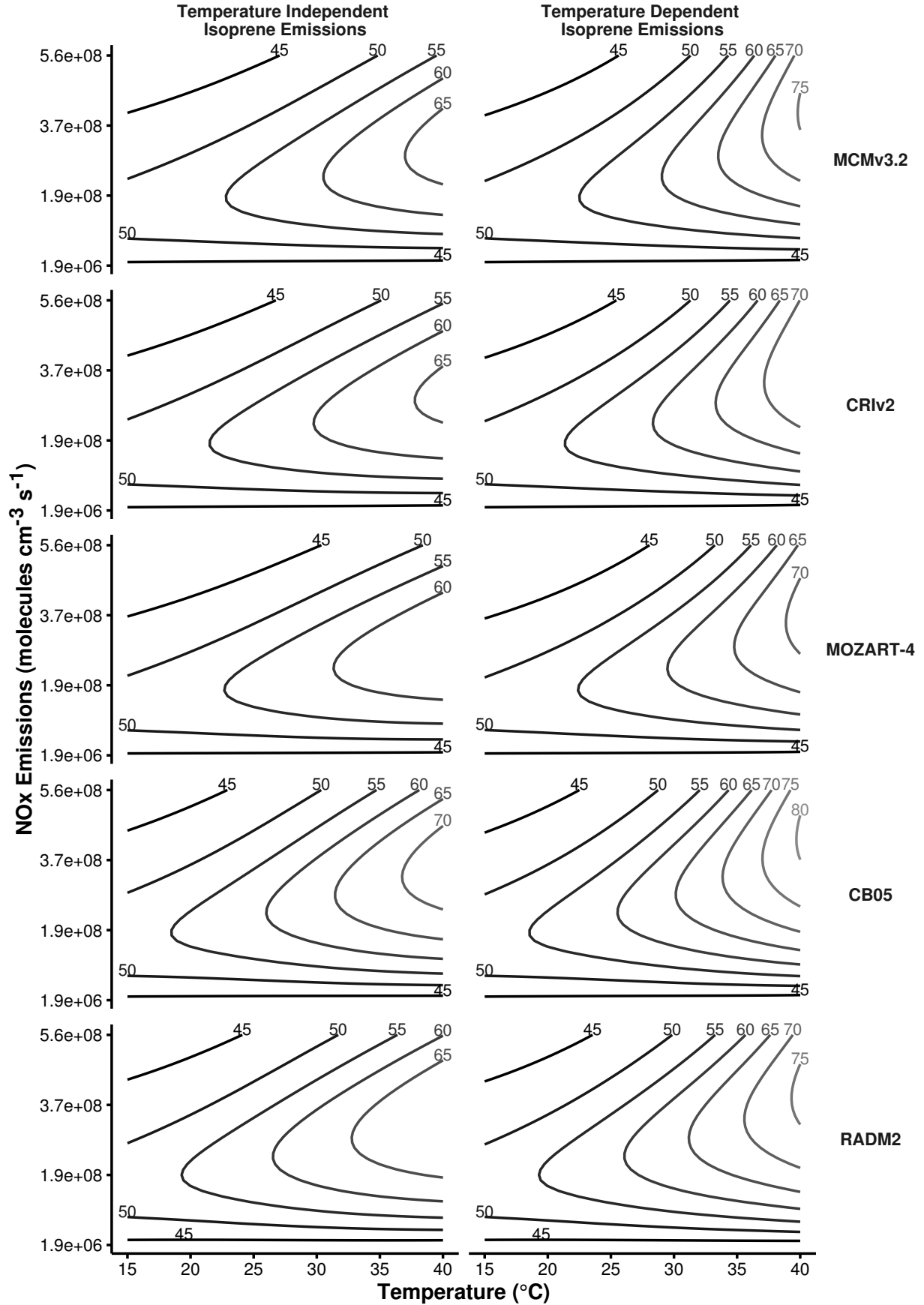
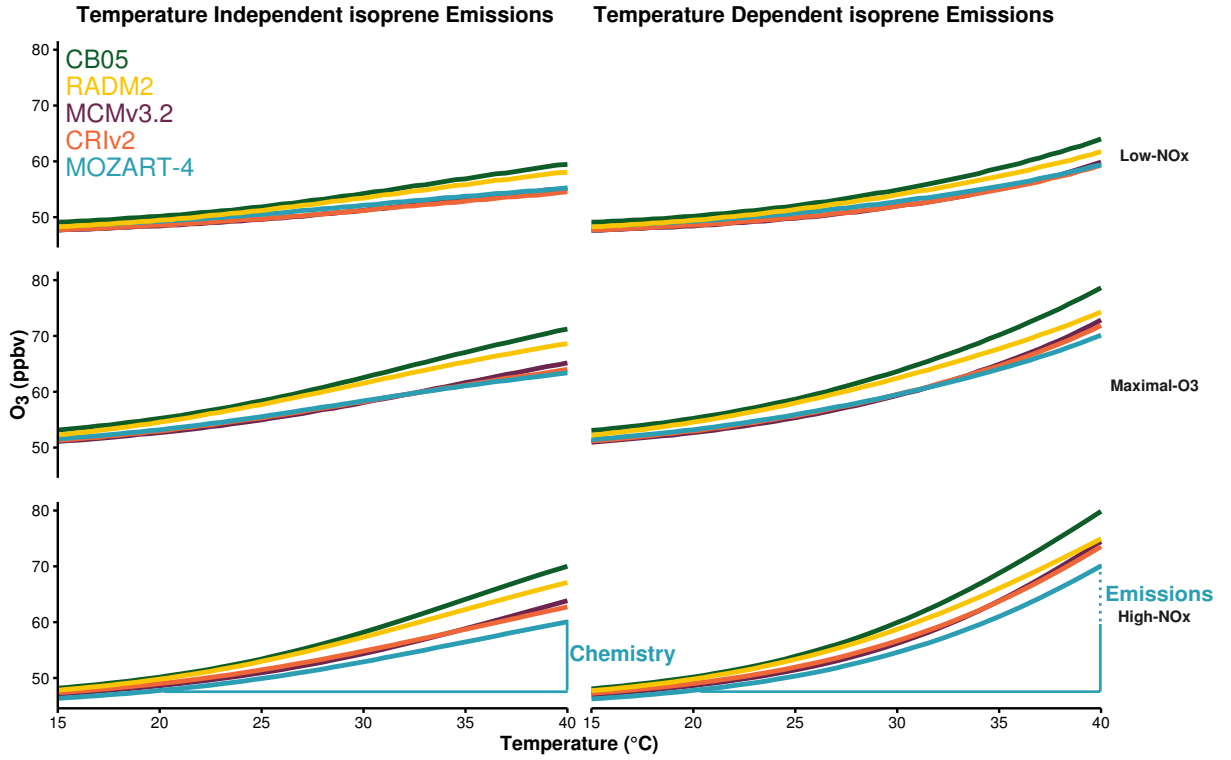


Figure 3: Ozone mixing ratios (ppbv) at each temperature are allocated to different NO_x -regimes of Fig. 2. The differences in ozone mixing ratios due to chemistry (solid line) and emissions (dotted line) are represented graphically for MOZART-4 with High- NO_x conditions and summarised in Table 2, the approach was used to calculate the differences with each chemical mechanism.



and moderate emissions of NO_x regardless of the source of isoprene emissions. Conversely, the least amount of ozone is produced with low emissions of NO_x over the whole temperature range (15 – 40 °C) when using both a temperature-independent and temperature-dependent source of isoprene emissions.

The contours of ozone mixing ratios as a function of NO_x and temperature can be split into three NO_x regimes (Low- NO_x , Maximal- O_3 and High- NO_x), similar to the NO_x regimes defined for the non-linear relationship of ozone with VOC and NO_x . The Low- NO_x regime corresponds with regions with little increase in ozone with temperature, also called the NO_x -sensitive regime. The High- NO_x (or NO_x -saturated) regime is when ozone levels increase rapidly with temperature and the contour ridges correspond to regions of maximal ozone production. This is the Maximal- O_3 regime. Pusede et al. (2014) showed that temperature can be used as a proxy for VOC, thus we assigned the ozone mixing ratios from each box model simulation to a NO_x regime based on the ratio of HNO_3 to H_2O_2 . This ratio was used by Sillman (1995) to designate ozone to NO_x regimes based on NO_x and VOC levels.

Fig. 3 illustrates the mean ozone mixing ratio at each temperature in the NO_x regimes for each chemical mechanism and each type of isoprene emissions (temperature independent and

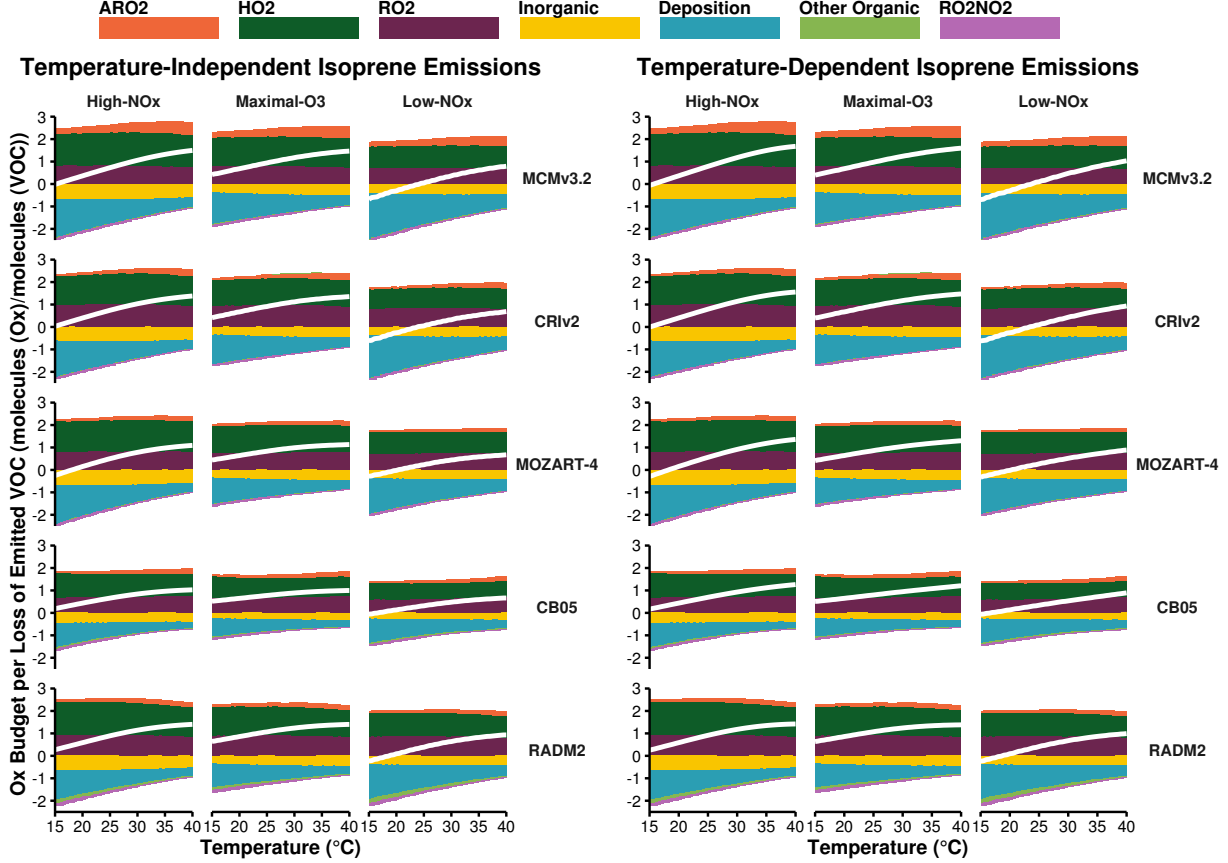
temperature dependent). We define the absolute increase in ozone at 40 °C from 20 °C due to faster reaction rates as the difference between ozone mixing ratios at 40 °C and 20 °C when using a temperature-independent source of isoprene emissions. When using a temperature-dependent source of isoprene emissions, the difference in ozone mixing ratios at 40 °C from 20 °C minus the increase due to faster chemistry, gives the absolute increase in ozone due to increased isoprene emissions. These differences are represented graphically in Fig. 3 and summarised in Table 2.

Table 2 shows that the absolute increase in ozone with temperature due to chemistry (i.e. faster reaction rates) is larger than the absolute increase in ozone due to increased isoprene emissions for each chemical mechanism and each NO_x regime. In all cases the absolute increase in ozone is largest under High-NO_x conditions and lowest with Low-NO_x conditions (Fig. 3 and Table 2). The increase in ozone mixing ratio at 40 °C from 20 °C due to faster reaction rates with High-NO_x conditions is almost double that under Low-NO_x conditions.

In the Low-NO_x regime, the increase of ozone with temperature using the reduced chemical mechanisms (CRIv2, MOZART-4, CB05 and RADM2) is similar to that from the MCMv3.2, larger differences occur in the Maximal-O₃ and High-NO_x regimes. All reduced chemical mechanisms except RADM2 have similar increases in ozone due to increased isoprene emissions to MCMv3.2 (Table 2). RADM2 produces 3 ppbv less ozone than the MCMv3.2 due to increased isoprene emissions in each NO_x regime, indicating that this difference is due the representation of isoprene degradation chemistry in RADM2.

The Tagged Ozone Production Potential (TOPP) defined in Butler et al. (2011) is a measure of the number of molecules of ozone produced per molecule of VOC emitted. Coates and Butler (2015) compared ozone production in different chemical mechanisms to the MCMv3.2 using TOPPs and showed that less ozone is produced per molecule of isoprene emitted using RADM2 than with MCMv3.2. The degradation of isoprene has been extensively studied and it is well-known that methyl vinyl ketone (MVK) and methacrolein are signatures of isoprene degradation (Atkinson, 2000). All chemical mechanisms in our study except RADM2 explicitly represent MVK and methacrolein (or in the case of CB05, a lumped species representing both these secondary degradation products). RADM2 does not represent methacrolein and the mechanism species representing ketones (KET) is a mixture of acetone and methyl ethyl ketone (MEK) (Stockwell et al., 1990). Thus the secondary degradation of isoprene in RADM2 is unable to represent the ozone production from the further degradation of the signature secondary degradation products of isoprene, MVK and methacrolein. Updated versions of RADM2, RACM

Figure 4: Day-time budgets of O_x normalised by the total oxidation rate of emitted VOC in the NO_x -regimes of Fig. 3. The white line indicates net production or consumption of O_x . The net contribution of reactions to O_x budgets are allocated to categories of deposition, inorganic reactions, peroxy nitrates (RO2NO2), reactions of NO with HO2, alkyl peroxy radicals (RO2) and acyl peroxy radicals (ARO2). All other reactions contributing to O_x budgets are allocated to the ‘Other Organic’ category.



(Stockwell et al., 1997) and RACM2 (Goliff et al., 2013), sequentially included methacrolein and MVK and with these updates the TOPP value of isoprene approached that of the MCMv3.2 (Coates and Butler, 2015).

3.2 Ozone Production Budgets

The day-time production and consumption budgets of O_x ($\equiv O_3 + NO_2 + O + O(^1D)$) normalised by the total rate of oxidation of the emitted VOC are displayed in Fig. 4. The O_x budgets are assigned to each NO_x regime for each chemical mechanism and type of isoprene emissions. The budgets are allocated to the net contribution of major sources, where ‘HO2’, ‘RO2’, ‘ARO2’ represent the reactions of NO with HO₂, alkyl peroxy radicals and acyl peroxy radicals respectively. ‘RO2NO2’ represents the effects of peroxy nitrates, ‘Deposition’ represents ozone deposition, ‘Inorganic’ is all other inorganic contributions to O_x production and any other remaining organic reactions are included in the ‘Other Organic’ category. Figure 4 also illustrates the net production

or consumption of O_x in each case.

The net O_x production efficiency increases from 20 °C to 40 °C by ~ 1 molecule of O_x per molecule of VOC oxidised using both temperature-dependent and temperature-independent isoprene emissions and in each NO_x condition (Fig. 4). The increase in net O_x production efficiency is due to a decrease in the consumption efficiency of O_x with temperature while the production efficiency of O_x remains constant with temperature (~ 2 molecules of O_x per molecule of VOC). The decrease in ozone deposition per VOC loss from 20 °C to 40 °C of ~ 1 molecule of O_x per loss of VOC mirrors the increase in net O_x production efficiency.

As the production efficiency of O_x remains constant with temperature in Fig. 4, the rate of O_x production is controlled by the loss rate of the VOCs. Faster oxidation of VOCs with temperature speeds up the production of peroxy radicals which increases ozone production when peroxy radicals react with NO to produce NO_2 . The review of Pusede et al. (2015) also acknowledged the importance of organic reactivity and radical production to the ozone-temperature relationship. Also, the modelling study of Steiner et al. (2006) noted that the increase in initial oxidation rates of VOCs with temperature leads to increased formaldehyde concentrations and in turn an increase of ozone as formaldehyde is an important source of HO_2 radicals.

Increased VOC reactivity with temperature is tied to faster reaction rates and increased levels of OH with temperature. An increase of ozone with temperature goes hand in hand with an increase of OH with temperature since ozone photolysis is typically the dominant source of OH radicals in the atmosphere.



Furthermore, enhanced formaldehyde production from the faster oxidation of VOCs increases HO_2 formation which speeds up the reaction rate of the reaction of HO_2 and NO.



(R3) is responsible for both ozone production (due to NO_2 production) and OH recycling further illustrating the strong coupling of ozone and OH production.

The net effect of peroxy nitrates on O_x budgets in our study is negligible, contributing to a loss of ~ 0.1 molecules of O_x per VOC oxidised at 40 °C from 20 °C. Peroxynitrates are produced

Table 3: Slopes (m_{O_3-T} in ppbv per $^{\circ}C$) of the linear fit to the ozone-temperature correlations in Fig. 5

(a) Slope of linear fit of the ERA-Interim observational data and WRF-Chem model output chemistry over central and eastern Germany and western and central Poland.

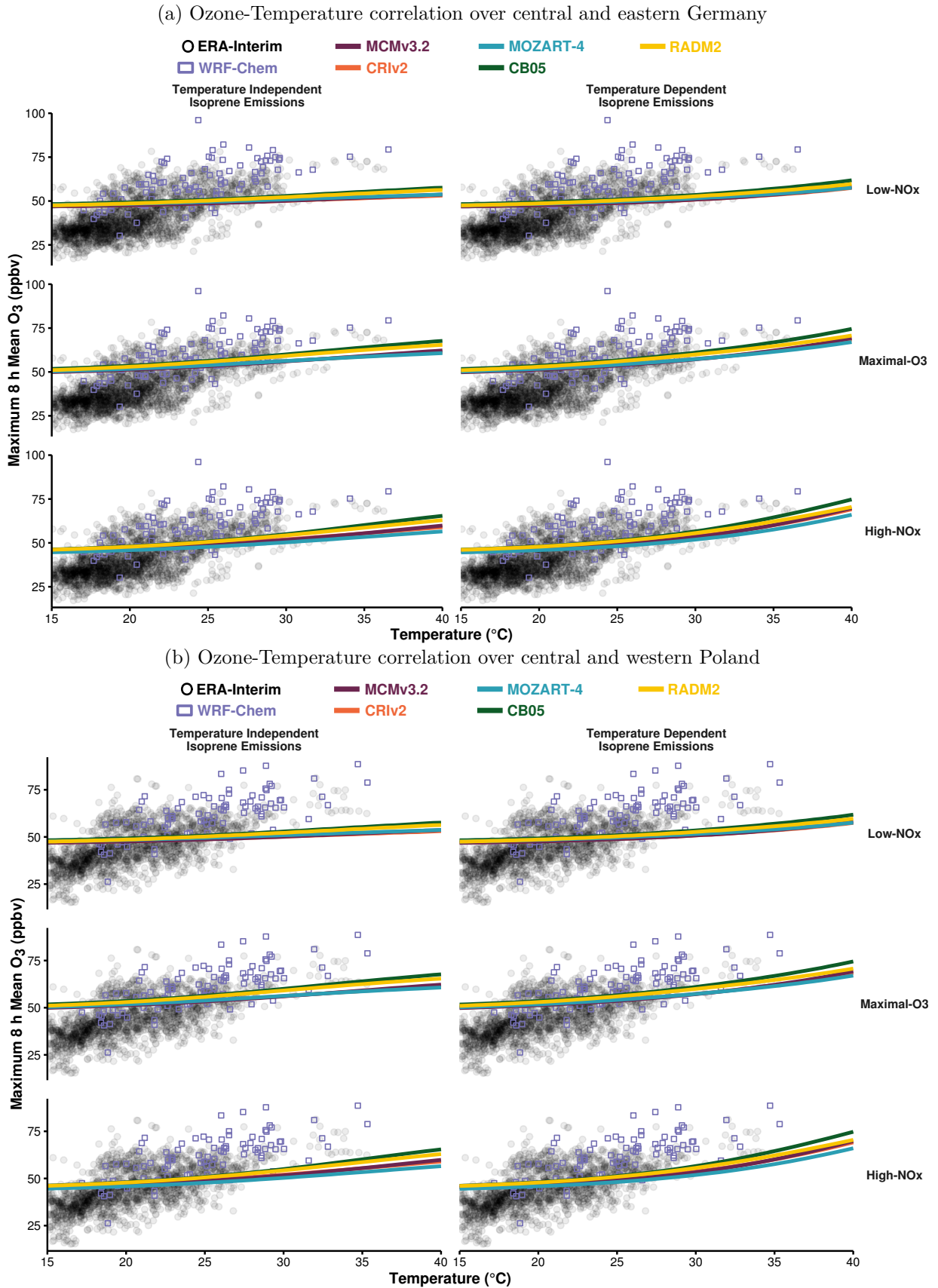
	Germany	Poland
ERA-Interim	2.15	1.94
WRF-Chem with MOZART-4	2.05	2.00

(b) Slope of linear fit of box model experiments for each chemical mechanism, source of isoprene emissions allocated to the three NO_x -regimes.

Mechanism	Isoprene Emissions	Low- NO_x	Maximal- O_3	High- NO_x
MCMv3.2	Temperature Independent	0.28	0.51	0.59
	Temperature Dependent	0.42	0.74	0.93
CRIV2	Temperature Independent	0.25	0.47	0.55
	Temperature Dependent	0.40	0.71	0.90
MOZART-4	Temperature Independent	0.25	0.44	0.49
	Temperature Dependent	0.38	0.65	0.81
CB05	Temperature Independent	0.39	0.67	0.79
	Temperature Dependent	0.52	0.89	1.12
RADM2	Temperature Independent	0.37	0.61	0.70
	Temperature Dependent	0.48	0.79	0.97

from the reactions of acyl peroxy radicals with NO_2 and are an important reservoir of peroxy radicals and NO_x . The decomposition rate of peroxy nitrates is strongly temperature dependent and at higher temperatures the faster decomposition rate of RO_2NO_2 leads to faster release of peroxy radicals and NO_x influencing ozone production. The lack of influence of RO_2NO_2 on O_x budgets in Fig. 4 is surprising as many other studies cite the decrease in the lifetime of peroxy nitrates as one of the most important factors for the increase of ozone with temperature. For example, Dawson et al. (2007) attributed the increase in ozone with temperature during a modelling study over the eastern US to the decrease in PAN lifetime with temperature. Steiner et al. (2006) also recognised that the decrease in PAN lifetime with temperature may contribute to the increase of ozone with temperature and concluded that the combined effects of increased VOC reactivity and PAN decomposition are most likely to increase the production of ozone with temperature.

Figure 5: The maximum 8 h mean ozone from the box model simulations allocated to the different NO_x regimes for each chemical mechanisms (solid lines). The box model ozone-temperature correlation is compared to the summer 2007 ERA-Interim data (black circles) and WRF-Chem output (purple boxes).



3.3 Comparison to Observations and 3D Model Simulations

This section compares the results from our idealised box model simulations to real-world observations and model output from a 3D model. Otero et al. (2016) showed that over the summer (JJA) months, temperature is the main meteorological driver of ozone production over many regions of central Europe using the observational data set of the ERA-Interim re-analysis. This data set includes the daily maximum temperature and daily maximum 8 h mean of ozone for the years 1998–2012 over Europe. Model output from the 3D WRF-Chem regional model using MOZART-4 chemistry set-up over the European domain for simulations of the year 2007 from Mar et al. (2016) was used to further compare the box model simulations to a model including more meteorological processes than our box model.

Figure 5 compares the observational (ERA-Interim) and WRF-Chem data from summer 2007 to the maximum 8 h mean ozone from the box model simulations for each chemical mechanism, NO_x regime and type of isoprene emissions. In Fig. 5, we limited the observational data to days where the observed daily maximum temperature corresponded to the temperature range in our study (15–40 °C). We selected two regions from the observations and WRF-Chem output, central and eastern Germany (Fig. 5a) and central and western Poland (Fig. 5b), where the summertime ozone values are driven by with temperature (Otero et al., 2016). Table 3 summarises the slopes ($m_{\text{O}_3\text{-T}}$) of the linear fits of all the ozone-temperature correlations displayed in Fig. 5 in ppbv of ozone per °C determining the rate of change of ozone with temperature.

The spread of the observed ozone-temperature values over both Germany and Poland are generally captured by the WRF-Chem simulation. However, WRF-Chem only reproduces the higher observed values of ozone with temperature and not the lower values. The rate of change of ozone with temperature from the WRF-Chem simulations is similar to the rate of change of ozone with temperature from the observed data (Table 3a).

The differences in ozone production between the different chemical mechanisms with the box model are small compared to the spread of the observational and WRF-Chem data. When using a temperature-dependent source of isoprene emissions in the box model, the rate of change of ozone with temperature in the box model approaches that of the observed data, although still less than half that of the observations. The box model simulations using a temperature-independent source of isoprene emissions do not reproduce the range of observed ozone-temperature values (Table 3).

A temperature-dependent source of isoprene with high- NO_x conditions produces the most

similar ozone-temperature slope to the observational data but this is still lower than the observed ozone-temperature slope by a factor of two. In particular, the box model simulations over-predict the ozone values at lower temperatures and under-predict the ozone values at higher temperatures compared to the observed data. Similarly, the rate of change of ozone with temperature in the box model is less-sensitive than WRF-Chem.

The main reason for the box model simulations being less sensitive to temperature than the observations and WRF-Chem simulation is related to the set-up of the box model. In our simulations, we focused on instantaneous production of ozone from a freshly-emitted source of VOC and do not consider stagnant atmospheric conditions characterised by high temperatures and low wind speeds slowing the transport of ozone and its precursors away from sources. Otero et al. (2016) showed that the previous day’s ozone was also an important driver for observed ozone production over Europe and Jacob et al. (1993) correlated high-ozone episodes in the summer over eastern US to regional stagnation.

Stagnation is an example of how the coupling of meteorological variables (in this case temperature and wind speed) impact ozone production. In observational studies, which consider the total derivative of ozone with temperature, the direct effects of temperature (e.g. increasing reaction rates, emissions from vegetation) and indirect (e.g. heatwaves characterised by high temperatures and low wind speed) effects of temperature on ozone are not easily separated. In other words, observational studies represent the total derivative of ozone with temperature while models consider the partial derivatives of the temperature-dependent processes influencing ozone (Rasmussen et al., 2013).

$$\frac{d[\text{O}_3]}{dT} = \frac{\partial[\text{O}_3]}{\partial[\text{BVOC}]} \frac{\partial[\text{BVOC}]}{\partial T} + \frac{\partial[\text{O}_3]}{\partial \text{Chemistry}} \frac{\partial \text{Chemistry}}{\partial T} + \frac{\partial[\text{O}_3]}{\partial \text{Stagnation}} \frac{\partial \text{Stagnation}}{\partial T} + \dots$$

3D models such as WRF-Chem that can simulate more realistic atmospheric conditions would play a valuable role for future work evaluating the ozone-temperature relationship at different NO_x conditions at a regional scale.

4 Conclusions

In this study, we determined the effects of temperature on ozone production using a box model over a range of temperatures and NO_x conditions with a temperature-independent and temperature-dependent source of isoprene emissions. These simulations were repeated using

reduced chemical mechanism schemes (CRIV2, MOZART-4, CB05 and RADM2) typically used in 3D models and compared to the near-explicit MCMv3.2 chemical mechanism.

Each chemical mechanism produced a non-linear relationship of ozone with temperature and NO_x with the most ozone produced at high temperatures and moderate emissions of NO_x . Conversely, lower NO_x levels led to a minimal increase of ozone with at all temperatures. Thus air quality in a future with higher temperatures predicted with climate change would benefit from dramatical reductions in NO_x emissions.

Faster reaction rates at higher temperatures were responsible for a greater absolute increase in ozone than increased isoprene emissions. The increase in VOC reactivity with temperature was the dominant chemical process increasing ozone production with temperature. The increase of ozone with temperature is coupled to increasing OH with temperature increasing VOC reactivity. Enhanced VOC reactivity at higher temperature leads to increased production of peroxy radicals, leading to ozone production, and recycling of OH.

The rate of change of ozone with temperature using observational data (ERA-Interim) over Europe was twice as high as when using the box model. This was consistent with the box model not representing stagnant atmospheric conditions that are inherently included in observational data and models including meteorology, such as WRF-Chem. Future work looking at the influence of temperature on ozone should include stagnant conditions to represent more realistic atmospheric conditions. Any modelling work addressing this should also consider a range of NO_x conditions as this strongly influenced the amount of ozone produced in our study.

Acknowledgements

The authors would like to thank Noelia Otero Felipe for providing the ERA-Interim data.

References

- Roger Atkinson. Atmospheric chemistry of VOCs and NO_x . *Atmospheric Environment*, 34(12-14): 2063–2101, 2000.
- A. Baklanov, K. Schlünzen, P. Suppan, J. Baldasano, D. Brunner, S. Aksoyoglu, G. Carmichael, J. Douros, J. Flemming, R. Forkel, S. Galmarini, M. Gauss, G. Grell, M. Hirtl, S. Joffre, O. Jorba, E. Kaas, M. Kaasik, G. Kallos, X. Kong, U. Korsholm, A. Kurganskiy, J. Kushta, U. Lohmann, A. Mahura, A. Manders-Groot, A. Maurizi, N. Moussiopoulos, S. T. Rao, N. Savage, C. Seigneur,

352 R. S. Sokhi, E. Solazzo, S. Solomos, B. Sørensen, G. Tsegas, E. Vignati, B. Vogel, and Y. Zhang.
 353 Online coupled regional meteorology chemistry models in Europe: current status and prospects.
 354 *Atmospheric Chemistry and Physics*, 14(1):317–398, 2014.

355 T.M. Butler, M.G. Lawrence, D. Taraborrelli, and J. Lelieveld. Multi-day ozone production
 356 potential of volatile organic compounds calculated with a tagging approach. *Atmospheric*
 357 *Environment*, 45(24):4082 – 4090, 2011.

358 William P. L. Carter. Development of a Database for Chemical Mechanism Assignments for
 359 Volatile Organic Emissions. *Journal of the Air & Waste Management Association*, 0, 2015.

360 William P. L. Carter, Arthur M. Winer, Karen R. Darnall, and James N. Pitts Jr. Smog chamber
 361 studies of temperature effects in photochemical smog. *Environmental Science & Technology*, 13
 362 (9):1094–1100, 1979.

363 J. Coates and T. M. Butler. A comparison of chemical mechanisms using tagged ozone production
 364 potential (TOPP) analysis. *Atmospheric Chemistry and Physics*, 15(15):8795–8808, 2015.

365 John P. Dawson, Peter J. Adams, and Spyros N. Pandis. Sensitivity of ozone to summertime
 366 climate in the eastern USA: A modeling case study . *Atmospheric Environment*, 41(7):1494 –
 367 1511, 2007.

368 K. M. Emmerson and M. J. Evans. Comparison of tropospheric gas-phase chemistry schemes for
 369 use within global models. *Atmospheric Chemistry and Physics*, 9(5):1831–1845, 2009.

370 L. K. Emmons, S. Walters, P. G. Hess, J.-F. Lamarque, G. G. Pfister, D. Fillmore, C. Granier,
 371 A. Guenther, D. Kinnison, T. Laepple, J. Orlando, X. Tie, G. Tyndall, C. Wiedinmyer, S. L.
 372 Baughcum, and S. Kloster. Description and evaluation of the Model for Ozone and Related
 373 chemical Tracers, version 4 (MOZART-4). *Geoscientific Model Development*, 3(1):43–67, 2010.

374 Wendy S. Goliff, William R. Stockwell, and Charlene V. Lawson. The regional atmospheric
 375 chemistry mechanism, version 2. *Atmospheric Environment*, 68:174 – 185, 2013.

376 A. Guenther, T. Karl, P. Harley, C. Wiedinmyer, P. I. Palmer, and C. Geron. Estimates of global
 377 terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from
 378 Nature). *Atmospheric Chemistry and Physics*, 6(11):3181–3210, 2006.

379 A. B. Guenther, X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons, and
 380 X. Wang. The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1):

381 an extended and updated framework for modeling biogenic emissions. *Geoscientific Model*
382 *Development*, 5(6):1471–1492, 2012.

383 Shiro Hatakeyama, Hajime Akimoto, and Nobuaki Washida. Effect of temperature on the
384 formation of photochemical ozone in a propene-nitrogen oxide (NO_x)-air-irradiation system.
385 *Environmental Science & Technology*, 25(11):1884–1890, 1991.

386 Daniel J. Jacob and Darrell A. Winner. Effect of climate change on air quality. *Atmospheric*
387 *Environment*, 43(1):51 – 63, 2009. Atmospheric Environment - Fifty Years of Endeavour.

388 Daniel J Jacob, Jennifer A Logan, Geraldine M Gardner, Rose M Yevich, Clarisa M Spivakovsky,
389 Steven C Wofsy, Sanford Sillman, and Michael J Prather. Factors regulating ozone over the
390 united states and its export to the global atmosphere. *Journal of Geophysical Research*, 98(D8),
391 1993.

392 M. E. Jenkin, S. M. Saunders, V. Wagner, and M. J. Pilling. Protocol for the development of the
393 Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile
394 organic compounds. *Atmospheric Chemistry and Physics*, 3(1):181–193, 2003.

395 M.E. Jenkin, L.A. Watson, S.R. Utembe, and D.E. Shallcross. A Common Representative
396 Intermediates (CRI) mechanism for VOC degradation. Part 1: Gas phase mechanism development.
397 *Atmospheric Environment*, 42(31):7185 – 7195, 2008.

398 Michael E. Jenkin, Sandra M. Saunders, and Michael J. Pilling. The tropospheric degradation of
399 volatile organic compounds: a protocol for mechanism development. *Atmospheric Environment*,
400 31(1):81 – 104, 1997.

401 J. J. P. Kuenen, A. J. H. Visschedijk, M. Jozwicka, and H. A. C. Denier van der Gon.
402 TNO-MACC_II emission inventory; a multi-year (2003–2009) consistent high-resolution european
403 emission inventory for air quality modelling. *Atmospheric Chemistry and Physics*, 14(20):
404 10963–10976, 2014.

405 K. A. Mar, N. Ojha, A. Pozzer, and T. M. Butler. WRF-Chem Simulations over Europe: Model
406 Evaluation and Chemical Mechanism Comparison. *In Preparation*, 2016.

407 N. Otero, J. Sillmann, J. L. Schnell, H. W. Rust, and T. Butler. Synoptic and meteorological
408 drivers of extreme ozone concentrations over europe. *Environmental Research Letters*, 11(2):
409 024005, 2016.

410 N. Passant. Speciation of UK emissions of non-methane volatile organic compounds. Technical
 411 report, DEFRA, Oxon, UK., 2002.

412 George Pouliot, Hugo A.C. Denier van der Gon, Jeroen Kuenen, Junhua Zhang, Michael D. Moran,
 413 and Paul A. Makar. Analysis of the emission inventories and model-ready emission datasets of
 414 Europe and North America for phase 2 of the AQMEII project. *Atmospheric Environment*, 115:
 415 345–360, 2015.

416 S. E. Pusede, D. R. Gentner, P. J. Wooldridge, E. C. Browne, A. W. Rollins, K.-E. Min, A. R.
 417 Russell, J. Thomas, L. Zhang, W. H. Brune, S. B. Henry, J. P. DiGangi, F. N. Keutsch, S. A.
 418 Harrold, J. A. Thornton, M. R. Beaver, J. M. St. Clair, P. O. Wennberg, J. Sanders, X. Ren,
 419 T. C. VandenBoer, M. Z. Markovic, A. Guha, R. Weber, A. H. Goldstein, and R. C. Cohen.
 420 On the temperature dependence of organic reactivity, nitrogen oxides, ozone production, and
 421 the impact of emission controls in San Joaquin Valley, California. *Atmospheric Chemistry and*
 422 *Physics*, 14(7):3373–3395, 2014.

423 Sally E. Pusede, Allison L. Steiner, and Ronald C. Cohen. Temperature and Recent Trends in
 424 the Chemistry of Continental Surface Ozone. *Chemical Reviews*, 115(10):3898–3918, 2015.

425 D. J. Rasmussen, Jianlin Hu, Abdullah Mahmud, and Michael J. Kleeman. The ozone–climate
 426 penalty: Past, present, and future. *Environmental Science & Technology*, 47(24):14258–14266,
 427 2013. PMID: 24187951.

428 Andrew Rickard, Jenny Young, M. J. Pilling, M. E. Jenkin, Stephen Pascoe, and S. M. Saunders.
 429 The Master Chemical Mechanism Version MCM v3.2. <http://mcm.leeds.ac.uk/MCMv3.2/>,
 430 2015. [Online; accessed 25-March-2015].

431 Juli I. Rubin, Andrew J. Kean, Robert A. Harley, Dylan B. Millet, and Allen H. Goldstein.
 432 Temperature dependence of volatile organic compound evaporative emissions from motor vehicles.
 433 *Journal of Geophysical Research: Atmospheres*, 111(D3), 2006. D03305.

434 S. M. Saunders, M. E. Jenkin, R. G. Derwent, and M. J. Pilling. Protocol for the development of
 435 the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic
 436 volatile organic compounds. *Atmospheric Chemistry and Physics*, 3(1):161–180, 2003.

437 Sanford Sillman. The use of NO_y, H₂O₂, and HNO₃ as indicators for ozone-NO_x-hydrocarbon

438 sensitivity in urban locations. *Journal of Geophysical Research: Atmospheres*, 100(D7):
439 14175–14188, 1995.

440 Sanford Sillman. The relation between ozone, NO_x and hydrocarbons in urban and polluted
441 rural environments. *Atmospheric Environment*, 33(12):1821 – 1845, 1999.

442 Sanford Sillman and Perry J. Samson. Impact of temperature on oxidant photochemistry in
443 urban, polluted rural and remote environments. *Journal of Geophysical Research: Atmospheres*,
444 100(D6):11497–11508, 1995.

445 D. Simpson, A. Benedictow, H. Berge, R. Bergström, L. D. Emberson, H. Fagerli, C. R. Flechard,
446 G. D. Hayman, M. Gauss, J. E. Jonson, M. E. Jenkin, A. Nyíri, C. Richter, V. S. Semeena,
447 S. Tsyro, J.-P. Tuovinen, Á. Valdebenito, and P. Wind. The EMEP MSC-W chemical transport
448 model – technical description. *Atmospheric Chemistry and Physics*, 12(16):7825–7865, 2012.

449 Allison L. Steiner, Shaheen Tonse, Ronald C. Cohen, Allen H. Goldstein, and Robert A. Harley.
450 Influence of future climate and emissions on regional air quality in California. *Journal of*
451 *Geophysical Research: Atmospheres*, 111(D18), 2006. D18303.

452 William R. Stockwell, Paulette Middleton, Julius S. Chang, and Xiaoyan Tang. The second
453 generation regional acid deposition model chemical mechanism for regional air quality modeling.
454 *Journal of Geophysical Research: Atmospheres*, 95(D10):16343–16367, 1990.

455 William R. Stockwell, Frank Kirchner, Michael Kuhn, and Stephan Seefeld. A new mechanism
456 for regional atmospheric chemistry modeling. *Journal of Geophysical Research: Atmospheres*,
457 102(D22):25847–25879, 1997.

458 E. von Schneidemesser, J. Coates, A. J. H. Visschedijk, H. A. C. Denier van der Gon, and T. M.
459 Butler. Variation of the NMVOC speciation in the solvent sector and the sensitivity of modelled
460 tropospheric ozone. *Atmospheric Environment*, Submitted for Publication, 2016.

461 Erika von Schneidemesser, Paul S. Monks, James D. Allan, Lori Bruhwiler, Piers Forster, David
462 Fowler, Axel Lauer, William T. Morgan, Pauli Paasonen, Mattia Righi, Katerina Sindelarova,
463 and Mark A. Sutton. Chemistry and the Linkages between Air Quality and Climate Change.
464 *Chemical Reviews*, 2015. PMID: 25926133.

465 Patrick Wagner and Wilhelm Kuttler. Biogenic and anthropogenic isoprene in the near-surface

466 urban atmosphere — A case study in Essen, Germany. *Science of The Total Environment*, 475:
467 104 – 115, 2014.

468 Greg Yarwood, Sunja Rao, Mark Yocke, and Gary Z. Whitten. Updates to the Carbon Bond
469 Chemical Mechanism: CB05. Technical report, U. S Environmental Protection Agency, 2005.