# The Influence of Temperature on Ozone Production under

## varying NO<sub>x</sub> Conditions – a modelling study

- J. Coates<sup>1</sup> and T. Butler<sup>1</sup>
- <sup>1</sup>Institute for Advanced Sustainability Studies, Potsdam, Germany

December 29, 2015

6 Abstract

7

1

2

3

## 8 1 Introduction

Surface-level ozone  $(O_3)$  is a secondary air pollutant formed during the photochemical degradation of volatile organic compounds (VOCs) in the presence of nitrogen oxides  $(NO_x \equiv NO + NO_2)$ . Due to the photochemical nature of ozone production, meteorological variables such as temperature strongly influence ozone production (Jacob and Winner, 2009). A study by Otero et al. (2016) indicated that temperature is a major meteorological driver for ozone in many areas of central Europe during the summertime.

Temperature primarily influences ozone production in two ways: speeding up the reaction rates of many chemical reactions leading to ozone production and increasing emissions of VOCs from biogenic sources (BVOCs). In general, emissions of anthropogenic VOCs (AVOCs) are not typically temperature dependent, although evaporative emissions of AVOCs tend to increase with temperature (Rubin et al., 2006). The review of Pusede et al. (2015) provides further details of the temperature-dependent processes impacting ozone production.

Many studies over the US (Sillman and Samson, 1995; Dawson et al., 2007; Pusede et al., 2014) noted that increased temperatures tend to lead to higher ozone levels, often exceeding local air quality guidelines. Some studies (Sillman and Samson, 1995; Dawson et al., 2007) included modelling experiments using regional chemical tranport models simulating the observed

increases in ozone with temperature. In these studies, the increase of ozone with temperature was attibuted to the decrease in the lifetime of PAN (peroxy acetyl nitrate) at higher temperatures and increased emissions of BVOCs, in particular isoprene, from vegetation.

Pusede et al. (2014) used an analytical model constrained by observations over San Joaquin Valley, California to infer a non-linear relationship of ozone production with temperature and NO<sub>x</sub>, similar to the well-known non-linear relationship of ozone production on NO<sub>x</sub> and VOC levels (Sillman, 1999). Morever, Pusede et al. (2014) showed that temperature can be used as a surrogate for VOC levels when considering the relationship of ozone across NO<sub>x</sub> gradients.

Environmental chamber studies have also been used to analyse the relationship of ozone with temperature for a particular mixture of VOCs. The chamber experiments of Carter et al. (1979) and Hatakeyama et al. (1991), also showed inreases in ozone with temperature linked to increased PAN decomposition at higher temperatures (T >303 K).

Despite many studies considering the effects of temperature on ozone production from an observational and chamber study perspective, there have not been (to our knowledge) modelling studies focusing on the detailed chemical processes of the influence of temperature on ozone production. Where regional modelling has accompanied observational studies, these have only considered the relationship of temperature on ozone under the same  $NO_x$  conditions. The review of Pusede et al. (2015) also highlights a lack of modelling studies looking at the non-linear relationship of ozone on temperature across  $NO_x$  gradients.

In this study, we use an idealised box model to determine how ozone levels vary with temperature and across  $NO_x$  gradients. We separate the effects of temperature-dependent chemistry and temperature-dependent BVOC emissions on ozone production by performing simulations using a temperature-independent source of isoprene followed by simulations using a temperature-dependent source of isoprene at differing temperatures and  $NO_x$  emissions.

## 49 2 Methodology

#### 50 2.1 Model Setup

We used the MECCA box model for this study to determine the important chemical processes for ozone production under different temperatures and  $NO_x$  conditions. The MECCA box model was set up as described in Coates and Butler (2015) and updated to include vertical mixing with the free troposphere and a diurnal cycle for the PBL height based on the data from the

BAERLIN 2014 campaign over Berlin, Germany (Bonn and et.al., 2016). The supplementary
 material includes further details of these updates.

Simulations were performed to broadly simulate urban conditions of central Europe with equinoctical conditions. The simulations started at 06:00 with a total run time of two days. Methane was fixed at 1.7 ppmv throughout the model run, carbon monoxide (CO) and ozone were initialised at 200 ppbv and 40 ppbv and then allowed to evolve freely throughout the the simulation. All VOC emissions were held constant until noon of first day simulating a plume of freshly-emitted VOC.

In order to determine whether increased emissions of BVOC or faster chemistry is more important for the increase of ozone with temperature, model runs were repeated using a temperature-dependent and temperature-independent source of BVOC emissions. MEGAN2.1 (Guenther et al., 2012) specified the temperature-dependent BVOC emissions of isoprene, Sect. 2.3 provides further details. We considered isoprene as representative of all BVOC emissions as isoprene emissions are the most important on the global scale (Guenther et al., 2006). In reality, many other BVOC are emitted from varying vegetation types (Guenther et al., 2006) and increased temperature can also increase AVOC emissions through increased evaporation (Rubin et al., 2006).

Rasmussen et al. (2013) indicated that changing the chemical mechanism used by a model may 72 also change the simulated ozone-temperature relationship and warranted further investigation. All simulations were repeated using different chemical mechanisms to investigate how well the relationship of ozone with temprature across  $NO_x$  gradients is represented. The reference chemical mechanism was the near-explicit Master Chemical Mechanism, MCMv3.2, (Jenkin et al., 1997), 76 (Jenkin et al., 2003), (Saunders et al., 2003), (Rickard et al., 2015). The reduced chemical mechanisms in our study were Common Representative Intermediates, CRIv2 (Jenkin et al., 2008), Model for ozone and related chemical tracers, MOZART-4 (Emmons et al., 2010), Regional Acid Deposition Model, RADM2 (Stockwell et al., 1990) and the Carbon Bond Mechanism, CB05 (Yarwood et al., 2005). Coates and Butler (2015) described these chemical mechanisms and the implementation of these chemical mechanisms in MECCA. These reduced chemical mechanisms were chosen as they are commonly used by modelling groups in 3D regional and global models. 83 Separate box model simulations were performed with systematically varying the temperature between 288 and 313 K (15 – 40  $^{\circ}$ C). The only source of NO<sub>x</sub> emissions in the box model was a constant source of NO emissions. Box model runs were performed with the NO emissions 86

Table 1: Total AVOC emissions in 2011 in tonnes from each SNAP category assigned from TNO-MACC\_III emission inventory and temperature-independent biogenic VOC emissions in tonnes from Benelux region assigned from EMEP. The allocation of these emissions to MCMv3.2, CRIv2, CB05, MOZART-4 and RADM2 species is found in the supplementary material.

	SNAP1	SNAP2	SNAP34	SNAP5	SNAP6	SNAP71
Belgium	4494	9034	22152	5448	42809	6592
Netherlands	9140	12173	29177	8723	53535	16589
Luxembourg	121	44	208	1371	4482	1740
Total	13755	21251	62648	15542	100826	24921
	SNAP72	SNAP73	SNAP74	SNAP8	SNAP9	BVOC
Belgium	2446	144	210	6448	821	7042
Netherlands	3230	1283	1793	10067	521	1462
Luxembourg	1051	6	324	643	0	2198
Total	6727	1433	2327	17158	1342	10702

systematically varied from  $5.0 \times 10^9$  to  $1.5 \times 10^{12}$  molecules (NO) cm<sup>-2</sup> s<sup>-1</sup> at each temperature used in this study.

#### 99 2.2 VOC Emissions

- 90 Typical emissions of urban AVOC over central Europe were taken from TNO-MACC\_III emission
- 91 inventory for the Benelux (Belgium, Netherlands and Luxembourg) region for the year 2011.
- 92 TNO-MACC\_III is the current version of the TNO-MACC\_II emission inventory created using
- the same methodology as Kuenen et al. (2014) and based upon improvements to the existing
- emission inventory during AQMEII-2 (Pouliot et al., 2015).
- Temperature-independent emissions of the biogenic VOC isoprene and monoterpenes were
- calculated as a fraction of the total AVOC emissions from each country in the Benelux region.
- 97 This data was obtained from the supplementary data available from the EMEP (European
- Monitoring and Evaluation Programme) model (Simpson et al., 2012). Temperature-dependent
- emissions of isoprene are detailed in Sect. 2.3.
- AVOC emissions were allocated to SNAP (Selected Nomenclature for Air Pollution) source categories. Table 1 shows the tonnes of NMVOC emissions from each SNAP category and the temperature-independent BVOC emissions. These categorised AVOC emissions were assigned to chemical species and groups based on the country specific profiles for Belgium, the Netherlands and Luxembourg provided by TNO. Most individual chemical species are represented by the MCMv3.2, otherwise the individual contributions of a group of NMVOC were further split into individual components using the detailed speciation of Passant (2002). For example, 'xylenes' are

one of the component chemical groups to many SNAP categories but the MCMv3.2 treats xylenes as the individual isomers (m-, o-, p-xylene) and the individual contributions of the individual isomers to a SNAP category was provided by Passant (2002). This approach was also used in von Schneidemesser et al. (2016) to allocate AVOC emissions from different solvent sector speciations to MCMv3.2 species.

Again similarly to von Schneidemesser et al. (2016), the NMVOC emissions were first assigned to chemcial species represented by the MCMv3.2 and then mapped to the mechanism species representing NMVOC emissions in each reduced chemical mechanism. The NMVOC emissions in the reduced chemical mechanisms were weighted by the carbon numbers of the MCMv3.2 species and the emitted mechanism species. The supplementary data outlines the primary NMVOC and calculated emissions with each chemical mechanism.

#### 2.3 Temperature Dependent Isoprene Emissions

124

125

126

127

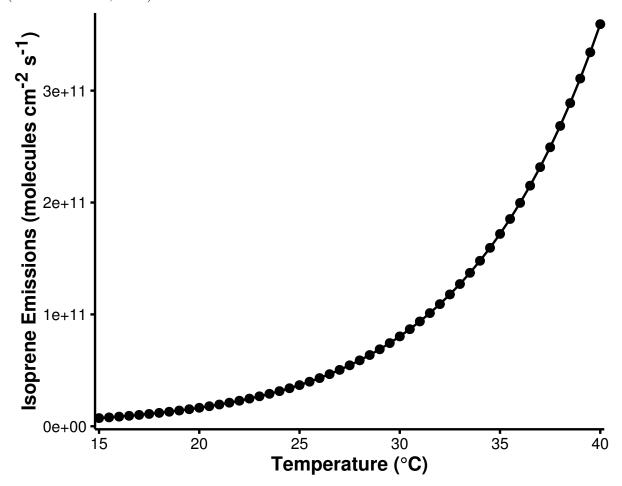
128

Temperature-dependent emissions of isoprene were estimated using the MEGAN2.1 model for calculating the emissions of VOC from vegetation (Guenther et al., 2012). Emissions from plants are dependent on variables including temperature, radiation and age but for the purpose of our study we are only interested in the effects of temperature, hence all variables except temperature were held constant.

The MEGAN2.1 parameters were chosen to give similar isoprene mixing ratios at 20 °C to the temperature-independent emissions of isoprene in order to compare the effects of increased isoprene emissions with temperature. The estimated emissions of isoprene with MEGAN2.1 using these assumptions, are illustrated in Fig. 1 and show the expected exponential increase in emissions with temperature (Guenther et al., 2006).

We compared simulated isoprene mixing ratios from our simulations to measured isoprene 129 emissions from the urban area of Essen, Germany (Wagner and Kuttler, 2014) to verify whether 130 our chosen variables for MEGAN2.1 produced realistic isoprene mixing ratios. At 20 °C, the 131 estimated emissions of isoprene lead to 0.07 ppbv of isoprene in our model while at 30 °C, the increased emissions of isoprene using MEGAN2.1 estimations lead to 0.35 ppbv of isoprene 133 in the model. In the measurement campaign over Essen, 0.1 ppbv of isoprene were measured 134 at temperatures of 20 °C and 0.3 ppbv of isoprene were measured at 30 °C. This comparison 135 indicates that the values chosen for calculating the temperature-dependent emissions of isoprene 136 with MEGAN2.1 lead to reasonable values of isoprene mixing ratios.

Figure 1: The estimated isoprene emissions (molecules isoprene  $cm^{-2} s^{-1}$ ) at each temperature step used in the study. Isoprene emissions were estimated using the MEGAN2.1 algorithm (Guenther et al., 2012).



## 3 Results and Discussion

## 3.1 Ozone as a Function of $NO_x$ and Temperature

Figure 2 depicts the maximum mixing ratio of ozone as a function of the total  $\mathrm{NO_x}$  emissions on the first day of simulations and temperature when using temperature-independent and temperature-dependent source of isoprene emissions for each chemical mechanism. A non-linear relationship of ozone mixing ratios with  $\mathrm{NO_x}$  and temperature is reproduced by each chemical mechanism. This non-linear relationship has a similar form to that determined by Pusede et al. (2014) using an analytical model constrained to observational measurements over the San Joaquin Valley in California.

Larger ozone mixing ratios are produced when using a temperature-dependent source of isoprene emissions in the ridges of the ozone contours and for higher emissions of  $\mathrm{NO_x}$  in Fig. 2.

The highest mixing ratios of ozone in Fig. 2 are produced at higher temperatures and high

Figure 2: Contours of maximum ozone mixing ratios as a function of the total  $\mathrm{NO_x}$  emissions on the first day and temperature for each chemical mechanism using a temperature-dependent and temperature-independent source of isoprene emissions.

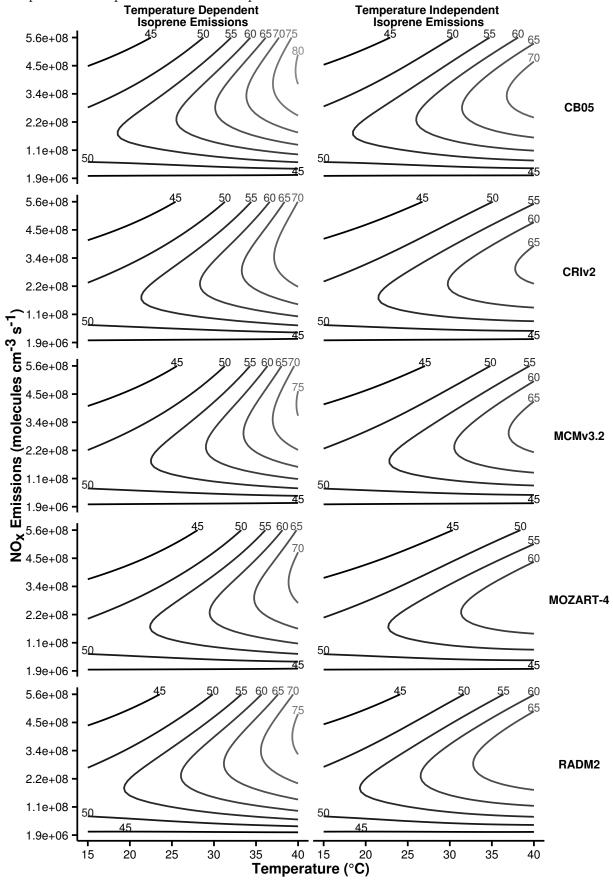
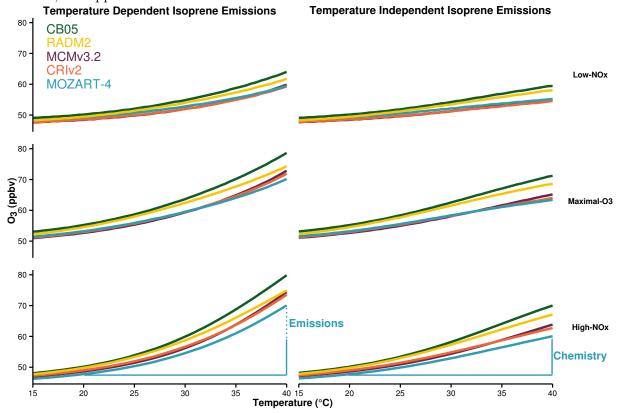


Figure 3: Ozone mixing ratios at each temperature are allocated to different  $NO_x$ -regimes of Fig. 2. The differences in ozone mixing ratios due to chemistry (solid line) and emissions (dotted line) are represented graphically for MOZART-4 with High- $NO_x$  conditions and summarised in Table 2, the approach was used to calculate the differences with each chemical mechanism.



emissions of  $NO_x$  regardless of the source of isoprene emissions. Conversely, the least amount of ozone is produced with low emissions of  $NO_x$  over the whole temperature range (15 – 40 °C) when using both a temperature-independent and temperature-dependent source of isoprene emissions. The contours of ozone mixing ratios as a function of  $NO_x$  and temperature can be split into three  $NO_x$ -regimes (Low- $NO_x$ , Maximal- $O_3$  and High- $NO_x$ ). The Low- $NO_x$  regime corresponds to the regions where there is little increase in ozone with temperature, also called  $NO_x$ -sensitive regime. The High- $NO_x$  (or  $NO_x$ -saturated) regime is when ozone levels increase rapidly with temperature and the ridges of the contours in Fig. 2 correspond to regions of maximal ozone production and we call this the Maximal- $O_3$  regime. Pusede et al. (2014) have shown that temperature can be used as a proxy for VOC, thus we used the ratio of  $HNO_3$  to  $H_2O_2$  used by Sillman (1995) to determine  $NO_x$ -regimes for the non-linear relationship of ozone with  $NO_x$  and VOC to assign the ozone mixing ratios for each simulation to the defined  $NO_x$  regimes. Fig. 3 illustrates the mean ozone mixing ratio at each temperature in the three  $NO_x$  regimes for each chemical mechanism and each source of isoprene emissions.

The absolute increase in ozone at 40 °C from 20 °C due to faster chemistry is the difference in

Table 2: Increase in ozone mixing ratio (ppbv) due to chemistry and emissions at 40 °C from reference temperature (20 °C) in the NO<sub>x</sub>-regimes of Fig. 3.

Chemical	Source of	Increase in Ozone at 40 °C from 20 °C (ppbv)			
Mechanism	Difference	$Low-NO_x$	${ m Maximal-O}_3$	$\operatorname{High-NO}_{\operatorname{x}}$	
MCMv3.2	Chemistry	6.8	12.5	15.2	
	Emissions	4.6	7.7	10.6	
CRIv2	Chemistry	6.0	11.1	13.7	
	Emissions	4.8	7.9	10.8	
MOZART-4	Chemistry	6.0	10.2	12.3	
	Emissions	4.1	6.7	10.0	
CB05	Chemistry	9.3	16.0	19.9	
	Emissions	4.6	7.4	9.8	
RADM2	Chemistry	8.6	14.1	17.3	
	Emissions	3.8	5.7	7.8	

ozone mixing ratios at 40 °C and 20 °C when using a temperature-independent source of isoprene emissions. When using a temperature-dependent source of isoprene emissions, the difference in ozone mixing ratios at 40 °C from 20 °C minus the increase due to faster chemistry, gives the absolute increase in ozone due to increased isoprene emissions. These differences are represented graphically in Fig. 3 and summarised in Table 2.

The results in Table 2 show that the absolute increase in ozone due to faster chemistry is larger than the absolute increase in ozone due to increased isoprene emissions for each cehmical mechanism and each  $NO_x$  regime. Both Fig. 3 and Table 2 highlight that the absolute increase in ozone at 40 °C from 20 °C is largest with High- $NO_x$  conditions. The increase in ozone mixing ratio at 40 °C from 20 °C due to faster chemistry with high- $NO_x$  conditions is almost double that with Low- $NO_x$  conditions. We shall explore which chemical processes are responsible for the increases in ozone mixing ratios at 40 °C from 20 °C by analysing  $O_x$  production budgets in Sect. 3.2.

Comparing the response of ozone mixing ratios to temperature in the reduced chemical mechanisms (CRIv2, MOZART-4, CB05 and RADM2) to the near-explicit MCMv3.2 chemical mechanism shows that the largest differences from the MCMv3.2 occur in the Maximal-O<sub>3</sub> and High-NO<sub>x</sub> regimes. Table 2 also indicates that all reduced chemical mechanisms, except RADM2, have similar increases in ozone due to increased isoprene emissions to MCMv3.2. RADM2 produces 3 ppbv less ozone than the MCMv3.2 due to increased isoprene emissions consistently in each NO<sub>x</sub> regime, indicating that this difference is due to how isoprene degradation chemistry is treated in RADM2.

The Tagged Ozone Production Potential (TOPP) is a measure of the number of molecules

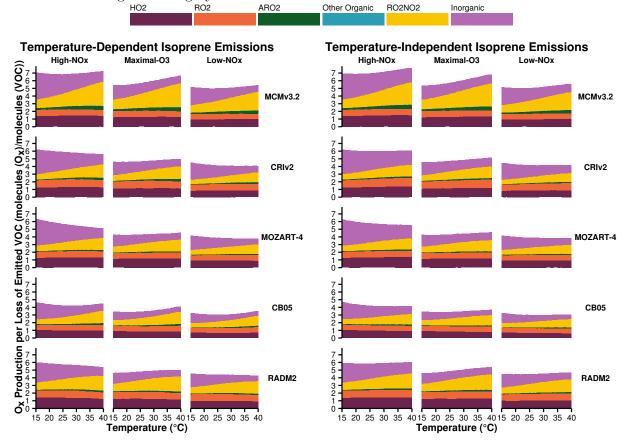
of ozone produced per molecule of VOC emitted originally defined in Butler et al. (2011) and 187 used to compare different chemical mechanisms to the near-explicit MCMv3.2 in Coates and 188 Butler (2015). Coates and Butler (2015) showed that less ozone is produced per molecule of 189 isoprene emitted using RADM2 than with MCMv3.2. The degradation of isoprene has been extensively studied and it is well-known that the secondary degradation species methyl vinyl 191 ketone (MVK) and methacrolein are signatures of isoprene degradation. All chemical mechanisms 192 except RADM2 used in our study explicitly represent MVK and methacrolein (in the case of 193 CB05, a lumped species representing both these secondary degradation products) production 194 during isoprene degradation. RADM2 does not represent methacrolein and the mechanism species representing ketones represents a mixture of acetone and methyl ethyl ketone (MEK). Thus the 196 secondary degradation of isoprene in RADM2 is unable to represent the ozone production from 197 the further degradation of the signature secondary degradation products of isoprene, MVK and 198 methacrolein. Updated versions of RADM2, RACM (Stockwell et al., 1997) and RACM2 (Goliff 199 et al., 2013), sequentially include methacrolein and MVK and with these updates the TOPP 200 values of isoprene approach the TOPP value of isoprene in the MCMv3.2 (Coates and Butler, 201 2015). 202

#### 203 3.2 Ozone Production Budgets

The total day-time production budgets of  $O_x$  ( $\equiv O_3 + NO_2 + O$ ) normalised by the total rate of 204 oxidation of the emitted VOC are displayed in Fig. 4. The  $O_x$  production budgets are assigned 205 to each  $NO_x$  regime for each chemical mechanism and source of isoprene emissions. The budgets 206 allocated to the major sources, where 'HO2', 'RO2', 'ARO2' represent the reactions of NO with 207  $\mathrm{HO}_2$ , alkyl peroxy radicals and acyl peroxy radicals respectively. 'RO2NO2' represents the 208 thermal decomposition of peroxy nitrates, 'Inorganic' represents all inorganic contributions to  $O_x$ 209 production (primarily the de-excitation of O<sup>1</sup>D to O) and any other remaining organic reactions 210 producing  $O_x$  are included in the 'Other Organic' category. 211

In Fig. 4 the number of molecules of  $O_x$  produced per molecule of emitted VOC oxidised in High-NO<sub>x</sub> conditions is similar when using either a temperature-dependent or a temperature-independent isoprene emissions for each chemical mechanism, the same is true for the Maximal-O<sub>3</sub> and Low-NO<sub>x</sub> regimes. Thus the increases in isoprene emissions with temperature are balanced by the faster oxidation rates of the increased amounts of emitted isoprene. The highest amount of  $O_x$  is produced in the High-NO<sub>x</sub> regime and the lowest amount of  $O_x$  is

Figure 4: Day-time  $O_x$  production budgets normalised by the total oxidation rate of emitted VOC in the  $NO_x$ -regimes of Fig. 3. The budgets are allocated to categories of inorganic reactions, peroxy nitrate (RO2NO2) decomposition, reactions of NO with HO2, alkyl peroxy radicals (RO2) and acyl peroxy radicals (ARO2). All other reactions contributing to  $O_x$  budgets are allocated to the 'Other Organic' category.



produced in the Low- $NO_x$  regime, mirroring the  $O_3$  mixing ratios in the different  $NO_x$  regimes in Fig.3. For example, MCMv3.2 produces seven molecules of  $O_x$  per molecule of emitted VOC oxidised in High- $NO_x$  conditions decreasing to about six and five molecules of  $O_x$  per molecule of emitted VOC oxidised in the Maximal- $O_3$  and Low- $NO_x$  regimes.

In each  $NO_x$  regime, each reduced chemical mechanism produces up to two molecules of  $O_x$  per molecule of emitted VOC oxidised less than the MCMv3.2 despite the reduced chemical mechanisms producing similar absolute amounts of ozone to the MCMv3.2 (Fig. 2 and Fig. 3). At high temperatures, most of the difference in the normalised  $O_x$  production using the reduced chemical mechanisms from the MCMv3.2 is due to differences in the contribution from peroxy nitrate (RO2NO2) decomposition. The difference in the maximum contribution of RO2NO2 to normalised  $O_x$  production of the total difference in normalised  $O_x$  production ranges from 65 % using CB05 and up to 86 % with CRIv2.

Thermal decomposition of RO2NO2 also has the largest contribution to the normalised  $O_x$ 

production at higher temperatures in Fig. 4 and will be analysed further in Sect. 3.2.1. While the contribution of RO2NO2 decomposition to normalised  $O_x$  production increases with temperature, the contribution of inorganic reactions to normalised  $O_x$  production decreases with temperature in all  $NO_x$  regimes, each chemical mechanism and regardless of the source of isoprene emissions. Whereas the contributions of the reaction of NO with the peroxy radicals (HO2, RO2 and ARO2 in Fig. 4) to the normalised production of  $O_x$  do not increase strongly with temperature indicating that these processes are strongly related to the faster degradation of the emitted VOC with temperature.

#### 239 3.2.1 Peroxy Nitrates

244

258

We shall now turn our focus to the peroxy nitrate  $(RO_2NO_2)$  contribution to the normalised  $O_x$  production as this category has a strongly temperature-dependent contribution in Fig. 4. Peroxy nitrates are formed from the reactions of alkyl and acyl peroxy nitrates with  $NO_2$  (Reaction R1) and are an important reservoir species for both peroxy radicals and  $NO_x$ .

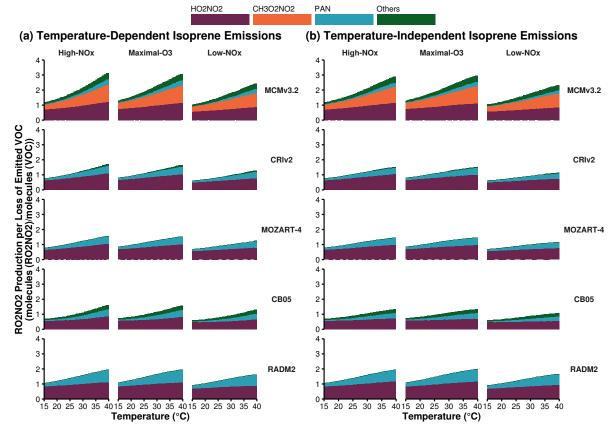
The chemical bond of  $RO_2NO_2$  is quite weak with thermal decomposition being the most

$$RO_2 + NO_2 + M \rightleftharpoons RO_2NO_2 + M$$
 (R1)

important chemical process and the rate of thermal decomposition increases strongly with 245 temperature. At low temperatures, RO<sub>2</sub>NO<sub>2</sub> can accumulate and be transported downwind of emissions of the sources of its precursors (VOC and NO<sub>x</sub>). After thermal decomposition of  $RO_2NO_2$ , the release of  $NO_2$  and peroxy radicals can promote production of  $O_3$  downwind 248 (Moxim et al., 1996). 249 Peroxy nitrates are formed from both alkyl and acyl peroxy radicals produced during the 250 secondary degradation of emitted VOC. The most important alkyl peroxy nitrates are pernitric 251 acid (HO<sub>2</sub>NO<sub>2</sub>) and methylperoxy nitrate (CH<sub>3</sub>O<sub>2</sub>NO<sub>2</sub>), while peroxy acetyl nitrate (PAN, 252  $\mathrm{CH_3C}(\mathrm{O})\mathrm{O_2NO_2})$  and peroxy propionyl nitrate (PPN,  $\mathrm{C_2H_5C}(\mathrm{O})\mathrm{O_2NO_2})$  are important acyl 253 peroxy nitrates. The alkyl peroxy nitrates have a weaker  $RO_2-NO_2$  bond than acyl peroxy 254 nitrates hence alkyl peroxy nitrates have a shorter lifetime than acyl peroxy nitrates. For example, 255  $\mathrm{CH_{3}O_{2}NO_{2}}$  has a lifetime of 0.5 seconds at 298 K while PAN has a lifetime of 51 minutes at 256 298 K (Orlando and Tyndall, 2012).

Each chemical mechanism used in our study represents  $HO_2NO_2$  and PAN, although in

Figure 5: Day-time  $\mathrm{RO_2NO_2}$  production budgets normalised by the total oxidation rate of emitted VOC in the  $\mathrm{NO_x}$ -regimes of Fig. 3. The total budgets are allocated to the most important peroxy nitrates and all other contributions included as 'Others'.



many reduced chemical mechanisms the PAN mechanism species represents  $CH_3C(O)O_2NO_2$  and other acyl peroxy nitrates. This representation of PAN in reduced chemical mechanisms can overestimate PAN levels compared to more detailed chemical mechanisms (Luecken et al., 1999). The near-explicit MCMv3.2 represent a diverse range of peroxy nitrates including  $CH_3O_2NO_2$  and about 280 acyl peroxy nitrates.

Figure 5 displays the normalised production budgets of  $RO_2NO_2$  by the total oxidation rate of the emitted NMVOC, similar to Fig. 4, for each chemical mechanism in each  $NO_x$  regime and when using a temperature-independent and temperature-dependent source of isoprene emissions. The large contribution of  $CH_3O_2NO_2$  to normalised  $RO_2NO_2$  production in MCMv3.2 is missing from the budgets of each reduced chemical mechanism as  $CH_3O_2NO_2$  is not represented in any of the reduced chemical mechanisms. In fact the number of molecules of  $RO_2NO_2$  produced per molecule of emitted VOC oxidised in each reduced chemical mechanism is similar to that of MCMv3.2 without the contribution of  $CH_3O_2NO_2$  for the separate  $NO_x$  regimes and regardless of isoprene source.

The contribution of RO<sub>2</sub>NO<sub>2</sub> to the normalised O<sub>x</sub> production in Fig. 4 is largest in the

Table 3: Slopes ( $m_{O3-T}$  in ppbv per  $^{\circ}C$ ) of the linear fit to the ozone-temperature correlations in Fig. 6

(a) Slope of linear fit of the ERA-Interim observational data and WRF-Chem model output using MOZART-4 and RADM2 chemistry over central and eastern Germany and western and central Poland.

	Germany	Poland
ERA-Interim	2.15	1.94
WRF-Chem with MOZART-4	2.05	2.00
WRF-Chem with RADM2	1.78	1.77

(b) Slope of linear fit of box model experiments for each chemical mechanism, source of isoprene emissions allocated to the three  $NO_x$ -regimes.

Mechanism	Isoprene Emissions	Low-NO <sub>x</sub>	$Maximal-O_3$	$\operatorname{High-NO}_{\mathbf{x}}$
MCMv3.2	Temperature Dependent	0.42	0.74	0.93
	Temperature Independent	0.28	0.51	0.59
CRIv2	Temperature Dependent	0.40	0.71	0.90
	Temperature Independent	0.25	0.47	0.55
MOZART-4	Temperature Dependent	0.38	0.65	0.81
	Temperature Independent	0.25	0.44	0.49
CB05	Temperature Dependent	0.52	0.89	1.12
	Temperature Independent	0.39	0.67	0.79
RADM2	Temperature Dependent	0.48	0.79	0.97
	Temperature Independent	0.37	0.61	0.70

MCMv3.2 than the reduced chemical mechanisms due to the representation of  $CH_3O_2NO_2$  in the MCMv3.2. If reduced chemical mechanisms represent  $CH_3O_2NO_2$  chemistry then this would improve the representation of the total  $RO_2NO_2$  production which would have the added effect of improving the representation of  $O_x$  production budgets.

#### 3.3 Comparison to Observations and 3-D Model Simulations

278

This section compares the results from the idealised box model simulations to real-world 279 observations and model output from a 3-D model verifying the applicability of our results 280 to more realistic atmospheric conditions. Otero et al. (2016) used the interpolated observational 281 data set of the ERA-Interim re-analysis data set of Schnell et al. (2015) which includes the 282 daily maximum temperature and daily maximum 8 h mean of ozone for the years 1998–2012 283 over Europe. In their study, Otero et al. (2016) showed that over the summer (JJA) months, 284 temperature is the main meteorological driver of ozone production over many regions of central 285 Europe. Katie provided model output from the 3-D WRF-Chem regional model set-up over the European domain for simulations of the year 2007 using MOZART-4 and RADM2 chemistry  $\underline{\phantom{a}}$ 

Katie

Figure 6: The maximum 8-h mean ozone from each box model experiment (temperature-dependent and temperature-independent isoprene emissions), allocated to the different  $NO_x$  regimes for each chemical mechanisms (solid lines). The box model ozone-temperature correlation is overlaid with the summer 2007 ERA-Interim data (black circles) and WRF-Chem output using MOZART-4 (purple boxes) and RADM2 (orange triangles).

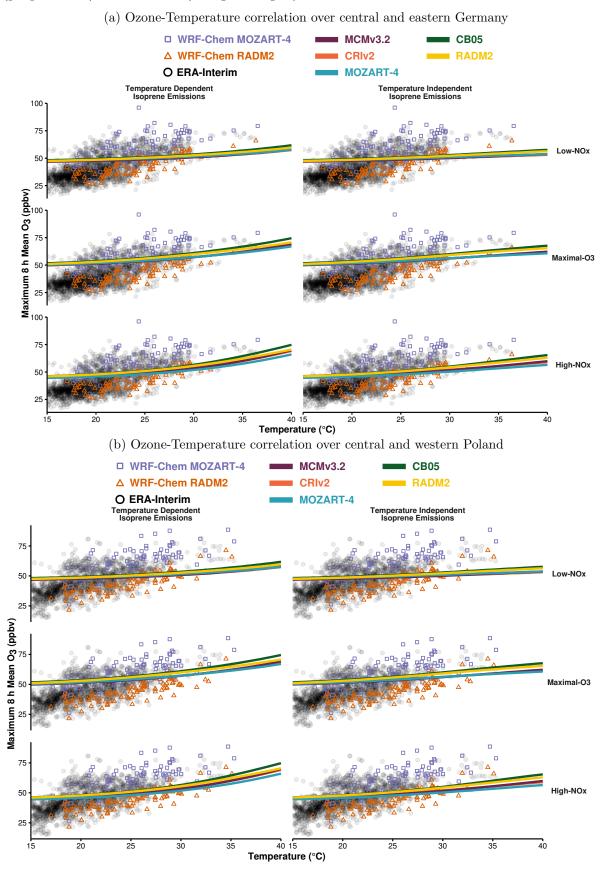


Figure 6 compares the summer data of 2007 from observations (ERA-Interim), WRF-Chem simulations and the maximum 8 h mean ozone from the box model simulations using a temperature-independent and temperature-dependent source of isoprene emissions for each chemical mechanism and allocated to the different  $NO_x$ -regimes. In Fig. 6, only days where the daily maximum temperature corresponded to the temperature range in our study (15–40 °C) were considered from the ERA-Interim observational data set. We selected two regions of the gridded domains of the observations and WRF-Chem output to central and eastern Germany (Fig. 6a) and central and western Poland (Fig. 6b) as the summertime ozone values in these regions are correlated with temperature (Otero et al., 2016). Table 3 summarises the slopes  $(m_{O3-T})$  of the linear fits of all the ozone-temperature correlations displayed in Fig. 6 in ppbv of ozone per °C.

The spread of the ERA-Interim ozone-temperature values over both Germany and Poland are generally captured by the combined WRF-Chem simulations using both MOZART-4 and RADM2 chemistry. However, the ozone-temperature outputs from WRF-Chem using MOZART-4 chemistry only reproduces the higher ozone values with temperature from ERA-Interim but not the lower values. On the other hand, WRF-Chem with RADM2 only reproduces the lower ozone values from the ERA-Interim data. The slope of the linear ozone-temperature correlation from the WRF-Chem simulations using MOZART-4 is closer to the ERA-Interim data than the WRF-Chem simulations using RADM2. However, the ozone values at the lower end of the temperature range (15–18 °C) are not reproduced with WRF-Chem using either MOZART-4 or RADM2 chemistry.

The box model simulations using a temperature-independent source of isoprene emissions do not reproduce the spread of ozone-temperature values of ERA-Interim, also indicated by the  $m_{\mathrm{O3-T}}$  values in Table 3. When using a temperature-dependent source of isoprene emissions in the box model, the ozone-temperature values from the ERA-Interim data over Germany and Poland are improved as seen by the m<sub>O3-T</sub> values in Table 3 but still lower than the observed slopes. High- $NO_x$  conditions with a temperature-dependent source of isoprene produce the most similar ozone-temperature slope to the observational data. However, this closest  $m_{O3-T}$  from the box model simulations is still lower by a factor of two than the observed ozone-temperature slope. In particular, the box model simulations over-predict the ozone values at lower temperatures and under-predict the ozone values at higher temperatures than the ERA-Interim data, regardless of the chemical mechanism. 

One reason for the box model simulations being less sensitive to temperature than the observations are related to the set-up of the experiments compared to observations. Observations consider the total effect of temperature on ozone, while models represent the individual effects of temperature on ozone separately. In other words, observational studies look at the total derivative of ozone with temperature while models look at the partial derivatives of the temperature-dependent processes influencing ozone.

$$\frac{d[\mathcal{O}_3]}{dT} = \frac{\partial[\mathcal{O}_3]}{\partial[\mathcal{B}\mathcal{V}\mathcal{O}\mathcal{C}]} \frac{\partial[\mathcal{B}\mathcal{V}\mathcal{O}\mathcal{C}]}{\partial T} + \frac{\partial[\mathcal{O}_3]}{\partial\mathcal{C}\text{hemistry}} \frac{\partial\mathcal{C}\text{hemistry}}{\partial T} + \frac{\partial[\mathcal{O}_3]}{\partial\mathcal{S}\text{tagnation}} \frac{\partial\mathcal{S}\text{tagnation}}{\partial T} + \dots$$

In our simulations, we focused on determining whether faster chemistry or increased BVOC 320 emissions are more important for the increase of ozone with temperature from a freshly-emitted 321 plume not under stagnant conditions. In stagnant atmospheric conditions, the ozone built-up from 322 the previous day is not transported away from the region and can lead to increased ozone levels 323 with the production of fresh ozone from new emissions. 3-D models such as WRF-Chem would play a valuable role for further work also considering stagnant conditions as these models also 325 include meteorological processes missing from our box model set-up. Despite these short-comings 326 of our box model set-up, the detailed analysis of the chemistry provided in this study should 327 complement any future analysis of the ozone-temperature relationship. 328

## 329 4 Conclusions

340

In this study, we have simulated idealised urban conditions with a box model over a range of 330 temperatures and  $NO_x$  conditions using a temperature-independent and temperature-dependent 331 source of isoprene emissions. These simulations were repeated using reduced chemical mechanism 332 schemes (CRIv2, MOZART-4, CB05 and RADM2) typically used in 3-D models and compared 333 to the near-explicit MCMv3.2 chemical mechanism. Each chemical mechanism produced a 334 non-linear relationship of ozone with temperature and  $NO_x$  with the most ozone produced under 335 high emissions of NO<sub>x</sub>. Conversely, lower NO<sub>x</sub> levels led to a minimal increase of ozone with temperature regardless of the source of isoprene. Thus future air quality under the influence of 337 higher temperatures due to climate change would benefit from reducing  $NO_x$  emissions. 338 Allocating the ozone from each box model simulation to separate NO<sub>x</sub> regimes indicated 339

that faster chemistry with temperature is responsible for a greater absolute increase in ozone

than increased isoprene emissions. The larger increase of ozone with temperature due to faster

 $_{342}$  chemistry was reproduced by each chemical mechanism and in each  $\mathrm{NO}_{\mathrm{x}}$  regime.

Production budgets of ozone were normalised by the total oxidation rate of emitted VOC and allocated to the major contributing processes: peroxy nitrate decomposition, reaction of NO with  $HO_2$ , alkyl and acyl peroxy radicals, inorganic reactions and any other organic reactions. The total normalised ozone production budget in each  $NO_x$  regime was larger in the MCMv3.2 than any reduced chemical mechanism.

The increase in thermal decomposition of peroxy nitrates with temperature has the largest contribution to ozone production with chemical mechanism and each  $NO_x$  regime. The contribution of peroxy nitrates is larger in MCMv3.2 due to the inclusion of methylperoxy nitrate chemistry that is not included in any other chemical mechanism used in this study. Including methylperoxy nitrate chemistry in reduced chemical mechanisms would increase the number of molecules of ozone produced per molecule of emitted VOC oxidised.

The slope (m<sub>O3-T</sub>) of the linear fit of ozone-temperature values from observational data 354 (ERA-Interim) over Europe was twice as high as the closest slope using the box model 355 (temperature-dependent emissions of isoprene and high emissions of NO<sub>x</sub>). Compared to 356 WRF-Chem output using MOZART-4 and RADM2 chemistry, the ozone-temperature values 357 from the box model are again less-sensitive to temperature than the WRF-Chem output. The box 358 model does not represent stagnant atmospheric conditions that are represented by observational 359 values and simulated by 3-D models that include meteorology, such as WRF-Chem. Future work 360 looking at the influence of temperature on ozone should include stagnant conditions to represent more realistic atmospheric conditions. Any modelling work addressing this should also consider 362 a range of  $NO_x$  conditions as this influences the amount of ozone produced. 363

## 64 References

B. Bonn and et.al. Mobile BAERLIN2014: Sources and sinks - The influence of land surface types and horizontal heterogeneity on air pollutant levels in Berlin. *In Preparation*, 2016.

T.M. Butler, M.G. Lawrence, D. Taraborrelli, and J. Lelieveld. Multi-day ozone production potential of volatile organic compounds calculated with a tagging approach. *Atmospheric Environment*, 45(24):4082 – 4090, 2011.

William P. L. Carter, Arthur M. Winer, Karen R. Darnall, and James N. Pitts Jr. Smog chamber

- studies of temperature effects in photochemical smog. Environmental Science & Technology, 13 (9):1094–1100, 1979.
- J. Coates and T. M. Butler. A comparison of chemical mechanisms using tagged ozone production
- potential (TOPP) analysis. Atmospheric Chemistry and Physics, 15(15):8795–8808, 2015.
- John P. Dawson, Peter J. Adams, and Spyros N. Pandis. Sensitivity of ozone to summertime
- climate in the eastern USA: A modeling case study. Atmospheric Environment, 41(7):1494 –
- зт 1511, 2007.
- L. K. Emmons, S. Walters, P. G. Hess, J.-F. Lamarque, G. G. Pfister, D. Fillmore, C. Granier,
- A. Guenther, D. Kinnison, T. Laepple, J. Orlando, X. Tie, G. Tyndall, C. Wiedinmyer, S. L.
- Baughcum, and S. Kloster. Description and evaluation of the Model for Ozone and Related
- chemical Tracers, version 4 (MOZART-4). Geoscientific Model Development, 3(1):43-67, 2010.
- Wendy S. Goliff, William R. Stockwell, and Charlene V. Lawson. The regional atmospheric
- chemistry mechanism, version 2. Atmospheric Environment, 68:174 185, 2013.
- A. Guenther, T. Karl, P. Harley, C. Wiedinmyer, P. I. Palmer, and C. Geron. Estimates of global
- terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from
- Nature). Atmospheric Chemistry and Physics, 6(11):3181–3210, 2006.
- A. B. Guenther, X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons, and
- 388 X. Wang. The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1):
- an extended and updated framework for modeling biogenic emissions. Geoscientific Model
- 390 Development, 5(6):1471-1492, 2012.
- 391 Shiro Hatakeyama, Hajime Akimoto, and Nobuaki Washida. Effect of temperature on the
- formation of photochemical ozone in a propene-nitrogen oxide (NOx)-air-irradiation system.
- 393 Environmental Science & Technology, 25(11):1884–1890, 1991.
- Daniel J. Jacob and Darrell A. Winner. Effect of climate change on air quality. Atmospheric
- Environment, 43(1):51-63, 2009. Atmospheric Environment Fifty Years of Endeavour.
- M. E. Jenkin, S. M. Saunders, V. Wagner, and M. J. Pilling. Protocol for the development of the
- Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile
- organic compounds. Atmospheric Chemistry and Physics, 3(1):181–193, 2003.

- M.E. Jenkin, L.A. Watson, S.R. Utembe, and D.E. Shallcross. A Common Representative
- 400 Intermediates (CRI) mechanism for VOC degradation. Part 1: Gas phase mechanism development.
- 401  $Atmospheric \ Environment, \ 42(31):7185-7195, \ 2008.$
- Michael E. Jenkin, Sandra M. Saunders, and Michael J. Pilling. The tropospheric degradation of
- volatile organic compounds: a protocol for mechanism development. Atmospheric Environment,
- 404 31(1):81 104, 1997.
- 405 J. J. P. Kuenen, A. J. H. Visschedijk, M. Jozwicka, and H. A. C. Denier van der Gon.
- 406 TNO-MACC II emission inventory; a multi-year (2003–2009) consistent high-resolution european
- 407 emission inventory for air quality modelling. Atmospheric Chemistry and Physics, 14(20):
- 408 10963-10976, 2014.
- 409 D.J. Luecken, G.S. Tonnesen, J.E. Sickles, and II. Differences in noy speciation predicted by
- three photochemical mechanisms. Atmospheric Environment, 33(7):1073 1084, 1999.
- W. J. Moxim, H. Levy, and P. S. Kasibhatla. Simulated global tropospheric PAN: Its transport
- and impact on NO<sub>x</sub>. Journal of Geophysical Research: Atmospheres, 101(D7):12621-12638, 1996.
- John J. Orlando and Geoffrey S. Tyndall. Laboratory studies of organic peroxy radical chemistry:
- an overview with emphasis on recent issues of atmospheric significance. Chem. Soc. Rev., 41:
- 415 6294-6317, 2012.
- N. Otero, J. Sillmann, J. L. Schnell, H. Rust, and T. M. Butler. Synoptic and meteorological
- drivers of extreme ozone concentrations over europe. Environmental Research Letters, page In
- Preparation, 2016.
- 419 N. Passant. Speciation of UK emissions of non-methane volatile organic compounds. Technical
- report, DEFRA, Oxon, UK., 2002.
- 421 George Pouliot, Hugo A.C. Denier van der Gon, Jeroen Kuenen, Junhua Zhang, Michael D. Moran,
- and Paul A. Makar. Analysis of the emission inventories and model-ready emission datasets of
- Europe and North America for phase 2 of the AQMEII project. Atmospheric Environment, 115:
- 424 345-360, 2015.
- S. E. Pusede, D. R. Gentner, P. J. Wooldridge, E. C. Browne, A. W. Rollins, K.-E. Min, A. R.
- Russell, J. Thomas, L. Zhang, W. H. Brune, S. B. Henry, J. P. DiGangi, F. N. Keutsch, S. A.
- Harrold, J. A. Thornton, M. R. Beaver, J. M. St. Clair, P. O. Wennberg, J. Sanders, X. Ren,

- T. C. VandenBoer, M. Z. Markovic, A. Guha, R. Weber, A. H. Goldstein, and R. C. Cohen.
- 429 On the temperature dependence of organic reactivity, nitrogen oxides, ozone production, and
- the impact of emission controls in San Joaquin Valley, California. Atmospheric Chemistry and
- 431 Physics, 14(7):3373–3395, 2014.
- 432 Sally E. Pusede, Allison L. Steiner, and Ronald C. Cohen. Temperature and Recent Trends in
- the Chemistry of Continental Surface Ozone. Chemical Reviews, 115(10):3898–3918, 2015.
- D. J. Rasmussen, Jianlin Hu, Abdullah Mahmud, and Michael J. Kleeman. The ozone-climate
- penalty: Past, present, and future. Environmental Science & Technology, 47(24):14258–14266,
- 436 2013. PMID: 24187951.
- Andrew Rickard, Jenny Young, M. J. Pilling, M. E. Jenkin, Stephen Pascoe, and S. M. Saunders.
- 438 The Master Chemical Mechanism Version MCM v3.2. http://mcm.leeds.ac.uk/MCMv3.2/,
- 439 2015. [Online; accessed 25-March-2015].
- Juli I. Rubin, Andrew J. Kean, Robert A. Harley, Dylan B. Millet, and Allen H. Goldstein.
- Temperature dependence of volatile organic compound evaporative emissions from motor vehicles.
- 442 Journal of Geophysical Research: Atmospheres, 111(D3), 2006. D03305.
- 443 S. M. Saunders, M. E. Jenkin, R. G. Derwent, and M. J. Pilling. Protocol for the development of
- the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic
- volatile organic compounds. Atmospheric Chemistry and Physics, 3(1):161–180, 2003.
- J. L. Schnell, M. J. Prather, B. Josse, V. Naik, L. W. Horowitz, P. Cameron-Smith, D. Bergmann,
- 447 G. Zeng, D. A. Plummer, K. Sudo, T. Nagashima, D. T. Shindell, G. Faluvegi, and S. A. Strode.
- <sup>448</sup> Use of north american and european air quality networks to evaluate global chemistry–climate
- modeling of surface ozone. Atmospheric Chemistry and Physics, 15(18):10581-10596, 2015.
- 450 Sanford Sillman. The use of NOy, H2O2, and HNO3 as indicators for ozone-NOx-hydrocarbon
- 451 sensitivity in urban locations. Journal of Geophysical Research: Atmospheres, 100(D7):
- 452 14175–14188, 1995.
- 453 Sanford Sillman. The relation between ozone, NOx and hydrocarbons in urban and polluted
- rural environments. Atmospheric Environment, 33(12):1821 1845, 1999.
- 455 Sanford Sillman and Perry J. Samson. Impact of temperature on oxidant photochemistry in

- urban, polluted rural and remote environments. Journal of Geophysical Research: Atmospheres,
- 457 100(D6):11497-11508, 1995.
- D. Simpson, A. Benedictow, H. Berge, R. Bergström, L. D. Emberson, H. Fagerli, C. R. Flechard,
- 459 G. D. Hayman, M. Gauss, J. E. Jonson, M. E. Jenkin, A. Nyíri, C. Richter, V. S. Semeena,
- S. Tsyro, J.-P. Tuovinen, Á. Valdebenito, and P. Wind. The EMEP MSC-W chemical transport
- model technical description. Atmospheric Chemistry and Physics, 12(16):7825–7865, 2012.
- William R. Stockwell, Paulette Middleton, Julius S. Chang, and Xiaoyan Tang. The second
- 463 generation regional acid deposition model chemical mechanism for regional air quality modeling.
- Journal of Geophysical Research: Atmospheres, 95(D10):16343–16367, 1990.
- William R. Stockwell, Frank Kirchner, Michael Kuhn, and Stephan Seefeld. A new mechanism
- 466 for regional atmospheric chemistry modeling. Journal of Geophysical Research: Atmospheres,
- 467 102(D22):25847-25879, 1997.
- E. von Schneidemesser, J. Coates, A. J. H. Visschedijk, H. A. C. Denier van der Gon, and T. M.
- Butler. Variation of the NMVOC speciation in the solvent sector and the sensitivity of modelled
- 470 tropospheric ozone. Atmospheric Environment, page In preparation, 2016.
- Patrick Wagner and Wilhelm Kuttler. Biogenic and anthropogenic isoprene in the near-surface
- urban atmosphere a case study in essen, germany. Science of The Total Environment, 475:104
- 473 115, 2014.
- 474 Greg Yarwood, Sunja Rao, Mark Yocke, and Gary Z. Whitten. Updates to the Carbon Bond
- <sup>475</sup> Chemical Mechanism: CB05. Technical report, U. S Environmental Protection Agency, 2005.