# The Influence of Temperature on Ozone Production under

# varying NO<sub>x</sub> Conditions – a modelling study

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6 Abstract

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# 8 1 Introduction

Surface-level ozone  $(O_3)$  is a secondary air pollutant formed during the photochemical degradation of volatile organic compounds (VOCs) in the presence of nitrogen oxides  $(NO_x \equiv NO + NO_2)$ . Due to the photochemical nature of ozone production, meteorological variables such as temperature strongly influence ozone production (Jacob and Winner, 2009). A study by Otero et al. (2016) indicated that temperature is a major meteorological driver for ozone in many areas of central Europe during the summertime.

Temperature primarily influences ozone production in two ways: speeding up the reaction rates of many chemical reactions leading to ozone production and increasing emissions of VOCs from biogenic sources (BVOCs). In general, emissions of anthropogenic VOCs (AVOCs) are not typically temperature dependent, although evaporative emissions of AVOCs tend to increase with temperature (Rubin et al., 2006). The review of Pusede et al. (2015) provides further details of the temperature-dependent processes impacting ozone production.

Many studies over the US (Sillman and Samson, 1995; Dawson et al., 2007; Pusede et al., 2014) noted that increased temperatures tend to lead to higher ozone levels, often exceeding local air quality guidelines. Some studies (Sillman and Samson, 1995; Dawson et al., 2007) included modelling experiments using regional chemical tranport models simulating the observed

increases in ozone with temperature. In these studies, the increase of ozone with temperature was attibuted to the decrease in the lifetime of PAN (peroxy acetyl nitrate) at higher temperatures and increased emissions of BVOCs, in particular isoprene, from vegetation.

Pusede et al. (2014) used an analytical model constrained by observations over San Joaquin Valley, California to infer a non-linear relationship of ozone production with temperature and NO<sub>x</sub>, similar to the well-known non-linear relationship of ozone production on NO<sub>x</sub> and VOC levels (Sillman, 1999). Morever, Pusede et al. (2014) showed that temperature can be used as a surrogate for VOC levels when considering the relationship of ozone across NO<sub>x</sub> gradients.

Environmental chamber studies have also been used to analyse the relationship of ozone with temperature for a particular mixture of VOCs. The chamber experiments of Carter et al. (1979) and Hatakeyama et al. (1991), also showed inreases in ozone with temperature linked to increased PAN decomposition at higher temperatures (T >303 K).

Despite many studies considering the effects of temperature on ozone production from an observational and chamber study perspective, there have not been (to our knowledge) modelling studies focusing on the detailed chemical processes of the influence of temperature on ozone production. Where regional modelling has accompanied observational studies, these have only considered the relationship of temperature on ozone under the same  $NO_x$  conditions. The review of Pusede et al. (2015) also highlights a lack of modelling studies looking at the non-linear relationship of ozone on temperature across  $NO_x$  gradients.

In this study, we use an idealised box model to determine how ozone levels vary with temperature and across  $NO_x$  gradients. We separate the effects of temperature-dependent chemistry and temperature-dependent BVOC emissions on ozone production by performing simulations using a temperature-independent source of isoprene followed by simulations using a temperature-dependent source of isoprene at differing temperatures and  $NO_x$  emissions.

# 49 2 Methodology

#### 50 2.1 Model Setup

We used the MECCA box model for this study to determine the important chemical processes for ozone production under different temperatures and  $NO_x$  conditions. The MECCA box model was set up as described in Coates and Butler (2015) and updated to include vertical mixing with the free troposphere and a diurnal cycle for the PBL height based on the data from the

BAERLIN 2014 campaign over Berlin, Germany (Bonn and et.al., 2016). The supplementary
 material includes further details of these updates.

Simulations were performed to broadly simulate urban conditions of central Europe with equinoctical conditions. The simulations started at 06:00 with a total run time of two days. Methane was fixed at 1.7 ppmv throughout the model run, carbon monoxide (CO) and ozone were initialised at 200 ppbv and 40 ppbv and then allowed to evolve freely throughout the the simulation. All VOC emissions were held constant until noon of first day simulating a plume of freshly-emitted VOC.

In order to determine whether increased emissions of BVOC or faster chemistry is more important for the increase of ozone with temperature, model runs were repeated using a temperature-dependent and temperature-independent source of BVOC emissions. MEGAN2.1 (Guenther et al., 2012) specified the temperature-dependent BVOC emissions of isoprene, Sect. 2.3 provides further details. We considered isoprene as representative of all BVOC emissions as isoprene emissions are the most important on the global scale (Guenther et al., 2006). In reality, many other BVOC are emitted from varying vegetation types (Guenther et al., 2006) and increased temperature can also increase AVOC emissions through increased evaporation (Rubin et al., 2006).

Rasmussen et al. (2013) indicated that changing the chemical mechanism used by a model may 72 also change the simulated ozone-temperature relationship and warranted further investigation. All simulations were repeated using different chemical mechanisms to investigate how well the relationship of ozone with temprature across  $NO_x$  gradients is represented. The reference chemical mechanism was the near-explicit Master Chemical Mechanism, MCMv3.2, (Jenkin et al., 1997), 76 (Jenkin et al., 2003), (Saunders et al., 2003), (Rickard et al., 2015). The reduced chemical mechanisms in our study were Common Representative Intermediates, CRIv2 (Jenkin et al., 2008), Model for ozone and related chemical tracers, MOZART-4 (Emmons et al., 2010), Regional Acid Deposition Model, RADM2 (Stockwell et al., 1990) and the Carbon Bond Mechanism, CB05 (Yarwood et al., 2005). Coates and Butler (2015) described these chemical mechanisms and the implementation of these chemical mechanisms in MECCA. These reduced chemical mechanisms were chosen as they are commonly used by modelling groups in 3D regional and global models. 83 Separate box model simulations were performed with systematically varying the temperature between 288 and 313 K (15 – 40  $^{\circ}$ C). The only source of NO<sub>x</sub> emissions in the box model was a constant source of NO emissions. Box model runs were performed with the NO emissions 86

Table 1: Total AVOC emissions in 2011 in tonnes from each SNAP category assigned from TNO-MACC\_III emission inventory and temperature-independent biogenic VOC emissions in tonnes from Benelux region assigned from EMEP. The allocation of these emissions to MCMv3.2, CRIv2, CB05, MOZART-4 and RADM2 species is found in the supplementary material.

	SNAP1	SNAP2	SNAP34	SNAP5	SNAP6	SNAP71
Belgium	4494	9034	22152	5448	42809	6592
Netherlands	9140	12173	29177	8723	53535	16589
Luxembourg	121	44	208	1371	4482	1740
Total	13755	21251	62648	15542	100826	24921
	SNAP72	SNAP73	SNAP74	SNAP8	SNAP9	BVOC
Belgium	2446	144	210	6448	821	7042
Netherlands	3230	1283	1793	10067	521	1462
Luxembourg	1051	6	324	643	0	2198
Total	6727	1433	2327	17158	1342	10702

systematically varied from  $5.0 \times 10^9$  to  $1.5 \times 10^{12}$  molecules (NO) cm<sup>-2</sup> s<sup>-1</sup> at each temperature used in this study.

#### 99 2.2 VOC Emissions

- 90 Typical emissions of urban AVOC over central Europe were taken from TNO-MACC\_III emission
- 91 inventory for the Benelux (Belgium, Netherlands and Luxembourg) region for the year 2011.
- 92 TNO-MACC\_III is the current version of the TNO-MACC\_II emission inventory created using
- the same methodology as Kuenen et al. (2014) and based upon improvements to the existing
- emission inventory during AQMEII-2 (Pouliot et al., 2015).
- Temperature-independent emissions of the biogenic VOC isoprene and monoterpenes were
- calculated as a fraction of the total AVOC emissions from each country in the Benelux region.
- 97 This data was obtained from the supplementary data available from the EMEP (European
- Monitoring and Evaluation Programme) model (Simpson et al., 2012). Temperature-dependent
- emissions of isoprene are detailed in Sect. 2.3.
- AVOC emissions were allocated to SNAP (Selected Nomenclature for Air Pollution) source categories. Table 1 shows the tonnes of NMVOC emissions from each SNAP category and the temperature-independent BVOC emissions. These categorised AVOC emissions were assigned to chemical species and groups based on the country specific profiles for Belgium, the Netherlands and Luxembourg provided by TNO. Most individual chemical species are represented by the MCMv3.2, otherwise the individual contributions of a group of NMVOC were further split into individual components using the detailed speciation of Passant (2002). For example, 'xylenes' are

one of the component chemical groups to many SNAP categories but the MCMv3.2 treats xylenes as the individual isomers (m-, o-, p-xylene) and the individual contributions of the individual isomers to a SNAP category was provided by Passant (2002). This approach was also used in von Schneidemesser et al. (2016) to allocate AVOC emissions from different solvent sector speciations to MCMv3.2 species.

Again similarly to von Schneidemesser et al. (2016), the NMVOC emissions were first assigned to chemcial species represented by the MCMv3.2 and then mapped to the mechanism species representing NMVOC emissions in each reduced chemical mechanism. The NMVOC emissions in the reduced chemical mechanisms were weighted by the carbon numbers of the MCMv3.2 species and the emitted mechanism species. The supplementary data outlines the primary NMVOC and calculated emissions with each chemical mechanism.

#### 2.3 Temperature Dependent Isoprene Emissions

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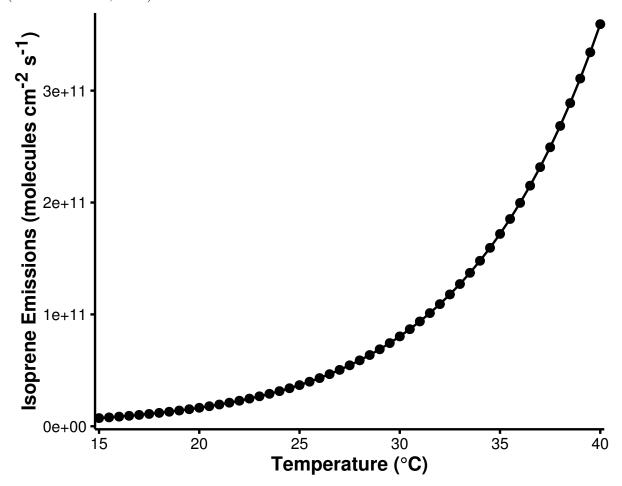
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Temperature-dependent emissions of isoprene were estimated using the MEGAN2.1 model for calculating the emissions of VOC from vegetation (Guenther et al., 2012). Emissions from plants are dependent on variables including temperature, radiation and age but for the purpose of our study we are only interested in the effects of temperature, hence all variables except temperature were held constant.

The MEGAN2.1 parameters were chosen to give similar isoprene mixing ratios at 20 °C to the temperature-independent emissions of isoprene in order to compare the effects of increased isoprene emissions with temperature. The estimated emissions of isoprene with MEGAN2.1 using these assumptions, are illustrated in Fig. 1 and show the expected exponential increase in emissions with temperature (Guenther et al., 2006).

We compared simulated isoprene mixing ratios from our simulations to measured isoprene 129 emissions from the urban area of Essen, Germany (Wagner and Kuttler, 2014) to verify whether 130 our chosen variables for MEGAN2.1 produced realistic isoprene mixing ratios. At 20 °C, the 131 estimated emissions of isoprene lead to 0.07 ppbv of isoprene in our model while at 30 °C, the increased emissions of isoprene using MEGAN2.1 estimations lead to 0.35 ppbv of isoprene 133 in the model. In the measurement campaign over Essen, 0.1 ppbv of isoprene were measured 134 at temperatures of 20 °C and 0.3 ppbv of isoprene were measured at 30 °C. This comparison 135 indicates that the values chosen for calculating the temperature-dependent emissions of isoprene 136 with MEGAN2.1 lead to reasonable values of isoprene mixing ratios.

Figure 1: The estimated isoprene emissions (molecules isoprene  $cm^{-2} s^{-1}$ ) at each temperature step used in the study. Isoprene emissions were estimated using the MEGAN2.1 algorithm (Guenther et al., 2012).



### 3 Results and Discussion

# 3.1 Ozone mixing ratios as function of NOx and Temperature

Figure 2 depicts the maximum mixing ratio of ozone as a function of the total  $NO_x$  emissions on the first day and temperature when using temperature-independent and temperature-dependent source of isoprene emissions for each chemical mechanism. A non-linear relationship of ozone mixing ratios with  $NO_x$  and temperature is reproduced by each chemical mechanism. This non-linear relationship has a similar form to that determined by Pusede et al. (2014) using an analytical model constrained to observational measurements over the San Joaquin Valley in California.

The highest mixing ratios of ozone in Fig. 2 are produced at higher temperatures and high- $NO_x$  conditions, also ozone mixing ratios increase when using a temperature-dependent source of isoprene emissions. Conversely, the least amount of ozone is produced with low- $NO_x$  conditions

Figure 2: Contours of maximum ozone mixing ratio as a function of the total  $\mathrm{NO_x}$  emissions on the first day and temperature for each chemical mechanism and using both a temperature-dependent and -independent source of isoprene emissions.

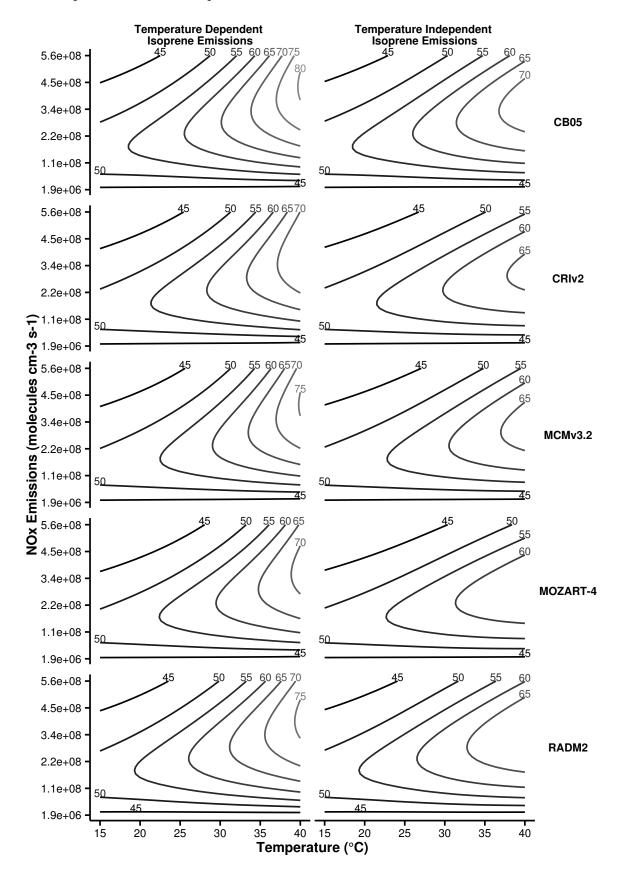
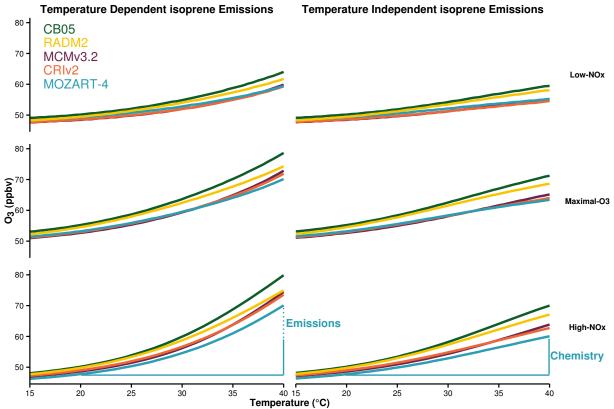


Figure 3: Ozone mixing ratios at each temperature are allocated to different  $NO_x$ -regimes of Fig. 2. The differences in ozone mixing ratios due to chemistry and emissions of Table 2 are represented graphically for MOZART-4, the approach was used to calculate the differences with each chemical mechanism.



over the whole temperature range  $(15 - 40 \, ^{\circ}\text{C})$  when using both a temperature-independent and temperature-dependent source of isoprene emissions.

The non-linear relationship of ozone with  $NO_x$  and temperature can be split into three  $NO_x$ -regimes (low- $NO_x$ , maximal- $O_3$  and high- $NO_x$ ) based on the ratio of  $HNO_3$  to  $H_2O_2$  used in Sillman (1995) to determine  $NO_x$ -regimes for the non-linear relationship of ozone with  $NO_x$  and VOC. The low- $NO_x$  regime corresponds to the lower-left most area in Fig. 2 where there is little increase in ozone with temperature, also called  $NO_x$ -sensitive conditions. The high- $NO_x$  regime is when ozone levels increase rapidly with temperature in Fig. 2, sometimes called  $NO_x$ -saturated conditions. Finally, the ridges of the contours in Fig. 2 correspond to maximal-ozone production and we call this the maximal- $O_3$  regime. The ozone mixing ratios obtained in each simulation were assigned to a  $NO_x$  regime based on the  $H_2O_2$ : $HNO_3$  of the simulation and Fig. 3 illustrates the mean ozone mixing ratio at each temperature in these  $NO_x$  regimes.

Calculating the difference in ozone mixing ratios at 40 °C from 20 °C when using a temperature-independent source of isoprene emissions gives the absolute increase in ozone due to faster chemistry. When using a temperature-dependent source of isoprene emissions, the

Table 2: Increase in ozone mixing ratio (ppbv) due to chemistry and emissions at 40 °C from reference temperature (20 °C) in the NO<sub>x</sub>-regimes of Fig. 2.

Chemical	Source of	Increase in Ozone at 40 °C from 20 °C (ppbv)			
Mechanism	Difference	$Low-NO_x$	${ m Maximal-O}_3$	$\operatorname{High-NO}_{\operatorname{x}}$	
MCMv3.2	Chemistry	6.8	12.5	15.2	
	Emissions	4.6	7.7	10.6	
CRIv2	Chemistry	6.0	11.1	13.7	
	Emissions	4.8	7.9	10.8	
MOZART-4	Chemistry	6.0	10.2	12.3	
	Emissions	4.1	6.7	10.0	
CB05	Chemistry	9.3	16.0	19.9	
	Emissions	4.6	7.4	9.8	
RADM2	Chemistry	8.6	14.1	17.3	
	Emissions	3.8	5.7	7.8	

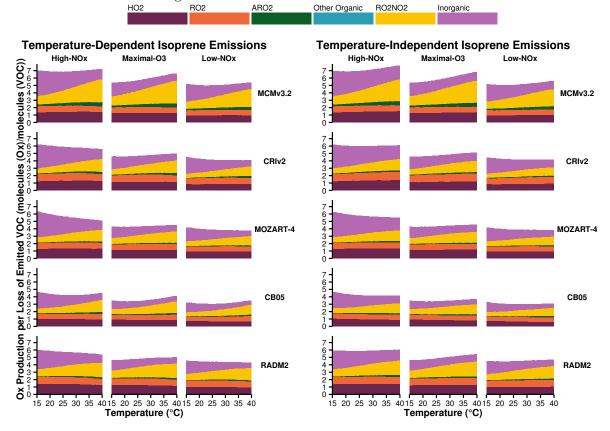
difference in ozone mixing ratios at 40 °C from 20 °C less the increase due to faster chemistry, gives the absolute increase in ozone due to increased isoprene emissions. These differences are represented graphically in Fig. 3 and summarised in Table 2.

Both Fig. 3 and Table 2 highlight that the absolute increase in ozone at 40 °C from 20 °C is largest with high-NO<sub>x</sub> conditions. The increase in ozone mixing ratio at 40 °C from 20 °C due to faster chemistry with high-NO<sub>x</sub> conditions is almost double that with low-NO<sub>x</sub> conditions. We shall explore which chemical processes are responsible for the increases in ozone mixing ratios at 40 °C from 20 °C by analysing  $O_x$  production budgets in Sect. 3.2.

Comparing the response of ozone mixing ratios to temperature in the reduced chemical mechanisms (CRIv2, MOZART-4, CB05 and RADM2) to the near-explicit MCMv3.2 chemical mechanism shows that the largest differences from the MCMv3.2 occur in the maximal- $O_3$  and high-NO<sub>x</sub> regimes. Table 2 also indicates that all reduced chemical mechanisms, except RADM2, have similar increases in ozone due to temperature-dependent isoprene emissions to MCMv3.2. RADM2 produces 3 ppbv less ozone than the MCMv3.2 due to temperature-dependent isoprene emissions consistently in each NO<sub>x</sub> regime, indicating that this difference is due to how isoprene degradation chemistry is treated in RADM2.

The Tagged Ozone Production Potential (TOPP) of isoprene is a measure of the number of molecules of ozone produced per molecule of isoprene emitted and Coates and Butler (2015) shows that less ozone is produced from isoprene degradation using RADM2 than with MCMv3.2. The degradation of isoprene has been extensively studied and it is well-known that the species methyl vinyl ketone (MVK) and methacrolein are signatures of isoprene degradation. All chemical mechanisms used in our study do explicitly include MVK and methacrolein (or in the case of CB05,

Figure 4: Day-time  $O_x$  production budgets normalised by the total oxidation rate of emitted VOC in the  $NO_x$ -regimes of Fig. 2. The budgets are allocated to the categories of inorganic reactions, peroxy nitrate (RO2NO2) decomposition, reactions of NO with HO2, alkyl peroxy radicals (RO2) and acyl peroxy radicals (ARO2). All other reactions contributing to  $O_x$  budgets are allocated to 'Other Organic'.



a lumped species representing both these secondary degradation products) production during isoprene degradation except RADM2. RADM2 does not include methacrolein and the ketone species included in RADM2 represents a mixture of acetone and methyl ethyl ketone (MEK), thus the secondary degradation of isoprene in RADM2 is unable to represent the ozone production from the further degradation of its signature degradation products MVK and methacrolein. More recent versions of RADM2, RACM (Stockwell et al., 1997) and RACM2 (Goliff et al., 2013), sequentially include methacrolein and MVK and with these updates the TOPP values of isoprene reported in Coates and Butler (2015) are similar to the TOPP value of isoprene in the MCMv3.2.

#### 3.2 Ozone Production Budgets

In order to determine which chemical processes are causing the increase in ozone with temperature (Sect. 3.1), we analyse the  $O_x$  production budgets in each  $NO_x$  regime (low- $NO_x$ , Maximal- $O_3$  and high- $NO_x$ ) defined in Sect. 3.1. We defined the  $O_x$  family to consist of  $O_3$ ,  $NO_2$  and  $O_3$ 

and Fig. 4 displays the total day-time  $O_x$  production budgets normalised by the total initial oxidation rates of the emitted NMVOC for each chemical mechanism within each  $NO_x$  regime. The  $O_x$  production budgets in Fig. 4 are allocated to the major sources, where 'HO2', 'RO2', 'ARO2' represent the reactions of NO with  $HO_2$ , alkyl peroxy radicals and acyl peroxy radicals respectively. 'RO2NO2' represents the thermal decomposition of peroxy nitrates, 'Inorganic' represents all inorganic contributions to the  $O_x$  production budgets (primarily the de-excitation of  $O^1D$  to O) and any other remaining organic reactions producing  $O_x$  are included in the 'Other Organic' category.

In Fig. 4 the number of molecules of  $O_x$  produced per molecule of NMVOC oxidised in 207  $\operatorname{High-NO}_{\mathbf{x}}$  conditions is similar when using temperature-dependent or temperature-independent 208 isoprene emissions for each chemical mechanism; the same is also true for the Maximal-O<sub>3</sub> 209 and  $Low-NO_x$  regimes. Thus the increases in isoprene emissions in the temperature-dependent 210 simulations are balanced by the faster oxidation rates of the emitted NMVOC. The highest 211 amount of  $O_x$  is produced in the High-NO<sub>x</sub> regime and the lowest amount of  $O_x$  is produced in 212 the Low- $NO_x$  regime, mirroring the  $O_3$  mixing ratios in the different  $NO_x$  regimes in Fig.3. For 213 example, when using MCMv3.2 seven molecules of  $O_x$  are produced per molecule of NMVOC 214 oxidation in High-NO $_{\rm x}$  conditions, decreased to about six and five molecules of O $_{\rm x}$  produced per 215 molecule of NMVOC oxidised in the Maximal- $O_3$  and Low- $NO_x$  regimes. In each  $NO_x$  regime, 216 all the reduced chemical mechanisms produce up to two molecules of  $O_x$  per molecule of emitted 217 NMVOC oxidised less than the MCMv3.2. 218

Turning to the individual contributions to the normalised production of  $O_x$  in Fig. 4, peroxy nitrate (RO2NO2) decomposition and inorganic reactions show a strong (and opposing) dependence on temperature in all  $NO_x$  regimes, each chemical mechanism and regardless of the source of isoprene emissions. Whereas the contributions of the reaction of NO with the peroxy radicals (HO2, RO2 and ARO2 in Fig. 4) to the normalised production budgets of  $O_x$  do not increase strongly with temperature indicating that these processes are strongly related to the faster oxidation of the emitted NMVOC with temperature.

#### 3.2.1 Peroxynitrates

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We shall now turn our focus to the peroxy nitrate (RO2NO2) contribution as this category has a strongly temperature-dependent contribution to the normalised  $O_x$  production. Peroxy nitrates are an important reservoir species for both peroxy radicals and  $NO_x$  that are formed from the

reactions of alkyl and acyl peroxy nitrates with  $NO_2$  (Reaction R1).

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$$RO_2 + NO_2 + M \rightleftharpoons RO_2NO_2 + M$$
 (R1)

The chemical bond of  $RO_2NO_2$  is quite weak with thermal decomposition being the most important chemical reaction and thermal decomposition depends strongly on temperature. At low temperatures,  $RO_2NO_2$  can accumulate and be transported downwind of emissions of the sources of its precursors (NMVOC and  $NO_x$ ), after thermal decomposition the release of  $NO_2$  and peroxy radicals can promote production of  $O_3$  downwind (Moxim et al., 1996).

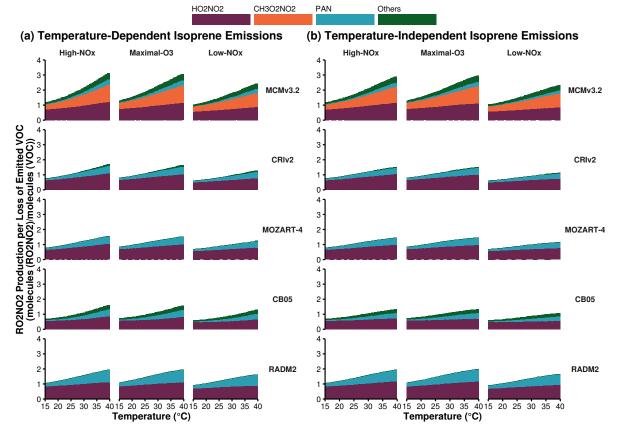
Peroxy nitrates are formed from both alkyl and acyl peroxy radicals, with the acyl peroxy radicals being more thermally stable than the alkyl peroxy nitrates. The most important alkyl peroxy nitrates are pernitric acid ( $HO_2NO_2$ ) and methylperoxy nitrate ( $CH_3O_2NO_2$ ), while peroxy acetyl nitrate (PAN,  $CH_3C(O)O_2NO_2$ ) and peroxy propionyl nitrate (PAN,  $C_2H_5C(O)O_2NO_2$ ) are important acyl peroxy nitrates.

The alkyl peroxy nitrates have a weaker  $\mathrm{RO_2-NO_2}$  bond than acyl peroxy nitrates hence alkyl peroxy nitrates have a shorter lifetime than acyl peroxy nitrates. For example,  $\mathrm{CH_3O_2NO_2}$  has a lifetime of 0.5 seconds at 298 K while PAN has a lifetime of 51 minutes at 298 K (Orlando and Tyndall, 2012).

Each chemical mechanism used in our study represents  $HO_2NO_2$  and PAN, although in many reduced chemical mechanisms PAN represents  $CH_3C(O)O_2NO_2$  and other acyl peroxy nitrates. This representation of PAN in reduced chemical mechanisms can lead to overestimations of PAN levels compared to more detailed chemical mechanisms (Luecken et al., 1999). The near-explicit MCMv3.2 represent a diverse range of peroxy nitrates including  $CH_3O_2NO_2$  and about 280 acyl peroxy nitrates.

Figure 5 displays the normalised production budgets of  $RO_2NO_2$  by the total loss rate of the emitted NMVOC, similar to Fig. 4 for each chemical mechanism in each  $NO_x$  regime and when using a temperature-independent and temperature-dependent source of isoprene emissions. The large contribution of  $CH_3O_2NO_2$  in MCMv3.2 is not mirrored in any reduced chemical mechanism as  $CH_3O_2NO_2$  is not represented in any of the reduced chemical mechanisms. In fact the number of molecules of  $RO_2NO_2$  per molecules of NMVOC oxidised in each reduced chemical mechanism is very similar to that of MCMv3.2 less the contribution of  $CH_3O_2NO_2$  for the separate  $NO_x$  regimes and regardless of isoprene source.

Figure 5: Day-time  $\mathrm{RO_2NO_2}$  production budgets normalised by the total oxidation rate of emitted VOC in the  $\mathrm{NO_x}$ -regimes of Fig. 2. The total budgets are allocated to the most important peroxy nitrates and all other contributions included as 'Others'.



The contribution of RO2NO2 to the normalised  $O_x$  production in Fig. 4 is largest in the MCMv3.2 than the reduced chemical mechanisms due to the representation of  $CH_3O_2NO_2$  in the MCMv3.2. If reduced chemical mechanisms represent  $CH_3O_2NO_2$  chemistry then this would improve the representation of the total  $RO_2NO_2$  production which would have the added effect of improving the representation of  $O_x$  production budgets.

#### 3.3 Comparison to Observations and Regional Model Simulations

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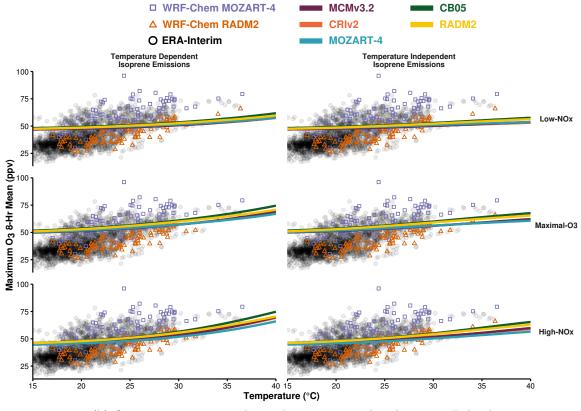
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Our next aim was to compare the results from the detailed box model simulations of this study to real-world observations. The study of Otero et al. (2016) used the interpolated data set of Schnell et al. (2015) of the ERA-Interim re-analysis data set (Dee et al., 2011) which includes the daily maximum temperature and daily maximum 8-h mean of ozone for the years 1998–2012. We were also in a position to compare our box model results to model simulations from the 3-D WRF-Chem regional model that was set-up over the European domain for the year 2007 using MOZART-4 and RADM2 chemistry. We have limited the ERA-Interim data to the summer (JJA) values as the Otero et al. (2016) study showed that summertime values of ozone over

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Figure 6: The maximum 8-h mean ozone from each box model experiment (temperature-dependent and temperature-independent isoprene emissions), allocated to the different  $NO_x$  regimes for each chemical mechanisms (solid lines). The box model ozone-temperature relationship is compared to the summer 2007 ERA-Interim data (black circles) and to WRF-Chem simulations using MOZART-4 (purple boxes) and RADM2 (orange triangles).

(a) Ozone-Temperature relationship over central and eastern Germany



(b) Ozone-Temperature relationship over central and western Poland

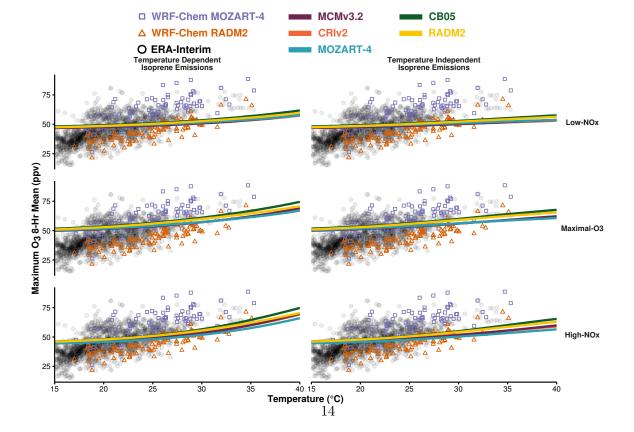


Table 3: Slopes ( $m_{O3-T}$  in ppbv ozone per  $^{\circ}$ C) of the linear fit to the ozone-temperature relationships displayed in Fig. 6

(a) Slope of linear fit of the ERA-Interim observational data and WRF-Chem model output using MOZART-4 and RADM2 chemistry over central and eastern Germany and western and central Poland.

	Germany	Poland
ERA-Interim	2.15	1.94
WRF-Chem with MOZART-4	2.05	2.00
WRF-Chem with RADM2	1.78	1.77

(b) Slope of linear fit of box model experiments for each chemical mechanism, type of isoprene emissions allocated to the three  $NO_x$ -regimes.

Mechanism	Isoprene Emissions	Low-NO <sub>x</sub>	${\bf Maximal \hbox{-} O_3}$	$\operatorname{High-NO}_{\mathbf{x}}$
MCMv3.2	Temperature Dependent	0.42	0.74	0.93
	Temperature Independent	0.28	0.51	0.59
CRIv2	Temperature Dependent	0.40	0.71	0.90
	Temperature Independent	0.25	0.47	0.55
MOZART-4	Temperature Dependent	0.38	0.65	0.81
	Temperature Independent	0.25	0.44	0.49
CB05	Temperature Dependent	0.52	0.89	1.12
	Temperature Independent	0.39	0.67	0.79
RADM2	Temperature Dependent	0.48	0.79	0.97
	Temperature Independent	0.37	0.61	0.70

certain parts of central Europe are primarily driven by temperature and the year 2007 to provide
a comparison with the WRF-Chem output.

Figure 6 compares the summer 2007 data from observations (ERA-Interim), WRF-Chem model output and the maximum 8-hr mean ozone from the box model simulations using a temperature-independent and temperature-dependent source of isoprene emissions for each chemical mechanism and allocated to the different NO<sub>x</sub>-regimes. We have selected two regions of the gridded domains for both the observations and WRF-Chem output to central and eastern Germany (Fig. 6a) and central and western Poland (Fig. 6b) as the summertime ozone values are correlated with temperature (Otero et al., 2016). Table 3 summarises the slopes (m<sub>O3-T</sub>) of the linear fits of all the data displayed in Fig. 6 in ppbv of ozone per °C.

The large spread of the ERA-Interim ozone values over both Germany and Poland at the different temperatures are well captured by the combined WRF-Chem simulations using both MOZART-4 and RADM2 chemistry. The ozone results from the WRF-Chem model using MOZART-4 chemistry re-produce the higher ozone values with temperature from ERA-Interim but not the lower values. On the other hand, the WRF-Chem simulations using RADM2

chemistry only reproduced the lower ozone values from the ERA-Interim data. The slopes of the WRF-Chem simulations using MOZART-4 chemistry are closer to the ERA-Interim data han the WRF-Chem simulations using RADM2 chemistry. However, the ozone values at the lower end of the temperature range (15–18 °C) are not simulated with the WRF-Chem model with either MOZART-4 or RADM2 chemistry.

The box model simulations using a temperature-independent source of isoprene emissions does not reproduce the spread of ozone-temperature data from ERA-Interim, also indicated by the  $m_{O3-T}$  values in Table 3. When using a temperature-dependent source of isoprene emissions in the box model, the ozone-temperature values from the ERA-Interim data over Germany and Poland are reproduced slightly better as indicated by the  $m_{O3-T}$  values in Table 3. The slopes from the box model results is best in the High-NO<sub>x</sub>.

However, even in the box model results that have the most similar  $m_{O3-T}$  to the ERA-Interim data, the slopes with the box model data are half that of the ERA-Interim data over both Germany and Poland. In particular, the box model ozone values at lower temperatures are over-predicted while the ozone values at higher temperatures are under-predicted to the ERA-Interim data.

One reason for the box model simulations being less sensitive to temperature than the observations are related to the set-up of the experiments. In our model set-up, we considered instantaneous ozone production from a freshly emitted emission plume at different temperatures whereas observational values would include ozone and temperature data resulting from other meteorological factors, in particular stagnant conditions. In stagnant conditions, the ozone built-up from the previous day is not transported away from the region and can lead to increased ozone levels with the production of fresh ozone from new emissions.

Observational studies look at the total effect of ozone with temperature, whereas a model can look at the temperature-dependent processes that influence ozone. In other words, observational studies look at the total derivative of ozone with temperature while models can look at the partial derivatives of the temperature-dependent processes influencing ozone.

$$\frac{d[O3]}{dT} = \frac{\partial[O3]}{\partial[BVOC]} \frac{\partial[BVOC]}{\partial T} + \frac{\partial[O3]}{\partial \text{Chemistry}} \frac{\partial \text{Chemistry}}{\partial T} + \frac{\partial[O3]}{\partial \text{Stagnation}} \frac{\partial \text{Stagnation}}{\partial T} + \dots$$

In our experiments, we have focused on determining whether chemistry or increased BVOC emissions are more important for the increase of ozone with temperature but further work including stagnation is also required. 3-D models such as WRF-Chem would play a valuable

role for such further work as these models represent meteorology which is missing from our box model. Despite these short-comings of our box model set-up, the detailed analysis of the chemistry provided in this study should complement any future analysis of the ozone-temperature relationship.

### 317 4 Conclusions

## References

- B. Bonn and et.al. Mobile BAERLIN2014: Sources and sinks The influence of land surface
- types and horizontal heterogeneity on air pollutant levels in Berlin. In Preparation, 2016.
- William P. L. Carter, Arthur M. Winer, Karen R. Darnall, and James N. Pitts Jr. Smog chamber
- studies of temperature effects in photochemical smog. Environmental Science & Technology, 13
- <sup>323</sup> (9):1094–1100, 1979.
- J. Coates and T. M. Butler. A comparison of chemical mechanisms using tagged ozone production
- potential (TOPP) analysis. Atmospheric Chemistry and Physics, 15(15):8795–8808, 2015.
- John P. Dawson, Peter J. Adams, and Spyros N. Pandis. Sensitivity of ozone to summertime
- climate in the eastern USA: A modeling case study. Atmospheric Environment, 41(7):1494 –
- 328 1511, 2007.
- D. P. Dee, S. M. Uppala, A. J. Simmons, P. Berrisford, P. Poli, S. Kobayashi, U. Andrae,
- 330 M. A. Balmaseda, G. Balsamo, P. Bauer, P. Bechtold, A. C. M. Beljaars, L. van de Berg,
- J. Bidlot, N. Bormann, C. Delsol, R. Dragani, M. Fuentes, A. J. Geer, L. Haimberger, S. B. Healy,
- H. Hersbach, E. V. Hólm, L. Isaksen, P. Kållberg, M. Köhler, M. Matricardi, A. P. McNally,
- B. M. Monge-Sanz, J.-J. Morcrette, B.-K. Park, C. Peubey, P. de Rosnay, C. Tavolato, J.-N.
- Thépaut, and F. Vitart. The era-interim reanalysis: configuration and performance of the data
- assimilation system. Quarterly Journal of the Royal Meteorological Society, 137(656):553-597,
- зз6 2011.
- L. K. Emmons, S. Walters, P. G. Hess, J.-F. Lamarque, G. G. Pfister, D. Fillmore, C. Granier,
- A. Guenther, D. Kinnison, T. Laepple, J. Orlando, X. Tie, G. Tyndall, C. Wiedinmyer, S. L.
- 339 Baughcum, and S. Kloster. Description and evaluation of the Model for Ozone and Related
- chemical Tracers, version 4 (MOZART-4). Geoscientific Model Development, 3(1):43-67, 2010.

- Wendy S. Goliff, William R. Stockwell, and Charlene V. Lawson. The regional atmospheric
- chemistry mechanism, version 2. Atmospheric Environment, 68:174 185, 2013.
- A. Guenther, T. Karl, P. Harley, C. Wiedinmyer, P. I. Palmer, and C. Geron. Estimates of global
- terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from
- Nature). Atmospheric Chemistry and Physics, 6(11):3181–3210, 2006.
- A. B. Guenther, X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons, and
- 347 X. Wang. The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1):
- an extended and updated framework for modeling biogenic emissions. Geoscientific Model
- Development, 5(6):1471-1492, 2012.
- 350 Shiro Hatakeyama, Hajime Akimoto, and Nobuaki Washida. Effect of temperature on the
- formation of photochemical ozone in a propene-nitrogen oxide (NOx)-air-irradiation system.
- Environmental Science & Technology, 25(11):1884–1890, 1991.
- Daniel J. Jacob and Darrell A. Winner. Effect of climate change on air quality. Atmospheric
- Environment, 43(1):51-63, 2009. Atmospheric Environment Fifty Years of Endeavour.
- M. E. Jenkin, S. M. Saunders, V. Wagner, and M. J. Pilling. Protocol for the development of the
- Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile
- organic compounds. Atmospheric Chemistry and Physics, 3(1):181–193, 2003.
- M.E. Jenkin, L.A. Watson, S.R. Utembe, and D.E. Shallcross. A Common Representative
- Intermediates (CRI) mechanism for VOC degradation. Part 1: Gas phase mechanism development.
- 360 Atmospheric Environment, 42(31):7185 7195, 2008.
- Michael E. Jenkin, Sandra M. Saunders, and Michael J. Pilling. The tropospheric degradation of
- volatile organic compounds: a protocol for mechanism development. Atmospheric Environment,
- 31(1):81-104, 1997.
- J. J. P. Kuenen, A. J. H. Visschedijk, M. Jozwicka, and H. A. C. Denier van der Gon.
- TNO-MACC\_II emission inventory; a multi-year (2003–2009) consistent high-resolution european
- emission inventory for air quality modelling. Atmospheric Chemistry and Physics, 14(20):
- 367 10963-10976, 2014.
- D.J. Luecken, G.S. Tonnesen, J.E. Sickles, and II. Differences in noy speciation predicted by
- three photochemical mechanisms. Atmospheric Environment, 33(7):1073 1084, 1999.

- W. J. Moxim, H. Levy, and P. S. Kasibhatla. Simulated global tropospheric PAN: Its transport
- and impact on NO<sub>x</sub>. Journal of Geophysical Research: Atmospheres, 101(D7):12621-12638, 1996.
- John J. Orlando and Geoffrey S. Tyndall. Laboratory studies of organic peroxy radical chemistry:
- an overview with emphasis on recent issues of atmospheric significance. Chem. Soc. Rev., 41:
- 374 6294–6317, 2012.
- N. Otero, J. Sillmann, J. L. Schnell, H. Rust, and T. M. Butler. Synoptic and meteorological
- drivers of extreme ozone concentrations over europe. Environmental Research Letters, page In
- Preparation, 2016.
- N. Passant. Speciation of UK emissions of non-methane volatile organic compounds. Technical
- report, DEFRA, Oxon, UK., 2002.
- George Pouliot, Hugo A.C. Denier van der Gon, Jeroen Kuenen, Junhua Zhang, Michael D. Moran,
- and Paul A. Makar. Analysis of the emission inventories and model-ready emission datasets of
- Europe and North America for phase 2 of the AQMEII project. Atmospheric Environment, 115:
- звз 345–360, 2015.
- S. E. Pusede, D. R. Gentner, P. J. Wooldridge, E. C. Browne, A. W. Rollins, K.-E. Min, A. R.
- Russell, J. Thomas, L. Zhang, W. H. Brune, S. B. Henry, J. P. DiGangi, F. N. Keutsch, S. A.
- Harrold, J. A. Thornton, M. R. Beaver, J. M. St. Clair, P. O. Wennberg, J. Sanders, X. Ren,
- T. C. VandenBoer, M. Z. Markovic, A. Guha, R. Weber, A. H. Goldstein, and R. C. Cohen.
- On the temperature dependence of organic reactivity, nitrogen oxides, ozone production, and
- the impact of emission controls in San Joaquin Valley, California. Atmospheric Chemistry and
- 390 *Physics*, 14(7):3373–3395, 2014.
- 391 Sally E. Pusede, Allison L. Steiner, and Ronald C. Cohen. Temperature and Recent Trends in
- the Chemistry of Continental Surface Ozone. Chemical Reviews, 115(10):3898–3918, 2015.
- D. J. Rasmussen, Jianlin Hu, Abdullah Mahmud, and Michael J. Kleeman. The ozone–climate
- penalty: Past, present, and future. Environmental Science & Technology, 47(24):14258-14266,
- <sup>395</sup> 2013. PMID: 24187951.
- Andrew Rickard, Jenny Young, M. J. Pilling, M. E. Jenkin, Stephen Pascoe, and S. M. Saunders.
- The Master Chemical Mechanism Version MCM v3.2. http://mcm.leeds.ac.uk/MCMv3.2/,
- <sup>398</sup> 2015. [Online; accessed 25-March-2015].

- Juli I. Rubin, Andrew J. Kean, Robert A. Harley, Dylan B. Millet, and Allen H. Goldstein.
- Temperature dependence of volatile organic compound evaporative emissions from motor vehicles.
- Journal of Geophysical Research: Atmospheres, 111(D3), 2006. D03305.
- R. Sander, A. Kerkweg, P. Jöckel, and J. Lelieveld. Technical note: The new comprehensive
- atmospheric chemistry module mecca. Atmospheric Chemistry and Physics, 5(2):445–450, 2005.
- S. M. Saunders, M. E. Jenkin, R. G. Derwent, and M. J. Pilling. Protocol for the development of
- the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic
- volatile organic compounds. Atmospheric Chemistry and Physics, 3(1):161–180, 2003.
- J. L. Schnell, M. J. Prather, B. Josse, V. Naik, L. W. Horowitz, P. Cameron-Smith, D. Bergmann,
- 408 G. Zeng, D. A. Plummer, K. Sudo, T. Nagashima, D. T. Shindell, G. Faluvegi, and S. A. Strode.
- 409 Use of north american and european air quality networks to evaluate global chemistry–climate
- modeling of surface ozone. Atmospheric Chemistry and Physics, 15(18):10581–10596, 2015.
- Sanford Sillman. The use of NOy, H2O2, and HNO3 as indicators for ozone-NOx-hydrocarbon
- sensitivity in urban locations. Journal of Geophysical Research: Atmospheres, 100(D7):
- 413 14175–14188, 1995.
- Sanford Sillman. The relation between ozone, NOx and hydrocarbons in urban and polluted
- rural environments. Atmospheric Environment, 33(12):1821 1845, 1999.
- 416 Sanford Sillman and Perry J. Samson. Impact of temperature on oxidant photochemistry in
- urban, polluted rural and remote environments. Journal of Geophysical Research: Atmospheres,
- 100(D6):11497-11508, 1995.
- D. Simpson, A. Benedictow, H. Berge, R. Bergström, L. D. Emberson, H. Fagerli, C. R. Flechard,
- 420 G. D. Hayman, M. Gauss, J. E. Jonson, M. E. Jenkin, A. Nyíri, C. Richter, V. S. Semeena,
- S. Tsyro, J.-P. Tuovinen, Á. Valdebenito, and P. Wind. The EMEP MSC-W chemical transport
- model technical description. Atmospheric Chemistry and Physics, 12(16):7825–7865, 2012.
- William R. Stockwell, Paulette Middleton, Julius S. Chang, and Xiaoyan Tang. The second
- generation regional acid deposition model chemical mechanism for regional air quality modeling.
- 425 Journal of Geophysical Research: Atmospheres, 95(D10):16343-16367, 1990.
- William R. Stockwell, Frank Kirchner, Michael Kuhn, and Stephan Seefeld. A new mechanism

- for regional atmospheric chemistry modeling. Journal of Geophysical Research: Atmospheres,
- 428 102(D22):25847-25879, 1997.
- E. von Schneidemesser, J. Coates, A. J. H. Visschedijk, H. A. C. Denier van der Gon, and T. M.
- Butler. Variation of the NMVOC speciation in the solvent sector and the sensitivity of modelled
- tropospheric ozone. Atmospheric Environment, page In preparation, 2016.
- Patrick Wagner and Wilhelm Kuttler. Biogenic and anthropogenic isoprene in the near-surface
- urban atmosphere a case study in essen, germany. Science of The Total Environment, 475:104
- -115, 2014.
- 435 Greg Yarwood, Sunja Rao, Mark Yocke, and Gary Z. Whitten. Updates to the Carbon Bond
- <sup>436</sup> Chemical Mechanism: CB05. Technical report, U. S Environmental Protection Agency, 2005.