

# Modelling palaeoecological time series using generalized additive models

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**9**

## **10 Abstract**

In the absence of annual laminations, time series generated from lake sediments or other similar stratigraphic sequences are irregularly spaced in time, which complicates formal analysis using classical statistical time series models. In lieu, statistical analyses of trends in palaeoenvironmental time series, if done at all, have typically used simpler linear regressions or (non-) parametric correlations with little regard for the violation of assumptions that almost surely occurs due to temporal dependencies in the data or that correlations do not provide estimates of the magnitude of change, just whether or not there is a linear or monotonic trend. Alternative approaches have used LoESS-estimated trends to justify data interpretations or test hypotheses as to the causal factors without considering the inherent subjectivity of the choice of parameters used to achieve the LoESS fit (e.g. span width, degree of polynomial). Generalized additive models (GAMs) are statistical models that can be used to estimate trends as smooth functions of time. Unlike LoESS, GAMs use automatic smoothness selection methods to objectively determine the complexity of the fitted trend, and as formal statistical models, GAMs, allow for potentially complex, non-linear trends, a proper accounting of model uncertainty, and the identification of periods of significant temporal change. Here, I present a consistent and modern approach to the estimation of trends in palaeoenvironmental time series using GAMs, illustrating features of the methodology with two example time series of contrasting complexity; a 150-year bulk organic matter  $\delta^{15}\text{N}$  time series from Small Water, UK, and a 3000-year alkenone record from Braya-Sø, Greenland. I discuss the underlying mechanics of GAMs that allow them to learn the shape of the trend from the data themselves and how simultaneous confidence intervals and the first derivatives of the trend are used to properly account for

32 model uncertainty and identify periods of change. It is hoped that by using GAMs greater at-  
33 tention is paid to the statistical estimation of trends in palaeoenvironmental time series leading  
34 to more a robust and reproducible palaeoscience.

## 35 1 Introduction

36 Palaeoecology and palaeolimnology have moved away from being descriptive disciplines,  
37 rapidly adopting new statistical developments in the 1990s and beyond (@ Smol et al., 2012).  
38 Less development has been observed in the area of trend estimation in palaeoenvironmental  
39 time series. The vast majority of data produced by palaeoecologists and palaeolimnologists  
40 is in the form of time-ordered observations on one or more proxies or biological taxa (Birks,  
41 2012b; Smol, 2008; Smol et al., 2012). Typically these data are arranged irregularly in time;  
42 in the absence of annual laminae or varves, the sediment core is sectioned at regular depth  
43 intervals (Glew et al., 2001), which, owing to variation in accumulation rates over time  
44 and compaction by overlying sediments, results in an uneven sampling in time. An under-  
45 appreciated feature of such sampling is that younger sediments often have larger variance  
46 than older sediments; each section of core represents fewer lake years in newer samples,  
47 relative to older samples. This variable averaging acts as a time-varying low-pass (high-cut)  
48 filter of the annual depositional signal.

49 Irregular intervals between samples means that the time-series analysis methods of autore-  
50 gressive or moving average processes, in the form of autoregressive integrated moving aver-  
51 age (ARIMA) models, are practically impossible to apply because software typically requires  
52 even spacing of observations in time. Dutilleul et al. (2012) and Birks (2012a), eschewing the  
53 term *time series*, prefer to call such data *temporal series* on account of the irregular spacing of  
54 samples, a distinction that I find unnecessary.

55 Where statistical approaches have been applied to trend estimation in palaeoenvironmental  
56 time series, a commonly-used method is LoESS (Birks, 1998, 2012a; Cleveland, 1979; Juggins  
57 and Telford, 2012). LoESS, locally weighted scatterplot smoother, as it's name suggests, was  
58 developed to smooth x-y scatterplot data (Cleveland, 1979). The method fits a smooth line  
59 through data by fitting weighted least squares (WLS) models to observations within a partic-  
60 ular, user-specified window of the focal point, whose width is typically expressed as a pro-  
61 portion  $\alpha$  of the  $n$  data points. Weights are determined by how close (in the x-axis only) an  
62 observation in the window is to the focal point giving greatest weight given to points closest  
63 to the focal point. The interim LoESS-smoothed value for the focal point is the predicted value  
64 from the weighted regression at the focal point. The interim values are updated using weights  
65 based on how far in the y-axis direction the interim smoothed value lies from the observed  
66 value plus the x-axis distance weights; this has the effect of down-weighting outlier observa-  
67 tions. The final LoESS is obtained by joining the smoothed values. The user has to choose how  
68 large a window to use, whether to fit degree 1 (linear) or degree 2 (quadratic) polynomials in  
69 the WLS model, and how to weight points in the x-axis. When used in an exploratory mode,  
70 the user has considerable freedom to choose the detail of the LoESS fit; the window width, for  
71 example, can be infinitely tweaked to give as close a fit to the data, as assessed by eye, as is de-  
72 sired. Using cross-validation (CV) to choose  $\alpha$  or the degree of polynomial in the WLS model

73 is complicated for a number of reasons, not least because the CV scheme used must involve  
74 the time ordering of the data (e.g. Bergmeir et al., 2018). This subjectivity is problematic how-  
75 ever once we wish to move beyond exploratory analysis and statistically identify trends to test  
76 hypotheses involving those trend estimates.

77 Running means or other types of filter (Juggins and Telford, 2012) have also been used exten-  
78 sively to smooth palaeoenvironmental time series, but, as with LOESS, their behaviour depends  
79 on a number of factors, including the filter width. Furthermore, the width of the filter causes  
80 boundary issues; with a centred filter, of width five, the filtered time series would be two data  
81 points shorter at both ends of the series because the filter values are not defined for the first  
82 and last two observations of the original series as these extra time points were not observed.  
83 Considerable research effort has been expended to identify ways to pad the original time series  
84 at one or both ends to maintain the original length in the filtered series, without introducing  
85 bias due to the padding (e.g. Mann, 2004, 2008; Mills, 2006, 2007, 2010).

86 These are not the only methods that have been used to estimate trends in stratigraphic series.  
87 Another common approach involves fitting a simple linear trend using ordinary least squares  
88 regression and use the resulting  $t$  statistic as evidence against the null hypothesis of no trend  
89 despite the statistical assumptions being almost surely violated due to dependence among  
90 observations. The Pearson correlation coefficient,  $r$ , is also often used to detect trends in palaeo  
91 time series (Birks, 2012a), despite the fact that  $r$  provides no information as to the magnitude of  
92 the estimated trend, and the same temporal autocorrelation problem that dogs ordinary least  
93 squares similarly plagues significance testing for  $r$  (Tian et al., 2011). Additionally, both the  
94 simple least squares trend line and  $r$  are tests for *linear* trends only, and yet we typically face  
95 data sets with potentially far more complex trends than can be identified by these methods.  
96 Instead, non-parametric rank correlation coefficients have been used (Birks, 2012a; Gautier,  
97 2001), and whilst these do allow for the detection of non-linear trends, trends are restricted to  
98 be monotonic, no magnitude of the trend is provided, and the theory underlying significance  
99 testing of Spearman's  $\rho$  and Kendall's  $\tau$  assumes independent observations.

100 Here, I describe generalized additive models (GAMs; Hastie and Tibshirani, 1986, 1990; Rup-  
101 pert et al., 2003; Wood, 2017; Yee and Mitchell, 1991) for trend estimation. GAMs, like simple  
102 linear regression, are a regression-based method for estimating trends, yet they are also, su-  
103 perficially at least, similar to LOESS. GAMs and LOESS estimate smooth, non-linear trends in  
104 time series and both can handle the irregular spacing of samples in time, yet GAMs do not  
105 suffer from the subjectivity that plagues LOESS as a method of formal statistical inference.

106 In the subsequent sections, I present an introduction to GAMs and discuss the issue of uncer-  
107 tainty in model-estimated trends, the topic of posterior simulation from a regression model  
108 and how to identify periods of significant environmental change using the first derivative of  
109 the estimated trend. Two non-standard types of spline — adaptive smoothers and Gaussian  
110 process splines — that are especially applicable to GAMs in the palaeoenvironmental setting  
111 are subsequently described, followed by an assessment of the impact of age-model uncer-  
112 tainty on trend estimation via GAMs. Finally, I briefly discuss the application of GAM trend  
113 analysis to multivariate species abundance and compositional data.

<sup>114</sup> **1.1 Example time series**

<sup>115</sup> To illustrate trend estimation in palaeoenvironmental data using GAMs, I use two proxy time  
<sup>116</sup> series; a 150-year bulk organic matter  $\delta^{15}\text{N}$  record from Small Water, and a 3000-year alkenone  
<sup>117</sup> record from Braya-Sø. Between them, the two examples, combine many of the features of  
<sup>118</sup> interest to palaeoecologists that motivate the use of GAMs; non-linear trends and the question  
<sup>119</sup> of when changes in the measured proxy occurred. The example analyses were all performed  
<sup>120</sup> using the *mgcv* package (version 1.8.23; Wood, 2017) and R (version 3.4.4; R Core Team, 2018),  
<sup>121</sup> and the supplementary material contains a fully annotated document showing the R code used  
<sup>122</sup> to replicate all the analyses described in the remainder of the paper.

<sup>123</sup> **1.1.1  $\delta^{15}\text{N}$  time series from Small Water**

<sup>124</sup> Figure 1a shows 48 nitrogen stable isotope measurements on the bulk organic matter of a sed-  
<sup>125</sup> iment core collected from Small Water, a small corrie lake located in the English Lake District,  
<sup>126</sup> UK. The data were collected to investigate disturbance of nitrogen (N) cycling in remote, oligo-  
<sup>127</sup> trophic lakes by N deposited from the atmosphere (Simpson, unpublished data). The data  
<sup>128</sup> are shown on a  $^{210}\text{Pb}$  time scale. Questions that might be asked about this series are; what is  
<sup>129</sup> the trend in  $\delta^{15}\text{N}$ ?, when do we first see evidence for a change in  $\delta^{15}\text{N}$ ?, and is the reversal in  
<sup>130</sup>  $\delta^{15}\text{N}$  values in the uppermost section of the core a real change?

<sup>131</sup> **1.1.2 Braya-Sø alkenone time series**

<sup>132</sup> The second example time series is a 3,000 year record of alkenone unsaturation,  $U_{37}^K$ , from  
<sup>133</sup> Braya-Sø, a meromictic lake in West Greenland (D'Andrea et al., 2011). Alkenones are long-  
<sup>134</sup> chained unsaturated organic compounds that are produced by a small number of planktonic  
<sup>135</sup> organisms known as haptophytes. The  $U_{37}^K$  unsaturation index is (Brassell, 1993)

$$U_{37}^K = \frac{[C_{37:2}] - [C_{37:4}]}{[C_{37:2}] + [C_{37:3}] + [C_{37:4}]}$$

<sup>136</sup> where  $[C_{37:x}]$  is the concentration of the alkenone with 37 carbon atoms and  $x$  double carbon  
<sup>137</sup> bonds. The relative abundance of these alkenones is known to vary with changes in water  
<sup>138</sup> temperature (Brassell, 1993; Chu et al., 2005; Toney et al., 2010; Zink et al., 2001), and as a  
<sup>139</sup> result  $U_{37}^K$  is used as a proxy for lake- and sea-surface temperatures. For further details on the  
<sup>140</sup> Braya-Sø  $U_{37}^K$  record and age model see D'Andrea et al. (2011). Here I use the 3,000 year  $U_{37}^K$   
<sup>141</sup> record from the PAGES 2K database (PAGES 2K Consortium, 2013). The data are presented in  
<sup>142</sup> Figure 1b.

<sup>143</sup> **2 Regression models for palaeoenvironmental time series**

<sup>144</sup> A linear model for a trend in a series of  $T$  observations  $y_t$  at observation times  $x_t$  with  $t =$   
<sup>145</sup>  $1, 2, \dots, T$  is

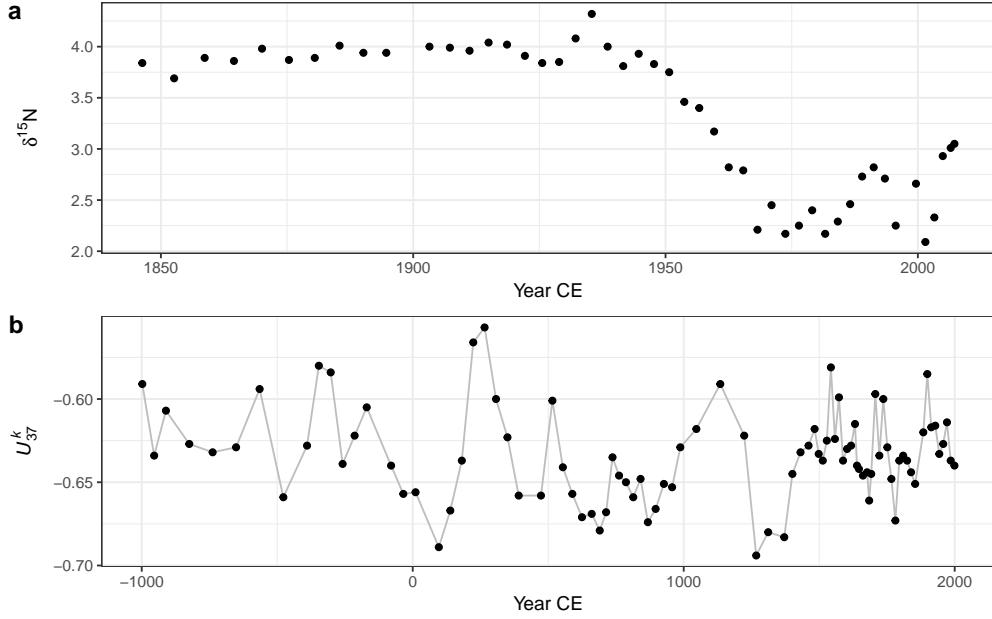


Figure 1: Example time series; a) Small Water bulk organic matter  $\delta^{15}\text{N}$  time series on a  $^{210}\text{Pb}$  time scale, and b) Braya-Sø  $U_{37}^K$  time series on a calibrated  $^{14}\text{C}$  time scale. The observations  $U_{37}^K$  time series have been joined by lines purely as a visual aid to highlight potential trends.

$$y_t = \beta_0 + \beta_1 x_t + \varepsilon_t, \quad (1)$$

where  $\beta_0$  is a constant term, the model *intercept*, representing the expected value of  $y_t$  where  $x_t$  is 0.  $\beta_1$  is the *slope* of the best fit line through the data; it measures the rate of change in  $y$  for a unit increase in  $x$ . The unknowns, the  $\beta_j$ , are commonly estimated using least squares by minimising the sum of squared errors,  $\sum_t \varepsilon_t^2$ . If we want to ask if the estimated trend  $\beta_1$  is statistically significant we must make further assumptions about the data (conditional upon the fitted model) or the model errors (residuals);  $\varepsilon_t \stackrel{iid}{\sim} \mathcal{N}(0, \sigma^2)$ . This notation indicates that the residuals  $\varepsilon_t$  are *independent* and *identically distributed* Gaussian random variables with mean equal to 0 and constant variance  $\sigma^2$ . In the time series setting, the assumption of independence of model residuals is often violated.

The linear model described above is quite restrictive in terms of the types of trend it can fit; essentially linear increasing or decreasing trends, or, trivially, a null trend of no change. This model can be extended to allow for non-linear trends by making  $y_t$  depend on polynomials of  $x_t$ , for example

$$\begin{aligned} y_t &= \beta_0 + \beta_1 x_t + \beta_2 x_t^2 + \cdots + \beta_p x_t^p + \varepsilon_t \\ &= \beta_0 + \sum_{p=1}^P \beta_p x_t^p + \varepsilon_t, \end{aligned} \quad (2)$$

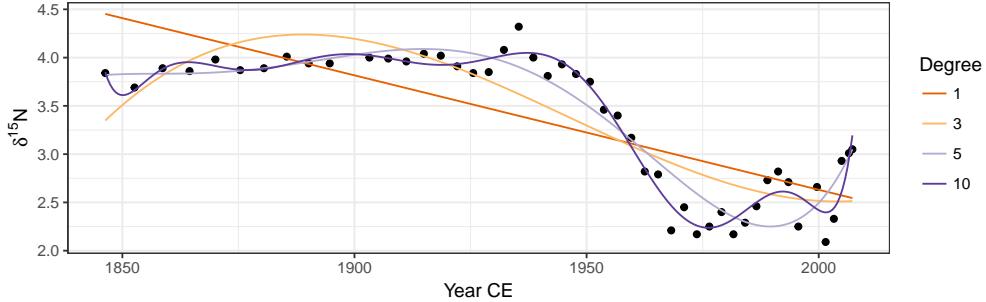


Figure 2: Linear models with various orders of polynomial of the covariate Year fitted using ordinary least squares to the  $\delta^{15}\text{N}$  time series from Small Water. The degree of polynomial is indicated, with the degree 1 line equal to a simple linear regression model.

159 where polynomials of  $x_t$  up to order  $P$  are used. This model allows for more complex trends  
 160 but it remains a fully parametric model and suffers from several problems, especially the be-  
 161 haviour of the fitted trend at the start and end of the observed series.

162 Linear models using a range of polynomials fitted to the Small Water data set are shown in  
 163 Figure 2. The low-order models ( $P \in \{1, 3\}$ ) result in very poor fit to the data. The model with  
 164  $P = 5$  does a reasonable job of capturing the gross pattern in the time series, but fails to adapt  
 165 quickly enough to the decrease in  $\delta^{15}\text{N}$  that begins  $\sim 1940$  CE, and the estimated trend is quite  
 166 biased as a result. The  $P = 10$ -th-order polynomial model is well able to capture this period  
 167 of rapid change, but it does so at the expense of increased complexity in the estimated trend  
 168 prior to  $\sim 1940$ . Additionally, this model ( $P = 10$ ) has undesirable behaviour at the ends of the  
 169 series, significantly overfitting the data, a commonly observed problem in polynomial models  
 170 such as these (Epperson, 1987; Runge, 1901). Finally, the choice of what order of polynomial to  
 171 fit is an additional choice left for the analyst to specify; choosing the optimal  $P$  is not a trivial  
 172 task when the data are a time series and residual autocorrelation is likely present.

173 Can we do better than these polynomial fits? In the remainder, I hope to demonstrate that  
 174 the answer to that question is emphatically “yes”! Below I describe a coherent and consistent  
 175 approach to modelling palaeoenvironmental time series using generalized additive models  
 176 that builds upon the linear regression framework.

### 177 3 Generalized additive models

178 The GAM version of the linear model (1) is

$$y_t = \beta_0 + f(x_t) + \varepsilon_t, \quad (3)$$

179 where the linear effect of time (the  $\beta_1 x_t$  part) has been replaced by a smooth function of time,  
 180  $f(x_t)$ . The immediate advantage of the GAM is that we are no longer restricted to the shapes  
 181 of trends that can be fitted via global polynomial functions such as (2). Instead, the shape of  
 182 the fitted trend will be estimated from the data itself.

183 The linear model is a special case of a broader class, known as the generalized linear model  
 184 (GLM; McCullagh and Nelder, 1989). The GLM provides a common framework for modelling  
 185 a wide range of types of data, such as count, proportions, or binary (presence/absence) data,  
 186 that are not conditionally distributed Gaussian. GLMs are, like the linear model, parametric  
 187 in nature; the types of trends that we can fit using a GLM are the linear or polynomial mod-  
 188 els. GAMs extend the GLM by relaxing this parametric assumption; in a GAM some, or all,  
 189 of the parametric terms, the  $\beta_p$ , are replaced by smooth functions  $f_j$  of the covariates  $x_j$ . For  
 190 completeness then, we can write (3) as a GLM/GAM

$$y_t \sim \text{EF}(\mu_t, \Theta) \quad (4a)$$

$$g(\mu_t) = \beta_0 + f(x_t) \quad (4b)$$

$$\mu_t = g^{-1}(\beta_0 + f(x_t)), \quad (4c)$$

191 where  $\mu_t$  is the expected value (e.g. the mean count or the probability of occurrence) of the  
 192 random variable  $Y_t$  ( $\mu_t \equiv \mathbb{E}(Y_t)$ ) of which we have observations  $y_t$ .  $g$  is the link function, an  
 193 invertible, monotonic function, such as the natural logarithm, and  $g^{-1}$  is its inverse. The link  
 194 function maps values from the response scale on to the scale of the linear predictor, whilst the  
 195 inverse of the link function provides the reverse mapping. For example, count data are strictly  
 196 non-negative integer values and are commonly modelled as a Poisson GLM/GAM using the  
 197 natural log link function. On the log scale, the response can take any real value between  $-\infty$   
 198 and  $+\infty$ , and it is on this scale that model fitting actually occurs (i.e. using equation (4b)).  
 199 However we need to map these unbounded values back on to the non-negative response scale.  
 200 The inverse of the log link function, the exponential function, achieves this and maps values  
 201 to the interval  $0-\infty$  (equation (4c)).

202 In (4a), we further assume that the observations are drawn from a member of the exponential  
 203 family of distributions — such as the Poisson for count data, the binomial for presence/absence  
 204 or counts from a total — with expected value  $\mu_t$  and possibly some additional parameters  $\Theta$   
 205 ( $y_t \sim \text{EF}(\mu_t, \Theta)$ ). Additionally, many software implementations of the above model also allow  
 206 for distributions that are not within the exponential family but which can be fitted using an  
 207 algorithm superficially similar to the one used to fit GAMs to members of the exponential  
 208 family (e.g. Wood et al., 2016). Common examples of such extended families include the  
 209 negative binomial distribution (for overdispersed counts) and the beta distribution (for true  
 210 proportions or other interval-bounded data).

### 211 3.1 Basis functions

212 It is clear from the plots of the data (Figure 1) that we require the fitted trends for the Small  
 213 Water  $\delta^{15}\text{N}$  and Braya-Sø  $U_{37}^K$  time series to be non-linear functions, but it is less clear how to  
 214 specify the actual shape require. Ideally, we'd like to learn the shape of the trend from the  
 215 data themselves. We will refer to these non-linear functions as *smooth functions*, or just *smooths*  
 216 for short, and we will denote a smooth using  $f(x_t)$ . Further, we would like to represent the  
 217 smooths in a way that (4) is represented parametrically so that it can be estimate within the

well-studied GLM framework. This is achieved by representing the smooth using a *basis*. A basis is a set of functions that collectively span a space of smooths that, we hope, contains the true  $f(x_t)$  or a close approximation to it. The functions in the basis are known as *basis functions*, and arise from a *basis expansion* of a covariate. Writing  $b_j(x_t)$  as the  $j$ th basis function of  $x_t$ , the smooth  $f(x_t)$  can be represented as a weighted sum of basis functions

$$f(x_t) = \sum_{j=1}^k b_j(x_t)\beta_j,$$

where  $\beta_j$  is the weight applied to the  $j$ th basis function.

The polynomial model is an example of a statistical model that uses a basis expansion. For the cubic polynomial ( $P = 3$ ) fit shown in Figure 2 there are in fact 4 basis functions:  $b_1(x_t) = x_t^0 = 1$ ,  $b_2(x_t) = x_t$ ,  $b_3(x_t) = x_t^2$ , and  $b_4(x_t) = x_t^3$ . Note that  $b_1(x_t)$  is constant and is linked to the model intercept,  $\beta_0$ , in the linear model (2), and further, that the basis function weights are the estimated coefficients in the model, the  $\beta_j$ .

As we have already seen, polynomial basis expansions do not necessarily lead to well-fitting models unless the true function  $f$  is itself a polynomial. One of the primary criticisms is that polynomial basis functions are global (Magee, 1998); the value of  $f$  at time point  $x_t$  affects the value of  $f$  at time point  $x_{t+s}$  even if the two time points are at opposite ends of the series. There are many other bases we could use; here I discuss one such set of bases, that of splines.

There are a bewildering array of different types of spline. In the models discussed below we will largely restrict ourselves to cubic regression splines (CRS) and thin plate regression splines (TPRS). In addition, I also discuss two special types of spline basis, an adaptive spline basis and a Gaussian process spline basis.

A cubic spline is a smooth curve comprised of sections of cubic polynomials, where the sections are joined together at some specified locations — known as *knots* — in such a way that at the joins, the two sections of cubic polynomial that meet have the same value as well as the same first and second derivative. These properties mean that the sections join smoothly and differentiably at the knots (Wood, 2017, 5.3.1).

The CRS can be parameterized in a number of different ways. One requires a knot at each unique data value in  $x_t$ , which is computationally inefficient. Another way of specifying a CRS basis is to parameterize in terms of the value of the spline at the knots. Typically in this parametrization there are many fewer knots than unique data, with the knots distributed evenly over the range of  $x_t$  or at the quantiles of  $x_t$ . Placing knots at the quantiles of  $x_t$  has the effect of placing a greater number of knots where the data is densest.

A CRS basis expansion comprised of 7 basis functions for the time covariate in the Small Water series, is shown in Figure 3a. The tick marks on the x-axis show the locations of the knots, which are located at the ends of the series and evenly in between. Notice that in this particular parametrization, the  $j$ th basis function takes a value of 1 at the  $j$ th knot and at all other knots a value of 0.

To estimate a model using this basis expansion each basis function forms a column in the model matrix  $\mathbf{X}$  and the weights  $\beta_j$  can be found using least squares regression (assuming a Gaussian

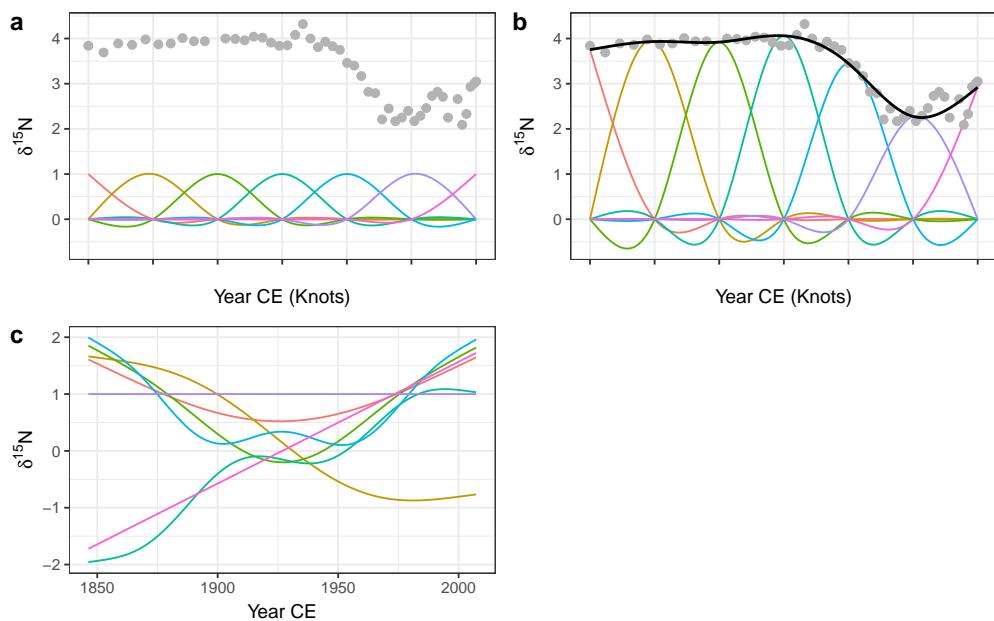


Figure 3: Basis functions for the time covariate and the Small Water  $\delta^{15}\text{N}$  time series. A rank (size) 7 cubic regression spline (CRS) basis expansion is show in a), with knots, indicated by tick marks on the x-axis, spread evenly through the rang of the data. b) shows the same CRS basis functions weighted by the estimated coefficients  $\beta_j$ , plus the resulting GAM trend line (black line drawn through the data). The grey points in both panels are the observed  $\delta^{15}\text{N}$  values. c) A rank 7 thin plate regression spline basis for the same data.

256 response). Note that in order to estimate a coefficient for each basis function the model has  
257 to be fitted without an intercept term. In practice we would include an intercept term in the  
258 model and therefore the basis functions are modified via an identifiability constraint (@ Wood,  
259 2017). This has the effect of making the basis orthogonal to the intercept but results in more  
260 complicated basis functions than those shown in in Figure 3a.

261 Having estimated the weight for each basis function, the  $j$ th basis function  $b_j$  is scaled  
262 (weighted) by its coefficient  $\beta_j$ . The scaled CRS basis functions for the Small Water time series  
263 are shown in Figure 3b. The solid line passing through the data points is formed by summing  
264 up the values of the scaled basis functions ( $b_j(x_t)\beta_j$ ) at any value of  $x_t$  (time).

265 Cubic regression splines, as well as many other types of spline, require the analyst to choose  
266 the number and location of the knots that parametrise the basis. Thin plate regression splines  
267 (TPRS) remove this element of subjectivity when fitting GAMs. Thin plate splines were in-  
268 troduced by Duchon (1977) and, as well as solving the knot selection problem, have several  
269 additional attractive properties in terms of optimality and their ability to estimate a smooth  
270 function of two or more variables, leading to smooth interactions between covariates. How-  
271 ever, thin plate splines have one key disadvantage over CRS; thin plate splines have as many  
272 unknown parameters as there are unique combinations of covariate values in a data set (Wood,  
273 2017, 5.5.1). It is unlikely that any real data problem would involve functions of such complex-  
274 ity that they require as many basis functions as data. It is much more likely that the true func-  
275 tions that we attempt to estimate are far simpler than the set of functions representable by 1  
276 basis function per unique data value. From a practical point of view, it is also highly inefficient  
277 to carry around all these basis functions whilst model fitting, and the available computational  
278 resources would become quickly exhausted for large time series with many observations.

279 To address this issue low rank thin plate regression splines (TPRS) have been suggested which  
280 truncate the space of the thin plate spline basis to some lower number of basis functions whilst  
281 preserving much of the advantage of the original basis as an optimally-fitting spline (Wood,  
282 2003). A rank 7 TPRS basis (i.e. one containing 7 basis functions) is shown in Figure 3c for the  
283 Small Water time series. The truncation is achieved by performing an eigen-decomposition  
284 of the basis functions and retaining the eigenvectors associated with the  $k$  largest eigenvalues.  
285 This is similar to the way principal components analysis decomposes a data set into axes of  
286 variation (eigenvectors) in decreasing order of variance explained. The truncated basis can  
287 preserve much of the space of functions spanned by the original basis but at the cost of using  
288 far fewer basis functions (Wood, 2003, 2017, 5.5.1). Note the horizontal TPRS basis function (at  
289  $\delta^{15}\text{N} = 1$ ) in Figure 3c; this basis function is confounded with the intercept term and, after the  
290 application of identifiability constraints, ends up being removed from the set of basis functions  
291 used to fit the model.

292 The truncation suggested by Wood (2003) is not without cost; the eigen-decomposition and  
293 related steps can be relatively costly for large data sets. For data sets of similar size to the two  
294 examples used here, the additional computational effort required to set up the TPRS basis over  
295 the CRS basis will not be noticeable. For highly resolved series containing more than  $\sim 1000$   
296 observations the truncation may be costly computationally. In such instances, little is lost by  
297 moving to the CRS basis with the same number of knots as the rank of the desired TPRS, with  
298 the benefit of considerably reduced set up time for the basis.

299 To fit a GAM using either of the two regression spline bases described above, the analyst is  
 300 generally only required to specify the size (rank) of the basis expansion required to rep-  
 301 resent or closely approximate the true function  $f$ . With practice and some knowledge of the  
 302 system from which the observations arise, it can be relatively easy to put an upper limit on  
 303 the expected complexity of the true trend in the data. Additionally, the number available data  
 304 points places a constraint on the upper limit of the size of basis expansion that can be used.  
 305 In practice, the size of the basis is an upper limit on the expected complexity of the trend,  
 306 and a simple test to check if the basis used was sufficiently large (Pya and Wood, 2016). This  
 307 test is available via the `gam.check()` function in *mgcv* for example, which essentially looks at  
 308 whether there is any additional nonlinearity or structure in the residuals that can be explained  
 309 by a further smooth of  $x_t$ . Should a smooth term in the fitted model fail this test the model can  
 310 be refitted using a larger basis expansion, say by doubling the value of  $k$  (the rank) used to fit  
 311 the original. Note also that a smooth might fail this test whilst using fewer effective degrees  
 312 of freedom than the maximum possible for the dimension of basis used. This may happen  
 313 when the true function lies at the upper limit of the set of functions encompassed by the size  
 314 of basis used. Additionally, a basis of size  $2k$  encompasses a richer space of functions of a  
 315 given complexity than a basis of size  $k$  (Wood, 2017); increasing the basis dimension used to fit  
 316 the model may unlock this additional function space resulting in a better fitting model whilst  
 317 using a similar number of effective degrees of freedom.

## 318 3.2 Smoothness selection

319 Having identified low rank regression splines as a useful way to represent  $f$ , we next need  
 320 a way to decide how wiggly the fitted trend should be. A backwards elimination approach  
 321 to sequentially remove knots or basis functions might seem appropriate, however such an  
 322 approach would likely fail as the resulting sequence of models would not be strictly nested,  
 323 precluding many forms of statistical comparison (Wood, 2017). Alternatively, we could keep  
 324 the basis dimension at a fixed size but guard against fitting very complex models through the  
 325 use of a wigginess penalty.  
 326 The default wigginess penalty used in GAMs is on the second derivative of the spline, which  
 327 measures the rate of change of the slope, or the curvature, of the spline at any infinitesimal  
 328 point in the interval spanned by  $x_t$ . The actual penalty used is the integrated squared second  
 329 derivative of the spline

$$\int_{\mathbb{R}} [f'']^2 dx = \beta^T \mathbf{S} \beta. \quad (5)$$

330 The right hand side of (5) is the penalty in quadratic form. The convenience of the quadratic  
 331 form is that it is a function of the estimated coefficients of  $f(x_t)$  where  $\mathbf{S}$  is known as the penalty  
 332 matrix. Notice that now both the weights for the basis functions and the wigginess penalty  
 333 are expressed as functions of the model coefficients.  
 334 Now that we have a convenient way to measure wigginess, it needs to be incorporated into  
 335 the objective function that will be minimised to fit the GAM. The likelihood of the model given

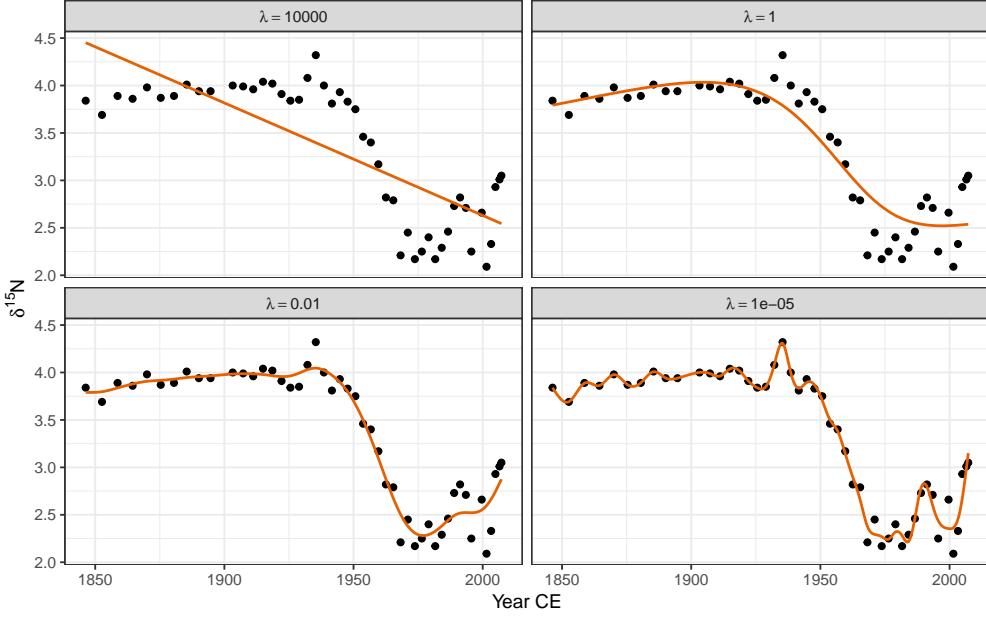


Figure 4: The effect of the smoothness parameter  $\lambda$  on the resulting wigginess of the estimated spline. Large values of  $\lambda$  penalize wigginess strongly, resulting in smooth trends (upper row), while smaller values allow increasingly wiggly trends. The aim of automatic smoothness selection is to find an optimal value of  $\lambda$  that balances the fit of the model with model complexity to avoid overfitting.

336 the parameter estimates  $\mathcal{L}(\beta)$  is combined with the penalty to create the penalized likelihood  
 337  $\mathcal{L}_p(\beta)$ :

$$\mathcal{L}_p(\beta) = \mathcal{L}(\beta) - \frac{1}{2}\lambda\beta^T S\beta.$$

338 The fraction of a half is there simply to make the penalised likelihood equal the penalised  
 339 sum of squares in the case of a Gaussian model.  $\lambda$  is known as the smoothness parameter  
 340 and controls the extent to which the penalty contributes to the likelihood of the model. In  
 341 the extreme case of  $\lambda = 0$  the penalty has no effect and the penalized likelihood equals the  
 342 likelihood of the model given the parameters. At the other extreme, as  $\lambda \rightarrow \infty$  the penalty  
 343 comes to dominate  $\mathcal{L}_p(\beta)$  and the wigginess of  $f(x_t)$  tends to 0 resulting in an infinitely smooth  
 344 function. In the case of a second derivative penalty, this is a straight line, and we recover the  
 345 simple linear trend from (1) when assuming a Gaussian response.

346 Figure 4 illustrates how the smoothness parameter  $\lambda$  controls the degree of wigginess in the  
 347 fitted spline. Four models are shown, each fitted with a fixed value of  $\lambda$ ; 10000, 1, 0.01, and  
 348 0.00001. At  $\lambda = 10000$  the model effectively fits a linear model through the data. As the value  
 349 of  $\lambda$  decreases, the fitted spline becomes increasingly wiggly. As  $\lambda$  becomes very small, the  
 350 resulting spline passes through most of the  $\delta^{15}\text{N}$  observations resulting in a model that is  
 351 clearly over fitted to the data.

352 To fully automate smoothness selection for  $f(x_t)$  we need to estimate  $\lambda$ . There are two main

353 ways that  $\lambda$  can be automatically chosen during model fitting. The first way is to choose  $\lambda$   
354 such that it minimises the prediction error of the model. This can be achieved by choosing  $\lambda$   
355 to minimise Akaike's information criterion (AIC) or via cross-validation (CV) or generalized  
356 cross-validation (GCV; Craven and Wahba, 1978). GCV avoids the computational overhead  
357 inherent to CV of having to repeatedly refit the model with one or more observations left out  
358 as a test set. Minimising the GCV score will, with a sufficiently large data set, find a model  
359 with the minimal prediction error (Wood, 2017). The second approach is to treat the smooth  
360 as a random effect, in which  $\lambda$  is now a variance parameter to be estimated using maximum  
361 likelihood (ML) or restricted maximum likelihood (REML; Wood, 2011; Wood et al., 2016).

362 Several recent results have shown that GCV, under certain circumstances, has a tendency to  
363 under smooth, resulting in fitted splines that are overly wiggly (Reiss and Ogden, 2009). Much  
364 better behaviour has been observed for REML and ML smoothness selection, in that order  
365 (Wood, 2011). REML is therefore the recommended means of fitting GAMs, though, where  
366 models have different fixed effects (covariates) they cannot be compared using REML, and  
367 ML selection should be used instead. In the sorts of data examples considered here there is  
368 only a single covariate  $x_t$  as our models contain a single estimated trend so REML smoothness  
369 selection is used throughout unless otherwise stated.

## 370 4 Fitting GAMs

### 371 4.1 Small Water

372 The trend in  $\delta^{15}\text{N}$  values is clearly non-linear but it would be difficult to suggest a suitable  
373 polynomial model that would allow for periods of relatively no change in  $\delta^{15}\text{N}$  as well as rapid  
374 change. Instead, a GAM is ideally suited to modelling such trends; the data suggest a smoothly  
375 varying change in  $\delta^{15}\text{N}$  between 1925 and 1975. It is reasonable to expect some autocorrelation  
376 in the model errors about the fitted trend. Therefore I fitted the following GAM to the  $\delta^{15}\text{N}$   
377 time series.

$$y_t = \beta_0 + f(x_t) + \varepsilon, \quad \varepsilon_t \sim (0, \Lambda\sigma^2) \quad (6)$$

378 Now the i.i.d. assumption has been relaxed and a correlation matrix,  $\Lambda$ , has been introduced  
379 that is used to model autocorrelation in the residuals. The  $\delta^{15}\text{N}$  values are irregularly spaced in  
380 time and a correlation structure that can handle the uneven spacing is needed (Pinheiro and  
381 Bates, 2000). A continuous time first-order autoregressive process (CAR(1)) is a reasonable  
382 choice; it is the continuous-time equivalent of the first-order autoregressive process (AR(1))  
383 and, simply stated, models the correlation between any two residuals as an exponentially de-  
384 creasing function of  $h(\phi^h)$ , where  $h$  is the amount of separation in time between the residuals  
385 (Pinheiro and Bates, 2000).  $h$  may be a real valued number in the CAR(1), which is how it can  
386 accommodate the irregular separation of samples in time.  $\phi$  controls how quickly the corre-  
387 lation between any two residuals declines as a function of their separation in time and is an  
388 additional parameter that will be estimated during model fitting. The model in (6) was fitted

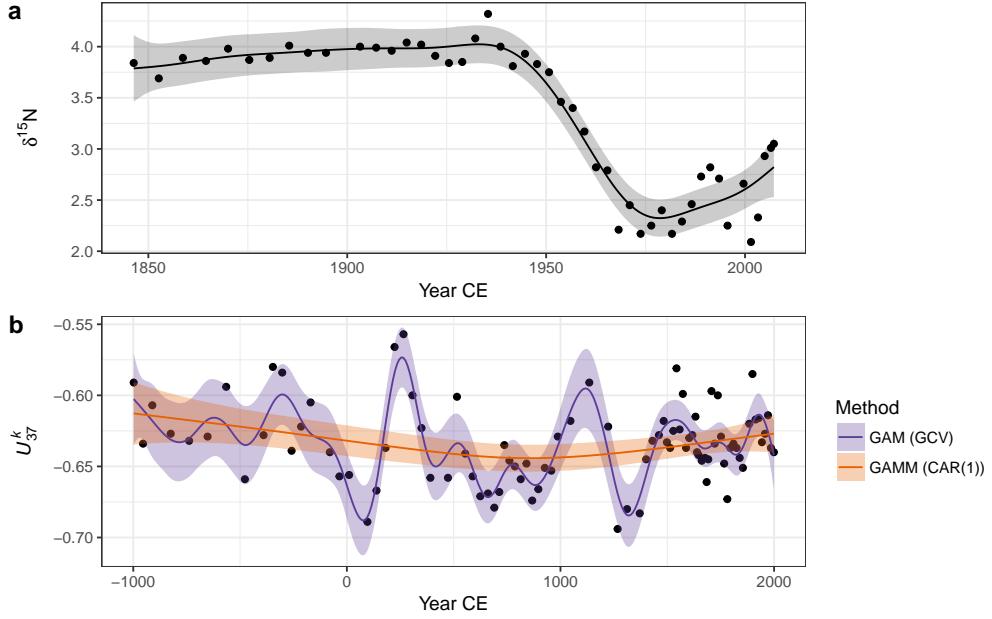


Figure 5: GAM-based trends fitted to the Small Water  $\delta^{15}\text{N}$  (a) and Braya-Sø  $U_{37}^K$  (b) time series. The shaded bands surrounding the estimated trends are approximate 95% across-the-function confidence intervals. For the  $U_{37}^K$  series, two models are shown; the orange fit is the result of a GAM with a continuous-time AR(1) process estimated using REML smoothness selection, while the blue fit is that of a simple GAM with GCV-based smoothness selection. The REML-based fit significantly oversmooths the  $U_{37}^K$  time series.

389 using the `gamm()` function (Wood, 2004) in the `mgcv` package (Wood, 2017) for R (R Core Team,  
 390 2017).

391 The fitted trend is shown in Figure 5a, and well-captures the strong pattern in the data. The  
 392 trend is statistically significant (estimated degrees of freedom = 7.95;  $F = 47.44$ , approximate  $p$   
 393 value =  $\ll 0.0001$ ). However further analysis of the fitted model is required to answer the other  
 394 questions posed earlier about the timing of change and whether features in the trend can be  
 395 distinguished from random noise. I discuss these issues shortly.

## 396 4.2 Braya-Sø

397 The  $U_{37}^K$  data present a more difficult data analysis challenge than the  $\delta^{15}\text{N}$  time series because  
 398 of the much more complex variation present. Fitting the same model as the Small Water ex-  
 399 ample, (6), to the  $U_{37}^K$  data resulted in the unsatisfactory fit shown as the very smooth line in  
 400 Figure 5b (labelled Gamm (CAR(1))). Further problems were evident with this model fit —  
 401 the covariance matrix of the model was non-positive definite, a sure sign of problems with the  
 402 fitted model. Refitting with a smaller basis dimension ( $k = 20$ ) for the trend term resulted in  
 403 a model with a positive-definite covariance matrix for the model variance-covariance terms,  
 404 but the estimated value of the CAR(1) parameter  $\phi = 0.2$  was exceedingly uncertain (95%  
 405 confidence interval 0 – 1!).

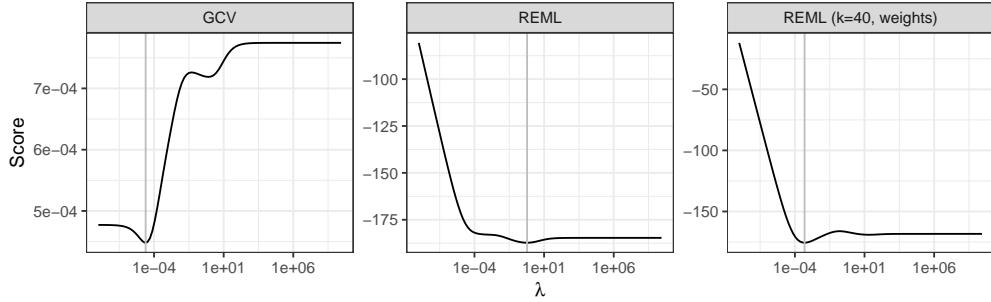


Figure 6: GCV and REML scores as a function of the smoothness parameter  $\lambda$ . From left to right, GAMs were estimated using GCV and REML smoothness selection, and REML using a basis dimension of 40 and observational weights to account for heterogeneity in the  $U_{37}^K$  times series. The selected value of  $\lambda$  for each model is indicated by the vertical grey line.

406 Fitting this model as a standard GAM with REML smoothness selection resulted in the same  
 407 fitted trend as the GAM with CAR(1) errors (not shown), whilst using GCV smoothness se-  
 408 lection resulted in a much more satisfactory fitted trend. There are two potential problems  
 409 with the GCV-selected trend: i) GCV is sensitive to the profile of the GCV score and has been  
 410 shown to under smooth data in situations where the profile is flat around the minimum GCV  
 411 score, and ii) the model fitted assumes that the observations are independent, an assumption  
 412 that is certainly violated in the  $U_{37}^K$  time series.

413 To investigate the first issue, the GCV and REML scores for an increasing sequence of values  
 414 of the smoothness parameter ( $\lambda$ ) were evaluated for the standard GAM (equation (4)) fit to the  
 415  $U_{37}^K$  time series. The resulting profiles are shown in Figure 6, with the optimal value of the  
 416 parameter shown by the vertical line. The GCV score profile suggests that the potential for  
 417 under smoothing identified by Reiss and Ogden (2009) is unlikely to apply here as there is a  
 418 well-defined minimum in profile.

419 To understand the reason why the GAM plus CAR(1) and the simple GAM with REML smooth-  
 420 ness selection performed poorly with the  $U_{37}^K$  time series we need to delve a little deeper into  
 421 what is happening when we are fitting these two models.

422 The primary issue leading to poor fit is that neither model accounts for the different variance  
 423 (known as (heteroscedasticity) of each observation in the  $U_{37}^K$  record. This seemingly isn't a  
 424 problem for GCV which minimizes prediction error. The sediments in Braya-Sø are not an-  
 425 nually laminated and therefore the core was sliced at regular depth intervals. Owing to com-  
 426 paction of older sediments and variation in accumulation rates over time, each sediment slice  
 427 represents a different number of "lake years". We can think of older samples as representing  
 428 some average of many years of sediment deposition, whilst younger samples are representa-  
 429 tive of fewer of these lake years. The average of a larger set of numbers is estimated more  
 430 precisely than the average of a smaller set, all things equal. A direct result of this variable av-  
 431 eraging of lake years is that some samples are more precise and therefore have lower variance  
 432 than other samples and yet the model assumed that the variance was constant across samples.

433 Accounting for heteroscedasticity within the model is relatively simple via the use of observa-  
 434 tional weights. The number of lake years represented by each slice is estimated by assigning a

date to the top and bottom of each sediment slice. The variance of each observation should be proportional to the inverse of the number of lake years each sample represents. In the `gam()` function used here, weights should be specified as the number of lake years each sample represents. Other software may require the weights to be specified in a different way.

A secondary problem is the size of the basis dimension used for the time variable. The main user selectable option when fitting a GAM in the penalised likelihood framework of Wood (2004) is how many basis functions to use. As described above, the basis should be large enough to contain the true, but unknown, function or a close approximation to it. For GCV selection the basis used contained 29 basis functions, whilst the CAR(1) model with REML smoothness selection would only converge with a basis containing 20 functions. The size of the basis appears to be sufficient for GCV smoothness selection, but following Wood (2011) REML smoothness selection is preferred. Using the test of Pya and Wood (2016), the basis dimension for the models with REML smoothness selection was too small. To proceed therefore, we must drop the CAR(1) term and increase the basis dimension to 39 functions (by setting `k = 40`; I set it larger than expected because the larger basis contains a richer family of functions and the excess complexity is reduced because of the smoothness penalty.)

With the larger basis dimension and accounting for the non-constant variance of the observations via weights, the model fitted using REML is indistinguishable from that obtained using GCV (Figure 5b). The trace of the REML score for this model shows a pronounced minimum at a much smaller value of  $\lambda$  than the original REML fit (Figure 6), indicating that a more wiggly trend provides a better fit to the Braya-Sø time series. This example illustrates that some care and understanding of the underlying principles of GAMs is required to diagnose potential issues with the estimated model. After standard modelling choices (size of basis to use, correct selection of response distribution and link function, etc.) are checked, it often pays to think carefully about the properties of the data and ensure that the assumptions of the model are met. Here, despite increasing the basis size, it was the failure to appreciate the magnitude of the effect on fitting of the non-constant variance that lead to the initially poor fit and the problems associated with the estimation of the CAR(1) process. I return to the issue of why the GAM plus CAR(1) model encountered problems during fitting later (see [Identifiability](#)).

### 4.3 Confidence intervals and uncertainty estimation

If we want to ask whether either of the estimated trends is statistically interesting or proceed to identifying periods of significant change, we must address the issue of uncertainty in the estimated model. What uncertainty is associated with the trend estimates? One way to visualise this is through a  $1 - \alpha$  confidence interval around the fitted trend, where  $\alpha$  is say 0.05 leading to a 95% interval. A 95% interval would be drawn at  $\hat{y}_t \pm (m_{1-\alpha} \times \text{SE}(\hat{y}_t))$ , with  $m_{1-\alpha} = 1.96$ , the 0.95 probability quantile of a standard normal distribution, and  $\text{SE}(\hat{y}_t)$  is the standard error of the estimated trend at time  $x_t$ . This type of confidence interval would normally be described as *pointwise*; the coverage properties of the interval being correct for a single point on the fitted trend, but, if we were to consider additional points on the trend, the coverage would logically be lower than  $1 - \alpha$ . This is the traditional frequentist interpretation of a confidence interval. However, the GAM described here has a Bayesian interpretation (Kimeldorf and Wahba, 1970;

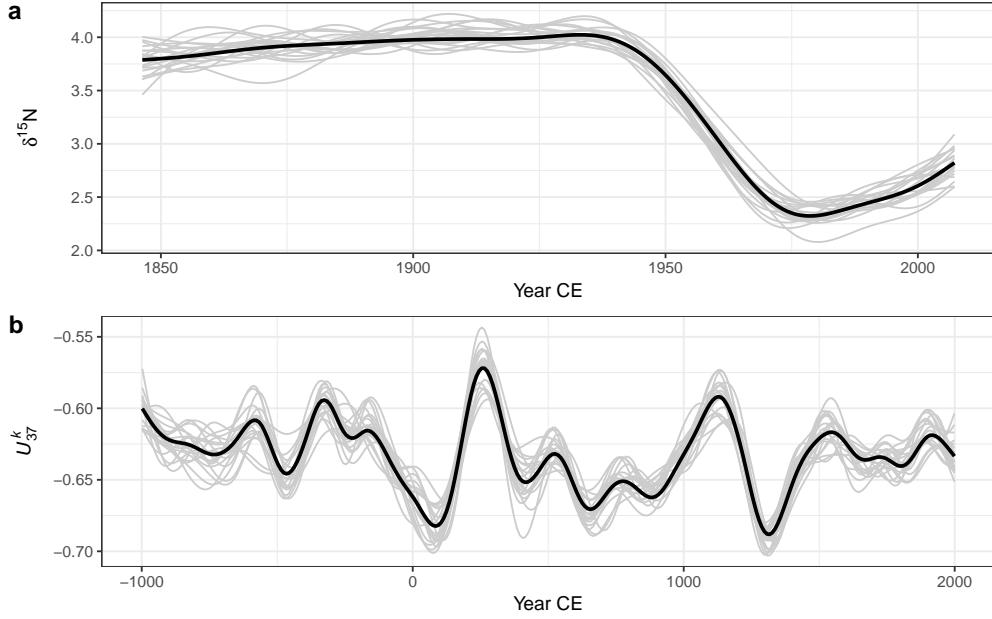


Figure 7: Estimated trends (thick black lines) and 20 random draws (grey lines) from the posterior distribution of the GAM fitted to the Small Water  $\delta^{15}\text{N}$  (a) and Braya-Sø  $U_{37}^K$  (b) time series.

476 Silverman, 1985; Wahba, 1983, 1990) and therefore the typical frequentist interpretation does  
 477 not apply. Nychka (1988) investigated the properties of a confidence interval created as de-  
 478 scribed above using standard errors derived from the Bayesian posterior covariance matrix for  
 479 the estimated mode parameters. Such intervals have the interesting property that they have  
 480 good *across-the-function* coverage when considered from a frequentist perspective. This means  
 481 that, when averaged over the range of the function, the Bayesian credible intervals shown in  
 482 Figure 5 have close to the expected 95% coverage. However, to achieve this some parts of the  
 483 function may have more or less than 95%-coverage. Marra and Wood (2012) recently explained  
 484 Nychka's (1988) surprising results and extended them to the case of generalized models (non-  
 485 Gaussian responses).

486 Whilst the *across-the-function* frequentist interpretation of the Bayesian credible intervals is use-  
 487 ful, it may be important to have an interval that contains the entirety of the true function with  
 488 some state probability  $(1 - \alpha)$ . Such an interval is known as a *simultaneous* interval. A  $(1 - \alpha)100\%$   
 489 simultaneous confidence interval contains *in their entirety*  $1 - \alpha$  of all random draws from the  
 490 posterior distribution of the fitted model.

491 Fitting a GAM involves finding estimates for coefficients of the basis functions. Together, these  
 492 coefficients are distributed multivariate normal with mean vector and covariance matrix spec-  
 493 ified by the model estimates of the coefficients and their covariances respectively. Random  
 494 draws from this distribution can be taken, where each random draw represents a new trend  
 495 that is consistent with the fitted trend but also reflects the uncertainty in the estimated trend.  
 496 This process is known as *posterior simulation*.

497 Figure 7 shows 20 random draws from the posterior distribution of the GAMs fitted to the

498 Small Water and Braya-Sø data sets. In the early period of the  $\delta^{15}\text{N}$  time series many of the  
499 posterior simulations exhibit short periods of increasing and decreasing trend, balancing out  
500 to the relatively flat trend estimated by the GAM (Fig. 7a). Reflecting this uncertainty, we  
501 might expect relatively wide simultaneous intervals during this period in order to contain the  
502 vast majority of the simulated trends. Conversely, the decreasing  $\delta^{15}\text{N}$  trend starting at  $\sim$ 1945  
503 is consistently reproduced in the posterior simulations, suggesting that this feature of the time  
504 series is both real and statistically significant, and that the rate of change in  $\delta^{15}\text{N}$  is relatively  
505 precisely estimated. We see a similar pattern in Figure 7b for the Braya-Sø record; the large  
506 peak in  $U_{37}^K$  at  $\sim$ 250CE and the strong decline at  $\sim$ 1200CE are well defined in the posterior  
507 simulations, whereas most of the localised trends that are smaller magnitude changes in  $y_t$   
508 are associated with posterior simulations that are less well constrained with the ends of the  
509 record in particular showing considerable variation in the strength, timing, and even sign of  
510 simulated trends, reflecting the greater uncertainty in estimated trend during these periods.  
511 For the random draws illustrated in Figure 7, a  $(1 - \alpha)100\%$  simultaneous interval should con-  
512 tain the entire function for on average 19 of the 20 draws.

513 There are a number of ways in which a simultaneous interval can be computed. Here I follow  
514 the simulation approach described by Ruppert et al. (2003) and present only the basic detail; a  
515 fuller description is contained in Appendix 1. The general idea is that if we want to create an  
516 interval that contains the whole of the true function with  $1 - \alpha$  probability, we need to increase  
517 the standard Bayesian credible interval by some amount. We could simulate a large number  
518 of functions from the posterior distribution of the model and then search for the value of  $m_{1-\alpha}$   
519 that when multiplied by  $\text{SE}(\hat{f}(x_t))$  yielded an interval that contained the whole function for  
520  $(1 - \alpha) 100\%$  of the functions simulated. In practice, the simulation method of Ruppert et al.  
521 (2003) does not involve a direct search, but yields the critical value  $m_{1-\alpha}$  required.

522 Simultaneous intervals computed using the method described are show in Figure 8 alongside  
523 the *across-the-function* confidence intervals for the trends fitted to both example data sets. As  
524 expected, the simultaneous interval is somewhat wider than the *across-the-function* interval.  
525 The critical value  $m_{1-\alpha}$  for the simultaneous interval of the estimated trend in  $\delta^{15}\text{N}$  is 3.08,  
526 whilst the same value for the  $U_{37}^K$  series is 3.42, leading to intervals that are approximately  
527  $\pm 50\%$  and  $\pm 75\%$  wider than the equivalent across-the-function intervals.

#### 528 4.4 Identifying periods change

529 In the simple linear trend model (1) whether the estimated trend constitutes evidence for or  
530 against a null hypothesis of no change rests on how large the estimated rate of change in  $y_t$   
531 is ( $\hat{\beta}_1$ ) relative to its uncertainty. This is summarised in the  $t$  statistic. As the rate of change  
532 in  $y_t$  is constant over the fitted trend — there is only a singe slope for the fitted trend  $\hat{\beta}_1$  — if  
533 the  $t$  statistic of the test that  $\hat{\beta}_1 = 0$  is unusually extreme this would be evidence against the  
534 null hypothesis of no change. Importantly, this applies to the whole time series as the linear  
535 model implies a constant rate of change throughout. More formally, the estimate  $\hat{\beta}_1$  is the first  
536 derivative of the fitted trend.

537 In the GAM, the fitted trend need not be linear; the slope of the trend is potentially different  
538 at every point in the time series. As such we might reasonably ask *where* in the series the

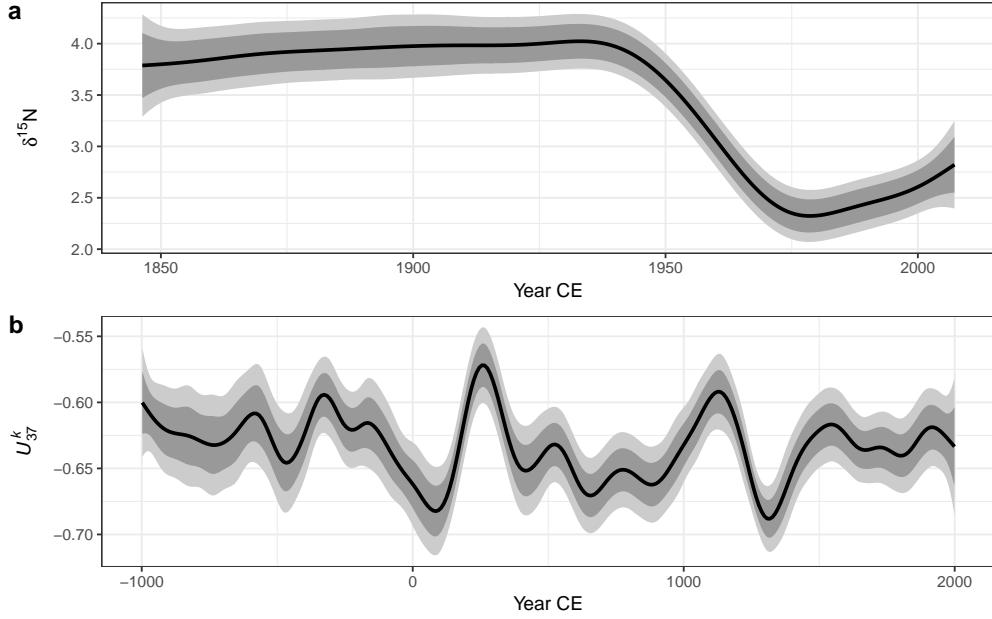


Figure 8: 95% simultaneous confidence intervals (light grey bands) and across-the-function confidence intervals (dark grey bands) on the estimated trends (black lines) for the Small Water  $\delta^{15}\text{N}$  (a) and Braya-Sø  $U_{37}^K$  (b) time series.

539 response  $y_t$  is changing, if at all? Mirroring the linear model we can answer this question by  
 540 determining whether or not the first derivative at any time point  $x_t$  of the fitted trend at any  
 541 time point is consistent with a null hypothesis of no change. We want to know whether or not  
 542 the first derivative is indistinguishable from a value of 0 — no trend — given the uncertainty  
 543 in the estimate of the derivative.

544 Derivatives of the fitted spline are not easily available analytically, but they can be estimated  
 545 using the method of finite differences. Two values of the estimated trend, separated by a very  
 546 small time-shift ( $\Delta_t$ ), are predicted from the model; the difference between the estimated values  
 547 for the two time points is an approximation of the true first derivative of the trend. As  $\Delta_t \rightarrow 0$   
 548 the approximation becomes increasingly accurate. In practice, the first derivative of the fitted  
 549 trend is evaluated using finite differences at a large number of points in the time series. An  
 550 approximate  $(1 - \alpha)100\%$  pointwise confidence interval can be calculated for the derivative  
 551 estimates using standard theory (i.e.  $\pm 1.96 \times \text{SE}(\hat{y}_t)$  for a 85% interval) and the covariance matrix  
 552 of the spline coefficients. A  $(1 - \alpha)100\%$  simultaneous interval for the derivatives can also be  
 553 computed using the method described [earlier](#). Periods of significant change are identified as  
 554 those time points where the (simultaneous) confidence interval on the first derivative does not  
 555 include zero.

556 Figure 9 shows the estimated first derivative of the fitted trend in the Small Water (9a) and  
 557 Braya-Sø (9b) time series. Although the estimated trend suggests a slight increase in  $\delta^{15}\text{N}$   
 558 from the start of the record to  $\sim 1940$ , the estimated trend is sufficiently uncertain that the si-  
 559 multaneous interval on the first derivative includes 0 throughout. We can understand why this  
 560 is so by looking at the posterior simulations in Figure 7a; there is considerable variation in the

shape of the simulated trends throughout this period. From ~1925 the derivative of the trend becomes negative, however it is not until ~1940 that the simultaneous interval doesn't include 0. At this point we have evidence to reject the null hypothesis of no change. This time point may be taken as the first evidence for change in  $\delta^{15}\text{N}$  in the Small Water core. The simultaneous interval on the first derivative of the trend in  $\delta^{15}\text{N}$  is bounded away from 0 between ~1940 and ~1975, covering the major decline in values evident in the observations. The simultaneous interval includes 0 from ~1975 onward, suggesting that, whilst quite pronounced, the recent increase in  $\delta^{15}\text{N}$  is not statistically significant. To determine whether or not the recent increase is real, we would require considerably more samples with which to (hopefully) more-precisely estimate the trend during this period. Alternatively, we might just have to wait until sufficient additional sedimentation has occurred to warrant recore Small Water and reestimating the trend in  $\delta^{15}\text{N}$ .

The estimated trend at Braya-Sø exhibited a number of oscillations in  $U_{37}^K$ . As we saw previously in Figures 7b and 8b, many of these are subject to significant uncertainty and it is important therefore to discern which, if any, of the oscillations in the response can be identified given the model uncertainty. In Figure 9b only two features of the estimated trend are considered significant based on the derivatives of the smooth; one centred on ~250CE and a second at ~1150CE. In both these periods, the simultaneous interval for the first derivative of the trend does not include zero. In the first case we detect the large peak and subsequent decline in  $U_{37}^K$  at ~250CE, whilst at ~1150CE the large trough is identified, but not the increasing trend immediately prior to this excursion to lower  $U_{37}^K$ . Recall that these intervals are simultaneous in nature, strongly guarding against false positives, and as such we can be confident in the estimation of these two features, whilst care must be taken to not over-interpret the remaining variations in the estimated trend.

## 4.5 Residual autocorrelation and model identification

The GAM fitted to the  $\delta^{15}\text{N}$  time series contained a CAR(1) process to model residual temporal autocorrelation in the residuals. The estimated magnitude of the autocorrelation is given by the parameter  $\phi$ . The estimated value of  $\phi$  for the  $\delta^{15}\text{N}$  series is 0.6 with 95% confidence interval 0.28–0.85, indicating moderate to strong residual autocorrelation about the fitted trend. The correlation function is an exponentially decreasing function of temporal separation ( $\Delta_t$ ), and whilst observations that are a few years apart are quite strongly dependent on one another, this dependence drops off rapidly as  $\Delta_t$  increases and is effectively zero when samples are separated by a decade or more (Figure 10).

Failure to account for the dependencies in the  $\delta^{15}\text{N}$  time series could lead to the estimation of a more wiggly trend than the one shown in Figure 5a which would negatively impact the confidence placed on the inferences we might draw from the fitted model. Importantly, failing to account for the strong dependency in the residuals would lead to smaller uncertainties in the estimated spline coefficients, which would propagate through to narrower confidence intervals on the fitted trend and on the first derivatives, and ultimately to the identification of significant periods of change. The end result would be a tendency toward anti-conservative identification of periods of change; the coverage probability would be lower than the antici-

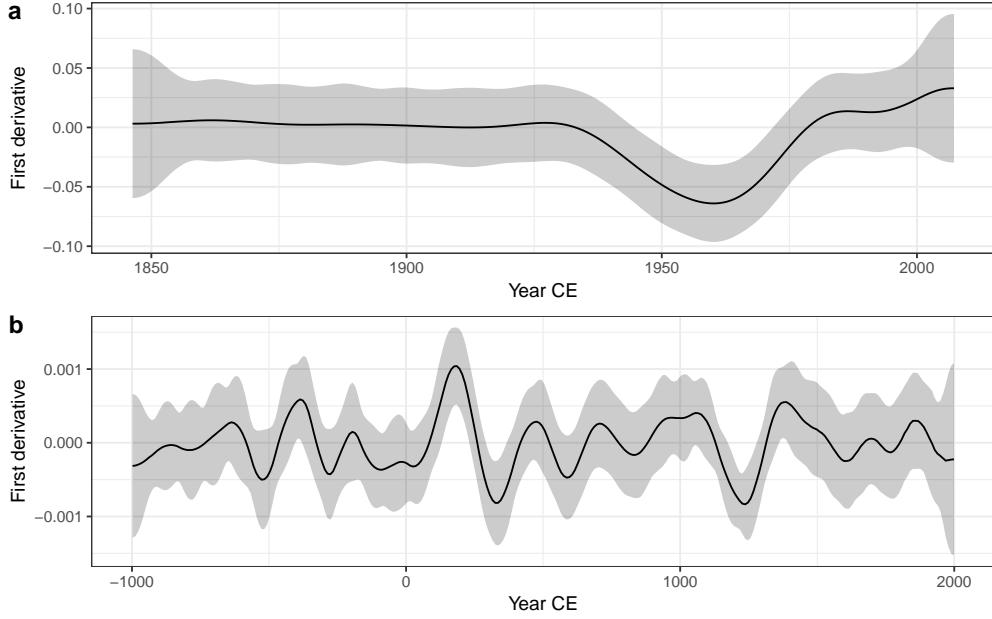


Figure 9: Estimated first derivatives (black lines) and 95% simultaneous confidence intervals of the GAM trends fitted to the Small Water  $\delta^{15}\text{N}$  (a) and Braya-Sø  $U_{37}^K$  (b) time series. Where the simultaneous interval does not include 0, the models detect significant temporal change in the response.

- 602 pated  $1 - \alpha$ , leading to a greater chance of false positive results.
- 603 Problems estimating the GAM plus CAR(1) model were encountered when this was fitted to  
604 the  $U_{37}^K$  time series; including both a smooth trend in the mean  $U_{37}^K$  and a CAR(1) process in  
605 the residuals lead to an unidentifiable model. What makes a model with a spline-based trend  
606 and an autocorrelation process like the CAR(1) potentially unidentifiable?
- 607 Consider again the basic GAM for a smooth trend, (3). In that equation the correlation ma-  
608 trix  $\Lambda$  was omitted for the sake of simplicity. As I did in (6), I reintroduce it and restate the  
609 distributional assumptions of this model

$$y_t = \beta_0 + f(x_t) + \varepsilon_t, \quad \varepsilon \sim (0, \Lambda \sigma^2) \quad (7)$$

- 610 In the basic GAM,  $\Lambda \equiv \mathbf{I}$  is an identity matrix, a matrix with 1s on the diagonal and 0s elsewhere

$$\begin{bmatrix} 1 & 0 & 0 & \dots & 0 \\ 0 & 1 & 0 & \dots & 0 \\ 0 & 0 & 1 & \dots & 0 \\ \vdots & \vdots & \vdots & \ddots & \vdots \\ 0 & 0 & 0 & \dots & 1 \end{bmatrix},$$

- 611 which is where the independence assumption of the model comes from; a model residual is  
612 perfectly correlated with itself (the 1s on the diagonal), but uncorrelated with any other resid-  
613 ual (the off-diagonal 0s). In the GAM plus CAR(1) model, an alternative correlation function

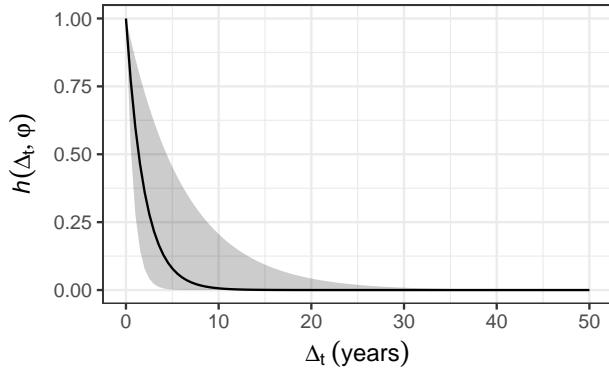


Figure 10: Estimated CAR(1) process from the GAM fitted to the Small Water  $\delta^{15}\text{N}$  time series.  $h(\Delta_t, \phi)$  is the correlation between residuals separated by  $\Delta_t$  years, where  $\hat{\phi} = 0.6$ . The shaded band is a 95% pointwise confidence interval on the estimated correlation  $h$ .

- for  $\Lambda$  was used — the CAR(1) with correlation parameter  $\phi$ . Fahrmeir and Kneib (2008) show that where the stochastic structure of  $f$  and  $\Lambda$  approach one another, i.e. where we have a potentially wiggly trend or strong autocorrelation as  $\phi \rightarrow 1$ , the two processes can quickly become unidentifiable (see also Fahrmeir et al., 2013). By unidentifiable, we mean that it becomes increasingly difficult to distinguish between a wiggly trend or strong autocorrelation because these two processes are very similar to one another in appearance. This leads to model estimation problems of the sort encountered with fitting the GAM plus CAR(1) model to the Braya-sø  $U_{37}^K$  series.
- Why might this be so? Autocorrelation is the tendency for a large (small) value of  $y_t$  at time  $x_t$  to be followed by a likewise large (small) value at time  $x_{t+1}$ . This leads to runs of values that are consistently greater (less) than the overall mean. Short runs would indicate weaker autocorrelation whilst longer runs are associated with stronger autocorrelation, and long runs of values greater (less) than the mean would be evident as non-linear trends in the time series. As a result, a wiggly trend and an autocorrelation function with large  $\phi$  are two ways to describe the same pattern of values in a time series, and without any further information to constrain either the model is unable to distinguish both components uniquely.
- Situations where it may be possible to uniquely identify separate wiggly trends and autocorrelation are exemplified by the Small Water  $\delta^{15}\text{N}$  time series. The non-linear trend and the autocorrelation operate at very different scales; the trend represents decadal-scale variation in mean  $\delta^{15}\text{N}$ , whilst the CAR(1) process represents the much smaller-scale tendency for values of the response to be followed in time by similar values. That such a pattern is observed in the Small Water core is the result of the high resolution of the sampling in time relative to the long-term trend. In contrast, the Braya-Sø record is sampled at far lower resolution relative to the fluctuations in the mean response, and consequently the data do not contain sufficient information to separate trend and autocorrelation.

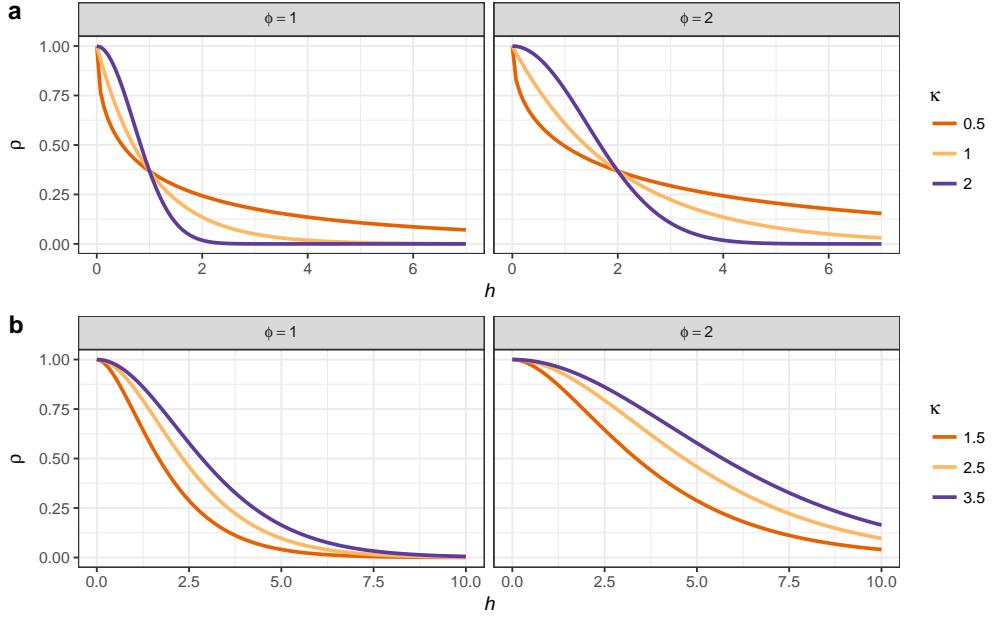


Figure 11: Power exponential (a) and Matérn (b) correlation functions for observation separation distance  $h$ . Two values of the effective range parameter ( $\phi$ ) are shown for each function. For the power exponential function,  $\kappa$  is the power in the power exponential function. For the Matérn correlation function,  $\kappa$  distinguishes the member of the Matérn family.

## 639 4.6 Gaussian process smooths

640 In the world of machine learning, Gaussian processes (Golding and Purse, 2016; Rasmussen  
 641 and Williams, 2006) are a widely-used method for fitting smooth non-parametric regression  
 642 models. A Gaussian process is a distribution over all possible smooth functions  $f(x)$ . In the  
 643 field of spatial statistics, Gaussian processes are known by name *kriging*.

644 With a Gaussian process we are interested in fitting a smooth temporal trend by modelling  
 645 the way the correlation between pairs of observations varies as a function of the distance,  $h$ , in  
 646 time that separates the observations. The correlation between pairs of observations decreases  
 647 with increasing separation, which is modelled using a correlation function,  $c(h)$ .

648 Several functions can be used to represent  $c(h)$ . Two common ones are the power exponential  
 649 function and the Matérn family of correlation functions. The power exponential function at  
 650 separation distance  $h$  is

$$c(h) = \exp\{(-h/\phi)^\kappa\}$$

651 where  $0 < \kappa \leq 2$ . The Matérn correlation function is actually a family of functions with closed-  
 652 forms only available for a subset of the family, distinguished by  $\kappa$ . When  $\kappa = 1.5$ , the Matérn  
 653 correlation function is

$$c(h) = (1 + h/\phi) \exp(-h/\phi)$$

654 whilst for  $\kappa = 2.5$  it is

$$c(h) = \{1 + h/\phi + (h/\phi)^2/3\} \exp(-h/\phi)$$

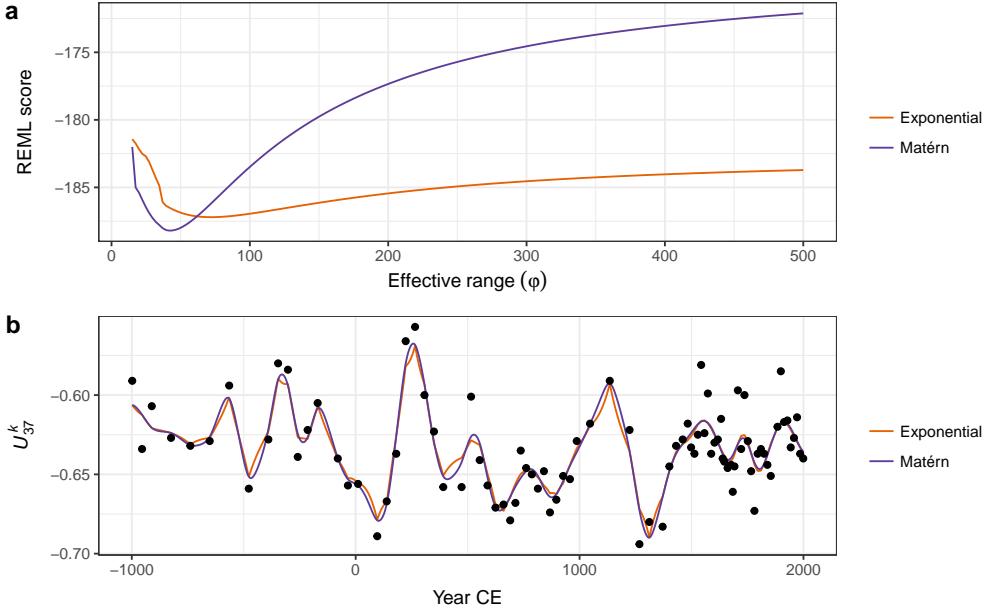


Figure 12: Gaussian process smooths fitted to the  $U_{37}^K$  time series. REML score traces for GAMs fitted using power exponential ( $\kappa = 1$ ) or Matérn ( $\kappa = 1.5$ ) correlation functions as a function of the effective range parameter ( $\phi$ ) are shown (a). The optimal model for each function is that with the lowest REML score. b) shows the resulting trends estimated using the respective correlation function with the value of  $\phi$  set to the optimal value.

and for  $\kappa = 3.5$

$$c(h) = \{1 + h/\phi + 2(h/\phi)^2/5 + (h/\phi)^3/15\} \exp(-h/\phi).$$

In all cases,  $\phi$  is the effective range parameter, which sets the distance beyond which the correlation function is effectively zero.

Figure 11 shows examples of two different correlation functions; the *power exponential* (Figure 11a), and the Matérn (Figure 11b) correlation functions. These functions are smooth and monotonic-decreasing, meaning that the value of the correlation function decreases with increasing separation ( $h$ ). When  $h = 0$ , the correlation is equal to 1 ( $c(0) = 1$ ); two samples taken at exactly the same time point are perfectly correlated. As  $h \rightarrow \infty$ , the correlation tends to zero ( $c(h) \rightarrow 0$ ); two samples separated by a large amount of time tend to be uncorrelated. Often we are interested in learning how large the separation in time needs to be before, on average, a pair of observations is effectively uncorrelated (i.e. where  $c(h)$  is sufficiently close to zero).

Gaussian processes and GAMs share many similarities and we can fit a Gaussian process using the techniques already described above for splines (Handcock et al., 1994; Kammann and Wand, 2003). It can be shown (e.g. Fahrmeir et al., 2013) that the Gaussian process model has the same penalised likelihood form as the GAM that we discussed earlier; the observations are the knots of the smoother and each has a basis function in the form of a correlation function. The equivalence is only true if the basis functions do not depend on any other parameters of the model, which is only achievable if the value of  $\phi$  is fixed and known (Fahrmeir et al., 2013). In general, however, we would like to estimate  $\phi$  as part of model fitting. To achieve this we

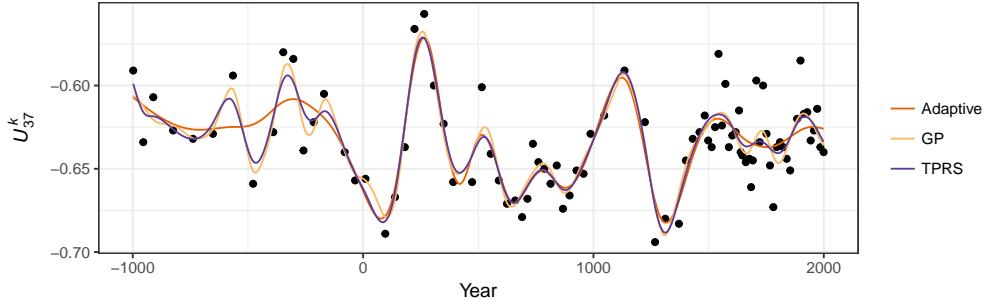


Figure 13: Comparison of trends estimated using i) adaptive smooth, ii) Gaussian process, and iii) thin plate regression spline bases for the  $U_{37}^K$  time series.

can maximise the profile likelihood or score statistic of the model over a range of values of  $\phi$  (Wood, 2017, 362–363). This involves proposing a value of  $\phi$  for the effective range of the correlation function and then estimating the resulting GAM by minimising the penalised log-likelihood conditional upon this value of  $\phi$  and repeating for a range of values for  $\phi$ . The model, and its corresponding value of  $\phi$ , with lowest penalised log-likelihood or score statistic is then retained as the estimated GAM. Figure 12a shows the REML score for models estimated using a Gaussian process smooth with a Matérn correlation function ( $\kappa = 1.5$ ) for a sequence of values of  $\phi$  between 15 and 1000 years. There is a clear minimum around 40 years separation, with the minimum REML score being observed at  $\phi = 41.81$ ). Also shown are the REML scores for models using the power exponential function ( $\kappa = 1$ ) with the minimum score observed at a somewhat higher effective range of  $\phi = 71.06$ .

Figure 12b shows the estimated trends for the  $U_{37}^K$  time series using Gaussian process smooths with exponential and Matérn correlations functions, both using  $\phi$  values at their respective optimal value as assessed using the REML score. The estimated trends are very similar to one another, although there is a noticeable difference in behaviour, with the power exponential ( $\kappa = 1$ ) version being noticeably less-smooth than the Matérn version. This difference is attributable to the shapes of the respective correlation functions; the Matérn approaches a correlation of 1 smoothly as  $h$  approaches 0, whilst the power exponential with  $\kappa = 1$  approaches a correlation of 1 increasingly quickly with decreasing  $h$ . The power exponential with  $\kappa = 2$ , like the Matérn, approaches  $\phi = 1$  smoothly, and consequently the trend estimated using this correlation function is qualitatively similar to that estimated using the Matérn correlation function.

## 4.7 Adaptive smoothing

Each of the spline types that I have discussed so far share a common feature; the degree of wiggliness over the time series is fixed due to the use of a single smoothness parameter,  $\lambda$ . The definition of wiggliness, as the integrated squared second derivative of the spline, ensures that the fitted smoother does not jump about wildly. This assumes that the data themselves are well described by a smoothly varying trend. If we anticipate abrupt change or step-like responses to environmental forcing this underlying assumption of the GAM would suggest that the method

- 703 is ill-suited to modelling palaeo time series in which such features are evident or expected.
- 704 While there is not much we can do within the GAM framework to model a series that contains  
 705 both smooth trends and step-like responses, an adaptive smoother can help address problems  
 706 where the time series consists of periods of rapid change in the mean combined with periods  
 707 of complacency or relatively little change. As suggested by their name, adaptive smoothers  
 708 can adjust to changes in the wigginess of the time series. This adaptive behaviour is achieved  
 709 by making the smoothness parameter  $\lambda$  itself depend smoothly on  $x_t$  (Ruppert et al., 2003,  
 710 17; Wood, 2017, 5.3.5); in other words, the adaptive smoother allows the wigginess of the  
 711 estimated trend to vary smoothly over time. Whilst this allows the estimated trend to adapt  
 712 to periods of rapid change in the response, adaptive smoothers make significant demands on  
 713 the data (Wood, 2017, 5.3.5); if we used  $m$  smoothness penalties to allow the wigginess to  
 714 vary over a time series, it would be like estimating  $m$  separate smooths from chunks of the  
 715 original series each of length  $n/m$ . In a practical sense, this limits the use of adaptive splines  
 716 in palaeoecology to proxies that are readily enumerated, such as the biogeochemical proxies  
 717 used in the two example data sets.
- 718 Figure 13 compares trends for the Braya-Sø time series estimated using GAMs with the three  
 719 main types of spline discussed; i) TPRS, ii) Gaussian process smooths, and iii) an adaptive  
 720 smoother using 45 basis functions and 5 smoothing parameters. There is a clear difference  
 721 in the behaviour of the adaptive and non-adaptive smoothers for the first 1000 years of the  
 722 record, with the adaptive smooth exhibiting much less variation compared with either the  
 723 TPRS or Gaussian process splines. Over the remaining two thirds of the series, there is much  
 724 closer agreement in the three smooths.
- 725 The behaviour of the TPRS and Gaussian process splines for these data is the result of requiring  
 726 a large amount of wigginess (a small  $\lambda$ ) to adapt to the large oscillations in  $U_{37}^K$  present around  
 727 year 250CE and again at ~900–1500CE. This large degree of wigginess allows the splines to  
 728 potentially over-fit individual data points much earlier in the record. Because the adaptive  
 729 smoother, in contrast, can adapt to these periods of rapid change in the response it is much  
 730 less susceptible to this “chasing” behaviour — we don’t need to waste effective degrees of  
 731 freedom in periods with little or no change just to be able to fit the data well when there is a  
 732 lot of change.
- 733 This potential for over-fitting in such situations is undesirable, yet if we recall Figure 9 and  
 734 the discussion around the use of the first derivative to identify periods of significant change,  
 735 we would not interpret the oscillations in the early part of the  $U_{37}^K$  record as being statistically  
 736 significant. Owing to the paucity of data in this part of the series the trends fitted using the  
 737 non-adaptive smoothers are subject to such a large degree of uncertainty that the alternative  
 738 of no trend through the first 1000 years of the record is also a plausible explanation of the data.  
 739 The trend estimated using the adaptive smooth reflects this. Therefore, should we conclude  
 740 that there is no trend in  $U_{37}^K$  and thence climate in this period? I believe that to be too-strong  
 741 a statement; those oscillations in  $U_{37}^K$  may be real responses to climate forcing but may simply  
 742 lack the statistical power to distinguish them from the null hypothesis of no trend through this  
 743 period. The adaptive smoother is only adjusting to the data available to it; just because it does  
 744 not detect a trend during this period does not lend itself to an interpretation of stable climate  
 745 forcing or complacency in the lake’s response to forcing. If there were particular interest in the

<sup>746</sup> climate of this particular period we might take from the Braya-Sø record that there is potential  
<sup>747</sup> early variation in climate forcing, but that additional data from this or other sites is required  
<sup>748</sup> before any definitive conclusion can be drawn.

## <sup>749</sup> 4.8 Accounting for age model uncertainty

<sup>750</sup> Thus far, the trend models that I have described and illustrated assumed that the time covariate  
<sup>751</sup> ( $x_t$ ) was fixed and known. In both examples, and more generally for most palaeoecological  
<sup>752</sup> records, this assumption is violated. Unless the record is annually laminated, assigning an  
<sup>753</sup> age to a sediment interval requires the development of an age model from observations of the  
<sup>754</sup> relationship between depth down the sediment core and estimates of the age of the sample  
<sup>755</sup> arrived at using any of a number of techniques, for example  $^{210}\text{Pb}$  or  $^{14}\text{C}$  radiometric dating.  
<sup>756</sup> This age-depth relationship is itself uncertain, usually being derived from a mathematical or  
<sup>757</sup> statistical model applied to point age estimates (e.g. Blaauw and Heegaard, 2012). Incorporating  
<sup>758</sup> this additional component of uncertainty complicates the estimation of statistical models  
<sup>759</sup> from palaeoenvironmental data. In this section I illustrate a simulation based approach to  
<sup>760</sup> quantify and account for age-model uncertainty as part of the trend estimation using a GAM  
<sup>761</sup> (see Anchukaitis and Tierney (2013) for a similar, non-GAM related idea).

<sup>762</sup> Figure 14a shows the estimated dates (in Years CE) for 12 levels in the Small Water core dated  
<sup>763</sup> using  $^{210}\text{Pb}$ . The vertical bars show the estimated age uncertainty of each level. The solid line  
<sup>764</sup> through the data points is an additive model fitted to the observations, with prior weights  
<sup>765</sup> given by the estimated age uncertainties. The fitted age-depth model is constrained to be  
<sup>766</sup> monotonically decreasing with increasing depth, following the method of (Pya and Wood,  
<sup>767</sup> 2015) using the *scam* package (Pya, 2017). Also shown are 25 simulations from the posterior  
<sup>768</sup> distribution of the monotonically-constrained GAM. Each simulation from the posterior dis-  
<sup>769</sup> tribution of the age-model is itself a potential age-depth model, which can be used to assign  
<sup>770</sup> dates to the Small Water core. The trend model in (4) can be fitted to the  $\delta^{15}\text{N}$  data using these  
<sup>771</sup> new dates as  $x_t$ , and the whole process repeated for a large number of simulations from the  
<sup>772</sup> age model.

<sup>773</sup> Figure 14b shows the trend in  $\delta^{15}\text{N}$  for the observed age-depth model, plus trends estimated  
<sup>774</sup> via the same model using 100 draws from the posterior distribution of the age model. In this  
<sup>775</sup> case, the age-depth model is relatively simple with little variation in the posterior draws, re-  
<sup>776</sup> sulting in trends that match closely that obtained from the estimated age-depth relationship.  
<sup>777</sup> Even so, this additional uncertainty suggests that the timing of the decline in  $\delta^{15}\text{N}$  covers the  
<sup>778</sup> interval ~1935–1945.

<sup>779</sup> The uncertainty in the trend estimates illustrated in Figure 14b only reflects the variation in  
<sup>780</sup> trends fitted to the uncertain dates of the sediment samples. To fully visualise the uncertainty  
<sup>781</sup> in the trend estimates, incorporating both age model uncertainty *and* uncertainty in the esti-  
<sup>782</sup> mated model coefficients themselves, 50 simulations from the posterior distribution of each  
<sup>783</sup> of the 100 estimated trends shown in Figure 14b were performed, resulting in 5,000 trend esti-  
<sup>784</sup> mates for the  $\delta^{15}\text{N}$  series. These are shown in Figure 14c, where the two obvious changes over  
<sup>785</sup> the same simulations without accounting for uncertainty in  $x_t$  (Figure 7a) are that the uncer-  
<sup>786</sup> tainty band traced out by the simulations is approximately 50% wider and, not surprisingly,

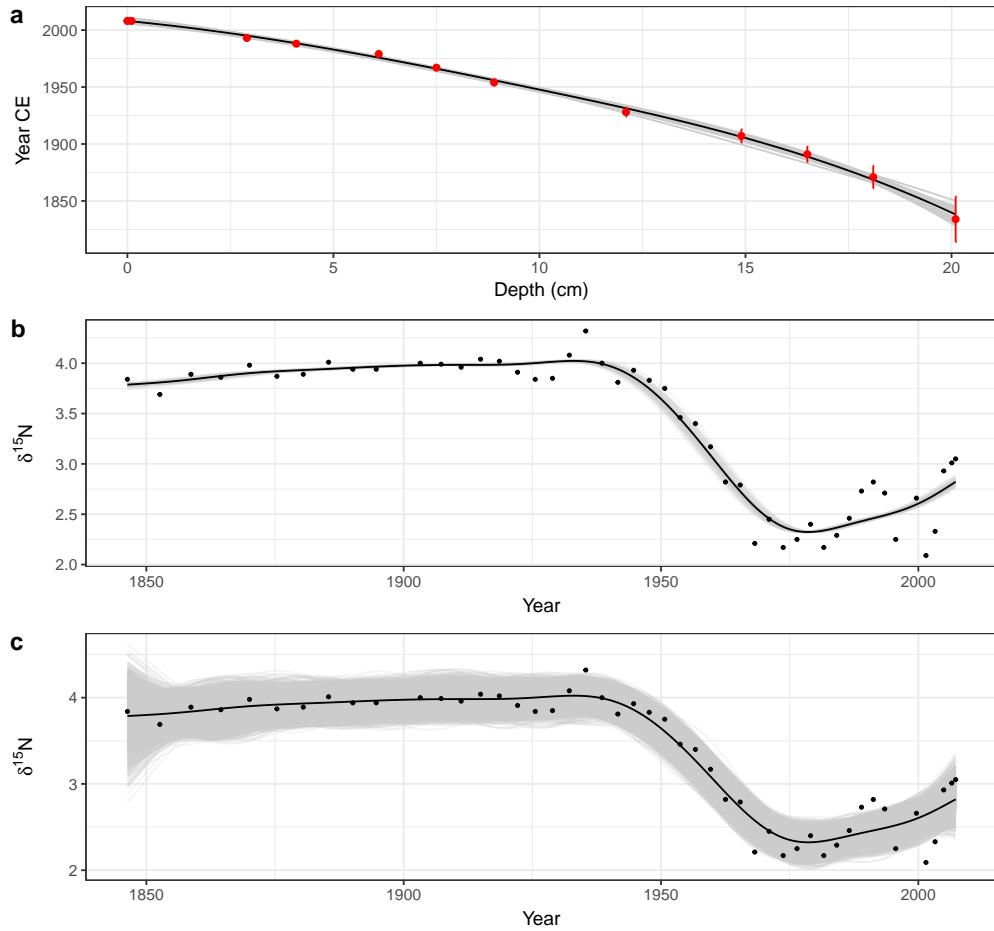


Figure 14: Accounting for uncertainty in age estimates whilst fitting a smooth trend to the Small Water  $\delta^{15}\text{N}$  time series. (a) Estimated age model using a monotonically-constrained spline fitted to  $^{210}\text{Pb}$  inferred ages for selected depths in the sediment core (red points). The uncertainty in the  $^{210}\text{Pb}$  inferred age is show by the red vertical bars. The fitted age model is illustrated by the solid black line. The faint grey lines are 25 random draws from the posterior distribution of the monotonically constrained GAM. The effect of age uncertainty on trend estimation is shown in b); for 100 simulations from the posterior distribution of the age model in a) a trend was estimated using a GAM with a thin plate regression spline basis and a CAR(1) process in the residuals. These trends are shown as grey lines. The combined effect of age model and fitted GAM uncertainty on the trends for the  $\delta^{15}\text{N}$  time series is shown in c). The grey lines in c) are based on 50 random draws from the model posterior distribution for each of the 100 trends shown in b). For both b) and c) the black line shows the trend estimated assuming the ages of each sediment sample are known and fixed.

787 the uncertainty in the estimated trend is most pronounced in the least accurately-dated sec-  
788 tion of the core. Despite this additional uncertainty however, the main result holds; a marked  
789 decline of ~1.5‰ that occurred between approximately 1930 and 1945, with mild evidence of  
790 a small increase in  $\delta^{15}\text{N}$  post 2000 CE.

## 791 4.9 Multivariate data

792 A large proportion of the palaeoenvironmental data generated today is multivariate in nature  
793 and yet the two examples used to illustrate GAMs were univariate. Can the approach de-  
794 scribed here be used for multivariate data? Yes, and no. With one main exception it is not  
795 possible to directly apply the GAM methodology described here to multivariate abundance  
796 data, where the aim is to model all species at once. The *mgcv* software, for example, is not able  
797 to estimate the penalized GAM for multiple non-Gaussian responses. The exception is for a  
798 small number of correlated Gaussian responses; these could be modelled as being distributed  
799 multivariate normal conditional upon the covariates. Such a model would estimate the ex-  
800 pected values of each response and the correlations between them. For example, we could  
801 jointly model  $\delta^{15}\text{N}$  and  $\delta^{13}\text{C}$  series using this approach.

802 Formal multivariate versions of GLM or GAMs are currently an important area of research  
803 within ecology (see Warton et al. (2015) for a recent review), where they go by the name joint  
804 species distribution models (JSDMs). Whilst undoubtedly powerful, our knowledge regard-  
805 ing JSDMs and their availability in software are still in their relative infancy and they require  
806 considerable expertise to implement. As such, JSDMs are currently beyond the reach of most  
807 palaeoecologists. Despite this, we should be watching JSDM research as developments are  
808 ongoing and a degree of method maturation occurring.

809 One currently available avenue for fitting a multivariate GAM is via regularized sandwich  
810 estimators and GLMs (Warton, 2011), which involves fitting separate GLMs (or GAMs) to  
811 each response variable and subsequently using resampling-based hypothesis tests to deter-  
812 mine which covariates are related to variation at the community level and for individual taxa  
813 (Wang et al., 2012; Warton, 2011; Warton et al., 2012). The *mvabund* package (Wang et al., 2012)  
814 implements this approach within R and can use *mgcv* to fit GAMs to each species.

815 A pragmatic although inelegant approach that has been used to estimate trends in multivariate  
816 palaeoecological data is to first summarise the response data using an unconstrained ordina-  
817 tion via a PCA, CA, or principal curve and then fit separate GAM models to the site (sample)  
818 scores of the first few ordination axes or principal curve (Beck et al., 2018; Bennion et al., 2015).  
819 Whilst this two-step approach is relatively easy to implement and builds on approaches that  
820 palaeoecologists already use to summarise multivariate stratigraphic data, it is best thought of  
821 as modelling changes in abundance or relative composition at the community level. It is less  
822 well suited to unpicking taxon-specific trends however, because the ordination step combines  
823 individual species information into latent variables (axes) that are linear combinations of *all*  
824 species and it is these latent variables that are then modelled using GAM.

## 825 5 Conclusions

826 Formal statistical estimation of trends in palaeoenvironmental data has been hampered by  
827 the nature of the data that comprise the time series; the uneven spacing of samples in time  
828 makes it, if not impossible, difficult to fit classical statistical time series models like ARIMA.  
829 This has lead palaeoecologists and palaeolimnologists to fall back on basic statistical methods  
830 such as linear parametric and non-parametric correlations or simple linear regression models,  
831 where the assumptions of the method are often grossly violated by the dependencies inherent  
832 to time series data. GAMs, whilst similar to the popular Loess smoother, provide a superior  
833 alternative approach to trend estimation in palaeoenvironmental time series. GAMs can es-  
834 timate non-linear trends, provide estimates of the magnitude of change as well as allow the  
835 identification of periods of change, can account for the lack of independence (either via auto-  
836 correlation processes or via the fitting of a wiggly trend), and provide a formal framework for  
837 statistical inference on each of these features.

838 In presenting the GAM with specific palaeoenvironmental examples and addressing the issues  
839 that arise in palaeoenvironmental time series, it is hoped that palaeoecologists and palaeolim-  
840 nologists will be motivated to give greater consideration to the estimation of trends and the  
841 identification of change in stratigraphic time series.

## 842 Conflict of interest statement

843 The author declares that the research was conducted in the absence of any commercial or  
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## 854 References

- 855 Anchukaitis, K. J., and Tierney, J. E. (2013). Identifying coherent spatiotemporal modes in time-  
856 uncertain proxy paleoclimate records. *Climate Dynamics* 41, 1291–1306. doi:[10.1007/s00382-013-1940-2](https://doi.org/10.1007/s00382-013-1940-2)

- 857 012-1483-0.
- 858 Beck, K. K., Fletcher, M.-S., Gadd, P. S., Heijnis, H., Saunders, K. M., Simpson, G. L., et al.  
859 (2018). Variance and Rate-of-Change as early warning signals for a critical transition in an  
860 aquatic ecosystem state: A test case from tasmania, australia. *Journal of Geophysical Research: Biogeosciences* 123, 2017JG004135. doi:[10.1002/2017JG004135](https://doi.org/10.1002/2017JG004135).
- 862 Bennion, H., Simpson, G. L., and Goldsmith, B. J. (2015). Assessing degradation and recovery  
863 pathways in lakes impacted by eutrophication using the sediment record. *Frontiers in Ecology  
864 and Evolution* 3. doi:[10.3389/fevo.2015.00094](https://doi.org/10.3389/fevo.2015.00094).
- 865 Bergmeir, C., Hyndman, R. J., and Koo, B. (2018). A note on the validity of cross-validation for  
866 evaluating autoregressive time series prediction. *Computational Statistics & Data Analysis* 120,  
867 70–83. doi:[10.1016/j.csda.2017.11.003](https://doi.org/10.1016/j.csda.2017.11.003).
- 868 Birks, H. J. B. (1998). Numerical tools in palaeolimnology — progress, potentialities, and prob-  
869 lems. *Journal of Paleolimnology* 20, 307–332. doi:[10.1023/A:1008038808690](https://doi.org/10.1023/A:1008038808690).
- 870 Birks, H. J. B. (2012a). “Introduction and overview of part III,” in *Tracking environmental change  
871 using lake sediments* (Springer, Dordrecht), 331–353. doi:[10.1007/978-94-007-2745-8\\\_10](https://doi.org/10.1007/978-94-007-2745-8\_10).
- 872 Birks, H. J. B. (2012b). “Overview of numerical methods in palaeolimnology,” in *Tracking envi-  
873 ronmental change using lake sediments* (Springer, Dordrecht), 19–92. doi:[10.1007/978-94-007-2745-8\\\_2](https://doi.org/10.1007/978-94-007-2745-8\_2).
- 875 Blaauw, M., and Heegaard, E. (2012). “Estimation of Age-Depth relationships,” in *Tracking  
876 environmental change using lake sediments* (Springer, Dordrecht), 379–413. doi:[10.1007/978-94-007-2745-8\\\_12](https://doi.org/10.1007/978-94-007-2745-8\_12).
- 878 Brassell, S. C. (1993). “Applications of biomarkers for delineating marine paleoclimatic fluctu-  
879 ations during the pleistocene,” in *Organic geochemistry: Principles and applications*, eds. M. H. En-  
880 gel and S. A. Macko (Boston, MA: Springer US), 699–738. doi:[10.1007/978-1-4615-2890-6\\\_34](https://doi.org/10.1007/978-1-4615-2890-6\_34).
- 881 Chu, G., Sun, Q., Li, S., Zheng, M., Jia, X., Lu, C., et al. (2005). Long-chain alkenone distribu-  
882 tions and temperature dependence in lacustrine surface sediments from china. *Geochimica et  
883 Cosmochimica Acta* 69, 4985–5003. doi:[10.1016/j.gca.2005.04.008](https://doi.org/10.1016/j.gca.2005.04.008).
- 884 Cleveland, W. S. (1979). Robust locally weighted regression and smoothing scatterplots. *Jour-  
885 nal of the American Statistical Association* 74, 829–836. doi:[10.1080/01621459.1979.10481038](https://doi.org/10.1080/01621459.1979.10481038).
- 886 Craven, P., and Wahba, G. (1978). Smoothing noisy data with spline functions. *Numerische  
887 Mathematik* 31, 377–403. doi:[10.1007/BF01404567](https://doi.org/10.1007/BF01404567).
- 888 Duchon, J. (1977). “Splines minimizing rotation-invariant semi-norms in sobolev spaces,”  
889 in *Constructive theory of functions of several variables* (Springer, Berlin, Heidelberg), 85–100.  
890 doi:[10.1007/BFb0086566](https://doi.org/10.1007/BFb0086566).
- 891 Dutilleul, P., Cumming, B. F., and Lontoc-Roy, M. (2012). “Autocorrelogram and periodogram  
892 analyses of palaeolimnological Temporal-Series from lakes in central and western north amer-  
893 ica to assess shifts in drought conditions,” in *Tracking environmental change using lake sediments*  
894 (Springer, Dordrecht), 523–548. doi:[10.1007/978-94-007-2745-8\\\_16](https://doi.org/10.1007/978-94-007-2745-8\_16).
- 895 D’Andrea, W. J., Huang, Y., Fritz, S. C., and Anderson, N. J. (2011). Abrupt holocene climate

- 896 change as an important factor for human migration in west greenland. *Proceedings of the National*  
897 *Academy of Sciences* 108, 9765–9769. doi:[10.1073/pnas.1101708108](https://doi.org/10.1073/pnas.1101708108).
- 898 Epperson, J. F. (1987). On the Runge example. *The American Mathematical Monthly* 94, 329–341.  
899 doi:[10.2307/2323093](https://doi.org/10.2307/2323093).
- 900 Fahrmeir, L., and Kneib, T. (2008). “On the identification of trend and correlation in temporal  
901 and spatial regression,” in *Recent advances in linear models and related areas* (Physica-Verlag HD),  
902 1–27. doi:[10.1007/978-3-7908-2064-5\1](https://doi.org/10.1007/978-3-7908-2064-5_1).
- 903 Fahrmeir, L., Kneib, T., Lang, S., and Marx, B. (2013). *Regression: Models, methods and applications*.  
904 Springer Berlin Heidelberg doi:[10.1007/978-3-642-34333-9](https://doi.org/10.1007/978-3-642-34333-9).
- 905 Gautheir, T. D. (2001). Detecting trends using spearman’s rank correlation coefficient. *Envi-  
906 ronmental Forensics* 2, 359–362. doi:[10.1080/713848278](https://doi.org/10.1080/713848278).
- 907 Glew, J. R., Smol, J. P., and Last, W. M. (2001). “Sediment core collection and extrusion,” in  
908 *Tracking environmental change using lake sediments: Basin analysis, coring, and chronological techniques*,  
909 eds. W. M. Last and J. P. Smol (Dordrecht: Springer Netherlands), 73–105. doi:[10.1007/0-306-47669-X\5](https://doi.org/10.1007/0-306-47669-X\5).
- 911 Golding, N., and Purse, B. V. (2016). Fast and flexible bayesian species distribution modelling  
912 using gaussian processes. *Methods in Ecology and Evolution*. doi:[10.1111/2041-210X.12523](https://doi.org/10.1111/2041-210X.12523).
- 913 Handcock, M. S., Meier, K., and Nychka, D. (1994). Kriging and splines: An empirical compar-  
914 ision of their predictive performance in some applications: Comment. *Journal of the American  
915 Statistical Association* 89, 401–403. doi:[10.2307/2290838](https://doi.org/10.2307/2290838).
- 916 Hastie, T. J., and Tibshirani, R. J. (1990). *Generalized additive models*. Boca Raton, Fl.: Chapman  
917 & Hall / CRC.
- 918 Hastie, T., and Tibshirani, R. (1986). Generalized additive models. *Statistical Science* 1, 297–310.
- 919 Juggins, S., and Telford, R. J. (2012). “Exploratory data analysis and data display,” in *Tracking  
920 environmental change using lake sediments* (Springer, Dordrecht), 123–141. doi:[10.1007/978-94-007-2745-8\5](https://doi.org/10.1007/978-94-<br/>921 007-2745-8\5).
- 922 Kammann, E. E., and Wand, M. P. (2003). Geoadditive models. *Journal of the Royal Statistical  
923 Society. Series C, Applied statistics* 52, 1–18. doi:[10.1111/1467-9876.00385](https://doi.org/10.1111/1467-9876.00385).
- 924 Kimeldorf, G. S., and Wahba, G. (1970). A correspondence between bayesian estimation on  
925 stochastic processes and smoothing by splines. *Annals of Mathematical Statistics* 41, 495–502.
- 926 Magee, L. (1998). Nonlocal behavior in polynomial regressions. *The American Statistician* 52,  
927 20–22. doi:[10.2307/2685560](https://doi.org/10.2307/2685560).
- 928 Mann, M. E. (2004). On smoothing potentially non-stationary climate time series. *Geophysical  
929 Research Letters* 31, L07214. doi:[10.1029/2004GL019569](https://doi.org/10.1029/2004GL019569).
- 930 Mann, M. E. (2008). Smoothing of climate time series revisited. *Geophysical Research Letters* 35,  
931 L16708. doi:[10.1029/2008GL034716](https://doi.org/10.1029/2008GL034716).
- 932 Marra, G., and Wood, S. N. (2012). Coverage properties of confidence intervals for generalized  
933 additive model components. *Scandinavian Journal of Statistics, Theory and Applications* 39, 53–74.

- 934 doi:[10.1111/j.1467-9469.2011.00760.x](https://doi.org/10.1111/j.1467-9469.2011.00760.x).
- 935 McCullagh, P., and Nelder, J. A. (1989). *Generalized linear models, second edition*. CRC Press.
- 936 Mills, T. C. (2006). Modelling current trends in Northern Hemisphere temperatures. *International Journal of Climatology* 26, 867–884. doi:[10.1002/joc.1286](https://doi.org/10.1002/joc.1286).
- 938 Mills, T. C. (2007). A note on trend decomposition: The “classical” approach revisited  
939 with an application to surface temperature trends. *Journal of applied statistics* 34, 963–972.  
940 doi:[10.1080/02664760701590418](https://doi.org/10.1080/02664760701590418).
- 941 Mills, T. C. (2010). “Skinning a cat”: Alternative models of representing temperature trends.  
942 *Climatic Change* 101, 415–426.
- 943 Nychka, D. (1988). Bayesian confidence intervals for smoothing splines. *Journal of the American  
944 Statistical Association* 83, 1134–1143. doi:[10.1080/01621459.1988.10478711](https://doi.org/10.1080/01621459.1988.10478711).
- 945 PAGES 2K Consortium (2013). Continental-scale temperature variability during the past two  
946 millennia. *Nature Geoscience* 6, 339–346. doi:[10.1038/ngeo1797](https://doi.org/10.1038/ngeo1797).
- 947 Pinheiro, J. C., and Bates, D. M. (2000). *Mixed-Effects models in S and S-PLUS*. Springer Science  
948 & Business Media.
- 949 Pya, N. (2017). *Scam: Shape constrained additive models*. Available at: <https://CRAN.R-project.org/package=scam>.
- 951 Pya, N., and Wood, S. N. (2015). Shape constrained additive models. *Statistics and Computing*  
952 25, 543–559. doi:[10.1007/s11222-013-9448-7](https://doi.org/10.1007/s11222-013-9448-7).
- 953 Pya, N., and Wood, S. N. (2016). A note on basis dimension selection in generalized additive  
954 modelling. *ArXiv e-prints*.
- 955 R Core Team (2017). *R: A language and environment for statistical computing*. Vienna, Austria: R  
956 Foundation for Statistical Computing Available at: <https://www.R-project.org/>.
- 957 R Core Team (2018). *R: A language and environment for statistical computing*. Vienna, Austria: R  
958 Foundation for Statistical Computing Available at: <https://www.R-project.org/>.
- 959 Rasmussen, C. E., and Williams, C. K. I. (2006). *Gaussian processes for machine learning*. MIT  
960 Press.
- 961 Reiss, P. T., and Ogden, R. T. (2009). Smoothing parameter selection for a class of semipara-  
962 metric linear models. *Journal of the Royal Statistical Society. Series B, Statistical methodology* 71,  
963 505–523. doi:[10.1111/j.1467-9868.2008.00695.x](https://doi.org/10.1111/j.1467-9868.2008.00695.x).
- 964 Runge, C. (1901). Über empirische funktionen und die interpolation zwischen äquidistanten  
965 ordinaten. *Zeitschrift fur Angewandte Mathematik und Physik* 46, 224–243.
- 966 Ruppert, D., Wand, M. P., and Carroll, R. J. (2003). *Semiparametric Regression*. Cambridge Uni-  
967 versity Press.
- 968 Silverman, B. W. (1985). Some aspects of the spline smoothing approach to non-parametric  
969 regression curve fitting. *Journal of the Royal Statistical Society. Series B, Statistical methodology* 47,

- 970 1–52.
- 971 Smol, J. P. (2008). *Pollution of lakes and rivers: A paleoenvironmental perspective*. Blackwell Pub.
- 972 Smol, J. P., Birks, H. J. B., Lotter, A. F., and Juggins, S. (2012). “The march towards the quantitative analysis of palaeolimnological data,” in *Tracking environmental change using lake sediments* (Springer, Dordrecht), 3–17. doi:[10.1007/978-94-007-2745-8\1](https://doi.org/10.1007/978-94-007-2745-8_1).
- 975 Tian, J., Nelson, D. M., and Hu, F. S. (2011). How well do sediment indicators record past  
976 climate? An evaluation using annually laminated sediments. *Journal of Paleolimnology* 45, 73–  
977 84. doi:[10.1007/s10933-010-9481-x](https://doi.org/10.1007/s10933-010-9481-x).
- 978 Toney, J. L., Huang, Y., Fritz, S. C., Baker, P. A., Grimm, E., and Nyren, P. (2010). Climatic and environmental controls on the occurrence and distributions of long chain  
979 alkenones in lakes of the interior united states. *Geochimica et Cosmochimica Acta* 74, 1563–1578.  
980 doi:[10.1016/j.gca.2009.11.021](https://doi.org/10.1016/j.gca.2009.11.021).
- 982 Wahba, G. (1983). Bayesian “confidence intervals” for the Cross-Validated smoothing spline.  
983 *Journal of the Royal Statistical Society. Series B, Statistical methodology* 45, 133–150.
- 984 Wahba, G. (1990). *Spline models for observational data*. SIAM.
- 985 Wang, Y., Naumann, U., Wright, S. T., and Warton, D. I. (2012). Mvabund- an R package for  
986 model-based analysis of multivariate abundance data. *Methods in Ecology and Evolution* 3, 471–  
987 474. doi:[10.1111/j.2041-210X.2012.00190.x](https://doi.org/10.1111/j.2041-210X.2012.00190.x).
- 988 Warton, D. I. (2011). Regularized sandwich estimators for analysis of high-dimensional  
989 data using generalized estimating equations. *Biometrics* 67, 116–123. doi:[10.1111/j.1541-0420.2010.01438.x](https://doi.org/10.1111/j.1541-0420.2010.01438.x).
- 991 Warton, D. I., Blanchet, F. G., O’Hara, R. B., Ovaskainen, O., Taskinen, S., Walker, S. C., et al.  
992 (2015). So many variables: Joint modeling in community ecology. *Trends in Ecology & Evolution*.  
993 doi:[10.1016/j.tree.2015.09.007](https://doi.org/10.1016/j.tree.2015.09.007).
- 994 Warton, D. I., Wright, S. T., and Wang, Y. (2012). Distance-based multivariate analyses confound location and dispersion effects. *Methods in Ecology and Evolution* 3, 89–101.  
995 doi:[10.1111/j.2041-210X.2011.00127.x](https://doi.org/10.1111/j.2041-210X.2011.00127.x).
- 997 Wood, S. N. (2003). Thin plate regression splines. *Journal of the Royal Statistical Society. Series B, Statistical methodology* 65, 95–114. doi:[10.1111/1467-9868.00374](https://doi.org/10.1111/1467-9868.00374).
- 999 Wood, S. N. (2004). Stable and efficient multiple smoothing parameter estimation for  
1000 generalized additive models. *Journal of the American Statistical Association* 99, 673–686.  
1001 doi:[10.2307/27590439](https://doi.org/10.2307/27590439).
- 1002 Wood, S. N. (2011). Fast stable restricted maximum likelihood and marginal likelihood estimation of semiparametric generalized linear models. *Journal of the Royal Statistical Society. Series B, Statistical methodology* 73, 3–36. doi:[10.1111/j.1467-9868.2010.00749.x](https://doi.org/10.1111/j.1467-9868.2010.00749.x).
- 1005 Wood, S. N. (2017). *Generalized Additive Models: An Introduction with R, Second edition*. CRC Press.
- 1006 Wood, S. N., Pya, N., and Säfken, B. (2016). Smoothing parameter and model selection  
1007 for general smooth models. *Journal of the American Statistical Association* 111, 1548–1563.

1008 doi:[10.1080/01621459.2016.1180986](https://doi.org/10.1080/01621459.2016.1180986).

1009 Yee, T. W., and Mitchell, N. D. (1991). Generalized additive models in plant ecology. *Journal of*  
1010 *Vegetation Science* 2, 587–602. doi:[10.2307/3236170](https://doi.org/10.2307/3236170).

1011 Zink, K.-G., Leythaeuser, D., Melkonian, M., and Schwark, L. (2001). Temperature de-  
1012 pendency of long-chain alkenone distributions in recent to fossil limnic sediments and  
1013 in lake waters<sup>11</sup>Associate editor: J. b. fein. *Geochimica et Cosmochimica Acta* 65, 253–265.  
1014 doi:[10.1016/S0016-7037\(00\)00509-3](https://doi.org/10.1016/S0016-7037(00)00509-3).

## 1015 Appendix 1 — Simultaneous intervals

1016 We proceed by considering a simultaneous confidence interval for a function  $f(x)$  at a set of  $M$   
1017 locations in  $x$ ; we'll refer to these locations, following the notation of Ruppert et al. (2003) by

$$\mathbf{g} = (g_1, g_2, \dots, g_M)$$

1018 The true function over  $\mathbf{g}$ ,  $\mathbf{f}_g$ , is defined as the vector of evaluations of  $f$  at each of the  $M$  locations

$$\mathbf{f}_g \equiv \begin{bmatrix} f(g_1) \\ f(g_2) \\ \vdots \\ f(g_M) \end{bmatrix}$$

1019 and the corresponding estimate of the true function given by the fitted GAM denoted by  $\hat{\mathbf{f}}_g$ .  
1020 The difference between the true function and our unbiased estimator is given by

$$\hat{\mathbf{f}}_g - \mathbf{f}_g = \mathbf{C}_g \begin{bmatrix} \hat{\beta} - \beta \\ \hat{\mathbf{u}} - \mathbf{u} \end{bmatrix},$$

1021 where  $\mathbf{C}_g$  is a matrix formed by the evaluation of the basis functions at locations  $\mathbf{g}$ , and the part  
1022 in square brackets is the bias in the estimated model coefficients, which we assume to be mean 0  
1023 and distributed, approximately, multivariate normal with mean vector  $\mathbf{0}$  and covariance matrix  
1024  $\mathbf{V}_b$

$$\begin{bmatrix} \hat{\beta} - \beta \\ \hat{\mathbf{u}} - \mathbf{u} \end{bmatrix} \stackrel{\text{approx.}}{\sim} N(\mathbf{0}, \mathbf{V}_b),$$

1025 where  $\mathbf{V}_b$  is the Bayesian covariance matrix of the GAM coefficients.

1026 Now, the  $(1 - \alpha)100\%$  simultaneous confidence interval is

$$\hat{\mathbf{f}}_g \pm m_{1-\alpha} \begin{bmatrix} \widehat{\text{st.dev}}(\hat{f}(g_1) - f(g_1)) \\ \widehat{\text{st.dev}}(\hat{f}(g_2) - f(g_2)) \\ \vdots \\ \widehat{\text{st.dev}}(\hat{f}(g_M) - f(g_M)) \end{bmatrix},$$

1027 where  $m_{1-\alpha}$  is the  $1 - \alpha$  quantile of the random variable

$$\sup_{x \in x} \left| \frac{\hat{f}(x) - f(x)}{\widehat{\text{st.dev}}(\hat{f}(x) - f(x))} \right| \approx \max_{1 \leq \ell \leq M} \left| \frac{\left( \mathbf{C}_g \begin{bmatrix} \hat{\beta} - \beta \\ \hat{\mathbf{u}} - \mathbf{u} \end{bmatrix} \right)_\ell}{\widehat{\text{st.dev}}(\hat{f}(g_\ell) - f(g_\ell))} \right|$$

1028 The sup refers to the *supremum* or the *least upper bound*; this is the least value of  $\mathcal{X}$ , the set of all  
1029 values of which we observed subset  $x$ , that is *greater* than all of the values in the subset. Often  
1030 this is the maximum value of the subset. This is what is indicated by the right-hand side of  
1031 the equation; we want the maximum (absolute) value of the ratio over all values in  $g$ .

1032 The fractions in both sides of the equation correspond to the standardized deviation between  
1033 the true function and the model estimate, and we consider the *maximum absolute* standardized  
1034 deviation. We don't usually know the distribution of the maximum absolute standardized  
1035 deviation but we need this to access its quantiles. However, we can closely approximate the  
1036 distribution via simulation. The difference here is that rather than simulating from the poste-  
1037 rior of the model as we did earlier [see section Confidence intervals](#), this time we simulate from  
1038 the multivariate normal distribution with mean vector  $\mathbf{0}$  and covariance matrix  $\mathbf{V}_b$ . For each  
1039 simulation we find the maximum absolute standardized deviation of the fitted function from  
1040 the true function over the grid of  $x$  values we are considering. Then we collect all these max-  
1041 ima, sort them and either take the  $1 - \alpha$  probability quantile of the maxima, or the maximum  
1042 with rank  $\lceil (1 - \alpha)/N \rceil$ .