

GEMS Workshop

Metastability and kinetics

Quartz (aggregate) dissolution exercise

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Chemical equilibrium





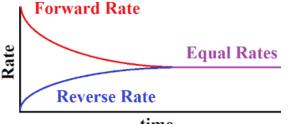
Thermodynamic equilibrium: rate of the forward reaction equals the rate of the reverse reaction, final state of the system (given by the configuration with the minimum Gibbs energy)

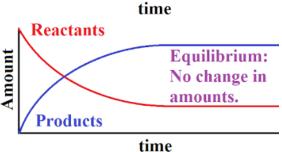
Kinetics: deals with the rate at which chemical reactions occur, pathway, speed, and mechanism leading to equilibrium

Metastability: an intermediate state that appears stable but is not the true equilibrium state, energy barriers; very slow transition; stuck when energy barrier too high (strong bonds)

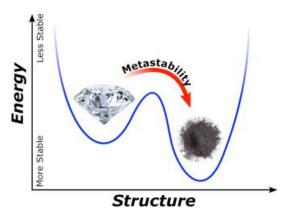
All can be modeled using thermodynamics (with constraints) in GEMS (e.g., dissolution, precipitation)











GEMS Thermodynamic equilibrium (complete equilibrium)



GEMS finds the **amounts of chemical components in phases** with stability indexes Ω in an **equilibrium state** (with minimum Gibbs energy) **defined by T, P, bulk elemental composition, thermodynamic data**.



- all reactions are infinitely fast
- slow reactions over infinite time

In many cases this is not attained at Earth surface conditions.

Time dependent metastability \rightarrow use the principle of partial equilibria (discretize the path to equilibrium as a sequence of partial equilibrium states)

- Lower metastability constraints allow stepwise simulation of mineral dissolution (min. amount imposed)
- Upper metastability constraints allow stepwise simulation of mineral precipitation (max. amount allowed)

Metastability constraints in recipe



Upper metastability, precipitation

1. Create

SysEq Recipe 0: 'PortlKin' at 1 bar 25 C, var. 0

Property	Name	Quantity	Units
xa_ xd_ bi	Aqua Portlandite	1 0.05 0.0004	G M M

2. Clone

SysEq Recipe 0: 'PortlKin' at 1 bar 25 C, var. 1

Prope	rty Name	Quantity	Units
xa_	Aqua	1	G
xd_	Portland	dite 0.05	M
bi_	0	0.0004	M
dul_	Portland	dite 0.01	M

Saturation Index

 $log10(\Omega) = 0 \rightarrow saturated (equilibrium)$

 $log 10(\Omega) < 0 \rightarrow undersaturated$

 $log10(\Omega) > 0 \rightarrow oversaturated$

Complete equilibrium (saturation):

pH = 12.47; IS = 0.05 m;

 $m_t(Ca) = 0.0204 m;$

n(Portlandite) = 0.0296

 Ω (Portlandite) = 1

 $log10(\Omega)$ SI/Fa(Portlandite) = 0 (-9.2e-7)

Partial equilibrium (oversaturation):

pH = 12.72; IS = 0.097 m;

 $m_t(Ca) = 0.04 m;$

n(Portlandite) = 0.01

 $\Omega(Portlandite) = 4.66$

 $log10(\Omega)$ SI/Fa(Portlandite) = 0.67

A 'dul_ = f(time)' would account for the precipitation rate

Metastability constraints in recipe



Lower metastability, dissolution

3. Clone

SysEq Recipe 1: 'PortlKin' at 1 bar 25 C, var. 0

Property	Name	Quantity	Units
xa_ xd_ bi_	Aqua Portlandite O	0.02 0.0004	G M M

4. Clone

SysEq Recipe 1: 'PortlKin' at 1 bar 25 C, var. 1

Property	Name	Quantity	Units
xa_	Aqua	1	G
xd_	Portlandite	0.02	M
bi_	0	0.0004	M
dll_	Portlandite	0.01	M

Saturation Index

 $log10(\Omega) = 0 \rightarrow saturated (equilibrium)$

 $log10(\Omega) < 0 \rightarrow undersaturated$

 $log10(\Omega) > 0 \rightarrow oversaturated$

Complete equilibrium (undersaturation):

$$pH = 12.46$$
; $IS = 0.05 m$;

$$m_t(Ca) = 0.02 m;$$

$$n(Portlandite) = 0.0$$

$\Omega(Portlandite) = 0.958$

$$log10(\Omega)$$
 SI/Fa Fa(Portlandite) = -0.018

Partial equilibrium (undersaturation):

$$pH = 12.2$$
; IS = 0.027 m;

$$m_t(Ca) = 0.01 m;$$

$$n(Portlandite) = 0.01$$

$$\Omega(Portlandite) = 0.183$$

$$log10(\Omega)$$
 SI/Fa Fa(Portlandite) = -0.738

A 'dll_ = f(time)' would account for the dissolution rate

Homework: Check out and test the "Kinetics" test project

(look at Process simulators and Phase definitions for calcite, portlandite)

Kinetic rate law



General equation (Lasaga, 1998), t (time s)

rate constant possible fitting parameters
$$r = \pm kS | 1 - \Omega^{\Theta}|^{\eta}$$
 net rate reactive surface saturation ratio (mol s⁻¹) area (m²)

$$\begin{aligned} \operatorname{dul}_{-t+\Delta t} &= \operatorname{dul}_{-t} + r_t \Delta t \text{ if } \log_{10} \Omega > 0 \\ \operatorname{dll}_{-t+\Delta t} &= \operatorname{dll}_{-t} - r_t \Delta t \text{ if } \log_{10} \Omega < 0 \end{aligned}$$

Kinetic rate law



The rate constant can be written as a function of T, activity (e.g., pH), ionic strength, partial pressure of a gas, etc.

Can be defined as a summation of different mechanisms.

For dissolution

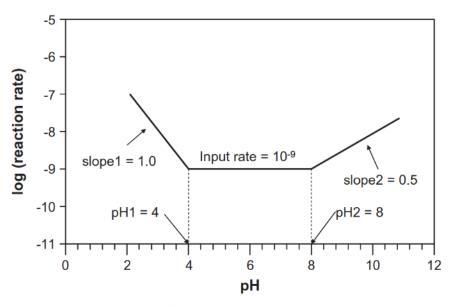
$$dll_{-t+\Delta t} = dll_{-t} - r_t \Delta t$$

$$r_t^{nu} = k^{nu} S |1 - \Omega^{\Theta}|^{\eta}$$

$$r_t^H = k^H S |1 - \Omega^{\Theta}|^{\eta}$$

$$r_t^{OH} = k^{OH} S |1 - \Omega^{\Theta}|^{\eta}$$

$$dll_{-t+\Delta t} = dll_{-t} - (r_t^{nu} + r_t^H + r_t^{OH})\Delta t$$



$$k = k_{25}^{nu} \exp\left[\frac{-E_a^{nu}}{R}\left(\frac{1}{T} - \frac{1}{298.15}\right)\right] +$$

$$k_{25}^{H} \exp \left[\frac{-E_{a}^{H}}{R} \left(\frac{1}{T} - \frac{1}{298.15} \right) \right] a_{\mathrm{H}}^{n_{\mathrm{H}}} +$$

$$k_{25}^{\text{OH}} \exp \left[\frac{-E_a^{\text{OH}}}{R} \left(\frac{1}{T} - \frac{1}{298.15} \right) \right] a_{\text{OH}}^{n_{\text{OH}}}$$

Quartz sand aggregate dissolution



2. Aggregate (quartz) dissolution (kinetics)

We define a kinetic rate law function in our jupyter notebook and we couple it with xgems to simulate a dissolution reaction, and sample the aqueous and solid changes

1st we need a GEMS chemical system from GEM-Selektor – we export it to be used in xgems, then we continue in the Jupyter Lab

Defining the initial system in GEM-Selektor (PC2 project)

SysEq: Please, enter a new record key:

Name of the modeling project

dynamic notential to minimize IG GVA

PC2:G:mortar:0:0:1:25:0:



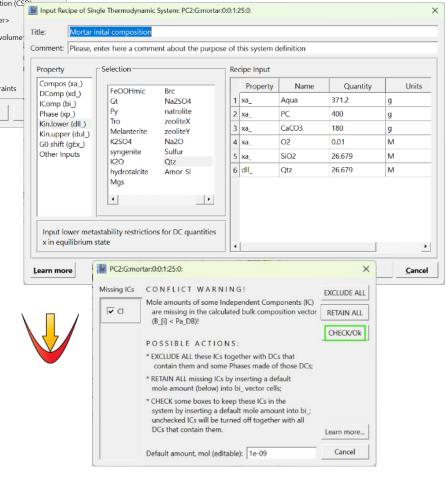


Tab. A-3: GEMS recipes of CPW systems at degradation Stage I

							G	Thermodynamic potential to minimize (G	i GVJ
Property	Compos or species name	Units	Stage-I- mt/sh (oxic)	Recipe variants: Porosity after hydration	Stage-I- mt/sh (reducing)	Recipe variants: Porosity after hydration	0 0 1 25	Name of the chemical system definition (CSD (recipe) variant number <integer> Volume of the system, dm3 (0 if no volur Pressure, bar, or 0 for Psat(H2O)g Temperature, C (>= 0)</integer>	Ti
xa_	Aqua (H ₂ O)	g	254.047 371.2	~ 11 % ~ 20 %	254.047 371.2 533.6	~ 11 % ~ 20 % ~ 30 %	0 <u>O</u> k	Variant number for additional constraints Reset From List Help	5
xa_	CEM-I- 52.5R	g	400		400		400		
xa_	CaCO ₃	g	180		180		180		
xa_	O ₂ ***	mol	-		-0.0017		-0.017		
xa_	SiO ₂	g	1,603		1,603		1,603		

Switch off `-`:

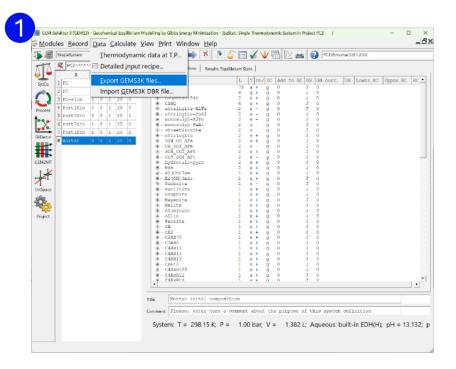
gibbsite, kaolinite hematite, goethite, ettringite-AlFe, ettringite-FeAl, SO4_CO3_AFt monosulph-AlFe, monosulph-FeAl, SO4_OH_AFm,

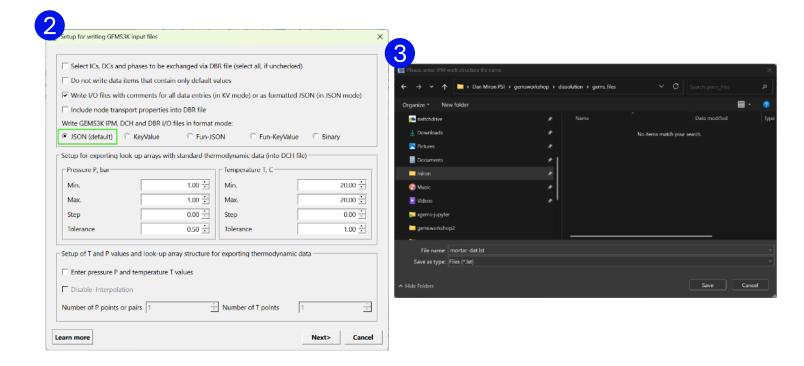


Export GEMS3K files



In gemsworkshop folder create subfolders dissolution and inside gems_files







For dissolution example create a new notebook aggregate_dissolution inside the leaching folder.

We use the following parameters for the rate law (neutral and basic)

Selected parameters for the dissolution rate model: regression curve model. Kinetic constants (k) are expressed in mol m⁻² s⁻¹ and the activation energy (E_a) in kJ mol⁻¹.

Mineral	Type of surface normalisation	k_{25}^{nu}	E_a^{nu}	k_{25}^{H+}	E_a^{H+}	n^{H+}	$k_{25}^{\mathrm{OH}-}$	E_a^{OH-}	$n^{\mathrm{OH}-}$
Albite	BET	5.1×10^{-20}	57	8.5×10^{-11}	58	0.34	1.4×10^{-10}	56	0.32
Biotite	BET	2.3×10^{-12}	49	1.1×10^{-09}	49	0.67	9.1×10^{-08}	49	0.79
Celestite	BET	2.2×10^{-08}	34	1.4×10^{-06}	33	0.10			
Chlorite	0.2% BET ^a	6.4×10^{-17}	16	8.2×10^{-09}	17	0.28	6.9×10^{-09}	16	0.34
C-S-H	BET	1.6×10^{-18}	23	5.9×10^{-08}	23	0.28			
Cristobalite	BET	6.4×10^{-14}	69				1.9×10^{-10}	69	0.34
Dolomite	BET	1.1×10^{-08}	31	2.8×10^{-04}	46	0.61			
Gibbsite	BET						3.1×10^{-06}	48	1.0
Illite	ESA	3.3×10^{-17}	35	9.8×10^{-12}	36	0.52	3.1×10^{-12}	48	0.38
Kaolinite	BET	1.1×10^{-14}	38	7.5×10^{-12}	43	0.51	2.5×10^{-11}	46	0.58
Microcline	BET	1.0×10^{-14}	31	1.7×10^{-11}	31	0.27	1.4×10^{-10}	31	0.35
Portlandite	BET	2.2×10^{-08}	75	8.0×10^{-04}	75	0.60			
Quartz	BET	6.4×10^{-14}	77				1.9×10^{-10}	80	0.34
Siderite	BET	2.1×10^{-09}	56	5.9×10^{-06}	56	0.60			
Montmorillonite	ESA	9.3×10^{-15}	63	5.3×10^{-11}	54	0.69	2.9×10^{-12}	61	0.34

Thank you for your attention!

Questions?



