

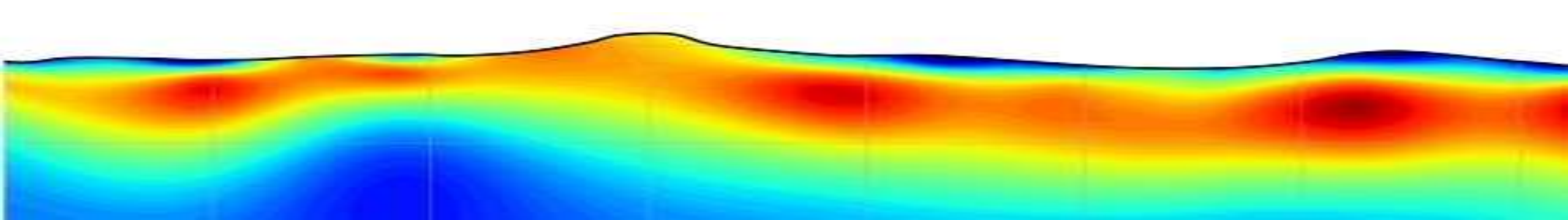
ESS302 Applied Geophysics II

Gravity, Magnetic, Electrical, Electromagnetic and Well Logging

EM Wrap-up

Instructor: Dikun Yang

Feb – May, 2020



well logging
(everything in borehole)

Maxwell Equations

$$\nabla \cdot \mathbf{D} = \rho$$

$$\nabla \cdot \mathbf{B} = 0$$

$$\nabla \times \mathbf{E} = -\frac{\partial \mathbf{B}}{\partial t}$$

$$\nabla \times \mathbf{H} = \mathbf{J} + \frac{\partial \mathbf{D}}{\partial t}$$

zero frequency

low frequency

high frequency

steady state

quasi-static state

EM wave

mechanical wave

magnetic

gravity

potential field

electrical

electromagnetic (induction)

electrical conductivity/resistivity

electromagnetic (geo-radar)

seismic

wave phenomena

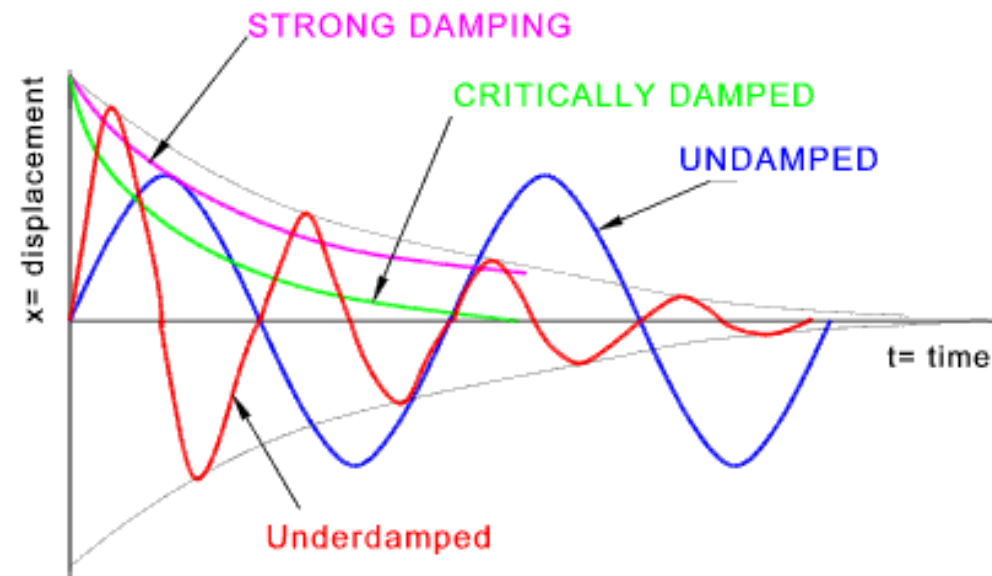
Quasi-static Maxwell's Equations

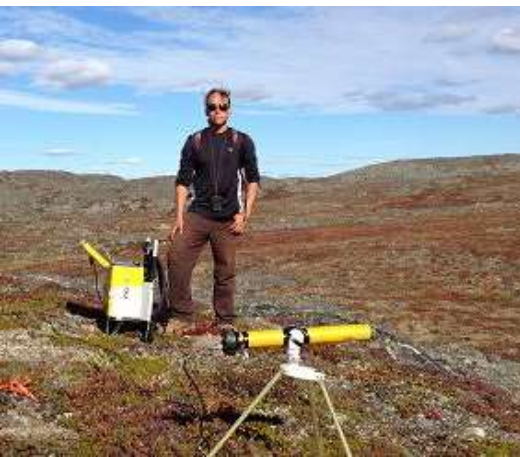
$$\nabla \times E = -\mu \frac{\partial H}{\partial t}$$

$$\nabla \times E = -i\omega\mu H$$

$$\nabla \times H = \sigma E + \epsilon \frac{\partial E}{\partial t}$$

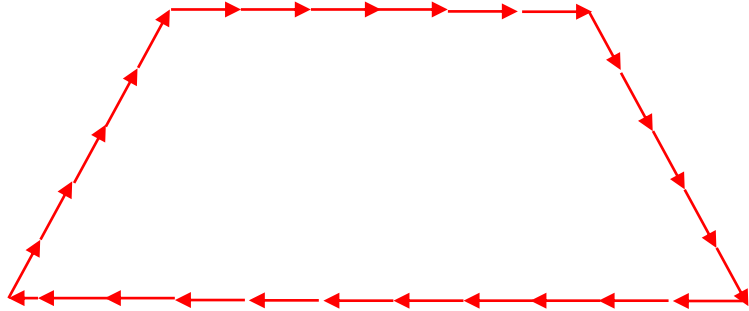
$$\nabla \times H = \sigma E + i\omega\epsilon E$$





Wires and Loops

Electrical dipole
(a *small* piece of wire)



Closed loop

- Magnetic field (dB/dt)
- Non-contact (divergence free)
- Inductive coupling



Grounded wire

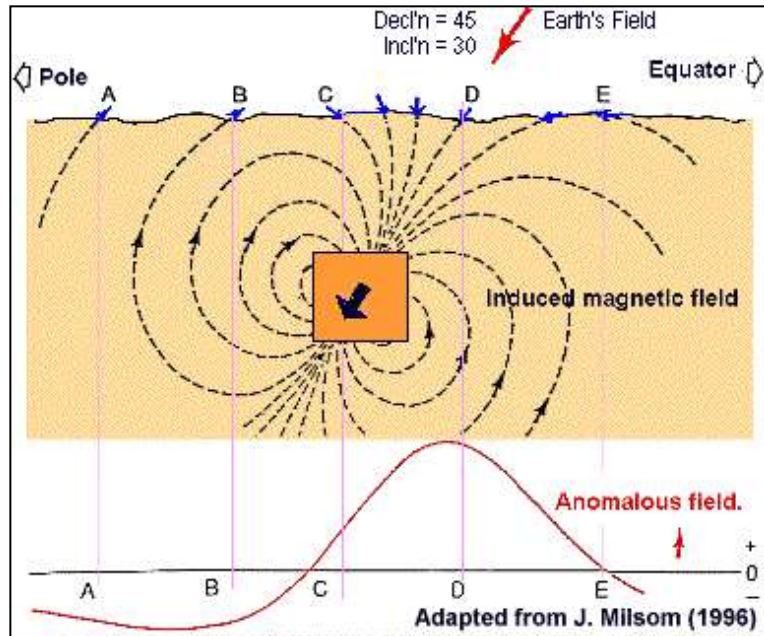
- Electrical field (E)
- End points in contact with ground
- Galvanic and inductive coupling

Loop-loop System in Frequency Domain



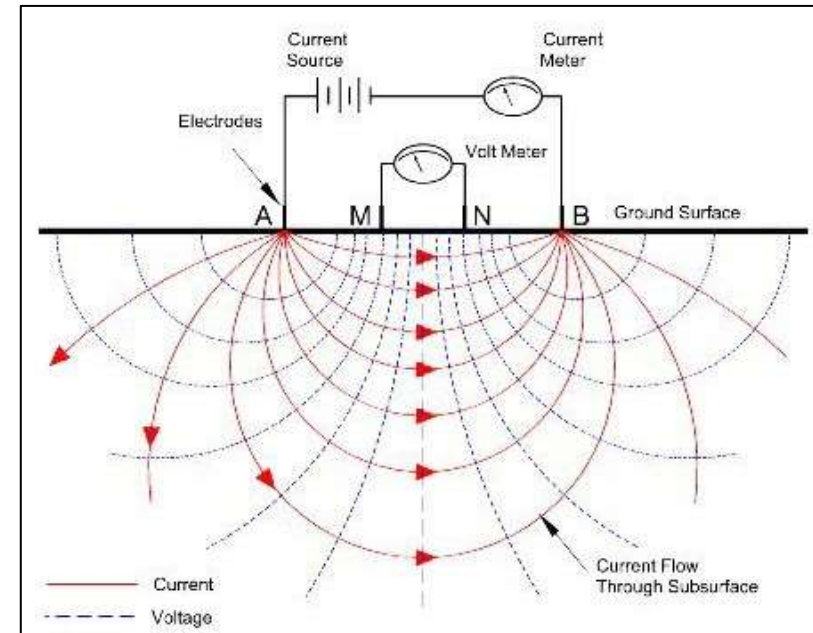
EM =

Magnetics



- Magnetic dipole
- Magnetic flux (B)

Electric Resistivity

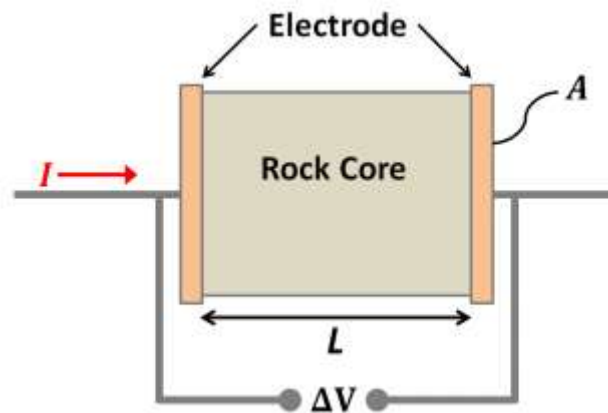


- Electric dipole
- Electric current (J)

Electrical energy transmission

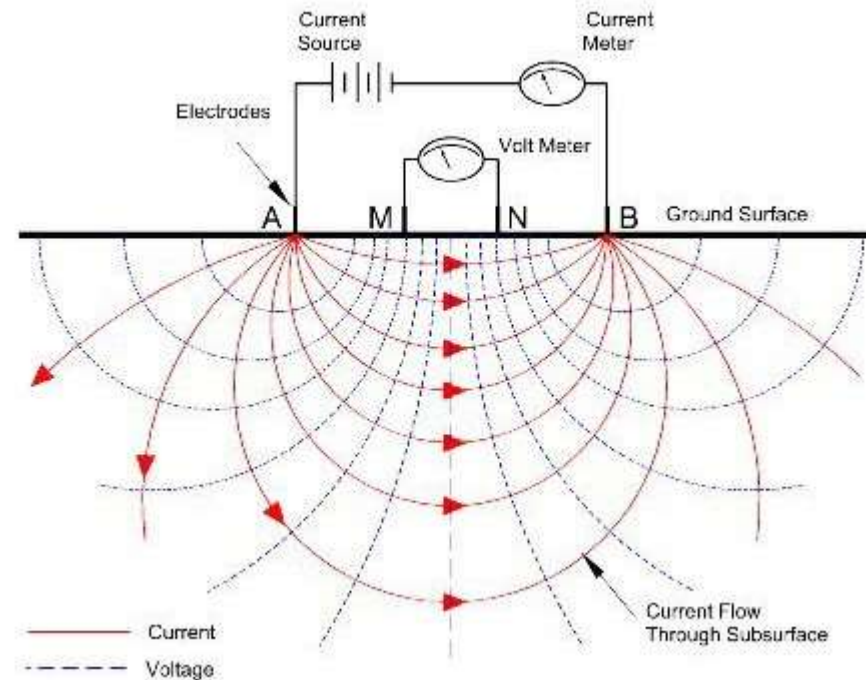
- Galvanic (electric current)

Ohm's law



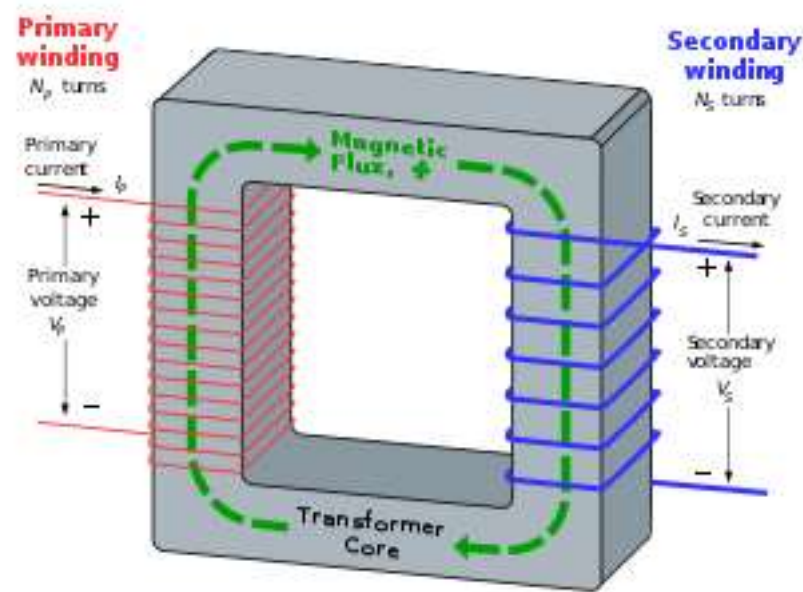
$$R = \frac{\Delta V}{I}$$

DC resistivity
(electric resistivity tomography)



Electrical energy transmission

- Inductive (magnetic flux B)



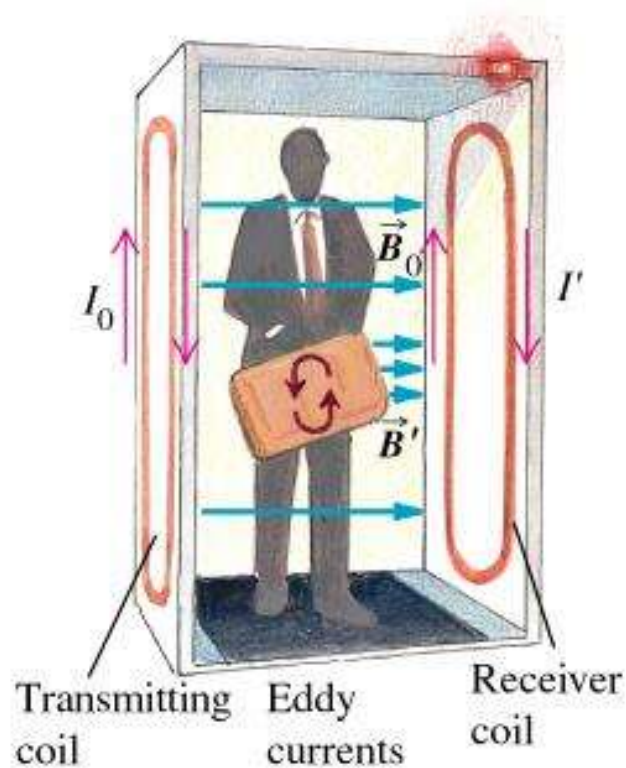
1. Change of current in the primary
2. Change of magnetic flux in the core
3. Induced current in the secondary

A transformer:

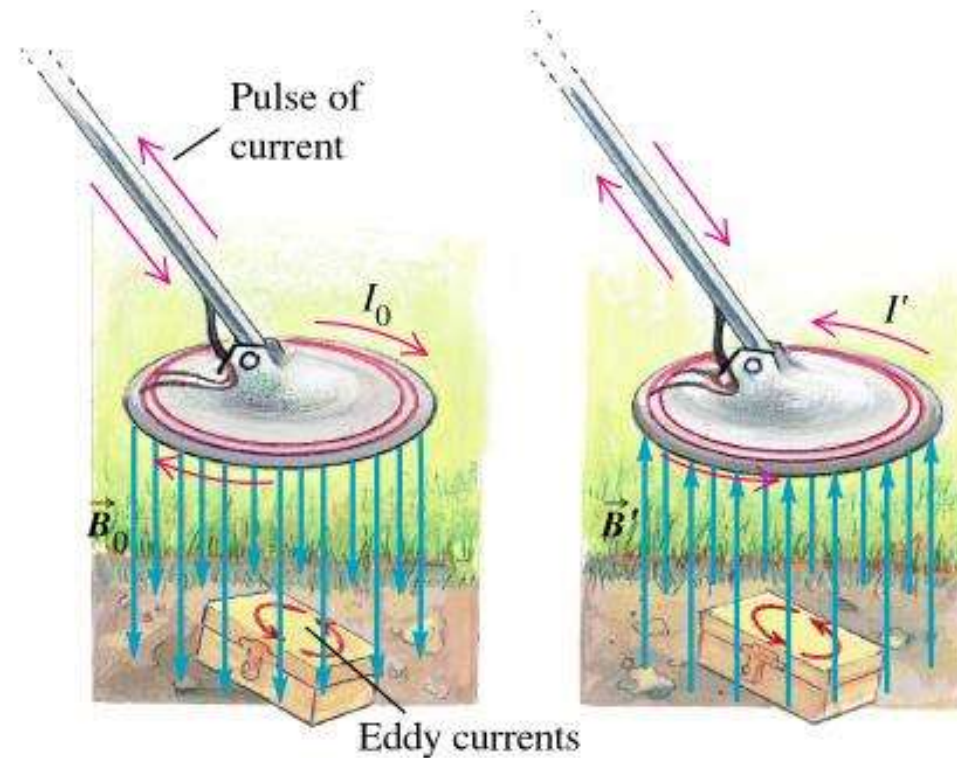
- No direct connection between primary and secondary windings
- Energy goes through in the forms of electric, magnetic then electric
- Magnetic flux linkage only in AC (requires non-stationary current)

Electrical energy transmission

- Inductive (magnetic flux B)



Security scan



Metal detector

Ampere's law

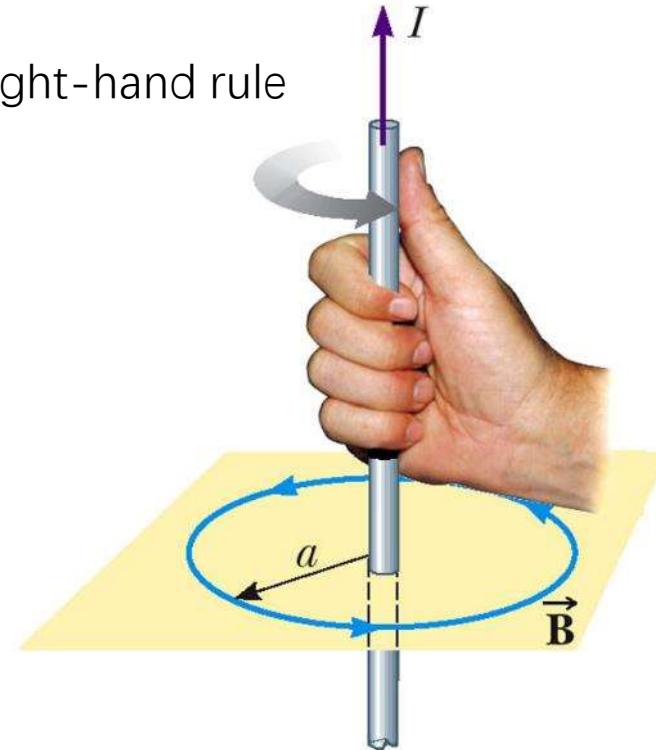
J generates B

$$\nabla \times \mu^{-1} \mathbf{B} = \mathbf{J} = \sigma \mathbf{E}$$



Current in wire causes a magnetic field to surround it (iron filings).

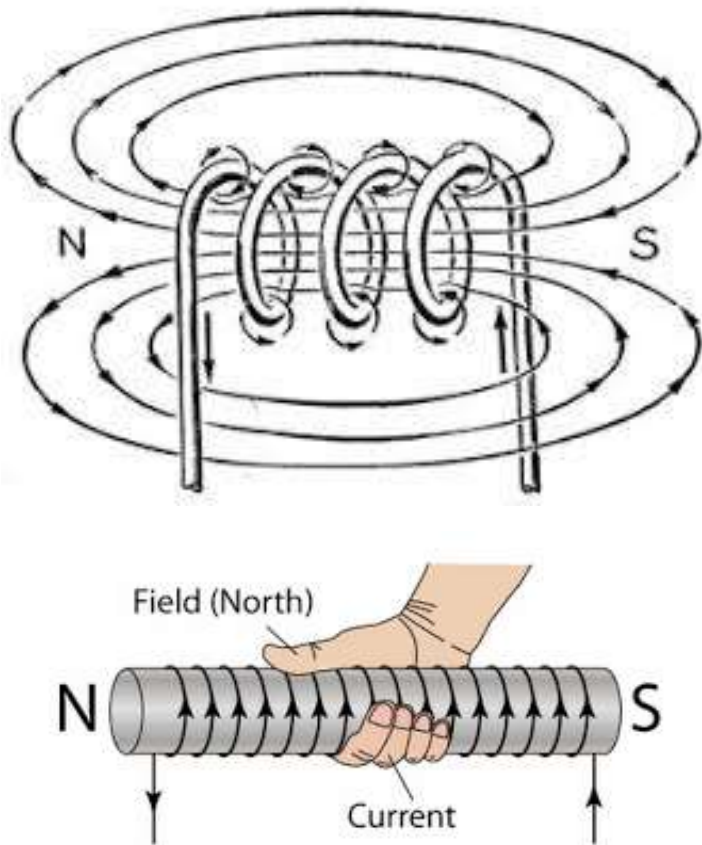
Right-hand rule



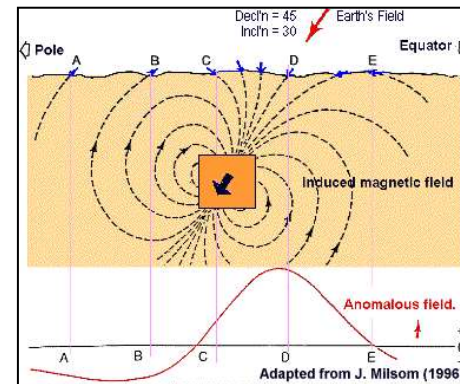
Ampere's law

J generates B

$$\nabla \times \mu^{-1} \mathbf{B} = \mathbf{J} = \sigma \mathbf{E}$$



A small solenoid generates a magnetic field that can be approximated by a magnetic dipole (or a small bar magnet)

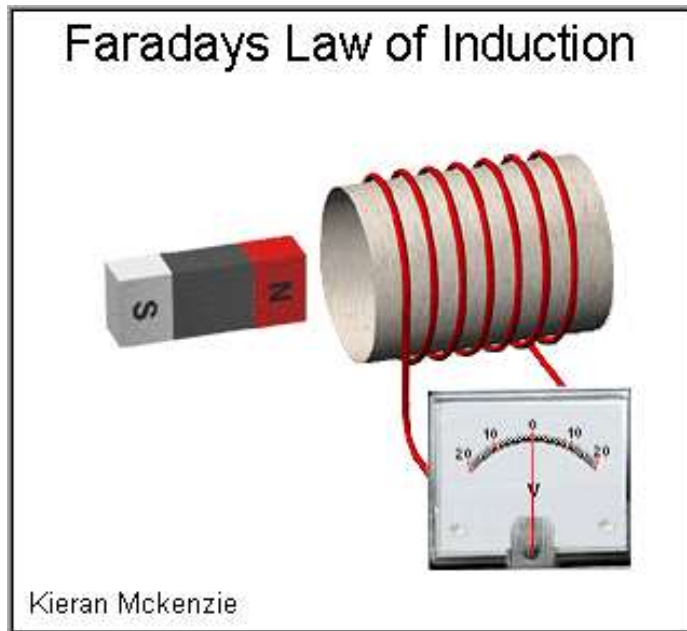


Still remember the magnetic dipole?

Faraday's law

Change of B generates J

$$\nabla \times \sigma^{-1} \mathbf{J} = \nabla \times \mathbf{E} = -\frac{\partial \mathbf{B}}{\partial t}$$

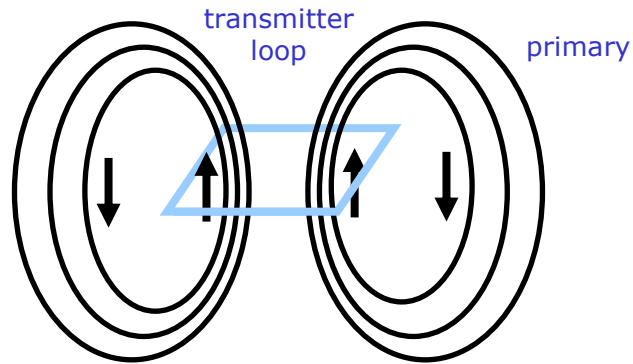


Induced current depends on

- How fast B changes
- How many B-field lines go through
- How conductive the object is

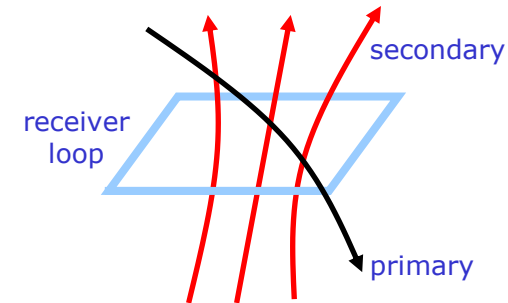
Communicate with the Earth without Contact

Transmitter loop



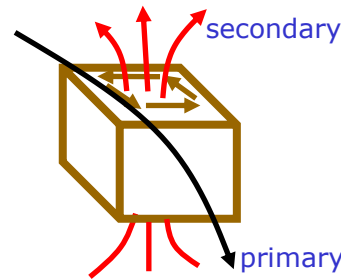
Ampere: time-varying current and changing primary magnetic field

Receiver loop



Faraday: measurable current induced in the loop by the changing secondary field

Target/Ground



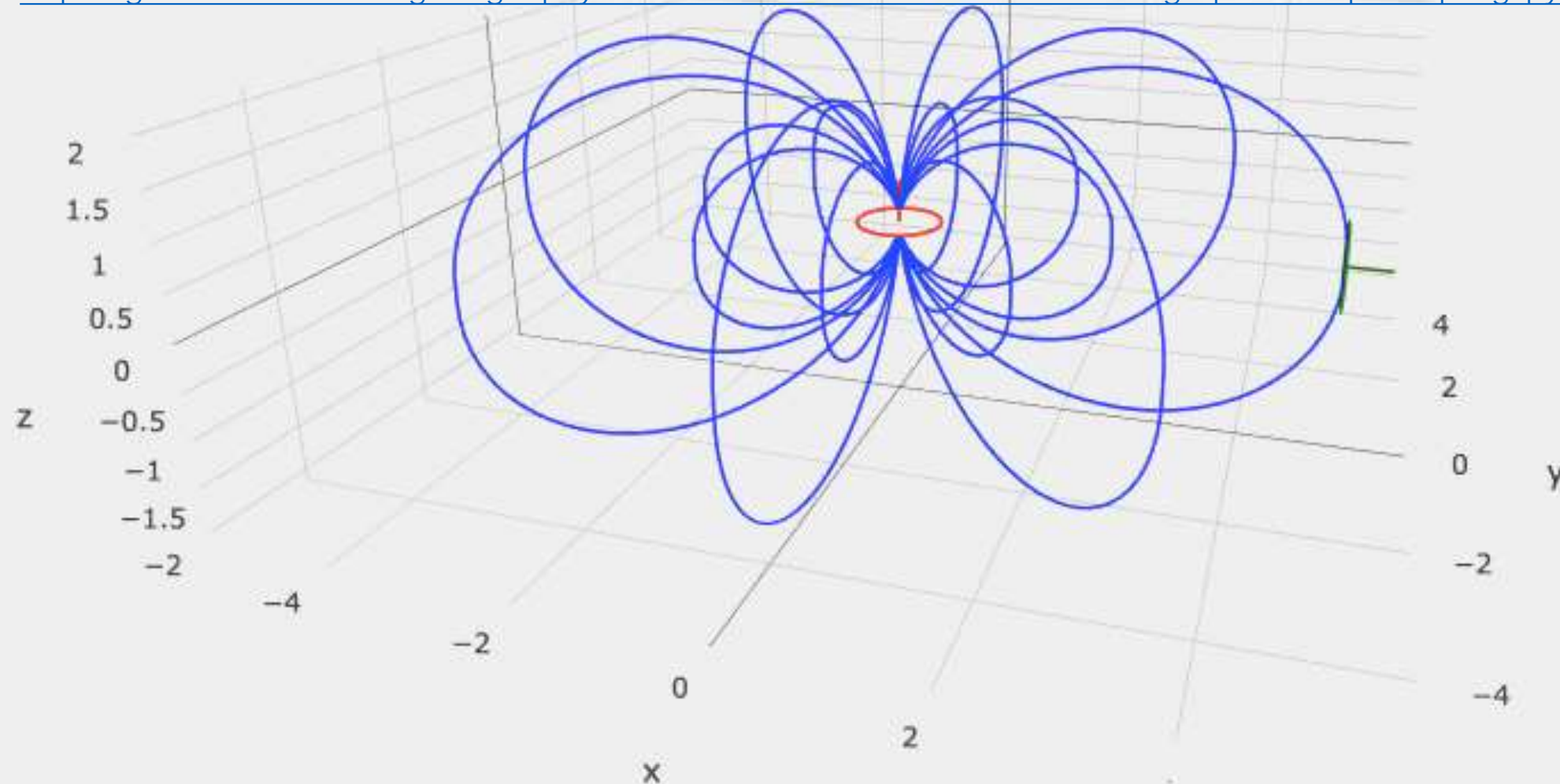
Faraday: current induced by the changing primary field;

Ampere: induced current generates a secondary magnetic field

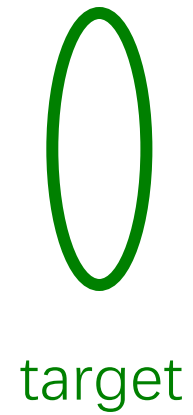
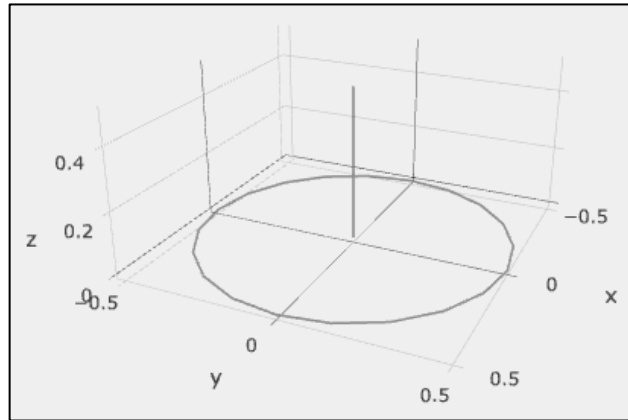
Notebook: Loop, dipole and field lines

- “MagDipole2LoopsCoupling.ipynb”

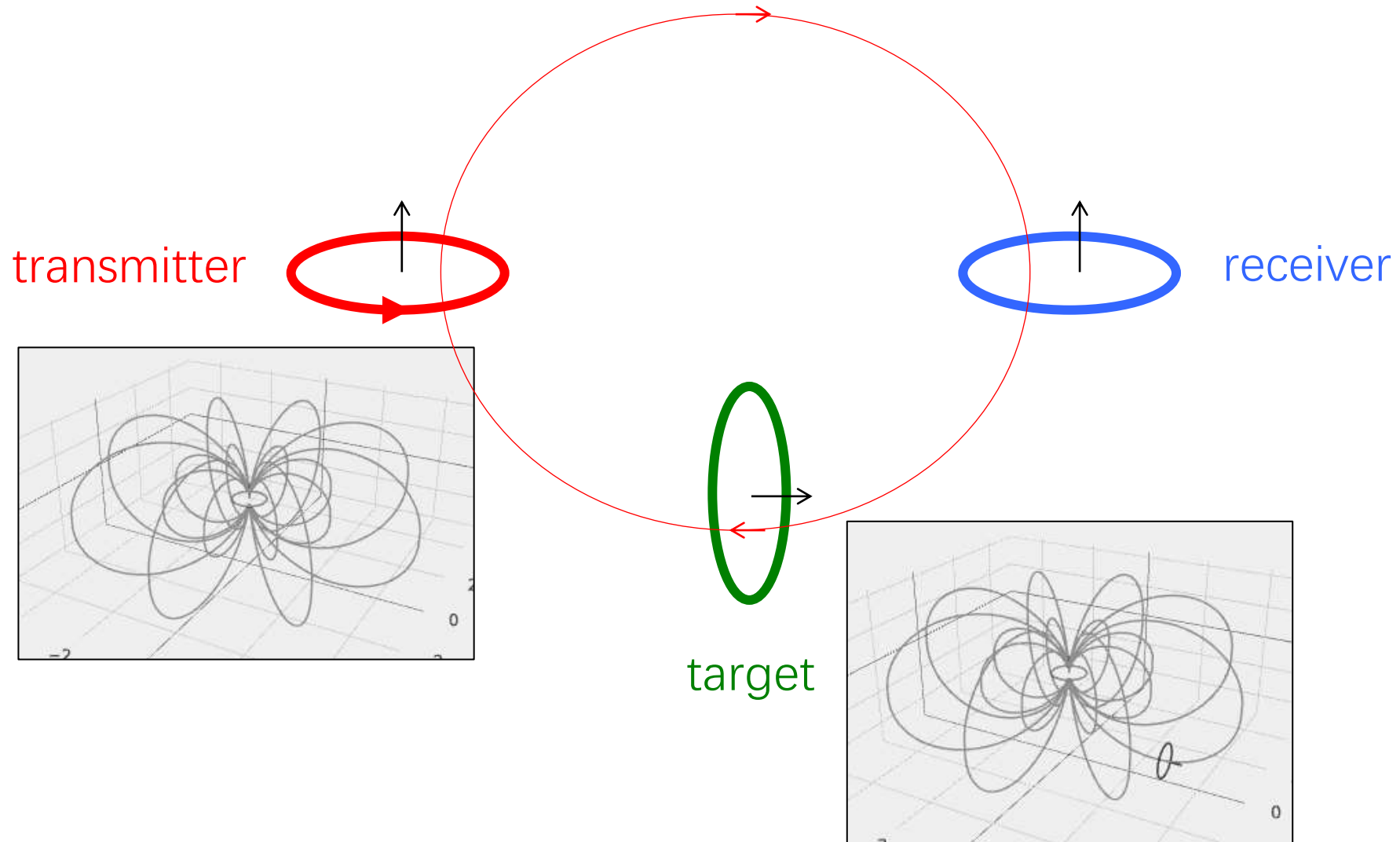
<https://github.com/sustechgem/geophysics-demo-notebooks/blob/master/MagDipole2LoopsCoupling.ipynb>



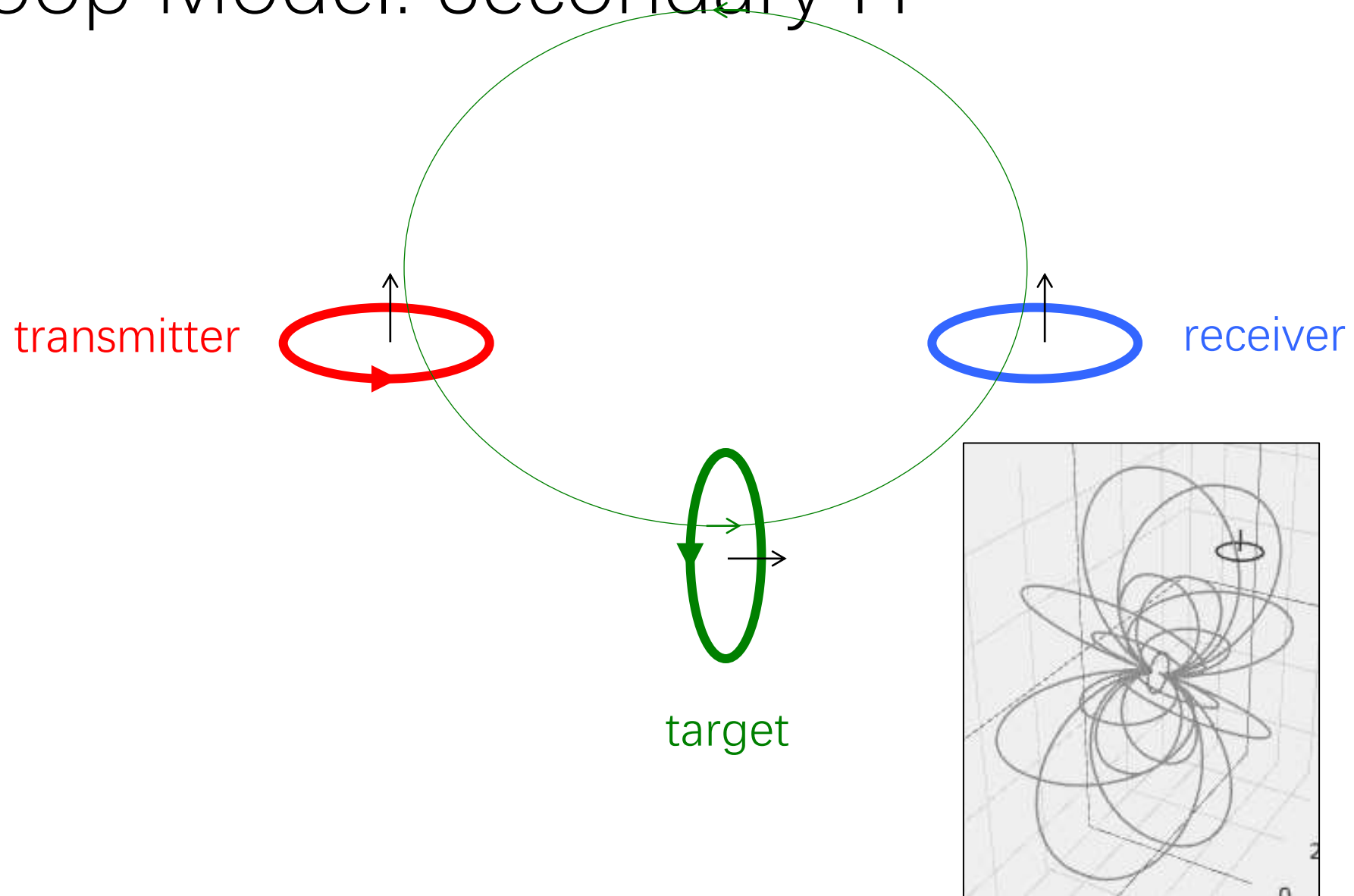
3-loop Model



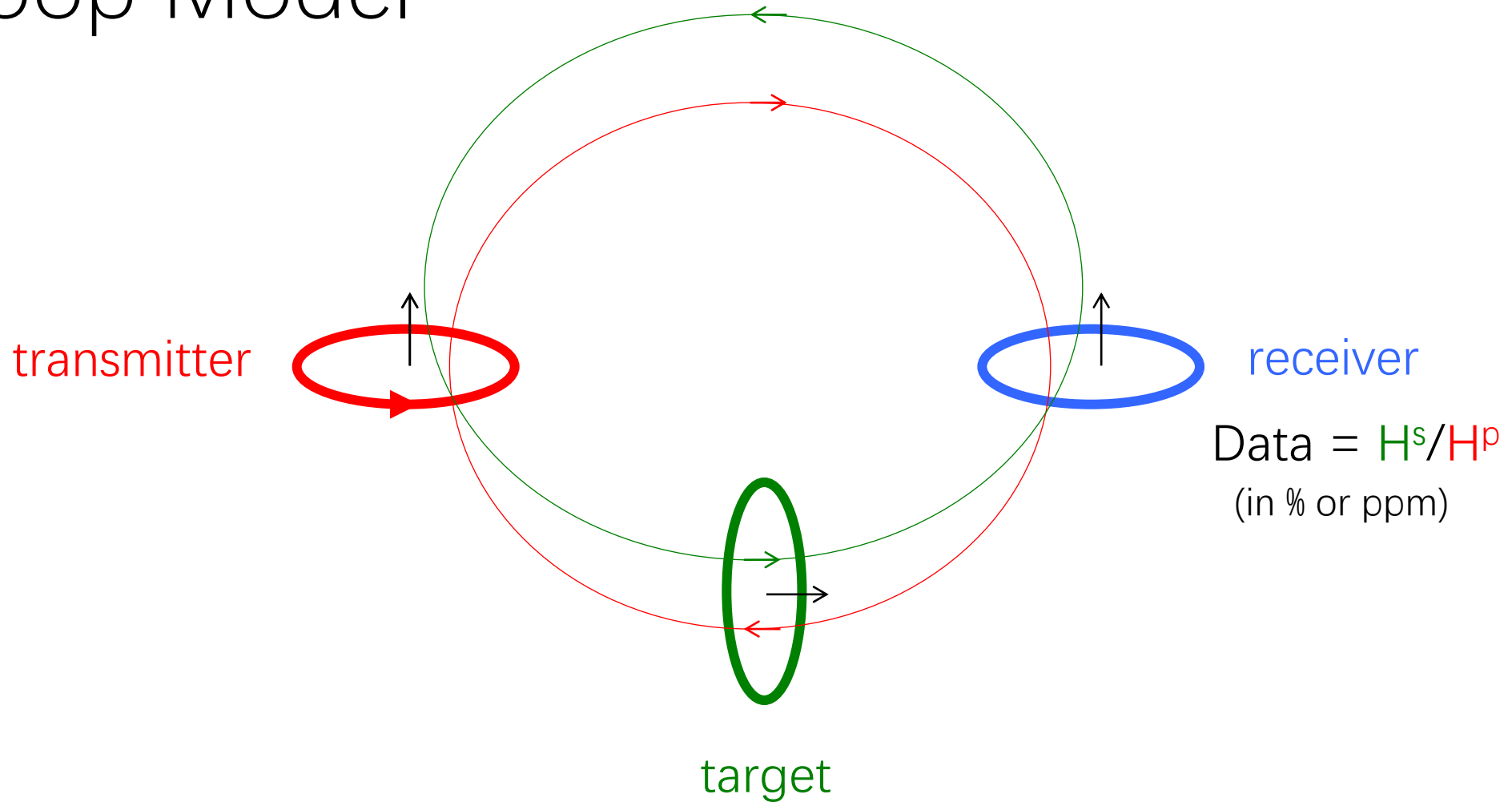
3-loop Model: Primary H^p



3-loop Model: Secondary H^s



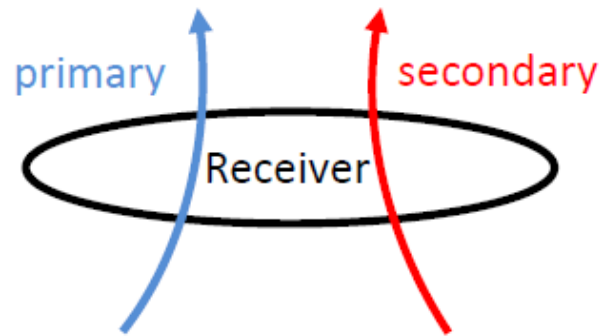
3-loop Model



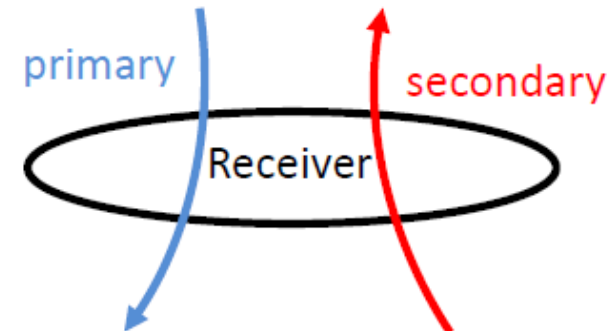
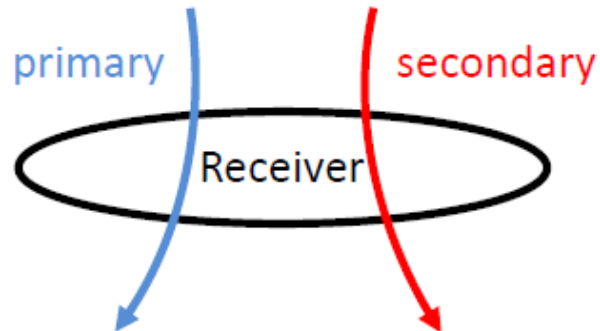
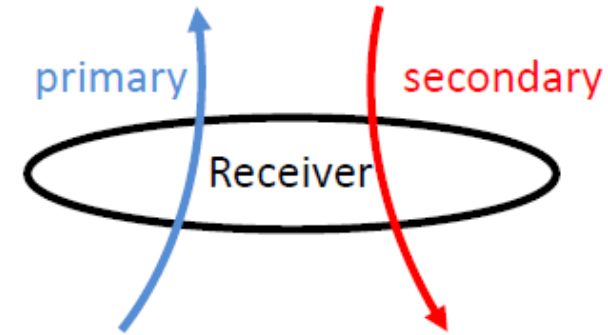
Question: Is the data positive or negative for the scenario on this page?
Hint: Think about the positive and negative anomalies in total field magnetics.

Data (Hs/Hp) Sign Convention

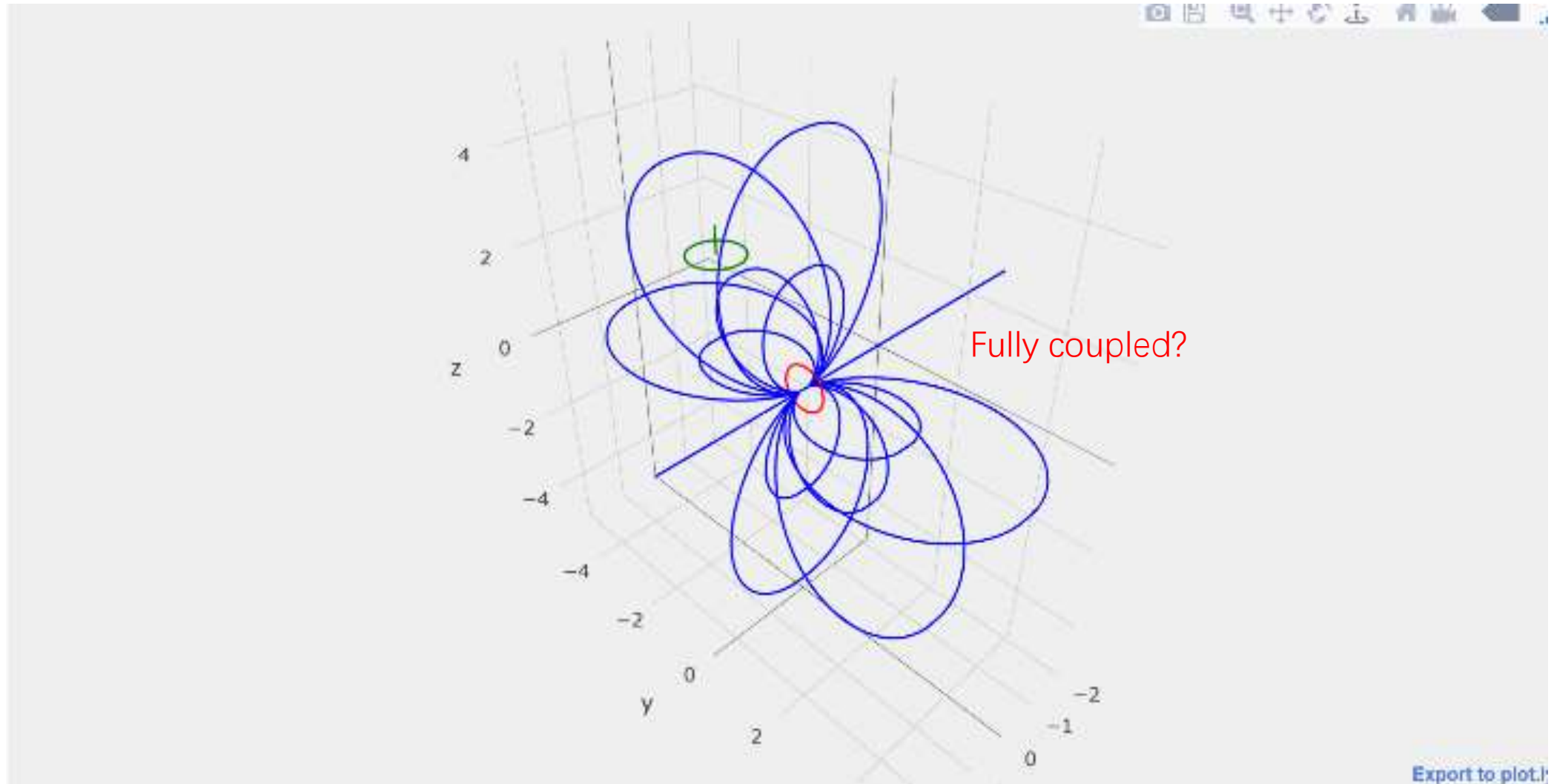
Positive
primary and secondary in same direction



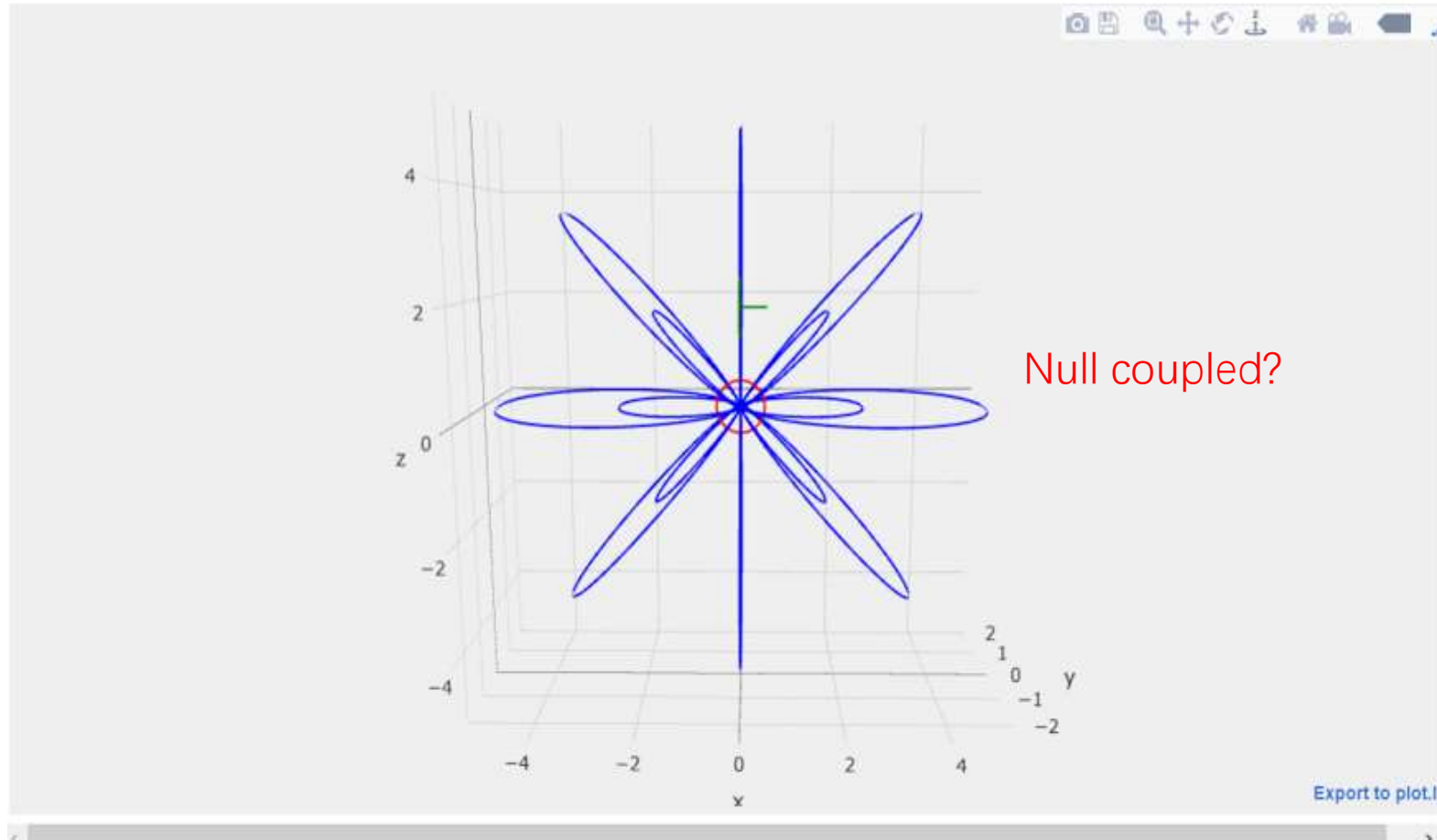
Negative
primary and secondary in opposite directions



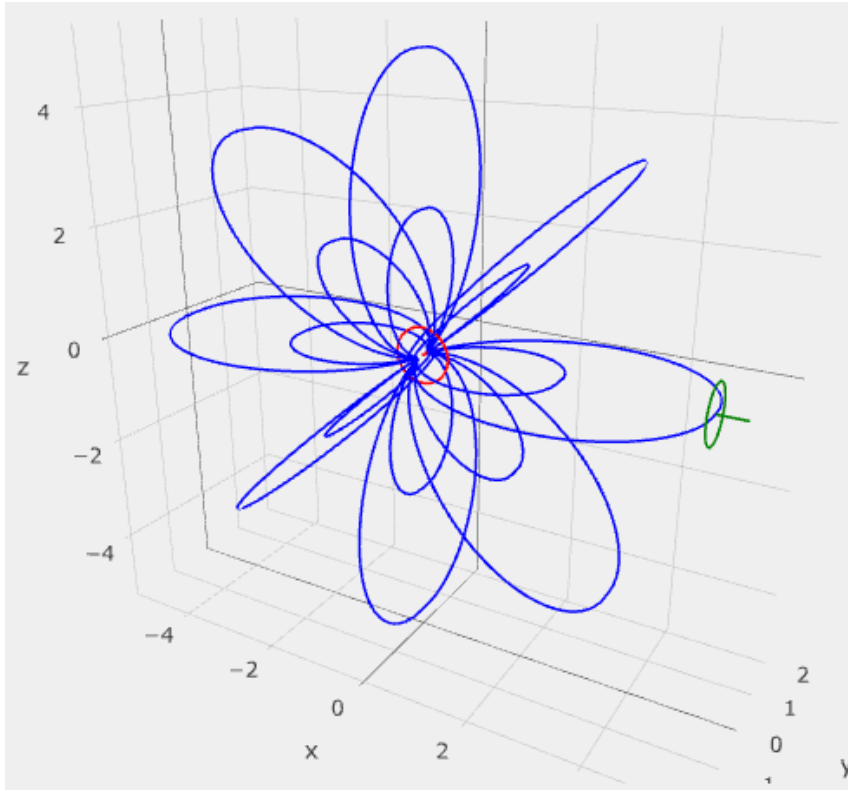
Coupling between Two Loops Through Magnetic Flux Linkage



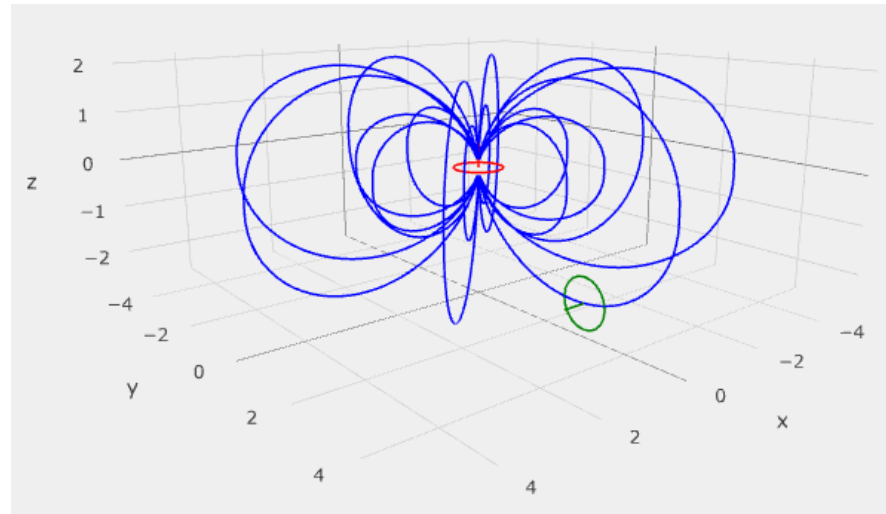
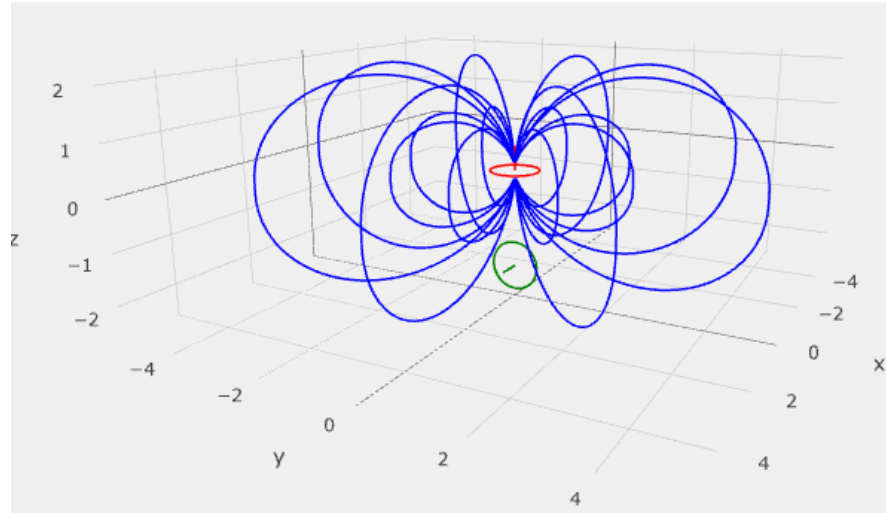
Coupling between Two Loops Through Magnetic Flux Linkage



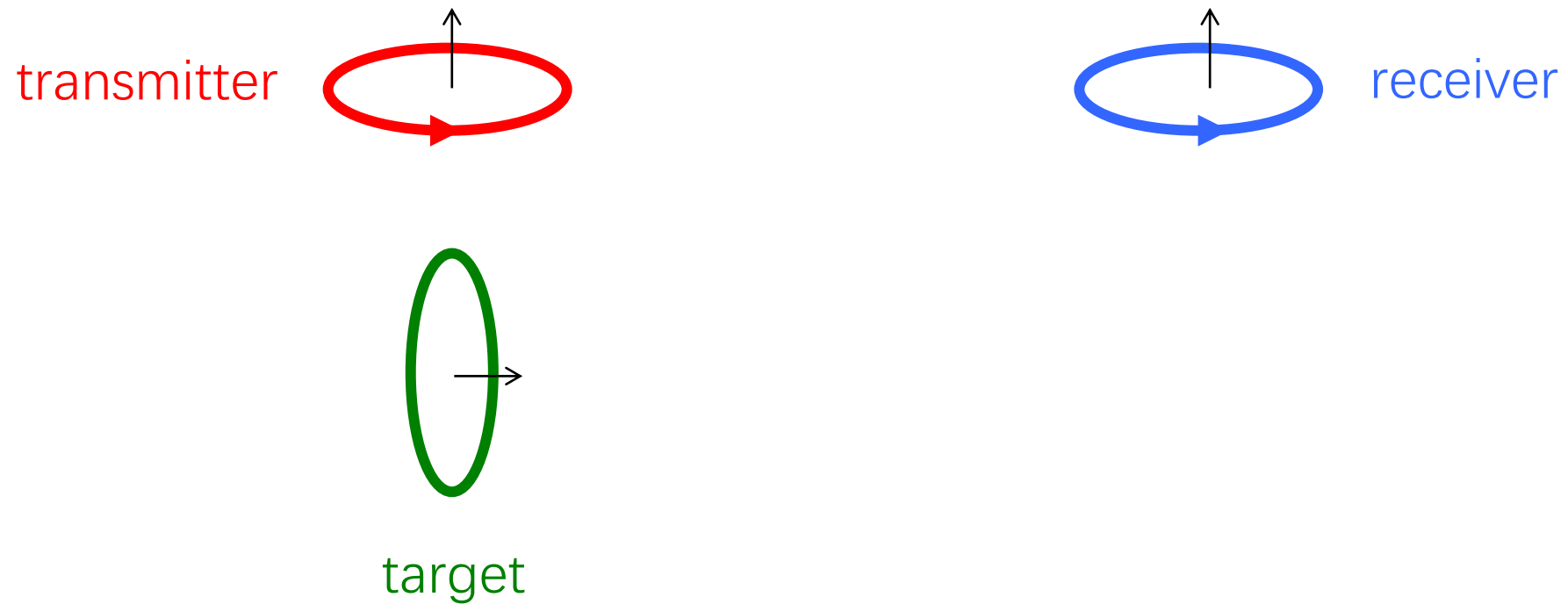
Coupling between Two Loops Through Magnetic Flux Linkage



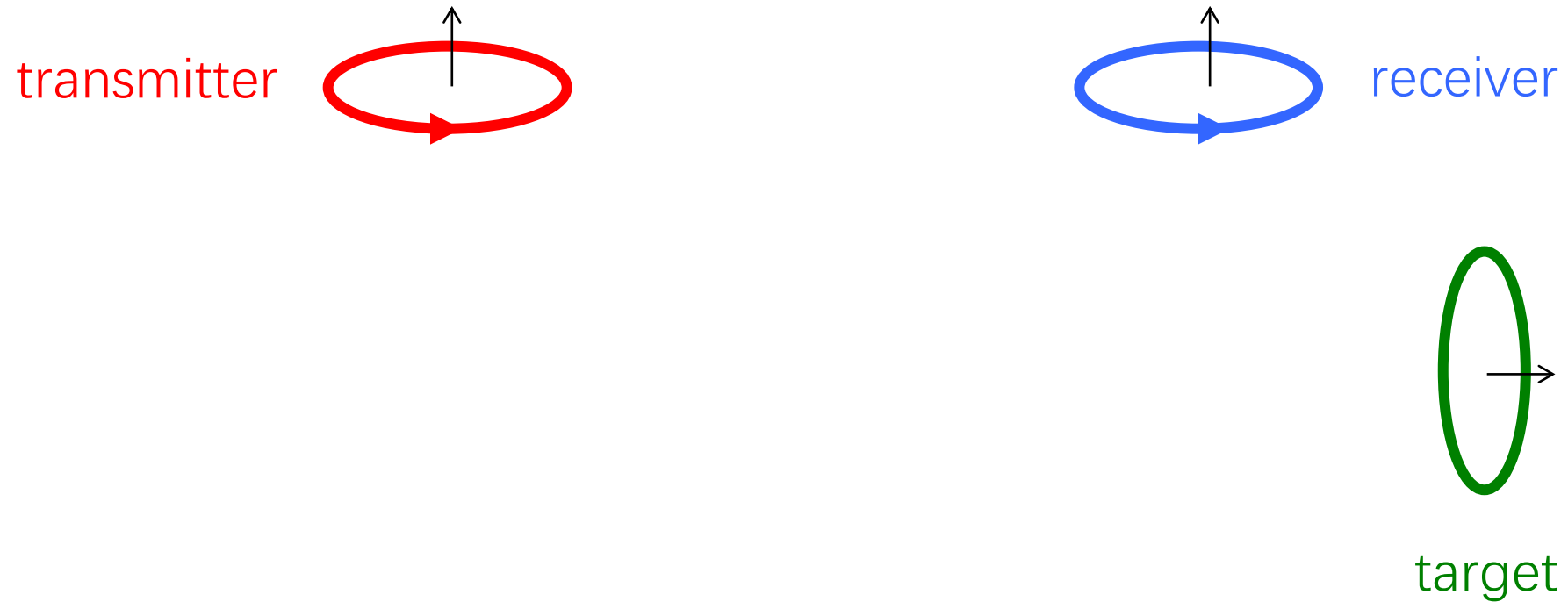
Null coupled



H^s/H^p : Positive or Negative?



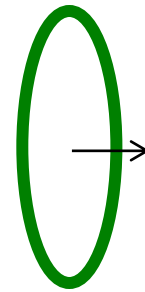
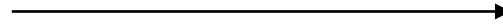
H^s/H^p : Positive or Negative?



Hs/Hp Profile



walk

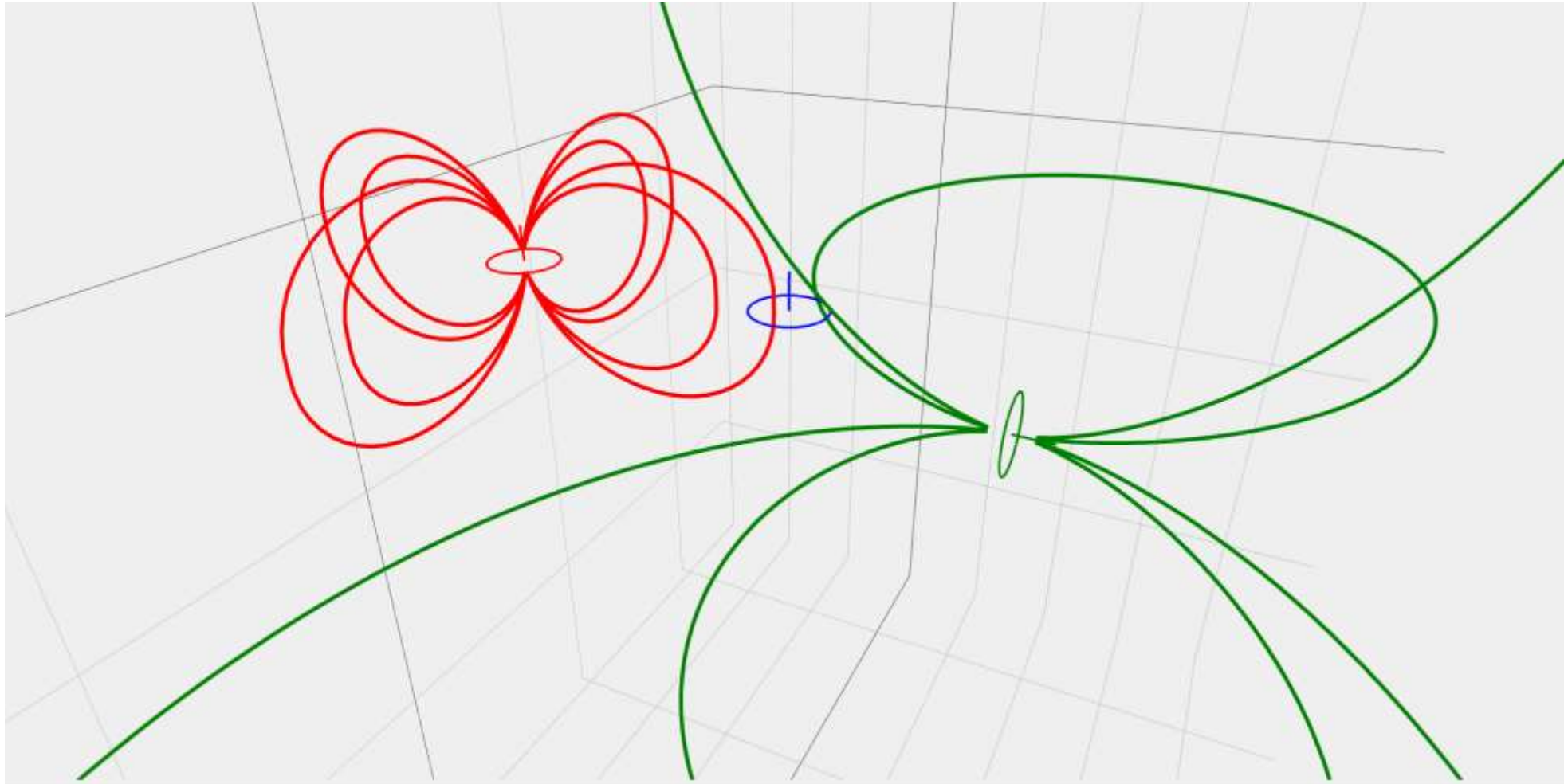


target

Verify using Demo Notebook

- “MagDipole3LoopsCoupling.ipynb”

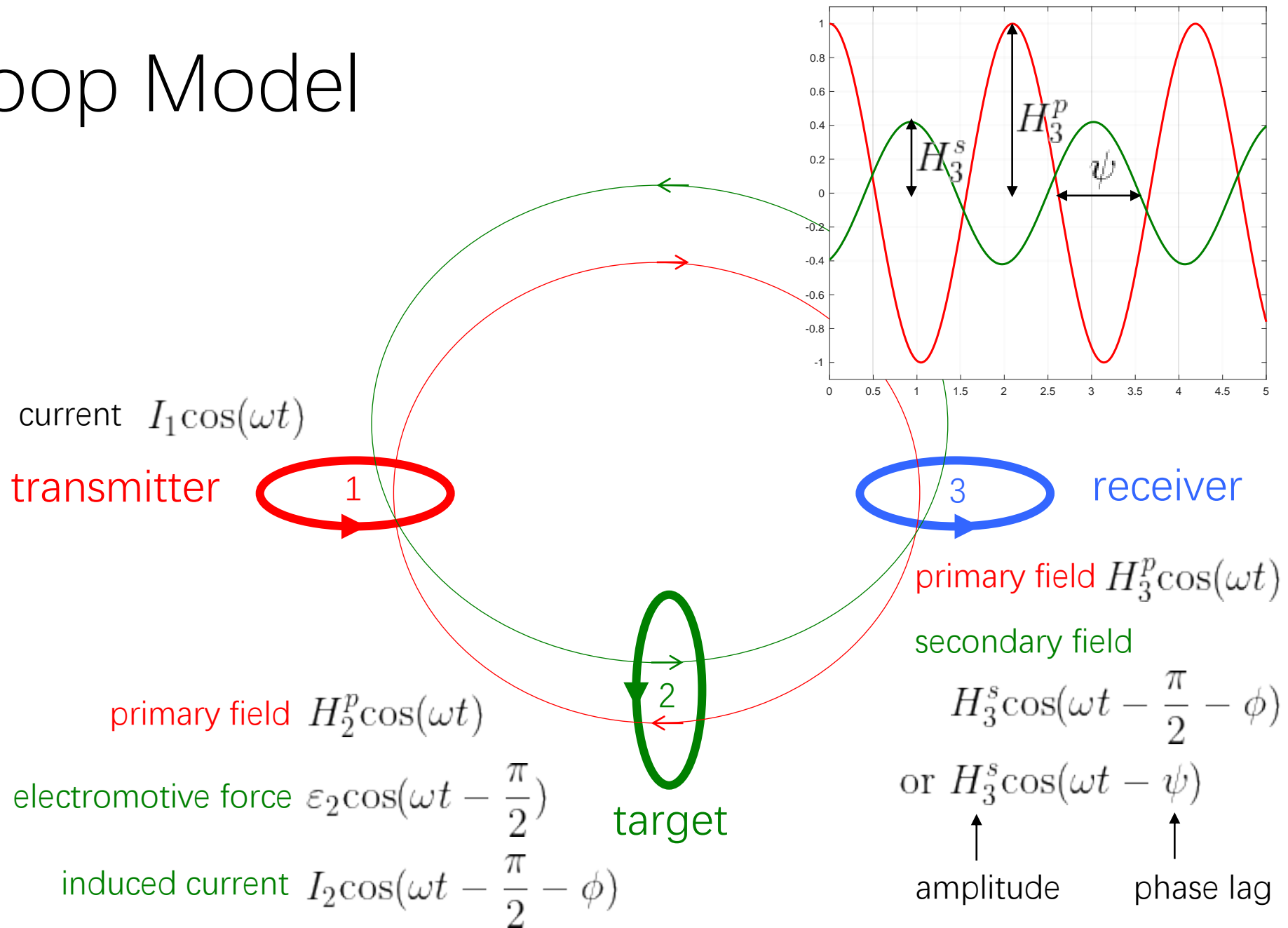
<https://github.com/sustechgem/geophysics-demo-notebooks/blob/master/MagDipole3LoopsCoupling.ipynb>



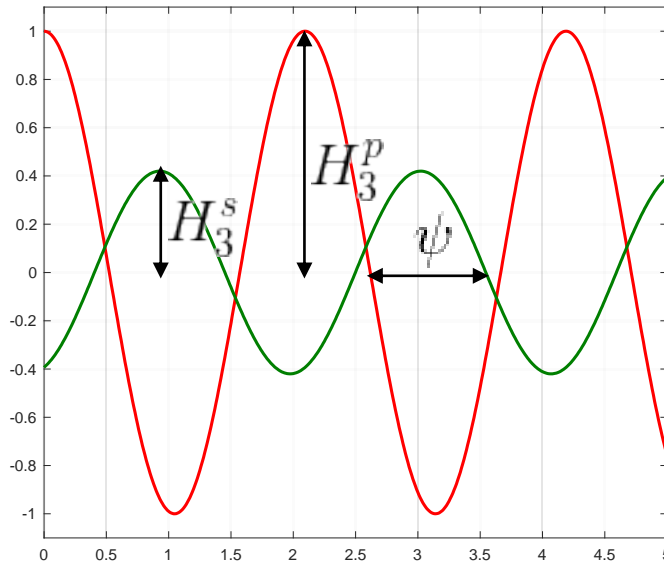
Drawing lines only helps qualitative understanding.

We need more math to do a quantitative interpretation.

3-loop Model



Decompose Secondary Field

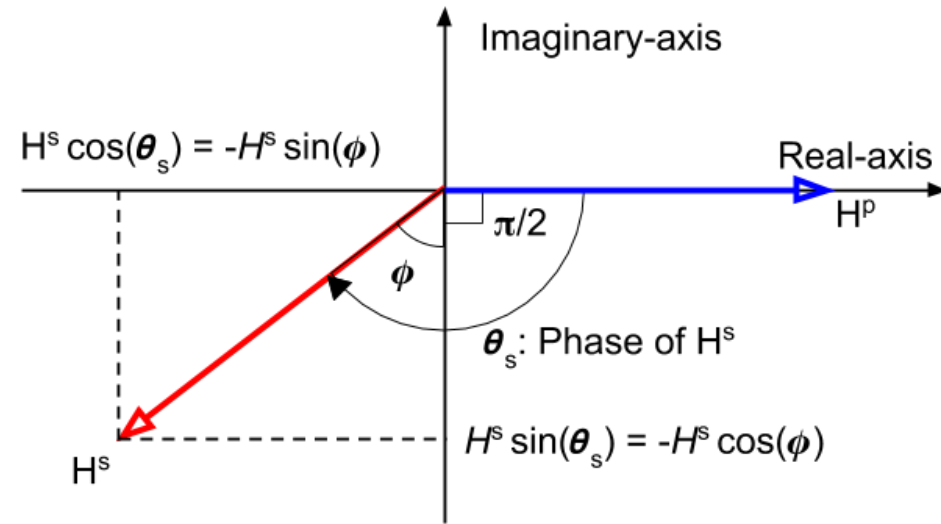


primary field $H_3^p \cos(\omega t)$

secondary field

$$H_3^s \cos(\omega t - \frac{\pi}{2} - \phi)$$

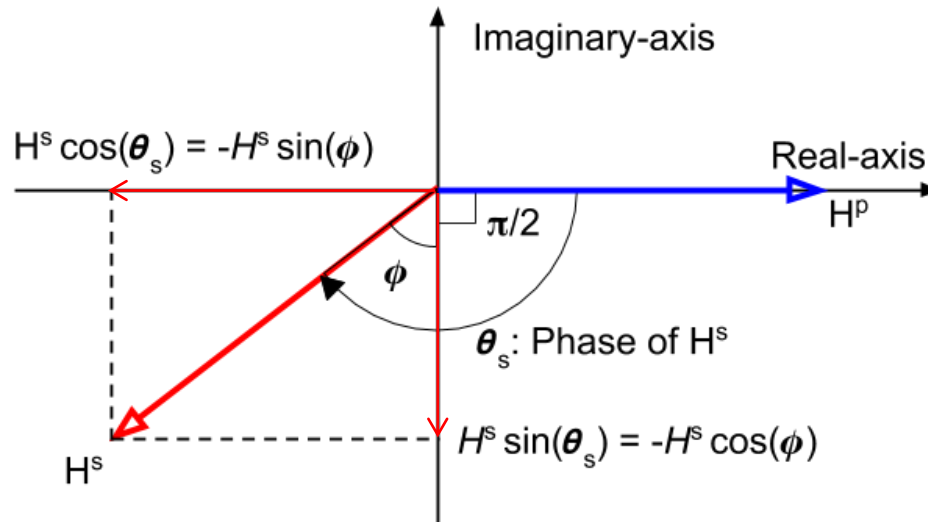
or $H_3^s \cos(\omega t - \psi)$



$$\phi = \tan^{-1}\left(\frac{\omega L}{R}\right) = \tan^{-1}(\alpha)$$

- H^s swings in the third quadrant: $0 < \phi < 90^\circ$
- ϕ depends on the induction number α
- α is a function of frequency ω , self inductance \mathbf{L} and resistance \mathbf{R} of Loop 2

Decompose Secondary Field



$$\phi = \tan^{-1}\left(\frac{\omega L}{R}\right) = \tan^{-1}(\alpha)$$

Question: What happens to the H^s (red arrow) for a very conductive or very resistive target?

Decompose H^s to two orthogonal components then normalize by H^p :

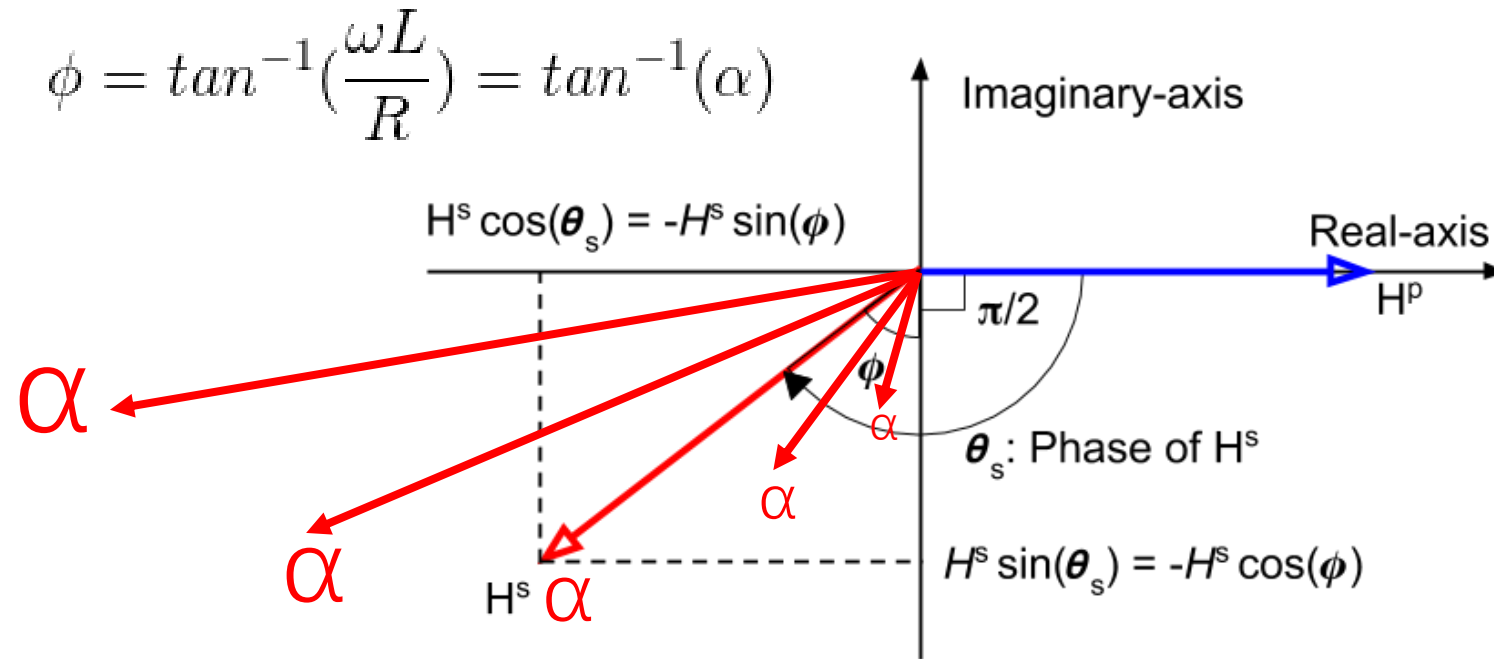
90° phase lag: called “out-of-phase”, “quadrature”, “imaginary”

$$\frac{H^s \cos(\phi)}{H^p}$$

180° phase lag: called “in-phase”, “real”

$$\frac{H^s \sin(\phi)}{H^p}$$

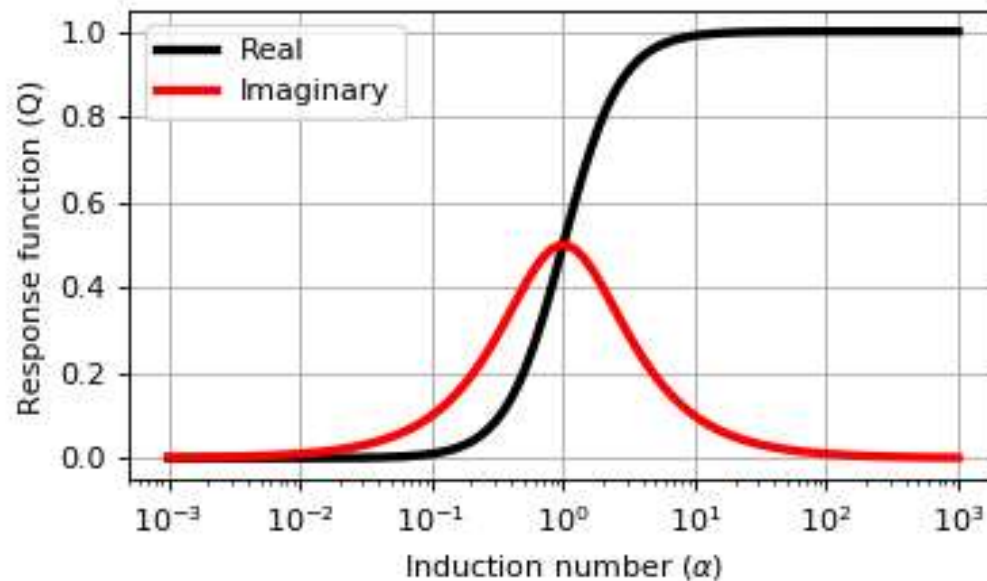
Response Function



Question: How would the real and imaginary data change with the induction number α ?

Response Function

$$Q(\alpha) = \frac{i\alpha}{1 + i\alpha} = \frac{\alpha^2 + i\alpha}{1 + \alpha^2} \quad \alpha = \frac{\omega L}{R}$$



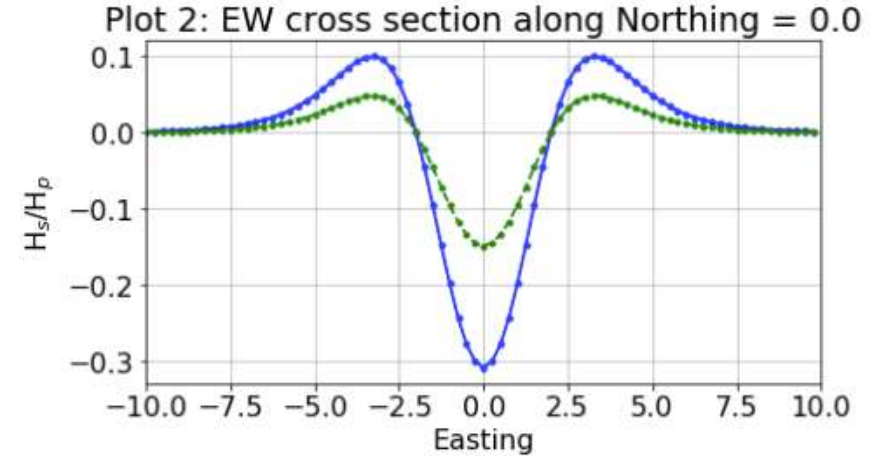
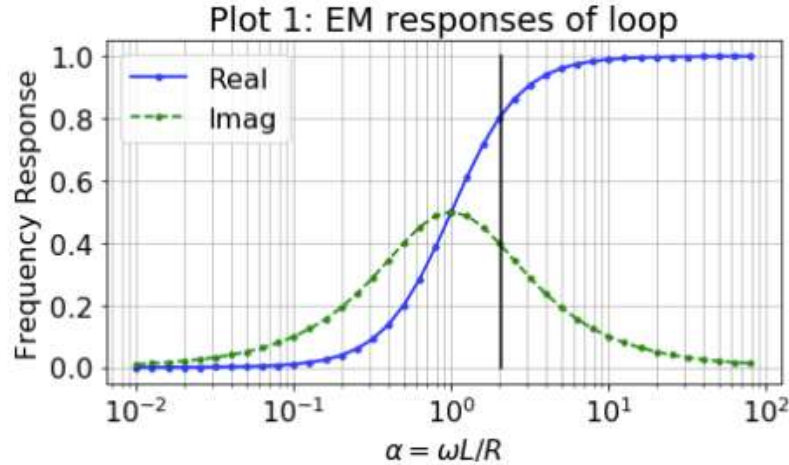
Resistive limit:

- low frequency
- low conductivity

Inductive limit:

- high frequency
- high conductivity

Expected Data From a Loop Target



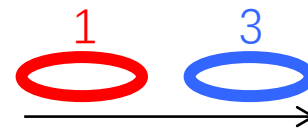
$$\frac{H_3^s}{H_3^p} = -\frac{M_{12}M_{23}}{M_{13}L} \left[\frac{\alpha^2 + i\alpha}{1 + \alpha^2} \right]$$

Coupling

- location, orientation
- overall magnitude

Induction

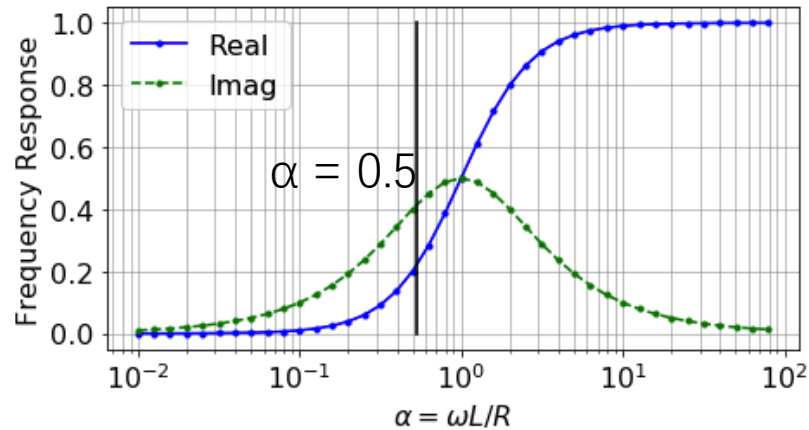
- properties of loop 2
- how much in Re & Im



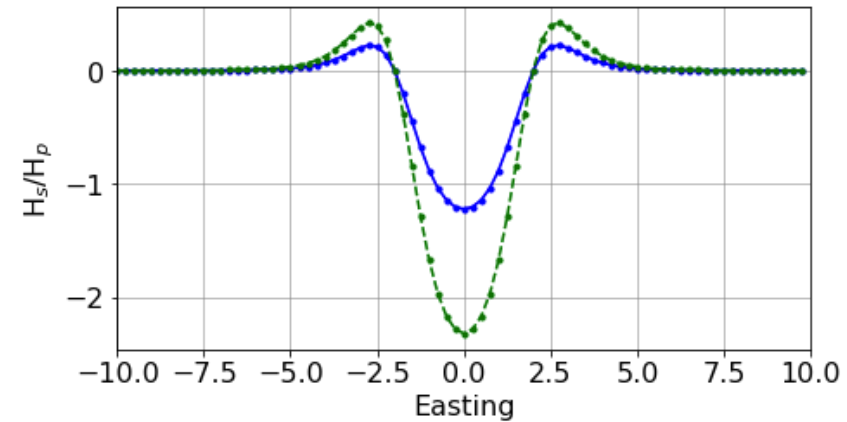
2

A Smaller Induction Number

Response function of different α



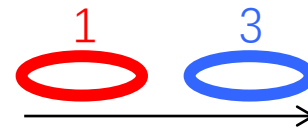
Data along a profile for $\alpha = 0.5$



$$\frac{H_3^s}{H_3^p} = -\frac{M_{12}M_{23}}{M_{13}L} \left[\frac{\alpha^2 + i\alpha}{1 + \alpha^2} \right]$$

Coupling
- location, orientation
- overall magnitude

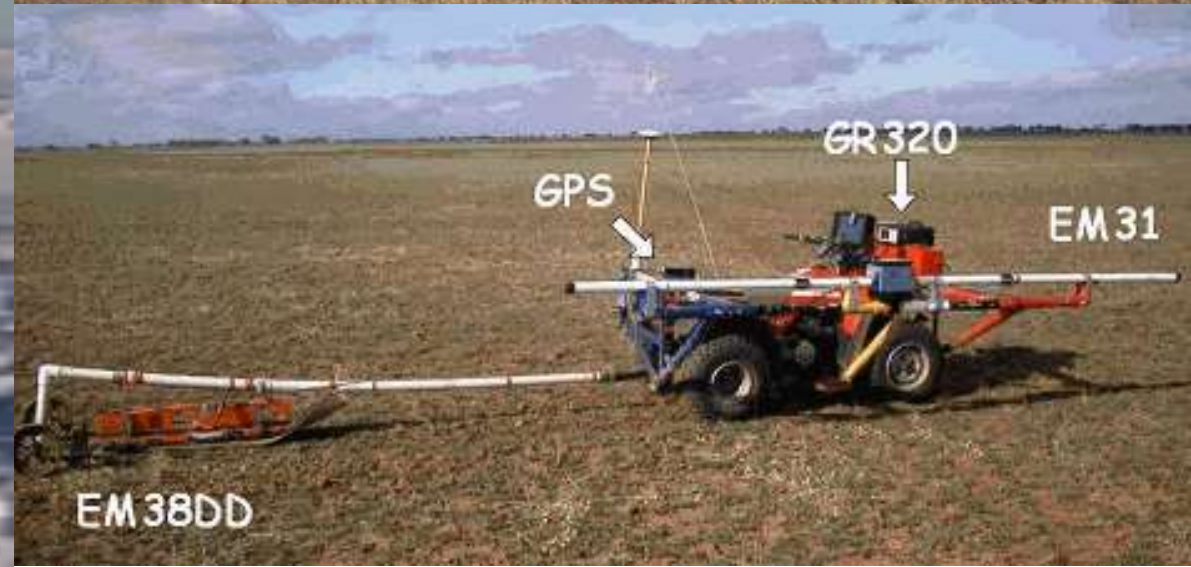
Induction
- properties of loop 2
- how much in Re & Im



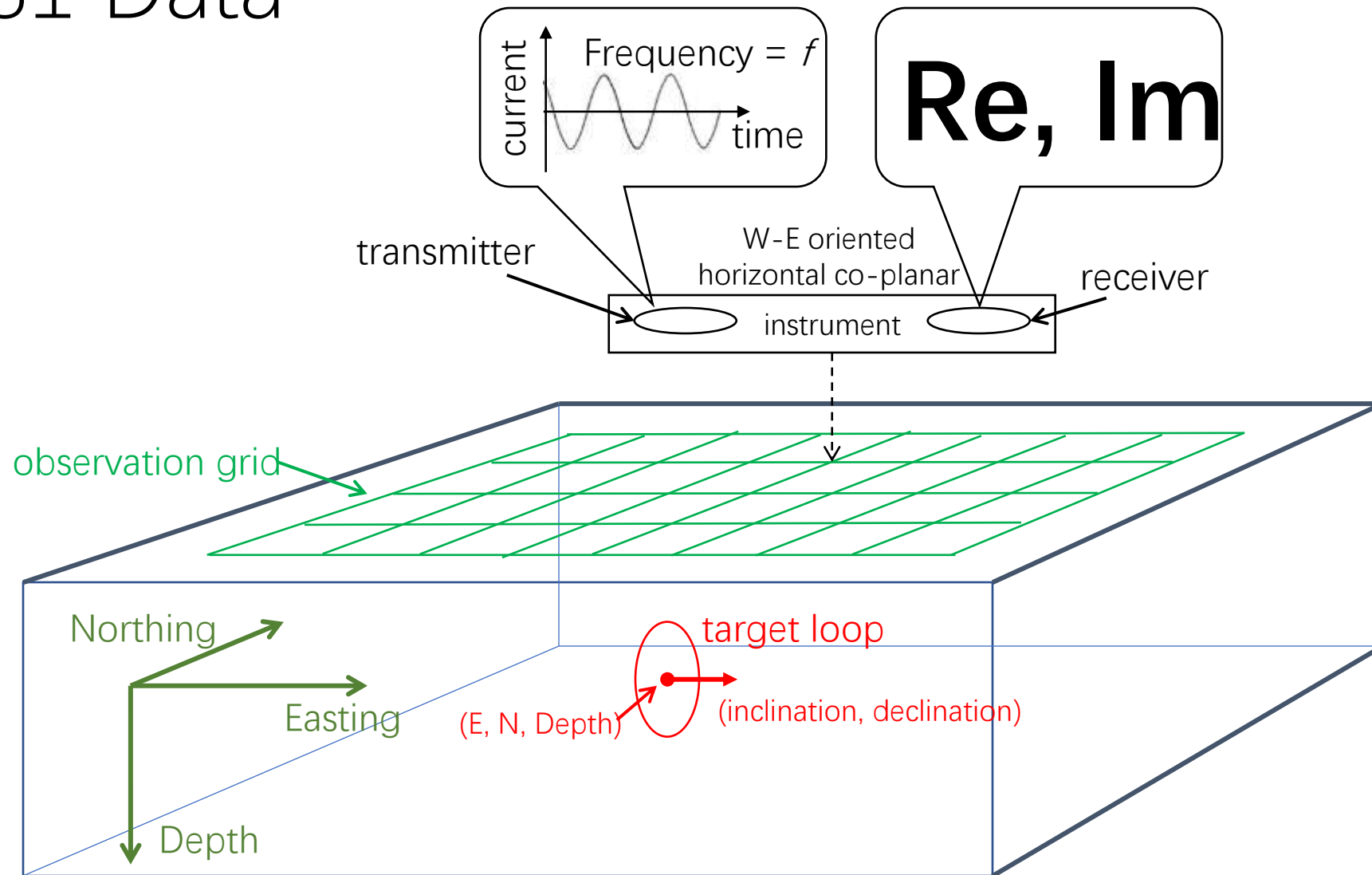
2
0

EM-31

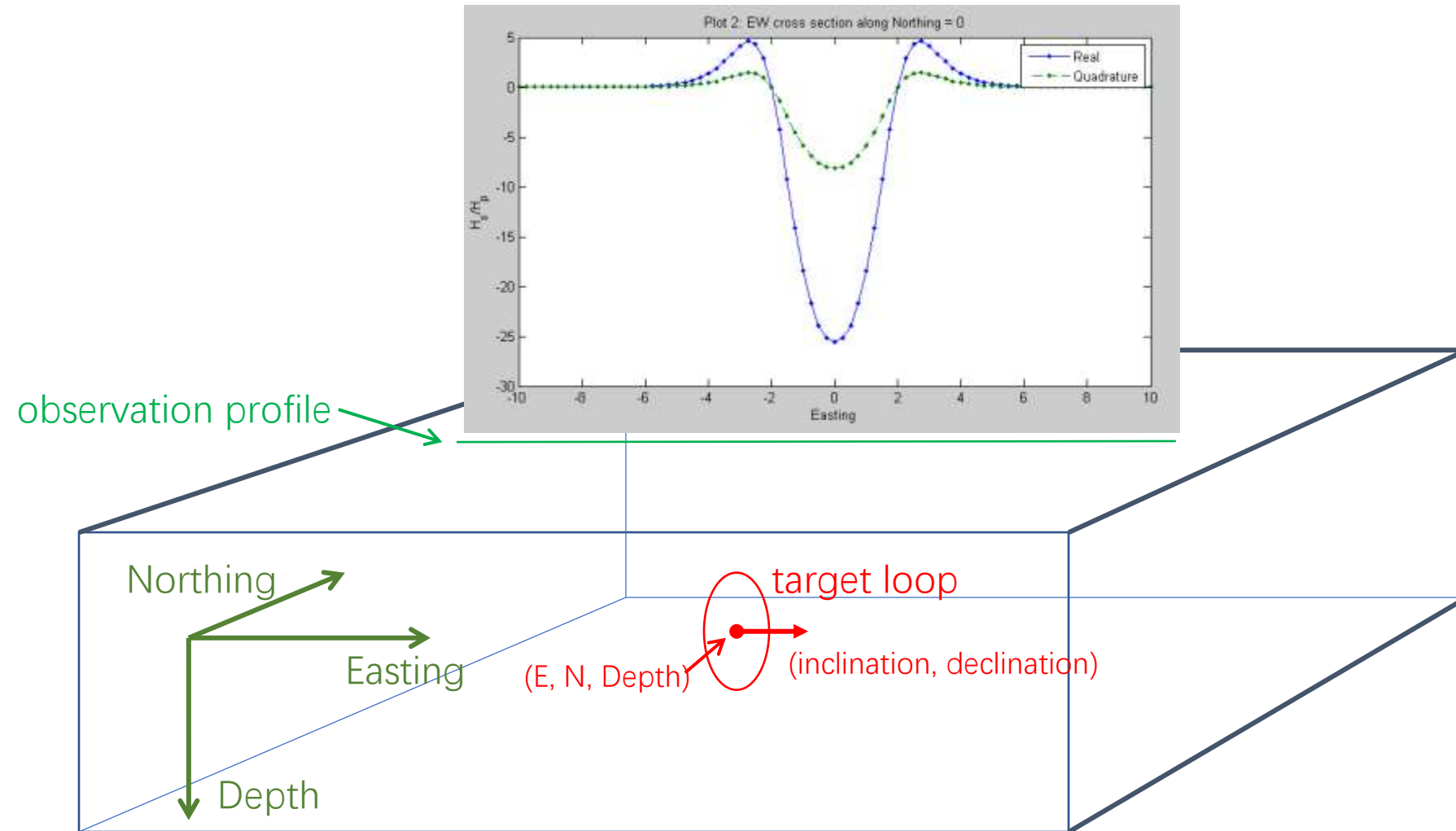
- Frequency = 9.8 kHz
- Tx-Rx spacing = 3.66 m
- Horizontal or vertical coplanar
- “Ground conductivity meter”



EM-31 Data

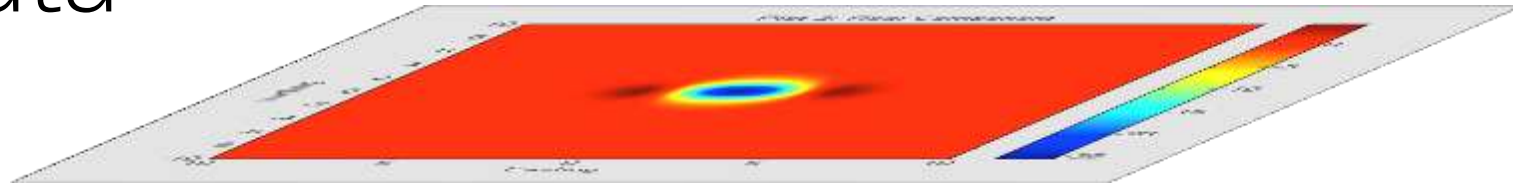


EM-31 Data

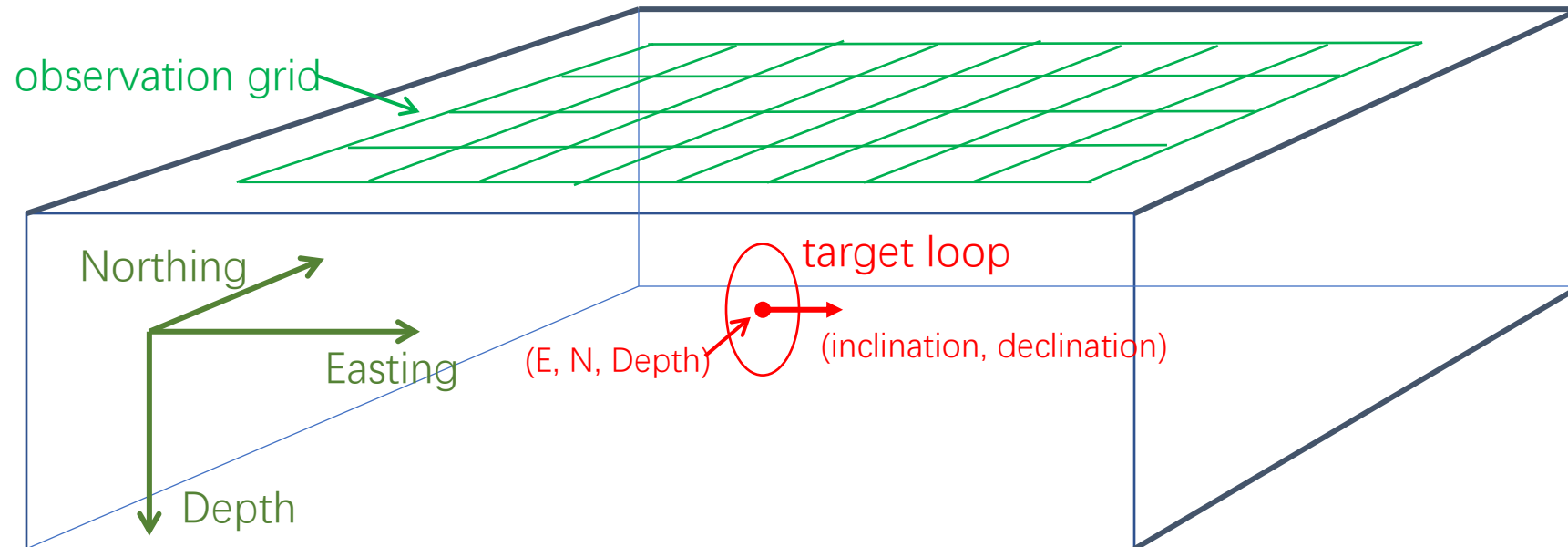
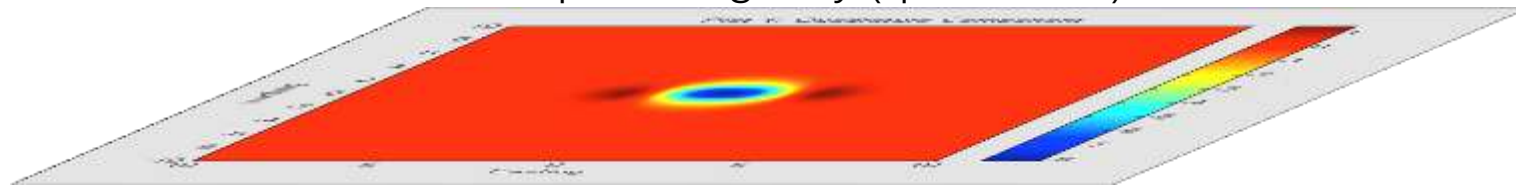


EM-31 Data

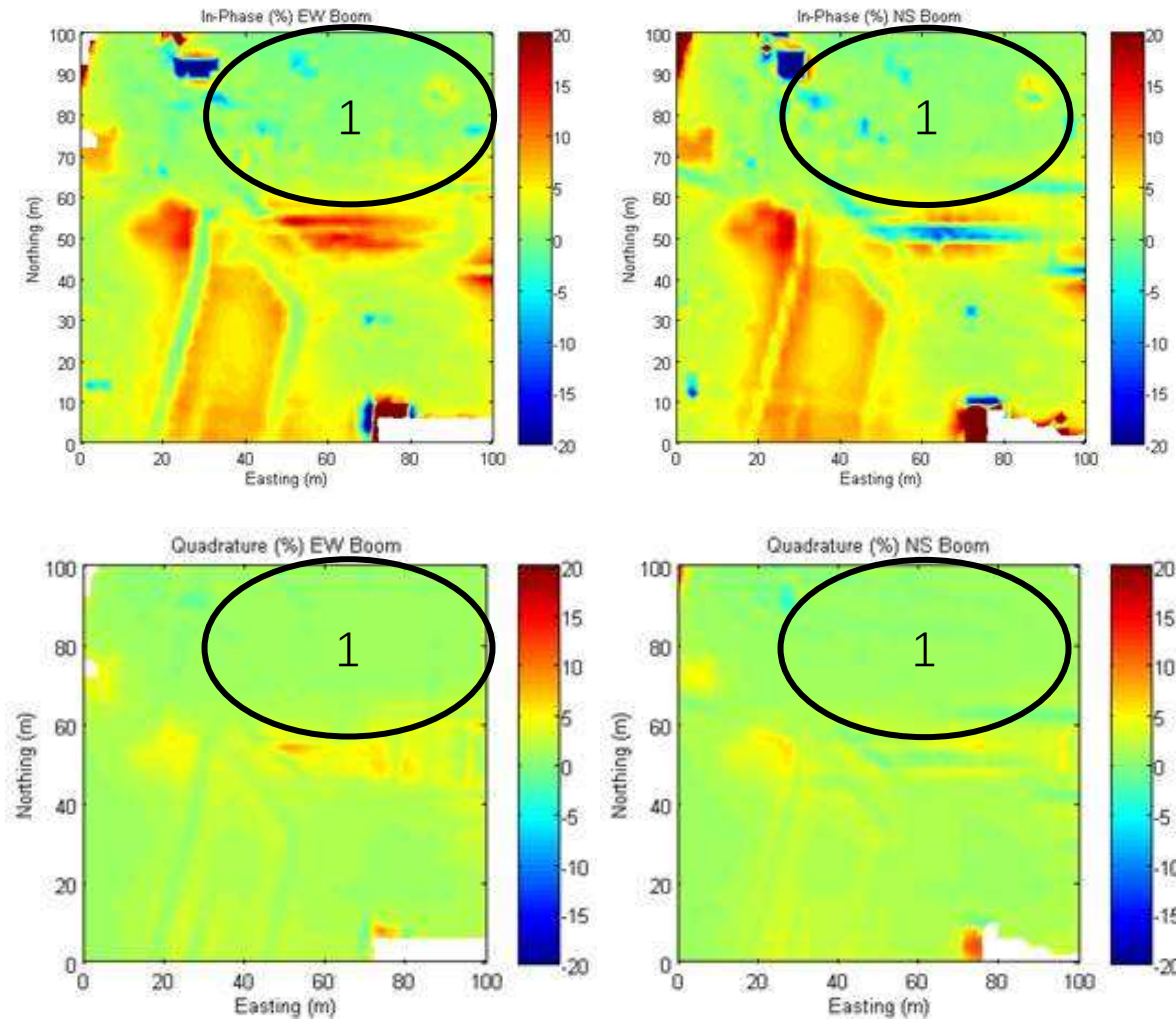
Map of real (in-phase) data



Map of imaginary (quadrature) data

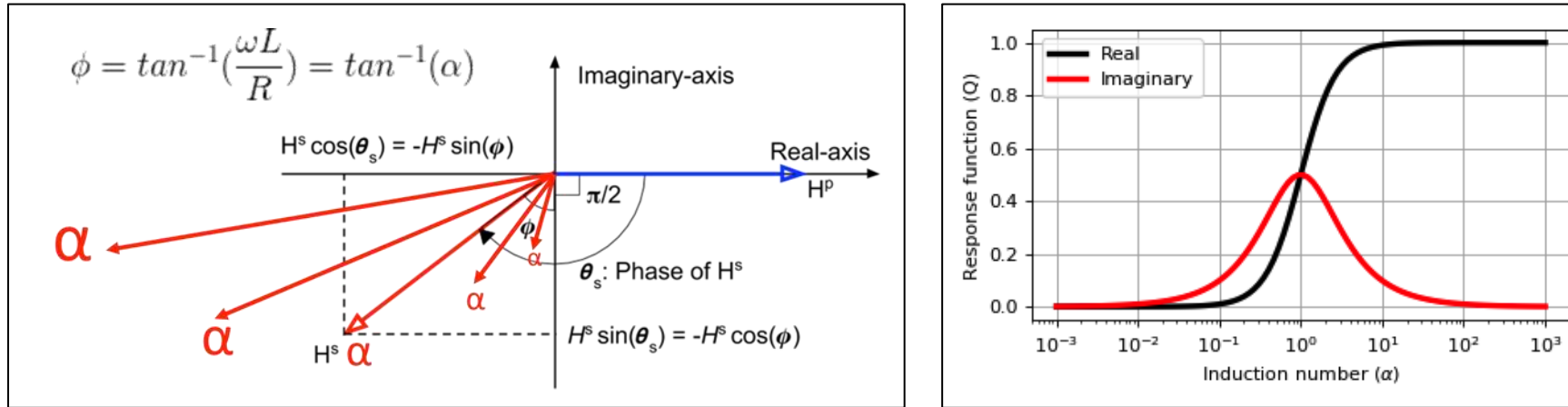


EM-31 Field Data



Data Feature 1:
Uniform, smooth and
small

EM-31 Data at Low Induction



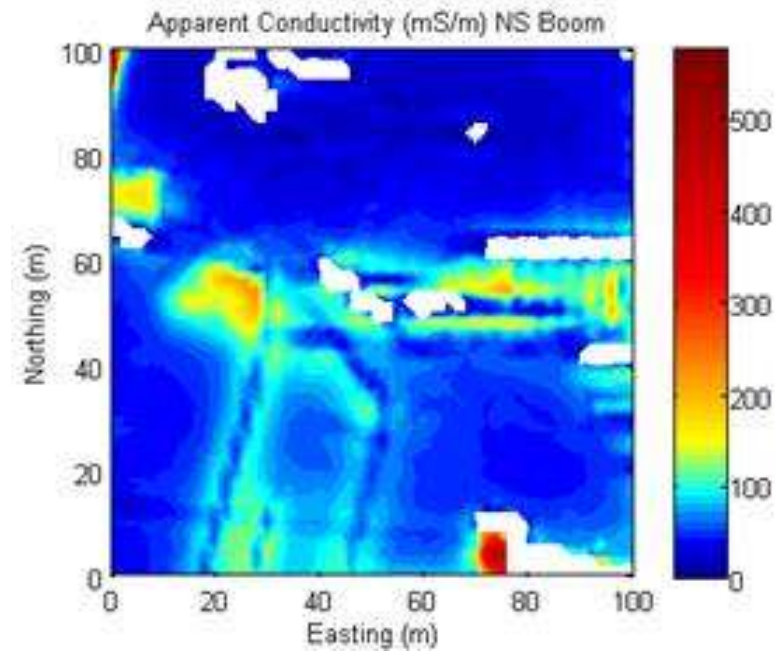
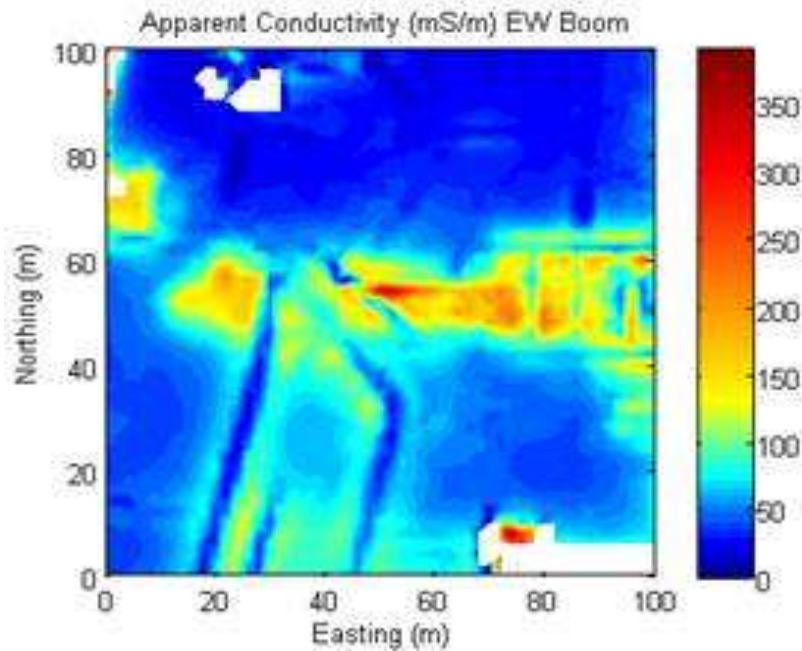
Small **Re** and small **Im** on the data maps, **α** big or small?

Low induction number:

- \mathbf{H}^s data mostly in quadrature, $\mathbf{Im} > \mathbf{Re} \approx 0$
- Very small induced current
- Subdivide the earth into many pieces; each piece interacts with Tx-Rx independently without interaction between any two pieces (**recall low induced magnetization in magnetics, easy calculation using superposition!**)

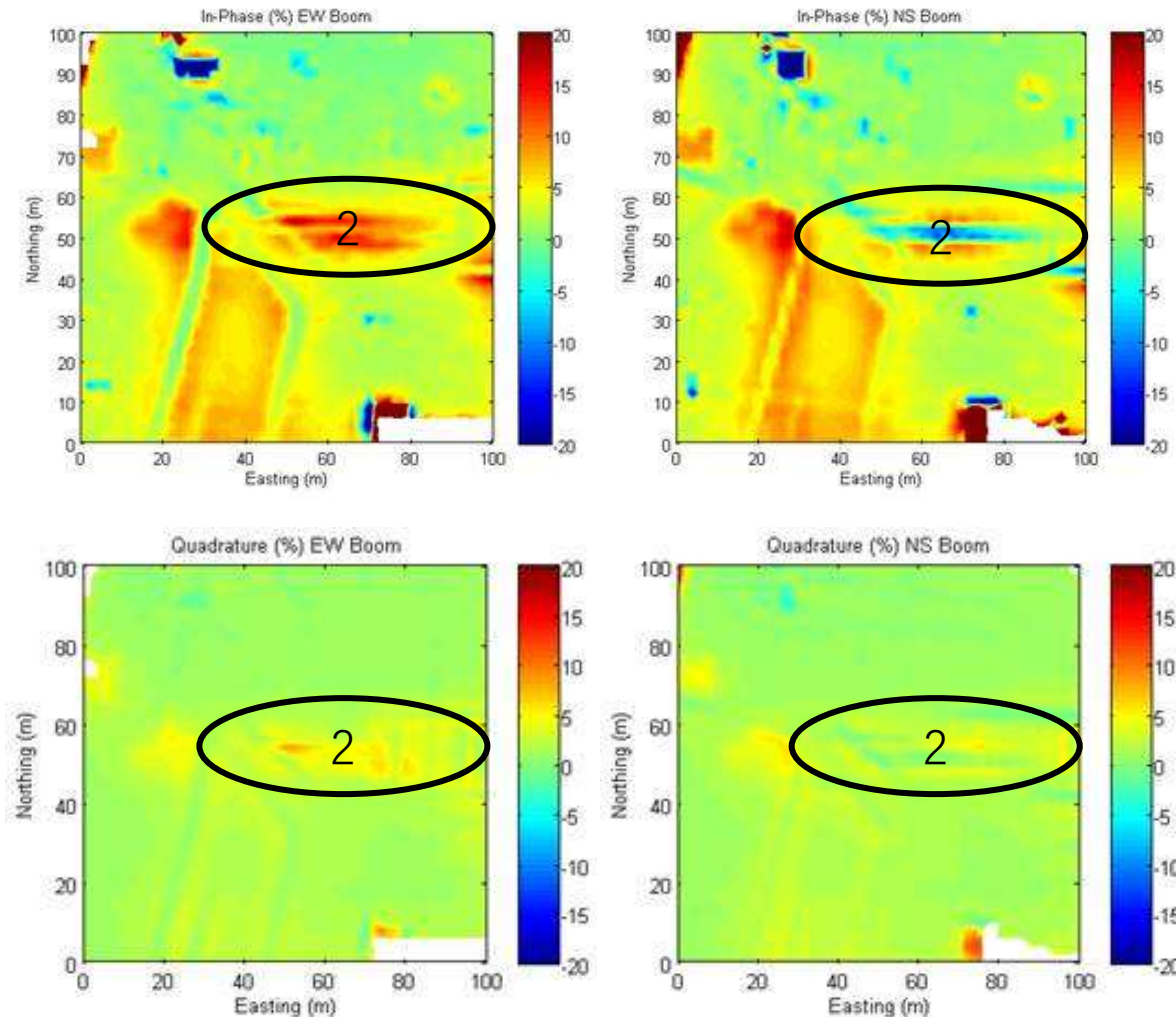
Apparent Conductivity

$$\sigma_a = \frac{4}{\omega \mu_0 s^2} \text{Im}$$



Question: Which area on the maps is the most likely to have a reliable estimate of the ground conductivity?

EM-31 Data Interpretation



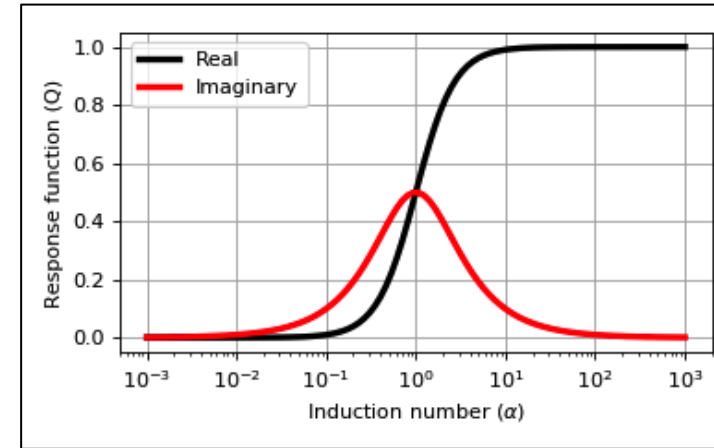
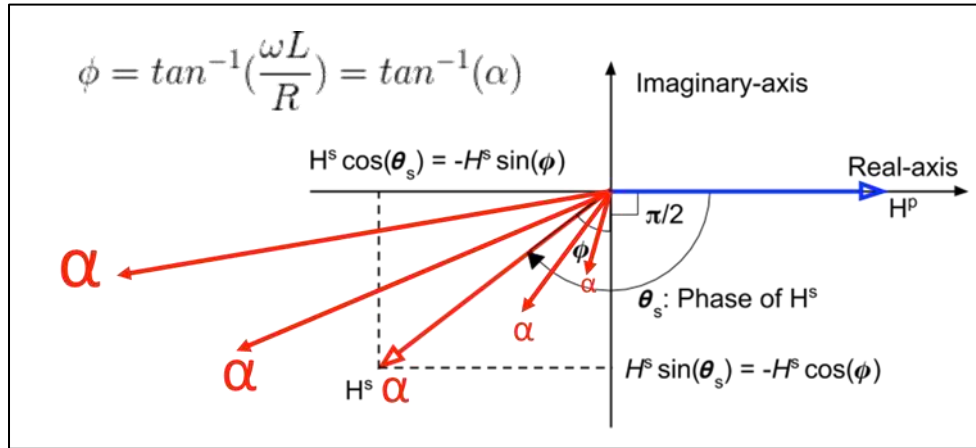
Data Feature 1:

Uniform, smooth and small

Data Feature 2:

Abrupt change
Positive and negative
Large **Re** and small **Im**

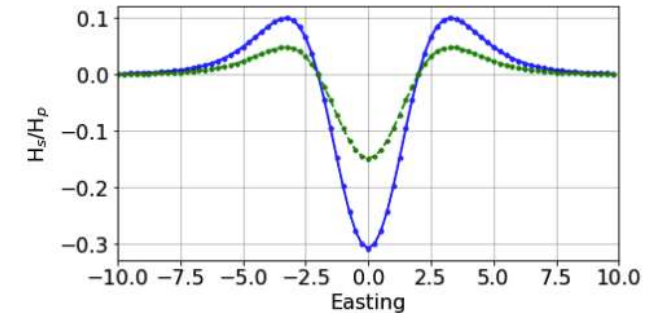
EM-31 Data at High Induction



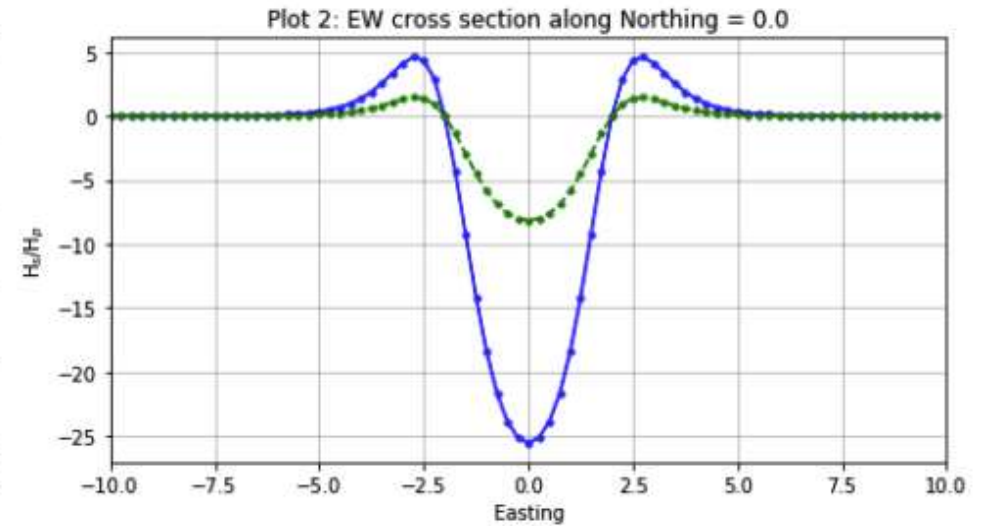
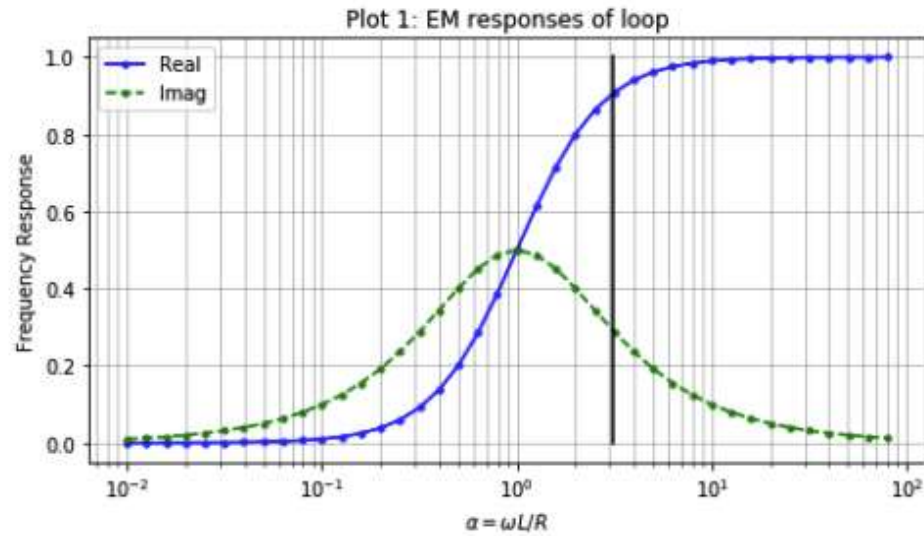
Large **Re** and small **Im** on the data maps, α big or small?

High induction number:

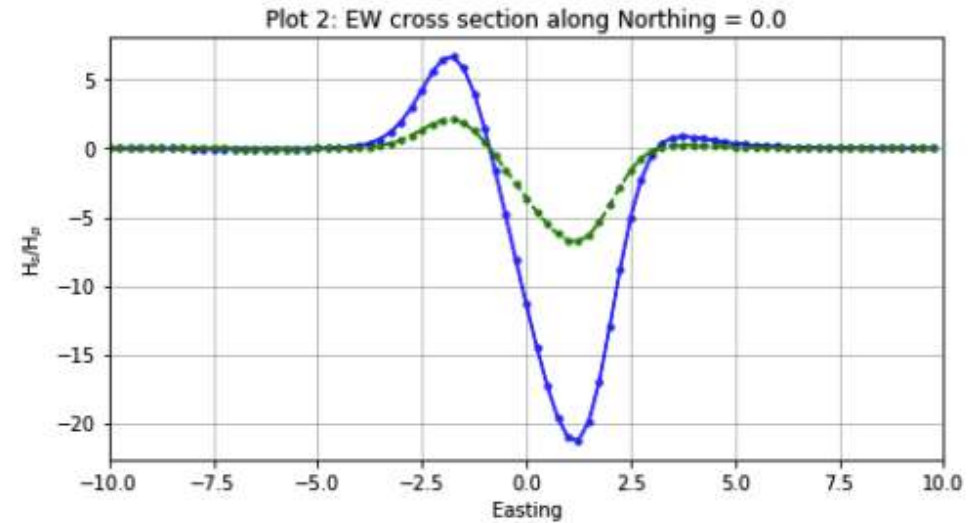
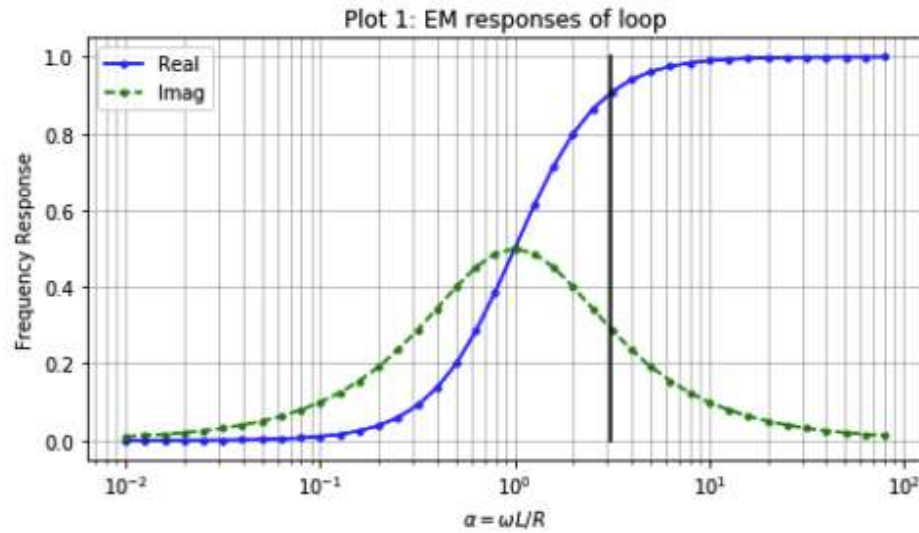
- H^s data mostly in in-phase, **Re** > **Im** ≈ 0
- Very strong induced current
- Cannot use apparent conductivity, but if the target is a good compact conductor, use the 3-loop model



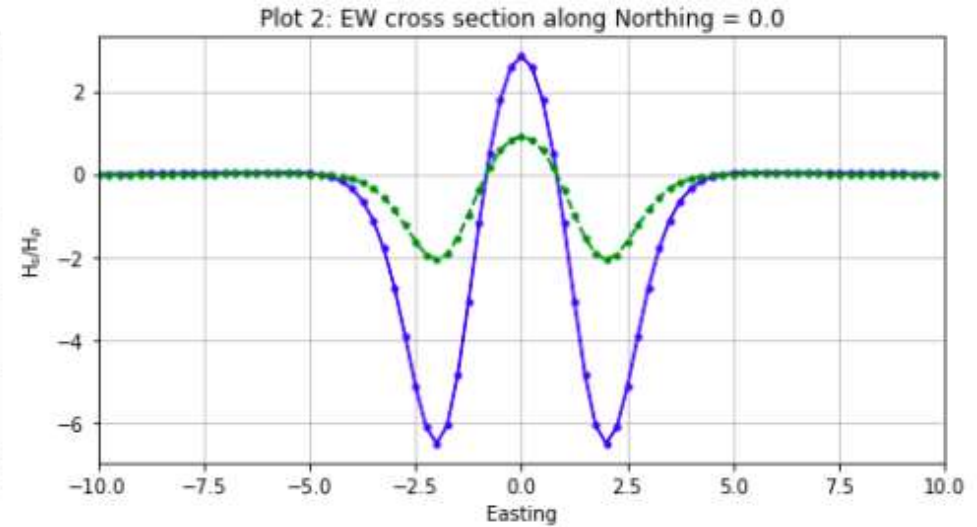
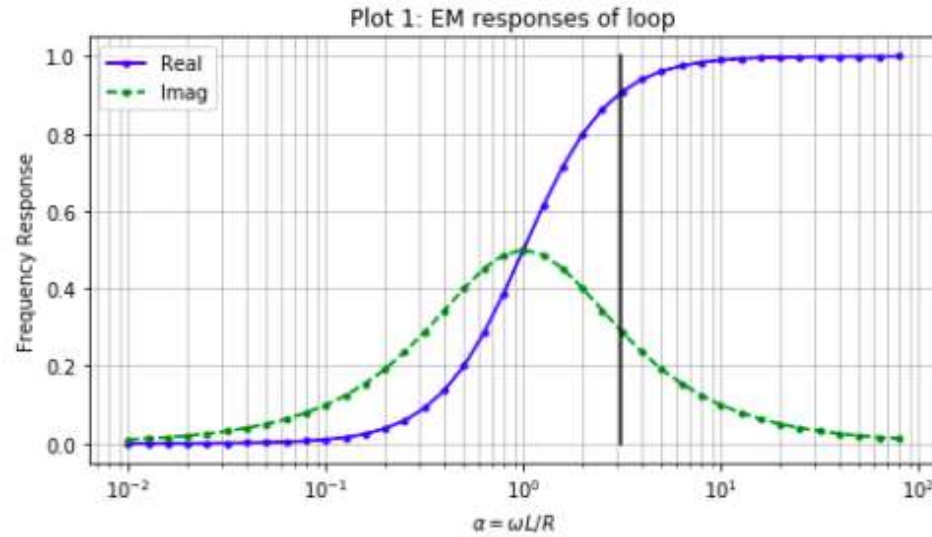
Vertical Target Loop



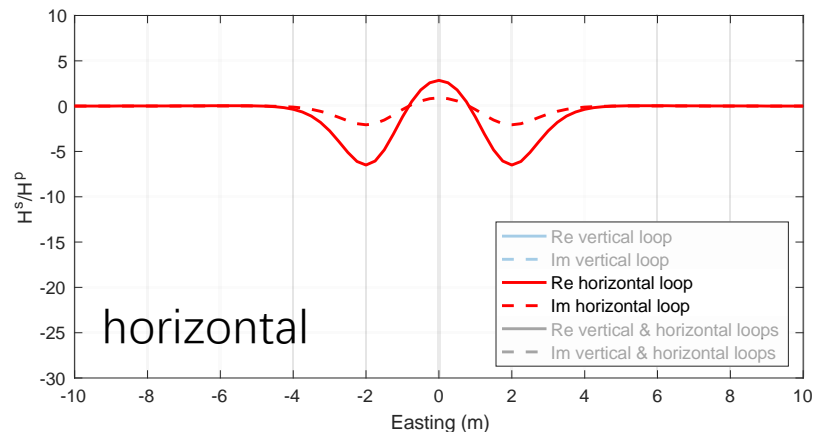
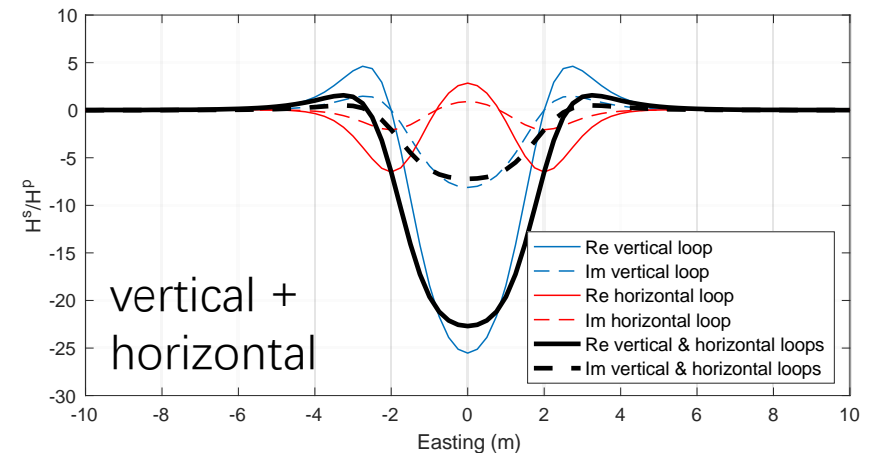
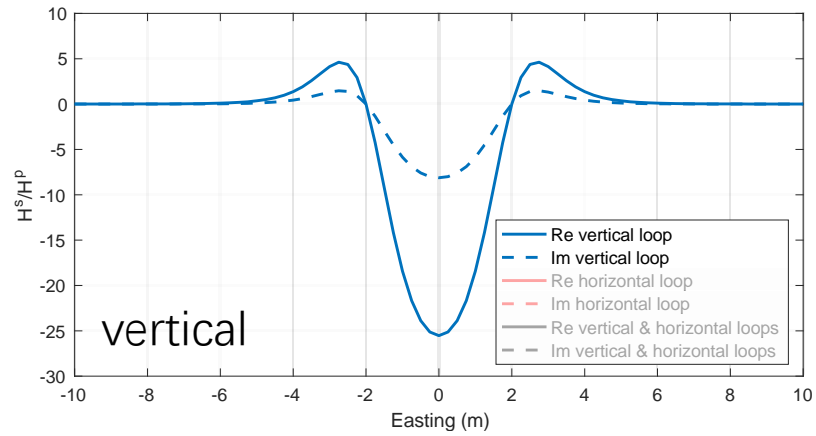
45 Degree Dipping Target Loop

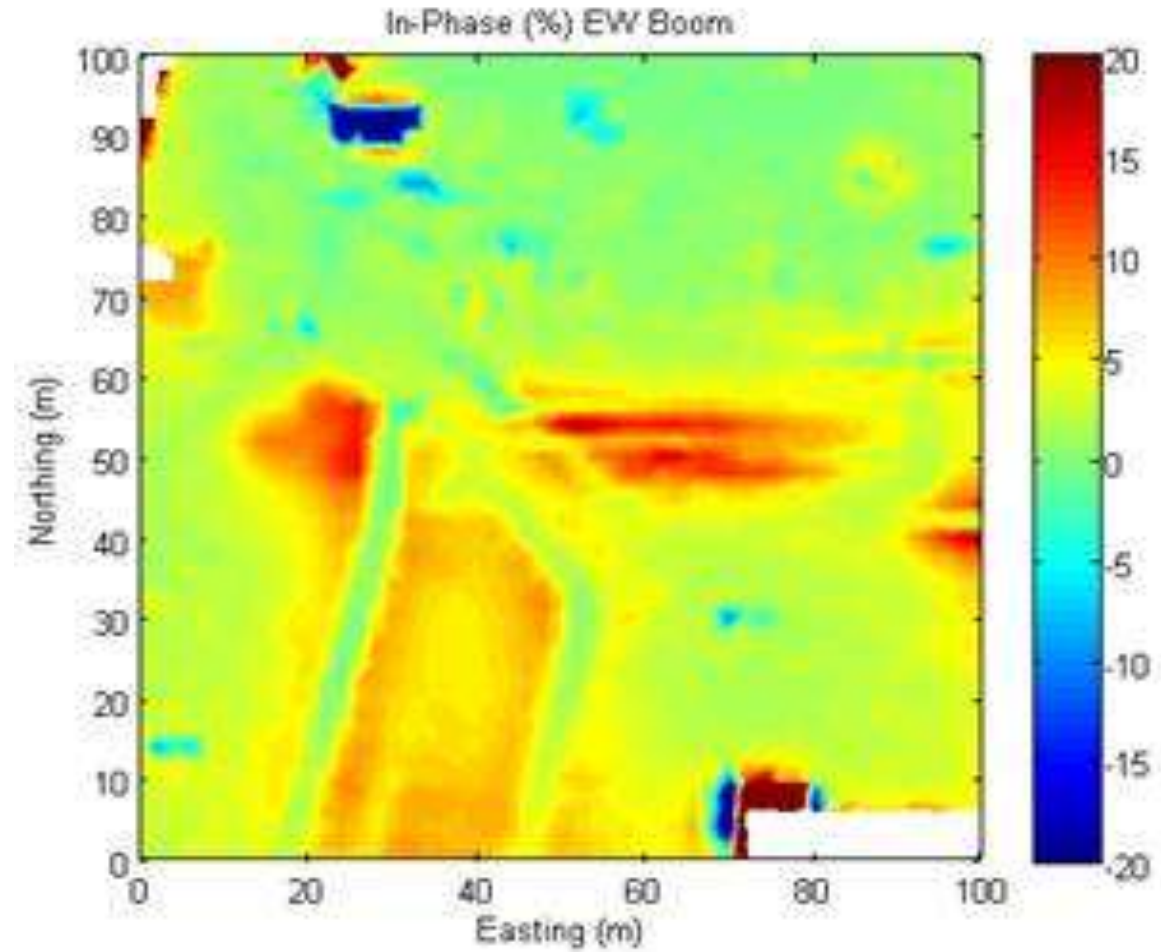
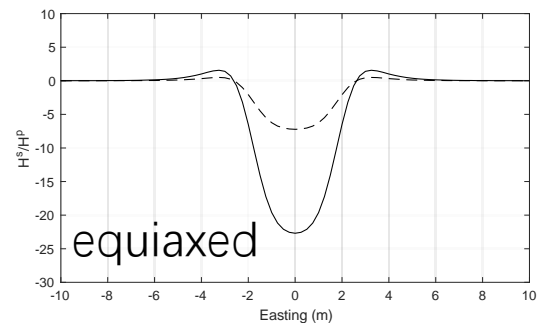
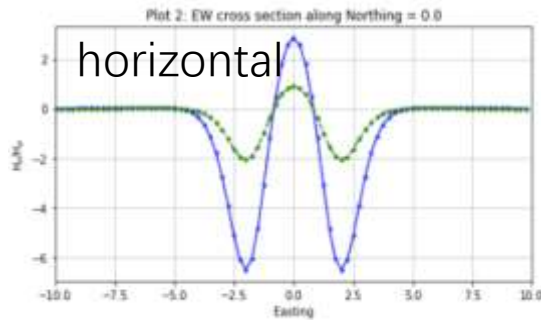
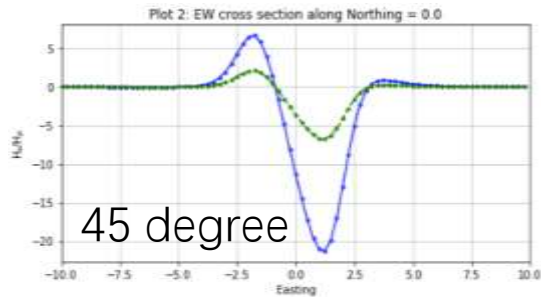
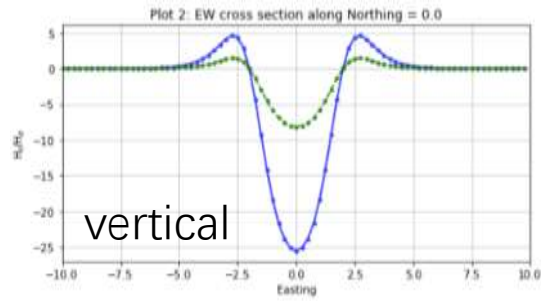


Horizontal Target Loop



Equiaxed Target





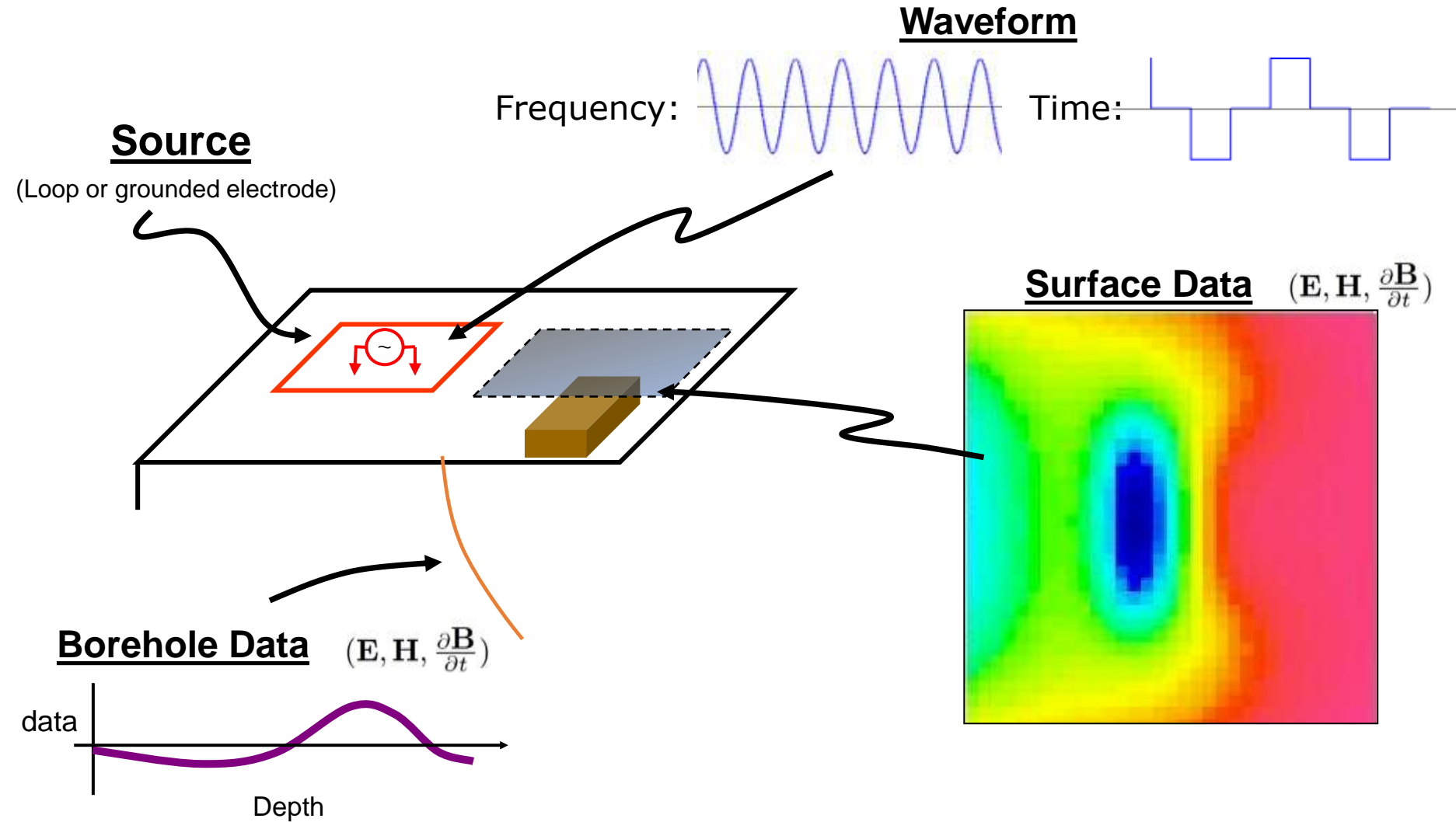
Question: Can you find those features on the data map and infer the geometry and orientations of the targets?

Summary

- EM induction: Quasi-static
- Loop-loop system in FD: Three loop model
 - Ampere's Law and Faraday's Law
 - Coupling
 - Induction number and response function
- EM-31 as an example
 - Positive or negative?
 - Compare in-phase with quadrature



EM Surveys



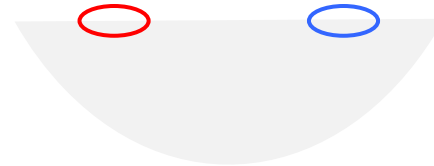
EM Surveys

- Type of source: magnetic dipole, electric dipole, plane wave (natural source)
- Frequency or time domain
- Source waveform: harmonics, square wave, pulse wave
- Operating frequencies or time channels
- Data: complex or real

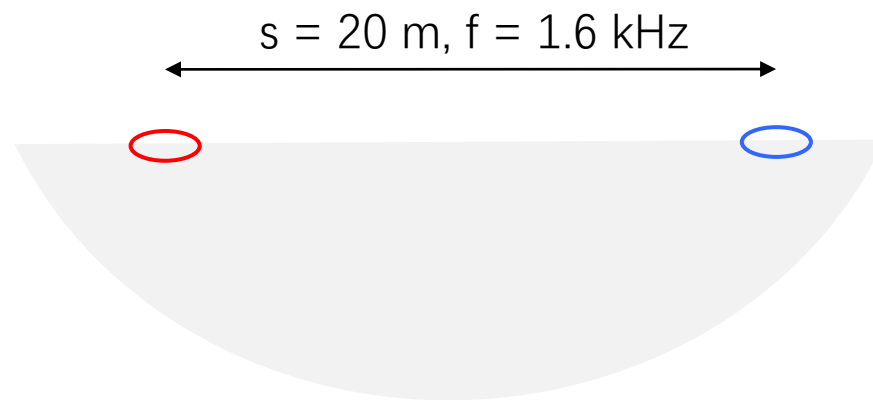
EM-41



$s = 10 \text{ m}, f = 6.4 \text{ kHz}$

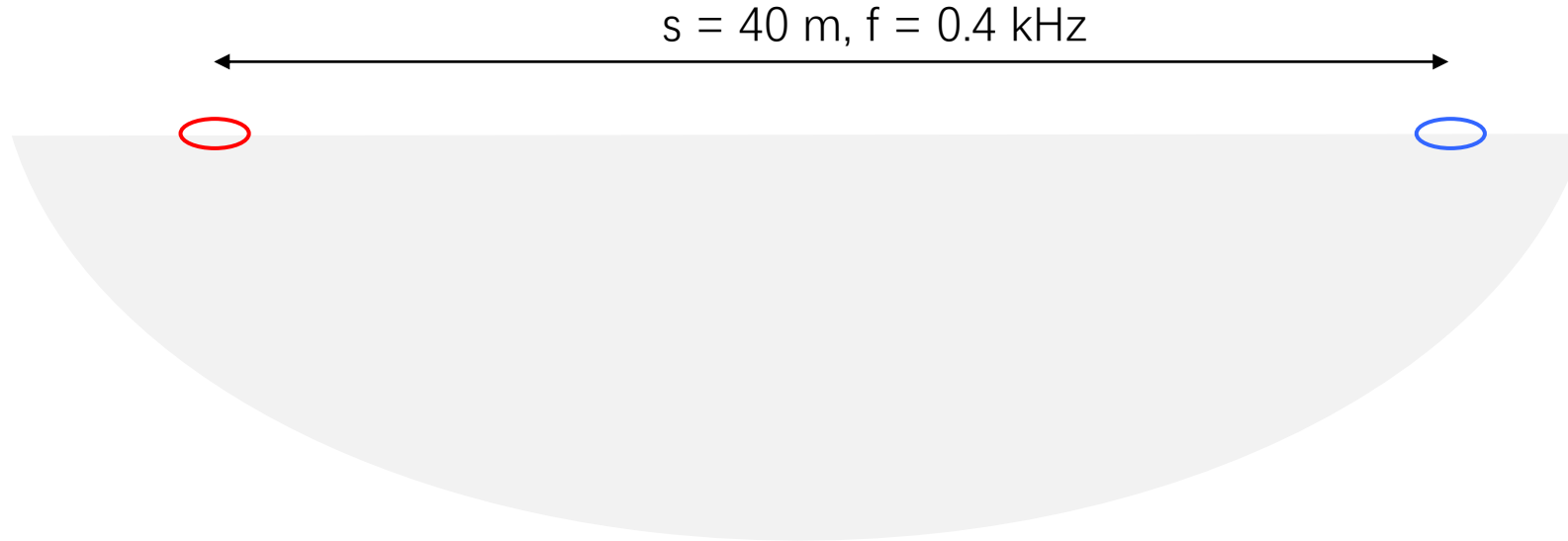


EM-41

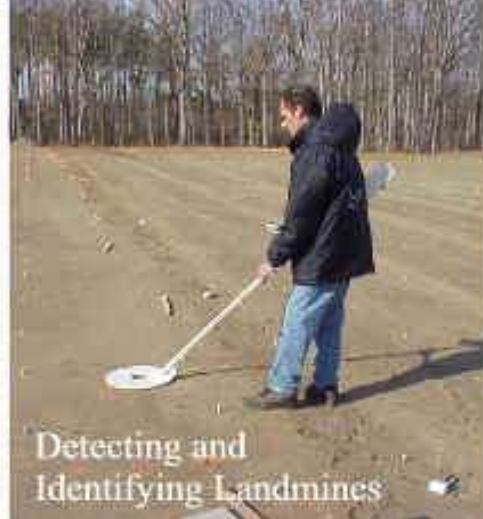


EM-41

- Variable depth of exploration down to 60 m
- HCP or VCP coil configuration
- Groundwater exploration in fractured and faulted bedrock

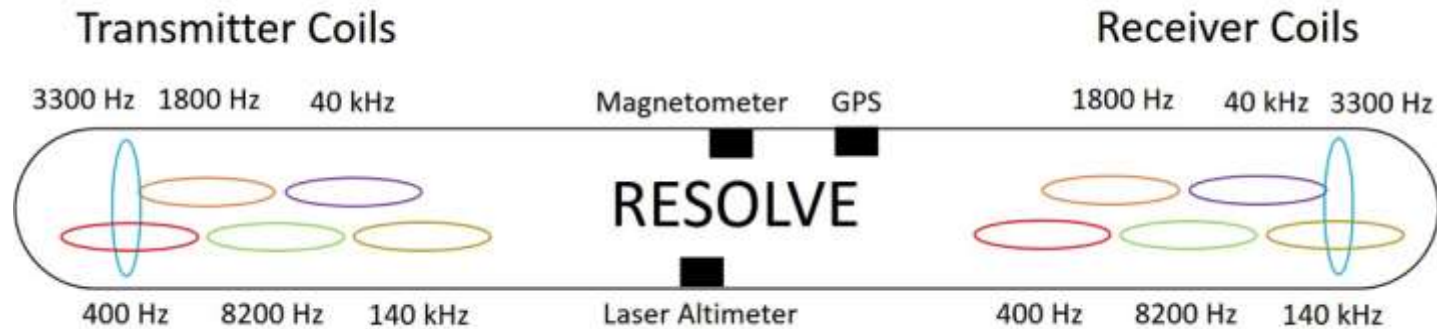
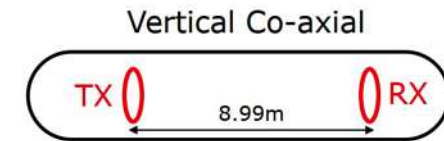
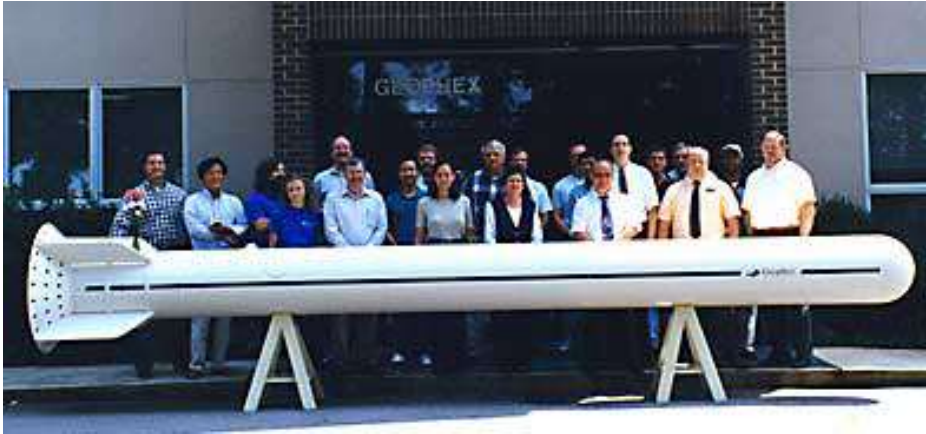


GEM3



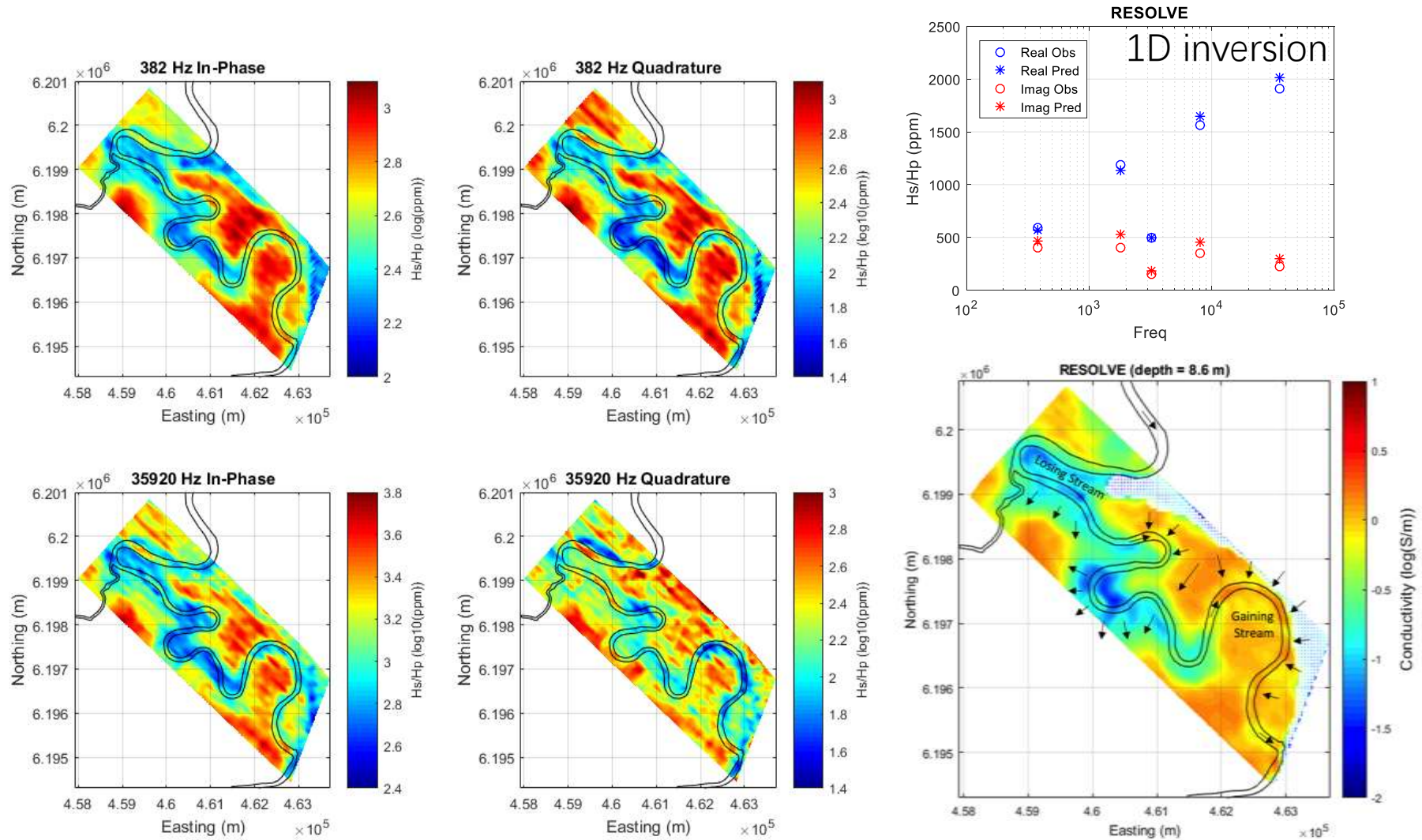
- Concentric Tx-Rx
- Frequency 60 Hz to 24 kHz
- Identify an object based on its spectral fingerprints

Airborne EM

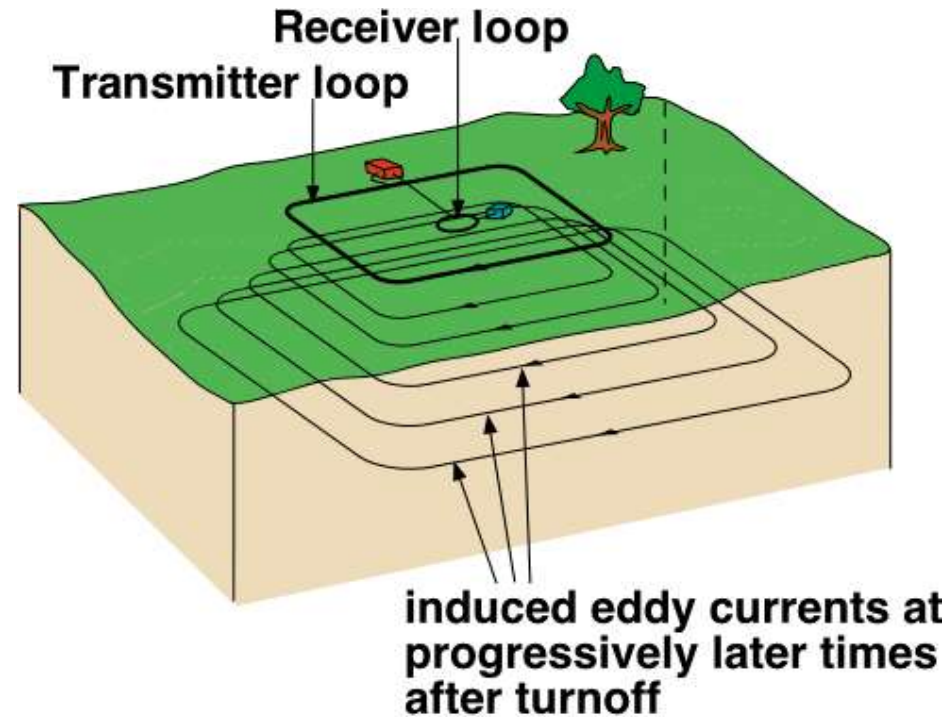
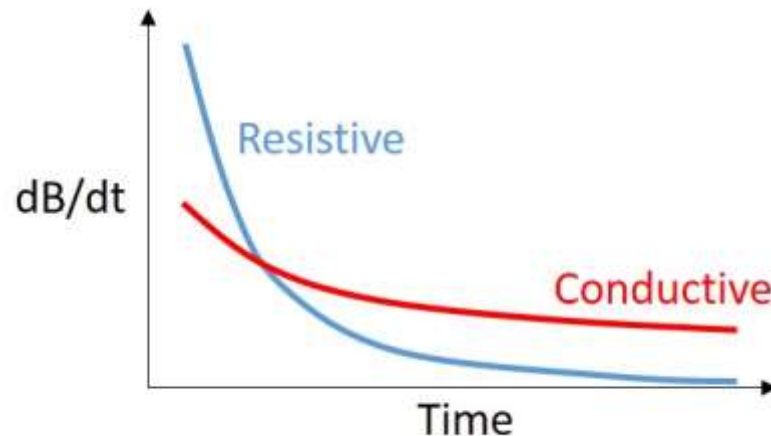
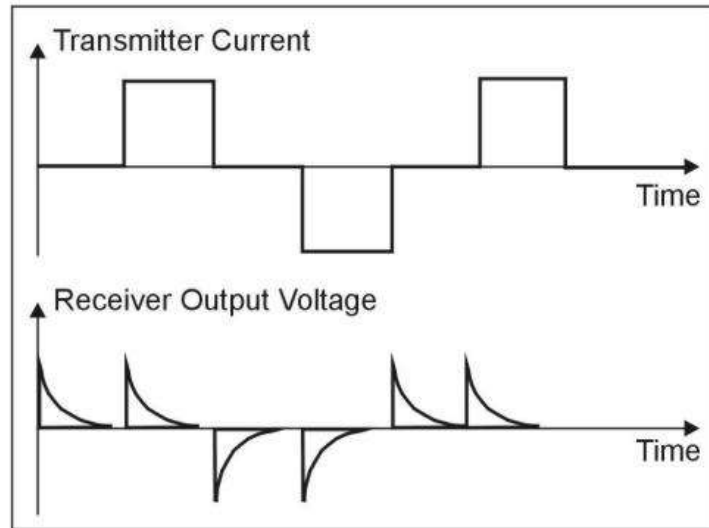


Skin depth: High frequency for shallow; low frequency for deep

Airborne EM – Groundwater Flow

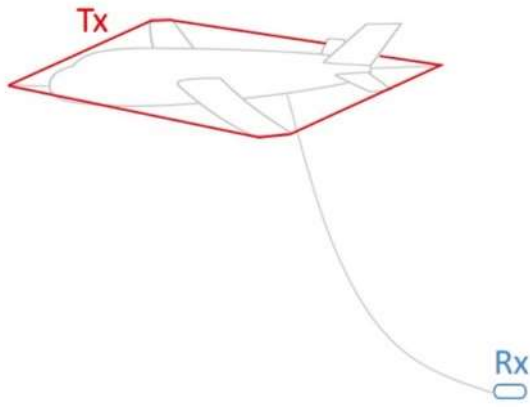


Time-domain (Transient) EM

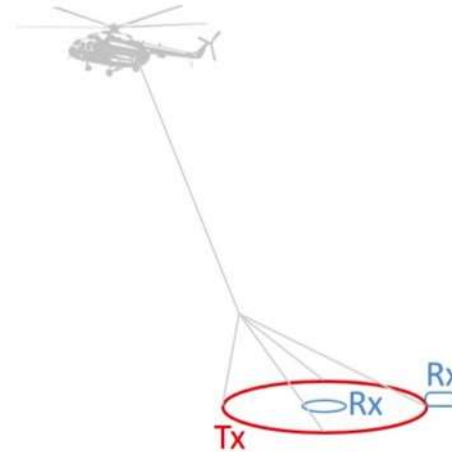


- Wider frequency bandwidth
- Deeper penetration
- Time channel: early for shallow, late for deep

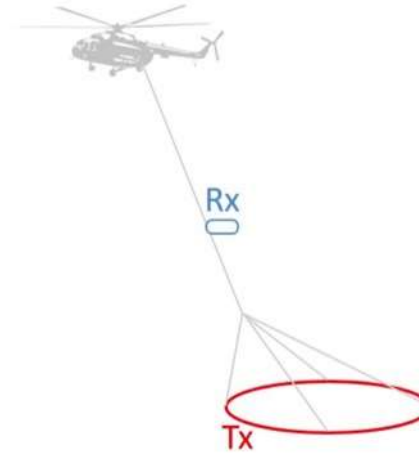
Airborne Time-domain EM (TEM)



(a) Fixed-wing

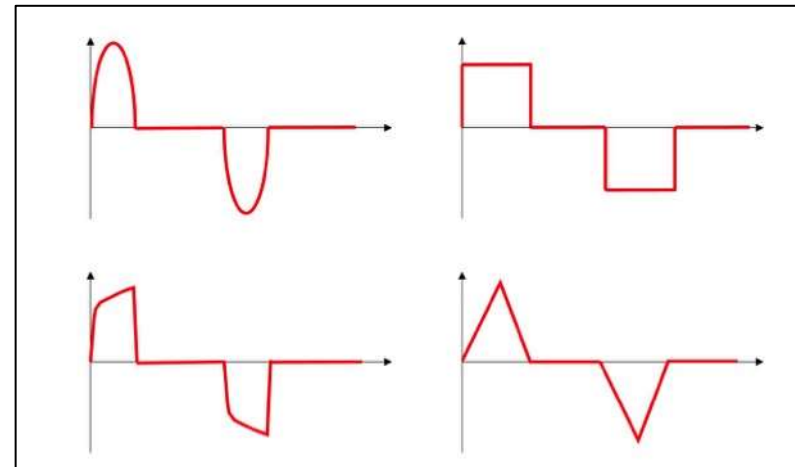


(b) Helicopter – rigid

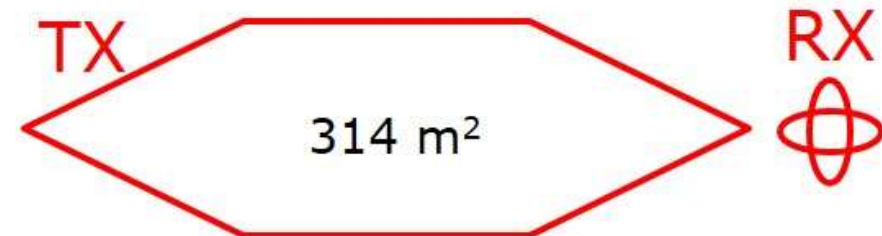
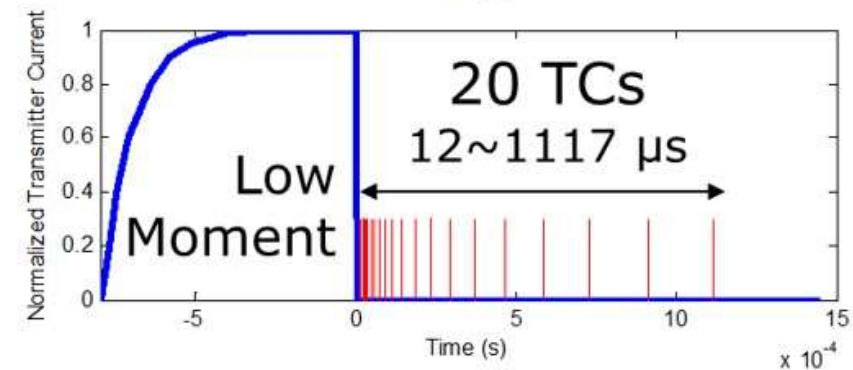
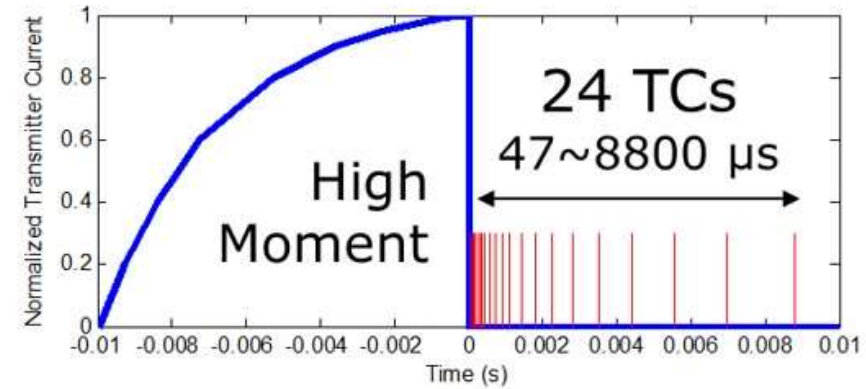


(c) Helicopter – nonrigid

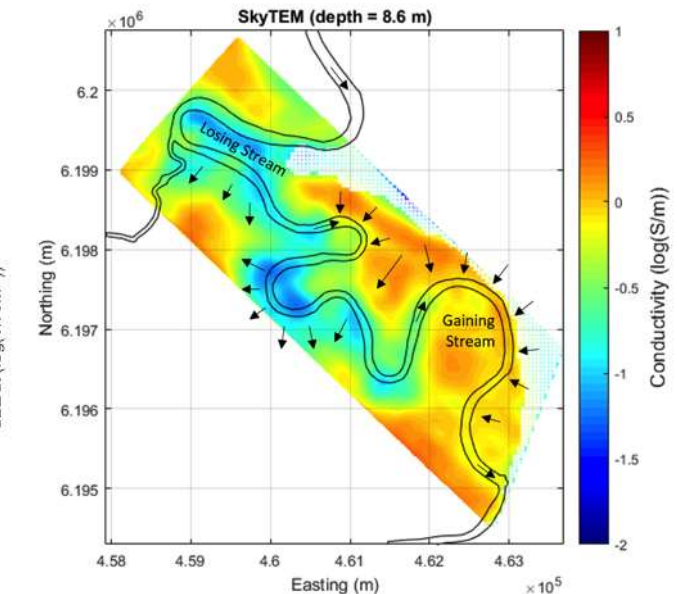
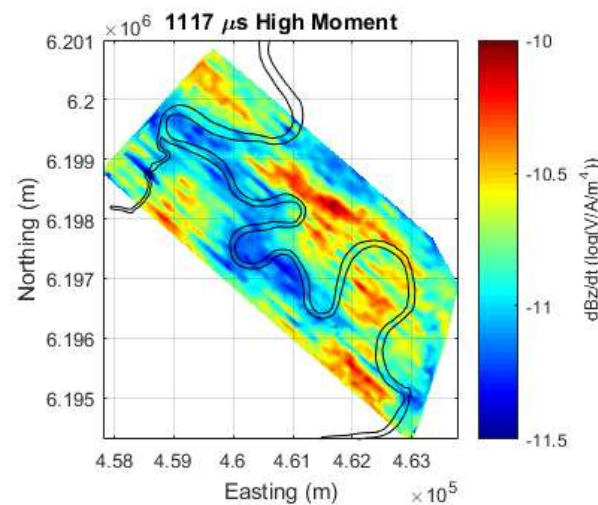
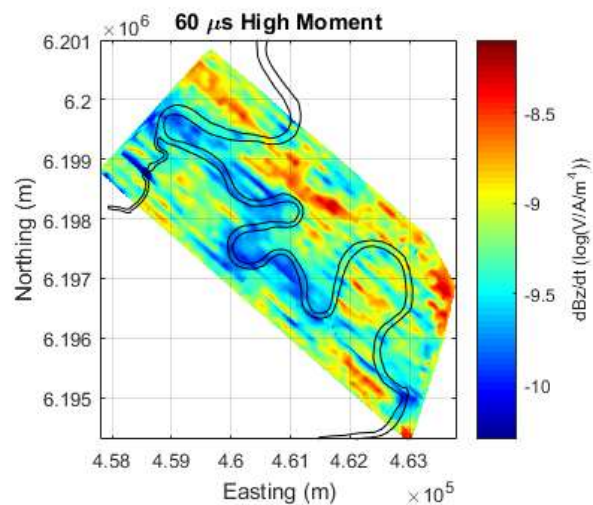
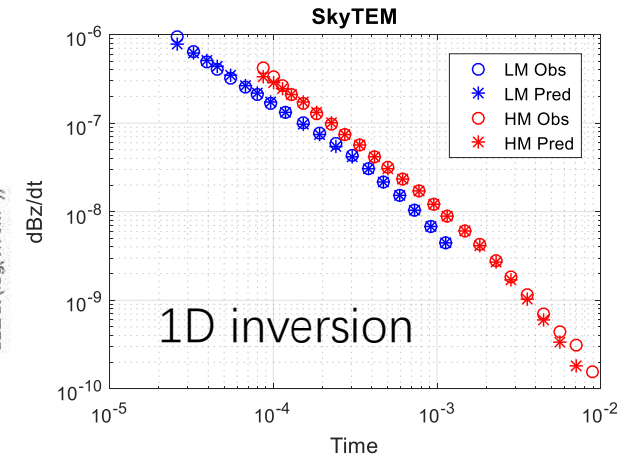
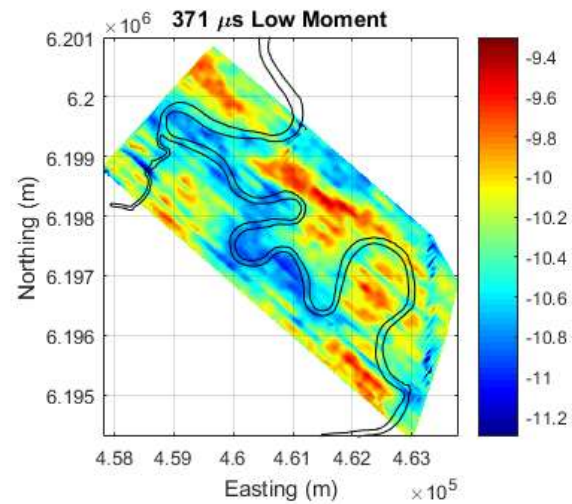
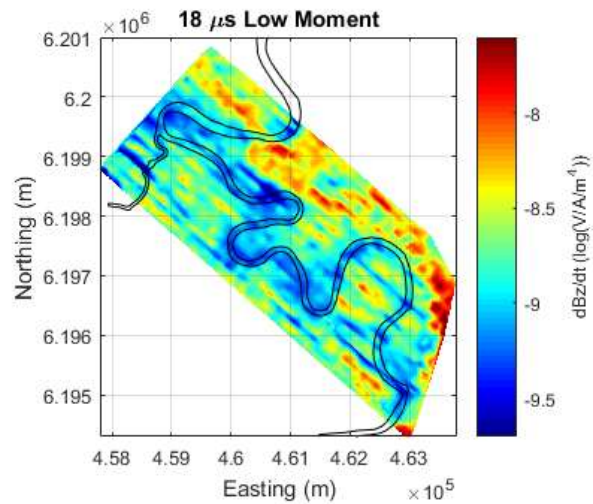
- Magnetic dipole Tx and Rx
- High efficiency
- Sensitive to conductors (water, minerals)
- Adjustable source moment
- Waveforms

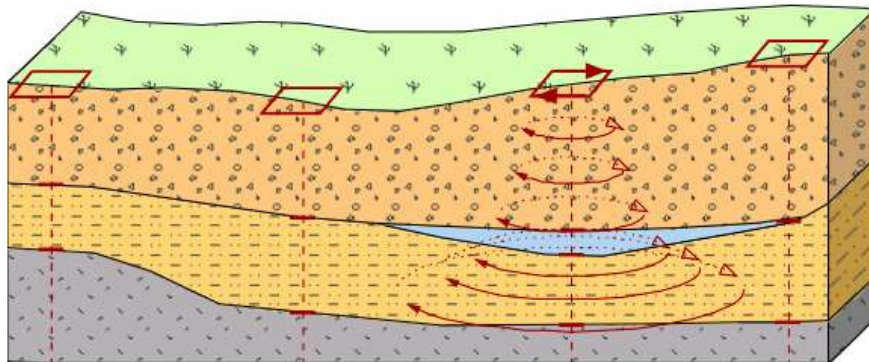


ATEM - SkyTEM



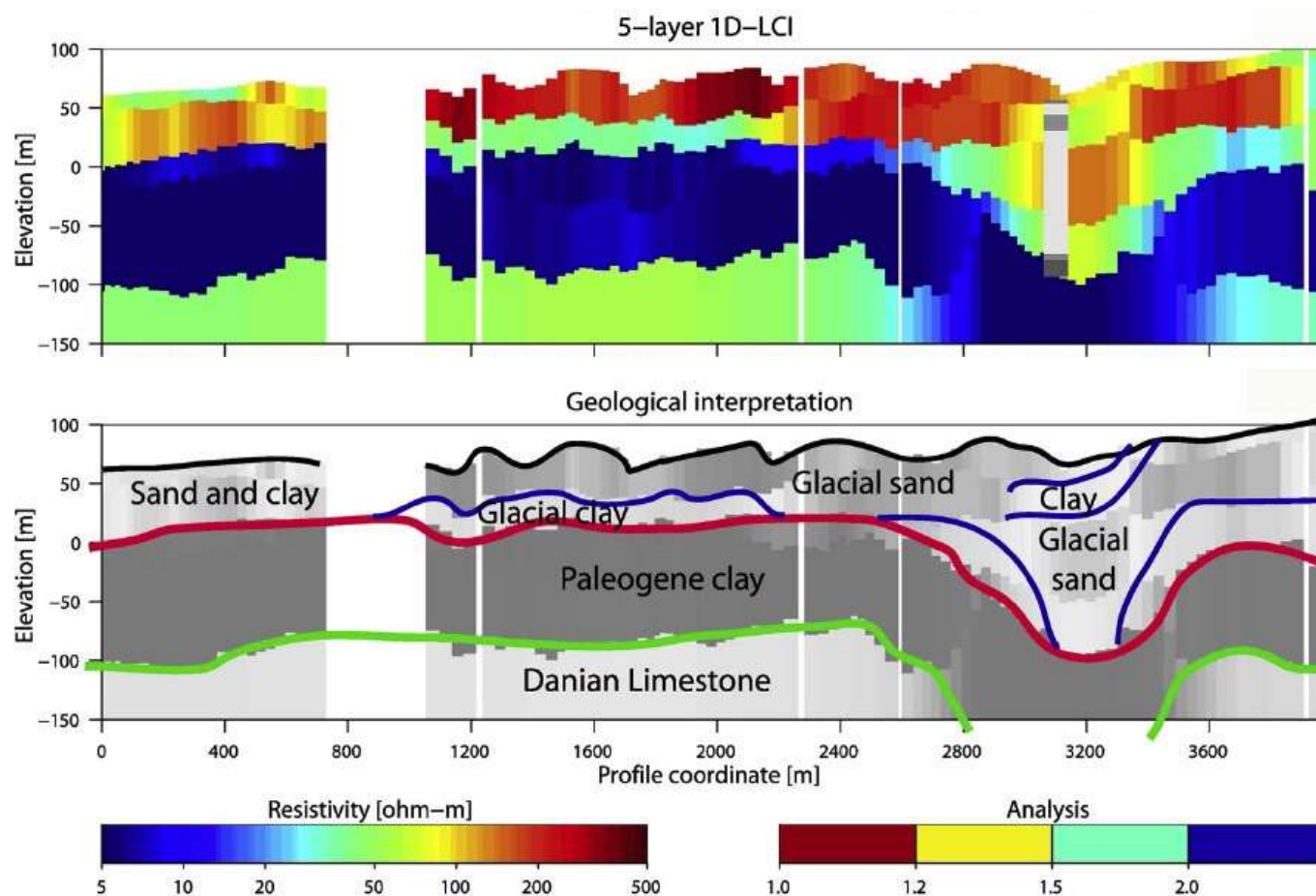
ATEM - Bookpurnong





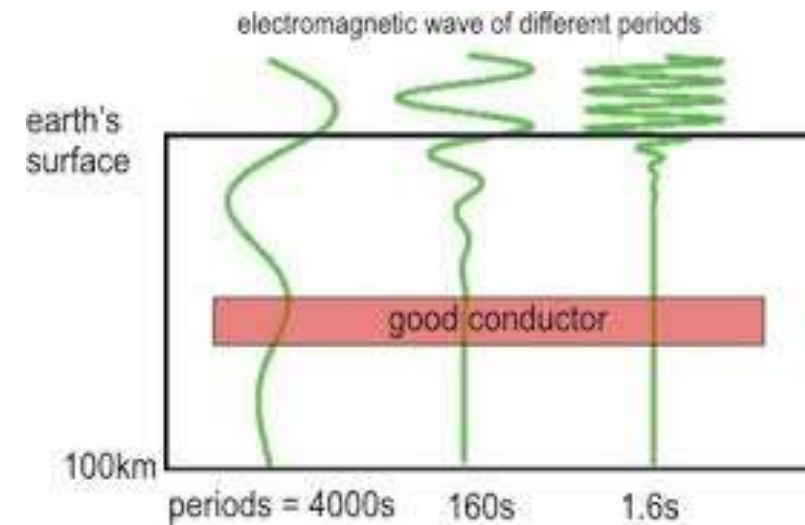
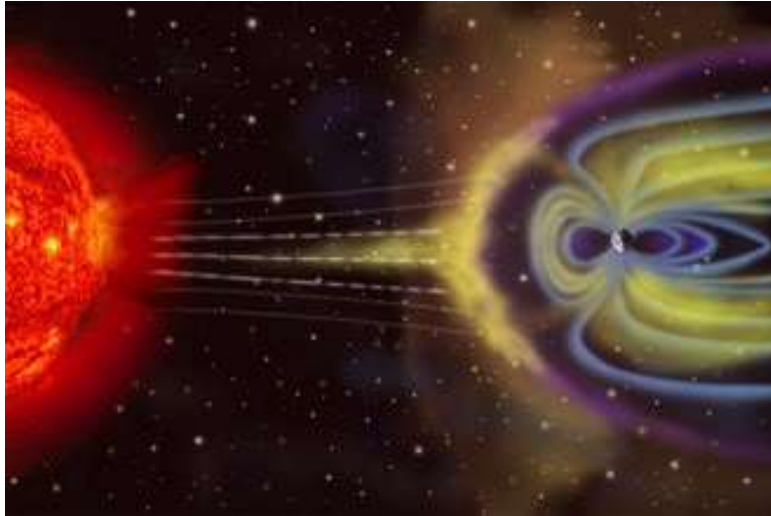
Surface TEM

- Concentric Tx-Rx
- Time decay curve at each station
- 1D layered inversion at each station
- Stitch 1D models to form a 2D section



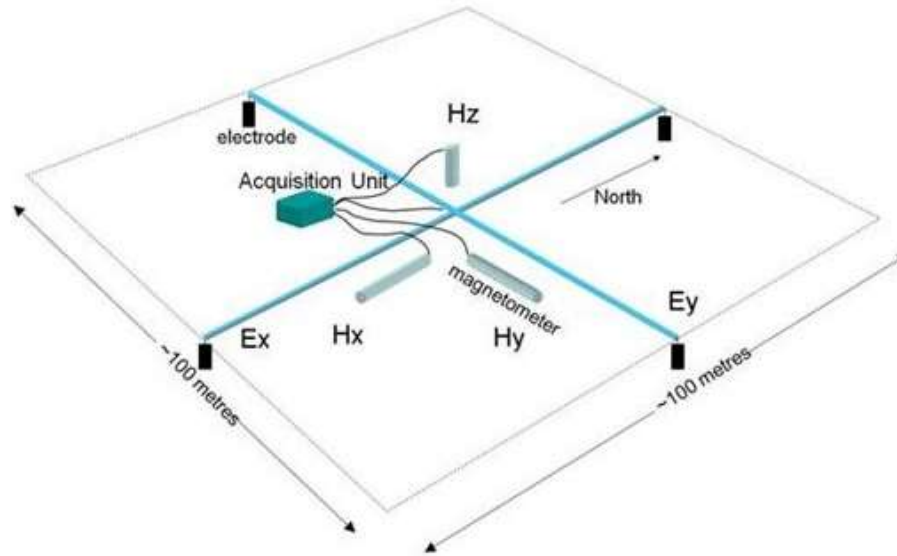
(NGA, EPA)

Natural Source EM



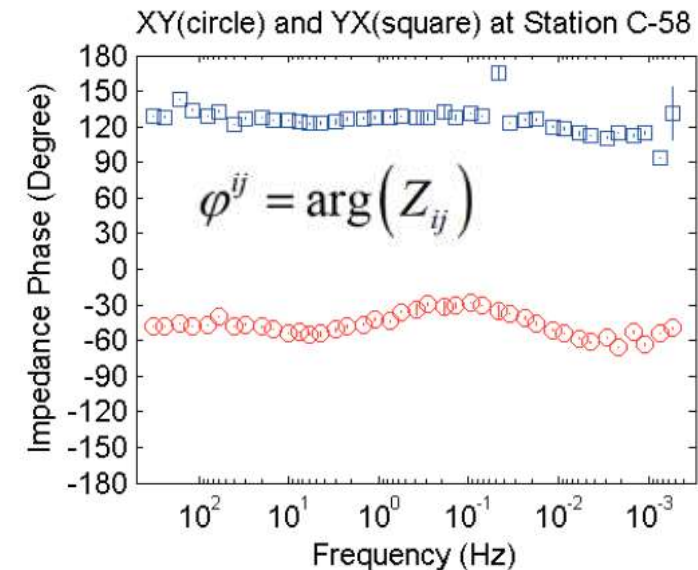
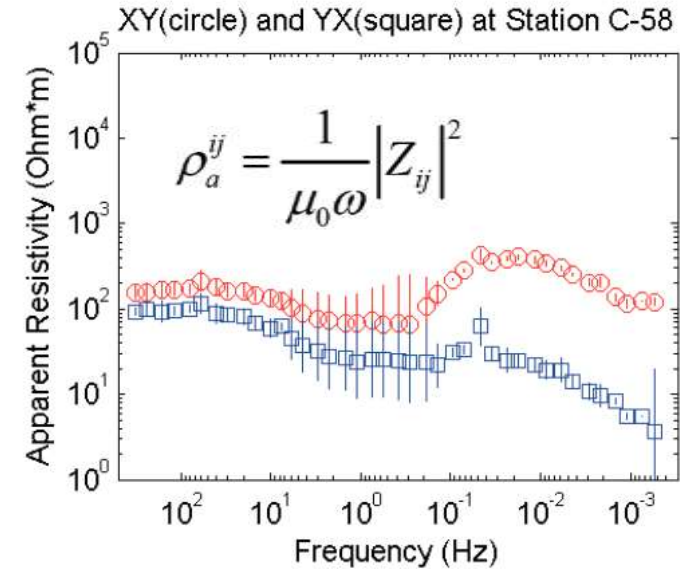
- Plane wave: horizontal E, H fields
- Frequency: 1 kHz – 10^{-4} Hz
- Depth of penetration: 10^1 – 10^5 m

Magnetotellurics (MT)

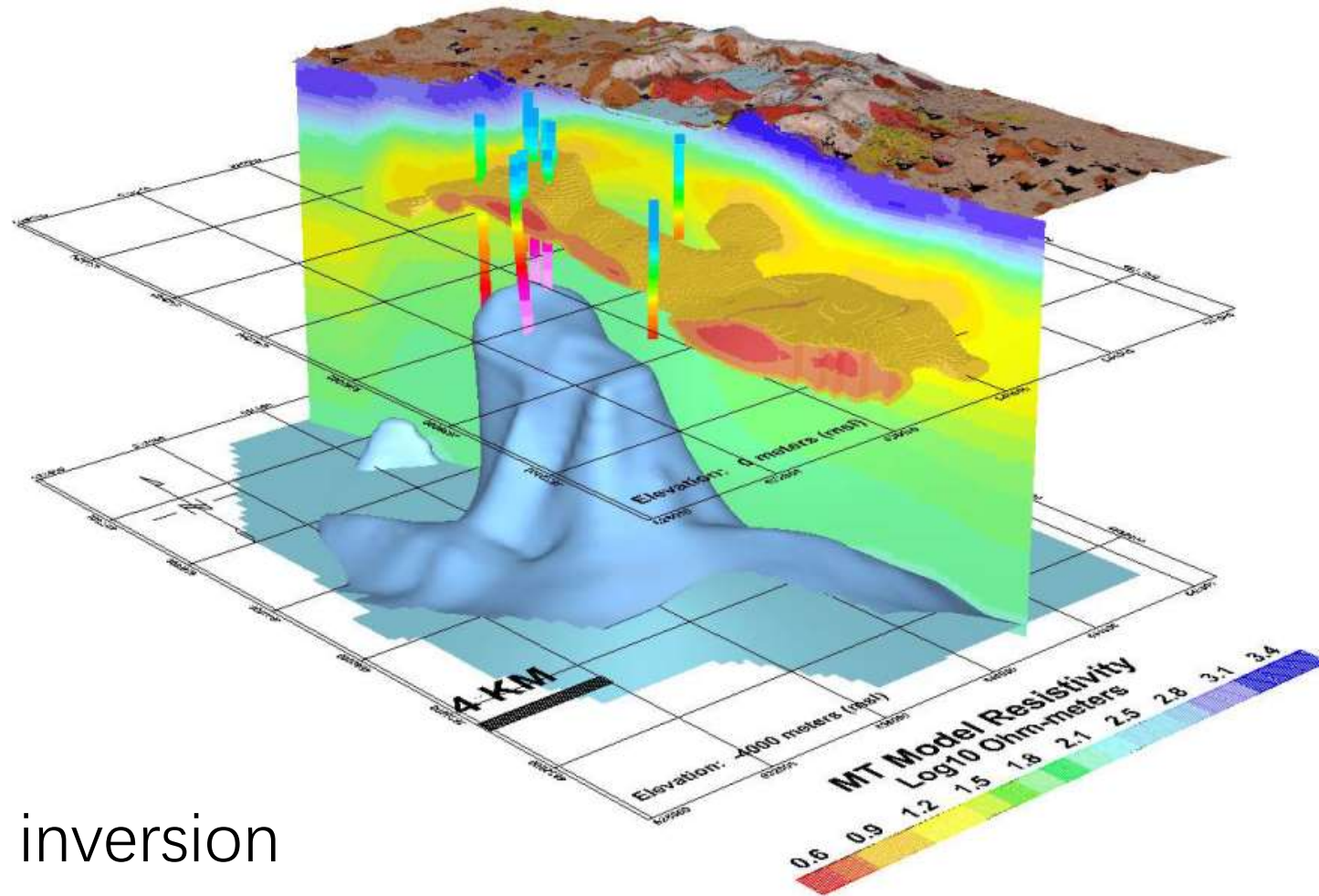


$$\begin{bmatrix} E_x \\ E_y \end{bmatrix} = \begin{bmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{bmatrix} \begin{bmatrix} H_x \\ H_y \end{bmatrix}$$

Impedance tensor element Z_{ij} is complex and a function of sounding frequency and the earth's conductivity at different depths.



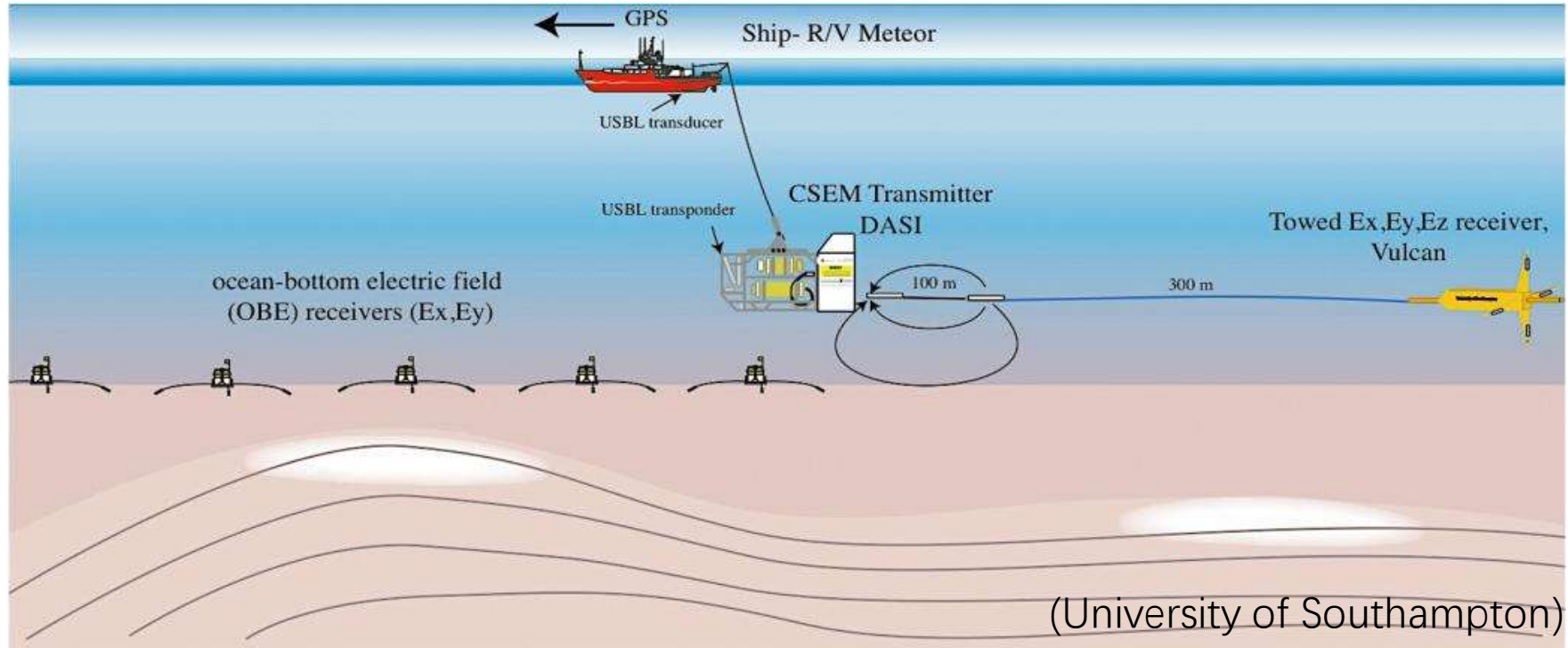
MT - Geothermal



3D inversion

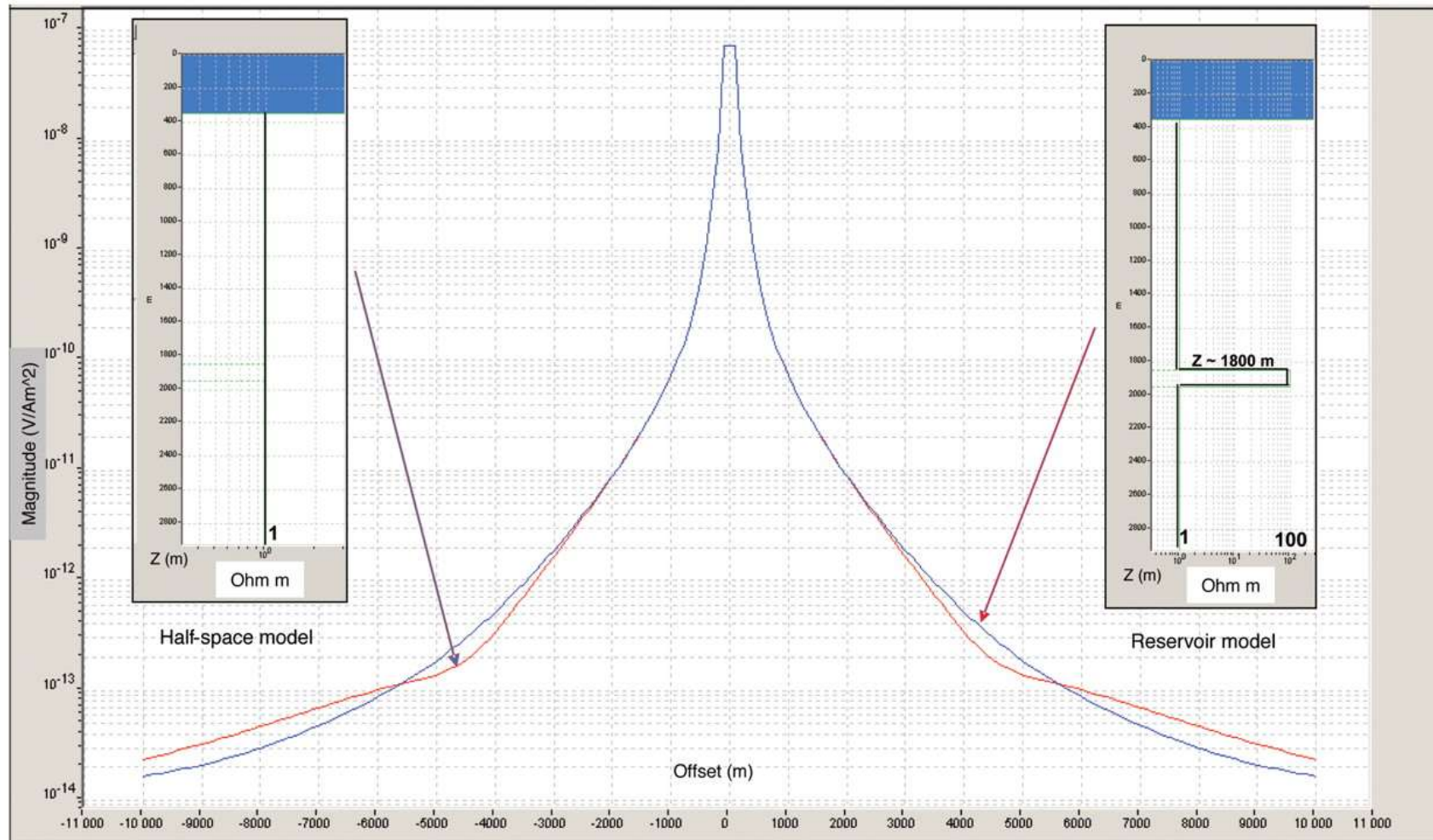
(Zonge)

Marine CSEM



- Electric dipole source
- Towed or ocean-bottom E-field receivers (electric dipoles too)
- Widely used in hydrocarbon exploration (resistors in a conductive background)

Marine CSEM



(Dell'Aversana)

Summary

- More EM surveys
 - Multi-frequency systems: EM-34, GEM3
 - Airborne EM: RESOLVE
 - Time domain EM: SkyTEM, concentric Tx-Rx
 - Natural source EM (MT)
 - Marine CSEM
- Applications
 - Groundwater/geothermal
 - Geologic mapping
 - Geotechnical, UXO
 - Petroleum