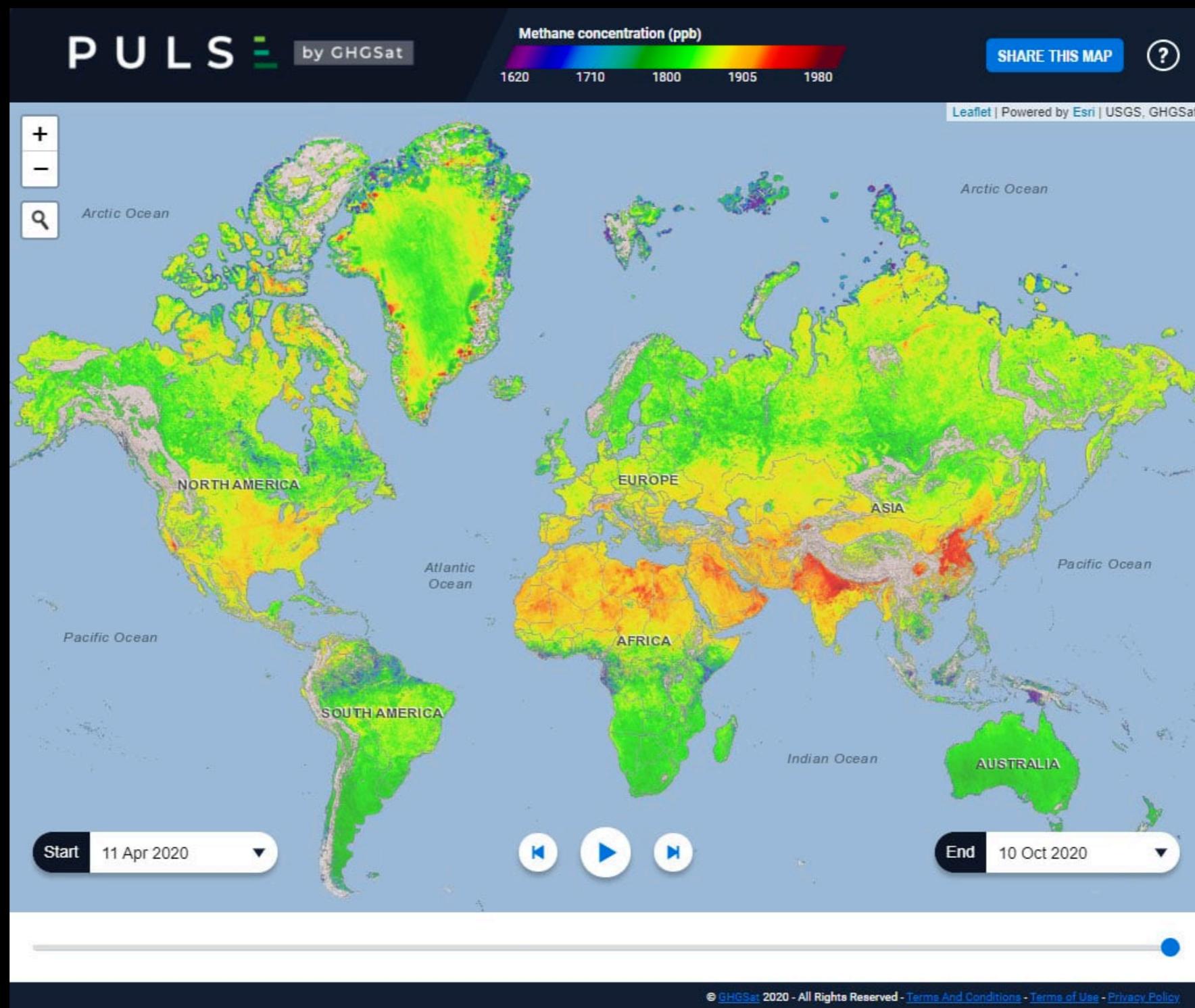


Data of the Week



<https://www.ghgsat.com/ghgsat-launches-pulse/>

Sentinel-2

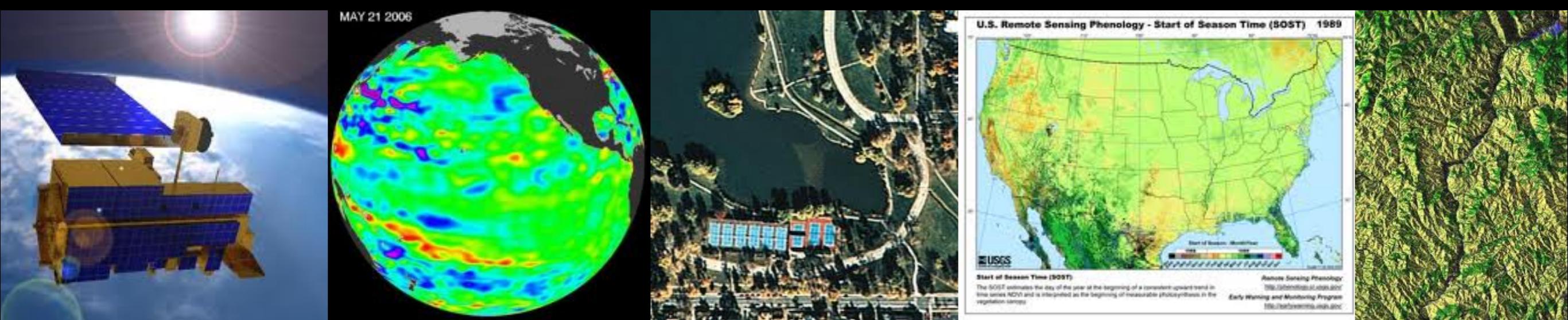
5 days revisit; Twin satellites — 2-3 days;
FOV (swath 290 km); four bands at 10 m, six bands at 20 m and
three bands at 60 m spatial resolution.

<https://www.sentinel-hub.com/explore/sentinel-playground>

<https://landsatlook.usgs.gov/sentinel2/viewer.html>

[https://sentinel.esa.int/web/sentinel/technical-guides/sentinel-2-msi/
msi-instrument](https://sentinel.esa.int/web/sentinel/technical-guides/sentinel-2-msi/msi-instrument)





Thermal remote sensing

Xi Yang

Department of Environmental Sciences
University of Virginia
xiyang@virginia.edu
390 Clark Hall

A new way of looking

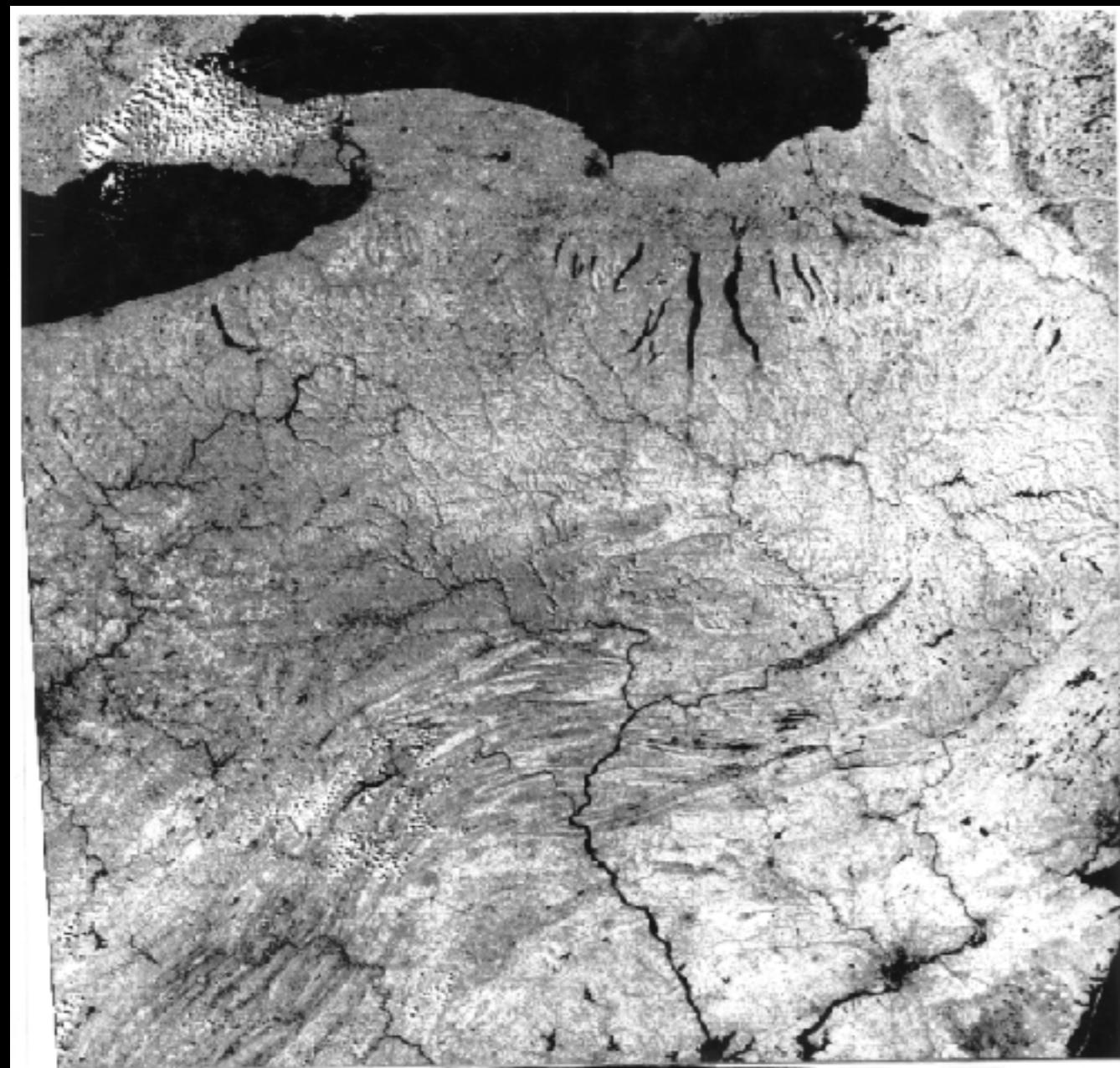


NIGHT VISION



THERMAL

Satellite RS of land surface temperature

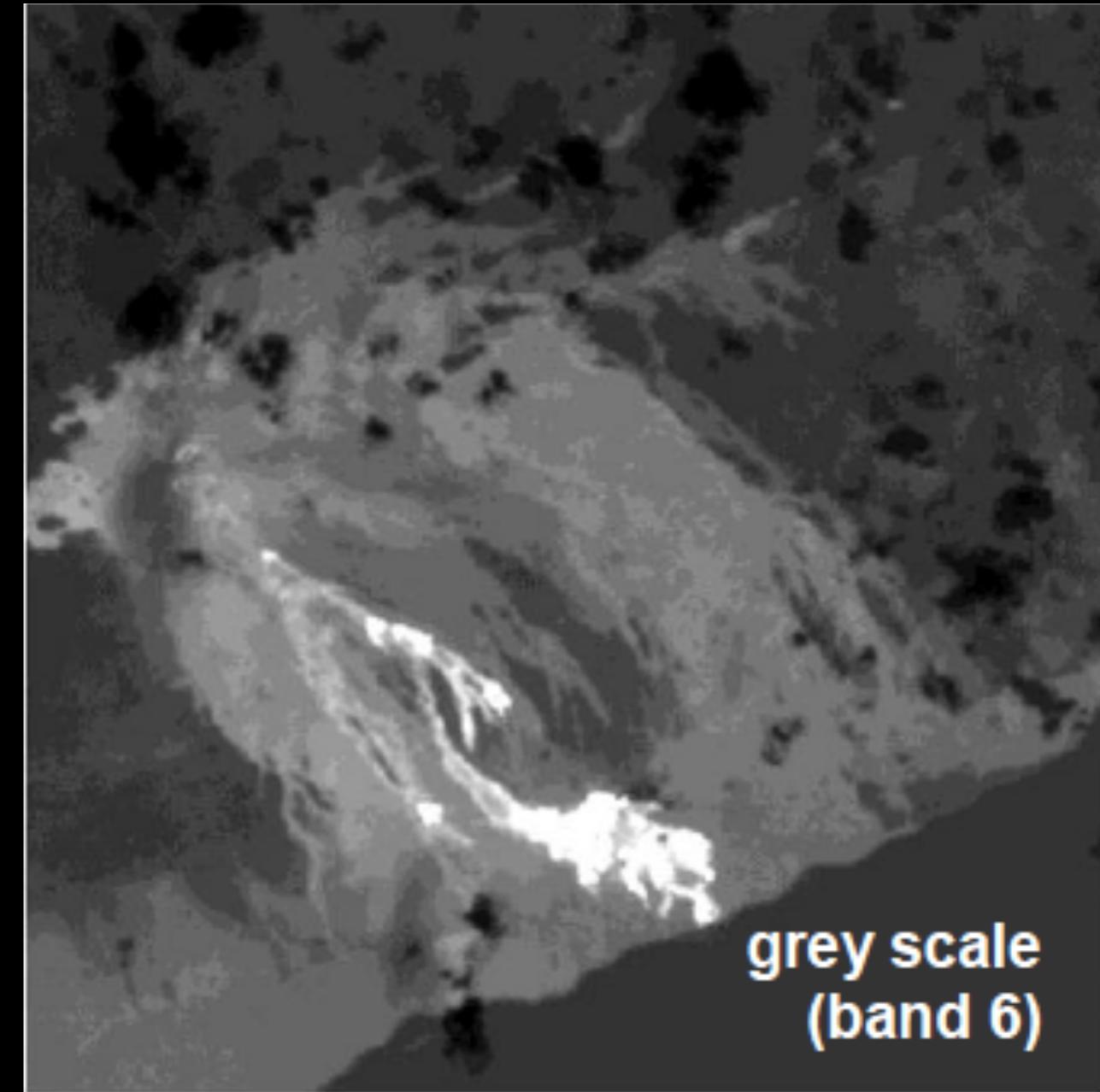


Day Time Visible (0.5 μm)



Day Time Thermal (10 μm)

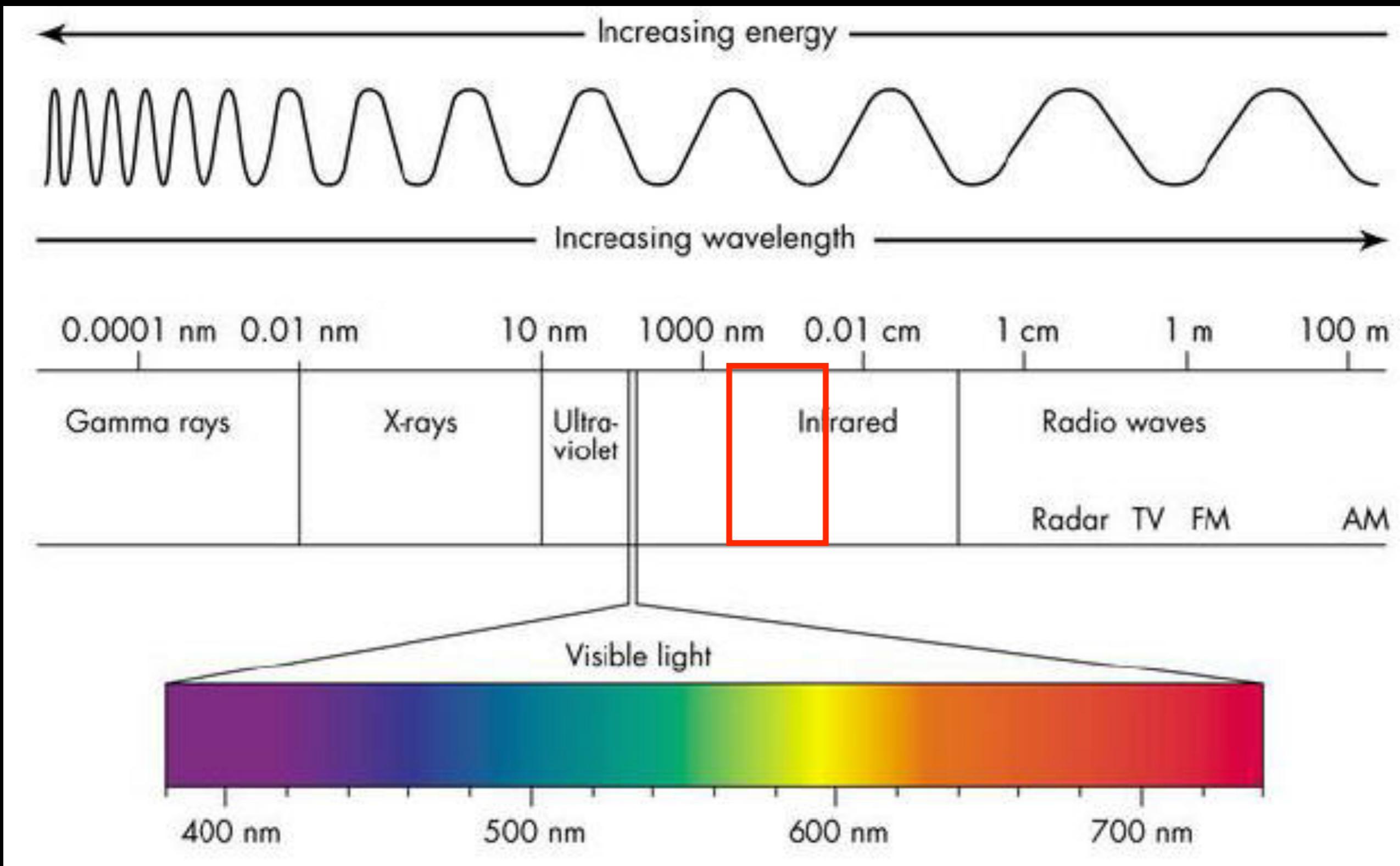
Satellite RS of land surface temperature



Kilauea Volcano, HI

Landsat 7 ETM+
February 14, 2000

Thermal infrared radiation: 3-14 um



EMR laws: Recap

- Stefan-Boltzmann Law (Blackbody)

$$M = \sigma T^4 \quad \sigma = 5.6697 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$$

- Gray body situation:

$$M = \epsilon \sigma T^4 \quad \epsilon = 0-1 \quad \sigma = 5.6697 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$$

- To be more precise, Kirchhoff's law is actually

$$\alpha = \epsilon$$

- Wien's Displacement Law: $\lambda_{\max} = \frac{2898}{T}$

Satellite sensors measure the radiance



$$M = \epsilon\sigma T^4$$

A constant

Radiance

Kinetic temperature

Emissivity

```
graph LR; A[A constant] --> M["M = εσT⁴"]; B[Radiance] --> M; C[Kinetic temperature] --> M; D[Emissivity] --> M;
```

Brightness temperature (unit: K): the apparent temperature of a surface assuming $\epsilon = 1$

Kinetic temperature: true temperature as measured by a thermometer

Same Kinetic Temperature, different emissivity

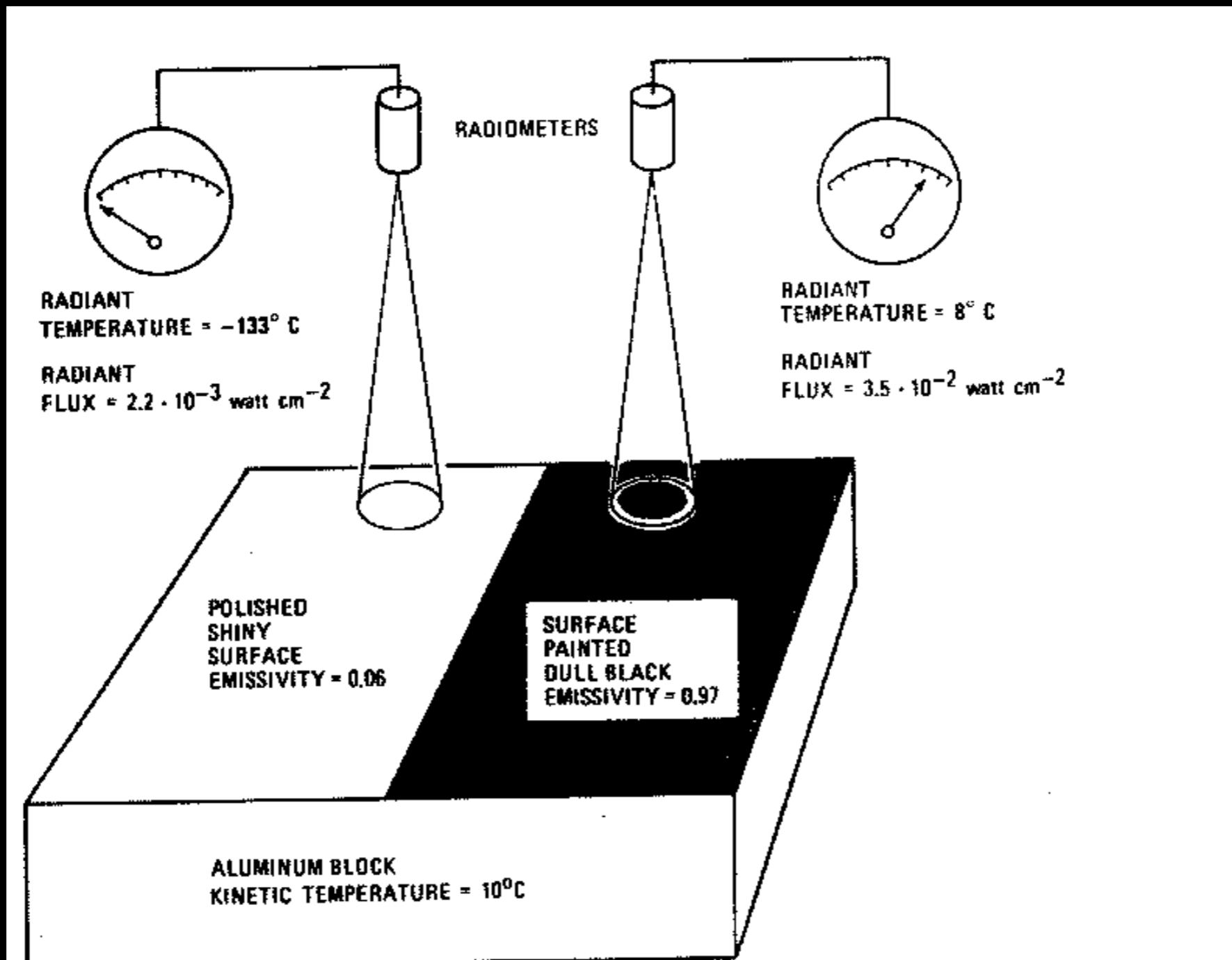


FIGURE 5.1

Effect of emissivity differences on radiant temperature. Kinetic temperature of aluminum block is uniformly 10°C . Different emissivities cause different radiant temperatures that are measured with radiometers.

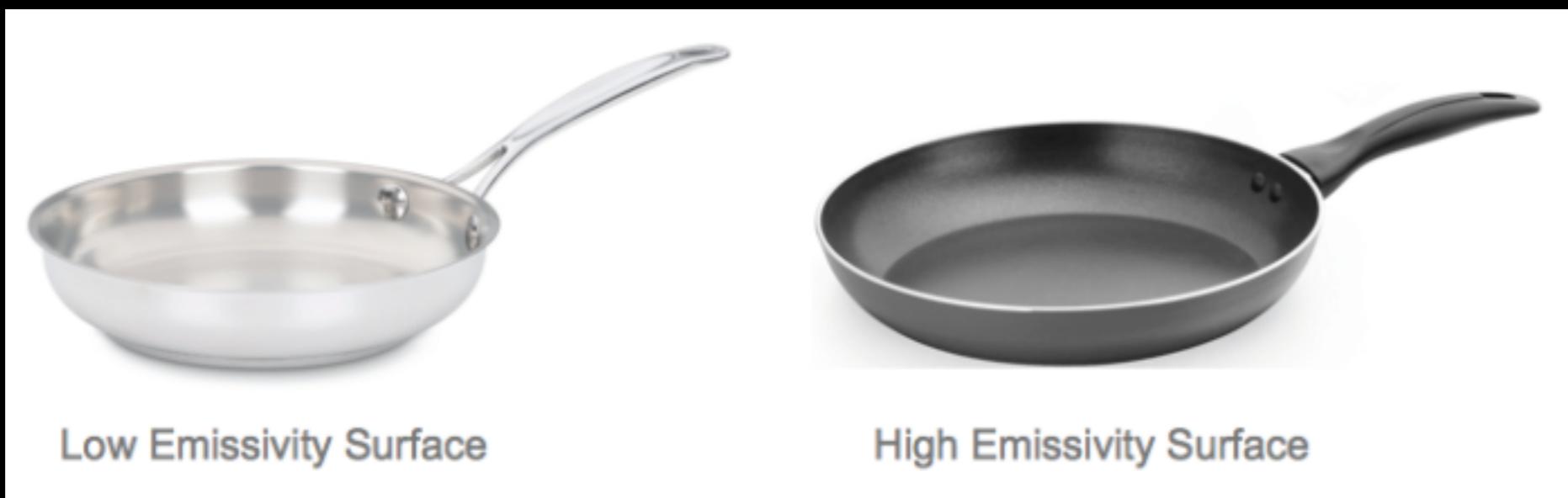
Emissivity of typical land cover

For thermal infrared radiation, there is little transmission. Thus we have:

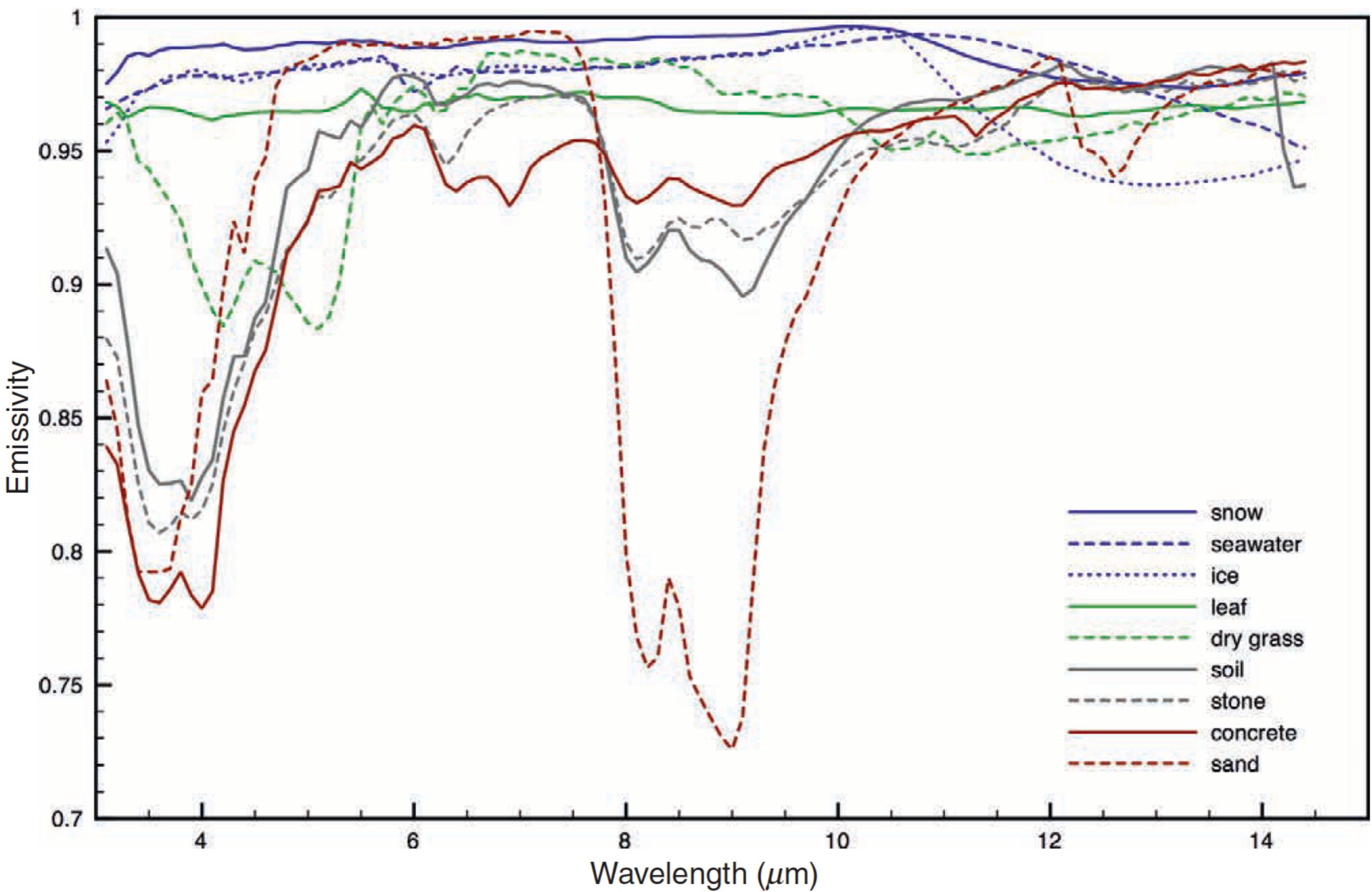
$$\epsilon + \rho = 1$$

Emissivity is usually a function of the material itself and the wavelength:

$$\epsilon = f(\text{material}, \lambda, \text{viewing_angle})$$



Emissivity of typical land cover

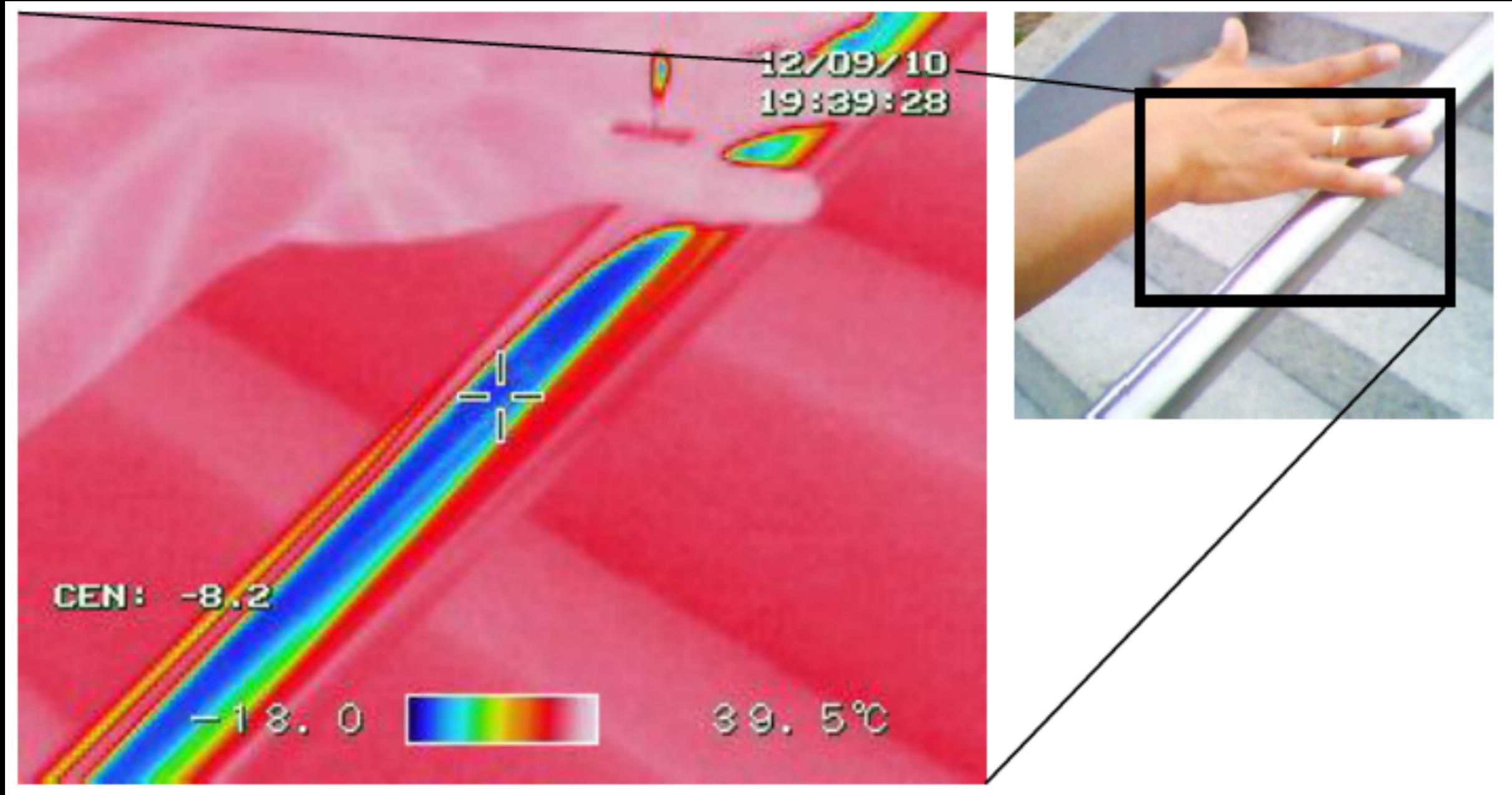


Emissivity of typical land cover

Table 1.1 Emissivity of different surfaces in the 8-14 μm wavelength range as compiled from different sources (own measurements, Lillesand et al. 2008, Sabins 1996)

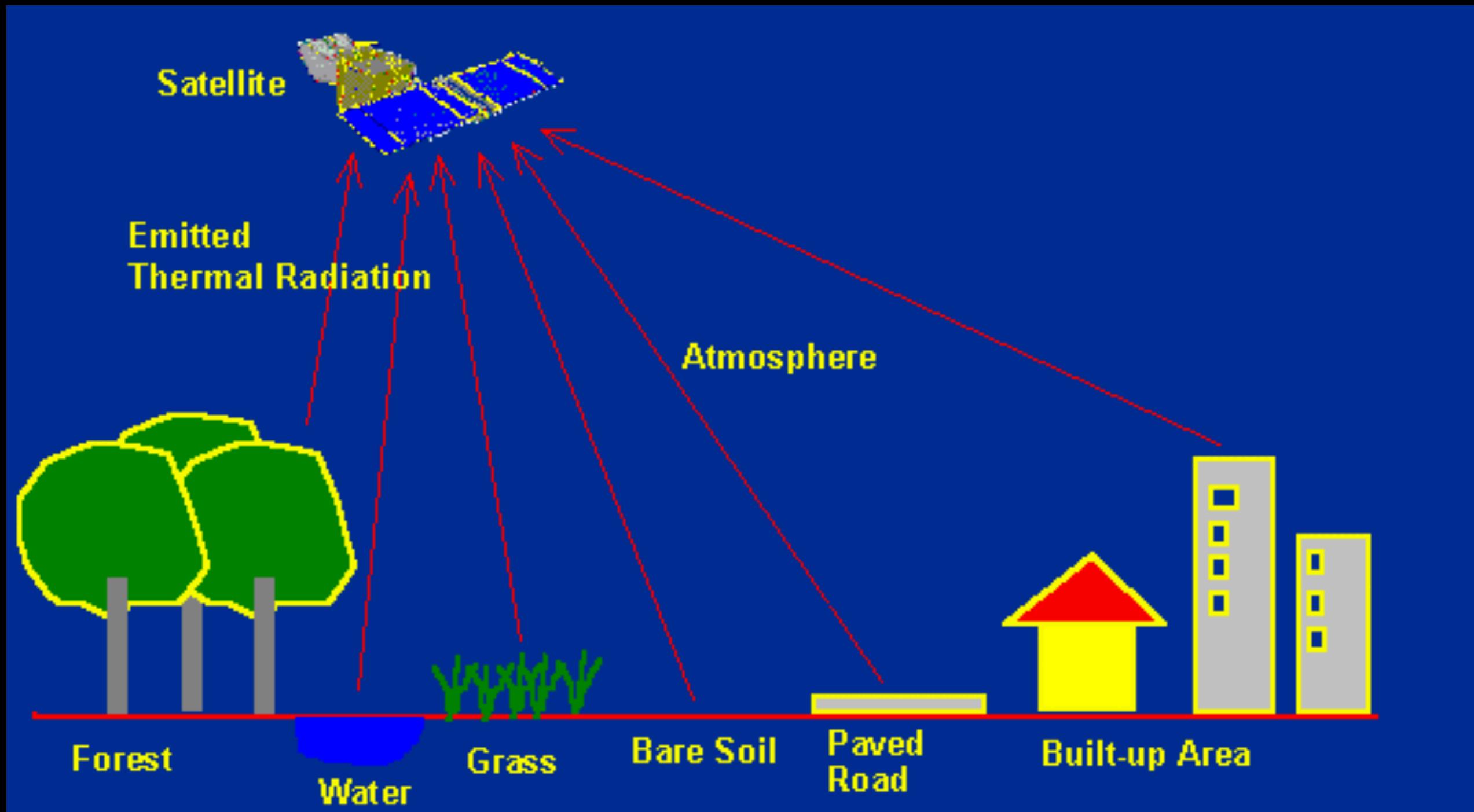
Surface	Emissivity at 8-14 μm
Carbon powder	0.98-0.99
Water	0.98
Ice	0.97-0.98
Plant leaves, healthy	0.96-0.99
Plant leaves, dry	0.88-0.94
Asphalt	0.96
Sand	0.93
Basalt	0.92
White paper	0.90
Wood	0.87
Granite	0.83-0.87
Polished metals, averaged	0.02-0.21
Aluminium foil	0.036

The importance of emissivity

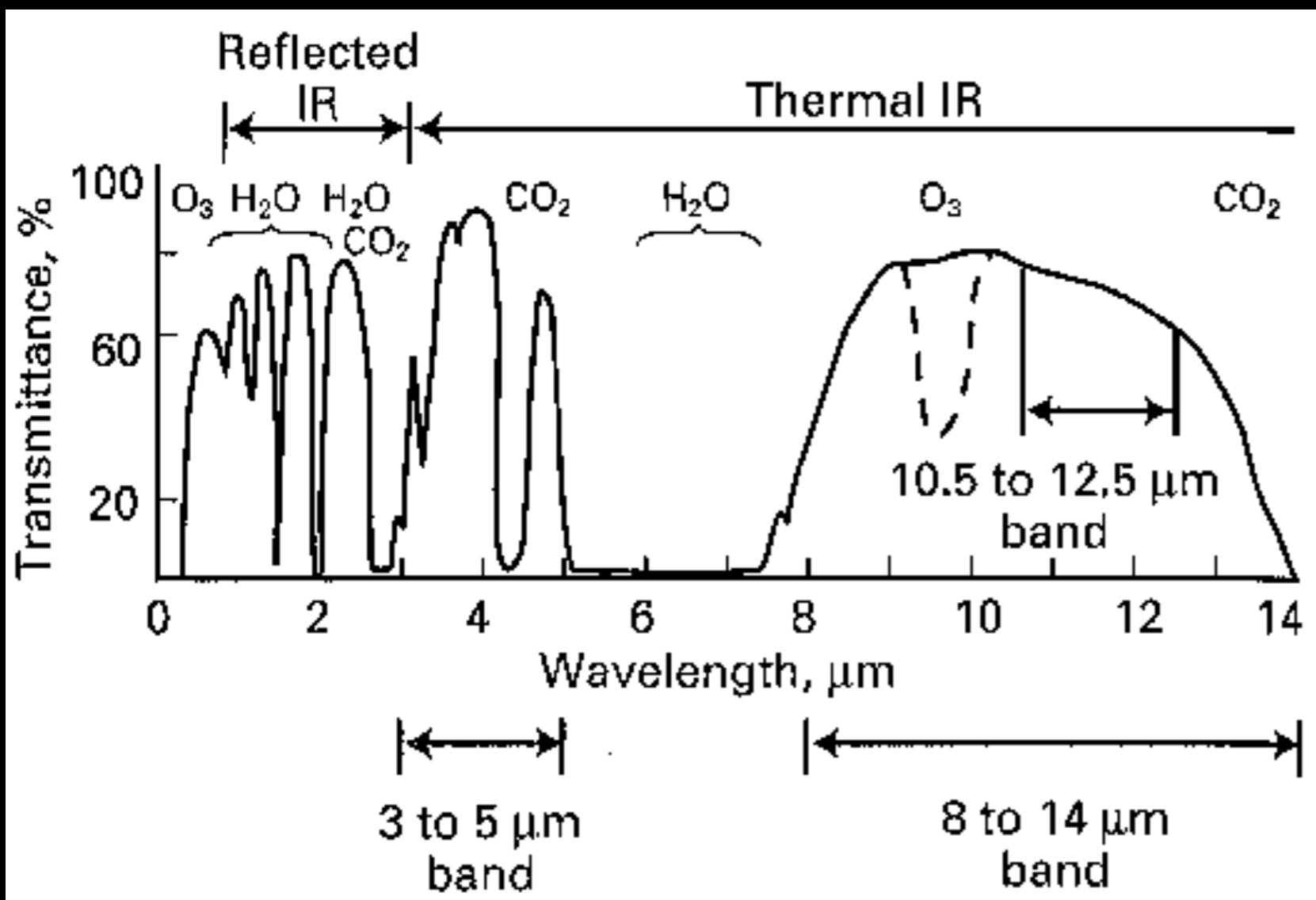


Brightness temperature or kinetic temperature?

Thermal radiation received by the satellite sensor



Atmospheric windows in the TIR region

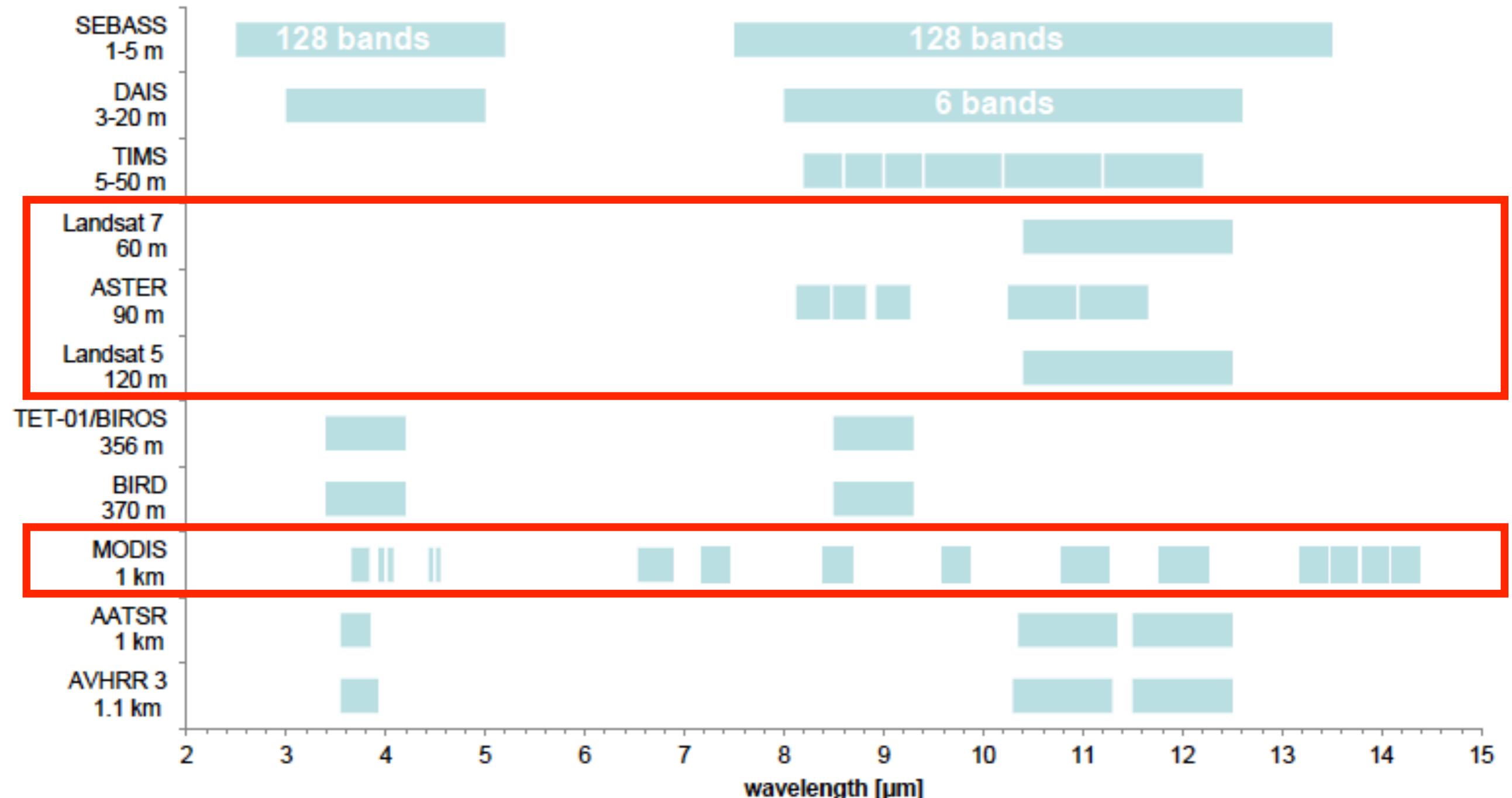


Where are the atmospheric windows in the TIR?

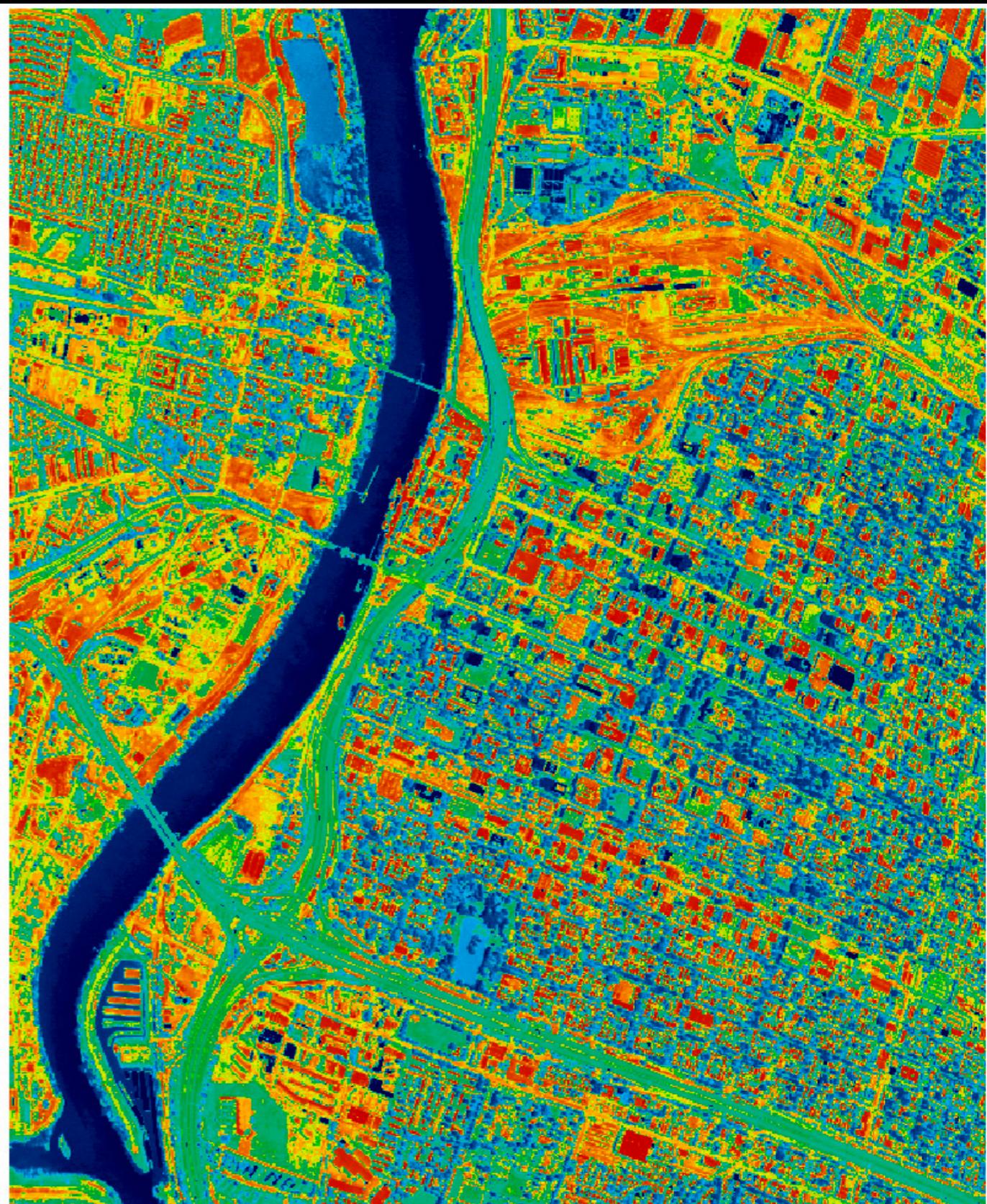
3-5 um & 8-14 um

Ozone absorption around 9-10 um

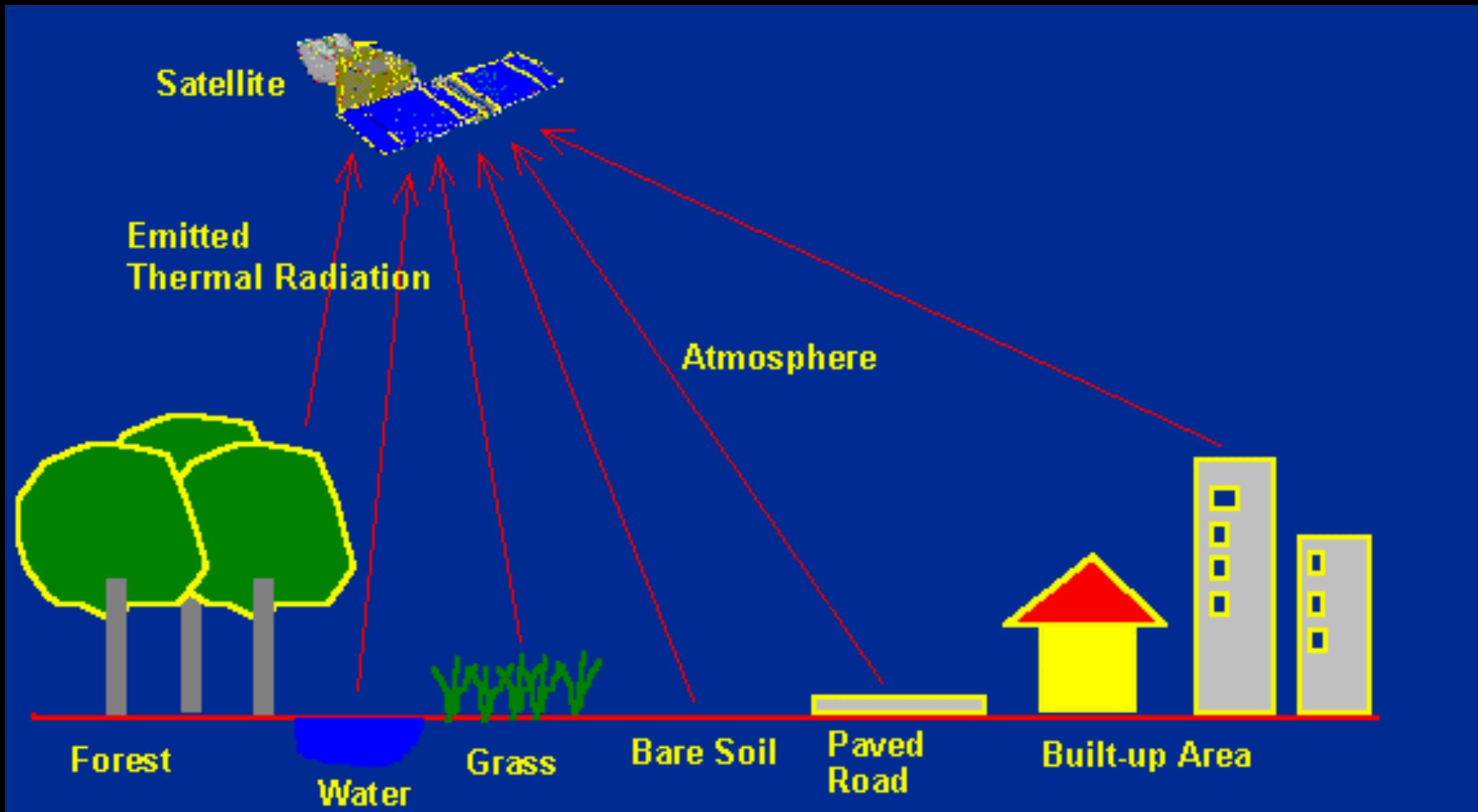
Satellites that measure TIR radiation



Sacramento, CA



What contribute to the radiation received by the satellite sensors?



Radiative Transfer in the TIR

$$L_{sat} = L_p + T_v \times \epsilon \times L_{surf}(T) + T_v \times (1 - \epsilon) \times F_d / \pi$$

Path Radiance

Radiance emitted from the surface

Radiance emitted from the atmosphere and then reflected by the surface

The diagram illustrates the components of satellite-derived surface radiance. The equation is:

$$L_{sat} = L_p + T_v \times \epsilon \times L_{surf}(T) + T_v \times (1 - \epsilon) \times F_d / \pi$$

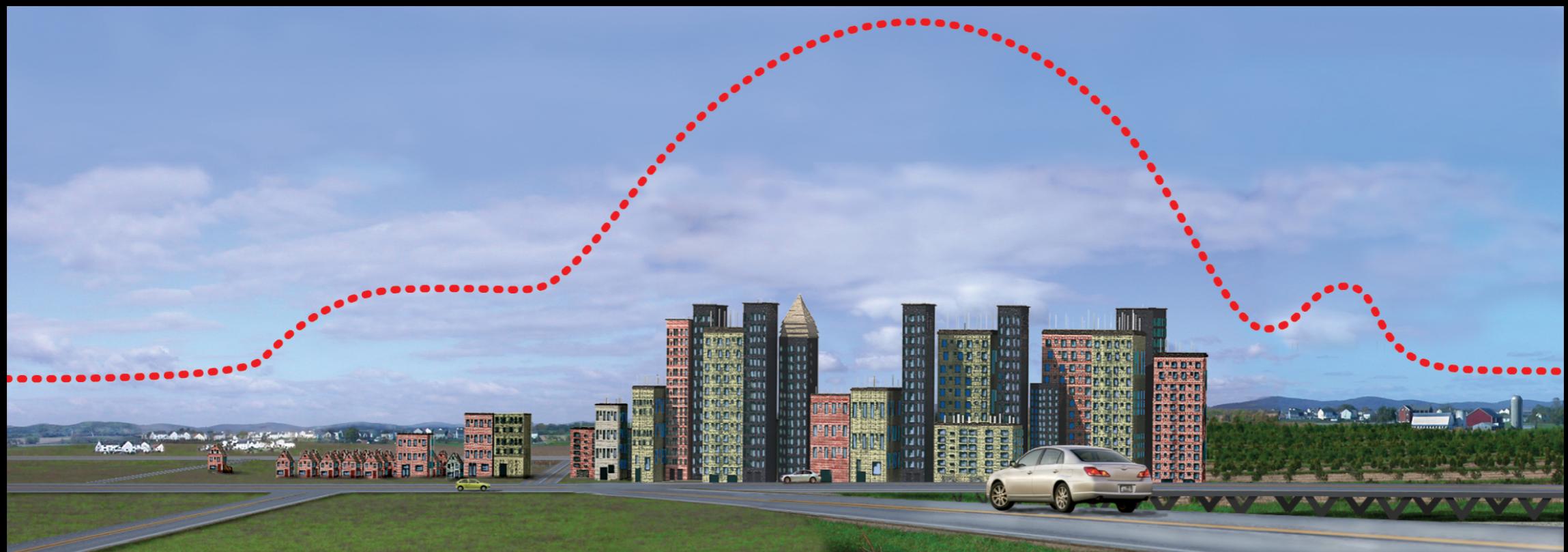
The terms are labeled as follows:

- Path Radiance: Points to L_p
- Radiance emitted from the surface: Points to $T_v \times \epsilon \times L_{surf}(T)$
- Radiance emitted from the atmosphere and then reflected by the surface: Points to $T_v \times (1 - \epsilon) \times F_d / \pi$

- L_p : thermal path radiance;
- T_v : transmittance of the atmosphere to thermal radiation;
- ϵ : emissivity of the land surface;
- $L_{surf}(T)$: surface radiation;
- F_d : downwelling thermal radiation

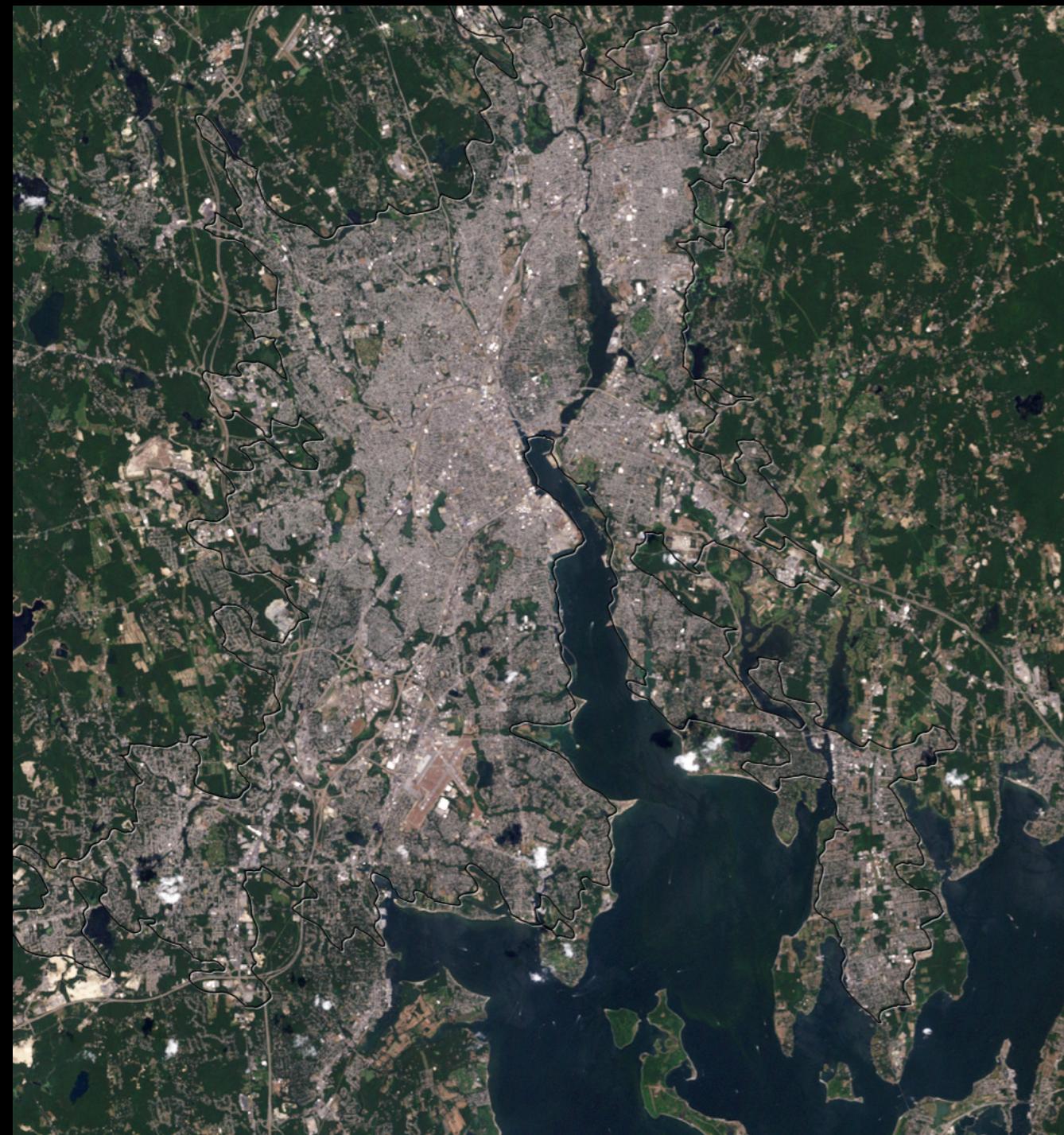
Case studies

- Urban Heat Island and TIR Change Detection
- Remote Sensing of Fire

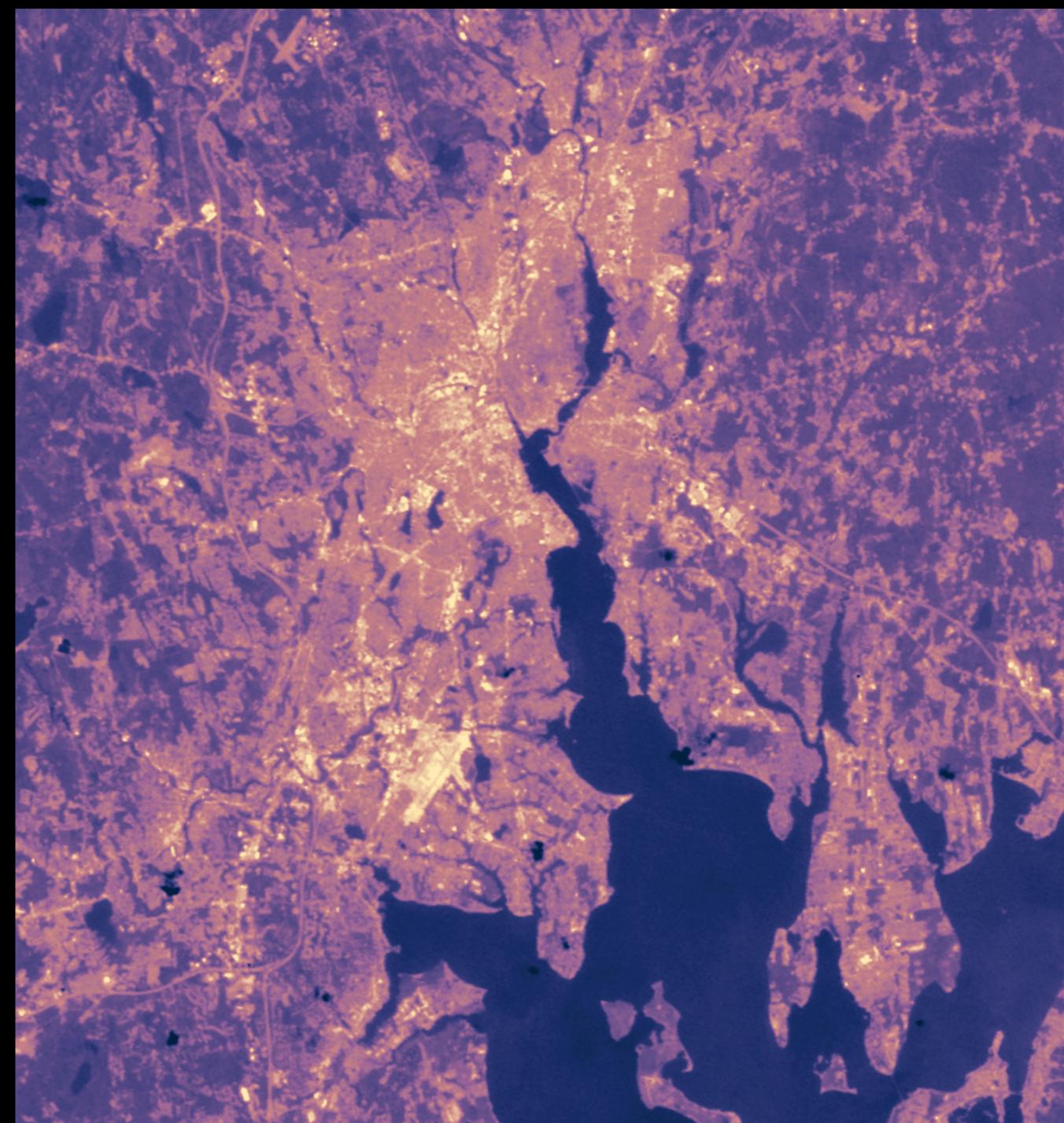


Urban Heat Island

Visible



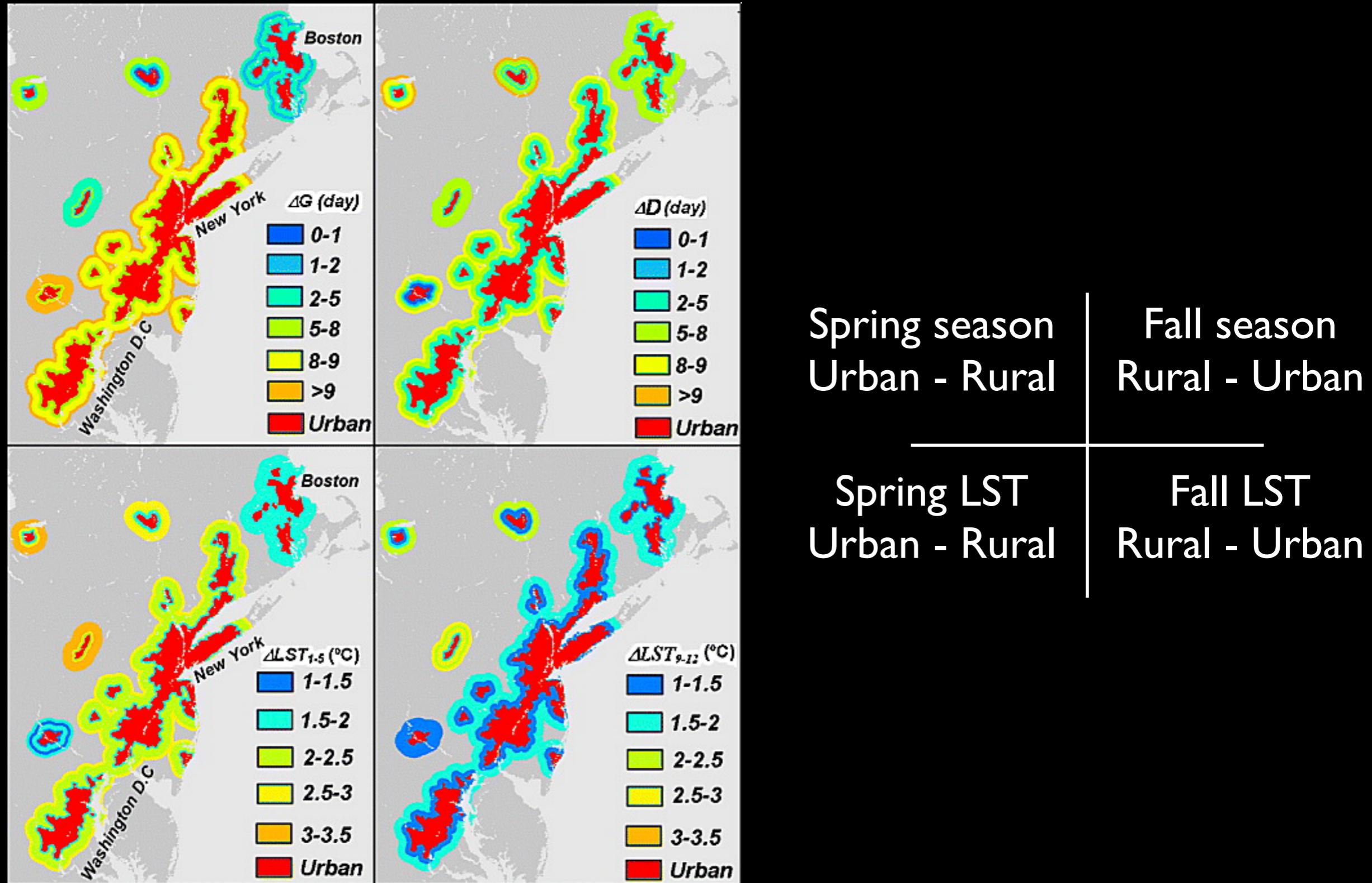
LST



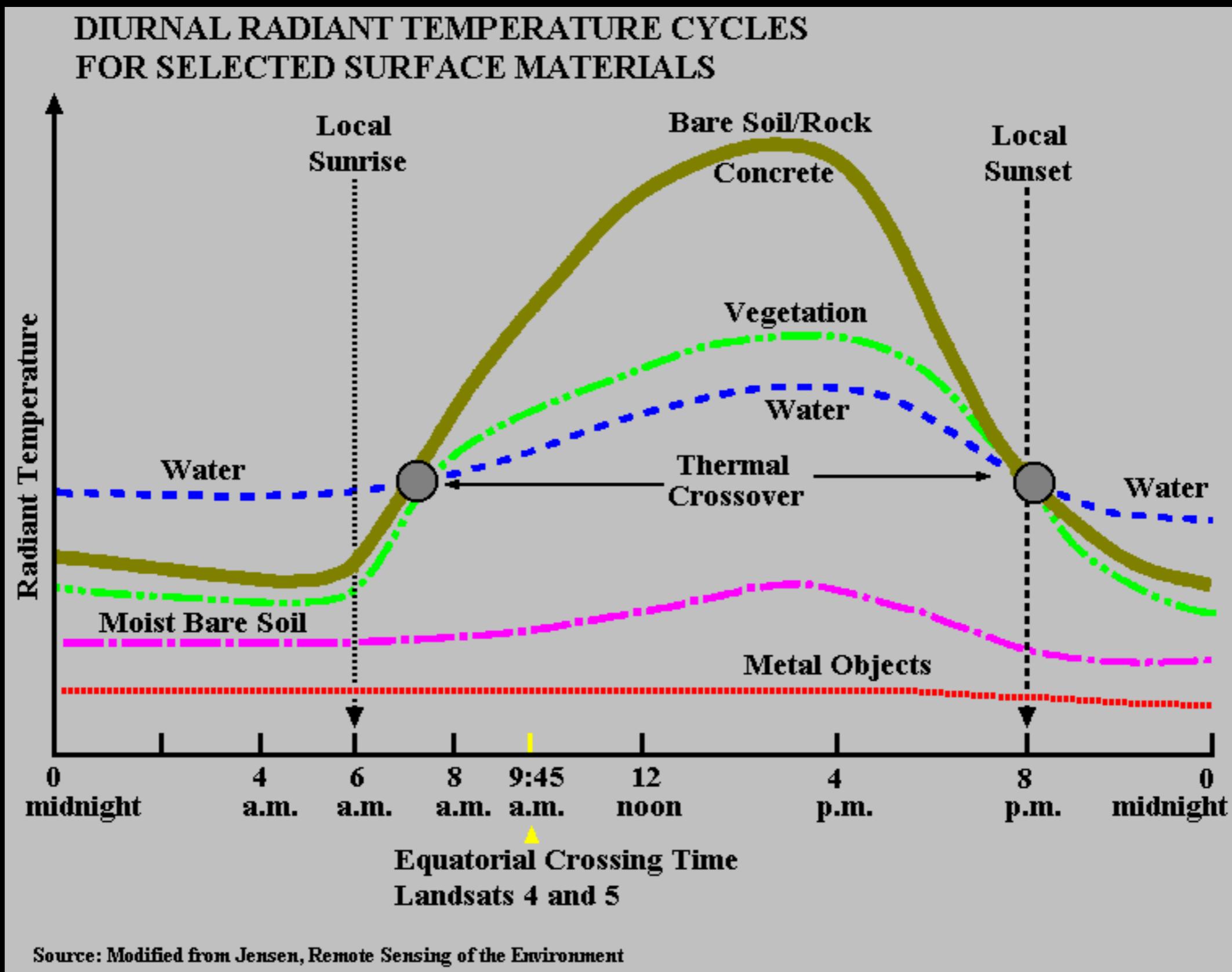
Brightness Temperature (°C)

20 30 40

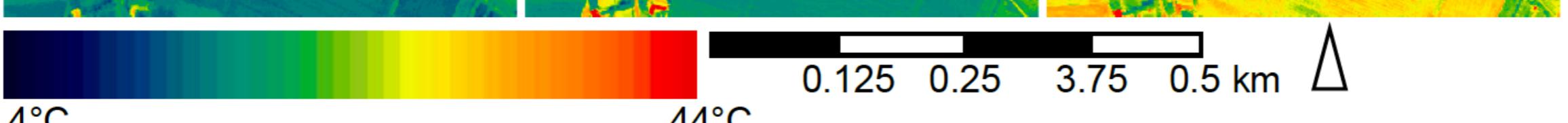
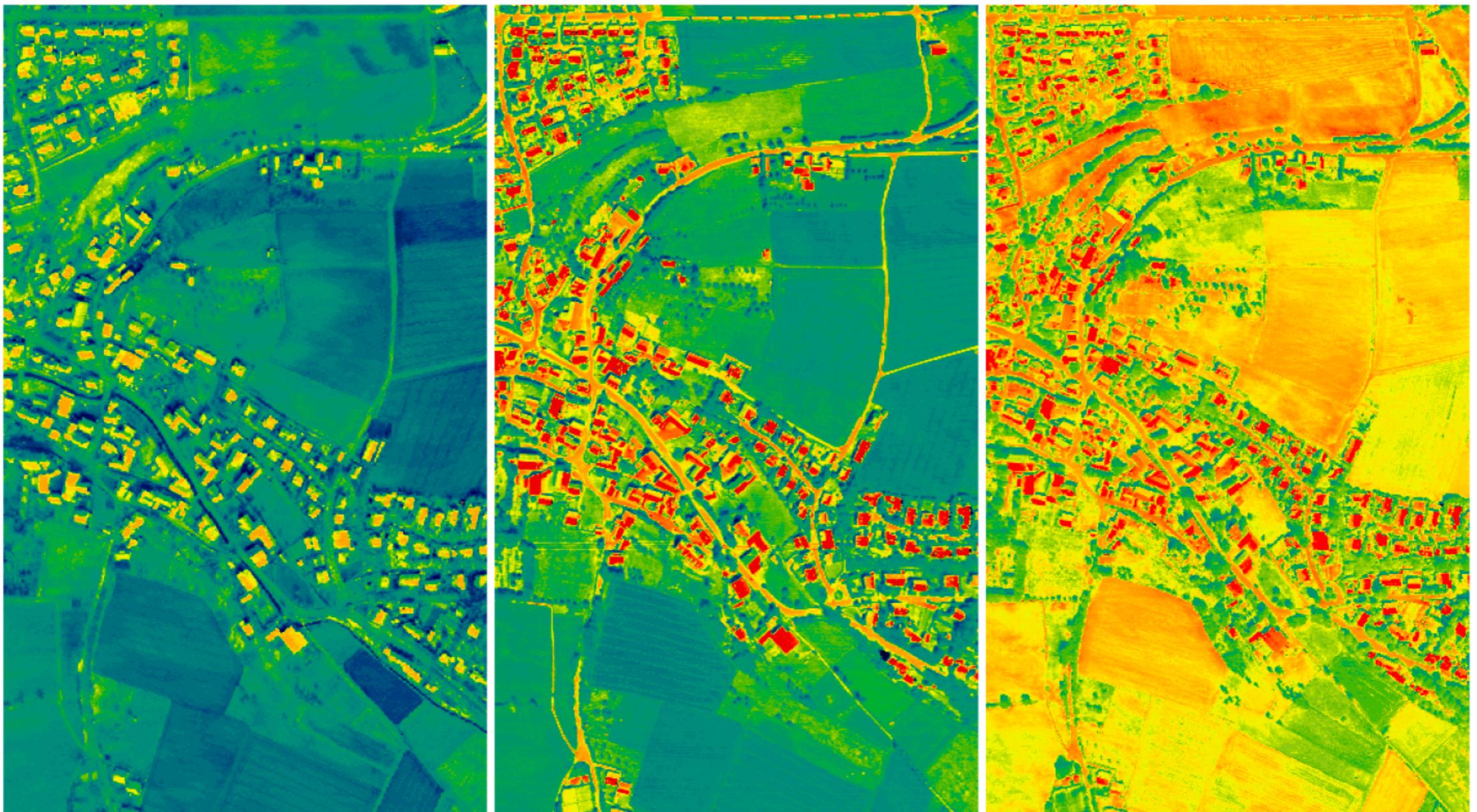
Urban Heat Island and its impact on vegetation



Diurnal patterns of surface temperature



TIR for change detection

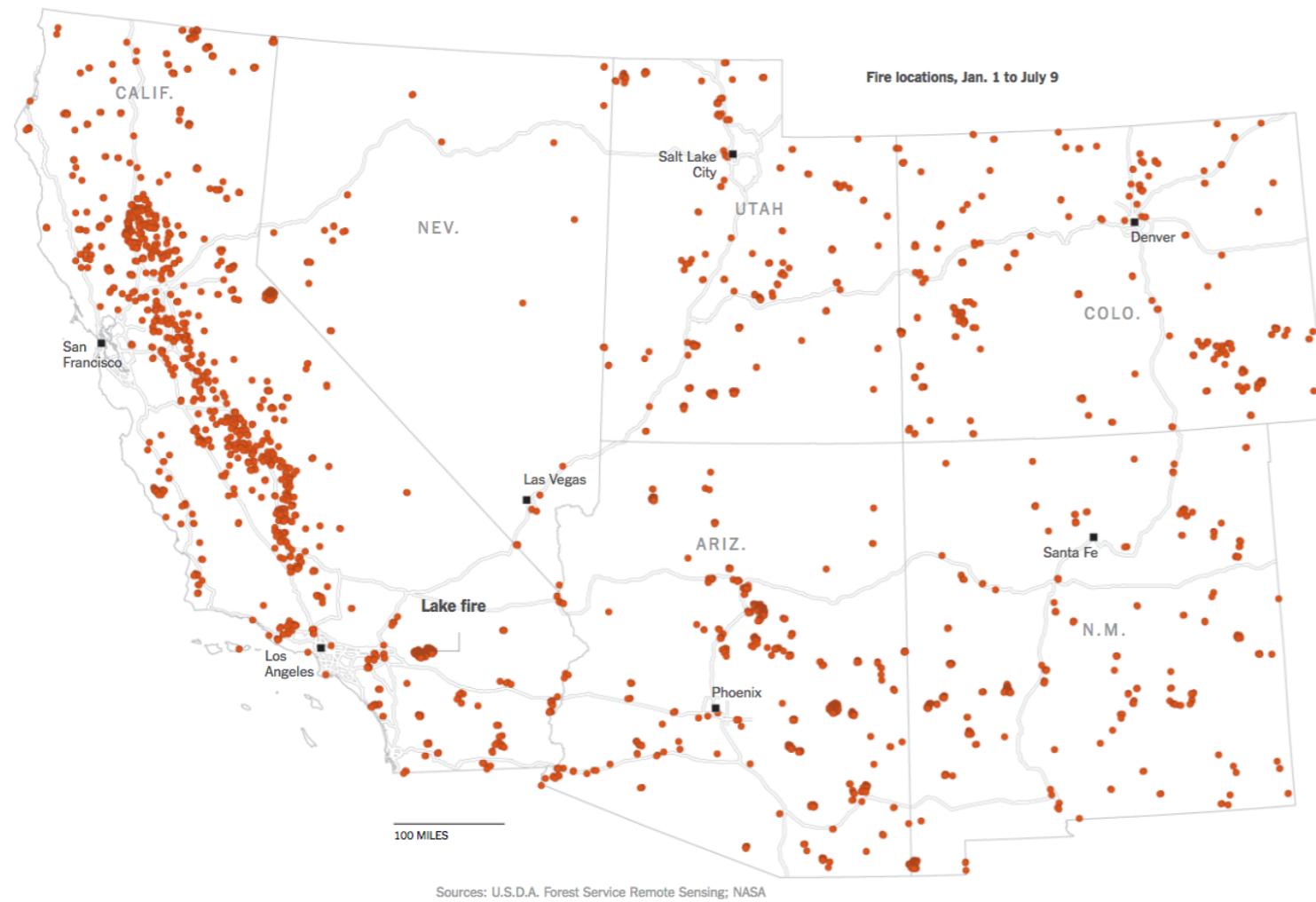


Winter

Spring

Summer

Forest Fire



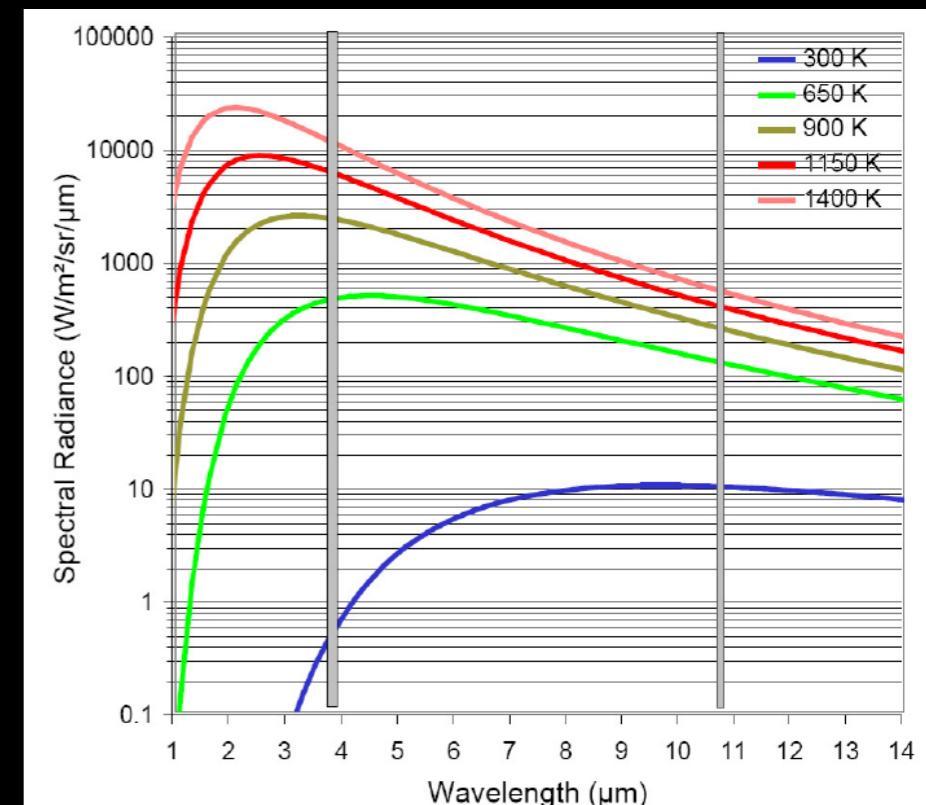
- Costly: the largest 10 fires in the U.S. last year costs \$320 million dollars
- Carbon emitter
- Difficult to monitor

<https://firms.modaps.eosdis.nasa.gov/firemap/>

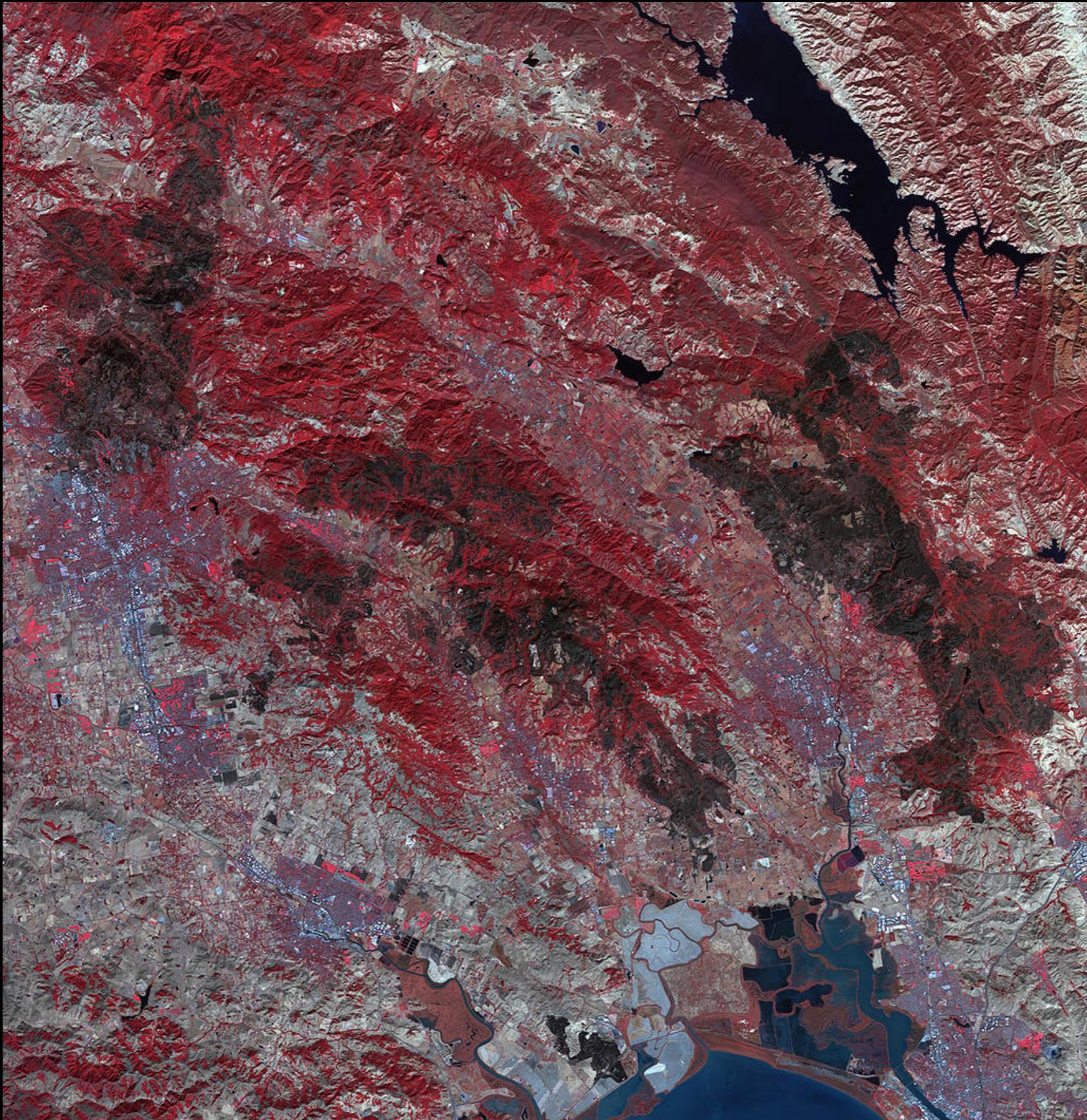


Remote Sensing of Forest Fire

- What aspects of fire are observable with remote sensing?
 - Direct optical signals (Light)
 - Thermal signals (heat) — which wavelengths do we consider?
 - Aerosols & smoke
 - Residue (char & ash)
 - Altered vegetation structure (burn scar)
- Active fire vs. Burned area



ASTER image (post-fire California)



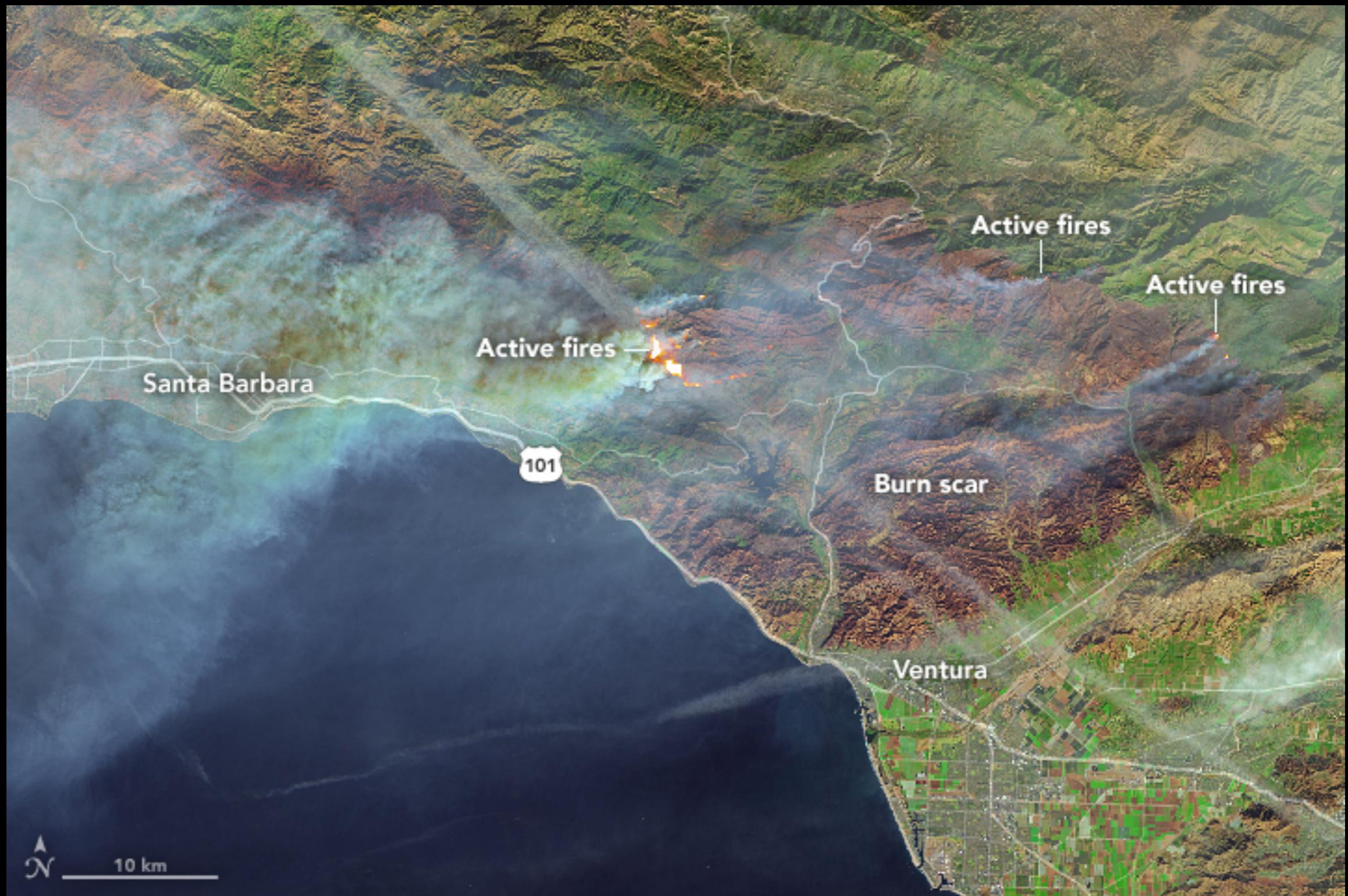
MODIS image



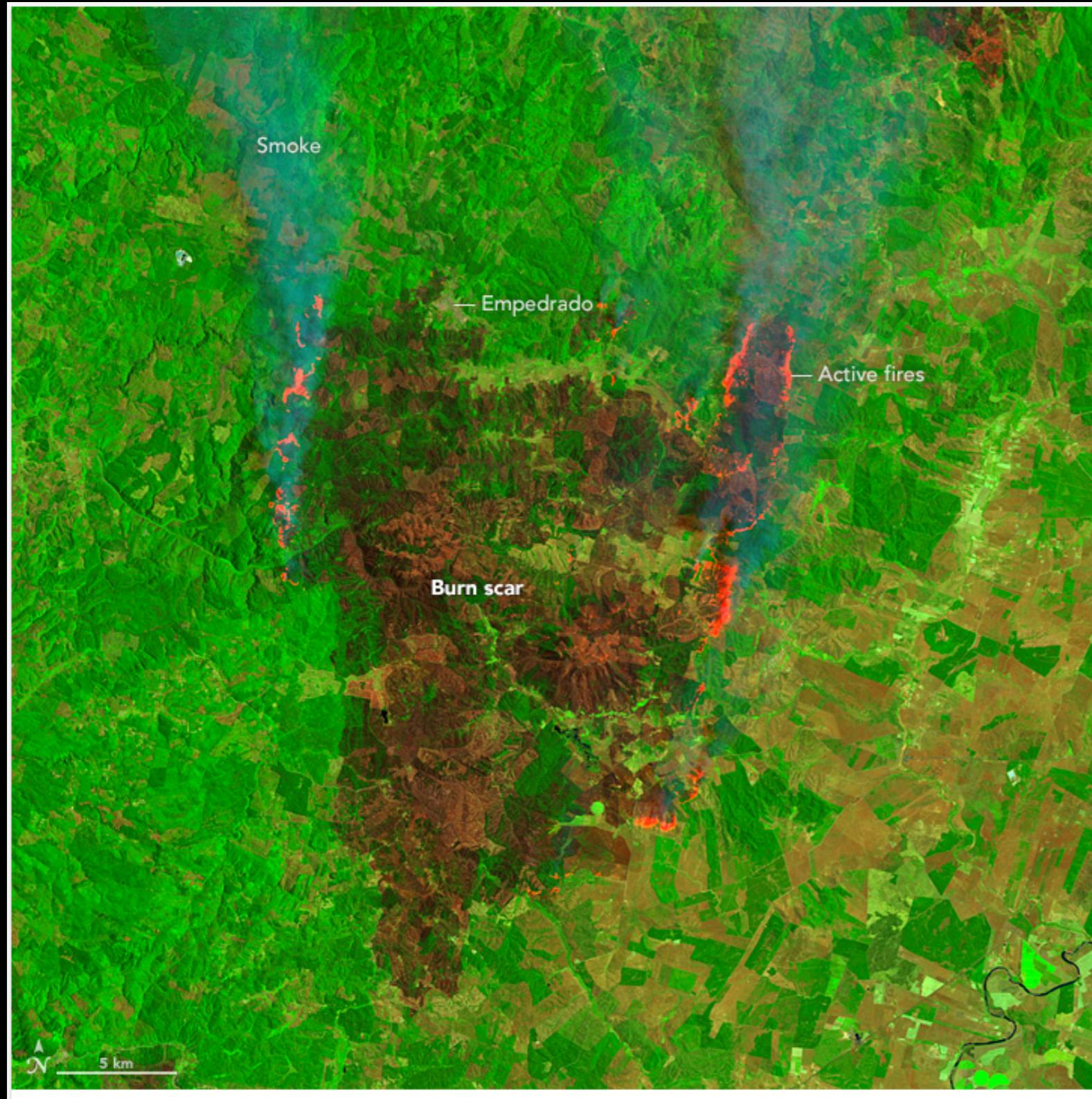
Landsat 8 OLI true color



Landsat 8 OLI Thermal Infrared + Visible



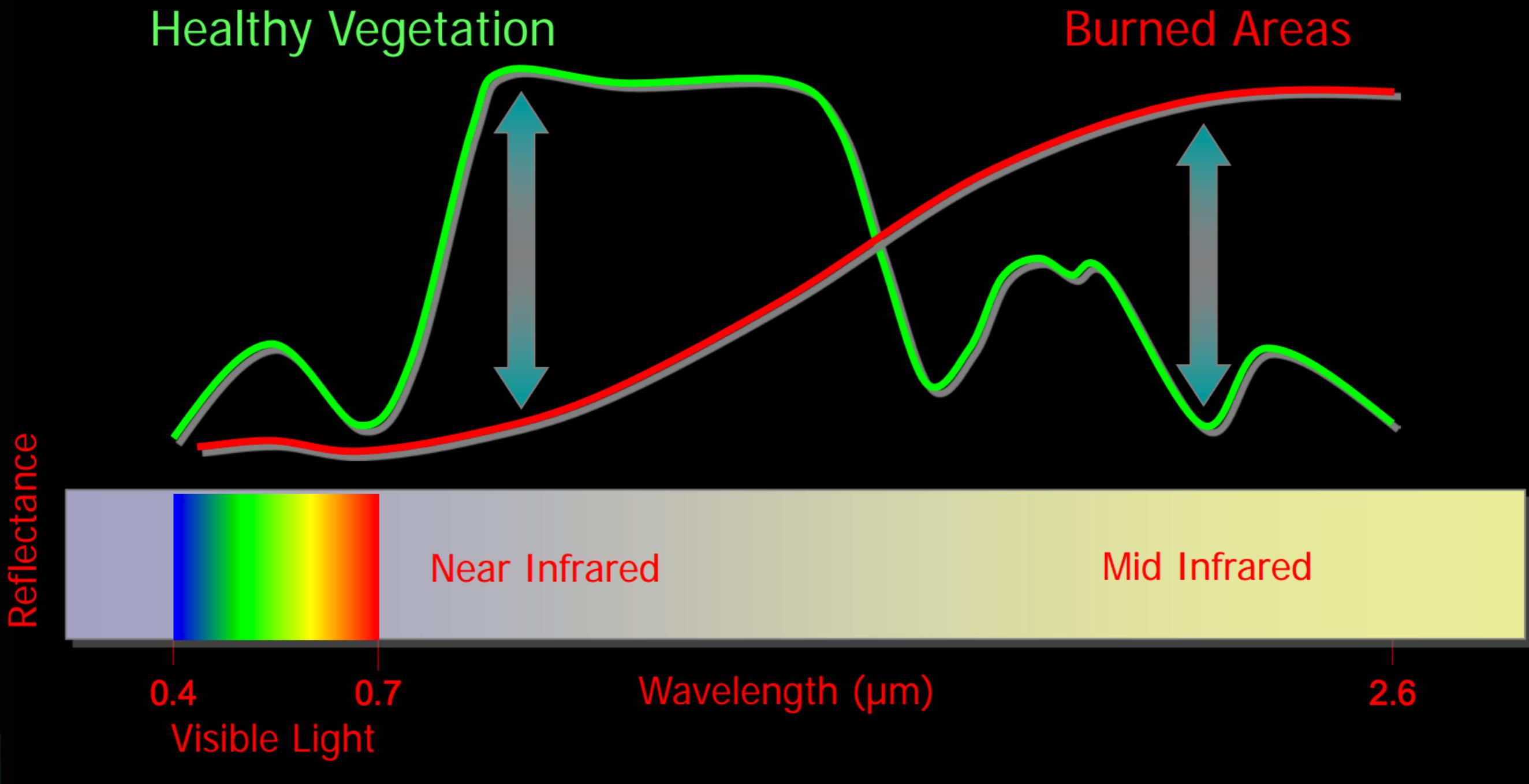
Burn scar vs. Active Fire



False color image:
Landsat 8, OLI
7-5-3

Find an index, like
NDVI, that identifies
the burn areas

Reflectance Spectroscopy



The Normalized Burn Ratio (NBR)
 $= (\text{NIR} - \text{SWIR}) / (\text{NIR} + \text{SWIR})$