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The authors



Gaspard Farge (presenting)

2nd year PhD student at the Institut de Physique du Globe de Paris, in the Geological Fluid dynamics team.

online presence : **ResearchGate**, **Twitter**

email: farge [at] ipgp [dot] fr

→ Claude Jaupart

Institut de Physique du Globe de Paris, in the Geological Fluid dynamics team.

→ Nikolaï Shapiro

Institut de Sciences de la Terre, Université Grenoble Alpes, CNRS (UMR5275), Grenoble, France

Schmidt Institute of Physics of the Earth, Russian Academy of Sciences, Moscow, Russia

















Model design

→ A permeable fault channel

We model the ascent of fluid along the permeable fault zone in a high permeability, 1D channel along the dip of the fault.

The channel is fed by a constant fluid flux q_{in} at its downdip end, from the dehydration of downdip slab segments.

At its updip end, the channel is connected to a highly permeable area (fractured continental crust), and the pressure is fixed to hydrostatic levels.

→ A valve mechanism

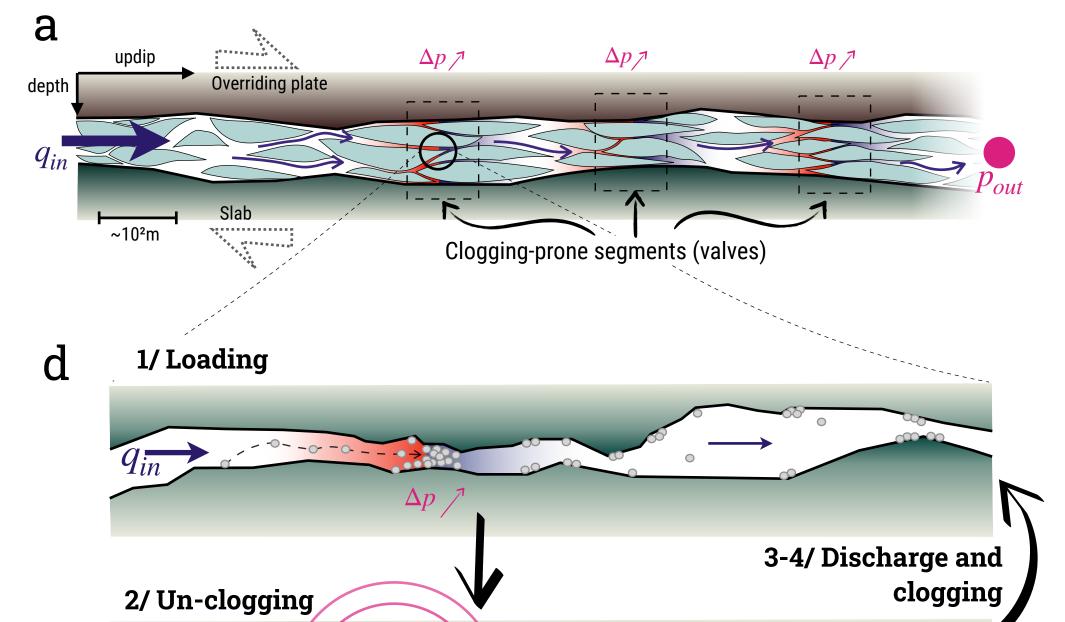
Fluid flow in the fault zone carries fine material (silica colloids or material eroded from the permeability walls) [5], that is subject to erosion and deposition processes, in response to the local velocity of the flow.

We model this phenomenon as permeability valves. In fixed locations, the permeability can open and close in response to the local pressure gradient: δp is high, then the valve opens, and δp drops, allowing the valve to close, with an hysteresis (e).

When a valve opens, it releases a strong fluid pulse because of the local pressure gradient build up while the valve is closed. This strong, rapid, localized pressure gradient can act as a source of seismic waves [6], or trigger seismic slip in local asperities.

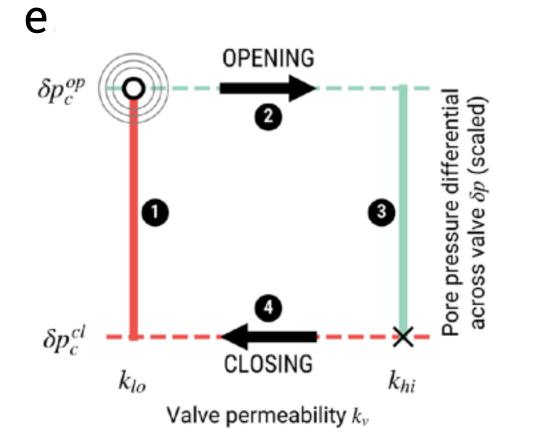
We therefore consider that each opening is accompanied by an LFE-like seismic event.

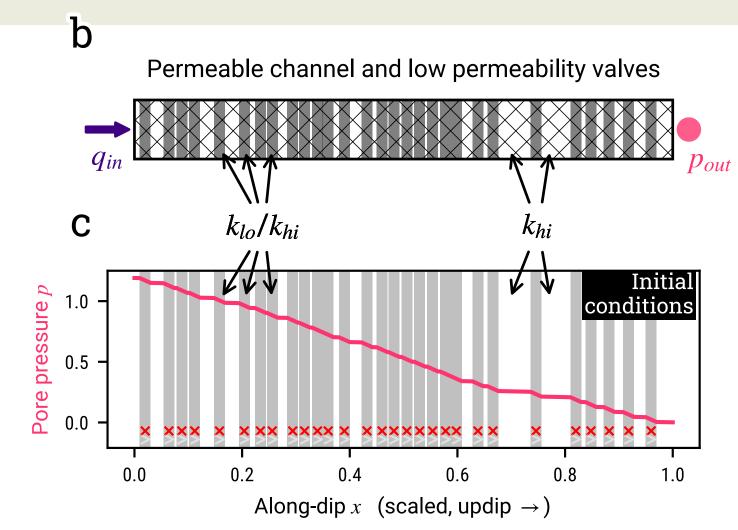
We solve for fluid pressure diffusion in a permeable channel with valves



A 1D permeable channel, with valves, along the dip of the fault

Schematic (a) and physical (b) representations of channel and permeability valves. (c) shows the initial fluid pressure conditions in a channel.





Valves: permeability heterogeneity and variability on an elementary scale

(d) Schematic representations of a valve cycle. (e) shows the hysteretical cycle that valve permeability follows in response to the pressure diff. across the







Go back

In this model,

Low-frequency earthquakes detections



the intermittence and migrations of tremor activity in subduction faults

Clogging and un-clogging of the subduction plumbing system may generate tremor-like patterns



Click where [+] for more! Go full screen for best results!

→ Background and motivation [+]

Variations of **fluid pressure** within subduction fault zones can **fuel fault-slip** events [1] and trigger the largest earthquakes [2].

In active subduction zones, fluid flow can be tracked through the activity of tremor and low-frequency earthquake (LFE) it seems to trigger.

Welcome to my display materials!

le build a model to explore how fluid transport in a fault shapes its seismicity.

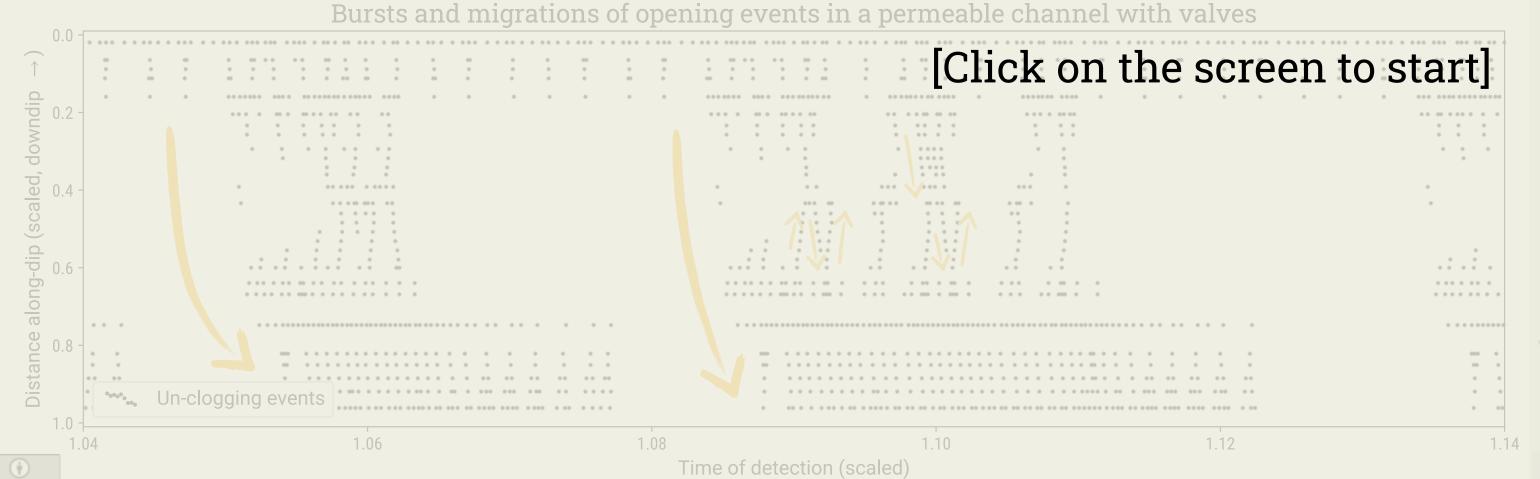
The online version is a bit slow, give it time before clicking on links, and refresh it if it crashes!

Have a good read!

Time of detection (dd/mm/yyyy)

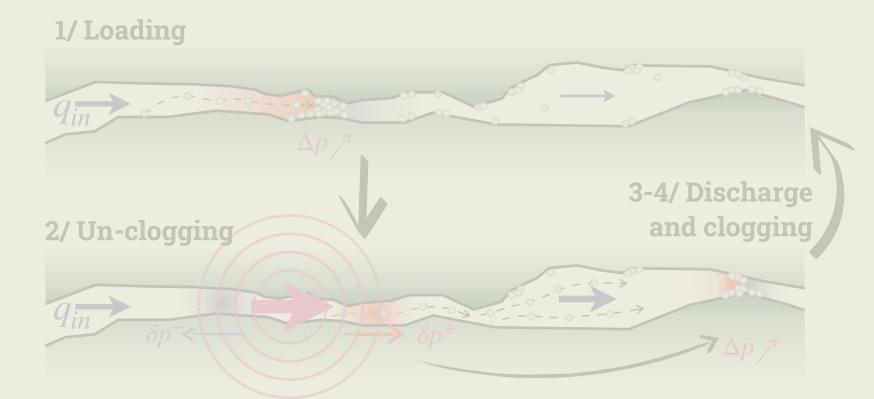
emerges from cascades and interactions of un-clogging and clogging in permeable channels

Bursts and migrations of low-frequency earthquake activity in Guerrero, Mexico



We solve for fluid pressure diffusion in a 1D, dynamically permeable channel along-dip. Locally, several **permeability valves** open/close in response to the pressure gradient.

The strong, rapid pressure transient at opening can act as (or trigger) a seismic source.



Results [+]

As seismic sources interact through rapid fluid pressure transients, events occur in cascades. Activity is **clustered** and **migrates** in the channel.

The input flux in the fault zone controls if activity occurs, and shapes its intermittence.

This simple conceptual model shows how variable and realistic patterns of seismicity can be driven by dynamic permeability in a constantly fed fluid transport system.

Background and motivation

→ Fluid in subduction zones

As it is subducted, the downgoing slab dehydrates, and free fluid then pushes its way towards the surface, at quasi-lithostatic pressures.

At 40 km depth, the fluid preferentially follows the fault interface, whose very anisotropic permeability is a barrier to the upward flow, but channels fluid along the fault plane.

Variations of **fluid pressure** induced by variations in fluid flow within subduction fault zone can fuel fault-slip fast and slow slip events [1] and thus be a trigger for the largest earthquakes [2].

→ Tremor and LFE activity as proxy for fluid transport processes

Tremor migrations move quickly on short distances, and slowly on larger distances, leading us to think that sources interact through diffusive interaction mainly, perhaps through fluid pressure transients.

In active subduction zones, the activity of swarms of low-frequency earthquakes (LFE)* and tremor is often thought to be caused by pulses of coupled slow fault slip and high fluid pressure [1, 3, among others] in a mechanism akin to a fault valve [1, after 4].

Based on the hypothesis that tremor activity reflects the dynamics of fluid transport processes in the fault,

we build a model to explore how the diffusion of fluid pressure in a fault with heterogeneous and variable permeability could shape its seismicity. **Summary**

Background and motivations

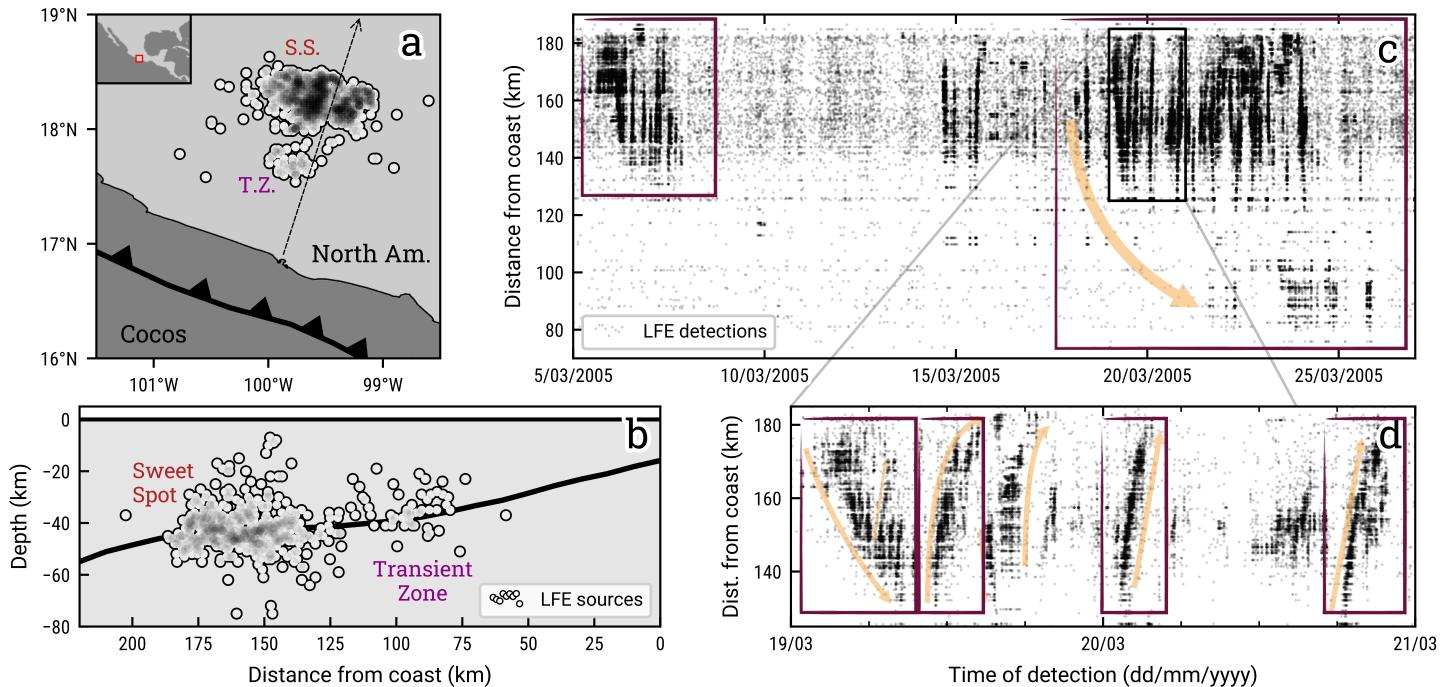
Model design

Results: patterns

Results: flux

Authors

Anatomy of low-frequency earthquakes activity in Guerrero, Mexico

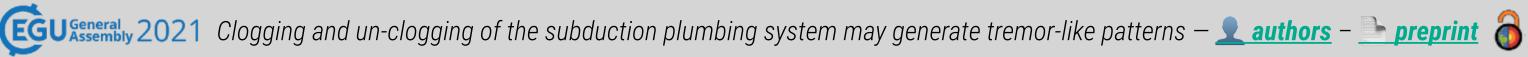


Tremor activity occurs is intermittent on both short and large time and space scales

it occurs in **bursts** and **migrates**

How does fluid transport processes shape these patterns?







^{*} Here we assume that LFEs are elementary tremor sources, and that they are generated by the same underlying processes. "tremor activity" and "low-frequency earthquakes activity" are used as synonyms.

openings

Results — (i)

→ Valve (source) interaction

The opening of a valve generates a strong pressure transient that diffuses to neighboring **valves**. Both updip and downdip valves are brought closer to failures as their δp increases.

This mechanism is the basis of constructive source interaction in our model, and generates fast cascades of valve openings, and therefore of seismic events.

→ Patterns of clustering and migration

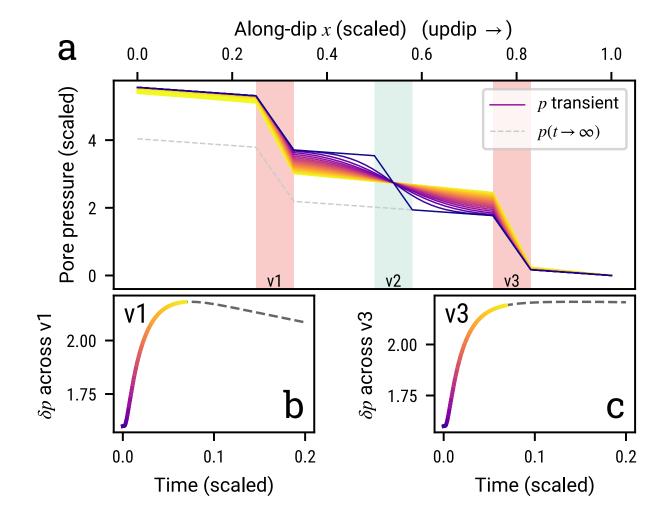
Go back

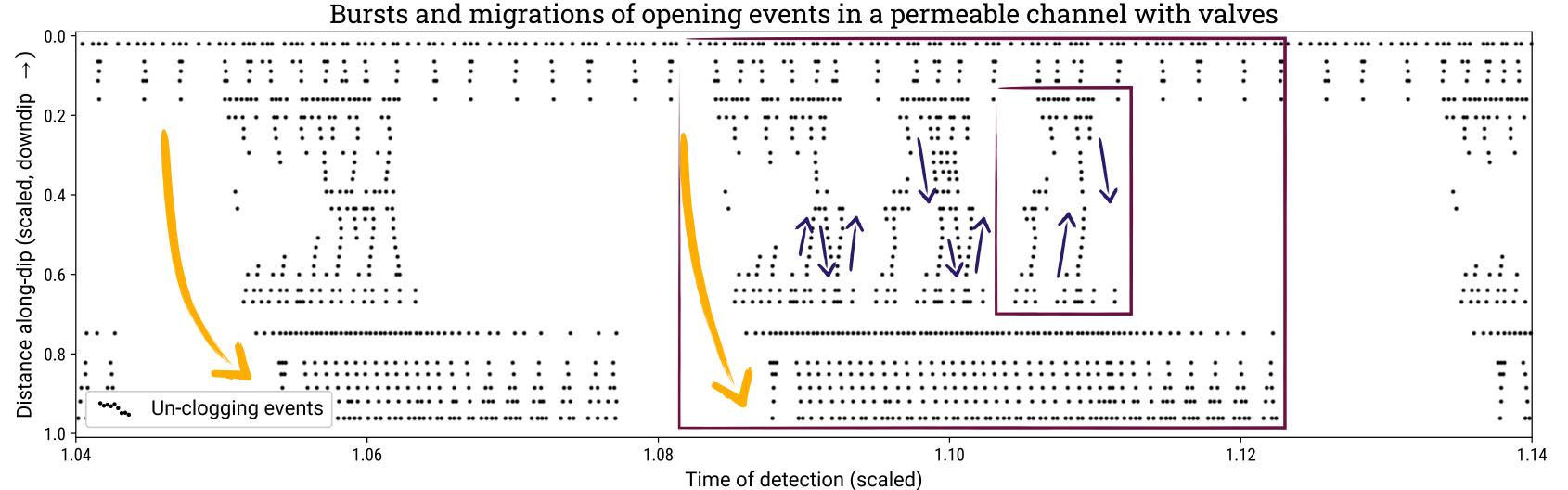
The synthetic seismic activity (opening activity) organizes around those rapid cascades of events. They are the building blocks of larger bursts and migrations, on both short and longer time and space scales.

Overriding plate **Cascading valve**

The opening of v2 sends a bidirectional pressure transient that raises the δp at v1 and v3, bringing them closer to failure

In the model, seismic sources interact with each other through pressure transients





patterns of **clustering** and **migration** emerge from these elementary cascades







Results — (ii)

→ Input flux controls the activity

In a range of input flux — depending of the valves' sensitivity to the pressure gradient and permeability — valves can both open when closed and close back when open. In this flux range, activity is permanent and variable.

Within the domain of flux allowing permanent activity, different regimes can be observed:

- Low flux, valves are mostly closed, and open in bursts: activity is low and time-clustered
- Intermediary flux, valves open and close rapidly, randomly but at a constant rate: activity is high and Poisson-like
- High flux, valves are mostly open, and when they close large bursts occur as they impede the flow: activity is low, and quasi-periodic

Under a **constant source** of fluid pressure,

the proposed fluid transport processes, generate and sustain

> realistic patterns of intermittence and migration of seismic activity

> > Go back

In a range of input flux q_{in} , activity is permanent and

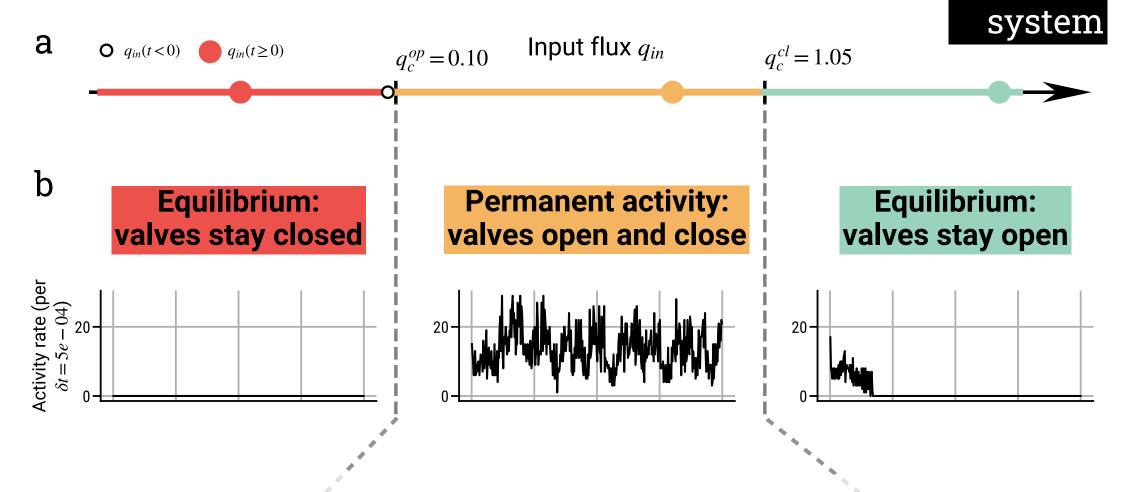
Model design

Background and

motivations

intermittent

Summary



Results: patterns

Results: flux

Authors

29 valves

within this range, q_{in} governs the activity style: time-clustered, constant,

or quasi-periodic

