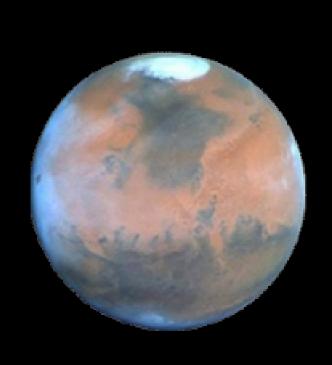
Energy Balance and Climate

EES 3310/5310
Global Climate Change
Jonathan Gilligan

Class #3: Friday January 10 2020

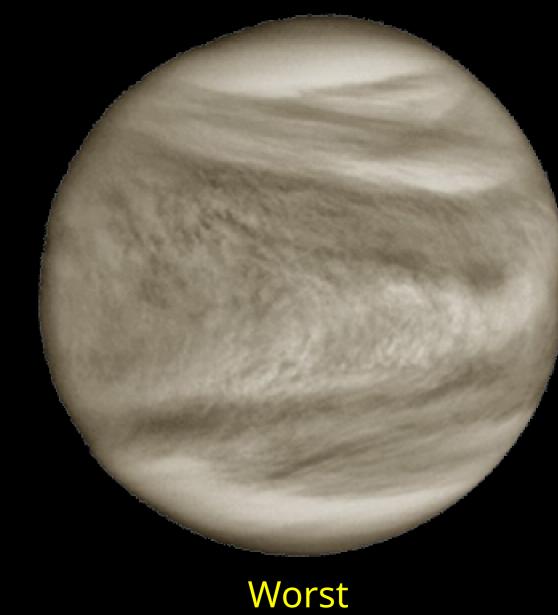
Looking for a Good Home



 $\begin{array}{c} \text{Bad} \\ -28^{\circ}\text{F} \end{array}$



Good 71°F



800°F

Basic Concepts

Vocabulary

- Energy, Heat:
 - Heat = energy flowing spontaneously from hot to cold
- Power: speed at which energy flows or transforms

Power, Flux = Heat flow/Time

Heat, Energy = Power \times Time

• Intensity: Concentration of power

Intensity = Power/Area

Power = Intensity \times Area

Temperature of a planet

• Basic principle:

Steady temperature if and only if

$$Power_{in} = Power_{out}$$

- How can heat get in or out?
 Electromagnetic radiation

Electromagnetic Waves

- Color and brightness
- Color:
 - Two ways to measure color:
 - Wavelength (λ)
 - Wavenumber ($n=1/\lambda$)
- Archer mostly uses wavenumber
 - Math is simpler that way
- Brightness:
 - Intensity (power/area, Watts/square meter)

Colors

Color	wavelengths	wavenumbers	
infrared	> 0.70 µm	< 14,000	
red	~ 0.70−0.64 µm	~ 14,000-16,000	
orange	~ 0.64−0.59 µm	~ 16,000-17,000	
yellow	~ 0.59−0.56 µm	~ 17,000-18,000	
green	~ 0.56-0.49 µm	~ 18,000-20,000	
blue	~ 0.49-0.45 µm	~20,000-22,000	
violet	~ 0.45-0.40 µm	~22,000-25,000	
ultraviolet	< 0.40 µm	> 25,000	

All you need to think about is shortwave vs. longwave radiation.

Shortwave and longwave:

- Shortwave:
 - Near-infrared, visible, ultraviolet
 - lacksquare $\lambda < 3 \mu \mathrm{m}$
 - n > 3,300cm⁻¹ (cycles per centimeter)
- Longwave:
 - Mid-infrared, far-infrared
 - lacksquare $\lambda > 3 \mu \mathrm{m}$
 - $n < 3,300 \text{cm}^{-1}$

More on this in the next class ...

4 Laws of Radiation

- 1. All objects continually radiate energy
- 2. Hotter objects are brighter
- 3. Hotter objects radiate at shorter wavelengths
- 4. Objects that are good absorbers are also good emitters
 - Black objects emit & absorb the most
 - Transparent and white objects emit & absorb the least

Blackbody Radiation

Blackbody Radiation

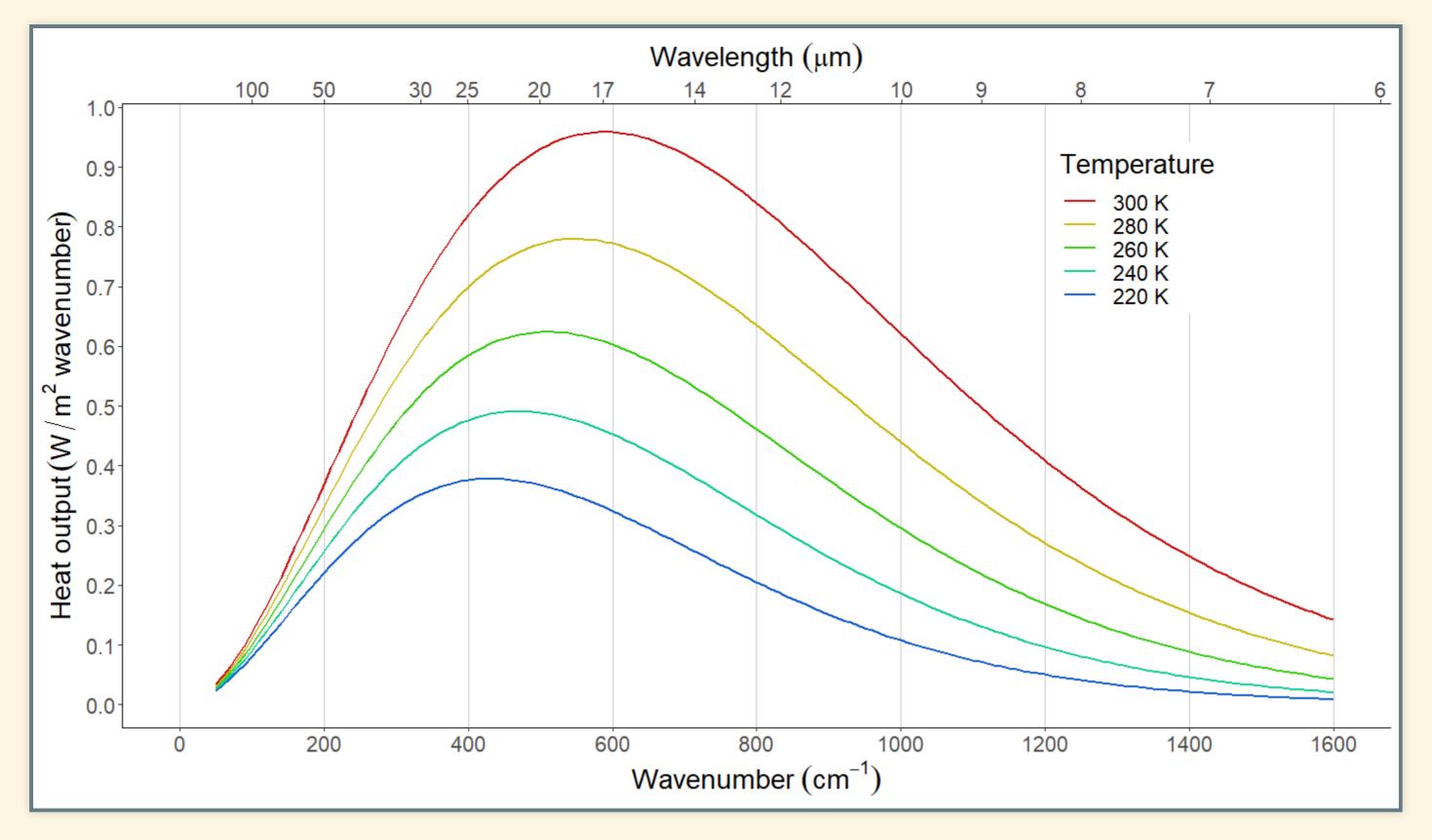
Emissivity (ε) measures how black something is:

- $oldsymbol{arepsilon} arepsilon = 1$ for perfectly black
- $\varepsilon = 0$ for perfectly white or transparent
- In between for gray.
- Black, white, and gray: ε is the same for all wavelengths.
- Colored objects: ε is a function of wavelength.
- For simplicity: start by assuming everything is black, white, or gray.

Remember: Good emitters are good absorbers

Fundamental rule: Temperature and emissivity determine radiation.

Heating Up: What Changes??



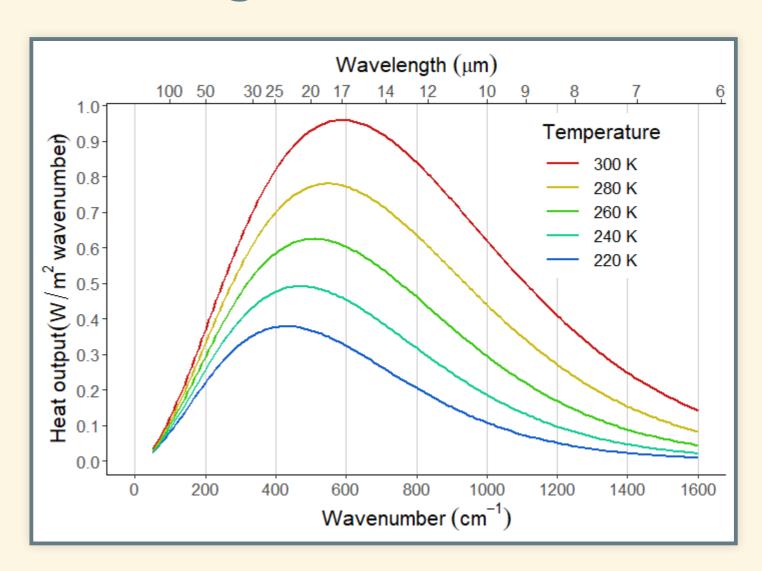
Heating Up: What Changes?

- Hotter temperature:
 - Brighter (greater intensity)
 - Bluer (greater wavenumber, shorter wavelength)

A curious thing:

A hot black object glows with color!

Total intensity = area under curve



Mathematical Description

Blackbody Radiation

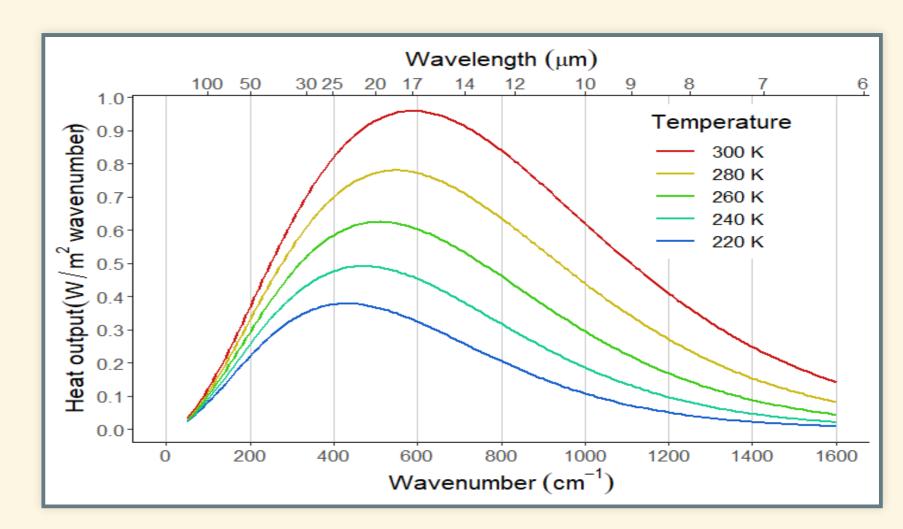
Intensity (brightness):

Stefan-Boltzmann law

$$I = \varepsilon \sigma T^4$$

after Josef Stefan and Ludwig Boltzmann

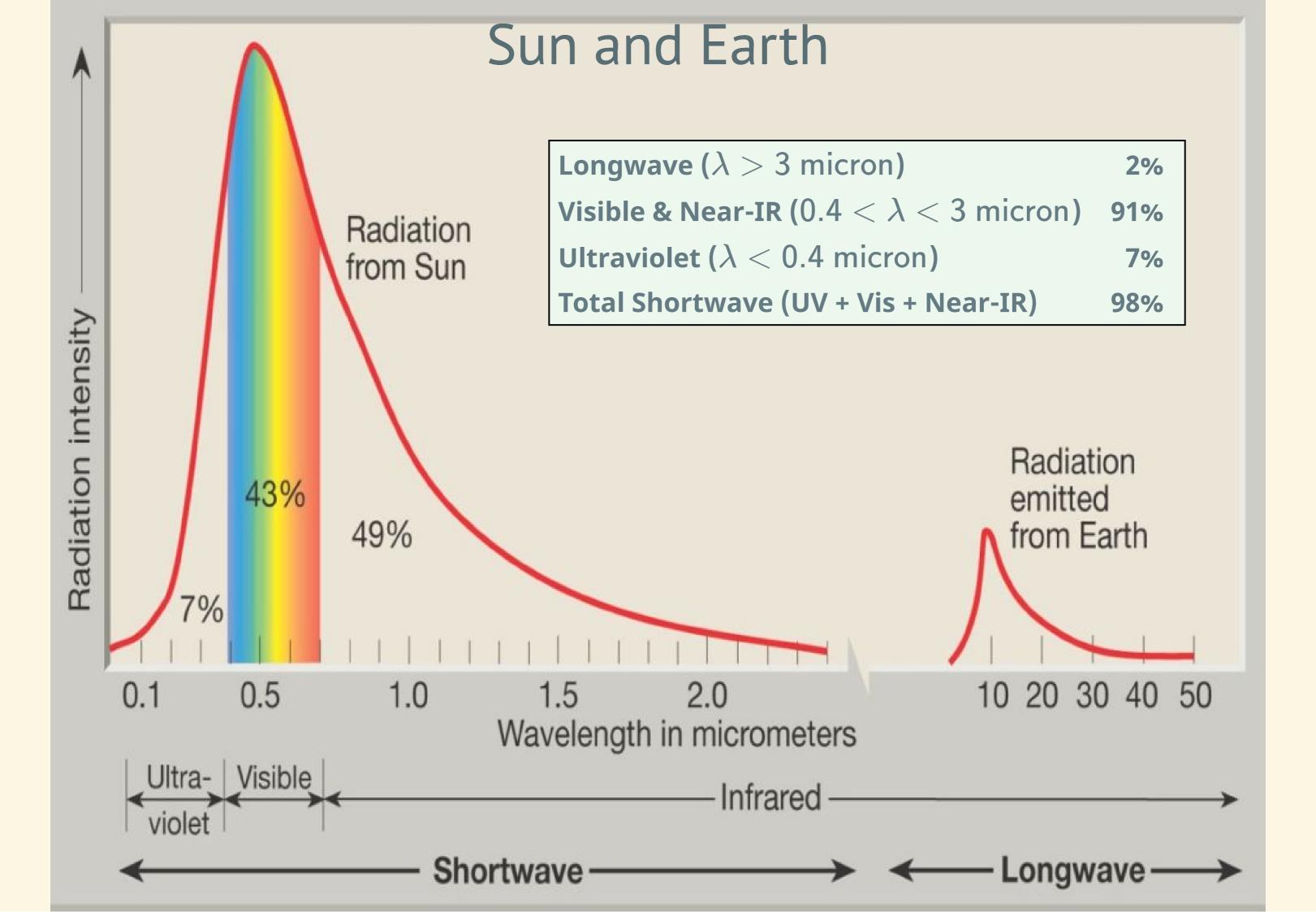
- ε = emissivity
 - Different for different objects.
- σ = Stefan-Boltzmann constant.
- T = absolute (Kelvin) temperature.



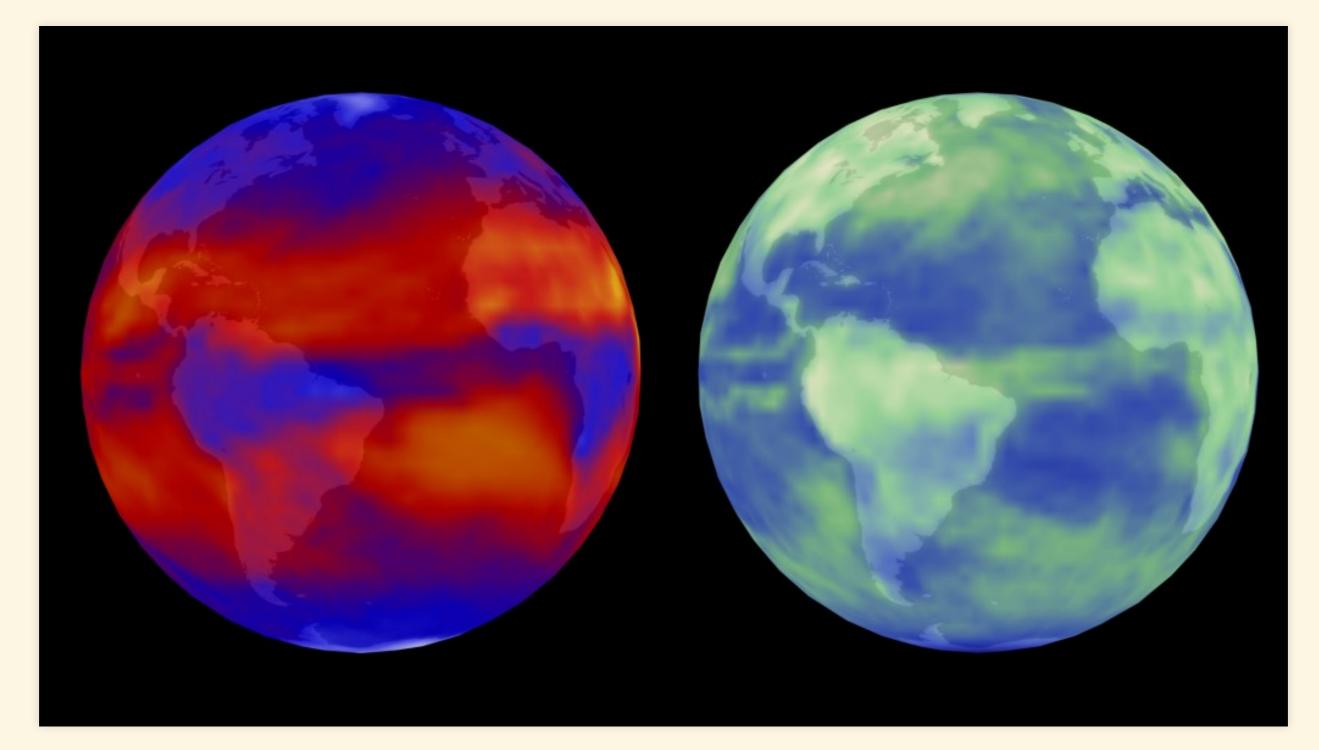
Color: Peak wavenumber proportional to (Kelvin) temperature.

Helpful Hint:

Fourth power on a calculator: press the x^2 button twice.



Earth and Radiation



False-color images of radiation from Earth, seen by NASA Terra satellite:

- Left: Thermal radiation (blue \rightarrow red \rightarrow yellow = dim \rightarrow bright)
- Right: Reflected sunlight (blue → green → white = dim → bright)

Efficiency of Light Bulbs

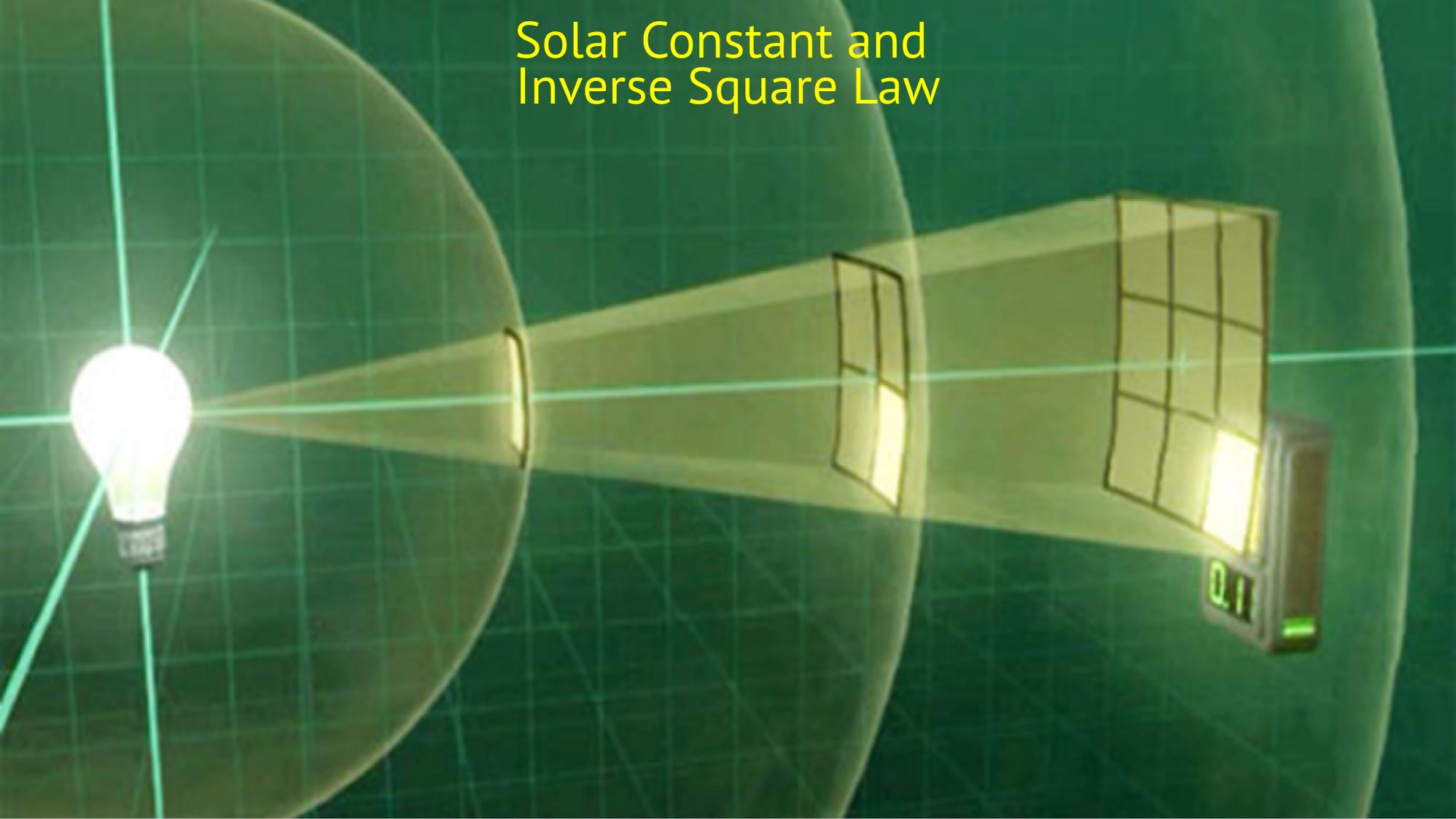
Type of Bulb	Efficiency
Standard 40W	1.8%
Standard 60W	2.1%
Standard 100W	2.6%
Quartz Halogen	3.5%
Ideal black body @ 7000K	14.0%
Compact Fluorescent	8-12%
LED	20-44%

- 7000K is the optimal temperature for a black body to emit visible light, but it will melt every known substance.
- Standard light bulbs operate at around 2000–3300 K.

Calculating Earth's Temperature: Bare-Rock Model

Basics Steady Temperature

- Heat in must balance heat out
- Total Heat Flux (Power) = Area × Intensity
 - Total heat flux in (F_{in}) :
 - Intensity depends on solar constant and albedo
 - Does not depend on earth's temperature
 - Total heat flux out (F_{out}):
 - Intensity depends on earth's temperature and emissivity
- Strategy:
 - 1. Figure out F_{in} .
 - 2. Figure out temperature T that makes $F_{\text{out}} = F_{\text{in}}$.

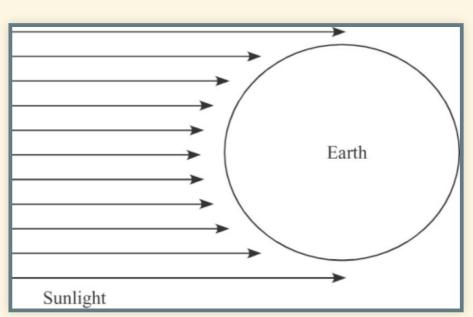


What is F_{in} ?

- $F_{in} = Area \times Intensity absorbed$
 - Intensity absorbed = $(1 \alpha) \times I_{in}$
 - $I_{in} = 1350 \text{ W/m}^2$
 - \circ Average albedo lpha= 0.30 (30% of sunlight is reflected)

What is area?

- Area = silhouette or shadow
- Circle: πr^2



What is F_{in} ?

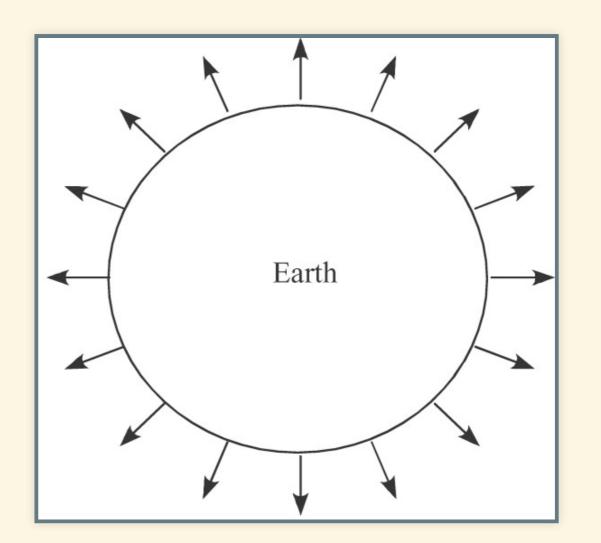
•
$$F_{\mathsf{in}} = \pi r_{\mathsf{Earth}}^2 imes (1-lpha) I_{\mathsf{in}}$$

$$\pi r^2 = 1.3 \times 10^{14} \text{m}^2$$

- $I_{in} = 1350 \, W/m^2$
- $F_{\rm in} = 1.2 \times 10^{17} \, {\rm Watt}$
 - 11,000 times total human energy production.

What is F_{out} ?

- $F_{\text{out}} = \text{Area} \times I_{\text{out}}$
 - $I_{\text{out}} = \varepsilon \sigma T^4$
 - $\circ \ arepsilon = 1$ (blackbody)
 - $\sigma = 5.67 \times 10^{-8} \text{ W/m}^2/\text{K}^4$
 - What is area?
 - Sphere: $4\pi r^2$
- $F_{\rm out} = 4\pi r_{\rm earth}^2 imes \varepsilon \sigma T^4$



Putting it all together

$$F_{
m out} = F_{
m in}$$
 $4\pi r^2 imes arepsilon au \sigma T^4 = \pi r^2 (1-lpha) I_{
m in}$
 $4\pi r^2 imes arepsilon \sigma T^4 = \pi r^2 (1-lpha) I_{
m in}$
 $4arepsilon \sigma T^4 = (1-lpha) I_{
m in}$

$$T^4 = rac{(1-lpha)I_{\mathsf{in}}}{4arepsilon\sigma}$$

$$T=\sqrt[4]{rac{(1-lpha)I_{
m in}}{4arepsilon\sigma}}$$

Temperature of Earth

- Steady Temperature:
 - Heat flux in must balance heat flux out ($F_{out} = F_{in}$).
 - *F*_{in}:
 - Does not depend on earth's temperature.
 - Depends on solar constant and earth's albedo.
 - **■** *F*_{out}:
 - Depends on earth's temperate
 - T adjusts until heat out = heat i

Helpful hint:

To take the fourth root on a calcula press the square-root key $(\sqrt{})$ twice

$$T=\sqrt[4]{rac{(1-lpha)I_{\mathsf{in}}}{4arepsilon\sigma}}$$

Temperature of Earth

$$T=\sqrt[4]{rac{(1-lpha)I_{\mathsf{in}}}{4arepsilon\sigma}}$$

Earth:

(Note: My numbers are slightly different from Archer's textbook)

- $I_{in} = 1350 \text{ W/m}^2$
- $\alpha = 0.30$
- $\varepsilon = 1$
- $\sigma = 5.67 \times 10^{-8} \text{ W/(m}^2 \text{K}^4)$
- Calculate *T*:
- $T = 254 \text{ K} = -19^{\circ} \text{C} = -2^{\circ} \text{F}.$

If the sun got 5% brighter, how much warmer would the earth become?

$$T=\sqrt[4]{rac{(1-lpha)I_{
m in}}{4arepsilon\sigma}}$$

- Normal: $l_{in} = 1350 \text{ W/m}^2$:
 - T = 254 K
- 5% Brighter: $I_{in} = 1.05 \times 1350 \text{ W/m}^2 = 1418 \text{ W/m}^2$:
 - T = 257 K
- $\Delta T = 3 \text{ K} = 6^{\circ} \text{F}$

Temperature of Earth

$$T=\sqrt[4]{rac{(1-lpha)I_{\mathsf{in}}}{4arepsilon\sigma}}$$

Earth:

(Note: My numbers are slightly different from Archer's textbook)

- $I_{in} = 1350 \text{ W/m}^2$
- $\alpha = 0.30$
- $\varepsilon=1$
- $\sigma = 5.67 \times 10^{-8} \text{ W/(m}^2 \text{K}^4)$
- $T = 254 \text{ K} = -19^{\circ} \text{C} = -2^{\circ} \text{F}.$

How does this compare to Earth's actual temperature?

Radiative Temperature

- Satellites orbiting in space can measure longwave radiation from earth
- To the satellites, the earth looks very much like a blackbody at the bare-rock temperature (254 K).
- Thus, scientists generally call the bare-rock temperature the **radiative temperature** because it describes the radiation coming off the earth.
- However, the surface temperature of the earth is around $295~\text{K}=71^{\circ}\text{F}$, which is significantly different from the radiative, or bare-rock, temperature.

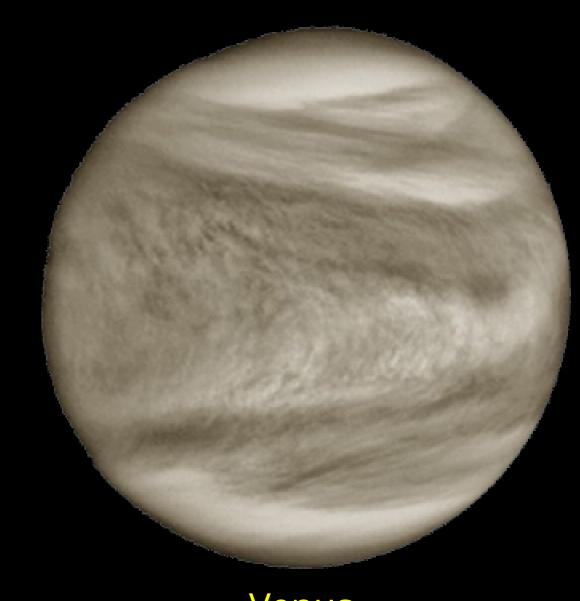
The Terrestrial Planets



Mars 240 K



Earth 295 K



Venus 700 K

Terrestrial Planets

	Earth	Mars	Venus
Distance from sun	1 AU	1.5 AU	0.72 AU
1/Distance ²	1.00	0.44	1.9
Solar constant	1350 W/m^2	$600 \mathrm{W/m^2}$	2604 W/m ²
Albedo	0.30	0.17	0.71
T _{bare rock}	254 K (-2°F)	216 K (-70°F)	240 K (-27°F)
T _{surface}	295 K (71°F)	240 K (-28°F)	700 K (800°F)
$\Delta_{\mathcal{T}}$	41 K (74°F)	24 K (42°F)	460 K (828°F)

- Total flux (power) radiated from sun doesn't change with distance.
- At a distance *r* total flux spreads over sphere of radius *r*
- Intensity = Total Flux / Area:
 - Proportional to $1/r^2$.
- At edge of Earth's atmosphere, solar intensity = 1350 W/m^2 .