Instructions for Lab #6

The Geochemical Carbon Cycle

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Carbon Cycle

For the following exercises, you will use the GEOCARB model, which simulates the earth's carbon cycle.

The GEOCARB model has two time periods:

- First, it runs for 5 million years with the "Spinup" settings in order to bring the carbon cycle and climate into a steady state. Only the last 1000 years of the spinup are saved.
- Then, at time zero, it abruptly changes the parameters to the "Simulation" settings and also dumps a "spike" of CO₂ into the atmosphere and runs for another 2 million years with the new parameters to see how the climate and carbon cycle adjust to the new parameters and the CO₂ spike.

The quantities that are graphed in the online version of the model include:

pCO2 is the concentration of CO_2 in the atmosphere, in parts per million.

WeatC is the rate of CO₂ being weathered from carbonate rocks and moved to the oceans.

BurC is the rate of carbonate being converted into limestone and buried on the ocean floor.

WeatS is the rate of SiO₂ being weathered from silicate rocks and moved to the oceans.

Degas is the rate at which CO₂ is released to the atmosphere by volcanic activity

tCO2 is the total amount of CO₂ dissolved in the ocean, adding all of its forms:

$$tco2 = [CO_2] + [H_2CO_3] + [HCO_3 -] + [CO_3^{2-}].$$

alk is the ocean alkalinity: the total amount of acid (H^+) necessary to neutralize the carbonate and bicarbonate in the ocean. The detailed definition is complicated, but to a good approximation, alk = $[HCO_3-]+2[CO_3^{2-}]$. This is not crucial for this lab.

CO3 is the concentration of dissolved carbonate (CO_3^{2-}) in the ocean, in moles per cubic meter.

d13Cocn is the change in the fraction of the carbon-13 (¹³C) isotope, relative to the more common carbon-12 (¹²C) isotope, in the various forms of carbon dissolved in the ocean water.

d13Catm is the change in the fraction of ¹³C, relative to ¹²C in atmospheric CO₂.

Tatm is the average air temperature.

Tocn is the average temperature of ocean water.

Running the GEOCARB model from R

I have provided functions for running the GEOCARB model from R:

To run the model:

You need to specify co2_spike (the spike in CO₂ at time zero, measured in billions of tons of carbon).

The other parameters will take default values if you don't specify them, but you can override those defaults by giving the parameters a value.

The arguments to the function are:

- filename an optional file to save the results of the run to. You can read them back in using the read_geocarb() function: r run_geocarb(spike = 1000, filename =
 "test_run.txt") data = read_geocarb("test_run.txt")
- degas_spinup and degas_sim the rates of CO₂ degassing from volcanoes for the spinup and simulation phases, in trillions of molecules per year.
- plants_spinup and plants_sim TRUE or FALSE values for whether to include the role of plants in weathering (their roots speed up weathering by making soil more permeable and by releasing CO₂ into the soil), and land_area is the total area of dry land, relative to today.
- land_area_spinup and land_area_sim The amount of land area, compared to today (1.0 means the same amount of land as today).
- delta_t2x The climate sensitivity (the amount warming for each time CO₂ is doubled), in degrees Celsius.
- million_years_ago Simulate past climates when the sun was not as bright as today. The value of this variable is how many million years ago the year zero of the simulation should be. This is not currently working because of a bug in the web version of GEOCARB.

mean_latitude_continents The mean latitude, in degrees, of the continents.

The default values are: degas = 7.5, plants = TRUE, and land_area = 1 for both the spinup and the simulation.

The default value for delta_t2x is 3.0, million_years_ago is 0, and mean_latitude_continents is 30, which corresponds to today's world.

mean_latitude_continents and land_area allow you to explore conditions in earth's past, where the continents had different areas and were located in different parts of the world.

million_years_ago is meant to allow you to explore how the silicate weathering thermostat worked in earth's past, when the sun was a lot less intense than it is today. However, this part of the model is not working now.

After you run run_geocarb, you would read the data in with read_geocarb(filename). This function will return a data frame with the columns year, co2_total, co2_atmos, alkalinity_ocean, delta_13C_ocean, delta_13C_atmos, carbonate_ocean, carbonate_weathering, silicate_weathering, total_weathering, carbon_burial, degassing_rate, temp_atmos, and temp ocean.

Refresher on Plotting Several Variables

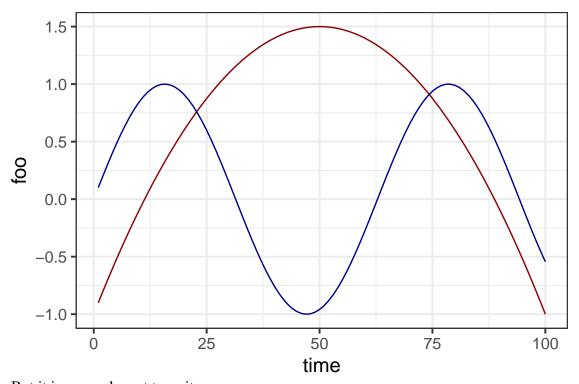
You may want to go back to the documentation for Lab #2 and refresh your memory about the pivot_longer() function for manipulating data frames and tibbles, and the different ways we can use ggplot to plot several variables on the same plot.

Suppose you have a tibble with columns time, foo and bar, as shown below:

time	foo	bar
1	-0.90	0.10
2	-0.80	0.20
3	-0.71	0.30
4	-0.62	0.39
5	-0.52	0.48
6	-0.44	0.56

Now, suppose you want to plot foo and bar on the same graph. You can do

```
ggplot(df, aes(x = time)) + geom_line(aes(y = foo), color = "darkred") +
  geom_line(aes(y = bar), color = "darkblue")
```

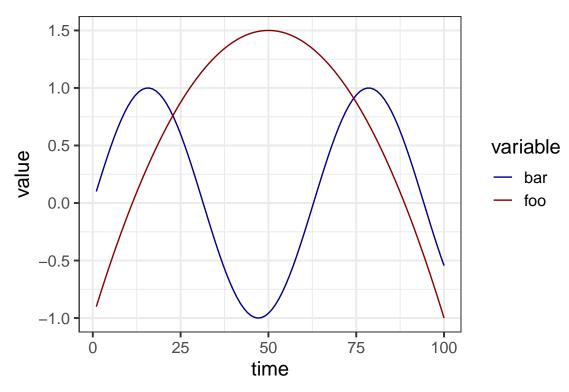


But it is more elegant to write

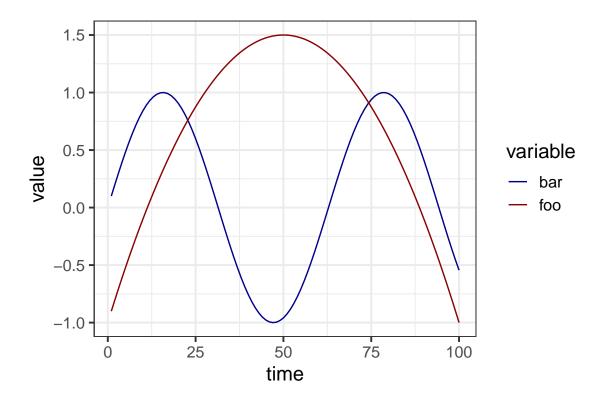
time	variable	value
1	foo	-0.9010000
1	bar	0.0998334
2	foo	-0.8040000

time	variable	value
2	bar	0.1986693
3	foo	-0.7090000
3	bar	0.2955202

Now you can plot this:



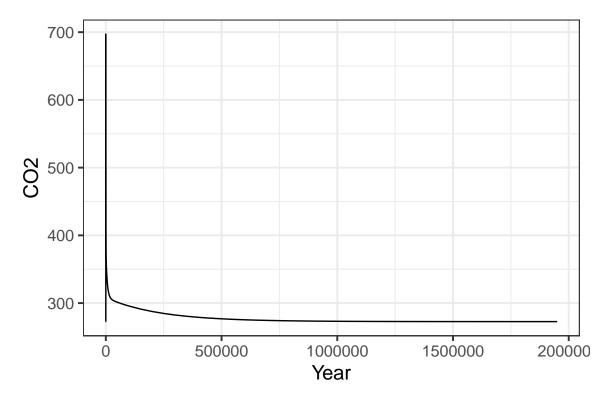
And you can put this all together in a single expression using pipes:



Modifying the Axes of a Plot

Sometimes you have a lot of data and you just want to plot a small part of it. Consider the following GEOCARB model run:

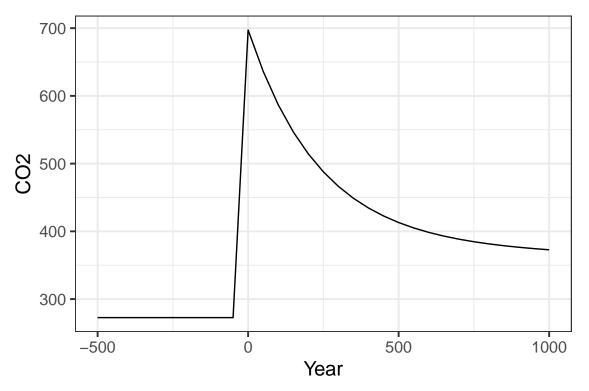
```
geocarb_data = run_geocarb(1000)
ggplot(geocarb_data, aes(x = year, y = co2_atmos)) + geom_line() +
labs(x = "Year", y = "CO2")
```



This shows us all 2 million years of the model run, but we can't see the detail of what's happening near year zero. There are several ways we can zoom our plot in to look only at the region near year zero:

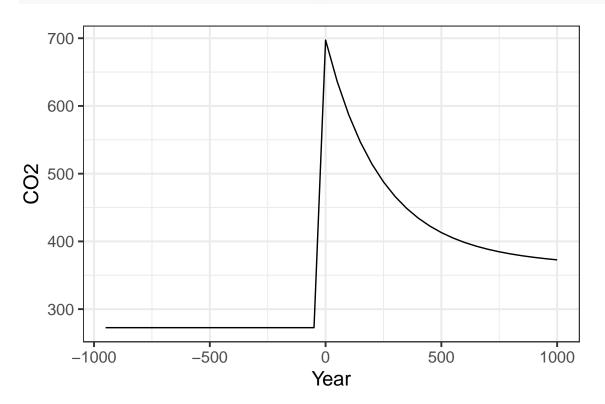
1. Use the xlim and ylim functions to set limits:

```
ggplot(geocarb_data, aes(x = year, y = co2_atmos)) + geom_line() +
    xlim(-500, 1000) + labs(x = "Year", y = "CO2")
```



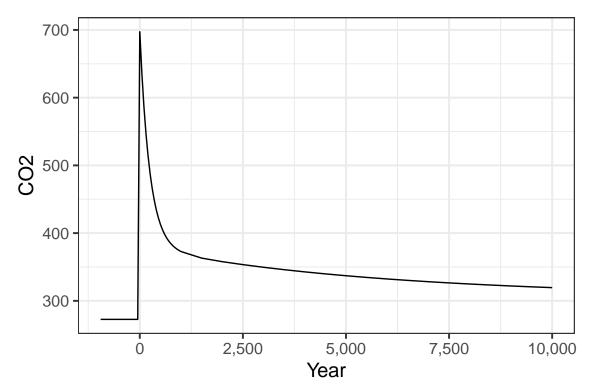
If you only want to change one limit, and leave the other at its default, you can put NA for the limit you want to leave alone:

```
ggplot(geocarb_data, aes(x = year, y = co2_atmos)) + geom_line() +
xlim(NA, 1000) + labs(x = "Year", y = "CO2")
```



2. Use the scale_x_continuous and scale_y_continuous functions

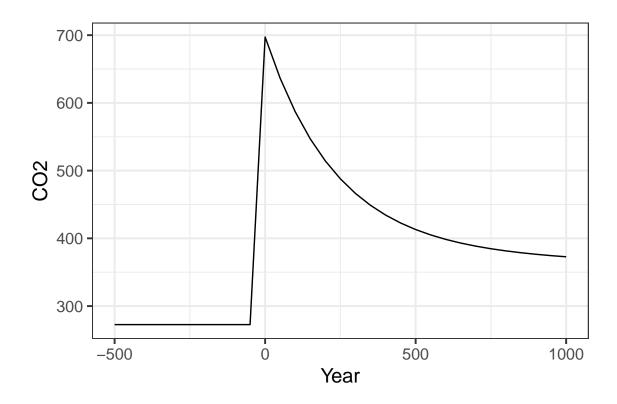
```
ggplot(geocarb_data, aes(x = year, y = co2_atmos)) + geom_line() +
    scale_x_continuous(limits = c(NA, 1E4), labels = label_comma()) +
    labs(x = "Year", y = "CO2")
```



Using the scale_x_continuous and scale_y_continuous functions lets you also modify the way numbers are formatted on the axis. Here, I used the label_comma() function to insert commas in the thousands and millions places. Other label commands include label_percent. You can read more about these at the web page for the scales package.

3. Another approach to limiting the extent of the plot is to filter the data before you call ggplot

```
geocarb_data %>% filter(year >= -500, year <= 1000) %>%
ggplot(aes(x = year, y = co2_atmos)) + geom_line() +
labs(x = "Year", y = "CO2")
```



Exercises for Lab #6

Exercise 1: Weathering as a Function of CO₂

In the steady state, the rate of weathering must balance the rate of CO₂ degassing from the Earth, from volcanoes and deep-sea vents.

Write this exercise up as a discussion of what happens when the rate of volcanic degassing changes. This rate has changed many times in Earth's history.

- Discuss how CO₂ and temperature change both in the first thousand years after the rate of degassing changes, and also in the longer term, over the course of one or two million years.
- Discuss what causes the amount of CO₂ in the atmosphere to stabilize after the degassing rate changes.
- Discuss how the size of the change in degassing rage affects the amount of change in CO₂
 and temperature between the original climate and where they finally stabilize with the new
 degassing rate.
- Explain the role of the silicate weathering thermostat in stabilizing the amount of CO₂, and what determines the final stable value of CO₂.

Details:

Run a simulation with co2_spike set to zero, and set the model to increase the degassing rate at time zero (i.e., set degas sim to a higher value than degas spinup). Leave degas sim at 7.5

and start out by setting degas sim = 10.

- a) Does an increase in CO₂ degassing drive atmospheric CO₂ up or down? How long does it take for CO₂ to stabilize after the degassing increases at time zero?
 - For the purposes of this exercise, consider that CO_2 has stabilized when the change in co2 atmos for a time-step in the model is less than 1 ppm.
 - Hint: Look back to the discussion of the lag() function in Lab #2. The expression co2_atmos lag(co2_atmos) will tell you the change in co2_atmos from the previous row to the current one in a tibble or data frame.
- b) Check that the model balances silicate weathering against CO₂ degassing when the CO₂ in the atmosphere stabilizes. Use ggplot to make a graph illustrating this balance. What is causing the silicate weathering rate to change?
 - **Hint:** This is a good place to use the pivot_longer function to make a plot with two different variables, as I described above.
- c) Repeat this run with a range of degassing values for the simulation phase and make a table or a graph of the equilibrium CO₂ concentration versus the degassing rate.
 - Does the weathering rate always balance the degassing rate when the CO₂ concentration stabilizes?
- d) Take the last row from each of the simulations you ran in part (c). This gives the values of all the variables 1.95 million years after the simulation began. Combine these into a single data frame, or tibble, and plot the silicate weathering rate versus the atmospheric CO₂ concentration. What does the relationship look like?
- e) Take the same data you used in part (d) and plot the silicate weathering rate versus the atmospheric temperature.
 - What does this relationship look like?

Exercise 2: The long-term fate of fossil fuel CO₂

Write this exercise up as a discussion of what happens if 2000 billion tons of carbon is released into the atmosphere as CO_2 . What do you learn from GEOCARB about where that CO_2 ends up and how the earth removes it from the atmosphere. Discuss how long the removal takes and what the implications are for how we should think about CO_2 in comparison to other kinds of pollution.

Details

Use the GEOCARB model in its default configuration.

- a) Run the model with no CO₂ spike at the transition. What happens to the weathering rates (Silicate, Carbonate, and Total) at the transition from spinup to simulation (i.e., year zero)? This is not a trick question. The answer should be obvious and simple.
- b) Now set the CO₂ spike at the transition to 2000 (2000 billion tons of carbon).
 - What happens to the weathering at the transition? How does weathering change over time after the transition?
 - How long does it take for CO₂ to roughly stabilize (stop changing)?
- c) In the experiment from (b), how do the rates of total weathering and carbonate burial change over time?
 - Plot what happens from shortly before the transition until 10,000 years afterward.
 Hint: See the discussion at the beginning of the lab instructions where I describe how to plot only a certain range of the data.
 - Now plot the carbon burial and total weathering for the range 1 million years to 2 million years. How do the two rates compare?

Exercise 3 (Graduate Students Only): How the Land Plants Changed the Carbon Cycle

The roots of plants accelerate weathering by two processes: First, as they grow, they open up the soil, making it more permeable to air and water. Second, the roots pump CO₂ down into the soil.

Write this exercise up as a report on the effects plants have on atmospheric CO₂ concentrations. If you turn off the plants during the spinup and then turn them on during the simulation, this simulates vascular land plants (plants with roots, stems, etc.) suddenly appearing in a world where they did not previously exist. How would this have affected the global carbon cycle and the composition of the atmosphere?

Details

Run a simulation with no CO₂ spike at the transition and with no plants in the spinup, but with plants present in the simulation.

- a) What happens to the rate of weathering when plants are introduced in year zero? Does it go up or down right after the transition? WHat happens later on?
- b) What happens to atmospheric CO_2 , and why?
- c) When the CO₂ concentration changes, where does the carbon go?