

# The Ultraviolet Environment of a Tropical Megacity in Transition: Mexico City 2000-2019

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## Abstract

Tropical regions experience naturally high levels of UV radiation, but urban pollution can reduce these levels substantially. We analyzed 20 years of measurements of the UV Index (UVI) at several ground-level locations in the Mexico City Metropolitan Area and compared these with UVI values estimated from satellite overpasses observing ozone and clouds (but not local pollution). The ground-based measurements were systematically lower than the satellite-based estimates, by ca. 40% in 2000 and 30% in 2019. Calculations with a radiative transfer model and observed concentrations of air pollutants explained well the difference between satellite- and ground-based UVI, and showed specific contributions from boundary layer and free tropospheric aerosols, O<sub>3</sub>, NO<sub>2</sub>, and SO<sub>2</sub>, in decreasing order of importance. Such large changes in UV radiation between 2000 and 2019 have important implications ranging from human health (skin cancer and cataract induction) to air pollution control (photochemical smog formation).

Keywords: Ultraviolet radiation, UV Index, Megacities, Air pollution, Urban photochemistry.

Synopsis: Improvements in Mexico City's air quality during 2000-2019 were accompanied by increases in UV radiation, augmenting human exposure and photochemical smog formation.

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## 14 Introduction

15 Ultraviolet (UV) radiation is an important component of the urban environment, affecting human populations  
16 directly through UV exposure of skin and eyes<sup>1;2;3</sup> and less directly (but with great impact) by driving the  
17 formation of photochemical smog, including tropospheric ozone and other oxidants, as well as secondary  
18 aerosols containing nitrates, sulfates, and organics.<sup>4;5;6</sup> These pollutants, along with others of primary origin  
19 commonly found in urban atmospheres (e.g., black carbon, sulfur dioxide), can in turn scatter and/or absorb  
20 UV radiation, alter its vertical distribution, and so modify the photochemical rate of their own formation.  
21 Such feedback complicates the calculation of both the UV radiation field (including at the surface), and the  
22 evolution of photochemical smog in the urban boundary layer.

23 The question of how air pollution alters the urban UV environment (and *vice versa*) is not new, but studies  
24 have relied mostly on numerical models,<sup>7;8;9</sup> with relatively fewer available observations (e.g., *McKenzie et*  
25 *al.*<sup>10</sup>; *Panicker et al.*<sup>11</sup>; *Palancar et al.*<sup>12</sup>; reviewed by Bais et al.<sup>13</sup>). Increases in UV have been estimated in  
26 association with decadal emission reductions, e.g. in China,<sup>14;15;16</sup> and that have led to less-than-expected  
27 reductions in photochemical smog, in part due to stronger UV photochemistry.<sup>17;18;19</sup> Emission reductions  
28 have also occurred globally during the 2020 COVID-19 pandemic,<sup>20;21</sup> but ground-level ozone in some pol-  
29 luted areas has actually increased,<sup>22;23</sup> due at least in part to the increased UV radiation. Unfortunately,  
30 the observational data base of relevant UV radiation remains rather sparse to evaluate such model-derived  
31 hypotheses.

32 The environment of Mexico City is of particular interest for several reasons: (1) Nearly 23 million people  
33 inhabit the Mexico City Metropolitan Area (MCMA), and the UV environment has direct implications for  
34 their health, both in terms of skin/eye UV exposure and *via* photochemical smog formation. (2) As a tropical  
35 megacity, it is to some extent representative of the situation of many others, with year-round intense midday  
36 UV irradiance, a shallower atmosphere due to the city's high elevation of 2240 m above sea level, and a  
37 transition toward newer and cleaner technologies, leading to gradual improvements in air quality. (3) Air  
38 quality within MCMA has undergone extensive scrutiny, with a well-established monitoring network since  
39 1986,<sup>24</sup> numerous intensive field campaigns to study the meteorology, emissions, and photochemistry of smog  
40 formation,<sup>25;26;27</sup> and numerical modeling incorporating the evolving knowledge.<sup>28;29;30;31</sup> This extensive  
41 body of knowledge provides the foundation for understanding our study.

42 Here, we analyze two decades of continuous measurements of the UV Index at multiple locations within the

MCMA, collected by the Secretariat of the Environment (Secretaría del Medio Ambiente, SEDEMA) of the Mexico City government as part of an intensive monitoring network over the MCMA.<sup>32</sup> The UV Index is defined as:

$$UVI = 40 \int_{250nm}^{400nm} E(\lambda, t) \cdot S_{er}(\lambda) d\lambda \quad (1)$$

where  $E(\lambda, t)$  is the solar spectral irradiance in units of  $W \cdot m^{-2} \cdot nm^{-1}$  and  $S_{er}(\lambda)$  is the erythral sensitivity of human skin.<sup>33;34</sup> Multiplication by 40 was chosen historically to express the UVI in small integer numbers, but is otherwise scientifically arbitrary.

The UVI is recognized by the World Health and Meteorological Organizations (WHO and WMO) as a standardized metric of UV radiation<sup>33</sup> for global public information. An advantage of using the UVI as (one) metric of UV radiation is that it is being increasingly observed or calculated and disseminated, enabling more objective comparisons among seasons and locations. The UVI observations from Mexico City, considered here, are an important element of this global picture.

While the UVI at the surface cannot be translated directly into photolysis frequencies for various photolabile molecules, the spectral weighting of the UVI (ca. 300-320 nm) is approximately similar to that for the photolysis of ozone to singlet oxygen atoms. Other UV wavelengths are of course also important, e.g. for the photolysis of nitrogen dioxide, and may be affected differently depending on the pollutant. With these considerations and a few other caveats, UVI trends examined here can also be used to infer accompanying trends in photolysis frequencies and influences on photochemical smog formation.

## Methods

### Ground-based measurements

The Mexico City Metropolitan Area is located at 19.4°N, 99.1°W, 2240 meters above sea level (asl), surrounded by mountain ridges exceeding 5000 m asl, with complex topography and thermal inversions that inhibit winds and favor intense air pollution.<sup>35;36;37</sup> Air quality monitoring and surface meteorological measurements in the MCMA are conducted continuously by the Automated Atmospheric Monitoring Network (RAMA, by its Spanish acronym) of the Atmospheric Monitoring System (SIMAT, by its Spanish acronym) of the Mexico City government. Since the year 2000, UV radiometers (model 501-A, Solar Light Company Inc., Glenside, PA) detecting wavelengths between 280-400 nm have been measuring erythemally-weighted

69 solar radiation. The calibration of the UV sensors was carried out annually by comparing against a factory-  
 70 calibrated reference sensor. The output voltages from the measuring sensors were compared during at least  
 71 one week against the UV readings from the reference sensor to derive the calibration factor. New calibration  
 72 factors typically differed from the old ones by 2% or less. Reference sensors were also calibrated by the  
 73 manufacturer and updated periodically (between 3 to 5 years) to avoid any bias due to aging. Long term  
 74 calibration drift was avoided by the yearly re-calibrations. Although at the beginning only a few stations  
 75 were in operation and have been changing, currently 11 stations are recording erythemal irradiances, which  
 76 are then multiplied by 40 (see Eq. 1) to give UV Indices. Table 1 describes the location of the stations where  
 77 UV Index has been measured. Figure 1 shows how the radiometers of the SIMAT have been distributed  
 78 over MCMA, prioritizing the sites with more density of population. Near real-time data for each station  
 79 are available on the SIMAT official website <http://www.aire.cdmx.gob.mx/default.php>. Daily maximum  
 80 values ( $UVI_{\max}$ ) were extracted from each of the stations around solar noon from the time interval from  
 81 11:00 h-15:00 h CST (Central Standard Time). This database with 7305 continuous days of measurements  
 82 during the period 2000-2019 was analyzed.

Station	Environment	Lat (°N)	Lon (°W)	El (masl)
CHO	semi-urban zone	19.27	99.89	2253
CUT	ecological park	19.72	99.20	2263
FAC	urban	19.48	99.24	2299
HAN	urban	19.42	99.08	2235
LAA	urban	19.48	99.15	2255
MER	downtown	19.42	99.12	2245
MON	rural	19.46	98.90	2252
MPA	rural	19.18	98.99	2594
PED	residential	19.33	99.20	2326
SAG	urban	19.53	99.03	2241
SFE	residential	19.36	99.26	2599
TLA	urban	19.53	99.20	2311
UNAM	University city	19.33	99.18	2294

Table 1: SIMAT stations, environmental descriptors and geographical positions. Abbreviations names: Chalco (CHA), Cuautitlán (CUT), FES Acatlán (FAC), Hangares (HAN), Laboratorio de Análisis Ambiental (LAA), Merced (MER), Montecillo (MON), Milpa Alta (MPA), Pedregal (PED), San Agustín (SAG), Santa Fe (SFE), National Autonomous University of Mexico (UNAM) and Tlalnepantla (TLA).

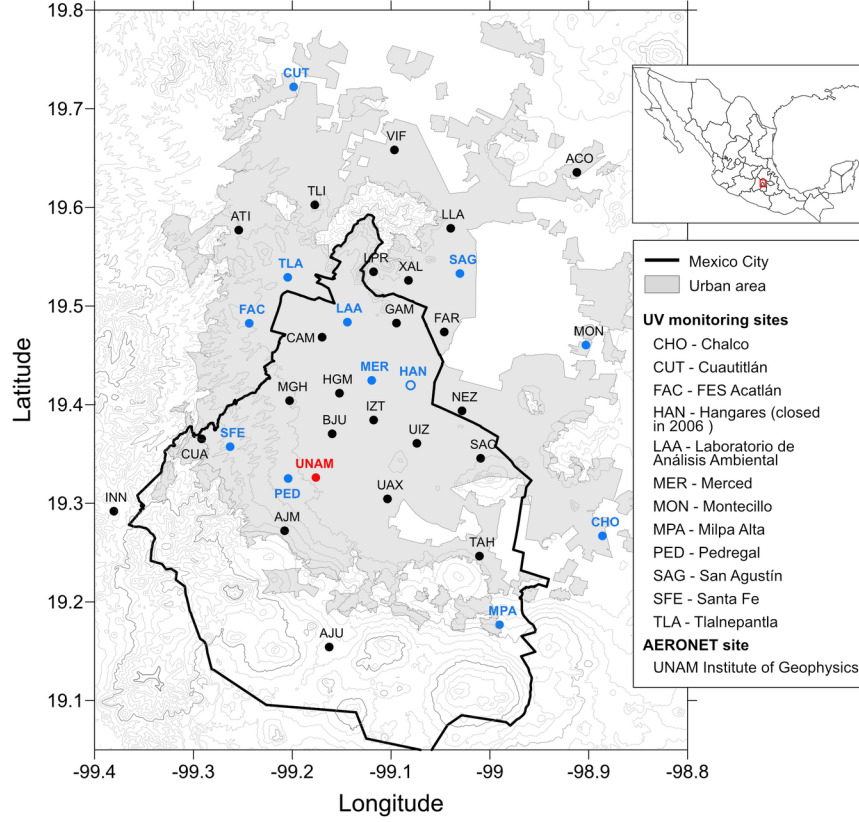


Figure 1: Map with the location of the SIMAT continuous monitoring stations over MCMA. Sites denoted by the blue solid dots correspond to SIMAT stations with UV measurements, while the black dots indicate SIMAT stations without UV measurements, the site on open blue dot represents a discontinued site, the red dot corresponds to the location of the AERONET site. The location of Mexico City and the acronyms of the UV and AERONET site are shown at the upper and lower frames at the right respectively.

With the aim to explore the relationship between UVI and air pollutants levels, the hourly averages for ozone ( $O_3$ ), carbon monoxide (CO), nitrogen dioxide ( $NO_2$ ), sulfur dioxide ( $SO_2$ ) and particle matter with diameter sizes  $\leq 10 \mu m$  ( $PM_{10}$ ), were downloaded from the SIMAT<sup>38</sup>. For purposes of assessing the influence on the UV Index at solar noon, only values obtained between 11h and 15h CST were considered for the trends analysis. Pollutant measurements are conducted by the SIMAT using regulatory-grade commercial instruments. Measurement principles include ultraviolet photometry (model 400E, Teledyne-API) for  $O_3$ , chemiluminescence (model 200E, Teledyne-API) for  $NO_2$ , UV fluorescence (model 100E, Teledyne API) for  $SO_2$ , and infrared absorption (model 300E, Teledyne-API) for CO. The  $PM_{10}$  continuous mass concentration was measured with Tapered Element Oscillating Microbalance (TEOM 1400AB or TEOM 1405 DF, Thermo

Scientific) monitors. Gaseous pollutant levels are reported in ppb concentration units for O<sub>3</sub>, SO<sub>2</sub> and NO<sub>2</sub>, and in ppm for CO. Particulate matter mass concentration is reported in  $\mu\text{g m}^{-3}$  at local conditions for temperature and pressure.

Aerosol optical depth at 340 nm was obtained from the Institute of Geophysics of the National Autonomous University of Mexico (UNAM); measurements were conducted with a CIMEL sun photometer model CE-318 of the AErosol RObotic NETwork (AERONET<sup>39</sup>). The data Product Level 2.0 and 1.5 (only in 2019) were selected, and annual averages AOD<sub>340</sub> were calculated from continuous measurements during at least 7 months. A previous study of the AOD behavior from 2000 to 2014 demonstrated that the gaps of data did not significantly change the trends over the analyzed period.<sup>40</sup> From 2014 to 2019 there was only 8% of the missing data, when the instrument was sent for calibration in 2018.

## Satellite data

Estimates of the UV Index from satellite-based measurements of clouds and O<sub>3</sub> were used for comparing to the ground-based measurements. These data were provided by the Ozone Monitoring Instrument (OMI) on board of AURA-NASA satellite.<sup>41</sup> OMI was created in cooperation between the Netherlands Agency for Aerospace Programmes (NIVR), the Finnish Meteorological Institute (FMI) and NASA. OMI (hereafter OMI-Aura/NIVR-FMI-NASA) performs observations over a geographical dimension of  $13 \times 24\text{km}^2$  at nadir. For Mexico City, the satellite overpass time is between 19:00h - 21:00h UTC and data are specific for the coordinates and elevation of Mexico City. Measurements of the ozone profile and cloud cover are used via a radiative transfer model to estimate the UVI at the ground. The OMI web site reports UVI values for both the overpass time, and corrected to local solar noon.

While the early OMI estimates of the UVI did not consider aerosols in the boundary layer, Arola et al.<sup>42</sup> suggested correcting the clean-skies UV index with a reduction factor based on a monthly global climatology of BL aerosols. For Mexico City a reduction of about 8% is applied, starting ca. 2013, to the OMI UVI data.

## TUV model

Calculations of the UV Index were also made with the Tropospheric Ultraviolet Visible (TUV v5.3) model.<sup>43</sup> The model atmosphere was represented by 80 vertical layers, each 1 km thick, starting at the 2.24 km asl elevation of MCMA, and for which the first three km constitute the atmospheric boundary layer (BL).<sup>44</sup> The ozone profile above the BL is from the US Standard Atmosphere, but rescaled to a value of 259.6 DU. Totals including the BL contributions (of 13.7 DU in the year 2000 and 9.8 DU in 2019, see below) were

121 273.1 DU in 2000 and 269.2 DU. The climatological  $O_3$  column for this latitude and season is about 270 DU  
122 (plus or minus ca. 5 DU) so in good agreement with the values used here.

123 Pollutants within the BL (including  $O_3$ ) are assumed to be well mixed, in agreement with observations from  
124 the MILAGRO field campaign that showed the disappearance of vertical gradients in the profiles of gases<sup>45;46</sup>  
125 and aerosols<sup>47;48</sup> by late morning. The UV-absorbing gases considered here are  $O_3$ ,  $NO_2$ , and  $SO_2$ , specified  
126 in ppb.

127 BL aerosols are modeled by prescribing the AOD at 340 nm (from AERONET observations), scaled to other  
128 wavelengths inversely with wavelength (Angstrom coefficient = 1.0), asymmetry factor of 0.7, and a single  
129 scattering albedo of 0.85 at UV wavelengths, following the determinations made in Mexico City by Corr et  
130 al.<sup>49</sup> and Palancar et al.<sup>12</sup>.

131 Above the atmospheric boundary layer, the model was taken to be free of aerosols,  $NO_2$ , or  $SO_2$ . In one  
132 sensitivity study, a total AOD of 0.7 was redistributed placing 0.2 in the free troposphere (decreasing verti-  
133 cally with an exponential scale height of 4 km, and 0.5 remaining in the BL). The calculated UVI differed  
134 by less than 1%. Thus, as long as the total AOD is known, knowledge of the exact vertical aerosol profile  
135 is not critical towards ground-level UVI – but would obviously affect the vertical structure of photolysis  
136 frequencies.

137 Radiative transfer calculations were carried out with the pseudo-spherical 4-stream option, at 1 nm steps  
138 between 280 and 400 nm.

## 139 Results and Discussion

140 Figure 2 shows the diurnal variation of the UVI for several specific cloud-free days, for different seasons and  
141 several locations (CHO, MER, MON, PED, SAG, SFE and TLA). UV Index from TUV model was used  
142 as reference of the behavior under clear sky days. According to the comparison of measurements minute  
143 by minute, the dates with prolonged fluctuations along the day were discarded. However, during the rainy  
144 period (from June to October) at least a brief clouds presence is common, as shown at CHO station around  
145 noon in 13 June 2017. Peak values range from 8 during autumn/winter to above 12 in spring/summer, in  
146 correspondence to the respective December and June solstices. Although the stations are all within a 25 km  
147 radius, substantial differences among them are notable. Survey of the locations revealed that shadowing from  
148 nearby structures is not an issue. The good agreement in the morning, followed by more divergence in the



149 afternoon, is consistent with the development of photo-chemical pollution hotspots during the day. Previous  
 150 studies (e.g., Castro et al.<sup>50</sup> and Palancar et al.<sup>12</sup>) have shown that surface UV radiation in Mexico City is  
 151 attenuated significantly by aerosols. The measurements shown in Fig. 2 are consistent with this increasing  
 152 pollution during the course of the day, with highest aerosol loading (and highest variability) attained in the  
 153 afternoon. Further support for the role of pollution in suppressing the UVI comes from the observation made  
 154 at the Santa Fe (SFE) site which in Fig. 2 are seen to be systematically higher, e.g. by over 10% in autumn  
 155 afternoons, compared to the other stations. The SFE station is displaced to the west from the majority of the  
 156 other stations, and remains in the outskirts. This site is also approximately 300 m higher than Mexico City  
 157 downtown, so that the cleaner atmosphere results from both vertical and horizontal variations in pollution  
 158 levels.<sup>51</sup> It is indeed expected to have higher values of the UVI, in agreement with the observations.

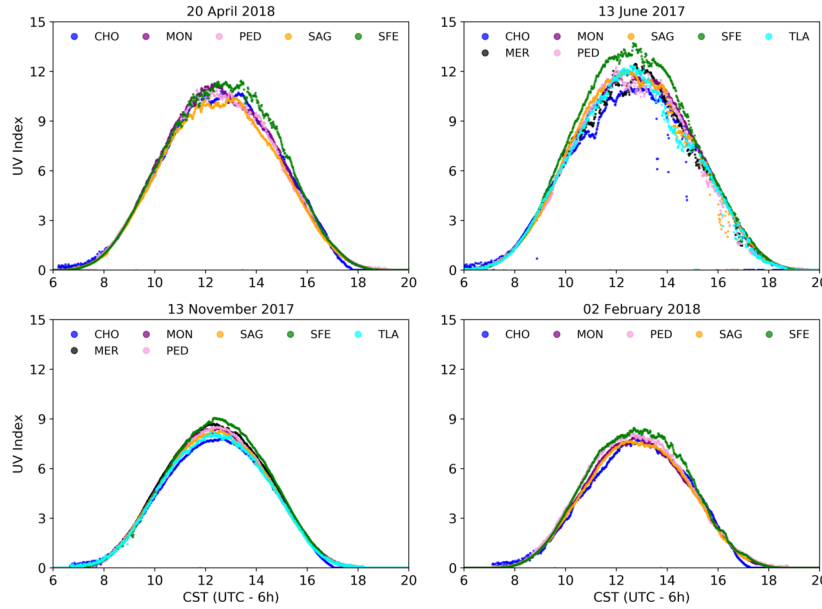


Figure 2: UV Index measured over MCMA by SIMAT stations each minute along the day under practically cloudless conditions for representative days of the year.

159 The daily maximum UV Index of each station is denoted by  $UVI_{\max}$ , all of them were counted in the  
 160 period 2000-2019. As shown in Figure 3, these values ranged from 1 to 16, with a majority (61%) of the  
 161 days experienced  $UVI_{\max}$  values between 6 and 10, and remarkably few, less than 1%, in the higher 13-16  
 162 range. The lowest values are likely due to winter days with heavy cloud cover and low sun angles, and UV  
 163 attenuation by pollutants could also be amplified under such conditions, due to longer photon path lengths

at low sun and within clouds. The extreme sparsity of high UVI values remains surprising, and may be an indication of the rarity of extremely unpolluted days within the city.

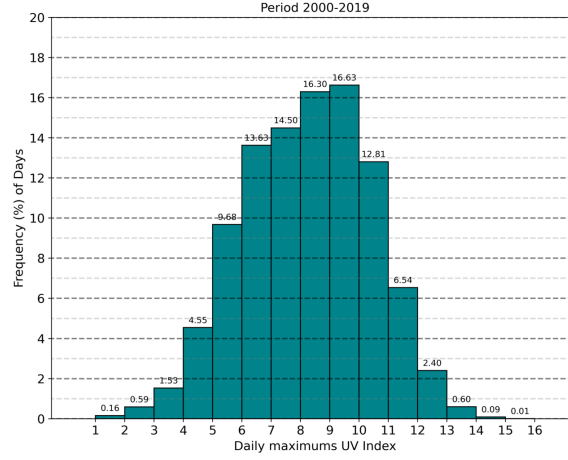


Figure 3: Frequency distribution of daily maximums UV Index values in Mexico City during 2000 -2019.

Similar patterns are found when considering the monthly average of the  $UVI_{\max}$  values as shown in Figure 4. The long-term averages present a seasonal variation (as in Fig. 2) that follows approximately the cosine of the noontime solar zenith angle. Notably, values rarely if ever exceed 12 (as in Fig. 3). The lowest average UVI (near 7) take place in winter while from March to August the values seem to be flattened in the range 10-11. The rather low monthly UV Index values, mainly could be a consequence of the presence of clouds in the rainy season. However, urban aerosol pollution sources, biomass burning for agriculture and wood cooking also contribute to poor air quality between March-May<sup>52</sup>. Using the maximum  $UVI_{\max}$  from all of stations every day (one daily point), the monthly averages ( $\overline{UVI}_m$ ) were calculated along the period 2000-2019. Long term trends in  $\overline{UVI}_m$  are shown in Figure 5. A clear upward trend is seen, with a slope for the linear fit of 0.9%/year or +1.5 UVI units over the two decades.

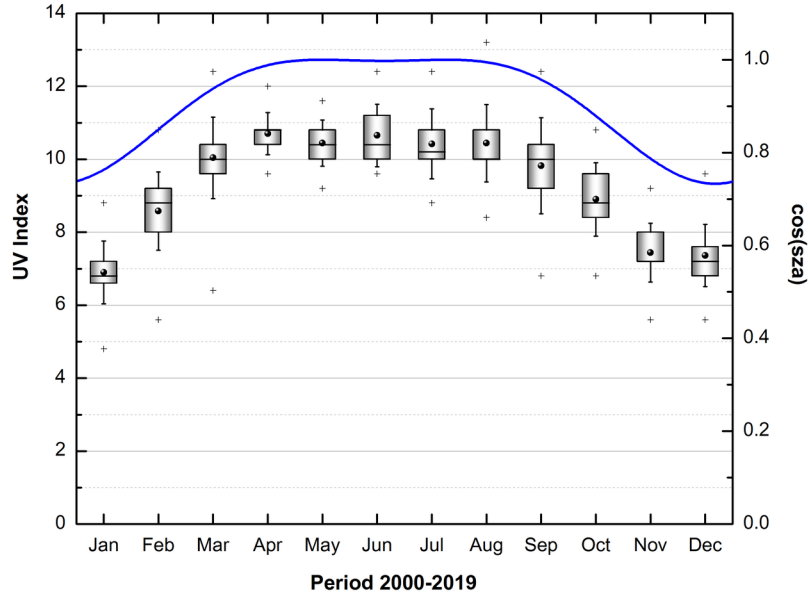


Figure 4: Boxplot of the monthly averages of the maximum UV Index values (black dot) in MCMA for the period 2000-2019: median (central bold line), 25<sup>th</sup> and 75<sup>th</sup> percentiles (box edges), standard deviation (the whiskers), the minimum and maximum values (plus sign) and the cosine of the solar zenith angle at solar noon (blue curve).

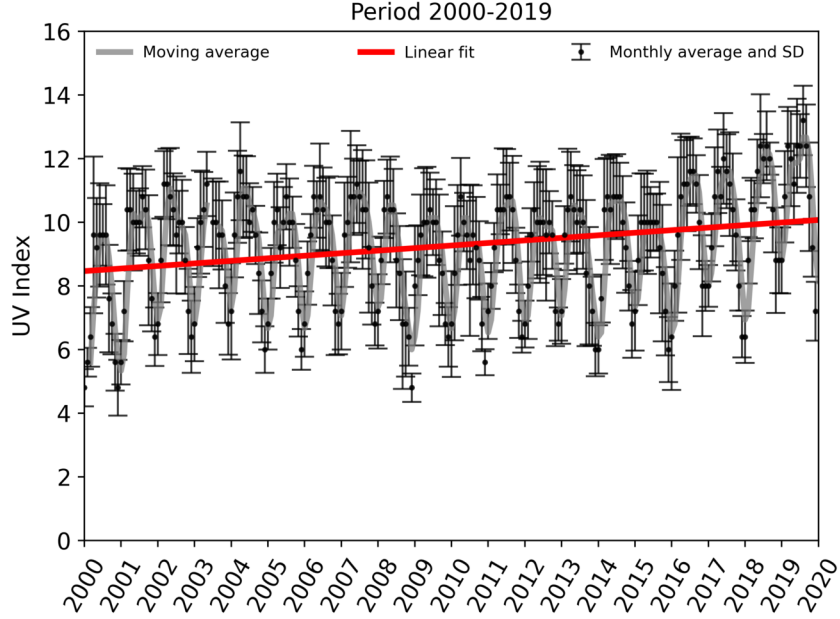


Figure 5: Moving average function (gray curve) quarterly applied to monthly average UV Index, standard deviation of all data within a given month (black dots and dash line) and linear fit (red line).

The UV Index computed from satellite-based observations (OMI-Aura/NIVR-FMI-NASA) over the period 2005-2019 is mapped in Figure 6. The satellite-derived UVIs vary from 8 in winter to 16 in summer, both values being substantially higher than the ground-based observations (ca. 7 for winter and 11 for summer, see Fig. 4). We hypothesize that this large difference between satellite-based estimation and ground-based observation of the UV index is due to the intense air pollution of Mexico City. A rather similar behavior was detected in Santiago city, Chile.<sup>53</sup>

Close inspection of Figure 6 shows that the maximum values, those from June and July of each year, show a slight but systematic decrease starting ca. 2013. As mentioned in section , this is due to the post-processing of OMI UVI data to account for absorption by BL aerosols (Arola et al.)<sup>42</sup>, using a climatological reduction of about 8% for Mexico City.

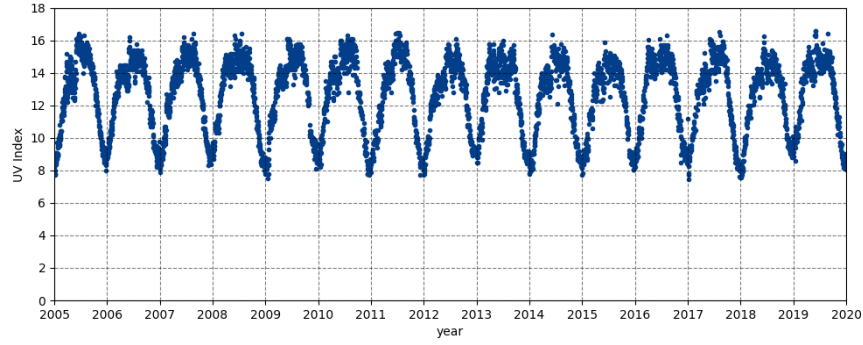


Figure 6: UV Index at solar noon for clear sky recorded by OMI-Aura/NIVR-FMI-NASA, from 2005 to 2019.

## Effect of pollutants on UV radiation

Trends and averages in aerosol optical depth  $AOD_{340}$  and criteria pollutants  $PM_{10}$ ,  $CO$ ,  $NO_2$ ,  $O_3$  and  $SO_2$  observed at the SIMAT stations over 2000-2019, are shown in Figure 7 and summarized in Table 2 together with the  $UVI_{max}$ . Similar trends in pollutants have been noted before<sup>54;55;56</sup> and reflect the long-term success of emission reduction policies and programs.

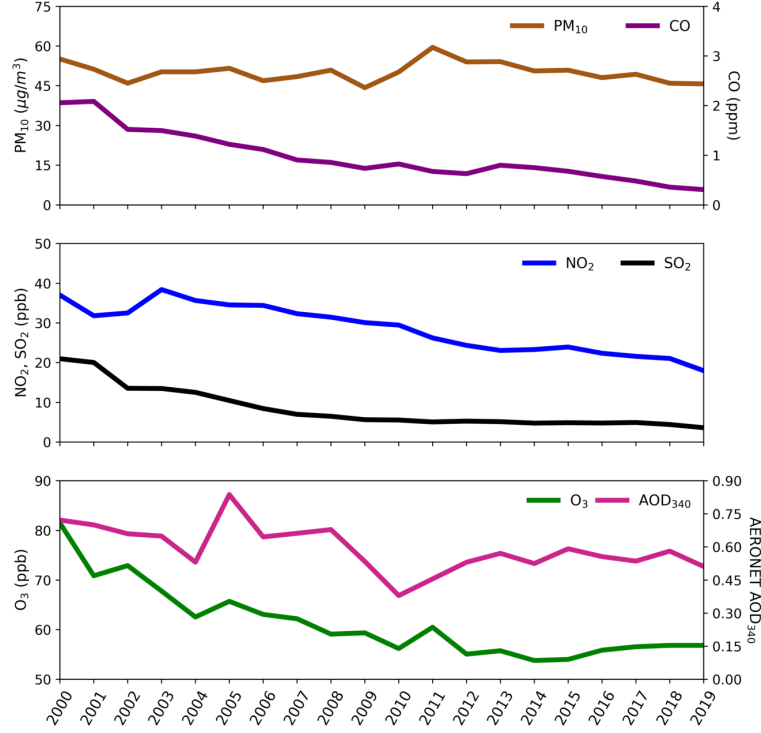


Figure 7: Air quality trends in MCMA for the period 2000-2019 from annual averages obtained between 11h to 15h CST every day: PM<sub>10</sub> (brown curve), CO (purple curve), NO<sub>2</sub> (blue curve), SO<sub>2</sub> (black curve), O<sub>3</sub> (green curve) and AOD<sub>340</sub> (pink curve).

Variable	$\frac{\Delta \text{variable}}{\Delta t}$	Avg <sub>2000-2019</sub>	$\Delta(\%/year)$
UVI	0.08	9.2	0.9
PM <sub>10</sub>	-0.10	50.2	-0.2
CO	-0.08	1.0	-8.2
NO <sub>2</sub>	-0.96	28.6	-3.4
O <sub>3</sub>	-1.05	61.3	-1.7
AOD <sub>340</sub>	-0.01	0.6	-1.6
SO <sub>2</sub>	-0.76	8.4	-9.1

Table 2: UV Index and criteria pollutants: slope for the period 2000-2019, averages in units of  $\mu\text{g}/\text{m}^3$  (PM<sub>10</sub>), ppm (CO), ppb (SO<sub>2</sub>, NO<sub>2</sub> and O<sub>3</sub>), dimensionless (UV Index and AOD<sub>340</sub>) and annual percentage change (%/year).

191 The observed changes in the concentrations of these air pollutants have significant implications for surface  
 192 UV radiation, as can be demonstrated with the TUV radiative transfer model. Table 3 summarizes UVI  
 193 values for the June-July time period, estimated by the three methods: OMI satellite-derived UVI, RAMA  
 194 ground-based observations, and TUV modeling using air pollution estimates. Two groups of values can be  
 195 readily identified: (1) RAMA daily record values, OMI with or without BL aerosols, and TUV for very clean  
 196 conditions, all with UVI values around 15-16; and (2) RAMA average daily maxima and TUV UVI using  
 197 MCMA pollutants as input, which are in good agreement for both 2018/19 and 2000/01 but much lower  
 198 than the first group. Compared to the OMI estimate that included the climatological aerosol correction  
 199 (15.3), observed RAMA values were lower by 35% in 2000 and by 20% in 2019. Similarly, TUV values were  
 200 35% lower in 2000 and 22% lower in 2019 relative to the TUV values of 15.6 for a pristine atmosphere.

Conditions for estimation in June-July	UV Index
RAMA maximum reached in period 2000-2019	15.0
RAMA average maxima 2018, 2019	12.3
RAMA average maxima 2000, 2001	9.9
OMI clear, local noon, no corrected for BL aerosols	16.6
OMI with 8% reduction for BL aerosol (Figure 1 from Arola et al.)	15.3
TUV "zero" pollution AOD 0, O <sub>3</sub> 0 ppb, NO <sub>2</sub> 0 ppb, SO <sub>2</sub> 0 ppb	16.1
TUV pristine pollution AOD 0.05, O <sub>3</sub> 10 ppb, NO <sub>2</sub> 0 ppb, SO <sub>2</sub> 0 ppb	15.6
TUV 2019: AOD 0.5, O <sub>3</sub> 50 ppb, NO <sub>2</sub> 20 ppb, SO <sub>2</sub> 1 ppb	12.1
TUV 2000: AOD 0.7, O <sub>3</sub> 70 ppb, NO <sub>2</sub> 40 ppb, SO <sub>2</sub> 20 ppb	10.2

Table 3: UV Index estimates for June-July. TUV model calculation illustrating the effect of pollutants on  
 the UV Index in Mexico City. The contribution of each pollutant is evaluated by subtraction from the full  
 mixture, rather than individual additions to clean air, to represent more faithfully the interaction between  
 scattering and absorption, e.g. scattering-induced changes in photon path lengths through absorbing gases.

201 The agreement between RAMA observations and TUV model estimates is excellent but also probably a bit

fortuitous. Clouds on average reduce the irradiance impinging on the surface, but scattering from them can also cause transient enhancements (especially if the direct sunbeam is not blocked) that could be recorded as daily maxima – with cancellation between these cloud effects resulting in improved agreement with the cloud-free model. We cannot exclude that some of the observed trend in UVI is due to changes in cloud cover. However, the modeled fractional UVI reductions due to pollutants, shown in Table 3, are in such good agreement with the observed UVI reductions, that a compelling case can be made for a dominant role of air pollutants in the long-term UVI trends.

Table 4 shows the contributions to UVI reductions from individual pollutants. Aerosols are seen to be the major factor in both time periods, followed by  $O_3$ ,  $NO_2$ , and  $SO_2$ . The 2000-2019 UVI increase is seen to result in comparable proportions from fewer aerosols, less  $SO_2$ , and the combined reductions in  $O_3$  and  $NO_2$ . Comparable UV reductions, of 30-40% due to aerosols, were reported by Panicker et al.<sup>11</sup> over Pune, India from April 2004 to March 2005, with sensitivity coefficients (i.e. change in UVI per unit change in AOD) similar to those found here in Table 3.

Comparisons of OMI with ground-based UV measurements have been reviewed recently by Zhang et al.<sup>57</sup> and Vitt et al.<sup>58</sup>. OMI-derived UV generally overestimates ground-level measurements by 1-10% in relatively clean conditions (e.g. rural U.S), by 10-30% in Southern Europe and by 40% or more in Santiago, Chile<sup>53</sup> and Thailand<sup>59</sup>. The overestimations appear related in large part to incomplete accounting of UV absorption by BL aerosol, although other factors such as the correction to solar noon may also introduce some bias.<sup>57</sup> Over Europe, ground-based UVI observations for several decades are systematically lower than those estimated from satellites even after consideration of climatological aerosol distributions, showing the importance of local pollution not resolved from space.<sup>58</sup> However, difference between satellite-derived and ground-based UVI was less than 1.0 UVI units in over 90% of the cases, in contrast to the difference of 3-5 units found for Mexico City (Table 3).

UV reductions by air pollutants are expected to be most severe near the surface, while chemical reactions leading to photochemical smog occur through the vertical extent of the BL. Figure 8 shows the vertical profiles of photolysis coefficients (the reciprocals of photolytic lifetimes) for two key reactions, the photolysis of  $O_3$  to yield excited oxygen atoms  $O(1D)$ , and the photolysis of  $NO_2$ . In the absence of optically active pollutants, these coefficients would be nearly independent of altitude in the BL. However, the presence of pollutants leads to a strong decrease toward the surface, with notably more severe reductions in 2000 compared to 2019. While the specific values shown in the figure are only illustrative for typical conditions,



232 routine daily air quality modeling should carefully account for the long term variations in these photolysis  
 233 coefficients.

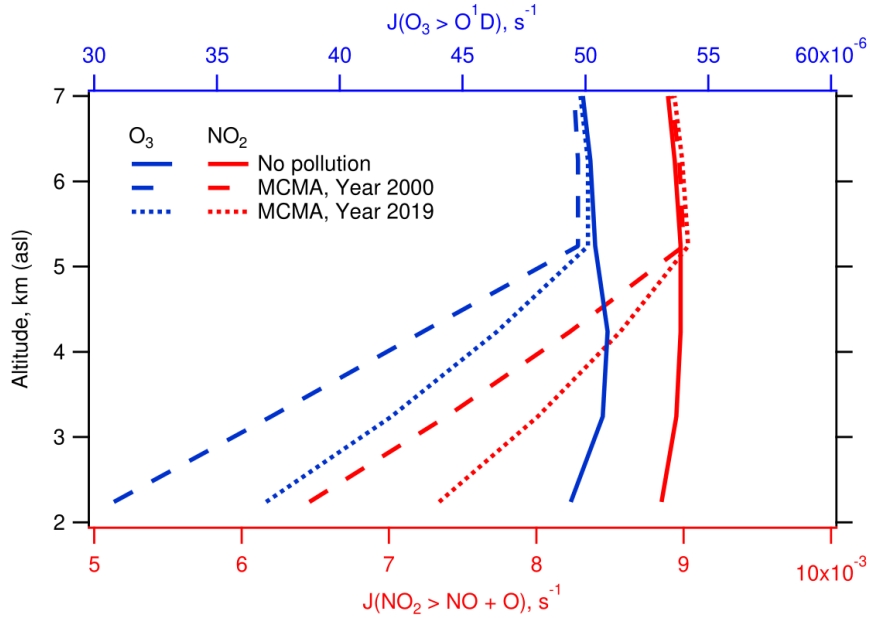


Figure 8: Vertical profiles of the photolysis coefficients for the reaction  $O_3 + h\nu \rightarrow O(^1D) + O_2$  (top axis, blue)  $NO_2 + h\nu \rightarrow O + NO$  (bottom axis, red) for zero pollution (solid) and Mexico City in the years 2000 (dashed) and 2019 (dotted)

234 An issue that is beginning to gain relevance in radiative balance models is the influence of a group of organic  
 235 compounds capable of strongly absorbing in the UV region (brown carbon).<sup>60</sup> Some of these compounds are  
 236 related with emissions from local and regional wildfires.<sup>61</sup> Mexico City is frequently exposed to regional fire  
 237 smoke transport during the dry part of the year (November to May)<sup>62</sup>, that sporadically modify the optical  
 238 properties of the aerosols.<sup>63</sup> This could partly explain the relatively minor reductions in  $PM_{10}$  (see Fig. 7),  
 239 compared to the larger reductions of  $CO$ ,  $NO_2$ , and  $O_3$  that are more directly related to urban activities, as  
 240 well as some of the seasonal asymmetry seen in Fig. 3.

241 The UVI is specific to wavelengths mainly in the 300-320 nm range, and so the question remains whether  
 242 these results can be applied at longer UV wavelengths, e.g. those important for  $NO_2$  photolysis ( $<420$  nm).  
 243 Absorption by  $SO_2$  and  $O_3$  vanishes, while absorption by  $NO_2$  increases and typical aerosols optical depth  
 244 decrease. These changes can easily be modeled, but unfortunately far fewer measurements of these longer  
 245 wavelengths are available in Mexico City or elsewhere.

Two decades of observations in Mexico City demonstrate unequivocally that air pollution reduces UV radiation at the ground. The ground-based observations are well below estimates derived from satellite-based observations, and below model calculations do not consider optically important pollutant aerosols, tropospheric ozone, and to a lesser extent NO<sub>2</sub> and SO<sub>2</sub>. When typical observed values of these pollutants are included in a model (e.g. TUV), the differences between satellite-derived and ground-based measured values are explained and can be attributed quantitatively to individual observed pollutants. Long term improvements in air quality, over two decades, are accompanied by statistically significant increases in the observed UVI, again in good agreement with the model-predicted changes.

Pollutant	Year 2000		Year 2019		2019-2000
	poll. level (a)	UVI (b)	poll. level	UVI	(c)
AOD	0.7	-3.7	0.5	-2.7	1.0
O <sub>3</sub> ppb (DU)	70 (14)	-1.4	50 (10)	-1.0	0.4
NO <sub>2</sub> ppb	40	-0.9	20	-0.5	0.4
SO <sub>2</sub> ppb	30	-0.8	1	-0.04	0.8

Table 4: (a) Values used one at the time, with the others held at zero. (b) UVI deviation from the zero-pollution value of 16.1 (from Table 3). (c) 2019-2000 UVI change due to changes in each pollutant.

The reductions in surface UV radiation with respect to an ideally clear atmosphere – by nearly 40% in 2000 and still 20% in 2020 – are large, both within the context of human UV exposure and air quality mitigation. In urban areas where ozone production scales proportionally with UV levels and with volatile organic compound (VOC) emissions (the VOC-limited regime), a 10% increase in BL average photolysis rates means that VOC emissions will need to be reduced by 10% to meet the same goals, or else successful reductions in aerosols would lead to unwanted UV-driven increases in O<sub>3</sub>. Such UV changes must be considered carefully in air quality mitigation strategies. For human exposure, a 20% increase in UV irradiances over two decades should be seen as a non-negligible public health issue requiring some reassessment of preventive behaviors to minimize the risk of skin cancer, cataract, and other UV-related health effects. The efforts that Mexico City has made to improve air quality have achieved positive results in the levels of most air pollutants. Nevertheless, they caused an increase in UV radiation that reaches the surface.

A limitation of the present work is our focus on daily maximum values, which largely exclude cloud cover. Absorption within clouds can be enhanced by the long path lengths of multiply scattered photons (e.g.,

Mayer et al.<sup>64</sup>), so that accurate quantification of UV effects of clouds in polluted environments remains a significant challenge and interesting opportunity for future work.

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