The Ultraviolet Radiation Environment of a Tropical Megacity in Transition: Mexico City 2000-2019

Adriana Ipiña,*^{*,†,‡} Gamaliel López-Padilla,[¶] Armando Retama,[§] Rubén D.

Piacentini,[†] and Sasha Madronich^{||}

†Instituto de Física Rosario (CONICET-UNR), Rosario, Argentina ‡Centro de Ciencias de la Atmósfera, Universidad Nacional Autónoma de México, Mexico City, Mexico

¶Facultad de Ciencias Físico Matemáticas, Universidad Autónoma de Nuevo León, San Nicolás de los Garza, México

 $\S Independent\ researcher,\ Mexico\ City,\ Mexico$ $\|National\ Center\ for\ Atmospheric\ Research,\ Boulder,\ Colorado,\ USA$

E-mail: ipina@ifir-conicet.gov.ar

Abstract

2

8

Tropical regions experience naturally high levels of UV radiation, but urban pollution can reduce these levels substantially. We analyzed 20 years of measurements of the UV Index (UVI) at several ground-level locations in the Mexico City Metropolitan Area and compared these data with UVI values derived from satellite observations of ozone and clouds (but not local pollution). The ground-based measurements were systematically lower than the satellite-based estimates, by ca. 35% in 2000 and 20% in 2019. Calculations with a radiative transfer model and observed concentrations of

air pollutants explained well the difference between satellite- and ground-based UVI, and showed specific contributions from aerosols in the boundary layer and the free troposphere, O₃, NO₂, and SO₂, in decreasing order of importance. Such large changes in UV radiation between 2000 and 2019 have important implications ranging from human health (skin cancer and cataract induction) to air pollution control (photochemical smog formation).

5 Introduction

10

11

12

13

14

Ultraviolet (UV) radiation is an important component of the urban environment, affecting 16 human populations directly through UV exposure of skin and eyes ¹⁻³ and less directly (but 17 with great impact) by driving the formation of photochemical smog, including tropospheric 18 ozone and other oxidants, as well as secondary aerosols containing nitrates, sulfates, and 19 organics. 4-6 These pollutants, along with others of primary origin commonly found in urban atmospheres (e.g., black carbon, sulfur dioxide), can in turn scatter and/or absorb UV 21 radiation, alter its vertical distribution, and so modify the photochemical rate of their own formation. Such feedback complicates the calculation of both the UV radiation field (including at the surface), and the evolution of photochemical smog in the urban boundary layer. 25 The question of how air pollution alters the urban UV environment (and vice versa) is 26 not new, but studies have relied mostly on numerical models, ⁷⁻⁹ with relatively few available 27 observations (e.g., McKenzie et al. 10, Panicker et al. 11, Palancar et al. 12, reviewed by Bais 28 et al. 13). Increases in UV have been estimated as a consequence of emission reductions over the last decades, e.g. in China, 14-16 and have led to less-than-expected reductions in photochemical smog, in part due to stronger UV photochemistry. 17-19 Emission reductions 31 have also occurred globally during the 2020 COVID-19 pandemic, ^{20,21} but ground-level ozone in some polluted areas has actually increased, ^{22,23} due at least in part to the increased UV radiation. Unfortunately, the observational data base of relevant UV radiation remains rather sparse to evaluate such model-predicted feedbacks between chemistry and radiation.

The environment of Mexico City is of particular interest for several reasons: (1) Nearly 36 23 million people inhabit the Mexico City Metropolitan Area (MCMA), and the UV en-37 vironment has direct implications for their health, both in terms of skin/eye UV exposure 38 and via photochemical smog formation. (2) As a tropical megacity, it is to some extent 39 representative of the situation of many others, with year-round intense midday UV irradi-40 ance, a shallower atmosphere due to the city's high elevation of 2240 m above sea level, 41 and a transition toward newer and cleaner technologies, leading to gradual improvements 42 in air quality. (3) Air quality within MCMA has undergone extensive scrutiny, with a well-43 established monitoring network operating since 1986, ²⁴ numerous intensive field campaigns to study the meteorology, emissions, and photochemistry of smog formation, ^{25–27} and numerical modeling incorporating the evolving knowledge. ^{28–31} This extensive body of knowledge 46 provides the foundation for understanding our study. 47

Here, we analyze two decades of continuous measurements of the UV Index at multiple locations within the MCMA, collected by the Secretariat of the Environment (Secretaria del Medio Ambiente, SEDEMA)¹ of the Mexico City government as part of an intensive monitoring network over the MCMA. The UV Index is defined as:

$$UVI = 40 \int_{250nm}^{400nm} E(\lambda, t) \cdot S_{er}(\lambda) d\lambda$$
 (1)

where $E(\lambda, t)$ is the solar spectral irradiance in units of W·m⁻²·nm⁻¹ and $S_{er}(\lambda)$ is the erythemal sensitivity of human skin. 32,33 Multiplication by 40 was chosen historically to express the UVI in small integer numbers, but is otherwise scientifically arbitrary.

The UVI is recognized by the World Health and Meteorological Organizations (WHO and WMO) as a standardized index of UV radiation³² for global public information. An advantage of using the UVI as (one) measure of UV radiation is that it is being increasingly observed or calculated and disseminated, enabling more objective comparisons among seasons

¹https://www.sedema.cdmx.gob.mx/

and locations. The UVI observations from Mexico City, considered here, are an important element of this global picture.

While the UVI at the surface cannot be translated directly into photolysis frequencies for various photo-labile molecules, the spectral weighting of the UVI (ca. 300-320 nm) is approximately similar to that for the photolysis of ozone to singlet oxygen atoms. Other UV wavelengths are of course also important, e.g. for the photolysis of nitrogen dioxide, and may be affected differently depending on the pollutant. With these considerations and a few other caveats, UVI trends examined here can also be used to infer accompanying trends in photolysis frequencies and influences on photochemical smog formation.

68 Methods

59 Ground-based measurements

The Mexico City Metropolitan Area is located at 19.4°N, 99.1°W, 2240 meters above sea level (m asl), surrounded by mountain ridges exceeding 5000 m asl, with complex topography and 71 thermal inversions that inhibit winds and favor intense air pollution. 34-36 Air quality monitoring and surface meteorological measurements in the MCMA are conducted continuously by the Automated Atmospheric Monitoring Network (RAMA, by its Spanish acronym) of the Atmospheric Monitoring System (SIMAT, by its Spanish acronym) of the Mexico City government. Since the year 2000, UV radiometers (model 501-A, Solar Light Company Inc., Glenside, PA) detecting wavelengths between 280-400 nm have been measuring erythemally-77 weighted solar radiation. Calibration of the UV sensors has been carried out annually by comparing against a factory-calibrated reference sensor. The output voltages from the net-79 work's sensors have been compared for at least one week against the UV readings from the 80 reference sensor to derive new calibration factors, which typically differed from the old ones 81 by 2% or less. Reference sensors were also calibrated by the manufacturer and updated periodically (between 3 to 5 years) to reduce systematic errors due to drifts in sensitivity. Long

term calibration drifts are minimized by these yearly re-calibrations. Although at the beginning only a few stations were in operation and varied, currently 11 stations are recording 85 erythemal irradiances, which are then multiplied by 40 (see Eq. 1) to give UV Indices. Table 1 describes the location of the stations where the UV Index has been measured. Figure 1 87 shows how the radiometers of the SIMAT have been distributed over MCMA, prioritizing 88 sites with higher population density. Near real-time data for each station are available on 89 the SIMAT official website http://www.aire.cdmx.gob.mx/default.php. Daily maximum 90 values (UVI_{max}) were extracted in each of the stations around solar noon from the time inter-91 val from 11:00 h-15:00 h CST (Central Standard Time). This database with 7305 continuous days of measurements during the period 2000-2019 was analyzed.

Table 1: SIMAT stations and AERONET site*, environmental descriptors and geographical positions. Abbreviations names: Chalco (CHA), Cuautitlán (CUT), FES Acatlán (FAC), Hangares (HAN), Laboratorio de Análisis Ambiental (LAA), Merced (MER), Montecillo (MON), Milpa Alta (MPA), Pedregal (PED), San Agustín (SAG), Santa Fe (SFE), Tlalnepantla (TLA) and National Autonomous University of Mexico (UNAM*)

Station	Environment	Lat (°N)	Lon (°W)	El (m asl)
СНО	semi-urban	19.27	98.89	2253
CUT	ecological park	19.72	99.20	2263
FAC	urban	19.48	99.24	2299
HAN	urban	19.42	99.08	2235
LAA	urban	19.48	99.15	2255
MER	downtown	19.42	99.12	2245
MON	rural	19.46	98.90	2252
MPA	rural	19.18	98.99	2594
PED	residential	19.33	99.20	2326
SAG	urban	19.53	99.03	2241
SFE	residential	19.36	99.26	2599
TLA	urban	19.53	99.20	2311
$UNAM^*$	University city	19.33	99.18	2294

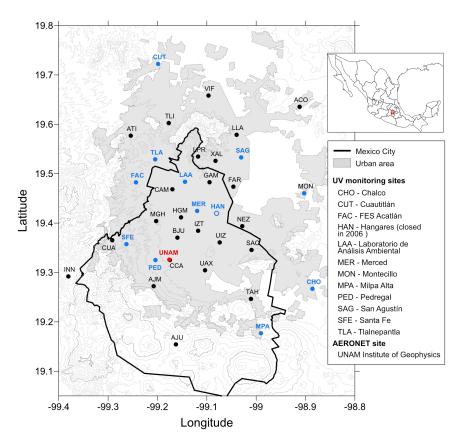


Figure 1: Map with the location of the SIMAT continuous monitoring stations over MCMA. Sites denoted by the blue solid dots correspond to SIMAT stations with UV measurements, while the black dots indicate SIMAT stations without UV measurements, the site indicated by an open blue dot represents a discontinued site, the red dot shows the location of the AERONET site. The location of Mexico City and the acronyms of the UV and AERONET site are shown at the upper and lower frames at the right respectively.

To explore the relationship between UVI and air pollutants levels, the hourly averages 94 for ozone (O_3) , carbon monoxide (CO), nitrogen dioxide (NO_2) , sulfur dioxide (SO_2) and 95 particle matter with diameter sizes $\leq 10~\mu m$ (PM₁₀), were downloaded from the SIMAT.³⁷ 96 For purposes of assessing the influence on the UV Index at solar noon, only values obtained 97 between 11h and 15h CST were considered for the trends analysis. Pollutant measurements 98 are conducted by the SIMAT using regulatory-grade commercial instruments. Measurement 99 principles include ultraviolet photometry (model 400E, Teledyne-API) for O₃, chemilumines-100 cence (model 200E, Teledyne-API) for NO₂, UV fluorescence (model 100E, Teledyne API) 101 for SO_2 , and infrared absorption (model 300E, Teledyne-API) for CO. The PM_{10} continuous 102

mass concentration was measured with Tapered Element Oscillating Microbalance (TEOM 1400AB or TEOM 1405 DF, Thermo Scientific) monitors. Gaseous pollutant levels are reported in ppb concentration units for O_3 , SO_2 and NO_2 , and in ppm for CO. Particulate matter mass concentration is reported in μg m⁻³ at local temperature and pressure.

Aerosol optical depth at 340 nm was obtained from the Institute of Geophysics of the 107 National Autonomous University of Mexico (UNAM). Measurements were conducted with 108 a CIMEL sun photometer model CE-318, which is an automatic sun-sky-scanning spectral 109 radiometer of the AErosol RObotic NETwork (AERONET³⁸). The data Product Level 110 2.0 and 1.5 (only in 2019) were selected, and annual averages AOD₃₄₀ were calculated for 111 the period 2000-2019, except for the year 2011 which contained large data gaps and was 112 therefore not used. A previous study of the AOD behavior from 2000 to 2014 demonstrated 113 that data gaps did not significantly change the trends calculated for this. ³⁹ Only 8% of the 114 data between the years 2014 and 2019 were missing. The largest data gap occurred in 2018 115 when the instrument was returned for calibration. 116

117 Satellite data

The UV Index data derived from satellite-based measurements were used for comparison 118 with the ground-based measurements. These data were provided by the Ozone Monitoring 119 Instrument (OMI) on board of AURA-NASA satellite. 40 OMI is operated collaboratively by 120 the Netherlands Agency for Aerospace Programmes (NIVR), the Finnish Meteorological In-121 stitute (FMI) and NASA. OMI performs observations over a geographical area of 13×24km² 122 at nadir. For the coordinates (19.33°N, 99.18°W) and elevation (2268 m asl) of Mexico City, 123 the satellite overpass time is between 19:00h - 21:00h UTC (20:00h - 22:00h CET). The OMI 124 instrument measures reflected and backscattered radiances of the atmosphere-surface system 125 at several UV wavelengths. 41 The ozone column is estimated by differential absorption of 126 several adjacent wavelengths in the O₃ Huggins band ⁴², and is used in a radiative transfer model to estimate clear-sky UVI at the surface. Clouds are estimated from the observed 128

reflectivity at a wavelength not affected by O₃ absorption (360 nm) and the clear-sky UVI is adjusted for their presence. A correction for absorbing aerosols is made using a global aerosol climatology. The UV Index used for Mexico City corresponds to values at local noon time and clear sky from OMUVB Level 2 OVP data product, version 1.3 and collection 3.

134 TUV model

Calculations of the UV Index were also made with the Tropospheric Ultraviolet Visible (TUV 135 v5.3) model. 45 The model atmosphere is represented by 80 vertical layers, each 1 km thick, 136 starting at the 2.24 km as elevation of MCMA, and for which the first three km constitute 137 the atmospheric boundary layer (BL). 46 The ozone profile above the BL is from the US Standard Atmosphere, but rescaled to a value of 259.6 DU. Total ozone columns including 139 the BL contributions (of 13.7 DU in the year 2000 and 9.8 DU in 2019, see below) were 140 273.1 DU in 2000 and 269.2 DU in 2019, respectively. The climatological O₃ column for this 141 latitude and season is about 270 DU (plus or minus ca. 5 DU) so in good agreement with 142 the values used here. 143

Pollutants within the BL (including O₃) are assumed to be well mixed, in agreement with observations from the MILAGRO field campaign that showed the disappearance of vertical gradients in the profiles of gases ^{47,48} and aerosols ^{49,50} by late morning. The UV-absorbing gases considered here are O₃, NO₂, and SO₂, specified in ppb.

BL aerosols are modeled by prescribing the AOD at 340 nm (from AERONET observations), scaled to other wavelengths inversely with wavelength (Angstrom coefficient = 1.0), asymmetry factor of 0.7, and a single scattering albedo of 0.85 at UV wavelengths, following the determinations made in Mexico City by Corr et al. ⁵¹ and Palancar et al. ¹².

Above the atmospheric boundary layer, the model atmosphere was specified to be free of aerosols, NO₂, or SO₂. In one sensitivity study, a total AOD of 0.7 was redistributed placing 0.2 in the free troposphere (decreasing vertically with an exponential scale height of 4 km, and 0.5 remaining in the BL). The calculated UVI differed by less than 1%. Thus, as long as the total AOD is known, knowledge of the exact vertical aerosol profile is not critical for ground-level UVI, though it would obviously affect the vertical structure of photolysis frequencies.

Radiative transfer calculations were carried out with the pseudo-spherical 4-stream option, at 1 nm steps between 280 and 400 nm.

Results and Discussion

Figure 2 shows the diurnal variation of the UVI for typical cloud-free days, for different sea-162 sons and several locations (CHO, MER, MON, PED, SAG, SFE and TLA). The selection of 163 cloud-free days was aided by comparison of the measurements with the TUV model predic-164 tions for clear skies, with prolonged deviations and fluctuations usually indicative of cloudy 165 conditions. However, during the rainy period (from June to October) at least a brief episode 166 of clouds, is seen at CHO station around noon in 13 June 2017. Peak values range from 8 167 during autumn/winter to above 12 in spring/summer, in correspondence to the respective 168 December and June solstices. Although the stations are all within a 25 km radius, substantial 169 differences among them are notable. Survey of the locations revealed that shadowing from 170 nearby structures is not an issue. The good agreement in the morning, followed by more divergence in the afternoon, is consistent with the development of photo-chemical pollution 172 hotspots during the day. Previous studies (e.g., Castro et al. ⁵² and Palancar et al. ¹²) have shown that surface UV radiation in Mexico City is attenuated significantly by aerosols. The 174 measurements shown in Fig. 2 are consistent with this increasing pollution during the course 175 of the day, with highest aerosol loading (and highest variability) attained in the afternoon. 176 Further support for the role of pollution in suppressing the UVI comes from the observation 177 made at the Santa Fe (SFE) site which in Fig. 2 are seen to be systematically higher, e.g. 178 by over 10% in autumn afternoons, compared to the other stations. The SFE station is displaced to the west from the majority of the other stations, and remains in the outskirts.

This site is also approximately 300 m higher than Mexico City downtown, so that the cleaner atmosphere at SFE is a result of the station's location farther away from pollution sources, both horizontally and vertically. ⁵³ It is indeed expected to have higher values of the UVI, in agreement with the observations.

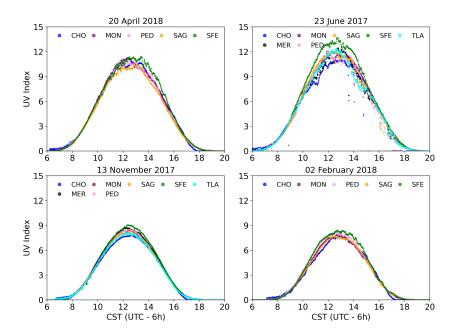


Figure 2: UV Index measurements at 1-minute resolution from several SIMAT stations within the MCMA under mostly cloudless conditions for representative days of the year.

The daily maximum UV Index values of each station, UVI_{max} , are summarized in Figure 3 185 for the entire period 2000-2019. These values are seen to range from 1 to 16, with a majority 186 (61%) of the days experiencing UVI_{max} values between 6 and 10, and remarkably few, less 187 than 1%, in the 13-16 range. The lowest values are likely due to winter days with heavy 188 cloud cover and low sun angles, and UV attenuation by pollutants could also be amplified 189 under such conditions, due to longer photon path lengths at low sun and within clouds. The 190 extreme sparsity of high UVI values remains surprising, and may be an indication of the 191 rarity of extremely unpolluted days within the city. 192

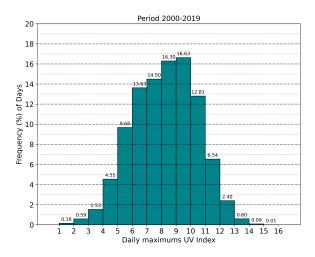


Figure 3: Frequency distribution of daily maximums UV Index values in Mexico City during 2000-2019.

Similar patterns are found when considering the monthly average of the UVI_{max} values 193 as shown in Figure 4. The long-term averages present a seasonal variation (as in Fig. 2) that 194 follows approximately the cosine of the noontime solar zenith angle. Notably, values rarely 195 if ever exceed 12 (as in Fig. 3). The lowest average UVI (near 7) occurs in winter while 196 from March to August the values seem to be flattened in the range 10-11. The rather low 197 monthly UV Index values could be a consequence of the presence of clouds in the rainy season. 198 However, urban aerosol pollution sources, biomass burning for agriculture and wood cooking 199 also contribute to poor air quality between March-May 54 . Using the maximum UVI_{max} from 200 all of stations every day (one daily point), the monthly averages (\overline{UVI}_m) were calculated 201 along the period 2000-2019 (except for June 2003 due to lack of measurements). Long term 202 trends in \overline{UVI}_m are shown in Figure 5. A clear upward trend is seen, with a slope for the 203 linear fit of 0.9%/year or +1.5 UVI units over the two decades. 204

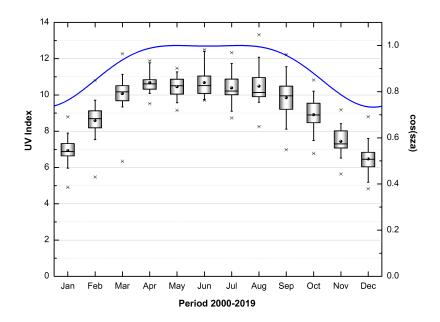


Figure 4: Monthly averages of the maximum UV Index values (black dot) in MCMA for the period 2000-2019: median (central bold line), interquartile range (box edges), 10% and 90% percentiles (whiskers), the minimum and maximum values (asterisk) and the cosine of the solar zenith angle at solar noon (blue curve).

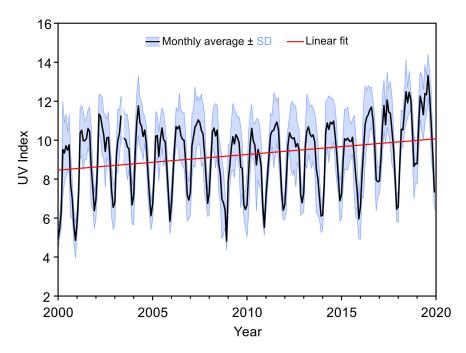


Figure 5: Monthly average UV Index (black line), the standard deviation (blue fill area) and linear fit (red line).

The UV Index computed from OMI satellite-based observations over the period 2005-2019 205 is shown in Figure 6. The satellite-derived UVIs vary from 8 in winter to 16 in summer, 206 these values being substantially higher than the ground-based observations (ca. 7 for winter 207 and 11 for summer, see Fig. 4). The maximum values are reached in June and July of each 208 year for both data sets. We hypothesize that this large difference between satellite-based 209 estimation and ground-based observation of the UV index is due to the intense air pollution 210 of Mexico City. A rather similar behavior was detected in the city of Santiago, Chile. 55 211 Although the OMI data are corrected for climatological aerosol absorption ⁴⁴, this correction 212 likely underestimates the absorption within the polluted Mexico City BL. 213

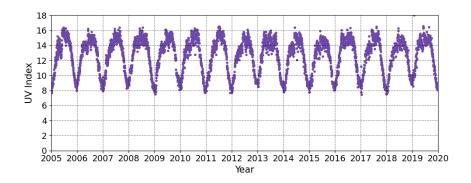


Figure 6: UV Index at solar noon for clear sky recorded by OMI-Aura/NIVR-FMI-NASA, from 2005 to 2019.

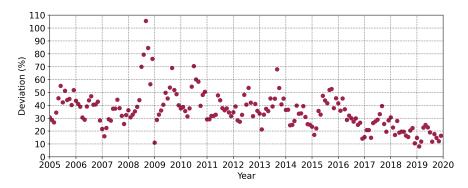


Figure 7: Deviation (%) of satellite-derived UVI from the ground-based monthly means UVI for the period 2005-2019.

Figure 7 shows the deviation of the OMI-derived UVI from the ground-based monthly

214

averages. The mean deviation was 35% in the period 2005-2019. Comparable UV reductions, 215 of 30-40% due to aerosols, were reported by Panicker et al. 11 over Pune, India from April 216 2004 to March 2005, with sensitivity coefficients (i.e. change in UVI per unit change in 217 AOD) similar to those found here in Table 3. 218 Comparisons of OMI with ground-based UV measurements have been reviewed recently 219 by Zhang et al. ⁵⁶ and Vitt et al. ⁵⁷. OMI-derived UV generally overestimates ground-level 220 measurements by 1-10% in relatively clean conditions (e.g. rural U.S), by 10-30% in Southern 221 Europe and by 40% or more in Santiago, Chile⁵⁵ and Thailand⁵⁸. The overestimations appear 222 related in large part to incomplete accounting of UV absorption by BL aerosol, although 223

other factors such as the correction to solar noon may also introduce some bias. ⁵⁶ Over Europe, ground-based UVI observations for several decades are systematically lower than those estimated from satellites even after consideration of climatological aerosol distributions, showing the importance of local pollution not resolved from space. ⁵⁷ However, the difference between satellite-derived and ground-based UVI was less than 1.0 UVI units in over 90% of the cases, in contrast to the difference of 3-5 units found here for Mexico City (Table 3).

230 Effect of pollutants on UV radiation in Mexico City

Trends and averages in aerosol optical depth AOD_{340} and criteria pollutants PM_{10} , CO, NO_2 , O_3 and SO_2 observed at the SIMAT stations over 2000-2019, are shown in Figure 8 and summarized in Table 2 together with the UVI_{max} . Similar trends in pollutants have been noted before $^{59-61}$ and reflect the long-term success of emission reduction policies and programs.

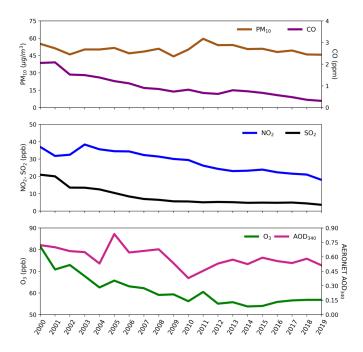


Figure 8: Air quality trends in MCMA for the period 2000-2019. Annual averages obtained between 11h to 15h CST every day: PM_{10} (brown curve), CO (purple curve), NO_2 (blue curve), SO_2 (black curve), O_3 (green curve) and AOD_{340} (magenta curve).

Table 2: Trends of UV Index and criteria pollutants over 2000-2019. Units: $\mu g/m^3$ for PM₁₀, ppm for CO, ppb for SO₂, NO₂ and O₃.

Variable	$\frac{absolute change}{year}$	Avg ₂₀₀₀₋₂₀₁₉	$\Delta(\%/year)$
UVI	0.08	9.2	0.9
PM_{10}	-0.10	50.2	-0.2
CO	-0.08	1.0	-8.2
NO_2	-0.96	28.6	-3.4
O_3	-1.05	61.3	-1.7
AOD_{340}	-0.01	0.6	-1.6
SO_2	-0.76	8.4	-9.1

The observed changes in the concentrations of these air pollutants have significant impli-236 cations for surface UV radiation, as can be demonstrated with the TUV radiative transfer 237 model. Table 3 summarizes UVI values for the June-July time period, estimated by the three 238 methods: OMI satellite-derived UVI, SIMAT ground-based observations, and TUV model-239 ing using air pollution estimates. Two groups of values can be readily identified: (1) SIMAT 240 daily record values, OMI with or without BL aerosols, and TUV for very clean conditions, all 241 with UVI values around 15-16; and (2) the average of daily maxima of SIMAT observations 242 and TUV UVI results using MCMA pollutants as input. Note that SIMAT and TUV value are in good agreement for both 2000-2001 and 2018-2019 and are considerably lower than for group (1). Compared to the OMI estimate that included the climatological absorbing 245 aerosol correction (15.3), observed SIMAT values were lower by 35% in 2000 and by 20% in 246 2019. Similarly, TUV values were 35% lower in 2000 and 22% lower in 2019 relative to the 247 TUV values of 15.6 for a pristine atmosphere.

Table 3: Comparison of June-July UVI maxima measured directly (SIMAT), inferred from satellite-based observations (OMI), and modeled for different pollution levels (TUV).

June-July	UV Index
SIMAT average maxima 2000-2001	9.9 ± 1.2
SIMAT average maxima 2018-2019	12.3 ± 1.3
SIMAT maximum value reached in period 2000-2019	15.0
OMI maximum value reached in period 2005-2019	16.6
TUV "zero" pollution	16.1
$AOD 0, O_3 0 ppb, NO_2 0 ppb, SO_2 0 ppb$	10.1
TUV pristine atmosphere	15.6
AOD 0.05 , O_3 10 ppb, NO_2 0 ppb, SO_2 0 ppb	10.0
TUV 2019:	12.1
$AOD~0.5,~O_3~50~ppb,~NO_2~20~ppb,~SO_2~1~ppb$	12.1
TUV 2000:	10.2
AOD 0.7 , O_3 70 ppb, NO_2 40 ppb, SO_2 20 ppb	10.2

The agreement between SIMAT observations and TUV model estimates is excellent but 249 also probably a bit fortuitous. Clouds on average reduce the irradiance impingent on the 250 surface, but scattering from them can also cause transient enhancements (these are only 251 possible if the disk of the Sun is visible) that could be recorded as daily maxima – with 252 cancellation between these cloud effects resulting in improved agreement with the cloud-free 253 model. We cannot exclude that some of the observed trend in UVI is due to changes in 254 cloud cover. However, the modeled UVI reductions due to pollutants, shown in Table 3, are 255 in such good agreement with the observed UVI reductions, that a compelling case can be 256 made for a dominant role of air pollutants in the long-term UVI trends. 257

Table 4 shows the contribution of different pollutants to the UVI increase between 2000 and 2019, computed with the TUV model. The year 2000 was taken as the base case, and each pollutant was decreased to its 2019 value individually while the others were maintained at their 2000 value. Decreasing aerosol AOD has been the largest driver of the UVI changes, but the absorption by gaseous pollutants has also decreased significantly. Additional future reductions in AOD, NO₂, and O₃ could increase the UVI further, while the low SO₂ levels reached by 2019 are expected to have negligible effects in the UVI.

UV reductions by air pollutants are expected to be most severe near the surface, while

chemical reactions leading to photochemical smog occur throughout the vertical extent of the 266 BL. Figure 9 shows the vertical profiles of photolysis coefficients (the reciprocals of photolytic 267 lifetimes) computed with the TUV model, for two key reactions, the photolysis of O₃ to yield 268 excited oxygen atoms O(1D), and the photolysis of NO_2 . In the absence of optically active 269 pollutants, these coefficients would be nearly independent of altitude in the BL. However, 270 the presence of pollutants leads to a strong decrease toward the surface, with notably more 271 severe reductions in 2000 compared to 2019. While the specific values shown in the figure 272 are only illustrative for typical conditions, routine daily air quality modeling should carefully 273 account for the long term variations in these photolysis coefficients. 274

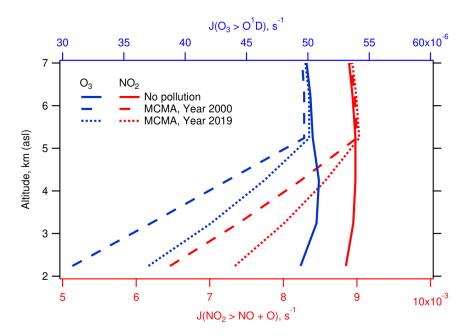


Figure 9: Vertical profiles of the photolysis coefficients for the reaction $O_3+h\nu \to O(1D)+O_2$ (top axis, blue) $NO_2+h\nu \to O+NO$ (bottom axis, red) for zero pollution (solid) and Mexico City in the years 2000 (dashed) and 2019 (dotted)

An issue that is beginning to gain relevance in radiative balance models is the influence of a group of organic compounds known as "brown carbon" that are capable of strongly absorbing in the UV region. ⁶² Some of these compounds are related to emissions from local and regional wildfires. ⁶³ Mexico City is frequently exposed to regional fire smoke transport during the dry part of the year (November to May) ⁶⁴, which sporadically modify the optical

275

276

277

278

properties of the aerosols. This could partly explain the relatively minor reductions in PM₁₀ (see Fig. 8), compared to the larger reductions of CO, NO₂, and O₃ that are more directly related to urban activities, as well as some of the seasonal asymmetry seen in Fig. 3.

The UVI is specific to wavelengths mainly in the 300-320 nm range, and so the question remains whether these results can be applied at longer UV wavelengths, e.g. those important for NO₂ photolysis (<420 nm). Absorption by SO₂ and O₃ vanishes for wavelengths larger than about 330 nm, while NO₂ photolysis is mostly driven by UV-A radiation and typical aerosols optical depth decrease with wavelength. These changes can easily be modeled, but unfortunately far fewer measurements of these longer wavelengths are available in Mexico City or elsewhere.

Two decades of observations in Mexico City demonstrate unequivocally that air pollution 291 reduces UV radiation at the ground. The ground-based observations are well below values 292 derived from satellite-based observations. Although ideally both data sets should match, 293 the satellite has a limitation to seeing boundary layer absorbers. Thus, they do not reveal 294 long-term UVI trends, due to substantial air quality improvements, that are evident in the 295 ground-based UVI data. When typical values of pollutant: aerosols, tropospheric ozone, NO_2 and SO_2 , are included in a model (e.g. TUV), the differences between satellite-derived 297 and ground-based measured values are explained and can be attributed quantitatively to 298 individual observed pollutants. Long term improvements in air quality over the last two 299 decades, were accompanied by statistically significant increases in the observed UVI, in 300 agreement with the model-predicted changes. 301

Table 4: Contributions of individual pollutants to the 2000-2019 UV Index changes. Baseline calculation for 2000 conditions (AOD = 0.7, $O_3 = 70$ ppb, $NO_2 = 40$ ppb, $SO_2 = 20$ ppb). Individual pollutant contributions are estimated by adjusting each pollutant, one at the time, to its 2019 value (AOD = 0.5, $O_3 = 50$ ppb, $NO_2 = 20$ ppb, $SO_2 = 1$ ppb), while maintaining the others at their 2000 value.

Pollutant	UVI Change
AOD	0.8
O_3	0.4
NO_2	0.4
SO_2	0.5

The reductions in surface UV radiation with respect to an ideally clear atmosphere – 302 by nearly 40% in 2000 and still 20% in 2020 – are large, both within the context of hu-303 man UV exposure and air quality mitigation. In urban areas where ozone production scales 304 proportionally with UV levels and with volatile organic compound (VOC) emissions (the 305 VOC-limited regime), a 10% increase in BL average photolysis rates means that VOC emis-306 sions will need to be reduced by 10% to meet the same goals, or else successful reductions 307 in aerosols would lead to unwanted UV-driven increases in O₃. Such UV changes must be 308 considered carefully in air quality mitigation strategies. For human exposure, a 20% increase 309 in UV irradiances over two decades should be seen as a non-negligible public health issue 310 requiring some reassessment of preventive behaviors to minimize the risk of skin cancer, 311 cataract, and other UV-related health effects. The efforts that Mexico City has made to 312 improve air quality have reduced levels of most air pollutants. Nevertheless, they caused an 313 increase in UV radiation that reaches the surface. 314

Finally, based on our results, the solar radiation monitoring network might be improved by adding observation sites far enough away to lessen the influence of the urban plume (e. g., Amecameca; latitude: 19.13°N, longitude: 99.76°W, elevation: 2420 m asl), and the adding sensors to existing monitoring sites at higher elevations (Ajusco, AJU, 2953 m asl or Instituto de Investigaciones Nucleares, INN, 3082 m asl). These sites would provide valuable observations to assess quantitatively the effects of urban pollution on solar radiation levels.

321 Acknowledgement

We wish to acknowledge the staff of SIMAT, from the Secretariat of Environment, for the
data and the continuous assistance during the realization of this project. Adriana Ipiña would
like to extend her thanks to Dirección General de Personal Académico, Universidad Nacional
Autónoma de México (DGAPA-UNAM) for the postdoctoral fellowship at Centro de Ciencias
de la Atmósfera of the UNAM. Rubén D Piacentini wishes to thank CONICET and National
University of Rosario, Argentina, for their partial support to the present work. The National
Center for Atmospheric Research is sponsored by the National Science Foundation.

329 References

- (1) Taylor, H. R.; West, S. K.; Rosenthal, F. S.; Muñoz, B.; Newland, H. S.; Abbey, H.;
 Emmett, E. A. Effect of Ultraviolet Radiation on Cataract Formation. New England
 Journal of Medicine 1988, 319, 1429–1433.
- 333 (2) Varotsos, C.; Feretis, E. Health effects on human eye resulting from the increased 334 ambient solar ultraviolet radiation. *Toxicological & Environmental Chemistry* **1997**, 335 61, 43–68.
- 336 (3) Lucas, R. M.; Yazar, S.; Young, A. R.; Norval, M.; de Gruijl, F. R.; Takizawa, Y.;

 Rhodes, L. E.; Sinclair, C. A.; Neale, R. E. Human health in relation to exposure to solar

 ultraviolet radiation under changing stratospheric ozone and climate. *Photochemical & Photobiological Sciences* 2019, 18, 641–680.
- 4) Leighton, P. A. *Photochemistry of Air Pollution*; Elsevier, 1961; pp v-vi.
- (5) Seinfeld, J. H.; Pandis, S. N.; Noone, K. Atmospheric Chemistry and Physics: From
 Air Pollution to Climate Change. *Physics Today* 1998, 51, 88–90.

- 343 (6) Finlayson-Pitts, B. J.; Pitts, J. N. Chemistry of the Upper and Lower Atmosphere;

 Academic Press, 2000; pp xvii–xviii.
- (7) Liu, S. C.; McKeen, S. A.; Madronich, S. Effect of anthropogenic aerosols on biologically
 active ultraviolet radiation. Geophysical Research Letters 1991, 18, 2265–2268.
- (8) Sabziparvar, A. A.; Forster, P. M. F.; Shine, K. P. Changes in ultraviolet radiation due
 to stratospheric and tropospheric ozone changes since preindustrial times. *Journal of Geophysical Research: Atmospheres* 1998, 103, 26107–26113.
- 9) Madronich, S.; Wagner, M.; Groth, P. Influence of Tropospheric Ozone Control on Exposure to Ultraviolet Radiation at the Surface. *Environmental Science & Technology*2011, 45, 6919–6923.
- McKenzie, R. L.; Weinreis, C.; Johnston, P. V.; Liley, B.; Shiona, H.; Kotkamp, M.;
 Smale, D.; Takegawa, N.; Kondo, Y. Effects of urban pollution on UV spectral irradiances. Atmospheric Chemistry and Physics 2008, 8, 5683–5697.
- (11) Panicker, A. S.; Pandithurai, G.; Takamura, T.; Pinker, R. T. Aerosol effects in the
 UV-B spectral region over Pune an urban site in India. Geophysical Research Letters
 2009, 36, L10802 1–5.
- Palancar, G. G.; Lefer, B. L.; Hall, S. R.; Shaw, W. J.; Corr, C. A.; Herndon, S. C.;
 Slusser, J. R.; Madronich, S. Effect of aerosols and NO₂ concentration on ultraviolet
 actinic flux near Mexico City during MILAGRO: measurements and model calculations.
 Atmospheric Chemistry and Physics 2013, 13, 1011–1022.
- (13) Bais, A. F.; McKenzie, R. L.; Bernhard, G.; Aucamp, P. J.; Ilyas, M.; Madronich, S.;
 Tourpali, K. Ozone depletion and climate change: impacts on UV radiation. *Photo-chemical & Photobiological Sciences* 2015, 14, 19–52.

- Heard, D.; Liu, D. Photochemical impacts of haze pollution in an urban environment.

 Atmospheric Chemistry and Physics 2019, 19, 9699–9714.
- of 2013–2017 trends in summer surface ozone in China. *Proceedings of the National Academy of Sciences* **2018**, *116*, 422–427.
- Wang, Y. et al. Contrasting trends of PM2.5 and surface-ozone concentrations in China from 2013 to 2017; Oxford University Press (OUP), 2020; Vol. 7; pp 1331–1339.
- wang, W.; Li, X.; Shao, M.; Hu, M.; Zeng, L.; Wu, Y.; Tan, T. The impact of aerosols on photolysis frequencies and ozone production in Beijing during the 4-year period 2012–2015. Atmospheric Chemistry and Physics 2019, 19, 9413–9429.
- 377 (18) Gao, J.; Li, Y.; Zhu, B.; Hu, B.; Wang, L.; Bao, F. What have we missed when studying
 378 the impact of aerosols on surface ozone via changing photolysis rates? *Atmospheric*379 *Chemistry and Physics* **2020**, *20*, 10831–10844.
- 380 (19) Ma, X.; Huang, J.; Zhao, T.; Liu, C.; Zhao, K.; Xing, J.; Xiao, W. Rapid increase in 381 summer surface ozone over the North China Plain during 2013–2019: a side effect of 382 particulate matters reduction control? Atmospheric Chemistry and Physics 2021, 21, 383 1–16.
- Bauwens, M.; Compernolle, S.; Stavrakou, T.; Müller, J.-F.; Gent, J.; Eskes, H.; Levelt, P. F.; A, R.; Veefkind, J. P.; Vlietinck, J.; Yu, H.; Zehner, C. Impact of coronavirus outbreak on NO2 pollution assessed using TROPOMI and OMI observations. **2020**, 47, e2020GL087978.
- ³⁸⁸ (21) Venter, Z. S.; Aunan, K.; Chowdhury, S.; Lelieveld, J. COVID-19 lockdowns cause global air pollution declines. *Proceedings of the National Academy of Sciences* **2020**, 117, 18984–18990.

- ³⁹¹ (22) Shi, X.; Brasseur, G. P. The Response in Air Quality to the Reduction of Chinese Economic Activities During the COVID-19 Outbreak. *Geophysical Research Letters*³⁹³ **2020**, 47, e2020GL088070.
- ³⁹⁴ (23) Le, T.; Wang, Y.; Liu, L.; Yang, J.; Yung, Y. L.; Li, G.; Seinfeld, J. H. Unexpected air pollution with marked emission reductions during the COVID-19 outbreak in China. ³⁹⁶ Science **2020**, 369, 702–706.
- 397 (24) Secretaría del Medio Ambiente, Gobierno de la Ciudad de México, México, Red
 398 Automática de Monitoreo Atmosférico (RAMA). http://www.aire.cdmx.gob.mx/
 399 descargas/datos/excel/RAMAxls.pdf.
- 400 (25) Doran, J. C. et al. The IMADA-AVER Boundary Layer Experiment in the Mexico City

 401 Area. Bulletin of the American Meteorological Society 1998, 79, 2497–2508.
- 402 (26) Molina, L. T.; Kolb, C. E.; de Foy, B.; Lamb, B. K.; Brune, W. H.; Jimenez, J. L.;
 403 Ramos-Villegas, R.; Sarmiento, J.; Paramo-Figueroa, V. H.; Cardenas, B.; Gutierrez404 Avedoy, V.; Molina, M. J. Air quality in North America's most populous city overview
 405 of the MCMA-2003 campaign. Atmospheric Chemistry and Physics 2007, 7, 2447–2473.
- 406 (27) Molina, L. T. et al. An overview of the MILAGRO 2006 Campaign: Mexico City
 407 emissions and their transport and transformation. Atmospheric Chemistry and Physics
 408 2010, 10, 8697–8760.
- 409 (28) Jazcilevich, A. D.; García, A. R.; Caetano, E. Locally induced surface air confluence
 410 by complex terrain and its effects on air pollution in the valley of Mexico. Atmospheric
 411 Environment 2005, 39, 5481–5489.
- (29) Tie, X.; Madronich, S.; Li, G.; Ying, Z.; Zhang, R.; Garcia, A. R.; Lee-Taylor, J.; Liu, Y.
 Characterizations of chemical oxidants in Mexico City: A regional chemical dynamical
 model (WRF-Chem) study. Atmospheric Environment 2007, 41, 1989–2008.

- 415 (30) Zhang, Y.; Dubey, M. K.; Olsen, S. C.; Zheng, J.; Zhang, R. Comparisons of
 416 WRF/Chem simulations in Mexico City with ground-based RAMA measurements dur417 ing the 2006-MILAGRO. Atmospheric Chemistry and Physics 2009, 9, 3777–3798.
- (31) Zavala, M.; Brune, W. H.; Velasco, E.; Retama, A.; Cruz-Alavez, L. A.; Molina, L. T.
 Changes in ozone production and VOC reactivity in the atmosphere of the Mexico City
 Metropolitan Area. Atmospheric Environment 2020, 238, 117747.
- (32) WHO,; WMO,; UNEP,; ICNIRP, Global solar UV index: a practical guide; 2002; pp A
 joint recommendation of the World Health Organization, World Meteorological Organization, United Nations Environment Programme, and the International Commission on Non–Ionizing Radiation Protection.
- (33) Webb, A. R.; Slaper, H.; Koepke, P.; Schmalwieser, A. W. Know Your Standard: Clarifying the CIE Erythema Action Spectrum. *Photochemistry and Photobiology* 2011, 87,
 427 483–486.
- 428 (34) Whiteman, C. D.; Zhong, S.; Bian, X.; Fast, J. D.; Doran, J. C. Boundary layer evo-429 lution and regional-scale diurnal circulations over the and Mexican plateau. *Journal of* 430 *Geophysical Research: Atmospheres* **2000**, *105*, 10081–10102.
- 431 (35) Fast, J. D.; de Foy, B.; Rosas, F. A.; Caetano, E.; Carmichael, G.; Emmons, L.;
 432 McKenna, D.; Mena, M.; Skamarock, W.; Tie, X.; Coulter, R. L.; Barnard, J. C.;
 433 Wiedinmyer, C.; Madronich, S. A meteorological overview of the MILAGRO field cam434 paigns. Atmospheric Chemistry and Physics 2007, 7, 2233–2257.
- (36) Carreón-Sierra, S.; Salcido, A.; Castro, T.; Celada-Murillo, A.-T. Cluster Analysis of
 the Wind Events and Seasonal Wind Circulation Patterns in the Mexico City Region.
 Atmosphere 2015, 6, 1006–1031.
- 438 (37) Sistema de Monitoreo Atmosférico, Ciudad de México. http://www.aire.cdmx.gob.

- mx/default.php?opc=%27aKBhnmI=%27&opcion=Zg, Accessed on Wed, December 02, 2020.
- (38) Holben, B.; Eck, T.; Slutsker, I.; Tanré, D.; Buis, J.; Setzer, A.; Vermote, E.; Reagan, J.;
 Kaufman, Y.; Nakajima, T.; Lavenu, F.; Jankowiak, I.; Smirnov, A. AERONET—A
 Federated Instrument Network and Data Archive for Aerosol Characterization. Remote
 Sensing of Environment 1998, 66, 1–16.
- (39) Carabali, G.; Estévez, H. R.; Valdés-Barrón, M.; Bonifaz-Alfonzo, R.; Riveros Rosas, D.; Velasco-Herrera, V. M.; Vázquez-Gálvez, F. A. Aerosol climatology over
 the Mexico City basin: Characterization of optical properties. Atmospheric Research
 2017, 194, 190–201.
- 449 (40) NASA EOS/Aura Validation Data Center (AVDC) Correlative data, Field of
 450 View Predictions, Data Subsets, GEOMS, DCIO. https://avdc.gsfc.nasa.gov/pub/
 451 most_popular/overpass/OMI/.
- (41) Tanskanen, A.; Krotkov, N.; Herman, J.; Arola, A. Surface ultraviolet irradiance from
 OMI. IEEE Transactions on Geoscience and Remote Sensing 2006, 44, 1267–1271.
- Veefkind, J. P.; de Haan, J. F.; Brinksma, E. J.; Kroon, M.; Levelt, P. F. Total Ozone
 From the Ozone Monitoring Instrument (OMI) Using the DOAS Technique. *IEEE* Transactions on Geoscience and Remote Sensing 2006, 44, 1239–1244.
- 457 (43) Tanskanen, A. et al. Validation of daily erythemal doses from Ozone Monitoring Instrument with ground-based UV measurement data. *Journal of Geophysical Research:* 459 Atmospheres 2007, 112, D24S44 1–15.
- 460 (44) Arola, A. et al. A new approach to correct for absorbing aerosols in OMI UV. Geophys-461 ical Research Letters **2009**, 36, L22805 1–5.

- 462 (45) Madronich, S. Intercomparison of NO2 photodissociation and U.V. Radiometer Measurements. *Atmospheric Environment* **1987**, *21*, 569–578.
- 464 (46) Shaw, W. J.; Pekour, M. S.; Coulter, R. L.; Martin, T. J.; Walters, J. T. The daytime
 465 mixing layer observed by radiosonde profiler, and lidar during MILAGRO. Atmospheric
 466 Chemistry and Physics Discussions 2007, 7, 15025–15065.
- (47) Velasco, E.; Márquez, C.; Bueno, E.; Bernabé, R. M.; Sánchez, A.; Fentanes, O.; Wöhrnschie schimmel, H.; Cárdenas, B.; Kamilla, A.; Wakamatsu, S.; Molina, L. T. Vertical distribution of ozone and VOCs in the low boundary layer of Mexico City. Atmospheric Chemistry and Physics 2008, 8, 3061–3079.
- 471 (48) Greenberg, J.; Guenther, A.; Turnipseed, A. Tethered balloon-based soundings of ozone
 472 aerosols, and solar radiation near Mexico City during MIRAGE-MEX. Atmospheric
 473 Environment 2009, 43, 2672–2677.
- 474 (49) Rogers, R. R.; Hair, J. W.; Hostetler, C. A.; Ferrare, R. A.; Obland, M. D.; Cook, A. L.;
 475 Harper, D. B.; Burton, S. P.; Shinozuka, Y.; McNaughton, C. S.; Clarke, A. D.; Re476 demann, J.; Russell, P. B.; Livingston, J. M.; Kleinman, L. I. NASA LaRC airborne
 477 high spectral resolution lidar aerosol measurements during MILAGRO: observations
 478 and validation. Atmospheric Chemistry and Physics 2009, 9, 4811–4826.
- (50) Lewandowski, P. A.; Eichinger, W. E.; Holder, H.; Prueger, J.; Wang, J.; Kleinman, L. I.
 Vertical distribution of aerosols in the vicinity of Mexico City during MILAGRO-2006
 Campaign. Atmospheric Chemistry and Physics 2010, 10, 1017–1030.
- (51) Corr, C. A.; Krotkov, N.; Madronich, S.; Slusser, J. R.; Holben, B.; Gao, W.; Flynn, J.;
 Lefer, B.; Kreidenweis, S. M. Retrieval of aerosol single scattering albedo at ultraviolet
 wavelengths at the T1 site during MILAGRO. Atmospheric Chemistry and Physics
 2009, 9, 5813–5827.

- 486 (52) Castro, T.; Madronich, S.; Rivale, S.; Muhlia, A.; Mar, B. The influence of aerosols on 487 photochemical smog in Mexico City. *Atmospheric Environment* **2001**, *35*, 1765–1772.
- 488 (53) SEDEMA, Calidad del aire en la Ciudad de México, Informe 2017; 2018.
- 489 (54) Retama, A.; Baumgardner, D.; Raga, G. B.; McMeeking, G. R.; Walker, J. W. Sea490 sonal and diurnal trends in black carbon properties and co-pollutants in Mexico City.

 491 Atmospheric Chemistry and Physics 2015, 15, 9693–9709.
- 492 (55) Cabrera, S.; Ipiña, A.; Damiani, A.; Cordero, R. R.; Piacentini, R. D. UV index values
 493 and trends in Santiago Chile (33.5°S) based on ground and satellite data. *Journal of*494 *Photochemistry and Photobiology B: Biology* **2012**, 115, 73–84.
- 495 (56) Zhang, H.; Wang, J.; García, L. C.; Zeng, J.; Dennhardt, C.; Liu, Y.; Krotkov, N. A.
 496 Surface erythemal UV irradiance in the continental United States derived from ground 497 based and OMI observations: quality assessment trend analysis and sampling issues.
 498 Atmospheric Chemistry and Physics 2019, 19, 2165–2181.
- Vitt, R.; Laschewski, G.; Bais, A.; Diémoz, H.; Fountoulakis, I.; Siani, A.-M.;
 Matzarakis, A. UV-Index Climatology for Europe Based on Satellite Data. Atmosphere
 2020, 11, 727.
- Janjai, S.; Wisitsirikun, S.; Buntoung, S.; Pattarapanitchai, S.; Wattan, R.; Masiri, I.;
 Bhattarai, B. K. Comparison of UV index from Ozone Monitoring Instrument (OMI)
 with multi-channel filter radiometers at four sites in the tropics: effects of aerosols and
 clouds. International Journal of Climatology 2013, 34, 453–461.
- (59) Parrish, D. D.; Singh, H. B.; Molina, L.; Madronich, S. Air quality progress in North
 American megacities: A review. Atmospheric Environment 2011, 45, 7015–7025.
- 508 (60) Secretaría del Medio Ambiente de la Ciudad de México, Calidad del aire en la Ciudad de México, informe 2017; 2018; pp 1–160.

- (61) Molina,; Velasco,; Retama,; Zavala, Experience from Integrated Air Quality Management in the Mexico City Metropolitan Area and Singapore. Atmosphere 2019, 10,
 512.
- (62) Laskin, A.; Laskin, J.; Nizkorodov, S. A. Chemistry of Atmospheric Brown Carbon.
 Chemical Reviews 2015, 115, 4335–4382.
- (63) Gadhavi, H.; Jayaraman, A. Absorbing aerosols: contribution of biomass burning and
 implications for radiative forcing. Annales Geophysicae 2010, 28, 103–111.
- 517 (64) Rios, B.; Raga, G. B. Smoke emissions from agricultural fires in Mexico and Central

 518 America. Journal of Applied Remote Sensing 2019, 13, 1.
- 519 (65) Barnard, J. C.; Volkamer, R.; Kassianov, E. I. Estimation of the mass absorption cross 520 section of the organic carbon component of aerosols in the Mexico City Metropolitan 521 Area. Atmospheric Chemistry and Physics 2008, 8, 6665–6679.

522 Graphical TOC Entry

523



Some journals require a graphical entry for the Table of Contents. This should be laid out "print ready" so that the sizing of the text is correct.

Inside the tocentry environment, the font used is Helvetica 8 pt, as required by *Journal of the American Chemical Society*.

The surrounding frame is 9 cm by 3.5 cm, which is the maximum permitted for *Journal of the American Chemical Society* graphical table of content entries. The box will not resize if the content is too big: instead it will overflow the edge of the box.

This box and the associated title will always be printed on a separate page at the end of the document.