

Transient Models

HWRS 482 - 3/19/20

$$\frac{\partial^2 h}{\partial x^2} + \frac{\partial^2 h}{\partial y^2} + \frac{\partial^2 h}{\partial z^2} = \frac{S}{T} \frac{\partial h}{\partial t}$$

↑ Steady state = 0
transient ≠ 0

- ① In transient simulations Storage terms matter

- heads changing over time which is a function of how easily water is released from storage

S_s = specific storage ← confined

S_y = specific yield ← unconfined

- ② Stress periods + time steps

- now we discretize space + time!

- Just like with space smaller time steps are desirable (less error)
But trade off with computational costs

- Goal: use largest time step possible impacting our results

- $\Delta t_c = \frac{S_a^2}{4K_b} \sim \Delta x = \Delta y = a$

\uparrow layer thickness
conductivity

$\Delta t_c \propto$ grid size (smaller Δt if grid resolution n)

$\Delta t_c \propto 1/K$ (smaller Δt if things move faster)

$\Delta t_c \propto$ storativity

↳ Large S means lots of water out

per $\Delta h \Rightarrow$ less sensitivity of Δh to storage changes

↳ smaller Δt if Δh is more sensitive.

$\Delta t_c \propto 1/b$

↳ thicker layers transmit more water

↳ smaller Δt for larger layers

- timesteps can often ↑ as time goes on if your system is stabilizing.
(i.e. when largest changes occur in early stress periods)

- growth ts can be used for this \Rightarrow set an initial Δt and $a_{max} \Delta t$ and some growth rate

Timestep vs Stress period

- Timestep = Δt = time between modelflow solutions
- Stress period = period of time where all of the stresses are constant (i.e. Boundaries (wells etc))
- Stress periods are broken up into timesteps, heads are calculated at the end of each timestep.

Vertical Layers

- Layers that are thin enough to capture heterogeneities
- if your layers are too thin you will have too many + that hurts computation
- All layers must have the same xy spacing
- Δz is determined by the elevation of our layer boundaries
 - ↳ Δz & K are used to calculate T
 - ↳ T can be different in each node in each layer (Δz can vary across rows & columns)
- # layers depends on vertical head gradients & variability in K
 - ↳ Need 3 or more layers to represent variability in vertical leakage.
 - ↳ for particle tracking you need 3 or more "transition" units to capture head radicals across sharp contrasts in K
- Vertical conductance: leakage between layers \leftarrow average K_z between 2 layers divided by distance between centers

Layer types

- (1) Confined \Rightarrow Saturated thickness constant
 - easiest / cheapest to calculate
 - problematic if head drops below top of layer
- (2) Confined / Unconfined -
 - simulating w/ unconfined as a confined
 - assuming saturated thickness not changing much in time
 - faster convergence more stable.
 - use s_y instead s_s
 - T is constant (i.e. always saturated)

(3) Unconfined

- head is below top of layer
- saturated thickness = elevation of cut
- T changes with WT, $S = s_y$

(4) Convertible (default)

- automatically switch confined \leftrightarrow unconfined with head changes
- T varies with saturated thickness
- varies between s_s & s_y