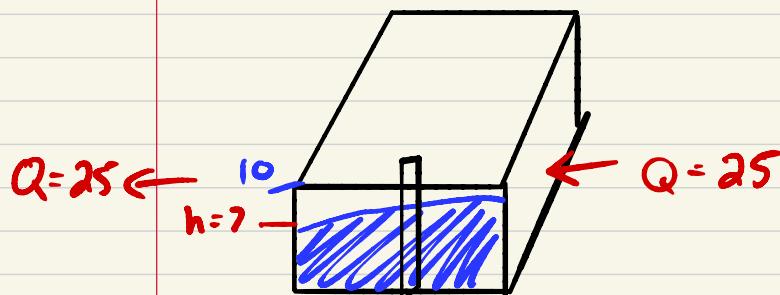


# Box Model HWI

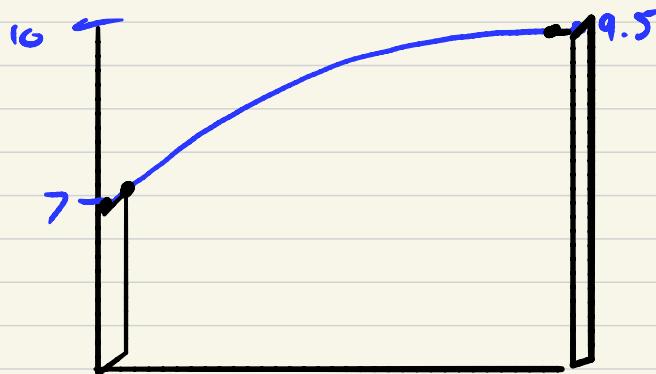
2-6-20



$$\text{inflow} - \text{outflow} = \Delta S$$

Steady state  
 $\Delta S = 0$

inflow = outflow



$$Q = KA \frac{dh}{dx}$$

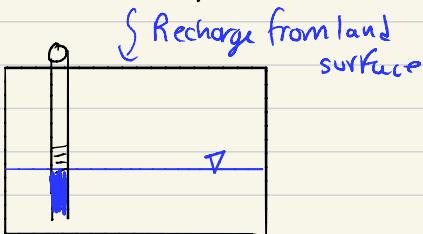
for  $Q$  to be constant

there must be an inverse relationship between  $A + \frac{dh}{dx}$

# Intro to Groundwater flow

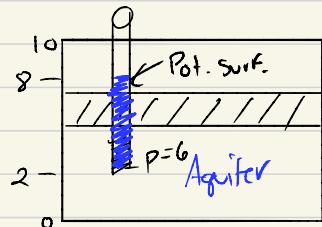
Aquifer: produces water in usable quantities  
Aquitard: confining unit

## Unconfined Aquifer



water level rises to water table

## Confined Aquifer



- upper layer is confining unit
- water will rise to potentiometric surface

## Key Hydrologic parameters

Storageability ( $s$ ) [-]: The volumetric change in water storage per unit surface area per unit change in head

Specific yield ( $S_y$ ) [-]: the change in storage caused by changes in water level

Specific storage ( $s_s$ ) [%] = the change in storage caused by compression or expansion of mineral skeleton (By changing pressure)

### Unconfined

- changes in storage by draining pore space

$$S = S_g + b S_s$$

↑  
Saturated thickness

\* Generally assume that  
 $S_g \gg S_s$

$$S = S_g$$

$$\sim 0.07 - 0.25$$

### Confined

- Storage change from saturated storage

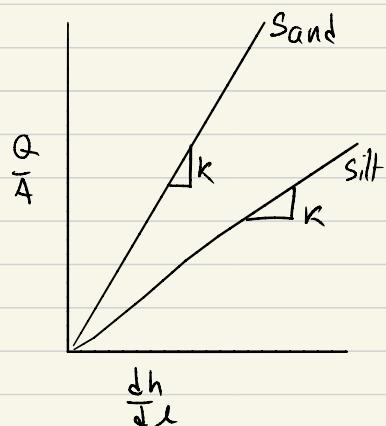
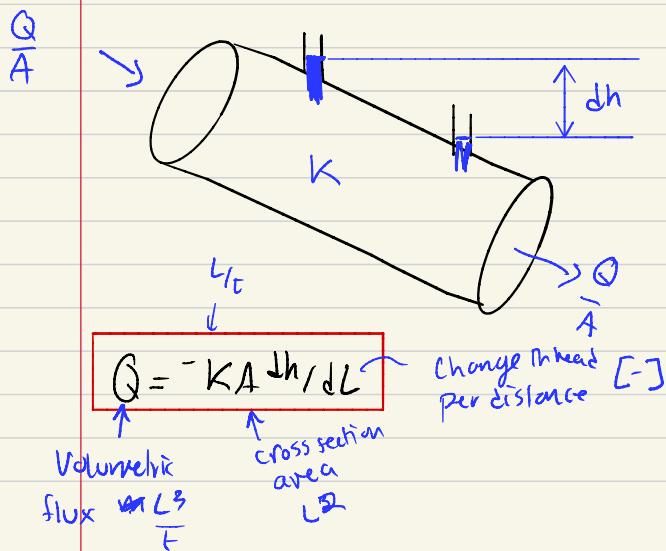
- Pot surf changes but not our water levels

$$S = b S_s$$

$$\sim 0.00005 - 0.005$$

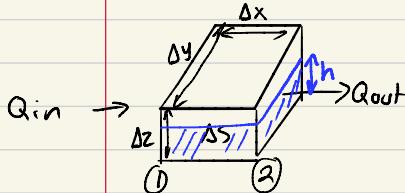
↑  
 small  $S_s$  values mean that large press. changes produce small changes in storage

### Saturated GW Flow: Darcy's Law (1856)



## Governing eqn

- 1) Conservation of mass:  $\text{out} - \text{in} = \Delta S$
- 2) Darcy's Law:  $Q = -kA \frac{dh}{dx}$



$$Q_{\text{in}} = -k \frac{\partial h}{\partial x} \Delta z \Delta y$$

$$Q_{\text{out}} = Q_{\text{in}} - \frac{\partial Q_{\text{in}}}{\partial x} \Delta x$$

$$Q_{\text{out}} = Q_{\text{in}} - \frac{\partial}{\partial x} \left( -k \frac{\partial h}{\partial x} \Delta y \Delta z \right) \Delta x$$

$\downarrow$

now apply conservation:

$$\begin{aligned} Q_{\text{out}} - Q_{\text{in}} &= \Delta S \\ \textcircled{1} \quad Q_{\text{out}} - Q_{\text{in}} &= \frac{\partial}{\partial x} \left( k \frac{\partial h}{\partial x} \Delta y \Delta z \right) \Delta x = \Delta S \end{aligned}$$

$$\textcircled{2} \quad \Delta S = S \Delta y \Delta x \frac{\partial h}{\partial t} * \text{because } S = \frac{\Delta S}{(\Delta x \Delta y) \frac{\partial h}{\partial t}}$$

↑ change in elevation  
unit area  
white space in head

\* Assuming Unconfined Aquifer

$$\Delta S = i h - \text{out} + R$$

combine egn ① + ②

$$\Delta S = \Delta S$$

$$S_y \Delta y \Delta x \frac{\partial h}{\partial t} = \frac{\partial}{\partial x} \left( k \frac{\partial h}{\partial x} h \Delta y \right) \Delta x + R \Delta y \Delta x$$

this is our  $\Delta S$

$$S_y \frac{\partial h}{\partial t} = \frac{\partial}{\partial x} \left( h \frac{\partial h}{\partial x} k \right) + R$$

$$\hookrightarrow h \frac{\partial^2 h}{\partial x^2} = \frac{1}{2} \frac{\partial^2 h^2}{\partial x^2}$$

because calculus

$$S_y \frac{\partial h}{\partial t} = \frac{1}{2} \frac{\partial}{\partial x} \left( \frac{\partial^2 h^2}{\partial x^2} \right) k + R$$

$$\boxed{\frac{2 S_y \frac{\partial h}{\partial t}}{K} = \frac{\partial^2 h^2}{\partial x^2} + \frac{2 R}{K}} \quad \leftarrow 1D \text{ flow, homogeneous unconfined aquifer}$$

$$\boxed{\frac{2 S_y \frac{dh}{dt}}{K} = \frac{d^2 h^2}{dx^2} + \frac{d^2 h^2}{dy^2} + \frac{d^2 h^2}{dz^2} + \frac{2R}{K}} \quad \leftarrow 3D \text{ flow}$$