Quasi-geostrophic dynamics for stratified fluids

$$\rho = \bar{\rho}(z) + \rho'(x, y, z, t) \quad \text{with} \quad |\rho'| \ll |\bar{\rho}|$$
$$p = \bar{p}(z) + p'(x, y, z, t)$$

Governing equations:

$$\frac{du}{dt} - f_0 v - \beta_0 y v = -\frac{1}{\rho_0} \frac{\partial p'}{\partial x}$$

$$\frac{dv}{dt} + f_0 u + \beta_0 y u = -\frac{1}{\rho_0} \frac{\partial p'}{\partial y}$$

$$0 = -\frac{\partial p'}{\partial z} - \rho' g$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0$$

$$\frac{\partial \rho'}{\partial t} + u \frac{\partial \rho'}{\partial x} + v \frac{\partial \rho'}{\partial y} + w \frac{d\bar{\rho}}{dz} = 0$$

Balance of large terms:

$$-f_0 v = -\frac{1}{\rho_0} \frac{\partial p'}{\partial x} + f_0 u = -\frac{1}{\rho_0} \frac{\partial p'}{\partial y}$$

$$u_g = -\frac{1}{f_0 \rho_0} \frac{\partial p'}{\partial y}$$

$$v_g = +\frac{1}{f_0 \rho_0} \frac{\partial p'}{\partial x}$$

Plug u_g and v_g into the small terms of the momentum equations, and neglect the vertical advection term:

$$\frac{du}{dt} - f_0 v - \beta_0 y v = -\frac{1}{\rho_0} \frac{\partial p'}{\partial x}$$

$$\frac{dv}{dt} + f_0 u + \beta_0 y u = -\frac{1}{\rho_0} \frac{\partial p'}{\partial y}$$

$$J(a,b) = \frac{\partial a}{\partial x} \frac{\partial b}{\partial y} - \frac{\partial b}{\partial x} \frac{\partial a}{\partial y}$$

$$- \frac{1}{\rho_0 f_0} \frac{\partial^2 p'}{\partial y \partial t} - \frac{1}{\rho_0^2 f_0^2} J\left(p', \frac{\partial p'}{\partial y}\right) - f_0 v - \frac{\beta_0}{\rho_0 f_0} y \frac{\partial p'}{\partial x} = -\frac{1}{\rho_0} \frac{\partial p'}{\partial x}$$

$$+ \frac{1}{\rho_0 f_0} \frac{\partial^2 p'}{\partial x \partial t} + \frac{1}{\rho_0^2 f_0^2} J\left(p', \frac{\partial p'}{\partial x}\right) + f_0 u - \frac{\beta_0}{\rho_0 f_0} y \frac{\partial p'}{\partial y} = -\frac{1}{\rho_0} \frac{\partial p'}{\partial y}$$

$$v_g = +\frac{1}{f_0 \rho_0} \frac{\partial p'}{\partial x}$$

$$u = u_g + u_a = -\frac{1}{\rho_0 f_0} \frac{\partial p'}{\partial y} - \frac{1}{\rho_0 f_0^2} \frac{\partial^2 p'}{\partial t \partial x}$$

$$-\frac{1}{\rho_0^2 f_0^3} J\left(p', \frac{\partial p'}{\partial x}\right) + \frac{\beta_0}{\rho_0 f_0^2} y \frac{\partial p'}{\partial y}$$

$$v = v_g + v_a = \frac{\rho_0 f_0}{\rho_0 f_0^2} \frac{\partial y}{\partial t} + \frac{\rho_0 f_0^2}{\rho_0 f_0^2} \frac{\partial t \partial x}{\partial y}$$

$$v = v_g + v_a = \frac{1}{\rho_0 f_0} \frac{\partial p'}{\partial x} - \frac{1}{\rho_0 f_0^2} \frac{\partial^2 p'}{\partial t \partial y}$$

$$-\frac{1}{\rho_0^2 f_0^3} J\left(p', \frac{\partial p'}{\partial y}\right) - \frac{\beta_0}{\rho_0 f_0^2} y \frac{\partial p'}{\partial x}$$

Substitution into the continuity equation:

$$\frac{\partial w}{\partial z} = \frac{1}{\rho_0 f_0^2} \left[\frac{\partial}{\partial t} \nabla^2 p' + \frac{1}{\rho_0 f_0} J(p', \nabla^2 p') + \beta_0 \frac{\partial p'}{\partial x} \right]$$

$$\frac{\partial \rho'}{\partial t} + u \frac{\partial \rho'}{\partial x} + v \frac{\partial \rho'}{\partial y} + w \frac{\partial \bar{\rho}}{\partial z} = 0$$

$$u_g = -\frac{1}{f_0 \rho_0} \frac{\partial p'}{\partial y}$$

$$v_g = +\frac{1}{f_0 \rho_0} \frac{\partial p'}{\partial x}$$

Plug u_g and v_g into the density equation:

$$\frac{\partial \rho'}{\partial t} + \frac{1}{\rho_0 f_0} J(p', \rho') - \frac{\rho_0 N^2}{g} w = 0$$

Divide the equation by $\frac{N^2}{g}$, and take the z-derivative:

$$0 = -\frac{\partial p'}{\partial z} - \rho' g$$

$$\begin{split} \frac{\partial}{\partial t \partial z} \left(\frac{g}{N^2} \rho' \right) + \frac{1}{\rho_0 f_0} \left[\frac{\partial p'}{\partial x} \frac{\partial}{\partial z} \left(\frac{g}{N^2} \frac{\partial \rho'}{\partial y} \right) - \frac{\partial p'}{\partial y} \frac{\partial}{\partial z} \left(\frac{g}{N^2} \frac{\partial \rho'}{\partial x} \right) \right] - \rho_0 \frac{\partial w}{\partial z} = 0 \\ - \frac{\partial}{\partial t \partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) + \frac{1}{\rho_0 f_0} \left\{ \frac{\partial p'}{\partial x} \frac{\partial}{\partial y} \left[\frac{\partial}{\partial z} \left(-\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right] - \frac{\partial p'}{\partial y} \frac{\partial}{\partial x} \left[\frac{\partial}{\partial z} \left(-\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right] \right\} - \rho_0 \frac{\partial w}{\partial z} = 0 \\ - \frac{\partial}{\partial t} \left[\frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right] - \frac{1}{\rho_0 f_0} J \left(p', \frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right) - \rho_0 \frac{\partial w}{\partial z} = 0 \\ \frac{\partial w}{\partial z} = -\frac{1}{\rho_0} \frac{\partial}{\partial t} \left[\frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right] - \frac{1}{\rho_0^2 f_0} J \left(p', \frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right) \end{split}$$

$$\frac{\partial w}{\partial z} = -\frac{1}{\rho_0} \frac{\partial}{\partial t} \left[\frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right] - \frac{1}{\rho_0^2 f_0} J \left(p', \frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right)$$

From the derivations based on the momentum equations:

$$\begin{split} \frac{\partial w}{\partial z} &= \frac{1}{\rho_0 f_0^{\ 2}} \left[\frac{\partial}{\partial t} \nabla^2 p' + \frac{1}{\rho_0 f_0} J(p', \nabla^2 p') + \beta_0 \frac{\partial p'}{\partial x} \right] \\ \frac{\partial}{\partial t} \left[\nabla^2 p' + \frac{\partial}{\partial z} \left(\frac{f_0^{\ 2}}{N^2} \frac{\partial p'}{\partial z} \right) \right] + \frac{1}{\rho_0 f_0} J\left(p', \nabla^2 p' + \frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial p'}{\partial z} \right) \right) + \beta_0 \frac{\partial p'}{\partial x} = 0 \end{split}$$

The geostrophic flows have streamfunction ψ :

$$u_{g} = -\frac{1}{\rho_{0}f_{0}} \frac{\partial p'}{\partial y} = -\frac{\partial \psi}{\partial y}$$

$$v_{g} = \frac{1}{\rho_{0}f_{0}} \frac{\partial p'}{\partial x} = \frac{\partial \psi}{\partial x}$$

$$+\beta_{0}y$$

$$\rho_{0}f_{0} \frac{\partial}{\partial t} \left[\nabla^{2}\psi + \frac{\partial}{\partial z} \left(\frac{f_{0}^{2}}{N^{2}} \frac{\partial \psi}{\partial z} \right) \right] + \rho_{0}f_{0}J \left(\psi, \nabla^{2}\psi + \frac{\partial}{\partial z} \left(\frac{f_{0}^{2}}{N^{2}} \frac{\partial \psi}{\partial z} \right) \right) + \rho_{0}f_{0}\beta_{0} \frac{\partial \psi}{\partial x} = 0$$

$$\frac{\partial}{\partial t} \left[\nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) + \beta_0 y \right] + J \left(\psi, \nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) \right) + J(\psi, \beta_0 y) = 0$$

$$\zeta \qquad \text{planetary vorticity}$$

$$q = \nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) + \beta_0 y \qquad \text{potential vorticity}$$

$$\frac{\partial q}{\partial t} + J(\psi, q) = 0$$

$$\frac{\partial q}{\partial t} + \frac{\partial \psi}{\partial x} \frac{\partial q}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial q}{\partial x} = 0$$

$$\frac{\partial q}{\partial t} + u_g \frac{\partial q}{\partial x} + v_g \frac{\partial q}{\partial y} = 0$$

$$\frac{dq}{dt} = 0$$

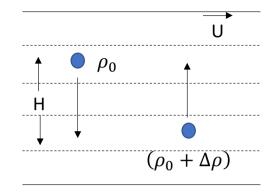
 $\frac{dq}{dt} = 0$ potential vorticity conservation

The importance of stratification: the Froude number

For per unit volume,

Potential energy change:

$$\Delta PE = (\rho_0 + \Delta \rho)gH - \rho_0 gH = \Delta \rho gH$$



Kinetic energy:

$$KE = \frac{1}{2}\rho_0 U^2 + \frac{1}{2}(\rho_0 + \Delta \rho) U^2 \approx \rho_0 U^2$$

$$\sigma = \frac{KE}{\Delta PE} = \frac{\rho_0 U^2}{\Delta \rho g H} \sim \frac{U^2}{N^2 H^2}$$
 Froude number: $Fr = \frac{U}{NH}$

- PE change consumes a small portion of the KE of the system, so it takes $\sigma > 1$, little cost to break stratification, stratification is unimportant
- PE change consumes all KE of the system, or KE is not sufficient to supply $\sigma \leq 1$, ΔPE , stratification cannot be broken and is important

relative vorticity planetary vorticity
$$q = \nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) + \beta_0 y$$
$$- \frac{g}{\rho_0 f_0} \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \rho' \right) \quad \text{vertical stretching}$$

$$0 = -\frac{\partial p'}{\partial z} - \rho' g \qquad \qquad p' = \rho_0 f_0 \psi \qquad \qquad \frac{\partial \psi}{\partial z} = -\frac{g}{\rho_0 f_0} \rho'$$

relative vorticity:
$$\frac{U}{L}$$
 vertical stretching: $\frac{f_0^2 UL}{N^2 H^2}$

$$\frac{\frac{U}{L}}{\frac{f_0^2 U L}{N^2 H^2}} = \frac{N^2 H^2}{f_0^2 L^2} = \frac{\frac{U^2}{f_0^2 L^2}}{\frac{U^2}{N^2 H^2}} = (\frac{R_0}{Fr})^2 \qquad Bu: \text{Burger number}$$

Bu < 1, rotation is more important, vertical stretching dominates the PV Bu > 1, stratification is more important, relative vorticity dominates the PV

Burger number:

$$Bu = (\frac{R_0}{Fr})^2$$
 a measure of relative importance of rotation and stratification
$$= \frac{U^2}{f^2 L^2} / \frac{U^2}{N^2 H^2}$$

$$= \frac{N^2 H^2}{f^2 L^2} = \frac{g' H}{f^2 L^2} = \frac{R^2}{L^2}$$

$$N^2 = -\frac{g}{\rho_0} \frac{d\rho}{dz} \simeq \frac{g}{\rho_0} \frac{\Delta \rho}{H} = \frac{g'}{H}$$

$$L < R$$
, $Bu > 1$, $Fr < R_0$, motion is more affected by stratification

$$L > R$$
, $Bu < 1$, $R_0 < Fr$, motion is more affected by rotation

$$Fr^2 = \frac{U^2}{N^2 H^2} = \frac{U^2}{g'H} \qquad Fr = \frac{U}{\sqrt{g'H}}$$

 $\sqrt{g'H}$: internal gravity wave speed

The final solutions:

$$u_{g} = -\frac{\partial \psi}{\partial y}$$

$$v_{g} = +\frac{\partial \psi}{\partial x}$$

$$u_{a} = -\frac{1}{f_{0}} \frac{\partial^{2} \psi}{\partial t \partial x} - \frac{1}{f_{0}} J\left(\psi, \frac{\partial \psi}{\partial x}\right) + \frac{\beta_{0}}{f_{0}} y \frac{\partial \psi}{\partial y}$$

$$v_{a} = -\frac{1}{f_{0}} \frac{\partial^{2} \psi}{\partial t \partial y} - \frac{1}{f_{0}} J\left(\psi, \frac{\partial \psi}{\partial y}\right) - \frac{\beta_{0}}{f_{0}} y \frac{\partial \psi}{\partial x}$$

$$w = -\frac{f_{0}}{N^{2}} \left[\frac{\partial^{2} \psi}{\partial t \partial z} + J\left(\psi, \frac{\partial \psi}{\partial z}\right)\right]$$

$$p' = \rho_{0} f_{0} \psi$$

$$\rho' = -\frac{\rho_{0} f_{0}}{g} \frac{\partial \psi}{\partial z}.$$

Energetics of quasi-geostrophic dynamics in stratified fluids

$$\frac{dq}{dt} = 0$$

$$\frac{d}{dt} (\nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) + \beta_0 y) = 0$$

For f-plane:

$$\frac{d}{dt}\nabla^2\psi + \frac{d}{dt}\left[\frac{\partial}{\partial z}\left(\frac{f_0^2}{N^2}\frac{\partial\psi}{\partial z}\right)\right] = 0$$

Multiply the equation by ψ and do a volume integration:

$$\frac{d}{dt}\iiint \frac{1}{2}\rho_0 |\nabla\psi|^2 dx dy dz + \frac{d}{dt}\iiint \frac{1}{2}\rho_0 \frac{f_0^2}{N^2} \left(\frac{\partial\psi}{\partial z}\right)^2 dx dy dz = 0$$

 KE PE

Available potential energy

$$PE(a) = \iint \left[\int_{0}^{H_{2}+a} \rho_{2}gzdz + \int_{H_{2}+a}^{H} \rho_{1}gzdz \right] dxdy$$

$$= \iint \left[\frac{1}{2} (\rho_{1} + \Delta \rho)g(H_{2} + a)^{2} + \frac{1}{2} \rho_{1}g[H^{2} - (H_{2} + a)^{2}] \right] dxdy$$

$$= \iint \frac{1}{2} \Delta \rho g H_{2}^{2} dxdy + \iint \frac{1}{2} \rho_{1}gH^{2} dxdy$$

$$= \iint \Delta \rho g H_{2}a dxdy + \iint \frac{1}{2} \Delta \rho g a^{2} dxdy$$

$$PE(a) - PE(a = 0) = \iint \frac{1}{2} \Delta \rho g a^{2} dxdy = \iint \frac{1}{2} \rho_{0}HN^{2}a^{2} dxdy$$

$$= \iiint \frac{1}{2} \rho_{0}N^{2}a^{2} dxdydz$$

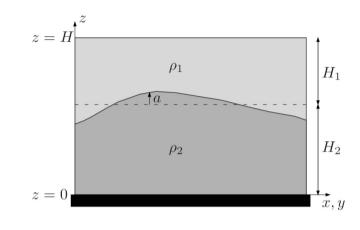
$$N^{2} = -\frac{g}{\rho_{0}} \frac{d\bar{\rho}}{dz} = \frac{g\Delta \rho}{\rho_{0}H}$$

$$\rho' = \bar{\rho}(z - a) - \bar{\rho}(z) = -a\frac{d\bar{\rho}}{dz} = \frac{\rho_0 N^2}{g} a$$

$$PE(a) - PE(a = 0) = \iiint \frac{1}{2} \rho_0 N^2 a^2 dx dy dz$$

$$= \iiint \frac{1}{2} \rho_0 N^2 (\frac{\rho' g}{\rho_0 N^2})^2 dx dy dz$$

$$= \iiint \frac{1}{2} \frac{1}{\rho_0 N^2} (-\rho_0 f_0 \frac{\partial \psi}{\partial z})^2 dx dy dz$$



$$\rho' = -\frac{\rho_0 f_0}{g} \frac{\partial \psi}{\partial z}$$

$$= \iiint \frac{1}{2} \rho_0 \frac{f_0^2}{N^2} (\frac{\partial \psi}{\partial z})^2 dx dy dz \quad \text{available potential energy (APE)}$$

APE: difference between existing PE and PE if the stratified system is not perturbed

$$\frac{d}{dt} \iiint \frac{1}{2} \rho_0 |\nabla \psi|^2 dx dy dz + \frac{d}{dt} \iiint \frac{1}{2} \rho_0 \frac{f_0^2}{N^2} \left(\frac{\partial \psi}{\partial z}\right)^2 dx dy dz = 0$$

Planetary waves in stratified fluids

For quasi-geostrophic model: $\frac{\partial q}{\partial t} + J(\psi, q) = 0$

$$\frac{\partial q}{\partial t} + J(\psi, q) = 0$$

$$q = \nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) + \beta_0 y$$

Linearization of the PV conservation equation, and take constant N^2 :

$$\frac{\partial}{\partial t} \left[\nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) + \beta_0 y \right] + J \left(\psi, \nabla^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) \right) + J(\psi, \beta_0 y) = 0$$

$$\frac{\partial}{\partial t} \left(\nabla^2 \psi + \frac{f_0^2}{N^2} \frac{\partial^2 \psi}{\partial z^2} \right) + \beta_0 \frac{\partial \psi}{\partial x} = 0$$

Assume the motion is bounded by a flat bottom and free surface, apply a wave solution of ψ as:

$$\varphi(x, y, z, t) = \phi(z)e^{i(kx+ly-\omega t)}$$

$$\frac{d^2\phi}{dz^2} - \frac{N^2}{f_0^2} \left(k^2 + l^2 + \frac{\beta_0 k}{\omega} \right) \phi = 0$$

Boundary conditions:

$$z = 0: w = -\frac{f_0}{N^2} \left[\frac{\partial^2 \psi}{\partial t \partial z} + J \left(\psi, \frac{\partial \psi}{\partial z} \right) \right] = 0 \frac{d\phi}{dz} = 0$$

$$z = h: p = \bar{p}(z) + p'(x, y, z, t)$$

$$p' = \rho_0 g \eta p' = \rho_0 f_0 \psi$$

$$w = \frac{\partial \eta}{\partial t} + u \frac{\partial \eta}{\partial x} + v \frac{\partial \eta}{\partial y} = \frac{1}{\rho_0 g} \frac{\partial p'}{\partial t} = \frac{f_0}{g} \frac{\partial \psi}{\partial t} = -\frac{f_0}{N^2} \frac{\partial^2 \psi}{\partial t \partial z}$$

$$\frac{\partial^2 \psi}{\partial t \partial z} + \frac{N^2}{g} \frac{\partial \psi}{\partial t} = 0$$

$$\frac{d\phi}{dz} + \frac{N^2}{g} \phi = 0$$

The solution for ϕ that satisfies both boundary conditions is:

$$\phi(z) = Acosmz$$

Substitution into the governing equation: $\frac{d^2\phi}{dz^2} - \frac{N^2}{f_0^2} \left(k^2 + l^2 + \frac{\beta_0 k}{\omega}\right) \phi = 0$

$$\omega = -\frac{\beta_0 k}{k^2 + l^2 + m^2 f_0^2 / N^2}$$

barotropic Rossby wave

$$\omega = -\frac{\beta_0 k}{k^2 + l^2 + m^2 f_0^2 / N^2}$$

$$\omega = -\beta_0 R^2 \frac{k}{1 + R^2 (k^2 + l^2)}$$

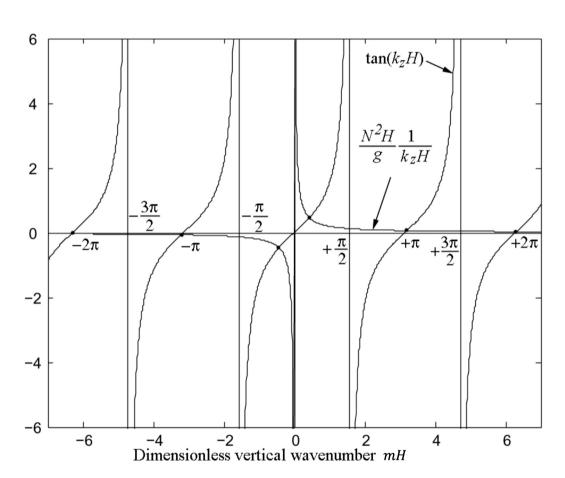
Substitution into the boundary condition at z = H: $\frac{d\phi}{dz} + \frac{N^2}{a} \phi = 0$

$$-AmsinmH + \frac{N^2}{g}AcosmH = 0$$

$$tanmH = \frac{N^2}{gm} = \frac{N^2H}{g} \frac{1}{mH}$$

$$tanmH = \frac{N^2H}{g} \frac{1}{mH}$$

$$\omega = -\frac{\beta_0 k}{k^2 + l^2 + m^2 f_0^2 / N^2}$$



$$\frac{N^2 H}{g} \sim \frac{g}{\rho_0} \frac{\Delta \rho}{H} \frac{H}{g} \sim \frac{\Delta \rho}{\rho_0}$$

For the first solution: mH is small

$$mH = \frac{N^2 H}{g} \frac{1}{mH}$$

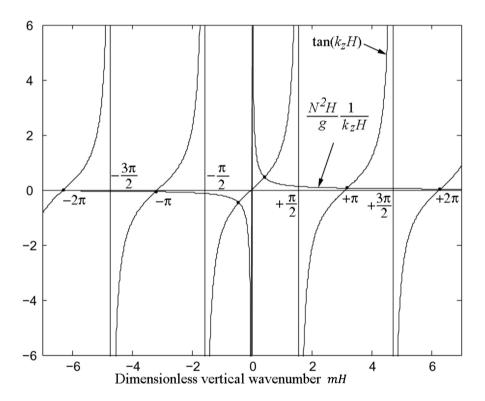
$$m = \frac{N}{\sqrt{gH}}$$

$$\omega = -\frac{\beta_0 k}{k^2 + l^2 + f_0^2/gH}$$

$$= -\beta_0 R^2 \frac{k}{1 + R^2(k^2 + l^2)}$$

barotropic Rossby wave

$$\omega = -\frac{\beta_0 k}{k^2 + l^2 + m^2 f_0^2 / N^2}$$



For larger solutions:

$$tanmH \sim 0$$

$$m_j = j \frac{\pi}{H}, j = 1, 2, 3, ...$$

baroclinic modes

$$\omega_j = -\frac{\beta_0 k}{k^2 + l^2 + (j\pi f_0/NH)^2}$$

$$c_x < 0$$
 westward propagation

$$R_j = \frac{1}{j} \frac{NH}{\pi f_0}$$

The maximum wave speed in the x direction is when

$$k^2 + l^2 \rightarrow 0$$
 long waves

$$|c_x|_{max} = \beta_0 R_j^2$$