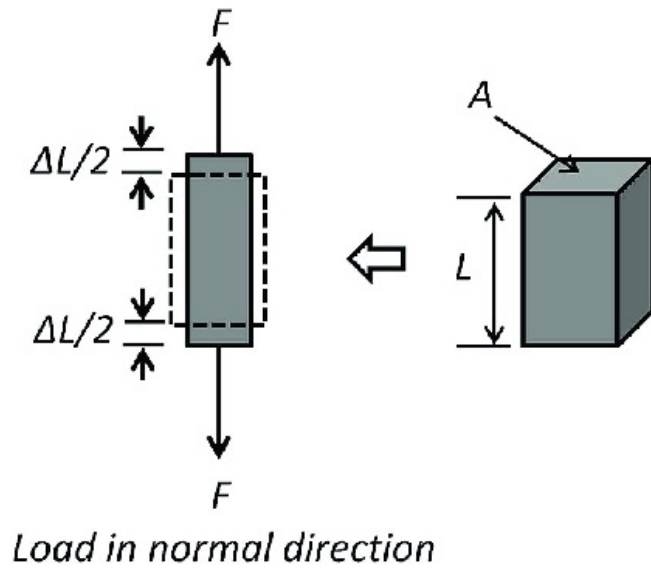


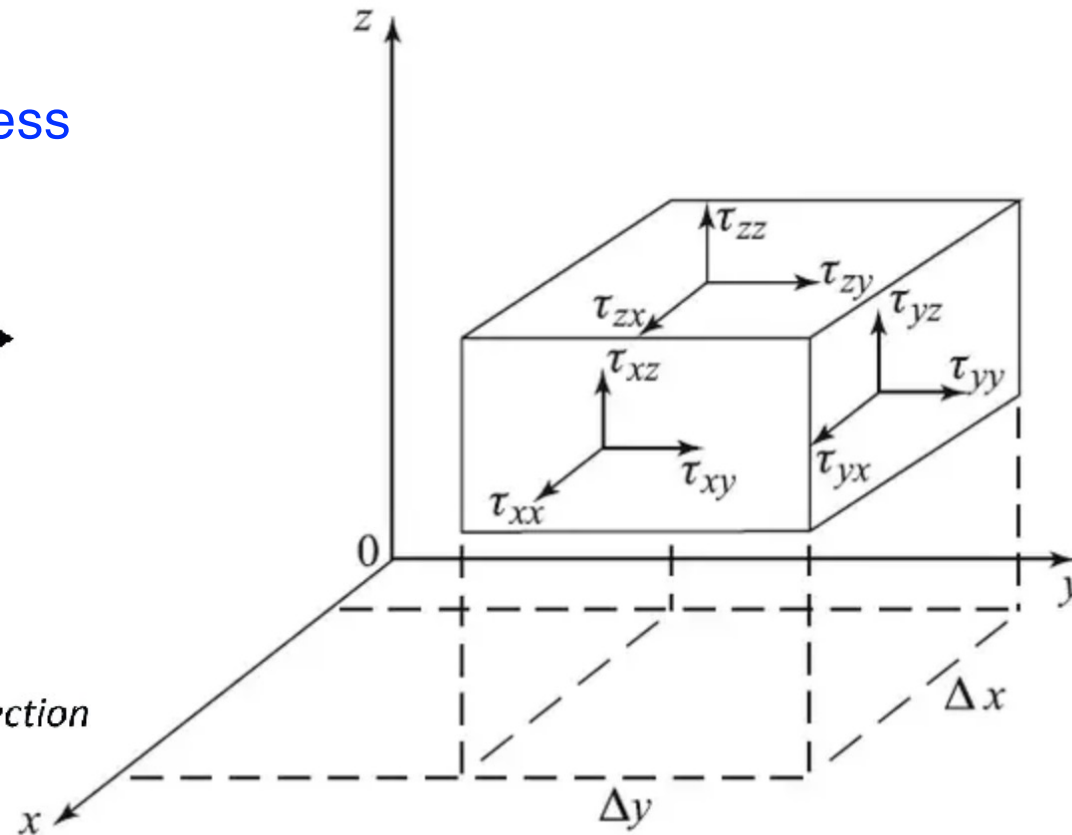
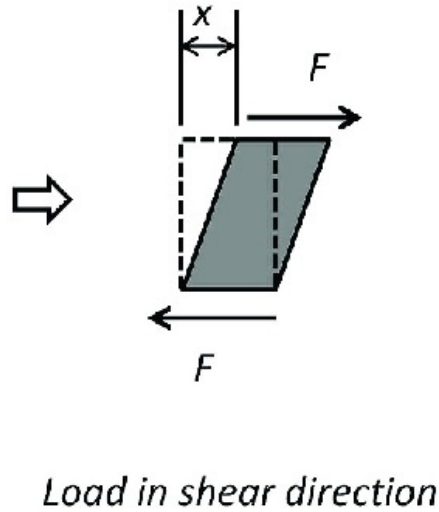
## Frictional force term

**Stress** – second-order tensor ( $\text{N m}^{-2}$ )

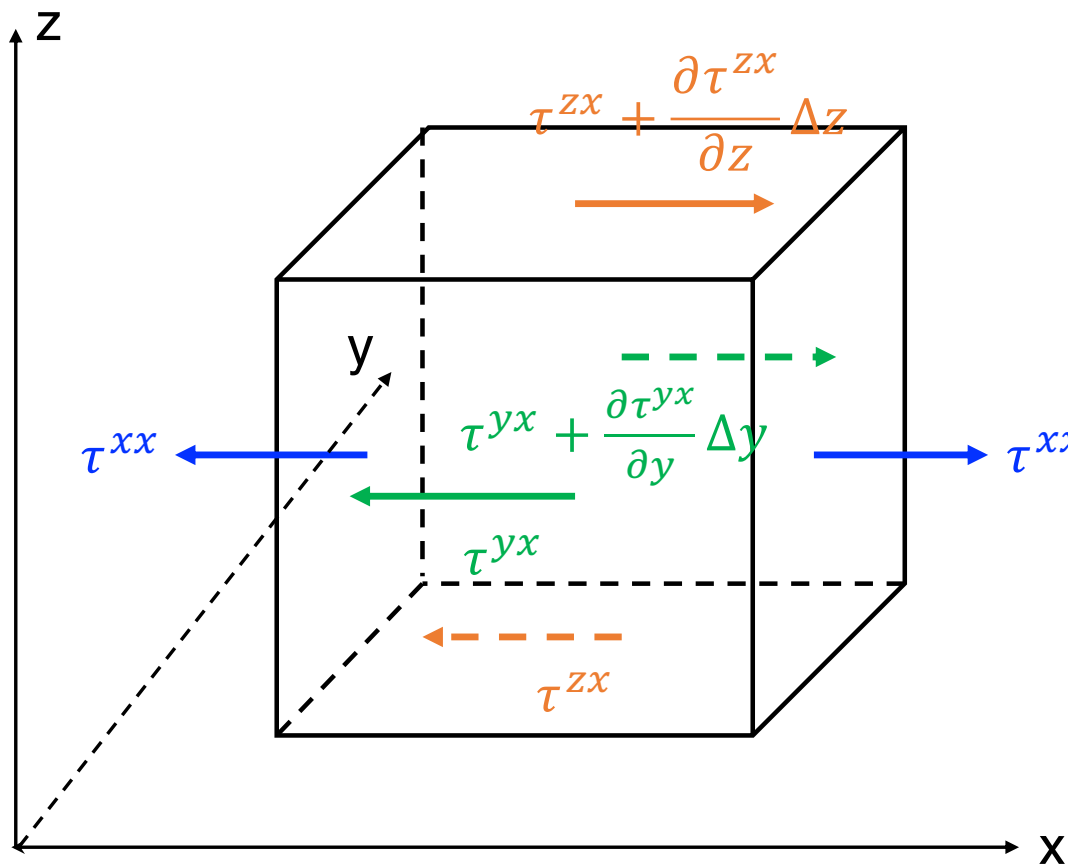
normal stress



shear stress



9 components



$$\begin{aligned}
 F^x = & \left( \tau^{xx} + \frac{\partial \tau^{xx}}{\partial x} \Delta x \right) \Delta y \Delta z - \tau^{xx} \Delta y \Delta z \\
 & + \left( \tau^{yx} + \frac{\partial \tau^{yx}}{\partial y} \Delta y \right) \Delta x \Delta z - \tau^{yx} \Delta x \Delta z \\
 & + \left( \tau^{zx} + \frac{\partial \tau^{zx}}{\partial z} \Delta z \right) \Delta x \Delta y - \tau^{zx} \Delta x \Delta y
 \end{aligned}$$

For per unit volume

$$F^x = \frac{\partial \tau^{xx}}{\partial x} + \frac{\partial \tau^{yx}}{\partial y} + \frac{\partial \tau^{zx}}{\partial z}$$

For Newtonian fluids, viscous stress is:  $\tau^{zx} = \mu \frac{\partial u}{\partial z}$

$\mu$ : dynamic viscosity coefficient

Assumption:  
 $\mu$  is constant

$$\begin{aligned} F^x &= \frac{\partial \tau^{xx}}{\partial x} + \frac{\partial \tau^{yx}}{\partial y} + \frac{\partial \tau^{zx}}{\partial z} \\ &= \frac{1}{\partial x} \left( \mu \frac{\partial u}{\partial x} \right) + \frac{1}{\partial y} \left( \mu \frac{\partial u}{\partial y} \right) + \frac{1}{\partial z} \left( \mu \frac{\partial u}{\partial z} \right) \\ &= \mu \nabla^2 u \end{aligned}$$

Laplacian operator

$$\nabla^2 = \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2} + \frac{\partial^2}{\partial z^2}$$

$$\begin{aligned} F^y &= \frac{\partial \tau^{xy}}{\partial x} + \frac{\partial \tau^{yy}}{\partial y} + \frac{\partial \tau^{zy}}{\partial z} \\ &= \frac{1}{\partial x} \left( \mu \frac{\partial v}{\partial x} \right) + \frac{1}{\partial y} \left( \mu \frac{\partial v}{\partial y} \right) + \frac{1}{\partial z} \left( \mu \frac{\partial v}{\partial z} \right) \\ &= \mu \nabla^2 v \end{aligned}$$

$$\begin{aligned} \frac{F^x}{\rho} &= \frac{\mu}{\rho} \nabla^2 u \\ &= \nu \nabla^2 u \end{aligned}$$

$\nu$  : kinematic viscosity coefficient

$$\begin{aligned} F^z &= \frac{\partial \tau^{xz}}{\partial x} + \frac{\partial \tau^{yz}}{\partial y} + \frac{\partial \tau^{zz}}{\partial z} \\ &= \frac{1}{\partial x} \left( \mu \frac{\partial w}{\partial x} \right) + \frac{1}{\partial y} \left( \mu \frac{\partial w}{\partial y} \right) + \frac{1}{\partial z} \left( \mu \frac{\partial w}{\partial z} \right) \\ &= \mu \nabla^2 w \end{aligned}$$

---

	$\mu$ (kg m <sup>-1</sup> s <sup>-1</sup> )	$\nu$ (m <sup>2</sup> s <sup>-1</sup> )
Air	$1.8 \times 10^{-5}$	$1.5 \times 10^{-5}$
Water	$1.1 \times 10^{-3}$	$1.1 \times 10^{-6}$
Mercury	$1.6 \times 10^{-3}$	$1.2 \times 10^{-7}$

---

## The momentum equations

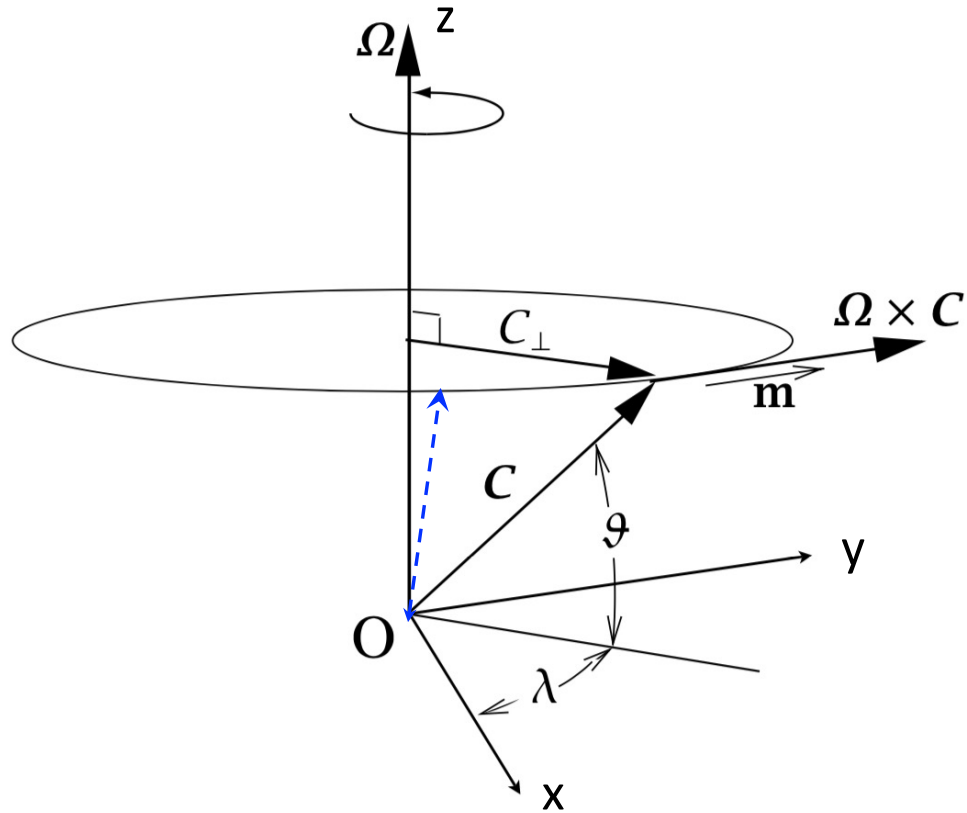
x direction: 
$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial x} + \underbrace{v \nabla^2 u}_{v \nabla^2 u \text{ (for constant } v)} + \underbrace{\frac{\partial}{\partial x} \left( v \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left( v \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left( v \frac{\partial u}{\partial z} \right)}_{\text{red terms}} + \dots$$

y direction: 
$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial y} + \underbrace{v \nabla^2 v}_{v \nabla^2 v} + \underbrace{\frac{\partial}{\partial x} \left( v \frac{\partial v}{\partial x} \right) + \frac{\partial}{\partial y} \left( v \frac{\partial v}{\partial y} \right) + \frac{\partial}{\partial z} \left( v \frac{\partial v}{\partial z} \right)}_{\text{red terms}} + \dots$$

z direction: 
$$\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + v \frac{\partial w}{\partial y} + w \frac{\partial w}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial z} + \underbrace{v \nabla^2 w}_{v \nabla^2 w} + \underbrace{\frac{\partial}{\partial x} \left( v \frac{\partial w}{\partial x} \right) + \frac{\partial}{\partial y} \left( v \frac{\partial w}{\partial y} \right) + \frac{\partial}{\partial z} \left( v \frac{\partial w}{\partial z} \right)}_{\text{red terms}} + \dots$$

Expressions for the force terms are given in the inertial frame.

So is the acceleration term.



**Fig. 2.1** A vector  $\mathbf{C}$  rotating at an angular velocity  $\boldsymbol{\Omega}$ . It appears to be a constant vector in the rotating frame, whereas in the inertial frame it evolves according to  $(d\mathbf{C}/dt)_I = \boldsymbol{\Omega} \times \mathbf{C}$ .

## Coriolis Force

The change in  $\mathbf{C}$  in  $\delta t$  with respect to the Inertial frame

$$\delta \mathbf{C} = |\mathbf{C}| \cos \vartheta \delta \lambda \mathbf{m},$$

$$\delta \lambda = |\boldsymbol{\Omega}| \delta t$$

Let  $\hat{\vartheta} = (\pi/2 - \vartheta)$

$$\delta \mathbf{C} = |\mathbf{C}| |\boldsymbol{\Omega}| \sin \hat{\vartheta} \mathbf{m} \delta t = \boldsymbol{\Omega} \times \mathbf{C} \delta t,$$

$$\left( \frac{d\mathbf{C}}{dt} \right)_I = \boldsymbol{\Omega} \times \mathbf{C}$$

## Non-constant vector in the rotating frame $\mathbf{B}$

The change of  $\mathbf{B}$  in the inertial frame:

$$(\delta \mathbf{B})_I = (\delta \mathbf{B})_R + (\delta \mathbf{B})_{rot}$$

With  $(\delta \mathbf{B})_{rot} = \boldsymbol{\Omega} \times \mathbf{B} \delta t$

$$\delta \mathbf{C} = \boldsymbol{\Omega} \times \mathbf{C} \delta t$$

$$\left( \frac{d\mathbf{C}}{dt} \right)_I = \boldsymbol{\Omega} \times \mathbf{C}$$

$$\left( \frac{d\mathbf{B}}{dt} \right)_I = \left( \frac{d\mathbf{B}}{dt} \right)_R + \boldsymbol{\Omega} \times \mathbf{B} \quad \mathbf{r} \text{ is a vector from the Earth center pointing to the Earth surface}$$

Substitute  $\mathbf{B}$  with  $\mathbf{v}_R$

$$\left( \frac{d\mathbf{r}}{dt} \right)_I = \left( \frac{d\mathbf{r}}{dt} \right)_R + \boldsymbol{\Omega} \times \mathbf{r}$$

$$\left( \frac{d\mathbf{v}_I}{dt} \right)_I = \left( \frac{d\mathbf{v}_R}{dt} \right)_R + \boldsymbol{\Omega} \times \mathbf{v}_R + \cancel{\frac{d\boldsymbol{\Omega}}{dt} \times \mathbf{r}} + \boldsymbol{\Omega} \times \left( \frac{d\mathbf{r}}{dt} \right)_I$$

$\mathbf{v}_I$

$$\mathbf{v}_I = \mathbf{v}_R + \boldsymbol{\Omega} \times \mathbf{r}$$

Centrifugal acceleration

$$\left( \frac{d\mathbf{v}_R}{dt} \right)_I = \left( \frac{d\mathbf{v}_R}{dt} \right)_R + \boldsymbol{\Omega} \times \mathbf{v}_R$$

$$\left( \frac{d\mathbf{v}_I}{dt} \right)_I = \left( \frac{d\mathbf{v}_R}{dt} \right)_R + \underline{2\boldsymbol{\Omega} \times \mathbf{v}_R} + \underline{\boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r})}$$

Coriolis acceleration

$$\left( \frac{d}{dt} (\mathbf{v}_I - \boldsymbol{\Omega} \times \mathbf{r}) \right)_I = \left( \frac{d\mathbf{v}_R}{dt} \right)_R + \boldsymbol{\Omega} \times \mathbf{v}_R$$

# The Coriolis acceleration

$$2\boldsymbol{\Omega} \times \mathbf{v}_R$$

$$\boldsymbol{\Omega} = \Omega \cos \varphi \mathbf{j} + \Omega \sin \varphi \mathbf{k}$$

$$\text{x: } 2\Omega \cos \varphi w - 2\Omega \sin \varphi v$$

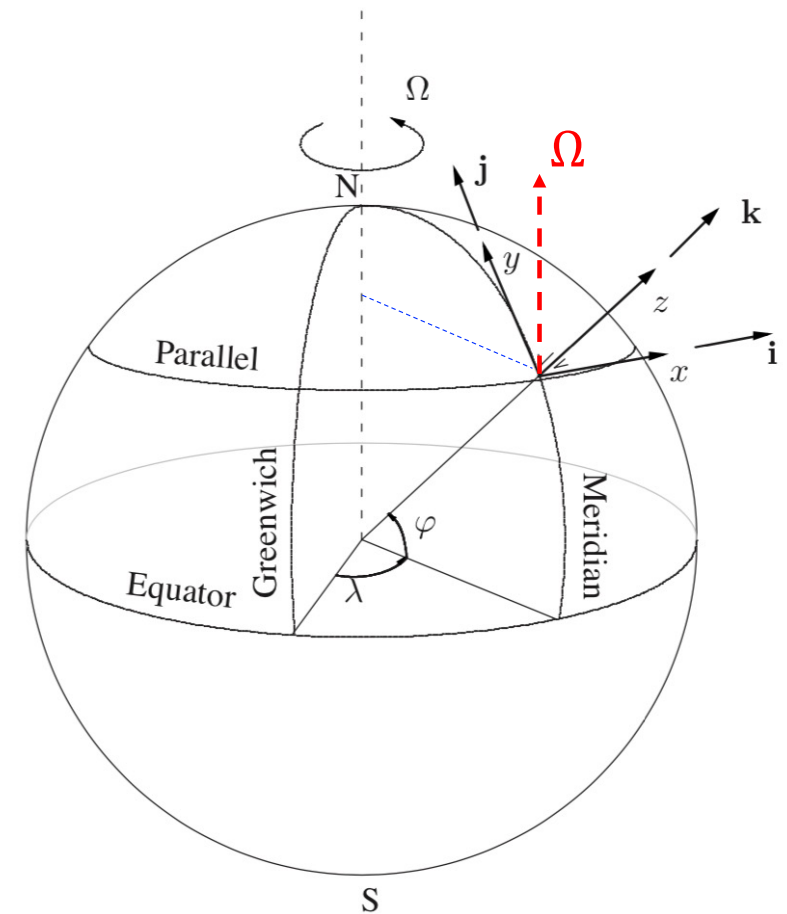
$$\text{y: } 2\Omega \sin \varphi u$$

$$\text{z: } -2\Omega \cos \varphi u.$$

$$f = 2\Omega \sin \varphi$$

$$f_* = 2\Omega \cos \varphi.$$

$f$ : Coriolis parameter



# The momentum equations

$$\left(\frac{d\mathbf{v}_I}{dt}\right)_I = \left(\frac{d\mathbf{v}_R}{dt}\right)_R + 2\boldsymbol{\Omega} \times \mathbf{v}_R + \boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r}) = \text{force terms}$$

Nonlinear advection term   Coriolis term   Pressure gradient term

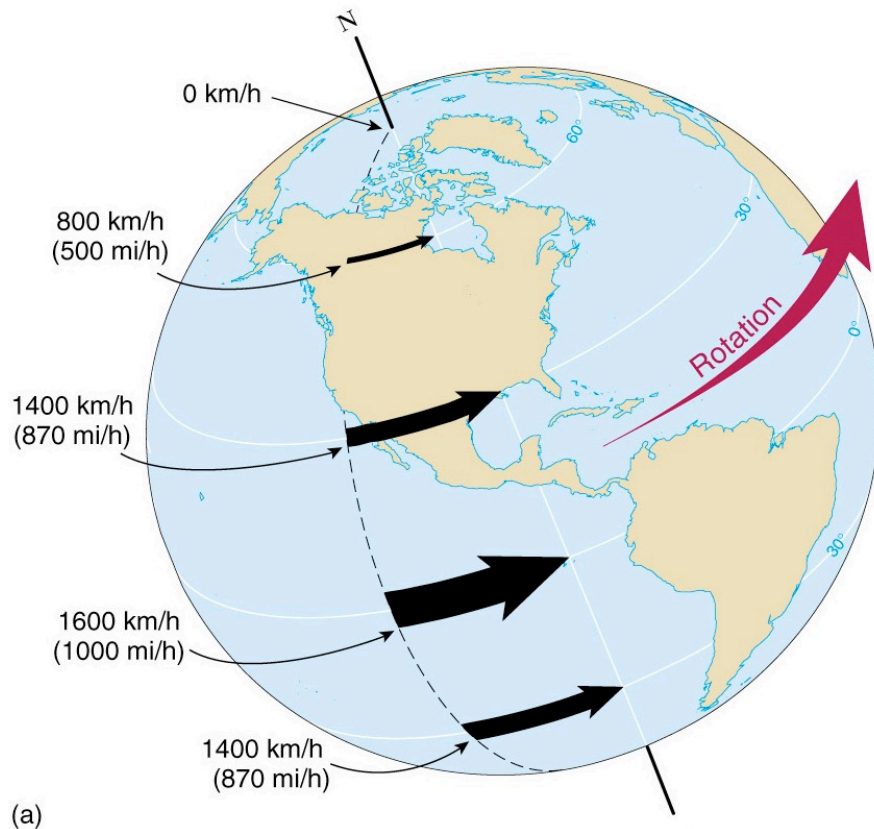
x direction:  $\underbrace{\frac{\partial u}{\partial t}}_{\text{Local acceleration}} + \underbrace{u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z}}_{\text{Nonlinear advection term}} - \underbrace{fv + f_* w}_{\text{Coriolis term}} = - \underbrace{\frac{1}{\rho} \frac{\partial p}{\partial x}}_{\text{Pressure gradient term}} + \underbrace{\nu \nabla^2 u}_{\text{Viscosity term}} + \dots$

y direction:  $\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} + \underbrace{fu}_{\text{Coriolis term}} = - \underbrace{\frac{1}{\rho} \frac{\partial p}{\partial y}}_{\text{Pressure gradient term}} + \underbrace{\nu \nabla^2 v}_{\text{Viscosity term}} + \dots$

z direction:  $\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + v \frac{\partial w}{\partial y} + w \frac{\partial w}{\partial z} - \underbrace{f_* u}_{\text{Coriolis term}} = - \underbrace{\frac{1}{\rho} \frac{\partial p}{\partial z}}_{\text{Pressure gradient term}} + \underbrace{\nu \nabla^2 w}_{\text{Viscosity term}} + \dots$

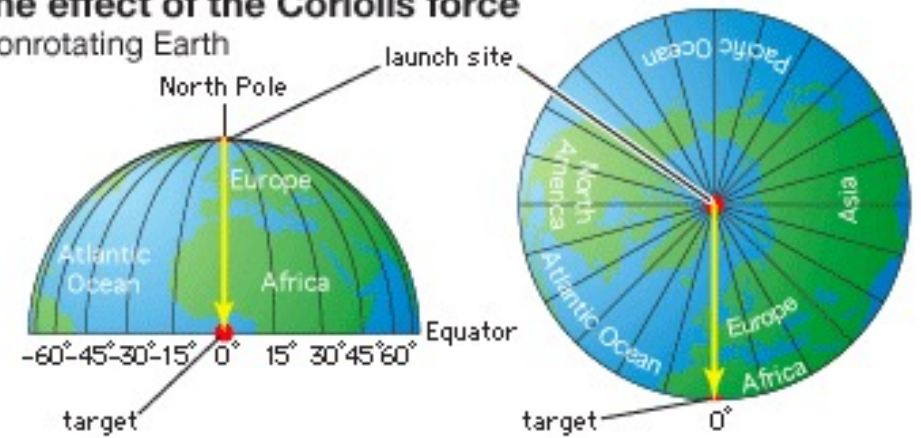


# The Coriolis Force

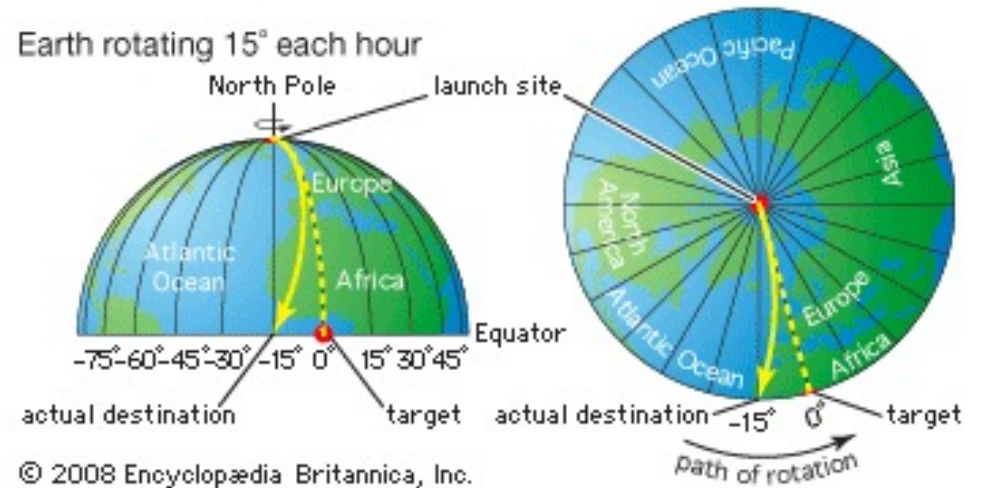


## The effect of the Coriolis force

### Nonrotating Earth



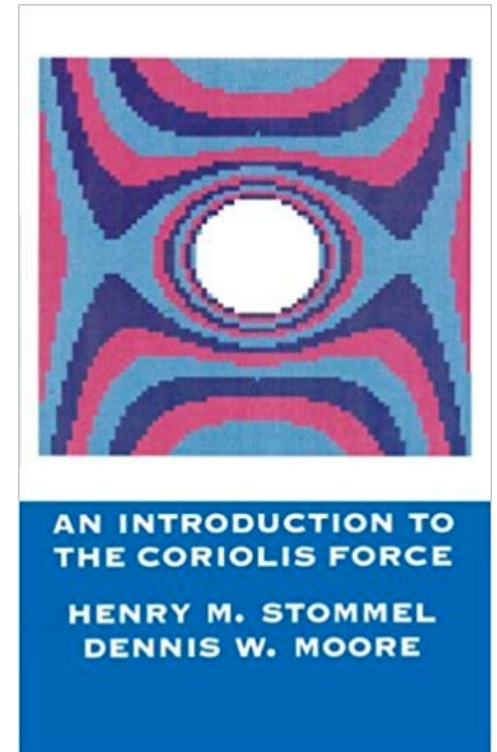
### Earth rotating 15° each hour



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# References about Coriolis force

- ***An introduction to the Coriolis Force***, H.M. Stommel and D.W. Moore, 1989
- Persson (2000). ***What is the Coriolis force?*** Weather, 55, 165–170.
- Persson (1998). ***How to understand the Coriolis force?*** Bulletin of the American Meteorological Society, 79(7), 1373–1385.



# $f$ -plane and $\beta$ -plane

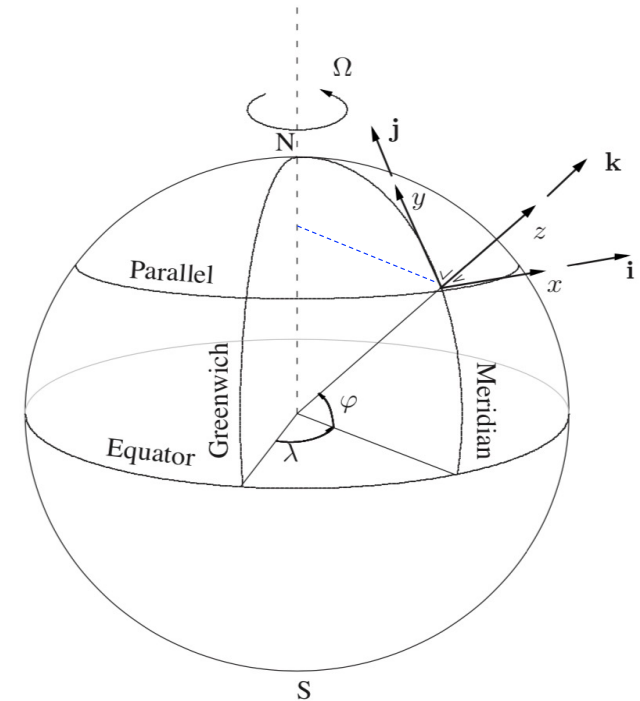
$$f = 2\Omega \sin \varphi = \frac{2\Omega(\sin \varphi_0 + \cos \varphi_0(\varphi - \varphi_0))}{f_0}$$

If  $\varphi - \varphi_0$  is small:

$f$ -plane:  $f = f_0$  is a constant.

If  $\varphi - \varphi_0$  cannot be neglected:

$$\beta\text{-plane: } f = f_0 + 2\Omega \cos \varphi_0 \frac{\varphi - \varphi_0}{y/a} = f_0 + \frac{2\Omega \cos \varphi_0}{\beta} y$$



## Centrifugal force term

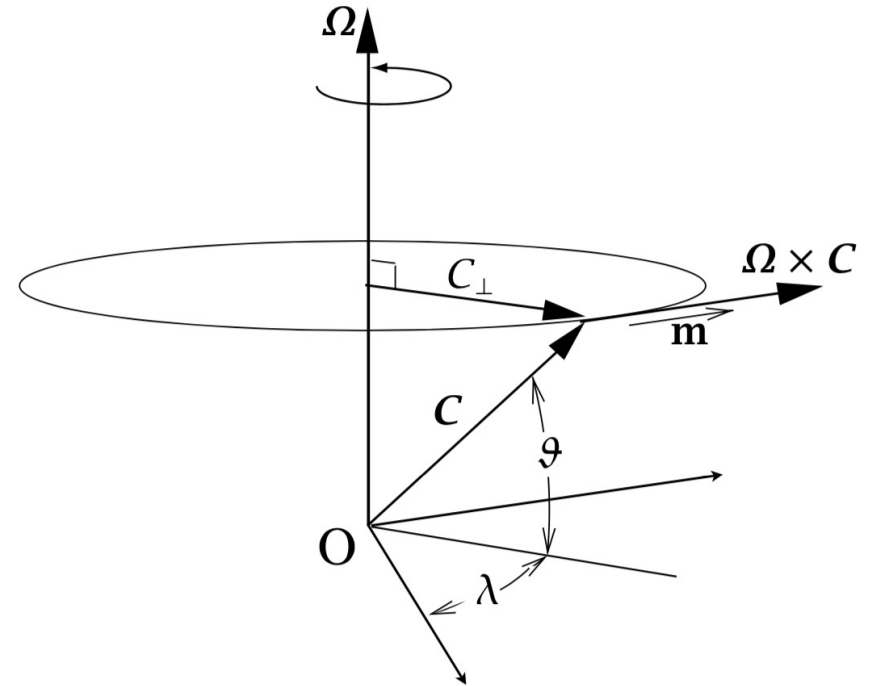
$$-\Omega \times (\Omega \times r)$$

## Substitute $\mathcal{C}$ with $r$ :

$$= -\boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r}_{\perp})$$

$$= -(\boldsymbol{\Omega} \cdot \mathbf{r}_\perp)\boldsymbol{\Omega} + (\boldsymbol{\Omega} \cdot \boldsymbol{\Omega})\mathbf{r}_\perp$$

$$= \Omega^2 r_{\perp}$$



# Gravity force term

**The Gravity Term in the Momentum Equation** The gravitational attraction of two masses  $M_1$  and  $m$  is:

$$\mathbf{F}_g = \frac{G M_1 m}{R^2}$$

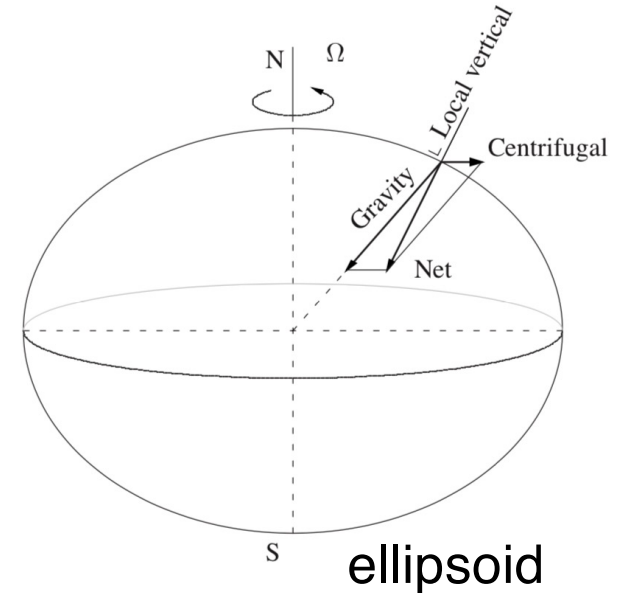
where  $R$  is the distance between the masses, and  $G$  is the gravitational constant. The vector force  $\mathbf{F}_g$  is along the line connecting the two masses.

The force per unit mass due to gravity is:

$$\frac{\mathbf{F}_g}{m} = \mathbf{g}_f = \frac{G M_E}{R^2} \quad (7.15)$$

where  $M_E$  is the mass of Earth. Adding the centrifugal acceleration to (7.15) gives gravity  $\mathbf{g}$  (Figure 7.5):

**effective gravity**  $\mathbf{g} = \mathbf{g}_f - \boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{R}) \quad (7.16)$



# The momentum equations

Nonlinear advection term
Coriolis term
Pressure gradient term

x direction:

$$\underbrace{\frac{\partial u}{\partial t}}_{\text{Local acceleration}} + \underbrace{u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z}}_{\text{Nonlinear advection term}} - \underbrace{fv + f_* w}_{\text{Coriolis term}} = - \underbrace{\frac{1}{\rho} \frac{\partial p}{\partial x}}_{\text{Pressure gradient term}} + \underbrace{\nu \nabla^2 u}_{\text{Viscosity term}}$$

y direction:

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} + \underbrace{fu}_{\text{Coriolis term}} = - \underbrace{\frac{1}{\rho} \frac{\partial p}{\partial y}}_{\text{Pressure gradient term}} + \underbrace{\nu \nabla^2 v}_{\text{Viscosity term}}$$

z direction:

$$\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + v \frac{\partial w}{\partial y} + w \frac{\partial w}{\partial z} + \underbrace{f_* u}_{\text{Coriolis term}} = - \underbrace{\frac{1}{\rho} \frac{\partial p}{\partial z}}_{\text{Pressure gradient term}} + \underbrace{\nu \nabla^2 w}_{\text{Viscosity term}} - \underbrace{g}_{\text{gravity term}}$$

# Continuity equation

$$\text{fluid loss} = \int_S \rho \mathbf{v} \cdot d\mathbf{S} = \int_V \nabla \cdot (\rho \mathbf{v}) dV, \quad \text{Divergence (Gaussian) theorem}$$

||

Vallis (Eq. 1.19)

$$-\frac{dM}{dt} = -\frac{d}{dt} \int_V \rho dV = -\int_V \frac{\partial \rho}{\partial t} dV$$

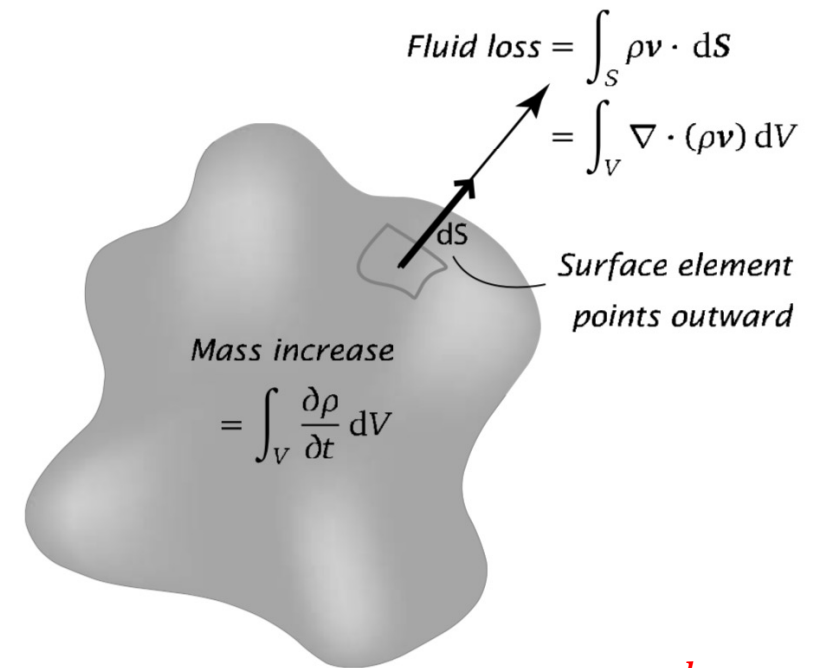
$$\int_V \left[ \frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{v}) \right] dV = 0$$

For any arbitrary volume  $V$ ,

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{v}) = 0 \quad \text{or} \quad \frac{d\rho}{dt} + \rho \nabla \cdot \mathbf{v} = 0$$

For incompressible fluids ( $\frac{d\rho}{dt} = 0$ )

$$\nabla \cdot \mathbf{v} = 0$$



# The thermodynamic equation

1<sup>st</sup> Law of Thermodynamics (energy conservation, and for per unit mass):

$$\frac{dI}{dt} = Q - W$$

$I$ : internal energy

$Q$ : rate of heat gain

$W$ : rate of work done by pressure force  
onto surrounding fluids

$C_v$ : specific heat capacity       $I = C_v T$

Fourier's Law of heat conduction:

$$Q = \frac{k_T}{\rho} \nabla^2 T$$

Rate of work done by pressure force:      specific volume

$$W = p \frac{d\alpha}{dt} \quad \boxed{\alpha = 1/\rho}$$



$$C_v \frac{dT}{dt} = \frac{k_T}{\rho} \nabla^2 T - p \frac{d\alpha}{dt}$$

$$= \frac{k_T}{\rho} \nabla^2 T + \frac{p}{\rho^2} \frac{d\rho}{dt}$$

$$\alpha = 1/\rho$$

From the continuity equation:

$$\frac{d\rho}{dt} = -\rho \nabla \cdot \mathbf{v}$$

$$C_v \frac{dT}{dt} = \frac{k_T}{\rho} \nabla^2 T - \frac{p}{\rho} \nabla \cdot \mathbf{v}$$

For incompressible fluids:

$K_T$ : thermal diffusivity coefficient

$$\frac{dT}{dt} = \frac{k_T}{\rho C_v} \nabla^2 T$$

$$\frac{dS}{dt} = K_S \nabla^2 S$$