# The equations of fluid motion with rotation

ATM2106

### Last time

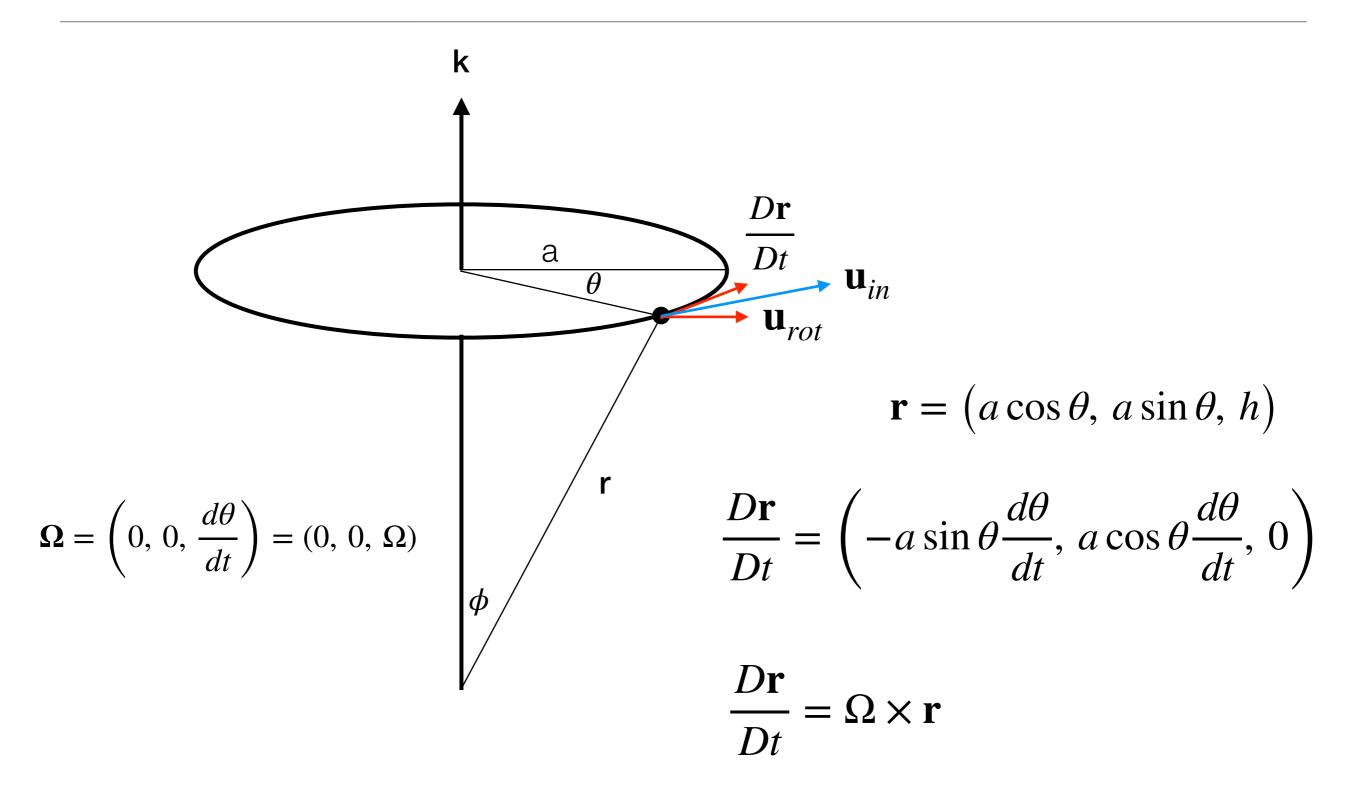
$$\frac{D\mathbf{u}}{Dt} + \frac{1}{\rho}\nabla p + g\hat{\mathbf{z}} = \mathcal{F}$$

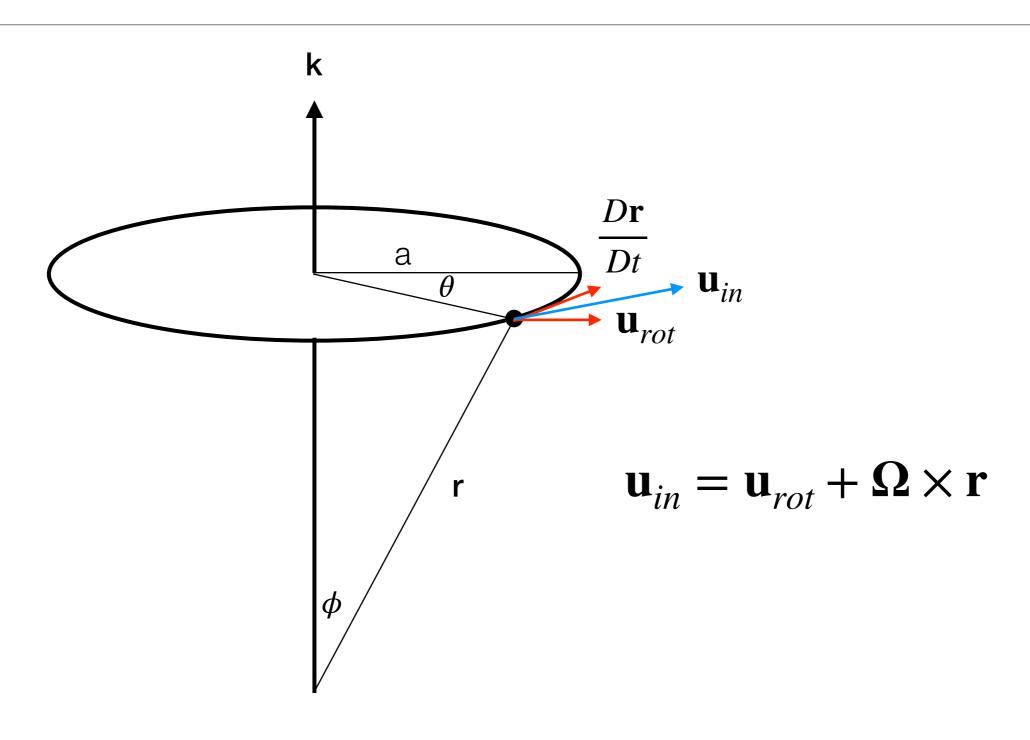
$$\frac{\partial \rho}{\partial t} + \frac{\partial}{\partial x}(\rho u) + \frac{\partial}{\partial y}(\rho v) + \frac{\partial}{\partial z}(\rho w) = 0$$

$$\frac{DQ}{Dt} = c_p \frac{DT}{Dt} - \frac{1}{\rho} \frac{Dp}{Dt}$$

## Today's topic

The equations of motion with rotation





A vector,  $\mathbf{A}$ , can be written with any three independent unit vectors.

$$\mathbf{A} = A_i \,\,\hat{\mathbf{i}} + A_j \,\,\hat{\mathbf{j}} + A_k \,\,\hat{\mathbf{k}} \longrightarrow \text{in the absolute frame}$$

$$= A_x \,\,\hat{\mathbf{x}} + A_y \,\,\hat{\mathbf{y}} + A_z \,\,\hat{\mathbf{z}} \longrightarrow \text{in the rotating frame}$$

Then, let's find out the Lagrangian differentiation of **A** with respect to the absolute (inertial) frame

$$\left(\frac{D\mathbf{A}}{Dt}\right)_{in} = \frac{D}{Dt} \left(A_i \,\hat{\mathbf{i}}\right) + \frac{D}{Dt} \left(A_j \,\hat{\mathbf{j}}\right) + \frac{D}{Dt} \left(A_k \,\hat{\mathbf{k}}\right)$$
$$= \hat{\mathbf{i}} \frac{DA_i}{Dt} + \hat{\mathbf{j}} \frac{DA_j}{Dt} + \hat{\mathbf{k}} \frac{DA_k}{Dt}$$

$$\left(\frac{D\mathbf{A}}{Dt}\right)_{in}$$
 can also be written as

$$\left(\frac{D\mathbf{A}}{Dt}\right)_{in} = \frac{D}{Dt} \left(A_x \,\hat{\mathbf{x}}\right) + \frac{D}{Dt} \left(A_y \,\hat{\mathbf{y}}\right) + \frac{D}{Dt} \left(A_z \,\hat{\mathbf{z}}\right) \qquad \frac{D\mathbf{r}}{Dt} = \mathbf{\Omega} \times \mathbf{r}$$

$$= \hat{\mathbf{x}} \frac{DA_x}{Dt} + \hat{\mathbf{y}} \frac{DA_y}{Dt} + \hat{\mathbf{z}} \frac{DA_z}{Dt} + A_x \frac{D\hat{\mathbf{x}}}{Dt} + A_y \frac{D\hat{\mathbf{y}}}{Dt} + A_z \frac{D\hat{\mathbf{z}}}{Dt}$$

$$= \hat{\mathbf{x}} \frac{DA_x}{Dt} + \hat{\mathbf{y}} \frac{DA_y}{Dt} + \hat{\mathbf{z}} \frac{DA_z}{Dt} + \mathbf{\Omega} \times \left(A_x \hat{\mathbf{x}} + A_y \hat{\mathbf{y}} + A_z \hat{\mathbf{z}}\right)$$

$$= \left(\frac{D\mathbf{A}}{Dt}\right)_{rot} + \mathbf{\Omega} \times \mathbf{A}$$

$$\left(\frac{D\mathbf{u}_{in}}{Dt}\right)_{in} = \left(\frac{D\mathbf{u}_{in}}{Dt}\right)_{rot} + \mathbf{\Omega} \times \mathbf{u}_{in} \qquad \mathbf{u}_{in} = \mathbf{u}_{rot} + \mathbf{\Omega} \times \mathbf{r}$$

$$= \left(\frac{D\left(\mathbf{u}_{rot} + \mathbf{\Omega} \times \mathbf{r}\right)}{Dt}\right)_{rot} + \mathbf{\Omega} \times \left(\mathbf{u}_{rot} + \mathbf{\Omega} \times \mathbf{r}\right)$$

$$= \left(\frac{D\mathbf{u}_{rot}}{Dt}\right)_{rot} + 2\mathbf{\Omega} \times \mathbf{u}_{rot} + \mathbf{\Omega} \times (\mathbf{\Omega} \times \mathbf{r})$$

Momentum equation we did last time can be applied to the parcel in the absolute frame

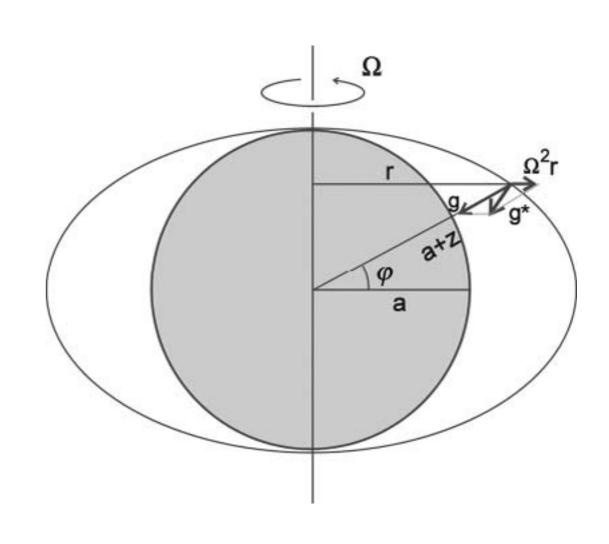
$$\left(\frac{D\mathbf{u}_{in}}{Dt}\right)_{in} + \frac{1}{\rho}\nabla p + g\hat{\mathbf{z}} = \mathcal{F}$$

$$\left(\frac{D\mathbf{u}_{rot}}{Dt}\right)_{rot} + \frac{1}{\rho}\nabla p + g\hat{\mathbf{z}} = -2\mathbf{\Omega} \times \mathbf{u}_{rot} - \mathbf{\Omega} \times (\mathbf{\Omega} \times \mathbf{r}) + \mathcal{F}$$

$$Centrifugal acceleration$$

$$Coriolis acceleration$$

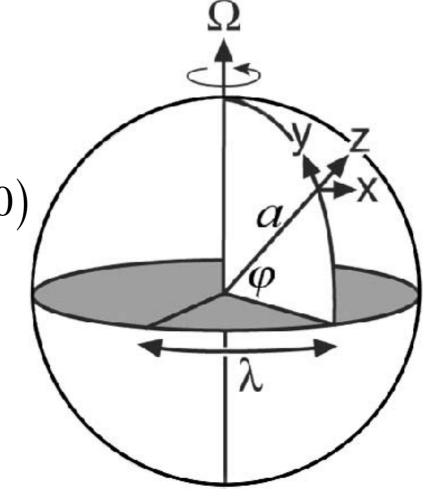
- Centrifugal acceleration:  $-\mathbf{\Omega} \times (\mathbf{\Omega} \times \mathbf{r})$ 
  - Modifies gravity acceleration
  - Effective gravity: g\*
  - $\Omega$  ~ 7.27 x 10<sup>-5</sup> s<sup>-1</sup> makes centrifugal acceleration small.
  - Geoid: The surface perpendicular to the effective gravity
  - We may use this surface to make centrifugal acceleration to disappear.



- Coriolis force:  $-2 \Omega \times \mathbf{u}$ 
  - $\Omega = (0, \Omega \cos \phi, \Omega \sin \phi)$
  - $\mathbf{\Omega} \times \mathbf{u} = (\Omega \cos \phi \ w \Omega \sin \phi \ v, \ \Omega \sin \phi \ u, \ -\Omega \cos \phi \ u)$
  - w is smaller than other terms.
  - $\Omega u$  is smaller than g.
  - $-2\mathbf{\Omega} \times \mathbf{u} \approx -(-2\mathbf{\Omega}\sin\phi \ v, \ 2\mathbf{\Omega}\sin\phi \ u, \ 0)$

$$=f\hat{\mathbf{z}}\times\mathbf{u}$$

$$(f = 2\Omega \sin \phi)$$



$$\frac{D\mathbf{u}}{Dt} + \frac{1}{\rho} \nabla p + g^* \hat{\mathbf{z}} + f \hat{\mathbf{z}} \times \mathbf{u} = \mathcal{F} \longrightarrow$$

- Hydrostatic approximation
- Vertical component of the frictional force is negligible compared with gravity

$$\frac{D\mathbf{u}}{Dt} + \frac{1}{\rho} \frac{\partial p}{\partial x} - fv = \mathcal{F}_x$$

$$\frac{D\mathbf{v}}{Dt} + \frac{1}{\rho} \frac{\partial p}{\partial y} + fu = \mathcal{F}_y$$

$$\frac{1}{\rho} \frac{\partial p}{\partial z} + g = 0$$
• Hydrostatic approximation

### Conservation of mass

The equation of continuity:

$$\frac{\partial \rho}{\partial t} + \frac{\partial}{\partial x}(\rho u) + \frac{\partial}{\partial y}(\rho v) + \frac{\partial}{\partial z}(\rho w) = 0$$

$$\frac{D\rho}{Dt} + \rho \frac{\partial u}{\partial x} + \rho \frac{\partial v}{\partial y} + \rho \frac{\partial w}{\partial z} = 0$$

$$\frac{D\rho}{Dt} + \rho \nabla \cdot \mathbf{u} = 0$$

https://youtu.be/RrWKSOvqV-0

### Conservation of mass:

- For compressible flow, the hydrostatic assumption allows us to write the unit volume as  $\delta x \delta y \delta p$
- Then the mass of the fluid parcel becomes

$$\delta M = \rho \, \delta x \, \delta y \, \delta z \, = -\frac{1}{g} \delta x \, \delta y \, \delta p$$

The mass is conserved in pressure coordinates, and

$$\nabla_p \cdot \mathbf{u}_p = \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial \omega}{\partial p} = 0$$

### Thermodynamic equation

- The first law of thermodynamics we dealt with for the dry adiabatic lapse rate is  $\delta Q = c_p dT \frac{dp}{\rho}$
- If we consider the first law of thermodynamics applied to a moving parcel of fluid,

$$\frac{DQ}{Dt} = c_p \frac{DT}{Dt} - \frac{1}{\rho} \frac{Dp}{Dt}$$

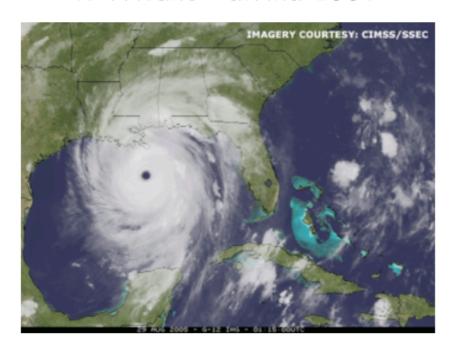
Diabatic heating rate (e.g. latent heating or cooling)

Temperature changes from the heat and/or expansion or compression

Extratropical cyclone 2010



Hurricane Katrina 2005



Ocean eddies in GFDL model

10°S

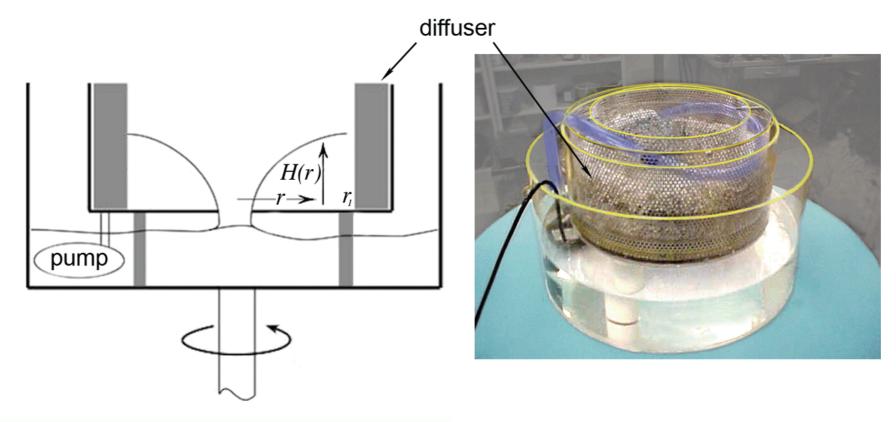
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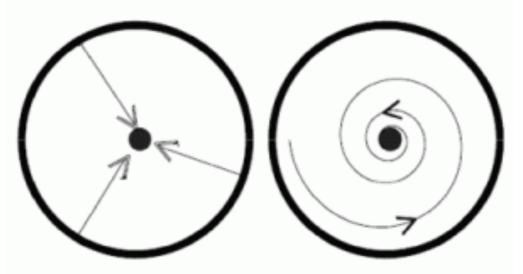
20°S

32.8
2.6
2.4
2.2
2.1
8
1.6
1.4
1.2
1.0
9
9.8
8.7
0.6
6.5
0.4
0.3
0.3
0.3
0.3
0.2
0.1
Tornado

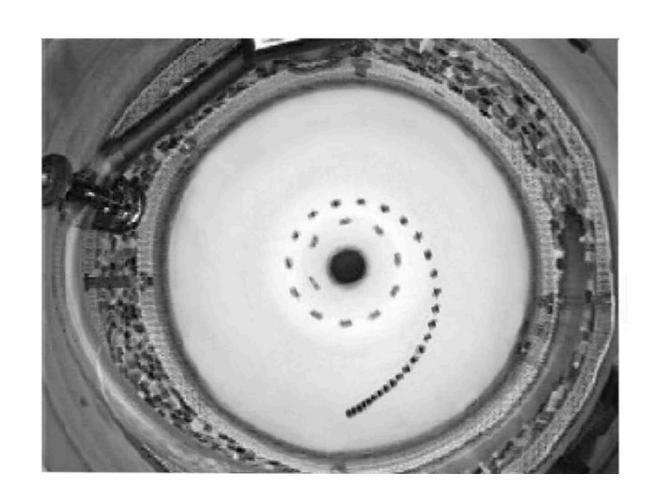


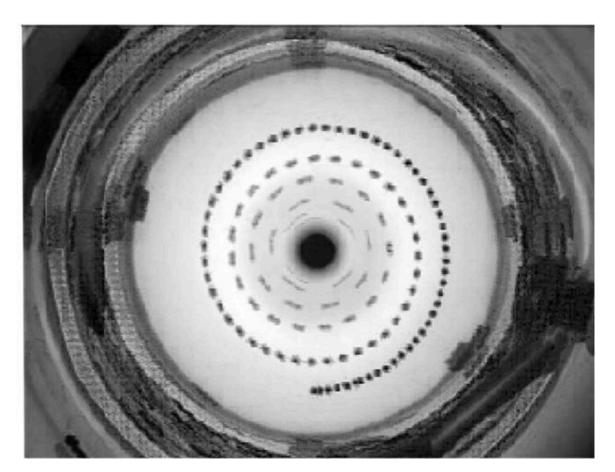
Balanced vortex





Flow patterns (left) in the absence of rotation and (right) when the apparatus is rotating in an anticlockwise direction.





Trajectories of particles in the radial inflow experiment viewed in the rotating frame. The positions are plotted every 1/30 s. On the left  $\Omega$ = 5 rpm (revolutions per minute). On the right  $\Omega$  = 10 rpm. Note how the pitch of the particle trajectory increases as  $\Omega$  increases, and how in both cases the speed of the particles increases as the radius decreases.

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#### **PHYSICS**

## The Bath-Tub Vortex in the Southern Hemisphere

It has long been thought that water draining from a tank would rotate counter-clockwise in the northern hemisphere and clockwise in the southern hemisphere, provided other influences were kept small compared with the influence of the rotation of the Earth. This idea has only recently been tested, by Shapiro in Watertown, Massachusetts, as part of a film on vorticity<sup>2-4</sup>, and later by Binnie in Cambridge, England<sup>1</sup>. Shapiro and Binnie both acquired confidence, after surmounting difficulties in their early experiments, that the counter-clockwise rotations observed in their later experiments were due to the rotation of the Earth.

Magnetic latitude (deg.)

Fig. 1. Plot of spread-F variation with sunspot number versus magnetic latitude

range of 40-120. It is found that the slopes of these linear portions vary widely from latitude to latitude, giving both positive and negative values. Positive slopes are obtained for the stations which show positive correlation of mean percentage occurrence of spread-F with sunspot numbers,

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