Scattering of surface waves by oceanic currents: the U2H map

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WISE SEMINAR, OCT. 2023

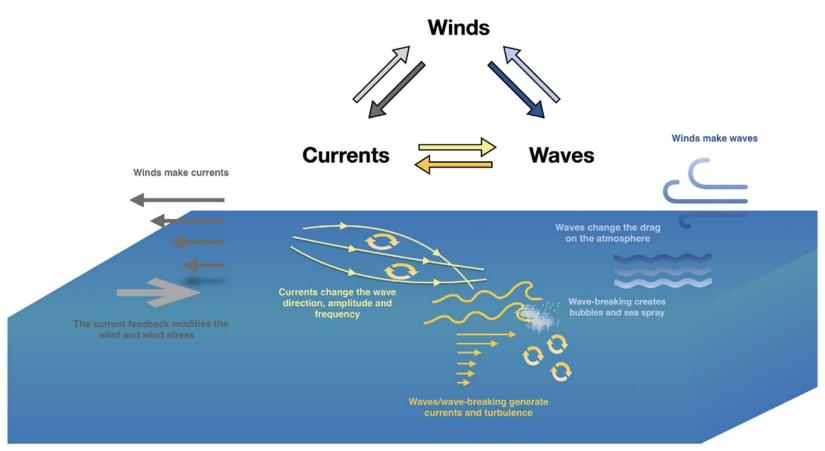




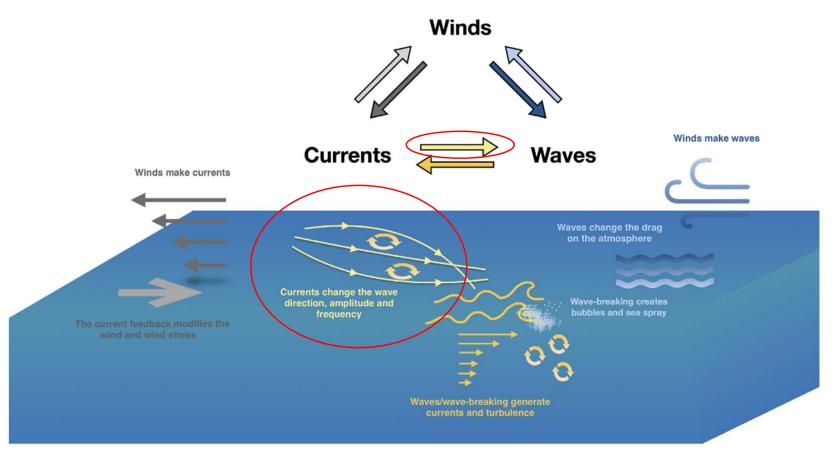








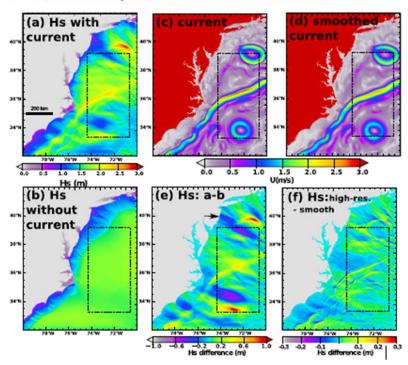
Villas Bôas and Pizzo (2021)



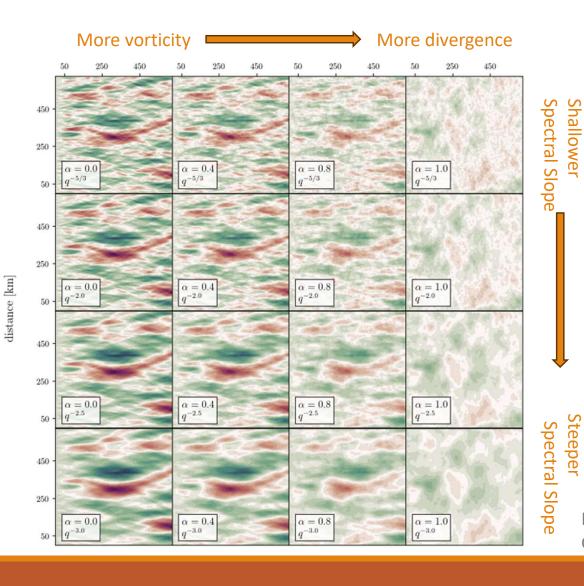
Villas Bôas and Pizzo (2021)

Small-scale open ocean currents have large effects on wind wave heights

Fabrice Ardhuin¹ (i), Sarah T. Gille² (i), Dimitris Menemenlis³ (i), Cesar B. Rocha² (i), Nicolas Rascle¹ (ii), Bertrand Chapron¹, Jonathan Gula¹ (ii), and Jeroen Molemaker¹



- Significant Wave Height (Hs)
 Hs = 4 × sea surface height rms
- Spatial variability of the Hs at scales between 10-100km is driven by currents



 Different impacts from rotational and divergent currents

$$\boldsymbol{U} = \boldsymbol{\nabla}\phi + \boldsymbol{\nabla}^{\perp}\psi\,,$$

- Rotational currents induce more changes than divergent currents, Hs more streaky
- Shallower spectral slope in KE
 -> finer structures in Hs

Bôas, A. B. V., Cornuelle, B. D., Mazloff, M. R., Gille, S. T., & Ardhuin, F. (2020).

The U2H map

- Map current velocity ("U") to Hs ("H")
- Assumptions:
 - 1. scale separation between waves and currents

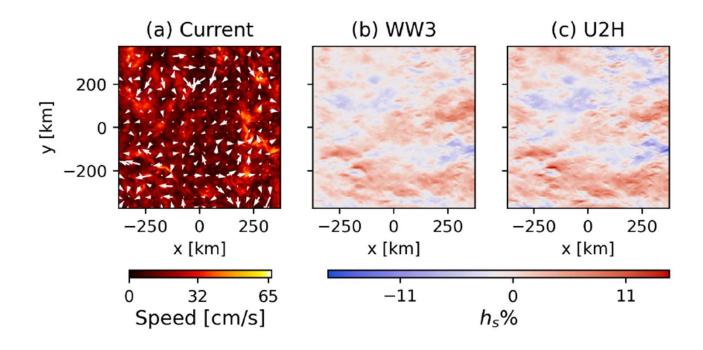
$$\partial_t \mathcal{A} + \nabla_k \omega \cdot \nabla_x \mathcal{A} - \nabla_x \omega \cdot \nabla_k \mathcal{A} = 0$$

where $\mathcal{A}=\mathcal{A}(\boldsymbol{x},\boldsymbol{k},t)$ wave-action density in position-wavevector space $\omega(\boldsymbol{x},\boldsymbol{k})=\sigma(k)+\boldsymbol{k}\cdot\boldsymbol{U}(\boldsymbol{x}), \,\, \text{deep-water surface gravity waves} \quad \sigma(k)=\sqrt{gk}$

2. weak current

$$\varepsilon = U/c \ll 1$$

It works well



Perturbation expansion

•
$$\partial_t \mathcal{A} + \nabla_k \omega \cdot \nabla_x \mathcal{A} - \nabla_x \omega \cdot \nabla_k \mathcal{A} = 0$$

- Small parameter $\varepsilon = U/c \ll 1$
- $\omega(\boldsymbol{x}, \boldsymbol{k}) = \sigma(k) + \varepsilon \boldsymbol{k} \cdot \boldsymbol{U}(\boldsymbol{x}).$ $\mathcal{A}(\boldsymbol{x}, \boldsymbol{k}, t) = \overline{A}(\boldsymbol{k}) + \varepsilon a(\boldsymbol{x}, \boldsymbol{k}, t) + O(\varepsilon^2).$
- $O(\varepsilon)$ terms:

Background, spatially uniform wave action

$$(\mu + c \cdot \nabla_x)a = (\nabla_k \bar{\mathcal{A}} \cdot \nabla_x)U \cdot k$$

Steady state; dissipation term to enforce causality

Linear relation

$$(\mu + c \cdot
abla_{m{x}})a = (
abla_{m{k}} ar{\mathcal{A}} \cdot
abla_{m{x}})U \cdot m{k}$$

action. Linear in **U**

Hs and wave action

- •Hs = $4 \times \text{sea}$ surface height rms
- Wave action × intrinsic frequency = KE+PE
- For deep water surface gravity waves, KE=PE

$$H_s(\boldsymbol{x},t) = 4 \left(g^{-1} \int \sigma(k) \mathcal{A}(\boldsymbol{x},\boldsymbol{k},t) \,\mathrm{d}\boldsymbol{k}\right)^{1/2}.$$

•Expand
$$H_s({m x}) = \bar{H}_s + arepsilon h_{
m s}({m x}) + O(arepsilon^2)$$

•Therefore
$$h_{\mathrm{s}}({m x}) = rac{8}{gar{H}_s} \int \sigma(k) a({m x},{m k}) \,\mathrm{d}{m k}$$
 .

hs (Changes in Hs) is linear in \mathcal{Q} . Therefore hs is linear in \boldsymbol{U} .

Linear relation

$$(\mu + \boldsymbol{c} \cdot \boldsymbol{\nabla}_{\boldsymbol{x}}) a = (\boldsymbol{\nabla}_{\boldsymbol{k}} \bar{\mathcal{A}} \cdot \boldsymbol{\nabla}_{\boldsymbol{x}}) \boldsymbol{U} \cdot \boldsymbol{k}$$
$$h_{s}(\boldsymbol{x}) = \frac{8}{g\bar{H}_{s}} \int \sigma(k) a(\boldsymbol{x}, \boldsymbol{k}) \, \mathrm{d}\boldsymbol{k}.$$

Solve in Fourier space

$$\hat{a}(\boldsymbol{q}, \boldsymbol{k}) \stackrel{\text{def}}{=} \int a(\boldsymbol{x}, \boldsymbol{k}) e^{-i\boldsymbol{q}\cdot\boldsymbol{x}} d\boldsymbol{x}$$

Spatial variations of:

wave phase

currents

$$\mathbf{k} = k \begin{pmatrix} \cos \theta \\ \sin \theta \end{pmatrix} \quad \mathbf{q} = q \begin{pmatrix} \cos \varphi \\ \sin \varphi \end{pmatrix}$$

Linear relation

$$(\mu + \boldsymbol{c} \cdot \boldsymbol{\nabla}_{\boldsymbol{x}}) a = (\boldsymbol{\nabla}_{\boldsymbol{k}} \bar{\mathcal{A}} \cdot \boldsymbol{\nabla}_{\boldsymbol{x}}) \boldsymbol{U} \cdot \boldsymbol{k}$$
$$h_{s}(\boldsymbol{x}) = \frac{8}{g\bar{H}_{s}} \int \sigma(k) a(\boldsymbol{x}, \boldsymbol{k}) d\boldsymbol{k}.$$

Solve in Fourier space

$$\hat{a}(\boldsymbol{q}, \boldsymbol{k}) \stackrel{\text{def}}{=} \int a(\boldsymbol{x}, \boldsymbol{k}) e^{-i\boldsymbol{q}\cdot\boldsymbol{x}} d\boldsymbol{x}$$

We get a computable form (after some maneuvers):

where
$$\hat{L}(q) = \hat{L}(q) \cdot \hat{U}(q)$$
, $\hat{L}(q) = \frac{16}{g\bar{H}_s} \left(e_q^{\perp} \sum_{n=-\infty}^{\infty} n(-\mathrm{i})^{|n|} 2\pi p_n \, \mathrm{e}^{n\mathrm{i}\varphi} - 2P \right)$ where $e_q^{\perp} \stackrel{\mathrm{def}}{=} (-\sin\varphi, \cos\varphi)$, $P \stackrel{\mathrm{def}}{=} \int \bar{\mathcal{A}}(k) k \, \mathrm{d}k$, $P = \oint \mathcal{P}(\theta) \begin{pmatrix} \cos\theta \\ \sin\theta \end{pmatrix} \mathrm{d}\theta$, $P(\theta) = \sum_{n=-\infty}^{\infty} p_n \, \mathrm{e}^{n\mathrm{i}\theta}$ $\hat{L}(q) = \hat{L}(\varphi)$

Spatial variations of:

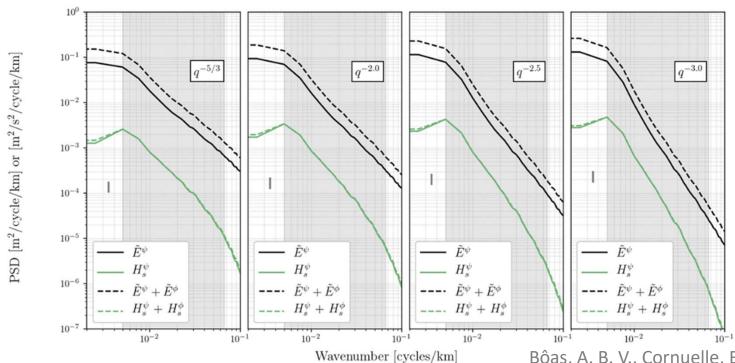
wave phase

currents

$$\mathbf{k} = k \begin{pmatrix} \cos \theta \\ \sin \theta \end{pmatrix} \quad \mathbf{q} = q \begin{pmatrix} \cos \varphi \\ \sin \varphi \end{pmatrix}$$

$$\hat{h}_s(\boldsymbol{q}) = \hat{\boldsymbol{L}}(\varphi) \cdot \hat{\boldsymbol{U}}(\boldsymbol{q}),$$

• Significant wave height should have similar spatial scales as currents

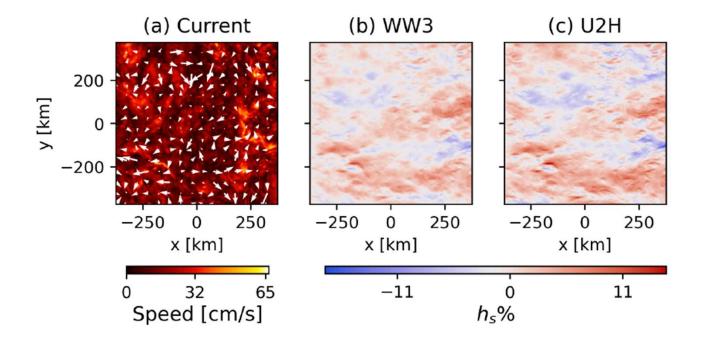


Bôas, A. B. V., Cornuelle, B. D., Mazloff, M. R., Gille, S. T., & Ardhuin, F. (2020).

$$\hat{h}_s(\boldsymbol{q}) = \hat{\boldsymbol{L}}(\varphi) \cdot \hat{\boldsymbol{U}}(\boldsymbol{q}),$$

 $\hat{m{L}}(arphi)$ is computable from $ar{\mathcal{A}}(m{k})$

Can compute significant wave height snapshots from current snapshots



Swells



Figure from Phillip Capper, Wikipedia

Long wavelengths (large group velocity)

$$\varepsilon = U/c \ll 1$$

Highly directional

How do Ocean Waves develop?

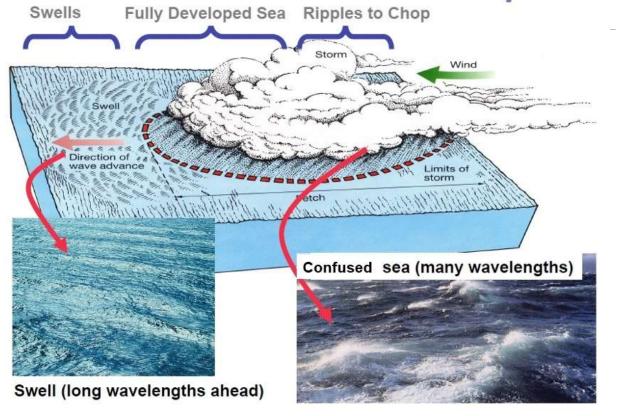


Figure from UBC EOSC 114 course

Swells



Figure from Phillip Capper, Wikipedia

Long wavelengths (large group velocity)

$$\varepsilon = U/c \ll 1$$

Highly directional

Another small parameter $\,\delta\ll 1\,$

$$\bar{\mathcal{A}}(k,\theta) = \delta^{-1}\bar{\mathcal{A}}(k,\Theta),$$

 $\Theta = \theta/\delta.$

U2H under the swell limit

Helmholtz decomposition

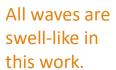
$$\boldsymbol{U} = \boldsymbol{\nabla}\phi + \boldsymbol{\nabla}^{\perp}\psi \,,$$

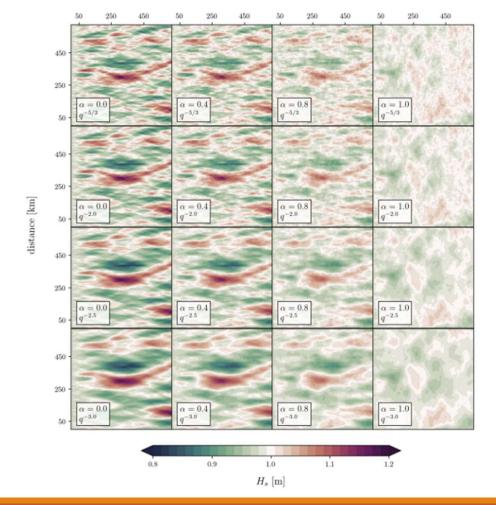
$$\hat{h}_s(\mathbf{q}) = iq\hat{L}_{\parallel}(\varphi)\hat{\phi}(\mathbf{q}) + iq\hat{L}_{\perp}(\varphi)\hat{\psi}(\mathbf{q}).$$

Under the swell limit $~\delta \ll 1~$, we find:

- $\hat{L}_{\perp} = O(\delta^{-2})\hat{L}_{\parallel}$
- → Rotational component of currents leave an asymptotically larger imprint on significant wave height than the divergent currents
- $\hat{m L}(arphi)$ dominant ($O(\delta^{-2})$) in narrow, $O(\delta)$ sectors around arphi= (swell propagation direction $\pm rac{\pi}{2}$)
 - → streaks: hs elongated in the direction of propagation of the waves







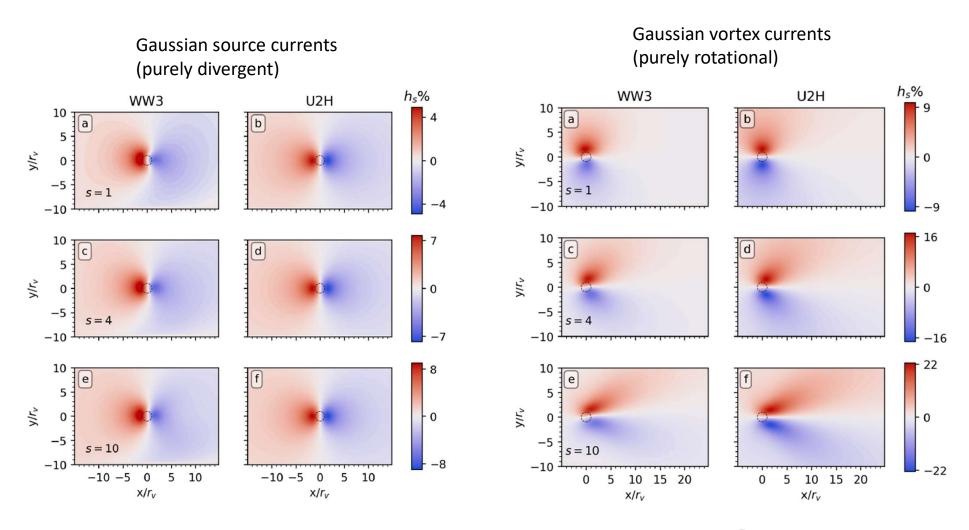
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Impact of divergent currents

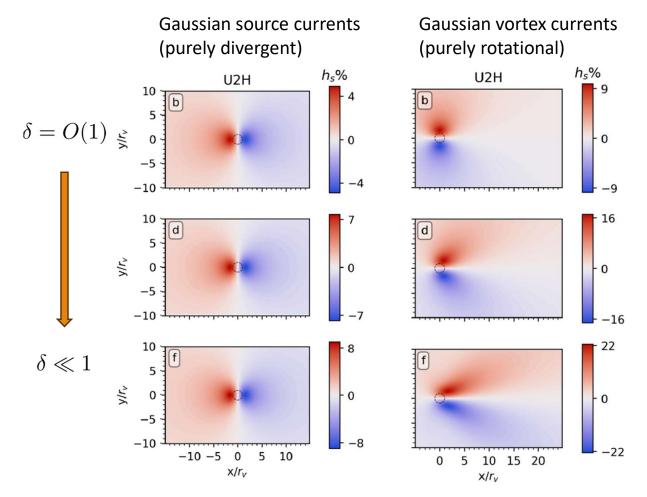
$$\boldsymbol{U} = \boldsymbol{\nabla}\phi + \boldsymbol{\nabla}^{\perp}\psi , \quad \hat{h}_s(\boldsymbol{q}) = iq\hat{L}_{\parallel}(\varphi)\hat{\phi}(\boldsymbol{q}) + iq\hat{L}_{\perp}(\varphi)\hat{\psi}(\boldsymbol{q}) .$$

$$\hat{L}_{\perp} = O(\delta^{-2})\hat{L}_{\parallel}$$

But for non-swells where $\,\delta\,$ is not small, $\,\hat{L}_{\parallel}\,$ can be comparable to $\,\hat{L}_{\perp}\,$



s is a parameter controlling the directional spreading; larger s correspond to smaller δ

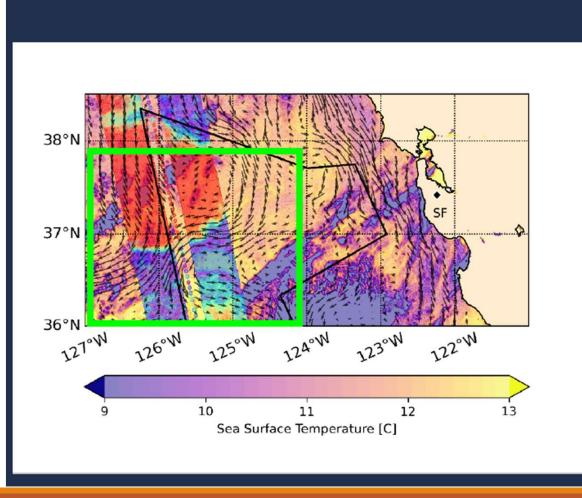


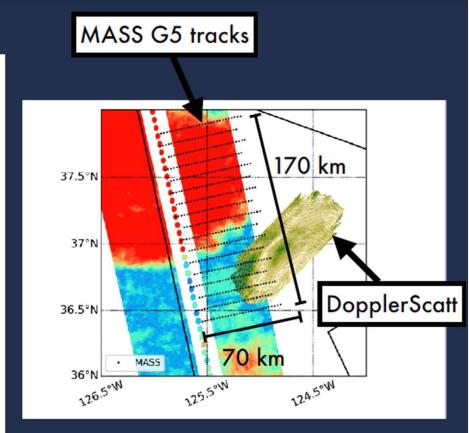
- Non-swells: divergent flows affect hs to a similar degree as rotational flow
- As δ decreases (swells), hs from divergent currents stop increasing at some point, but the maximum hs from
- rotational currents keep growing.
- Impacts of divergent flows are local.
- Rotational flows: streaks, nonlocal

Summary

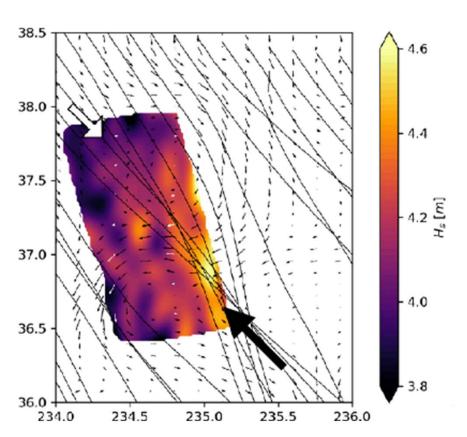
- Linear map from currents to significant wave height ("U2H")
- Assumptions: scale separation, weak current
- Significant wave height has similar spatial scales as currents
- For non-swells, both divergent and rotational currents can impact hs. Divergent currents' impacts are local.
- For swells,
 - 1). rotational currents has asymptotically larger impact.
 - 2). rotational currents' impacts are nonlocal.
 - 3). hs are elongated in the waves propagation direction.

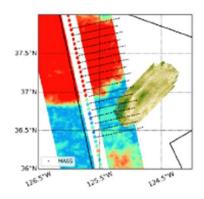
Overlap between S-MODE and SWOT calval campaign





Marechal, Villas Bôas, Lenain, and Pizzo (in prep).





Marechal, Villas Bôas, Lenain, and Pizzo (in prep).

Thank you!

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