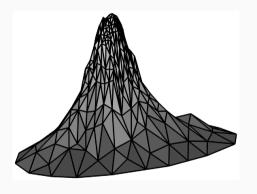
The Finite Volume Method

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Introduction

Motivation

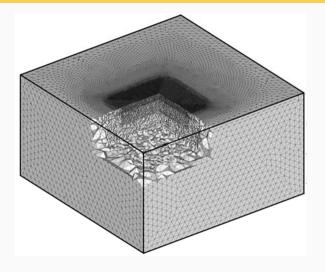


- Robust
- · Simple concept
- Irregular grids
- · Explicit method
- Find solution for strongly heterogeneous paramters and discontinuities

History

- First applications in **plasma physics** (Hermeline, 1993) and computational fluids (Versteeg and Malalasekera, 1995)
- Discrete version of the divergence theorem was used for seismic wave propagation (Dormy and Tarantola, 1995)
- Comparing and quantifying the accuracy of wave propagation of the finite volume method was done by Kaeser et al. 2001
- Leveque, 2001 presented the method as a natural consequence of conservation laws
- Dumbser et al. (2007) presented the arbitrary high-order scheme (ADER) for the finite-volume method

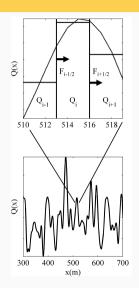
History



Finite Volume in a Nutshell

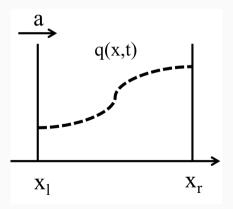
In its basic form it takes an entirely **local viewpoint** in the sense that the solution field q(x,t) is tracked inside a cell. The field is approximated by an average quantity Q_i^n inside cell C as:

$$Q_i^n = \frac{1}{dx} \int_C q(x, t) dx$$



Ingredients

To describe what is happening we put ourselve into a finite volume cell that we denote as \mathscr{C} and denote the boundaries as $x \in x_1, x_2$. We further assume a positive advection speed a.



The total mass of any quantity inside the cell is

$$\int_{x_1}^{x_2} q(x,t) dx$$

and a change in time can only be due to fluxes across the left and/or right cell boundaries. Thus

$$\partial_t \int_{x_1}^{x_2} q(x,t) dx = F_1(t) - F_2(t)$$

where $F_i(t)$ are rates (e.g., in g/s) at which the quantity flows through the boundaries.

If we assume advection with a constant transport velocity a this flux is given as a function of the values of q(x, t) as

$$F \rightarrow f(q(x,t)) = aq(x,t)$$

in other words

$$\partial_t \int_{x_1}^{x_2} q(x,t) dx = f(q(x_1,t)) - f(q(x_2,t))$$

This is called the integral form of a hyperbolic conservation law.

Using the definition of integration and antiderivates to obtain

$$\partial_t \int_{x_1}^{x_2} q(x,t) dx = -\int_{x_1}^{x_2} \partial_x f(q(x,t)) dx$$
$$\int_{x_1}^{x_2} \left[\partial_t q(x,t) + \partial_x f(q(x,t)) \right] dx = 0$$

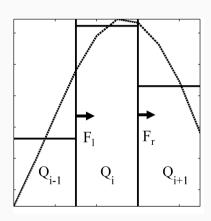
which leads to the well-known partial-differential equation of linear advection

$$\partial_t q(x,t) + \partial_x f(q(x,t)) = 0$$

Instead of working on the field q(x,t) itself we approximate the integral of q(x,t) over the cell $\mathscr C$ by

$$Q_i^n \approx \frac{1}{dx} \int_{\mathscr{C}} q(x, t^n) dx$$

This is the average value of q(x, t) inside the cell.



In order to find an extrapolation scheme to approximate the future state of our finite-volume cells we integrate the integral form of a hyperbolic conservation law.

$$\int_{C} q(x, t^{n+1}) dx - \int_{C} q(x, t^{n}) dx$$

$$= \int_{t_{n}}^{t_{n+1}} f(q(x_{L}, t)) dt - \int_{t_{n}}^{t_{n+1}} f(q(x_{R}, t)) dt$$

where we used the definitions $\mathscr{C} \to [x_1, x_2] = [x_L, x_R]$, rearranged terms, and divided by dx in order to recover the average cell values.

This equation is exact!

Using the following terms for the fluxes at the boundaries

$$F_{L,R}^n = \int_t^{t_{n+1}} f(q(x_{L,R},t)) dt$$

we obtain a time-discrete scheme for the average values of our solution field q(x,t)

$$Q_i^{n+1} = Q_i^n - \frac{dt}{dx}(F_R^n - F_L^n)$$

where the upper index n denotes time level $t_n = n * dt$ and the lower index i denotes cell \mathcal{C}_i of size dx.

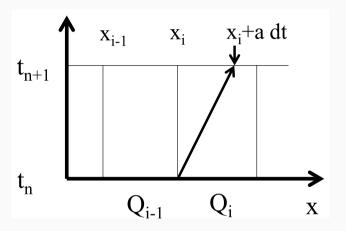


Figure 1: For the linear advection problem we can analytically predict where the tracer will be located after time dt. The value of $q(x_i, t^{n+1})$ will be exactly the same as $q(x_i - adt, t^n)$. We can use this information to predict the new cell average Q_i^{n+1} .

We thus seek to approximate the next cell update Q_i^{n+1} knowing that

$$Q_i^{n+1} \approx q(x_i, t^{n+1}) = q(x_i - adt, t^n)$$

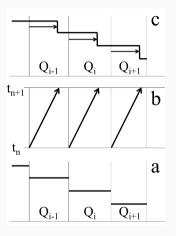
The new cell average analytically by adding the appropriate mass flowing via the left boundary by interpolation

$$Q_{i}^{n+1} = Q_{i-1}^{n} + \frac{dx - adt}{dx} (Q_{i}^{n} - Q_{i-1}^{n})$$

$$Q_{i}^{n+1} = Q_{i}^{n} (1 - \frac{adt}{dx}) + Q_{i-1}^{n} \frac{adt}{dx}.$$

After re-arranging we finally obtain a fully discrete scheme

$$Q_i^{n+1} = Q_i^n - \frac{adt}{dx}(Q_i^n - Q_{i-1}^n)$$



Stability criterion:

$$\left|\frac{adt}{dx}\right| \leq 1$$

Upwind scheme is of 1st order accuracy only and very dispersive. Therefore it is not accurate enough to be of any use for actual simulation tasks.

The Lax-Wendroff Scheme

Our goal is to find solutions to $\partial_t Q + a \partial_x Q = 0$. We start by using the Taylor expansion to extrapolate Q(x,t) in time to obtain

$$Q(x,t^{n+1}) = Q(x,t^n) + dt\partial_t Q(x,t^n) + \frac{1}{2}dt^2\partial_t^2 Q(x,t^n) + \dots$$

From the governing equation we can also state by additional differentiations

$$\partial_t^2 Q = -a\partial_x \partial_t Q$$
$$\partial_x \partial_t Q = \partial_t \partial_x Q = \partial_x (-a\partial_x Q)$$
$$\partial_t^2 Q = a^2 \partial_x^2 Q$$

noting that we just derived the 2nd order form of the acoustic wave equation.

The Lax-Wendroff Scheme

Replacing time derivatives by the equivalent expressions containing space derivatives only and obtain

$$Q(x,t^{n+1}) = Q(x,t^n) - dta\partial_x Q(x,t^n) + \frac{1}{2}dt^2a^2\partial_x^2 Q(x,t^n) + \dots$$

Using central differencing schemes for both space derivatives

$$\partial_X Q(x, t^n) pprox rac{Q_{i+1}^n - Q_{i-1}^n}{2dx}$$
 $\partial_X^2 Q(x, t^n) pprox rac{Q_{i+1}^n - 2Q_i^n + Q_{i-1}^n}{dx^2}$

we finally obtain a fully discrete second-order scheme

$$Q_{i}^{n+1} = Q_{i}^{n} - \frac{adt}{2dx}(Q_{i+1}^{n} - Q_{i-1}^{n}) + \frac{1}{2}(\frac{adt}{dx})^{2}(Q_{i+1}^{n} - 2Q_{i}^{n} + Q_{i-1}^{n})$$

known as the Lax-Wendroff scheme.

The Lax-Wendroff Scheme

Using standard finite-difference considerations without making use of flux concepts.

By extending the finite-volume method towards higher order to approximate the solution inside the finite volume as a piecewise linear function.

The choice of slope considered then determines the specific 2nd order numerical scheme that evolves.

The Lax-Wendroff scheme can also be interpreted as a finite-volume method by considering the flux functions

$$F_{L}^{n} = \frac{1}{2}a(Q_{i-1}^{n} + Q_{i}^{n}) - \frac{1}{2}\frac{dt}{dx}a^{2}(Q_{i}^{n} - Q_{i-1}^{n})$$

$$F_{R}^{n} = \frac{1}{2}a(Q_{i}^{n} + Q_{i+1}^{n}) - \frac{1}{2}\frac{dt}{dx}a^{2}(Q_{i+1}^{n} - Q_{i}^{n})$$

The Finite-Volume Method: Scalar

Advection

We proceed with implementing the two numerical schemes 1) the upwind method and 2) the Lax-Wendroff scheme. Recalling their formulations

$$Q_i^{n+1} = Q_i^n - \frac{adt}{dx}(Q_i^n - Q_{i-1}^n)$$

and

$$Q_i^{n+1} = Q_i^n - \frac{adt}{2dx}(Q_{i+1}^n - Q_{i-1}^n) + \frac{1}{2}(\frac{adt}{dx})^2(Q_{i+1}^n - 2Q_i^n + Q_{i-1}^n)$$

Using a spatial initial condition, a Gauss function

$$Q(x, t = 0) = e^{-1/\sigma^2(x-x_0)^2}$$

that is advected with speed c = 2500 m/s. The analytical solution to this problem is a simple translation of the initial condition to $x = x_0 + ct$, where t = idt is the simulation time at time step i.

Parameter	Value
X _{max}	75000 m
nx	6000
С	2500 m/s
dt	0.0025 s
dx	12.5 m
ϵ	0.9
σ (Gauss)	200 m
<i>x</i> ₀	1000 m

```
# Time extrapolation
for i in range(nt):
    # upwind method
   if method == 'upwind':
        for j in range (1, nx-1):
            # Forward (upwind) (c>0)
           dO[i] = O[i] - O[i-1]
        # Time extrapolation
        0 = 0 - dt/dx*c*d0
    # Lax wendroff method
   if method == 'Lax-Wendroff':
        for j in range(1, nx-1):
            # Forward (upwind) (c>0)
           d01[i] = 0[i+1] - 2*0[i] + 0[i-1]
           dQ2[i] = Q[i+1] - Q[i-1]
        # Time extrapolation
        Q = Q - 0.5*c*dt/dx*dQ2 + 0.5*(c*dt/dx)**2 *d01
```

Implemement periodic and absorbing boundary conditions with the statements

Periodic:
$$Q_1^n = Q_{nx-1}^n$$

Absorbing: $Q_{nx}^n = Q_{nx-1}^n$



Figure 2: Boundary conditions. Absorbing, or circular boundary conditions can be implemented by using ghost cells outside the physical domain $x \in [x_0, x_{max}]$.

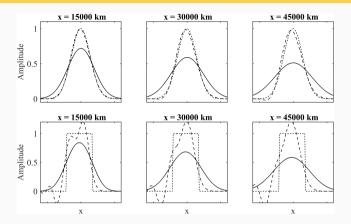


Figure 3: Top: Snapshots of an advected Gauss function (analytical solution, thick solid line) are compared with the numerical solution of the 1st order upwind method (thin solid line) and the 2nd order Lax-Wendroff scheme (dotted line) for increasing propagation distances. **Bottom:** The same for a box-car function.

Implemement periodic and absorbing boundary conditions with the statements

Periodic:
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Figure 4: Boundary conditions. Absorbing, or circular boundary conditions can be implemented by using ghost cells outside the physical domain $x \in [x_0, x_{max}]$.

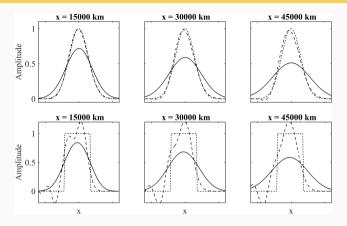


Figure 5: Top: Snapshots of an advected Gauss function (analytical solution, thick solid line) are compared with the numerical solution of the 1st order upwind method (thin solid line) and the 2nd order Lax-Wendroff scheme (dotted line) for increasing propagation distances. **Bottom:** The same for a box-car function.

The Finite-Volume Method for

Elastic Waves

Source-free version of the coupled first-order elastic wave equation

$$\partial_t v - \frac{1}{\rho} \partial_x \sigma = 0$$
$$\partial_t \sigma - \mu \partial_x v = 0.$$

We proceed by writing this equation in matrix-vector notation

$$\partial_t \mathbf{Q} + \mathbf{A} \partial_x \mathbf{Q} = 0$$

where $\mathbf{Q} = (\sigma, \mathbf{v})$ is the vector of unknowns and matrix \mathbf{A} contains the parameters

$$\mathbf{A} = \begin{pmatrix} 0 & -1/\rho \\ -\mu & 0 \end{pmatrix}$$

The problem hereby is that these eqations are coupled. What needs to be done is to demonstrate the *hyperbolicity* of the wave equation in this form, i.e. show that **A** is diagonalizable.

In the case of a quadratic matrix $\bf A$ with shape $m \times m$ leads to an eigenvalue problem. If we were able to obtain eigenvalues λ_p such that

$$\mathbf{A}\mathbf{x}_{p}=\lambda_{p}\mathbf{x}_{p}\;,\;\;p=1,...,m$$

we get a diagonal matrix of eigenvalues

$$\Lambda = \begin{pmatrix} \lambda_1 & & \\ & \ddots & \\ & & \lambda_m \end{pmatrix}$$

and the corresponding matrix \mathbf{R} containing the eigenvectors $\mathbf{x}_{\mathbf{p}}$ in each column

$$\mathbf{R} = (\mathbf{x}_1 | \mathbf{x}_2 | \dots | \mathbf{x}_p) .$$

The Jacobian matrix A can now be expressed with the definitions

$$\mathbf{A} = \mathbf{R} \wedge \mathbf{R}^{-1}$$
$$\Lambda = \mathbf{R}^{-1} \mathbf{A} \mathbf{R} .$$

Applying these definitions to above equation we obtain

$$\mathbf{R}^{-1}\partial_t\mathbf{Q} + \mathbf{R}^{-1}\mathbf{R}\Lambda\mathbf{R}^{-1}\partial_x\mathbf{Q} = 0$$

and introducing the solution vector $\mathbf{W} = \mathbf{R}^{-1}\mathbf{Q}$ results in

$$\partial_t \mathbf{W} + \Lambda \partial_x \mathbf{W} = \mathbf{0} \ .$$

What remains to be shown is that in our specific case **A** has real eigenvalues. These are easily determined as $\lambda_{1,2}=\pm\sqrt{\mu/\rho}=\pm c$, corresponding to the shear velocity c. For the eigenvectors we obtain

$$\mathbf{r}_{1,2} = \begin{pmatrix} \pm \rho \mathbf{c} \\ 1 \end{pmatrix}$$

interestingly enough containing as first elements values of the *seismic impedance* $Z = \rho c$ that is relevant for the reflection behaviour of seismic waves. Thus, the matrix **R** and its inverse are

$$\mathbf{R} = \begin{pmatrix} Z & -Z \\ 1 & 1 \end{pmatrix} , \mathbf{R}^{-1} = \frac{1}{2Z} \begin{pmatrix} 1 & Z \\ -1 & Z \end{pmatrix} .$$

The wave equation in the rotated eigensystem can be stated as

$$\partial_t \begin{pmatrix} w_1 \\ w_2 \end{pmatrix} + \begin{pmatrix} -c & 0 \\ 0 & c \end{pmatrix} \partial_x \begin{pmatrix} w_1 \\ w_2 \end{pmatrix} = 0$$

with the simple general solution $w_{1,2} = w_{1,2}^{(0)}(x \pm ct)$, where the upper index 0 stands for the initial condition.

The initial condition also fullfills $\mathbf{W}^{(0)} = \mathbf{R}^{-1}\mathbf{Q}^{(0)}$. We can therefore relate the so-called characteristic variables $w_{1,2}$ to the initial conditions of the physical variables as

$$w_1(x,t) = \frac{1}{2Z}(\sigma^{(0)}(x+ct) + Zv^{(0)}(x+ct))$$

$$w_2(x,t) = \frac{1}{2Z}(-\sigma^{(0)}(x-ct) + Zv^{(0)}(x-ct))$$

Obtaining the final analytical solution for velocity v and stress σ as

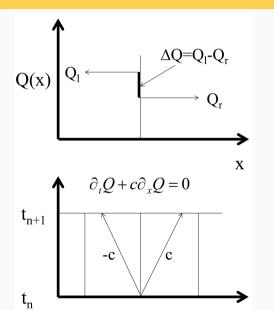
$$\sigma(x,t) = \frac{1}{2}(\sigma^{(0)}(x+ct) + \sigma^{(0)}(x-ct)) + \frac{Z}{2}(v^{(0)}(x+ct) - v^{(0)}(x-ct)) + \frac{1}{2Z}(\sigma^{(0)}(x+ct) - \sigma^{(0)}(x-ct)) + \frac{1}{2}(v^{(0)}(x+ct) + v^{(0)}(x-ct)).$$

In compact form this solution can be expressed as

$$\mathbf{Q}(x,t) = \sum_{p=1}^{m} \mathbf{w}_{p}(x,t) \mathbf{r}_{p}$$

meaning that any solution is a sum over weighted eigenvectors, a superposition of m waves.

Homogeneous Case



Riemann problem, homogeneous case. **Top:** A discontinuity ΔQ is located at x = 0 as initial condition to the advection equation (e.g., as initial stress discontinuity). **Bottom:** The discontinuity propagates along characteristic curves in the space-time domain. The figure illustrates adjacent cells and two time levels t_n and t_{n+1} . Two waves propagate in opposite direction modifying the values in the cells adjacent to x = 0.

The solution to our problem is a superposition of weighted eigenvectors \mathbf{r}_p , in our case p=1,2. Therefore, we can decompose the discontinuity jump into these eigenvectors according to

$$\Delta \mathbf{Q} = \mathbf{Q}_r - \mathbf{Q}_l = \alpha_1 \mathbf{r}_1 + \alpha_2 \mathbf{r}_2$$

$$\mathbf{R} \alpha = \Delta \mathbf{Q}$$

$$\alpha = \mathbf{R}^{-1} \Delta \mathbf{Q}$$

where ${\bf R}$ is the matrix of eigenvectors as defined above and α are unknown weights.

Decompose the solution into positive (right-propagating) and negative (left-propagating) eigenvalues

$$\Lambda^- = egin{pmatrix} -c & 0 \ 0 & 0 \end{pmatrix} \;,\;\; \Lambda^+ = egin{pmatrix} 0 & 0 \ 0 & c \end{pmatrix}$$

Then we can derive matrices \mathbf{A}^\pm - corresponding to the advection velocity in the scalar case

$$\mathbf{A}^+ = \mathbf{R} \Lambda^+ \mathbf{R}^{-1}$$

 $\mathbf{A}^- = \mathbf{R} \Lambda^- \mathbf{R}^{-1}$

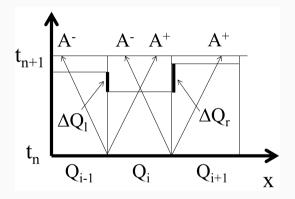


Figure 6: The constant average cell values of Q are illustrated for three adjacent cells i-1, i, i+1. The eigenvector decomposition leads to wave A^+ propagating from the left boundary with velocity c in to cell i and wave A^- propagating with velocity -c into cell i from the right boundary. This determines the flux of discontinuities $\Delta \mathbf{Q}_{l,r}$ into cell i by the amount $dt/dx\Delta \mathbf{Q}_{l,r}$.

We are ready to formulate an upwind finite-volume scheme for any multi-dimensional linear hyperbolic system:

$$\Delta \mathbf{Q}_{I} = \mathbf{Q}_{i} - \mathbf{Q}_{i-1}$$

$$\Delta \mathbf{Q}_{r} = \mathbf{Q}_{i+1} - \mathbf{Q}_{i}$$

$$\mathbf{Q}_{i}^{n+1} = \mathbf{Q}_{i}^{n} - \frac{dt}{dx} (\mathbf{A}^{+} \Delta \mathbf{Q}_{I} + \mathbf{A}^{-} \Delta \mathbf{Q}_{r}) .$$

We can relate this formulation to the very basic flux concept

$$\mathbf{F}_I = \mathbf{A}^+ \Delta \mathbf{Q}_I$$

 $\mathbf{F}_I = \mathbf{A}^- \Delta \mathbf{Q}_I$.

The 1st order upwind solution is of no practical use because of its strong dispersive behaviour.

⇒The 2nd order Lax-Wendroff scheme

The high-order scheme does not necessitate the separation into eigenvectors and the Jacobian matrix *A* can be used in its original form. The extrapolation scheme reads

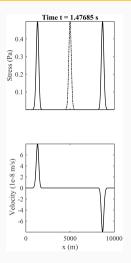
$$\mathbf{Q}_{i}^{n+1} = \mathbf{Q}_{i}^{n} - \frac{dt}{2dx} \mathbf{A} (\mathbf{Q}_{i+1}^{n} + \mathbf{Q}_{i-1}^{n}) + \frac{1}{2} \frac{dt}{dx}^{2} \mathbf{A}^{2} (\mathbf{Q}_{i-1}^{n} - 2\mathbf{Q}_{i}^{n} + \mathbf{Q}_{i+1}^{n}).$$

Example

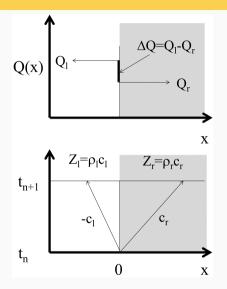
Parameter	Value
X _{max}	10000 m
nx	800
С	2500 m/s
ho	$2500 \ kg/m^3$
dt	0.025 s
dx	12.5 m
ϵ	0.5
σ (Gauss)	200 m
<i>X</i> ₀	5000 m

```
# Time extrapolation
for i in range(nt):
   if imethod == 'Lax-Wendroff':
       for j in range(1,nx-1):
           d01 = 0[:, i+1] - 0[:, i-1]
           d02 = 0[:, j-1] - 2*0[:, j] + 0[:, j+1]
           Qnew[:,i] = Q[:,i] - 0.5*dt/dx*(A @ dQ1)
           + 1./2.*(dt/dx)**2 * (A @ A) @ dO2 # Eq. 8.56
        # Absorbing boundary conditions
       Onew[:,0] = Onew[:,1]
       Qnew[:,nx-1] = Qnew[:,nx-2]
   elif imethod == 'upwind':
       for j in range(1,nx-1):
           d01 = 0[:,i] - 0[:,i-1]
           dOr = O[:, j+1] - O[:, j]
           Onew[:,i] = O[:,i] - dt/dx * (Ap @ dOl + Am @ dOr)
        # Absorbing boundary conditions
       Onew[:,0] = Onew[:,1]
       Onew[:.nx-1] = Onew[:.nx-2]
```

Result



The stress-velocity system is advected for an initial condition of Gaussian shape (top, dashed line, scaled by factor 1/2). **Top:** Stress snapshot at time t=1.5s. **Bottom:** Velocity snapshot at the same time. In both cases analytical solutions are superimposed.



Top: Single discontinuity separating two regions with different properties. **Bottom:** Velocities and impedances on both sides of the discontinuity. The Riemann problem solves the problem of how waves on both sides are partitioned.

Mathematically the solution still has to consist of a weighted sum over eigenvectors that now describe solutions in the left and right parts.

$$\Delta \mathbf{Q} = \alpha_1 \begin{pmatrix} -Z_l \\ 1 \end{pmatrix} + \alpha_2 \begin{pmatrix} Z_r \\ 1 \end{pmatrix}$$

for some unknown scalar values $\alpha_{1,2}$. This can be written as a linear system of the form

$$\mathbf{R}_{lr}\alpha = \Delta \mathbf{Q}$$

where α is a vector and the matrices with the eigenvector are

$$\mathbf{R}_{lr} = \begin{pmatrix} -Z_l & Z_r \\ 1 & 1 \end{pmatrix} , \ \mathbf{R}_{lr}^{-1} = \frac{1}{Z_l + Z_r} \begin{pmatrix} -Z_l & Z_r \\ 1 & 1 \end{pmatrix}$$

Using the \mathbf{R} matrix for cell boundaries separating cells with different (constant) properties. Again we separate into left- and right-propagating eigenvalues within cell i

$$\Lambda^- = egin{pmatrix} -c_i & 0 \ 0 & 0 \end{pmatrix} \;,\;\; \Lambda^+ = egin{pmatrix} 0 & 0 \ 0 & c_i \end{pmatrix}$$

and using the definitions

$$\mathbf{R}_{I} = \begin{pmatrix} -Z_{i-1} & Z_{i} \\ 1 & 1 \end{pmatrix} , \quad \mathbf{R}_{r} = \begin{pmatrix} -Z_{i} & Z_{i} + 1 \\ 1 & 1 \end{pmatrix}$$

for the eigenvectors describing the solutions around the left and right boundaries we can determine the corresponding advection terms as

$$\mathbf{A}^+ = \mathbf{R}_I \Lambda_I^+ \mathbf{R}_I^{-1}$$
$$\mathbf{A}^- = \mathbf{R}_r \Lambda_r^+ \mathbf{R}_r^{-1}$$

Leading to the 1st order upwind extrapolation scheme for the solution vector Q_i in the general heterogeneous case

$$\Delta \mathbf{Q}_{I} = \mathbf{Q}_{i} - \mathbf{Q}_{i-1}$$

$$\Delta \mathbf{Q}_{r} = \mathbf{Q}_{i+1} - \mathbf{Q}_{i}$$

$$\mathbf{Q}_{i}^{n+1} = \mathbf{Q}_{i}^{n} - \frac{dt}{dx} (\mathbf{A}^{+} \Delta \mathbf{Q}_{I} + \mathbf{A}^{-} \Delta \mathbf{Q}_{r})$$

⇒ too dispersive but interesting side effect!

Let us take the eigenvector (i.e., wave) propagating in the left domain. What does this imply for the wave propagating in the right domain?

$$\Delta \mathbf{Q} = \begin{pmatrix} Z_l \\ 1 \end{pmatrix}$$

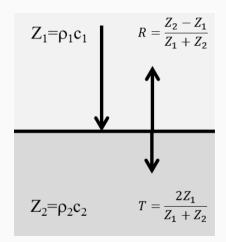
We need to determine how the waves are partitioned given the discontinuous material parameters described by matrix \mathbf{R}_{Ir} . This leads to the coefficients α as

$$\alpha = \mathbf{R}_{lr}^{-1} \Delta \mathbf{Q}$$

$$= \frac{1}{Z_l + Z_r} \begin{pmatrix} -1 & Z_r \\ 1 & Z_l \end{pmatrix} \begin{pmatrix} Z_l \\ 1 \end{pmatrix}$$

$$= \begin{pmatrix} \frac{Z_r - Z_l}{Z_l + Z_r} \\ \frac{2Z_l}{Z_l + Z_r} \end{pmatrix} = \begin{pmatrix} R \\ T \end{pmatrix}$$

that are the well known transmission and reflection coefficients for vertical incidence at a material discontinuity.



Reflection and transmission coefficients. Seismic waves incident
perpendicular to a material discontinuity are
reflected and transmitted according to
coefficients *R* and *T*. These coefficients
can be derived via the Riemann problem
used to develop flux schemes for
finite-volume methods.

Finally, we present the 2nd order scheme for the general heterogeneous results and a simulation example.

$$\mathbf{A}_{i-1/2} = \frac{1}{Z_{i-1} + Z_i} \begin{pmatrix} c_i Z_i - c_{i-1} Z_{i-1} & (c_{i-1} + c_i) Z_{i-1} Z_i \\ c_{i-1} + c_i & c_i Z_{i-1} - c_{i-1} Z_i \end{pmatrix}$$

$$\mathbf{A}_{i+1/2} = \frac{1}{Z_i + Z_{i+1}} \begin{pmatrix} c_{i+1} Z_{i+1} - c_i Z_i & (c_i + c_{i+1}) Z_i Z_{i+1} \\ c_i + c_{i+1} & c_{i+1} Z_i - c_i Z_{i+1} \end{pmatrix}.$$

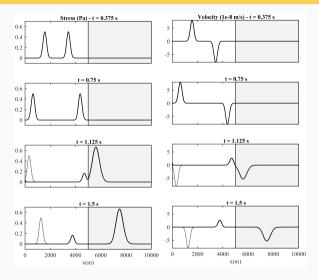
With these definitions we can formulate the Lax-Wendroff extrapolation scheme for elastic waves in heterogeneous material

$$\begin{aligned} \mathbf{Q}_{i}^{n+1} = & \mathbf{Q}_{i}^{n} - \frac{dt}{2dx} \left[\mathbf{A}_{i-1/2} (\mathbf{Q}_{i}^{n} - \mathbf{Q}_{i-1}^{n}) + \mathbf{A}_{i+1/2} (\mathbf{Q}_{i+1}^{n} - \mathbf{Q}_{i}^{n}) \right] \\ &+ \frac{1}{2} (\frac{dt}{dx})^{2} \left[\mathbf{A}_{i+1/2}^{2} (\mathbf{Q}_{i+1}^{n} - \mathbf{Q}_{i}^{n}) - \mathbf{A}_{i-1/2}^{2} (\mathbf{Q}_{i}^{n} - \mathbf{Q}_{i-1}^{n}) \right] \end{aligned}$$

Example

Parameter	Value
X _{max}	10000 m
nx	800
C_{I}	2500 m/s
C_r	5000 m/s
ρ	$2500 \ kg/m^3$
dt	0.025 s
dx	12.5 m
ϵ	0.5
σ (Gauss)	200 m
<i>X</i> ₀	5000 m

Result



Summary

- First-order finite-volume schemes are highly dispersive and not appropriate for the solution of wave propagation problems. The 2nd order Lax-Wendroff scheme does a much better job.
- The problem of elastic wave propagation can be formally cast as a 1st order hyperbolic problem. Therefore - only with slight modifications - the fundamental schemes developed for the scalar advection problem can be applied to elastic wave propagation.
- In the finite-volume method the problem of estimating partial derivatives (finite differences) is replaced by the requirement to accurately calculate fluxes across cell boundaries.
- A main advantage of the finite-volume method is the fact that the scheme can be easily applied to volume cells of any shape.
- Finite-volume schemes for arbitrary high-order reconstructions inside the cells and high-order time-extrapolation schemes have been developed but not used extensively in seismology.

Comprehension questions

- 1 What is the connection between finite-volume methods and conservation equations?
- 2 What is meant by finite *volumes*, is there any difference to a finite *element*?
- 3 If you look at the uwpind approach to the scalar advection problem (Eq. ??), why is the finite-volume method so closely linked to staggered-grid finite-difference schemes? Explain.
- 4 What are the main advantages of finite-volume methods compared with finite-difference methods?
- 5 Explain in a qualitative way what the Riemann problem is and why it is so essential for finite-volume schemes.
- 6 In what areas of natural sciences are finite-volume schemes mostly used. Explore the literature and try to give reasons.
- 7 What is numerical diffusion? Why is it relevant for finite-volume methods?
- 8 What is the connection between reflection/transmission coefficients of seismic waves and the finite-volume method?
- 9 The finite-volume method extrapolates cell averages. What strategies do you see to extend the method to high-order accuracy?