

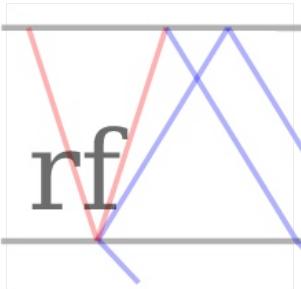
# Intrinsic attenuation and scattering

Why coda Q does not tell the full story

Using seismic envelopes to separate the effects of source,  
site and path

Tom Eulenfeld

# @trichter on GitHub and in the ObsPy forum



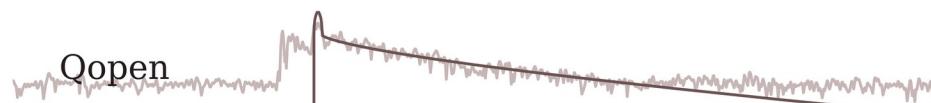
Framework for calculation  
of receiver functions  
Deconvolution, moveout,  
piercing points, ...

## obspyh5

Quick & dirty IO of  
waveforms preserving  
metadata, HDF5



dv/v with stretching technique  
CLI  
configuration in JSON file  
easy definition and house-keeping of  
different correlation and stretching schemes  
cc shorter than 1d possible



## obspycsv

Quick & dirty IO of  
earthquake catalogs to  
CSV format  
read EVENTTXT  
flatten ObsPy catalogs to  
NumPy arrays

# Contents

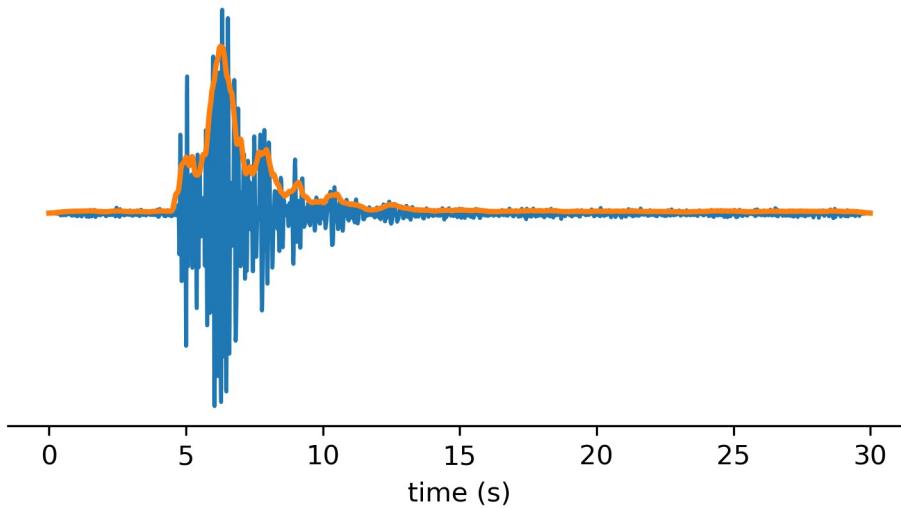
Introduction  
Motivation  
Qopen method  
Applications

- USArray
- Helsinki geothermal stimulation

# Introduction

# Envelope and radiative transfer

- Phase information in the coda cannot be modeled easily
- Only coda amplitude (resp. energy) is of interest
- Convert waveforms to envelopes (Hilbert transform)
- Transition of wave equation to equation of radiative transfer
- Opens field for Monte-Carlo particle simulations



# Intrinsic attenuation vs scattering, about Quality factors

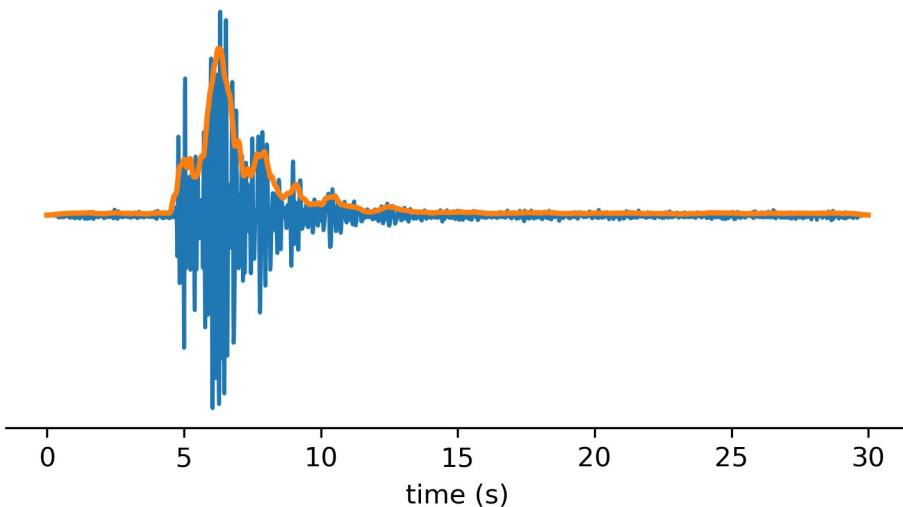
Definition Quality factor

$$Q := 2\pi \frac{\text{total energy}}{\text{energy loss per cycle}}$$

For direct wave:

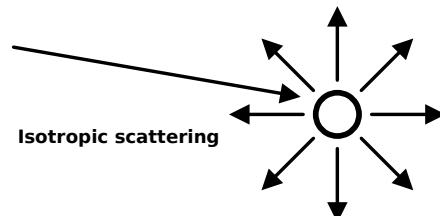
$$Q^{-1} = Q_{\text{intr}}^{-1} + Q_{\text{scatt}}^{-1}$$

$$Q_{\text{intr}}^{-1} = \frac{v}{2\pi f l_a} \quad Q_{\text{scatt}}^{-1} = \frac{v}{2\pi f l}$$

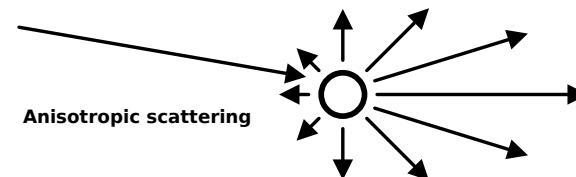


# Isotropic vs nonisotropic scattering – transport mean free path

- Mean free path  $l_0$ : Length in which 63% of the wave energy is scattered, mean length between two scattering events
- Transport mean free path  $l^*$ : Length in which the propagation direction of 63% of the wave energy becomes independent from its original propagation direction—the wave “forgets” its initial direction due to scattering



$$l^* = l_0 \quad t^* = t_0$$



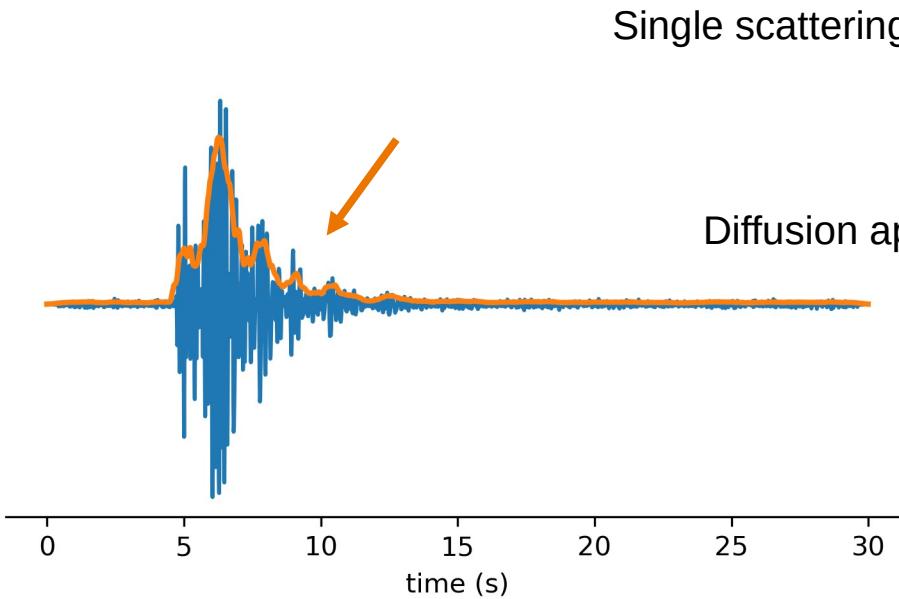
$$l^* > l_0 \quad t^* > t_0$$

$$t_0 = \frac{l_0}{v_s} \qquad t^* = \frac{l^*}{v_s}$$

# What about coda Q?

Obviously coda Q is not simply the sum of intrinsic and scattering Q as for the direct wave.

The interpretation of coda Q depends on the scattering regime in the coda!



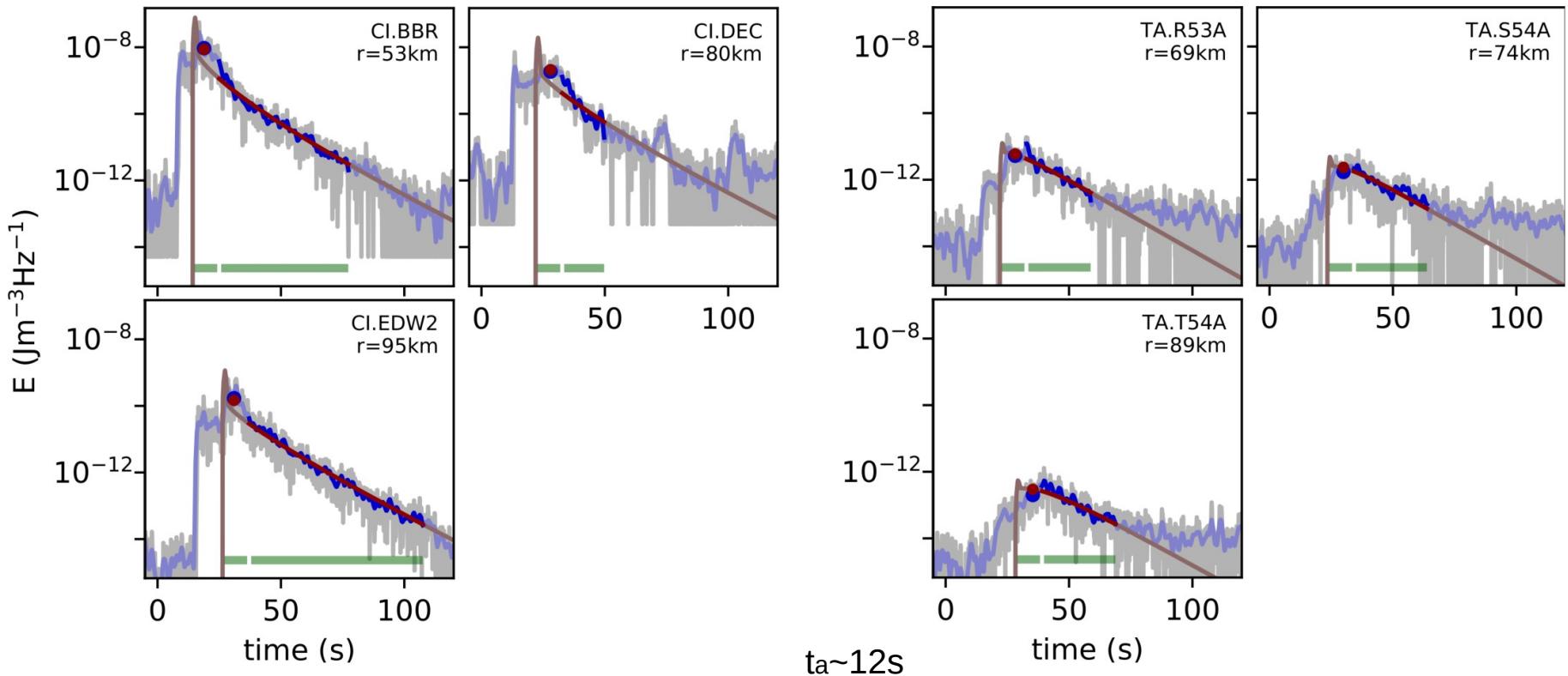
Single scattering approximation:

$$Q_{\text{coda}}^{-1} = Q_{\text{intr}}^{-1} + Q_{\text{scatt}}^{-1}$$

Diffusion approximation:

$$Q_{\text{coda}}^{-1} = Q_{\text{intr}}^{-1}$$

# Scattering regime in the coda – transport mean free time



$t^* \sim 40s \Rightarrow$  single and multiple scattering

$t^* \sim 7s \Rightarrow$  diffusion approximation valid

$$Q_{\text{coda}}^{-1} = Q_{\text{intr}}^{-1} + Q_{\text{scatt}}^{-1} ?$$

$$Q_{\text{coda}}^{-1} = Q_{\text{intr}}^{-1}$$

$\Rightarrow$  scattering regime can be determined with the shape of the envelope

# Motivation

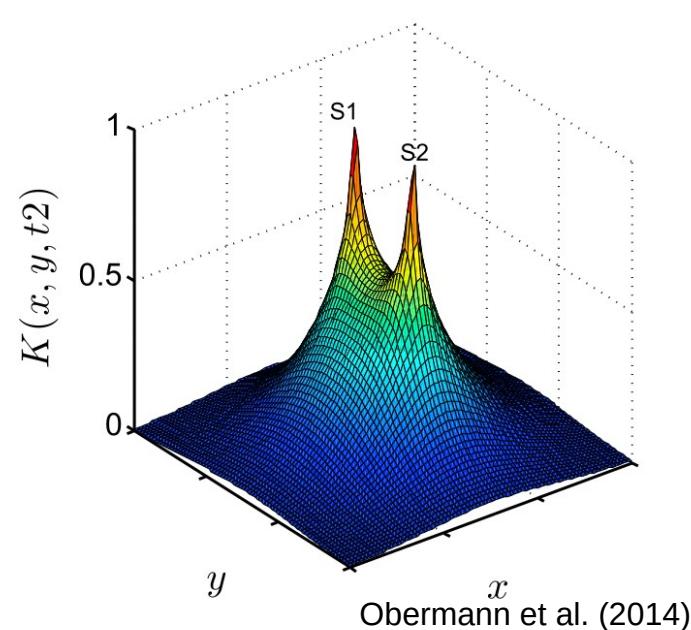
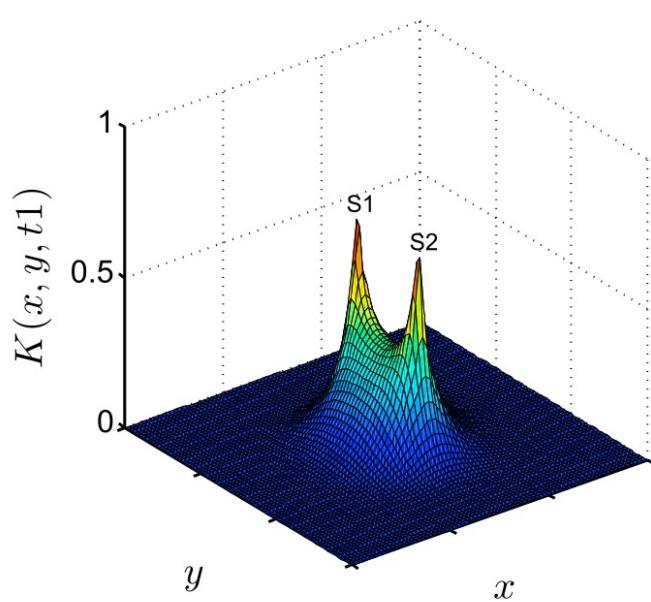
# Kernels for tomography of coda Q and dv/v observations

Observations of relative velocity change (dv/v) often use the coda

Coda Q can be determined for each station-earthquake pair (similar to first arrivals) and is therefore predestined for tomography.

=> Need for travel time kernel of the coda

=> Estimate of transport mean free path can confine the shape of the kernel  
(and check validity of assumptions leading to kernel estimate)



Obermann et al. (2014)

# Source, Site, Path

Seismogram is convolution of

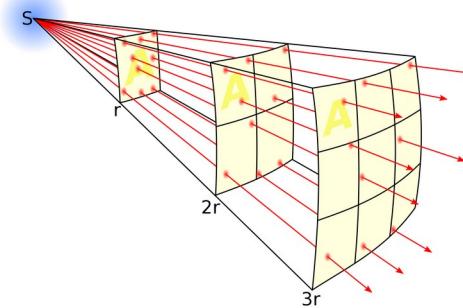
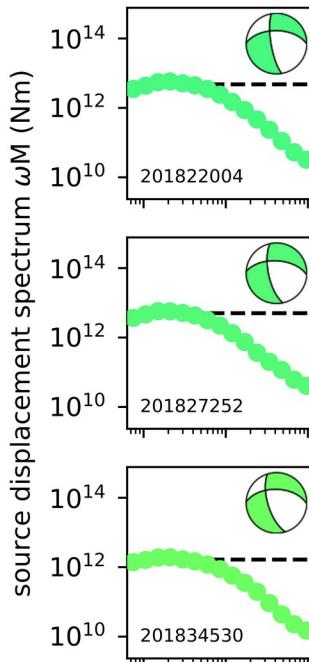
source function

x

propagation filter

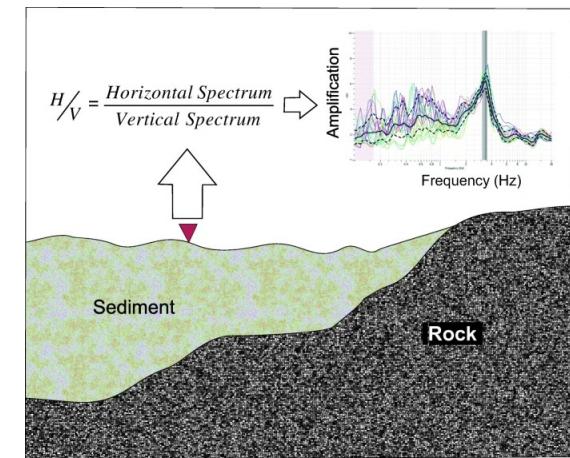
x

site response



- Moment tensor
- Moment rate function / source displacement spectrum
- Slip distribution

- Geometrical spreading
- Attenuation
- Scattering
- Reflections, conversions, ...

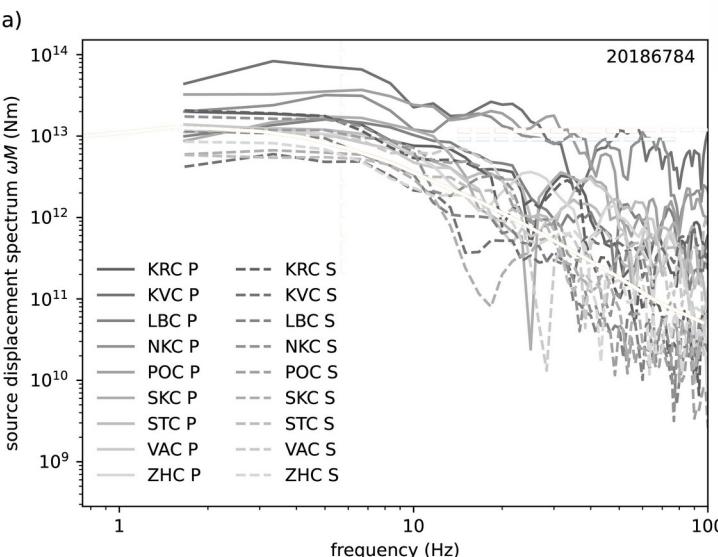


Yilmaz al. 2021

- H/V
- Vs30
- kappa

# Conventional method to calculate source spectrum

- Take spectra of waveforms around onset
- Correct for geometrical spreading and radiation pattern
- Optimize seismic moment  $M_0$ , corner freq  $f_c$  and attenuation  $Q$



$$\Omega(f) = \frac{\Omega_0 e^{-(\pi f t / Q)}}{\left[1 + (f / f_c)^{\gamma n}\right]^{1/\gamma}} \quad \text{Abercrombie 1995}$$

- Tradeoff between  $Q$  and  $f_c$
- $Q$  can be a function of frequency

Spectrum can be used to calculate stress drop.  
Self-similarity of differently sized earthquakes?

# Qopen method

Separation of intrinsic and scattering  
 $\mathbf{Q}$  by **envelope** inversion

Idea: Intrinsic attenuation and scattering strength can be separated and quantified with the temporal and spatial shape of the envelope!

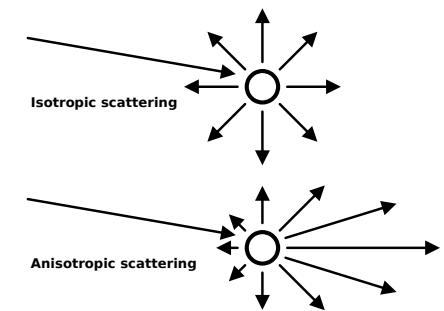
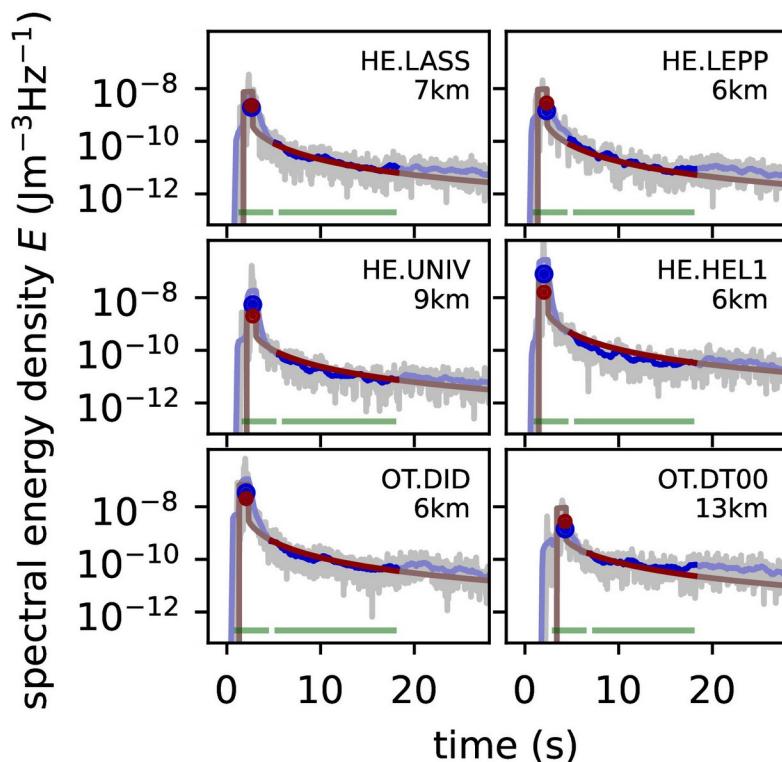
# Qopen method for shear waves

$$\overbrace{E_{\text{mod}ij}}^{\text{Modelled energy density envelope}} = \overbrace{R_i W G(g_0, r_i, t_j)}^{\text{Scattering Green's function}} e^{-\overbrace{bt_j}^{\text{Constant damping}}} \\ \begin{array}{l} \text{station index} \\ \text{sample index} \\ \text{site amplification factor} \\ \text{spectral source energy} \end{array} \quad \begin{array}{l} \text{distance} \\ \text{scattering parameter} \\ \rightarrow Q_{\text{sc}} \end{array} \quad \begin{array}{l} \text{absorption constant} \\ \rightarrow Q_i \end{array}$$

- $G$  accounts for geometrical spreading and scattering => here  $G$  is analytic
- Compare with observed envelopes of S wave + coda
- Invert for  $R_i$ ,  $W$ ,  $g_0$  and  $b$  (optimization in  $g_0$  + least squares log fit)
- Repeat the steps for all frequency bands
- Repeat with different earthquakes
- Assumptions:
  - homogeneous half space
  - point source (small EQ)
  - moment tensor ignored

# Imprint of anisotropic scattering

- Qopen assumes isotropic scattering, this is often a bad assumption
- In an anisotropic scattering environment the scattering strength estimated with Qopen relates to the transport mean free path (Gaebler et al. 2015)
- Model cannot predict correct envelope directly after the S body wave  
=> In the inversion the envelope inside the direct wave window needs to be averaged

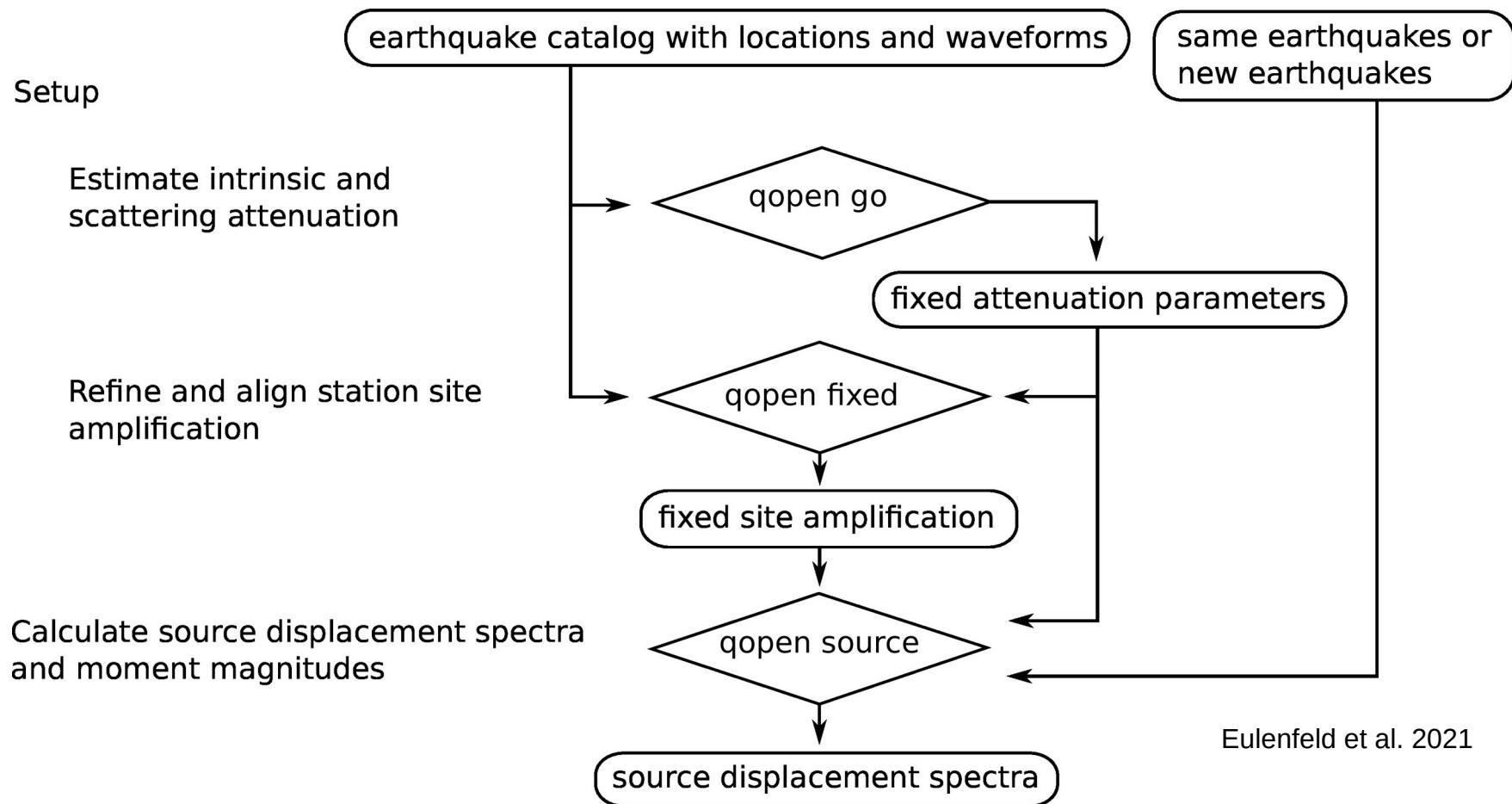


## Estimation of site response and source spectra

$$E_{\text{mod}ij} = \overbrace{R_i W G(g_0, r_i, t_j)}^{\substack{\text{Scattering} \\ \text{Green's function}}} e^{-\overbrace{bt_j}^{\substack{\text{absorption constant} \\ \rightarrow Q_i}}}$$

station index |  
sample index  
site amplification factor  
spectral source energy

distance  
scattering parameter  
 $\rightarrow Q_{\text{sc}}$



# Source spectra and seismic moments

Qopen inversion

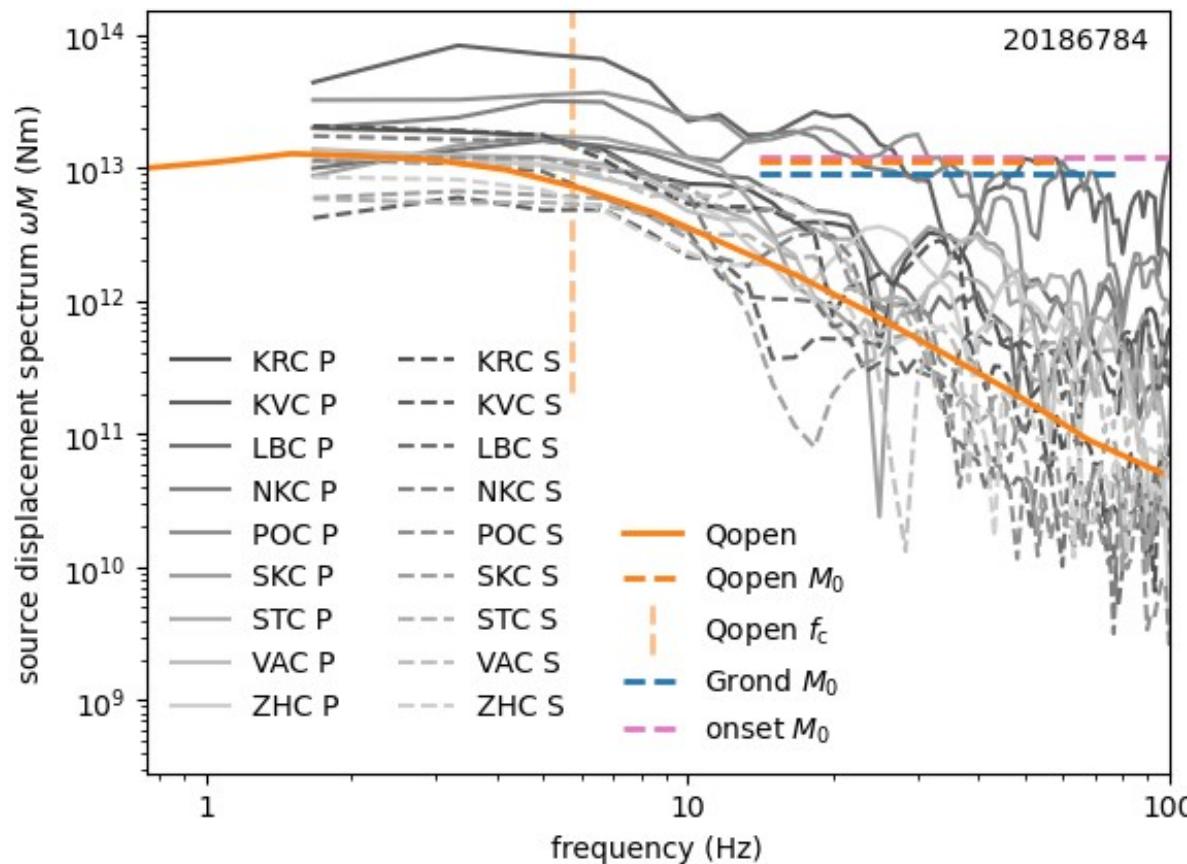
vs Grond moment tensor inversion

vs spectra from Fourier transform  
of body waves

Abercrombie 1995

$$\omega M(f) = M_0 \left( 1 + \left( \frac{f}{f_c} \right)^{\gamma n} \right)^{-\frac{1}{\gamma}}$$

source displacement spectrum      seismic moment  
high frequency fall-off (2 for omega-square model)  
corner sharpness  
corner frequency

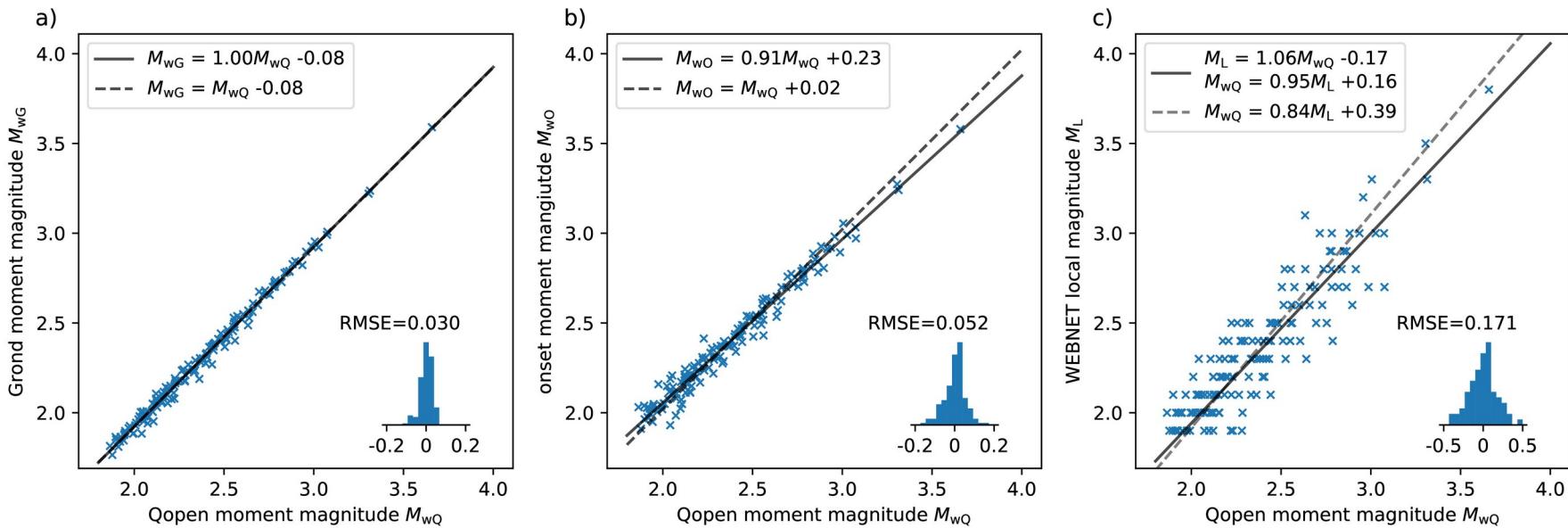


# Czech 2018 EQ swarm – moment magnitudes

Qopen inversion

vs Grond moment tensor inversion

vs spectra from Fourier transform of body waves



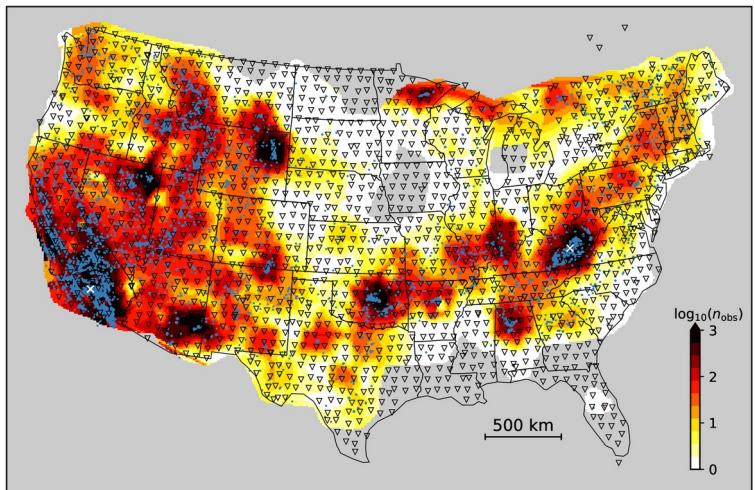
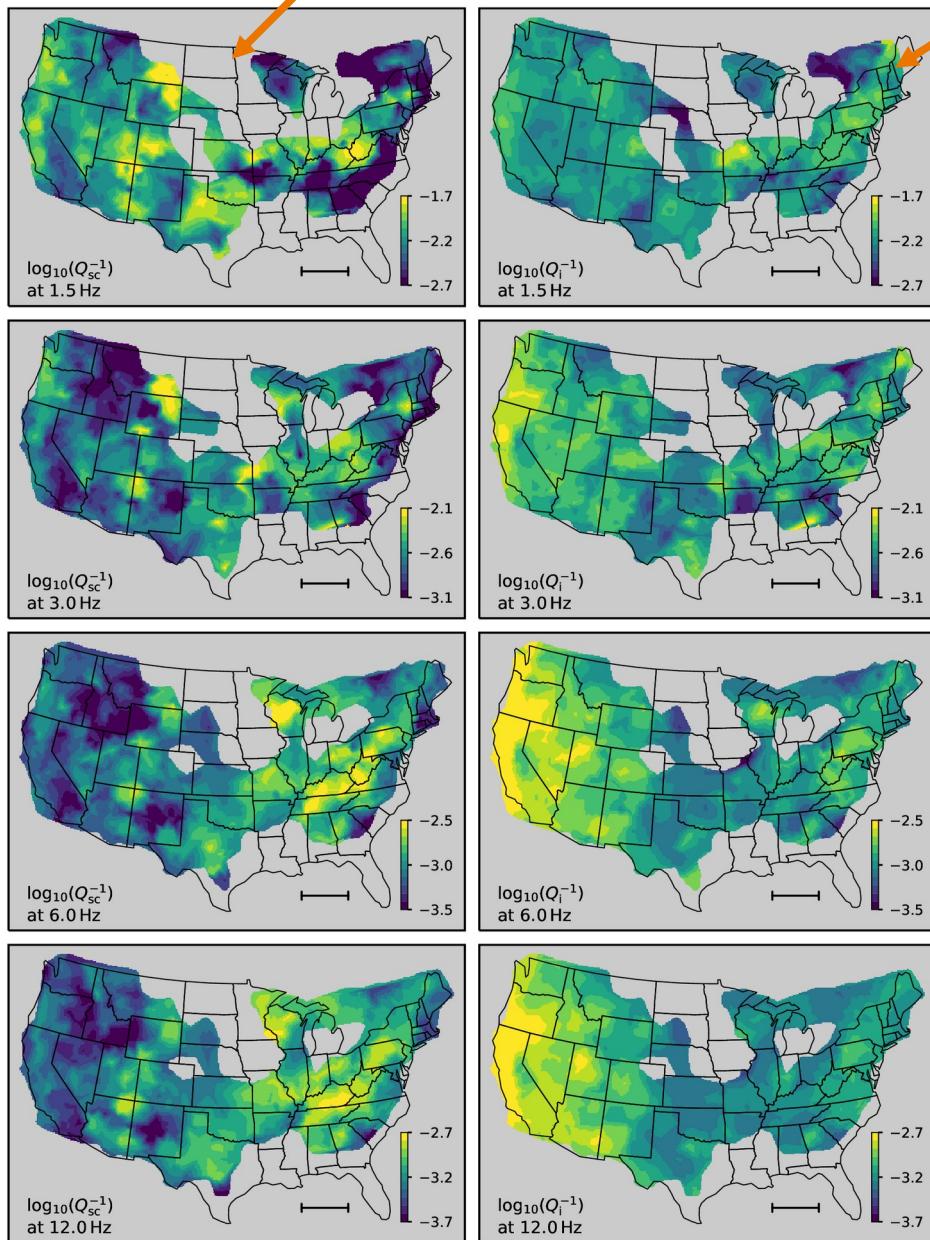
=> Robust estimation of moment magnitudes for small earthquakes

=> Can be used in high scattering environments with a lack of impulsive onsets

Code available at [github.com/trichter/qopen](https://github.com/trichter/qopen)

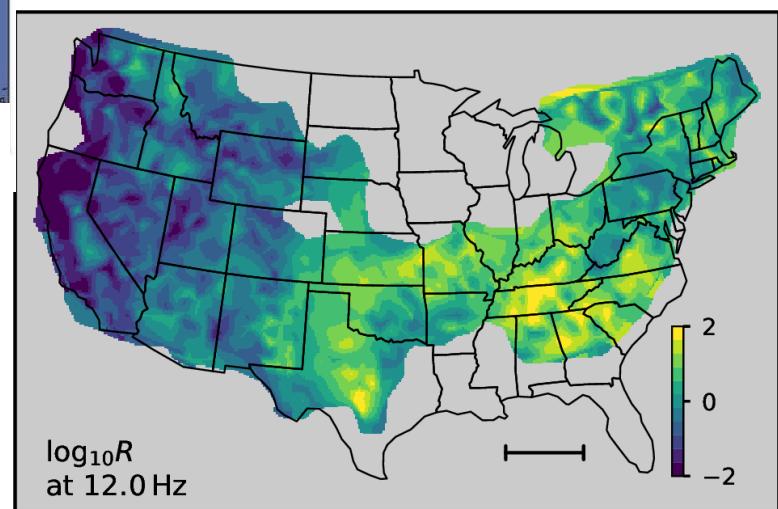
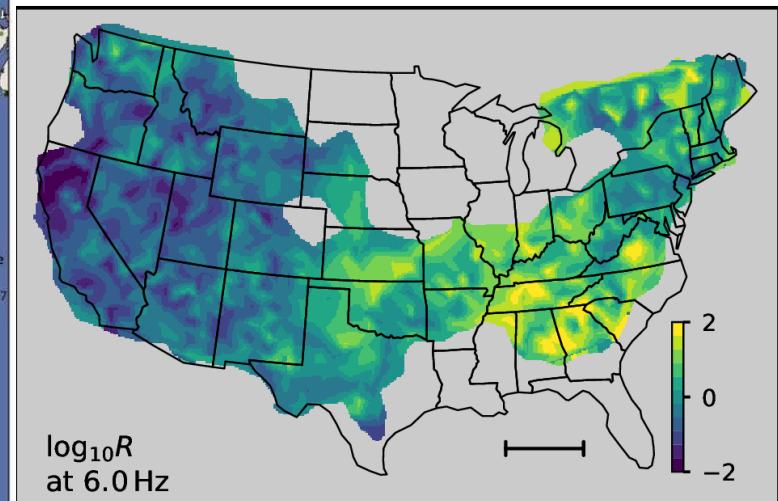
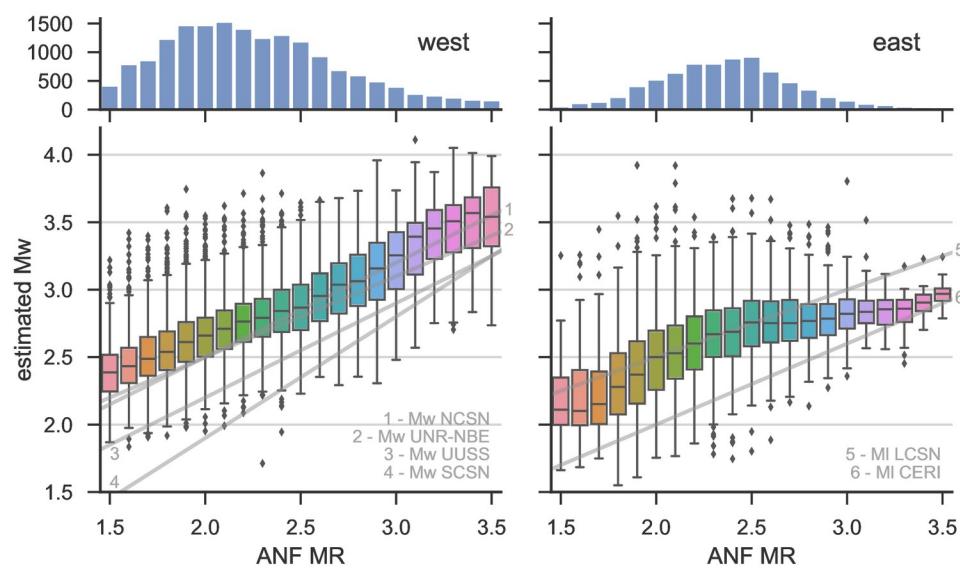
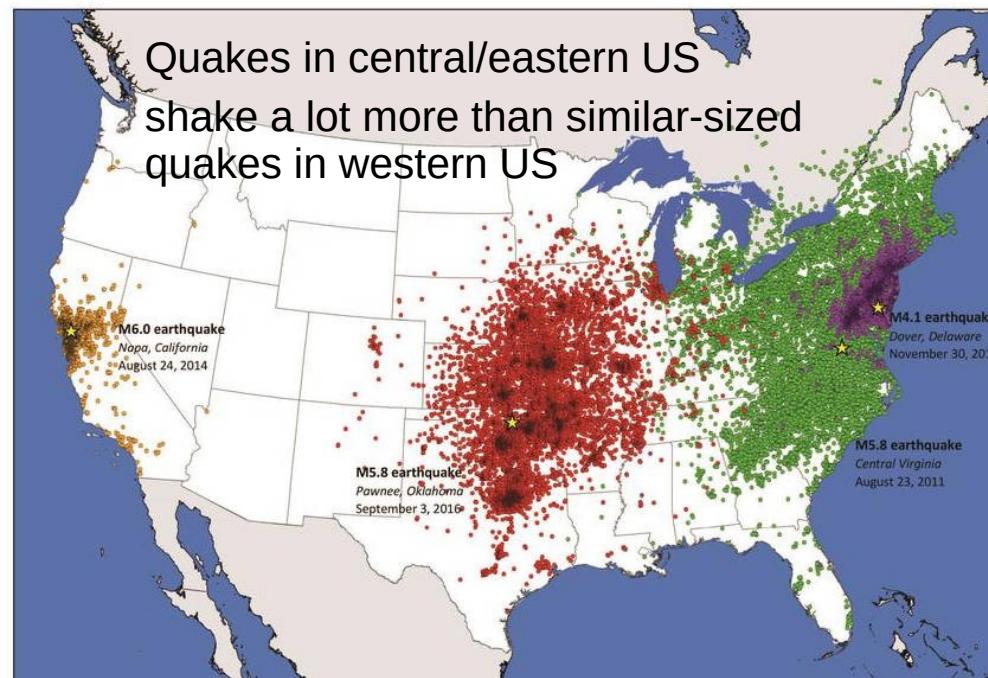
# Applications

# USArray – scattering strength (left) versus intrinsic attenuation (right)



Eulenfeld & Wegler 2017

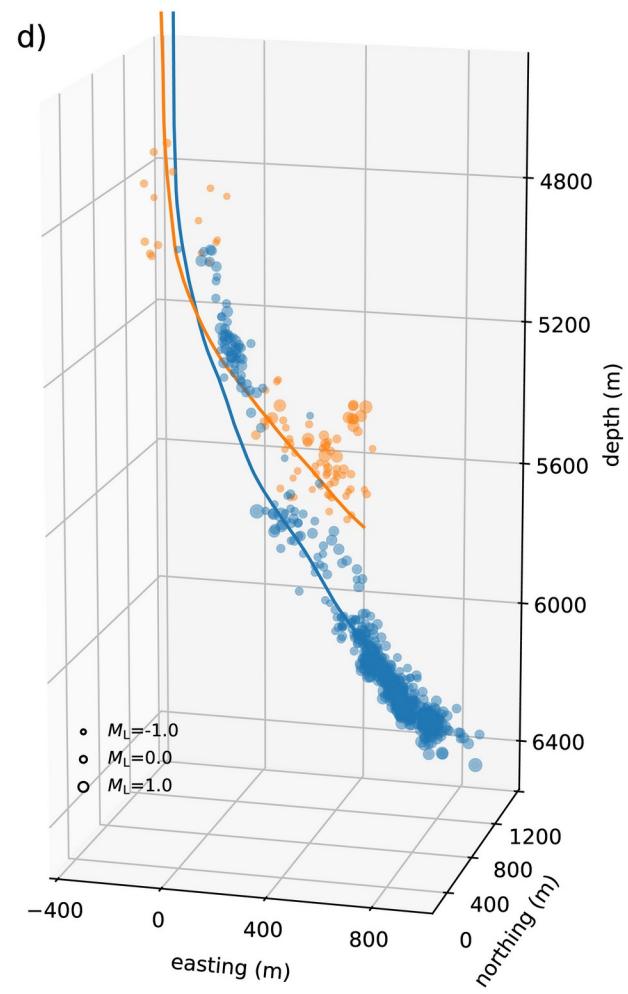
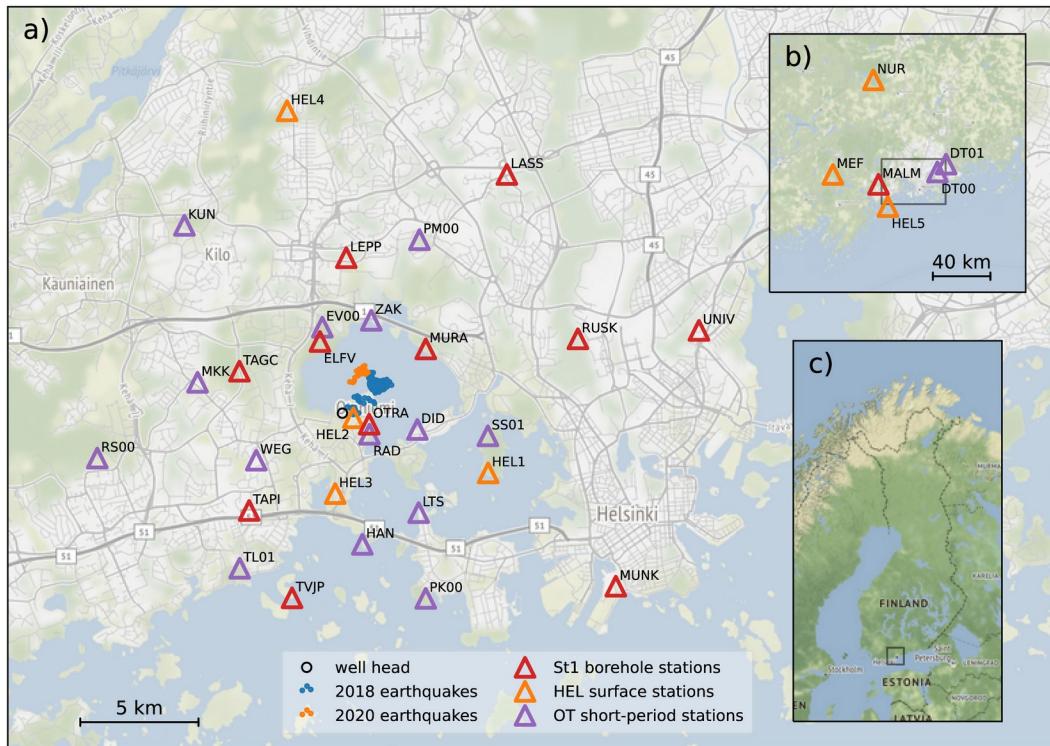
# Application USArray – high freq site amplification, magnitudes



Eulenfeld & Wegler 2017

# Helsinki 2018 and 2020 stimulation

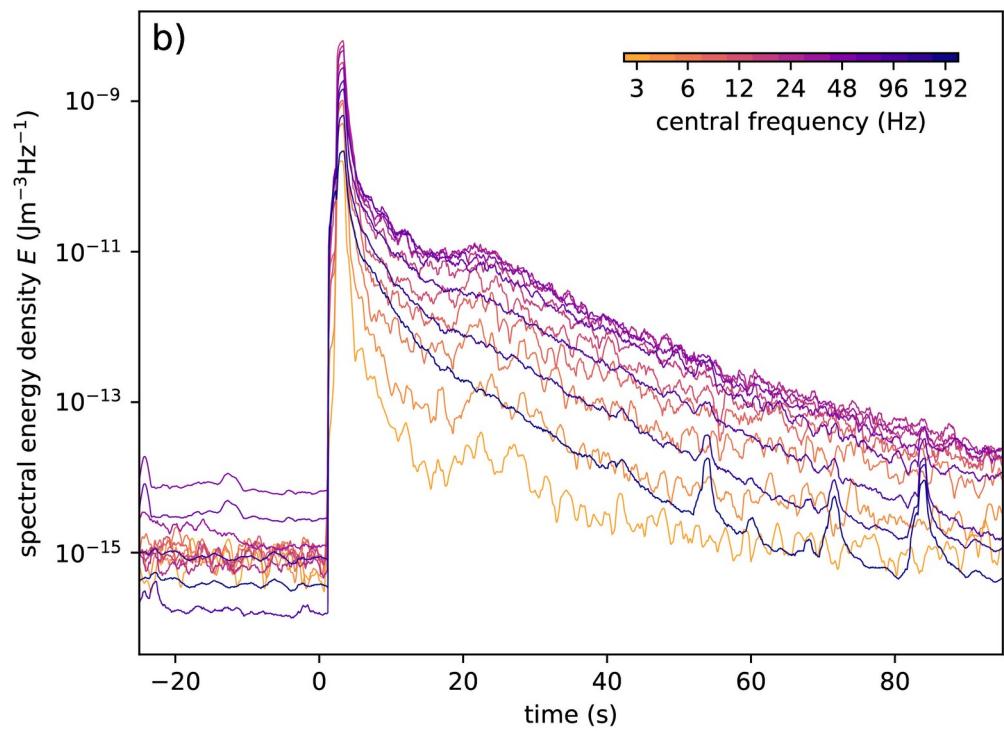
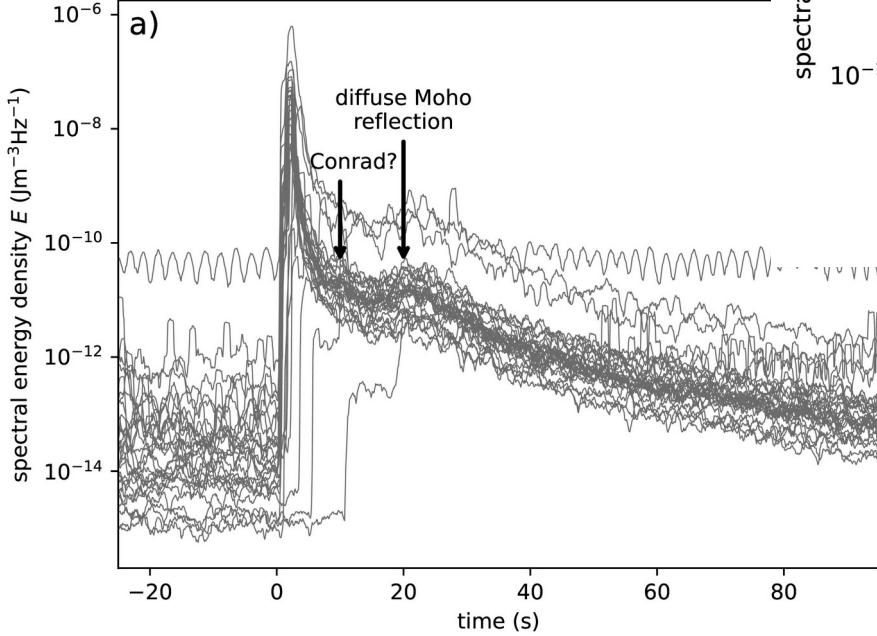
- 2018 stimulation induced ~450 earthquakes (blue) with  $0 \leq ML \leq 1.8$ ,  
90 MPa peak well-head pressure,  
18 000 m<sup>3</sup> volume
- 2020 stimulation induced ~25 earthquakes (orange)  
with  $0 \leq ML \leq 1.8$   
70 MPa, 2 900 m<sup>3</sup> volume



Eulenfeld et al. 2023

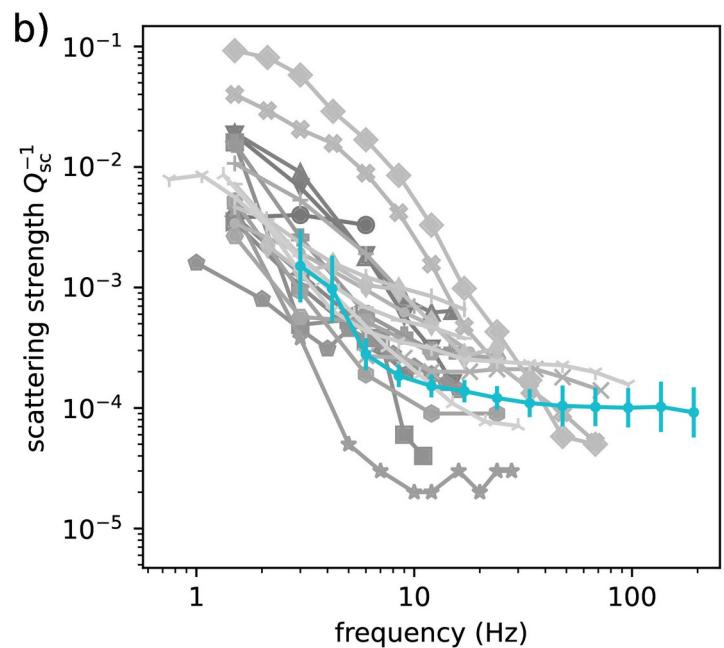
# Helsinki 2018 stimulation – example envelopes for 1 event

All stations 8 Hz – 16 Hz

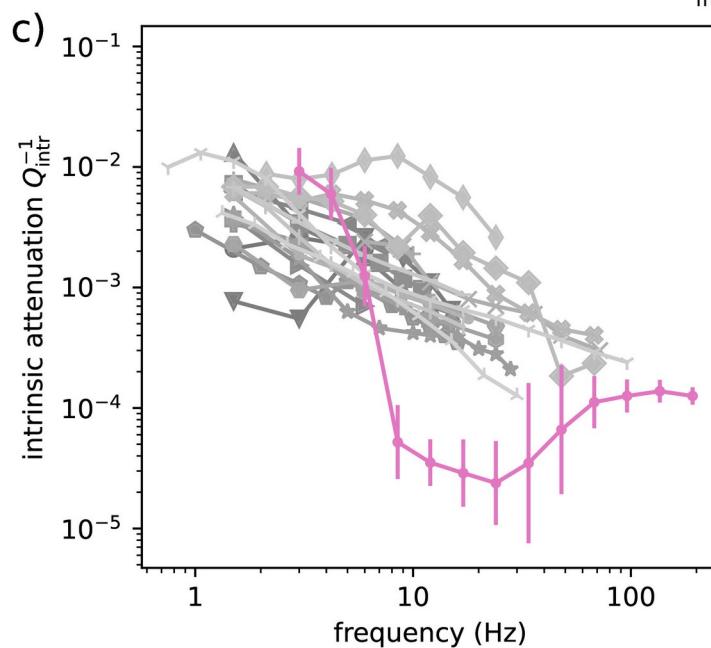


Single station,  
different frequency bands

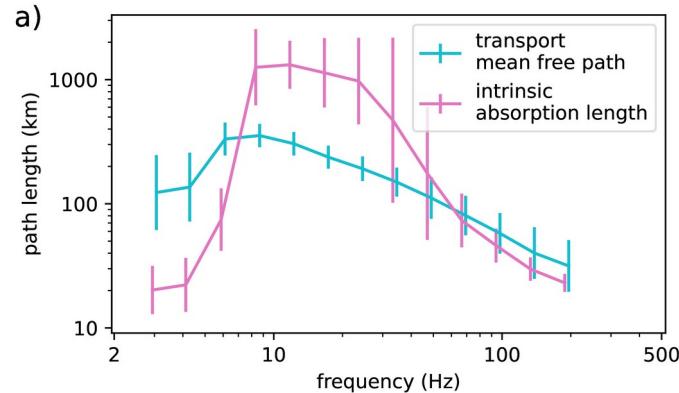
# Helsinki 2018 stimulation – Q values



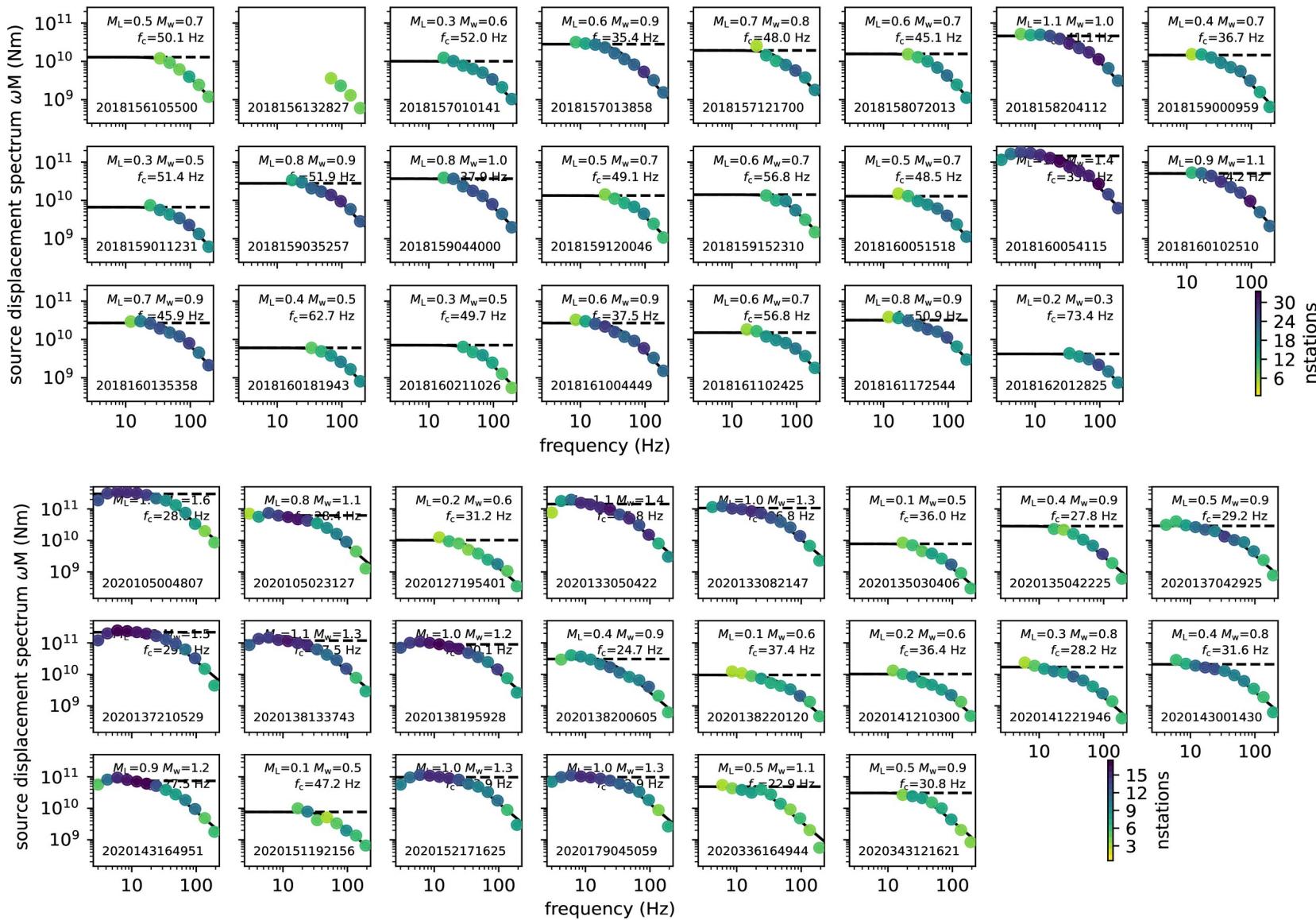
- Fehler et al. 1992, South-East Japan
- ▼ Mayeda et al. 1992, Hawaii
- ▲ Mayeda et al. 1992, Long Valley, California
- ◀ Canas et al. 1998, Canary Islands
- ▶ Hoshiba et al. 2001, Northern Chile
- Bianco et al. 2002, Southern Apennine, Italy
- ◆ Bianco et al. 2005, Northern Italy
- ✚ Giampicollo et al. 2006, Sicily, Italy
- ★ Ugalde et al. 2007, Western India
- ◆ Chung et al. 2009, Mt Fuji, Japan
- ◆ Chung et al. 2010, South Korea



- ✚ Ugalde et al. 2010, Galeras volcano, Colombia
- ✖ Fielitz & Wegler 2015, KTB, Germany
- Bachura & Fischer 2016, West Bohemia, Czechia
- ◆ Eulenfeld & Wegler 2016, Insheim, Upper Rhine Rift, Germany
- ◆ Eulenfeld & Wegler 2016, Landau, Upper Rhine Rift, Germany
- ◆ Eulenfeld & Wegler 2016, Unterhaching, Molasse basin, Germany
- ✚ Eulenfeld & Wegler 2017, Eastern USA
- ▬ Eulenfeld & Wegler 2017, Western USA
- ▬ Eulenfeld et al. 2021, Vogtland/West Bohemia, Germany/Czechia
- ▬ Hannemann et al. 2021, Eastern North Atlantic
- ✚ this study

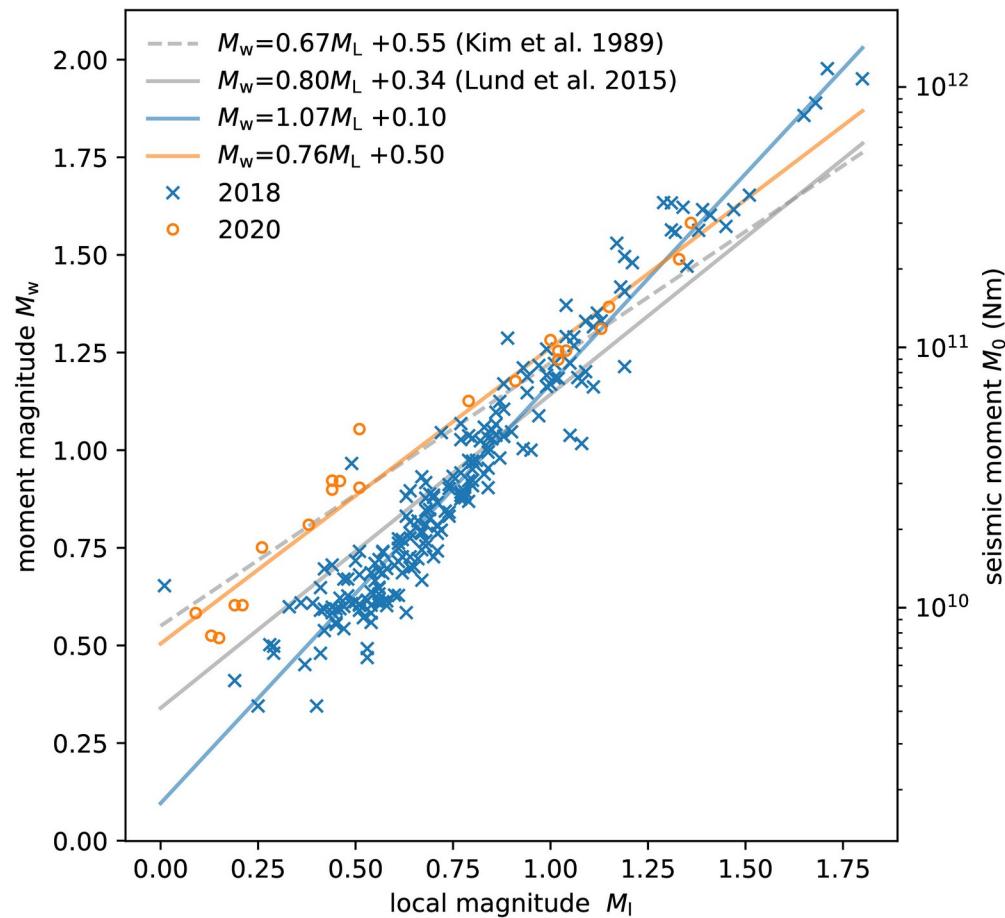


# Helsinki 2018/2020 – source displacement spectra



# Helsinki 2018/2020 – moment magnitudes

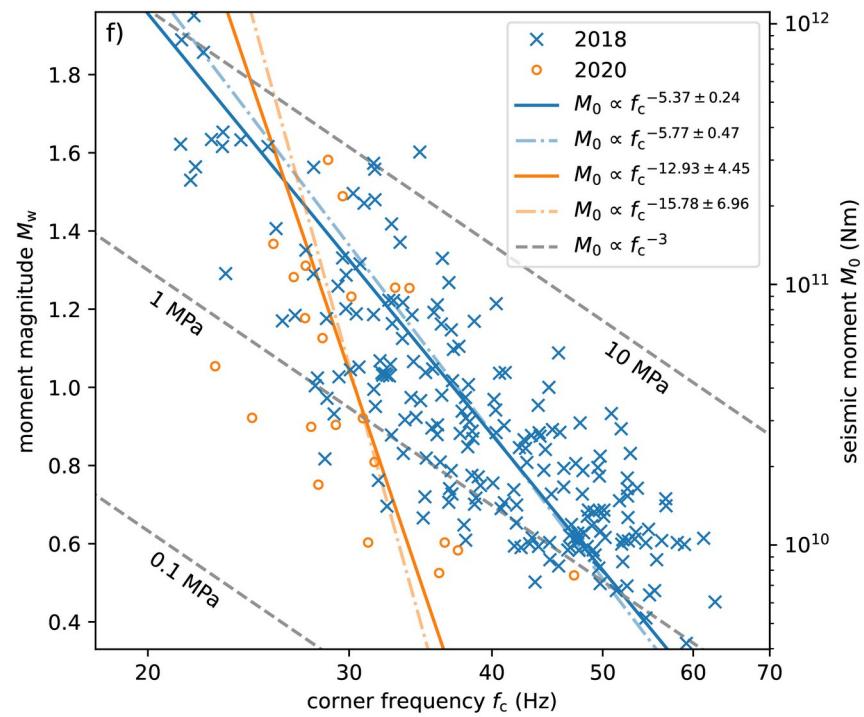
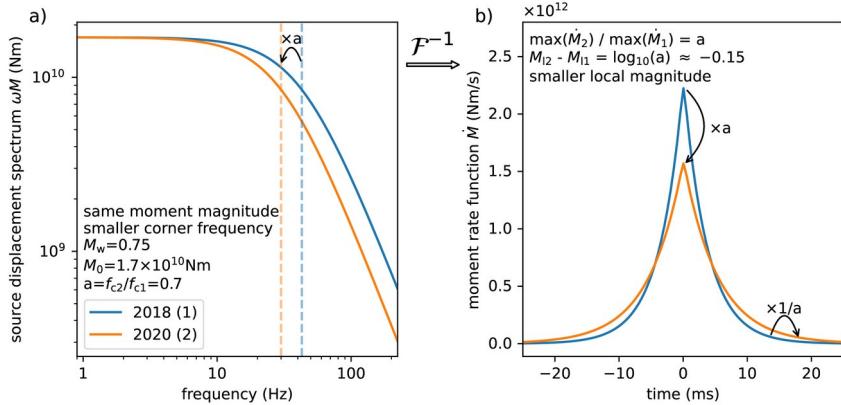
- $M_w$  versus  $ML$  relationship for the two stimulations
- 2020 events have systematically smaller  $ML$  for same  $M_w$  compared to 2018 events



# Helsinki 2018/2020 – source parameters

Eulenfeld et al. 2023

- Mw versus fc relationship for the two stimulations
- 2020 events have systematically smaller fc for same Mw compared to 2018 events
- Consistent with Mw-ML relationship



# Summary

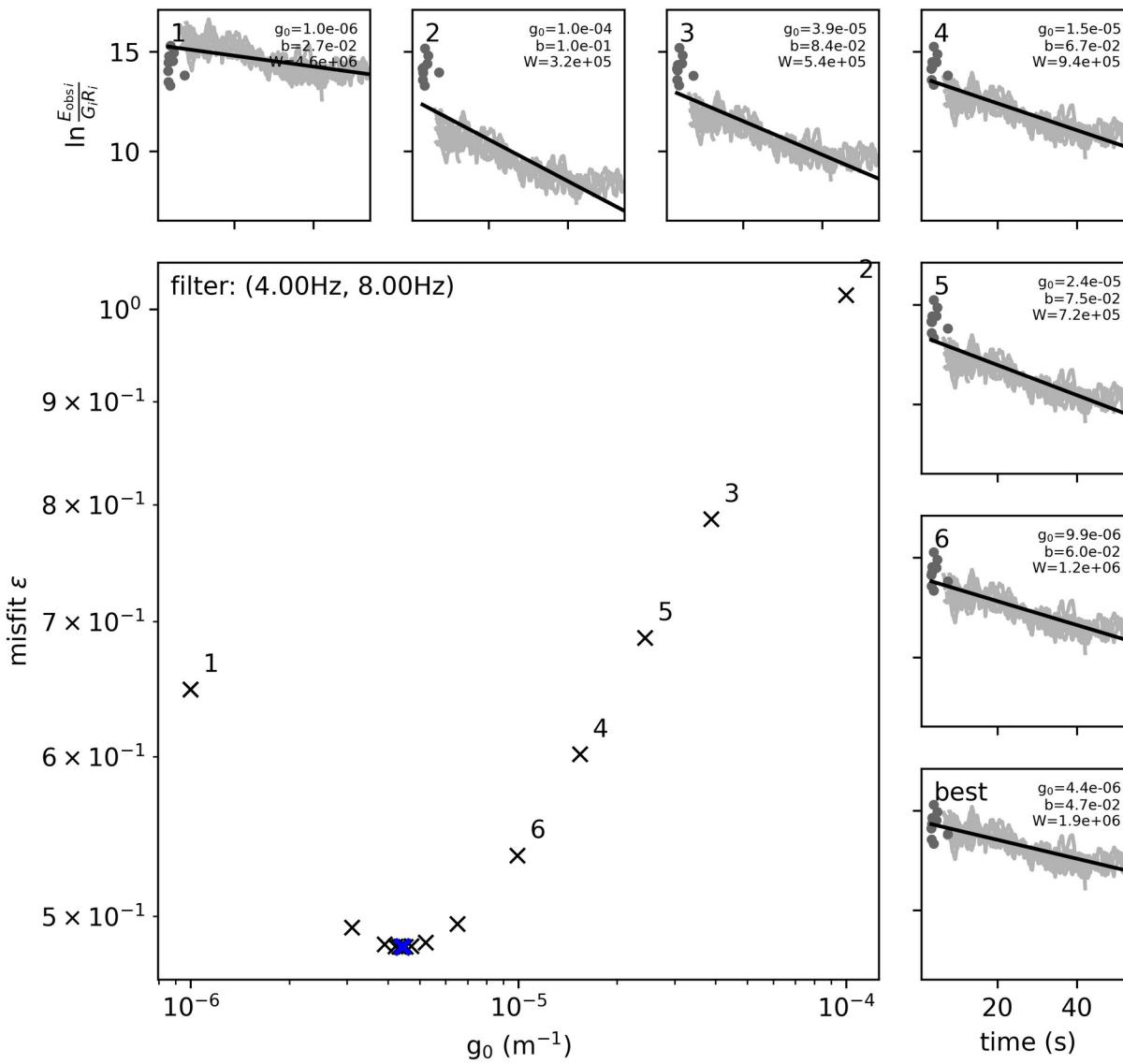
- Quickly estimate scattering and intrinsic attenuation parameters for your local data set
- Estimation of site responses (relative)
- Robust determination of moment magnitude and other source parameters

Thanks!

# References

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- Abercrombie R E (1995), Earthquake source scaling relationships from -1 to 5 ML using seismograms recorded at 2.5 km depth, Journal of Geophysical Research, doi: 10.1029/95JB02397

# Qopen optimization



Eulenfeld & Wegler 2016

# Helsinki 2018 stimulation – envelope fits example 16 – 32 Hz

