

Investigation of dipolar dominance in geodynamo simulations with different inner core sizes

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SUMMARY

The solid inner core of the Earth has been growing for approximately one billion years due to cooling of the Earth. The changing spherical shell geometry of the Earth's core is likely to influence on the geodynamo driven by convective motions in the fluid outer core. To understand the geometry effect on the dynamo regime through evolution of the core, we perform numerical simulations of geodynamo with three spherical shell radius ratios: $r_i/r_o = 0.15, 0.25$, and 0.35 , where r_i and r_o are the inner and outer core radii, respectively. To evaluate the morphology of the magnetic field, we examine two indices about dipole component dominance: (i) f_{dip} , dipolarity used to assess the relative strength of the dipole field at the core surface in numerical dynamo models; and (ii) $f_{mag-fit}$, the ratio of magnetic energy density for the dipole component to that extrapolated from the magnetic power spectrum for the high degree components. We investigate the field morphology estimated from f_{dip} , and $f_{mag-fit}$, and find that $f_{mag-fit}$ is valid to determine the dynamo regime, even if f_{dip} suggests a transition regime between dipolar and non-dipolar dominance. We also investigate the range of Rayleigh number for sustained dynamos based on f_{dip} and $f_{mag-fit}$, and find that the range of Rayleigh number for a dynamo characterized by a strong dipole field becomes narrower for a smaller inner core. The f_{dip} -dependence on the Rayleigh number for $r_i/r_o = 0.25$ and 0.35 is similar each other, whereas the $f_{mag-fit}$ -dependence for $r_i/r_o = 0.35$ is found to be relatively larger than that for $r_i/r_o = 0.25$. On the other hand, small values of f_{dip} and $f_{mag-fit}$ for $r_i/r_o = 0.15$ suggest that the dynamo regime is characterized not by a strong dipole field but by non-dipolar dominance. These results indicate that changes in the spherical shell radius ratio largely influence on the dynamo regime in numerical dynamos with a fixed temperature boundary condition.

1 INTRODUCTION

The Earth has an intrinsic magnetic field. The dynamo action of liquid iron alloy convection in the outer core generates the geomagnetic field. The fluid alloy gains buoyancy from the cooling of the Earth. Upon cooling, the inner core nucleated as liquid iron solidified from the center of the fluid core at high pressure. Compositional convection, which is associated with the growth of the inner core, is also a source of outer core convection. Recent thermochemical calculations suggest that the inner core which formed approximately one billion years ago and the inner core has been continually growing to its present size (Labrosse *et al.*, 2001). Although the size of the inner core has been changing across the geological time scale, the intensity of the geomagnetic field described by the virtual dipole moment (VDM) has maintained its present intensity for more than 3.5 billion years based on paleomagnetic observations (Biggin *et al.*, 2015); the geodynamo has been sustained during this period.

Previous studies have performed a number of numerical dynamo simulations under the assumption of the present geometry of the Earth's core; the aspect ratio of the inner core radius, r_i , to the outer core radius, r_o , is $r_i/r_o = 0.35$. For example, Christensen and Aubert (2006) revealed, in detail, sustained dynamo condi-

tions for various control parameters. They clarified dipolar dynamo cases and non-dipolar dynamo cases. The dynamo regime changes from stable dipolar to reversing non-dipolar with increase of the Rayleigh numbers (Kutzner and Christensen, 2002; Olson *et al.*, 2011).

In the smaller inner core setting, as compared with the present size, however, there have been a few attempts at a numerical dynamo. Some numerical dynamo studies have shown that the geometry effect on the dynamo regime is small. Hori *et al.* (2010) investigated the morphology of a magnetic field with fixed temperature (FT) and fixed heat flux (FF) boundary conditions for two spherical shell radius ratios: $r_i/r_o = 0.10$ and 0.35 . Regardless of the difference in radius ratios, they found that sustained dynamos were dipolar under the FF boundary condition and non-dipolar under the FT boundary condition. Driscoll (2016) carried out numerical simulations of geodynamo for eleven patterns of radius ratios in the range of $0.10 < r_i/r_o < 0.35$ and core power derived from a thermal evolution model. Driscoll (2016) found that the total magnetic energy in a spherical shell increased with increase of the ratios r_i/r_o and that sustained dynamos were characterized by a strong dipole magnetic field.

Other numerical dynamo studies have shown that the inner core size influences dipolar dominance. Heimpel *et al.* (2005) in-

investigated dynamo onset conditions for six spherical shell radius ratios: $0.15 < r_i/r_o < 0.65$. They found that the dipolar and total magnetic energy at the core–mantle boundary (CMB) decreases with decrease of r_i/r_o values for $r_i/r_o < 0.45$. Lhuillier *et al.* (2019) also reported on the effect of geometry on the dynamo regime. They performed chemically driven geodynamo simulations by changing ten patterns of radius ratios in the range of $0.10 < r_i/r_o < 0.44$. They found that sustained magnetic fields were dipolar for $r_i/r_o < 0.18$ and $r_i/r_o > 0.26$, whereas they were less dipolar for $0.20 < r_i/r_o < 0.22$. Although some studies have attempted to reveal the dependence of the dynamo regime on the spherical shell radius ratio, we do not yet fully understand how the morphology of the magnetic field is determined.

In recent numerical dynamos, dipolarity, f_{dip} , which is defined as the ratio of the dipole field strength to the total field strength at the CMB (Christensen and Aubert, 2006), has been widely used as an index for assessing the morphology of geomagnetic field. Christensen and Aubert (2006) mention that the magnetic field is dipolar-dominated when f_{dip} exceeds 0.35. This criterion for the dynamo regime is valid when dynamos are categorized into large and small f_{dip} groups (Soderlund *et al.*, 2012). However, this criterion is not valid when dynamos are not categorized by dipolarity (Aubert *et al.*, 2009).

The observed dipolar-dominated geomagnetic field can be expressed in terms of the magnetic power spectrum at the Earth's surface (Loves, 1974) and at the CMB (Langel and Estes, 1982). While Kono and Roberts (2002) compared a power spectrum of the observed geomagnetic field with that of numerical dynamos, there was a lack of quantitative evaluation of the dipolar dominance. As the dipolarity has no information of magnetic power spectrum distribution in higher degrees, we require not only the dipolarity, but also another index that represents dipolar dominance assessed from the spectrum distribution.

Although numerical dynamo simulations are useful tools to investigate magnetic field intensity and structure in the past Earth environment, previous studies have not yet established the criterion to evaluate dipolar dominance. The purpose of this study is to investigate the dynamo conditions of a sustained dipolar or non-dipolar dynamo for some spherical shell radius ratios based on an evaluation of dipolar dominance. Hence, we carried out numerical simulations of geodynamo for three spherical shell radius ratios, i.e., $r_i/r_o = 0.15, 0.25$, and 0.35 . To focus on how convection occurs, the Rayleigh number (Ra) was only treated as a variable. The Ra is a parameter related to buoyancy, which is the driving force of convection. By simulating a wider range in the Ra than previous studies, we can compare cases of a small inner core size setting with those of the present size. A combination of the dipolarity at the CMB, as well as the magnetic energy spectrum at the CMB in the spherical harmonic degree expansion, reveal the range in the Ra in the sustained dipolar or non-dipolar dynamo for each radius ratio.

2 METHOD

A numerical geodynamo model is given by an electrically conducting Boussinesq fluid in a rotating spherical shell. The governing equations of the geodynamo in the outer core are described by the momentum equation, heat equation, continuity equation, magnetic induction equation, and Gauss's law for the magnetic field, which

are respectively given as

$$\frac{\partial \mathbf{u}}{\partial t} + (\boldsymbol{\omega} \times \mathbf{u}) = -\nabla \left(P + \frac{1}{2} u^2 \right) - \nabla \times \nabla \times \mathbf{u} - \frac{2}{E} (\hat{\mathbf{z}} \times \mathbf{u}) + \frac{Ra}{Pr} T \frac{\mathbf{r}}{r_o} + \frac{1}{PmE} (\mathbf{J} \times \mathbf{B}), \quad (1)$$

$$\nabla \cdot \mathbf{u} = 0, \quad (2)$$

$$\frac{\partial T}{\partial t} + (\mathbf{u} \cdot \nabla) T = \frac{1}{Pr} \nabla^2 T, \quad (3)$$

$$\frac{\partial \mathbf{B}}{\partial t} = -\frac{1}{Pm} \nabla \times \nabla \times \mathbf{B} + \nabla \times (\mathbf{u} \times \mathbf{B}), \quad (4)$$

and

$$\nabla \cdot \mathbf{B} = 0, \quad (5)$$

where \mathbf{u} , P , T , \mathbf{B} , and $\hat{\mathbf{z}}$ are the velocity, pressure, temperature, magnetic field, and unit vector along the rotation axis, respectively. And, $\boldsymbol{\omega} = \nabla \times \mathbf{u}$ and $\mathbf{J} = \nabla \times \mathbf{B}$ are the vorticity and current density, respectively. In Eqs. (1) – (5), the length and temperature are normalized by the outer core thickness, $L = r_o - r_i$ and average temperature difference between the inner core boundary (ICB) and CMB, ΔT , respectively. The time is normalized by the kinematic viscosity diffusion time, L^2/ν , where ν is the kinetic viscosity, and the magnetic field is normalized by $\sqrt{\rho_0 \mu_0 \eta \Omega L}$, where ρ_0 , μ_0 , η , and Ω are the average density, magnetic permeability, magnetic diffusivity, and angular velocity of system's rotation, respectively. The Rayleigh number, Ra , Ekman number, E , Prandtl number, Pr , and magnetic Prandtl number, Pm are defined as follows:

$$Ra = \frac{\alpha g \Delta T L}{\nu \Omega}, E = \frac{\nu}{\Omega L^2}, Pr = \frac{\nu}{\kappa}, \text{ and } Pm = \frac{\nu}{\eta}. \quad (6)$$

We used the numerical dynamo code Calypso (Matsui *et al.*, 2014). In Calypso, the spherical harmonic expansion method is used in the horizontal discretization, and the second-order finite differences are used in the radial discretization. For the time integrations, the Crank-Nicolson method was used in the linear diffusive terms and the second order Adams-Bashforth method was used in the other terms.

At the initial condition, temperature perturbation was applied to all the sectorial modes. The initial magnetic field was set as an axial dipole based on Christensen *et al.* (2001). For the boundary condition, the temperatures at the CMB and ICB were fixed as $T(r_o) = 0$ and $T(r_i) = 1$. The mantle and inner core were assumed to be co-rotating, and a non-slip boundary ($\mathbf{u} = 0$) was applied to the CMB and ICB. The mantle and inner core were assumed to be electrically insulated, and the magnetic field at the boundaries was connected to the potential field. In the parameter setting, Ra was changed among the cases; the Ekman, Prandtl, and magnetic Prandtl numbers were fixed at $E = 10^{-3}$, $Pr = 1$, and $Pm = 5$ in all simulation cases. The truncation of the spherical harmonics and the radial grid points were set to $L_{max} = 47$ and $N_r = 63$, respectively. To eliminate aliasing in the spherical harmonic expansion, horizontal grids were set to $(N_\theta, N_\phi) = (72, 144)$. To investigate the effects of different inner core sizes, the spherical shell radius ratios of the inner core radius to the outer core radius were set as $r_i/r_o = 0.15, 0.25$, and 0.35 . First, we performed numerical simulations of non-magnetic thermal convection in rotating spherical shells to estimate the critical Rayleigh number, Ra_{crit} , for the onset of thermal convection. We then carried out numerical simulations of magnetohydrodynamic (MHD) dynamos driven by thermal convection. **Dynamo simulations with a lower Ekman number are necessary, however, we were forced to perform dynamo simulations with $E = 10^{-3}$ to inves-**

investigate various Rayleigh numbers and inner core radii for computational limitation like dynamo simulations of Lhuillier *et al.* (2019), in which the Ekman number is larger than $E = 10^{-3}$.

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