

1 **Assessing the Impact of Cumulus Convection and**
2 **Turbulence Parameterizations on Typhoon**
3 **Precipitation Forecast**

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10 **Key Points:**

- 11 • Applying the cumulus and RNA schemes improves the ability to catch heavy rain-
12 fall with higher recall scores.
- 13 • Employing the cumulus and RNA schemes can help maintain the compact struc-
14 ture and strength of typhoons.
- 15 • Considering the subgrid-scale turbulence can optimize the dissipation and backscat-
16 ter configuration to enhance deep convection.

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17 **Abstract**

18 Improving typhoon precipitation forecast with convection-permitting models re-
 19 mains challenging. This study investigates the influence of cumulus parameterizations
 20 and turbulence models, including the Reconstruction and Nonlinear Anisotropy (RNA)
 21 turbulence scheme, on precipitation prediction in multiple typhoon cases. Incorporat-
 22 ing the cumulus and RNA schemes increases domain-averaged precipitation, improves
 23 recall scores, and lowers relative error across various precipitation thresholds, which is
 24 substantial in three out of four studied typhoon cases. Applying appropriate cumulus
 25 parameterization schemes alone also contributes to enhancing heavy precipitation fore-
 26 casts. In Typhoon Hato, the RNA and Grell-3 schemes demonstrated a doubled recall
 27 rate for extreme rainfall compared to simulations without any cumulus scheme. The im-
 28 proved forecasting ability is attributed to the RNA's capacity to model dissipation and
 29 backscatter. The RNA scheme can dynamically reinforce typhoon circulation with up-
 30 gradient momentum transport in the lower troposphere and enhance the buoyancy by
 31 favorable heat flux distribution, which is conducive to developing heavy precipitation.

32 **Plain Language Summary**

33 Enhancing the forecast accuracy of typhoon-induced rainfall prediction with nu-
 34 matical weather prediction models is still challenging. This study focused on the impact
 35 of cumulus convection schemes and a new turbulence scheme named the Reconstruction
 36 and Nonlinear Anisotropy (RNA) scheme on the precipitation forecast performance when
 37 typhoons hit. We found that the convection and the RNA schemes help predict more
 38 rain on average and make our predictions more accurate, especially regarding heavy rain-
 39 fall. Still, it also leads to an overestimation of the precipitation. In addition, applying
 40 the cumulus and RNA scheme is beneficial in keeping the typhoon structure and inten-
 41 sity at a lower sea level pressure. This improvement in generating intense convections
 42 is due to the optimized configuration of the dissipation and backscattering caused by the
 43 subgrid-scale turbulence.

44 **1 Introduction**

45 Tropical cyclones frequently affect South China, causing extreme precipitation and
 46 winds leading to landslides and flooding, resulting in substantial economic damage and
 47 loss. Despite progress in numerical weather prediction, accurately forecasting typhoon
 48 precipitation intensities remains challenging. Cumulus parameterization, which repre-
 49 presents subgrid convection, is crucial in precipitation forecasting. Previous studies suggest
 50 that disabling the cumulus scheme is appropriate when the grid scale is less than 4 km,
 51 as the explicit microphysics scheme and model dynamics are expected to resolve cloud
 52 and precipitation processes (Weisman et al., 1997; Skamarock et al., 2008). However, whether
 53 cumulus parameterization should be applied at kilometer-scale resolution remains con-
 54 troversial in the tropical cyclone community, because this resolution falls within the grey
 55 zone, where both resolved and subgrid processes can contribute to turbulence (Wyngaard,
 56 2004; Gerard, 2007; Boulte et al., 2014; Shi, Chow, et al., 2019). Sun et al. (2013, 2014)
 57 performed sensitivity experiments to analyze the simulated Tropical Cyclone (TC) in-
 58 tensity for Typhoon Shanshan using different cumulus parameterization schemes under
 59 the grey-zone resolution by varying their resolution from 4 to 10 km. They suggest that
 60 a suitable cumulus scheme can enhance tropical cyclone convergence. Conversely, Yu and
 61 Lee (2011) discovered that simulations would overpredict the area-averaged precipita-
 62 tion rate without employing convective parameterization. Recent studies further indi-
 63 cated that the scale-aware cumulus scheme can improve precipitation prediction (Mahoney,
 64 2016; Gao et al., 2017). Shi and Wang (2022) demonstrated that simulations without
 65 cumulus schemes underestimate precipitation and overall performance for extreme rain-
 66 fall prediction. Given the uncertain impacts of utilizing cumulus schemes in high-resolution

67 simulations on typhoon precipitation prediction, further evaluation of the necessity and
 68 effect of cumulus convection in kilometer-scale simulations with additional typhoon cases
 69 is needed.

70 Previous studies pointed out that the equivalent potential temperature exhibits a
 71 significant horizontal gradient in TCs, indicating that subgrid-scale mixing should be con-
 72 sidered in high-resolution tropical cyclone simulations (Houze Jr, 2014). Although the
 73 traditional planetary boundary layer (PBL) scheme remains valid for subgrid-scale tur-
 74 bulence at the kilometer-scale resolution, with the grey zone bound being $\geq 100\text{m}$ for
 75 the PBL scheme, horizontal subgrid-scale turbulence is not accounted for in conventional
 76 PBL schemes or cumulus parameterizations, assuming the environment is horizontally
 77 homogeneous at subgrid scales.

78 In the Weather Research and Forecasting (WRF) model, the horizontal turbulence
 79 can be represented by a gradient-diffusion scheme, such as the two-dimensional Smagorin-
 80 sky model (Zhou et al., 2017). However, the Smagorinsky scheme does not allow backscat-
 81 ter, which is observed according to in situ measurements and LES simulation results (Shi
 82 et al., 2018; Carper & Porté-Agel, 2004). Chow et al. (2005) developed the dynamic re-
 83 construction model (DRM) of turbulence based on an explicit filtering framework, di-
 84 viding the subfilter-scale turbulence flux into resolvable subfilter-scale (RSFS) and subgrid-
 85 scale (SGS) components. Shi and Wang (2022) replaced the SGS part with the nonlin-
 86 ear backscatter and anisotropy (NBA) model and applied it to represent horizontal tur-
 87 bulence, their results for simulating Typhoon Vicente indicated that it can enhance the
 88 precipitation with the optimal configuration of dissipation and backscattering. Never-
 89 theless, studies examining the effects of cumulus schemes and horizontal turbulence on
 90 typhoon precipitation at the grey-zone scale are still limited. In this study, we further
 91 investigate the performance and necessity of considering vertical and horizontal turbu-
 92 lence mixing at kilometer-scale resolution by testing the impact of a cumulus parame-
 93 terization and RNA scheme on precipitation forecasting with multiple typhoon cases.

94 2 Methods and Experiment Design

95 2.1 Turbulence Schemes

96 The horizontal stress in the Smagorinsky scheme is represented as

$$\tau_{ij} = -K_h D_{ij} \quad (1)$$

97 the K_h and D_{ij} are the horizontal eddy viscosity and deformation tensor, respectively.
 98 In the WRF model, the turbulent scalar flux has a similar expression as Eq. (1), with
 99 the scalar diffusivity being divided by the turbulent Prandtl number $\text{Pr} = 1/3$.

100 In the RNA scheme, the subfilter-scale turbulence stress is (Shi, Chow, et al., 2019):

$$\tau_{ij} = \tau_{ij}^{\text{RSFS}} + \tau_{ij}^{\text{SGS}} \quad (2)$$

101 The RSFS is computed by adopting the explicit filtering-based RSFS model of DRM (Chow
 102 et al., 2005; Kirkil et al., 2012; Shi et al., 2018) as

$$\tau^{\text{RSFS}} = \overline{u_i^* u_j^*} - \overline{u_i^*} \overline{u_j^*} \quad (3)$$

103 Following the approximate deconvolution method(ADM) (Stolz et al., 2001; Stolz & Adams,
 104 1999), the u^* is:

$$u^* = \bar{u}_i + (I - G)\bar{u}_i + (I - G)[(I - G)\bar{u}_i] + \dots, \quad (4)$$

105 where I is the identity operator and G is the explicit filter. The reconstructed velocity
 106 retains the first term only and is estimated as the grid velocity; the overbar denotes a
 107 top-hat filter.

108 The nonlinear backscatter and anisotropy (NBA) model is adopted here to consider
 109 the backscattering effect (Kosović, 1997; Mirocha et al., 2010; Shi et al., 2018). There-
 110 fore, the SGS term in the DRM model is expressed as:

$$\tau_{ij}^{\text{SGS}} = -C_s'^2 l^2 \left[2(2S_{mn}S_{mn})^{1/2} S_{ij} + C_1 (S_{ik}S_{kj} - S_{mn}S_{mn}\delta_{ij}/3) + C_2 (S_{ij}R_{kj} - R_{ik}S_{kj}) \right] \quad (5)$$

111 where S_{ij} , R_{ij} , δ_{ij} represents the resolved strain rate tensor, resolved rotation rate tenor,
 112 and Kronecker delta, respectively. The constants followed Mirocha et al. (2010). We fur-
 113 ther conducted the simulations based on these different turbulence schemes together with
 114 the cumulus convection schemes.

115 2.2 Experiment Design

116 The study evaluates the impact of cumulus and turbulence parameterization schemes
 117 on precipitation forecasts for three typhoon events, including Typhoon Mujigae (2015),
 118 Typhoon Hato (2017) and Typhoon Mangkhut (2018) using the WRF model, and the
 119 impact on intense precipitation predictions was relatively significant for the first two cases.
 120 We conducted simulations on three nested domains with grid resolutions of 15 km, 5 km,
 121 and 1.67 km, respectively; the model top is at 50 hPa with 50 vertical levels. Support-
 122 ing Figure S1 displays the WRF domain configuration. The ECMWF Fifth-Generation
 123 Reanalysis (ERA5) was employed as the initial and boundary conditions for the WRF
 124 model. We conducted distinct simulations with or without employing the Grell-Freitas
 125 (Grell-3) cumulus scheme for cumulus convection associated with different horizontal tur-
 126 bulence schemes. For vertical turbulent mixing in the PBL, the ACM2 is applied in all
 127 the simulations (Pleim, 2007). Table 1 lists the eight simulations for each typhoon case,
 128 and Supporting Table 1 provides the detailed configuration of the simulations for each
 129 typhoon case. The Grell-Freitas scheme (Grell & Freitas, 2014; Freitas et al., 2020) is
 130 suitable both for the coarse and kilo-meter scale resolution as it is a scale-aware scheme
 131 based on the method described by Arakawa et al. (2011); the Grell-3 scheme is a con-
 132 ventional cumulus scheme based on the Grell–Devenyi ensemble scheme and can spread
 133 subsidence effects to neighboring grid columns and is also suitable for high-resolution ty-
 134 phoon simulations (Grell & Dévényi, 2002).

135 Specifically, for the GF-GF-R (G3-G3-R) simulations, the horizontal turbulence scheme
 136 which is referred to as the RNA scheme is in conjunction with the cumulus convective
 137 scheme; for the GF-GF-S (G3-G3-S) simulations, the Smagorinsky scheme was applied,
 138 while no horizontal turbulence scheme was used in the GF-GF-N (G3-G3-N) simulations.
 139 In the GF-N-N (G3-N-N) simulations, neither a cumulus nor a horizontal mixing tur-
 140 bulence scheme was activated in the two inner domains; in the GF-N-R (G3-N-R) sim-
 141 ulations, the RNA scheme is turned on for the horizontal turbulence scheme and no cu-
 142 mulus scheme is used. In this study, we evaluated the impact of these two convective schemes
 143 on typhoon precipitation forecasts. The cumulus convection scheme applied in the in-
 144 nermost domains is consistent with the scheme used in the outermost domain. Given that
 145 the impacts are comparatively subtle in Typhoon Mangkhut, we focus on Typhoon Mu-
 146 jigae and Typhoon Hato to show the effects of convection and turbulence schemes in the
 147 following analysis.

148 3 Results

149 3.1 Precipitation Forecast Evaluation

150 We compared the average precipitation from the ten simulations in the innermost
 151 domain for each typhoon case with the observation from the 1303 ground-based stations
 152 in Guangdong province. The GF-N-R (G3-N-R) simulations produced more domain-averaged
 153 precipitation compared to the GF-N-N (G3-N-N) by applying RNA scheme separately
 154 (Figure 1a). The GF-N-R simulation for Typhoon Mujigae generated 58mm accumulated

Table 1. Experiment design for each typhoon case with different schemes

Simulation	Cumulus Scheme(outer domain)	Cumulus Scheme(inner domains)	Horizontal Turbulence Scheme
GF-GF-R	Grell–Freitas	Grell–Freitas	RNA
GF-GF-S	Grell–Freitas	Grell–Freitas	Smagorinsky
GF-GF-N	Grell–Freitas	Grell–Freitas	None
GF-N-N	Grell–Freitas	None	None
GF-N-R	Grell–Freitas	None	RNA
G3-G3-R	Grell-3	Grell-3	RNA
G3-G3-S	Grell-3	Grell-3	Smagorinsky
G3-G3-N	Grell-3	Grell-3	None
G3-N-N	Grell-3	None	None
G3-N-R	Grell-3	None	RNA

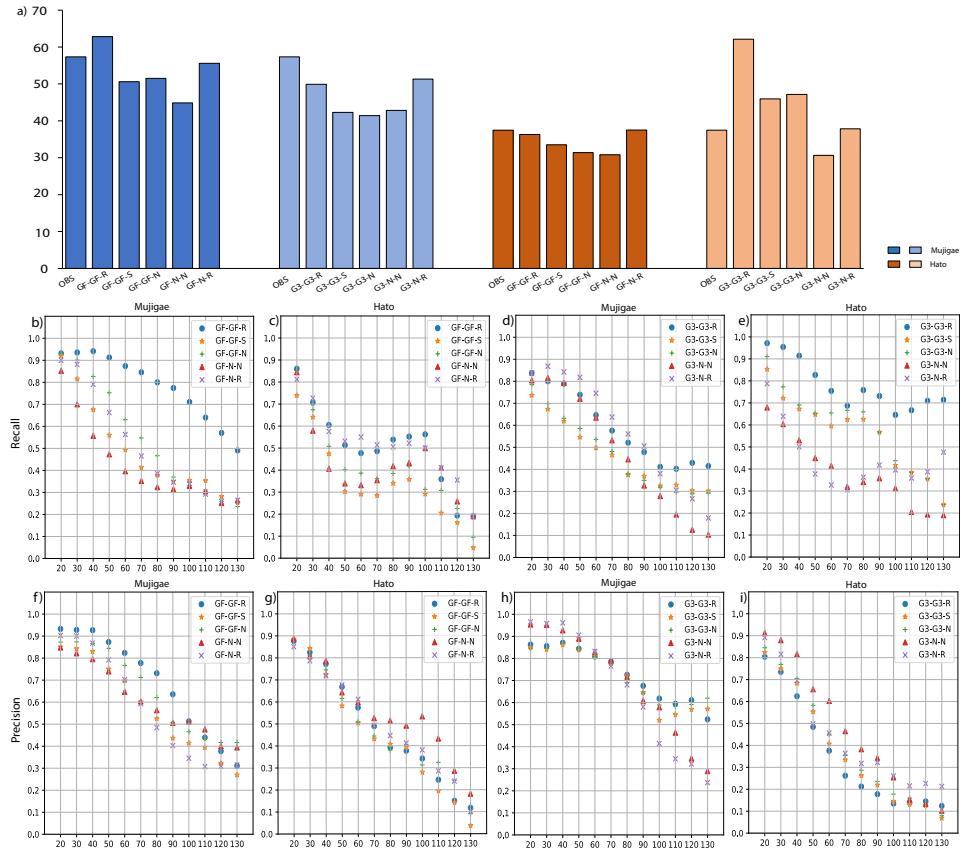


Figure 1. The averaged 24-hour accumulated precipitation in the innermost domain for Typhoon Mujigae and Typhoon Hato. The recall (b-e) and precision (f-i) scores for the 24-hour accumulated precipitation for simulations with different cumulus and turbulence schemes over observation at 1303 stations for Typhoon Mujigae (b,d,f,h) and Typhoon Hato (c,g,e,i) at different thresholds from 20 to 130 mm. (b)(c)(f)(g): Grell–Freitas scheme, (d)(h)(e)(i): Grell-3 scheme. Different turbulence schemes are shown using different symbols: blue dots represent the cumulus and RNA schemes; yellow stars represent the cumulus and Smagorinsky schemes; green crosses represent without applying a horizontal scheme, red triangles represent without applying cumulus and horizontal schemes and purple crosses represent applying the RNA scheme.

155 precipitation which is close to the observation. In addition, the GF-GF-R and G3-G3-
 156 R simulations, which account for the cumulus and horizontal subgrid-scale turbulence,
 157 produced higher domain-averaged precipitation amounts compared to the other simu-
 158 lations without applying the cumulus schemes or RNA scheme. Specifically, in the Mu-
 159 jigae case, the GF-GF-R simulation exhibited 48% more precipitation than the GF-N-
 160 N simulation and 24% more than the GF-GF-S (GF-GF-N) simulation. Moreover, ap-
 161 plying the Smagorinsky scheme did not significantly impact the typhoon precipitation
 162 amount as demonstrated by similar domain-averaged precipitation in the GF-GF-N and
 163 G3-G3-N simulations. We also analyzed the distribution of 12-hour accumulated pre-
 164 cipitation to examine the pattern of intense precipitation under different conditions, Sup-
 165 porting Figure S2 shows the results for Typhoon Mujigae. Although the overall typhoon
 166 structure in simulations using various schemes is similar, subtle differences exist in the
 167 rain band. The rain band is more compact, and the coverage of intense precipitation is
 168 more extensive in simulations that activate cumulus parameterization.

169 The recall and precision score for the 24-hour precipitation over 1303 stations in
 170 Guang Dong Province were calculated for the typhoon cases at different thresholds, from
 171 20mm to 130mm. Recall denotes the ratio of correctly predicted extreme events to the
 172 actual occurrence of extreme precipitation, which measures the fraction of true-positive
 173 stations experiencing extreme events; precision represents the ratio of correctly predicted
 174 extreme events to simulated occurrences of extreme precipitation. The Typhoon Muji-
 175 gae and Hato cases' precision and recall scores in simulations with different cumulus and
 176 horizontal turbulence schemes are shown in Figure 1b-i. Focusing on the RNA scheme
 177 effect on the recall scores for the two cases, we found simulations applying the RNA scheme
 178 produced higher recall scores compared to simulations without applying any cumulus or
 179 horizontal turbulence scheme, showing a higher ability to catch the precipitation events
 180 in most cases, especially for the extreme precipitation events. The application of the Grell-
 181 3 or Grell-Freitas cumulus scheme simultaneously associated with the RNA scheme gen-
 182 erated higher recall scores in most cases, especially in the threshold range of 40-100 mm,
 183 demonstrating the advantage in improving the hit rate of strong convection. For the Ty-
 184 phoon Mujigae case, the GF-GF-R simulation (Figure 1a) produced the highest recall
 185 score at all the thresholds compared with other simulations, displaying a three-times in-
 186 crease in recall compared to the GF-N-N simulation at the threshold of 80mm. For the
 187 Typhoon Hato case, the difference in the recall score between the GF-GF-R and the GF-
 188 GF-N was less than 0.1 when the precipitation was less than 40mm, and increased to 0.4
 189 when accumulated precipitation exceeded 90mm. In Figure 1d, in which the cumulus scheme
 190 is Grell-3, applying the RNA scheme showed significant advantages over simulations with-
 191 out applying the RNA scheme across all the thresholds. For Typhoon Mangkhut, apply-
 192 ing the cumulus and RNA turbulence schemes showed limited effects on the precipita-
 193 tion simulation (Supporting Figure S6).

194 The impacts of the configuration of the RNA scheme with different cumulus schemes
 195 are inconsistent. In the Typhoon Mujigae case, the GF-GF-R performs better than the
 196 G3-G3-R in the Mujigae case, the Grell-Freitas scheme shows a 60% increase in recall
 197 compared to the Grell-3 scheme for the Mujigae case as shown in Figure 1b,d. The op-
 198 posite result is found in the Typhoon Hato case. The simulation applying the Grell-3
 199 scheme with the RNA scheme shows higher recall scores; the simulated precipitation in
 200 simulations using the Grell-Freitas scheme is comparable to those without applying the
 201 cumulus scheme. The results can be attributed to the Grell-Freitas scheme's reduced sen-
 202 sitivity to model resolution, leading to proportionately less precipitation at finer reso-
 203 lutions. The Grell-3 scheme is more sensitive to model resolution and produces more pre-
 204 cipitation which aligns with findings from previous studies (Li et al., 2011). Concern-
 205 ing the precision score for Typhoon Mujigae (Figure 1f), the RNA scheme outperformed
 206 other schemes in predicting intense precipitation, accurately forecasting heavy rainfall
 207 lower than 90mm. However, the RNA scheme lowered the precision scores above 100mm.
 208 In conclusion, utilizing the cumulus and RNA schemes resulted in more accurate pre-

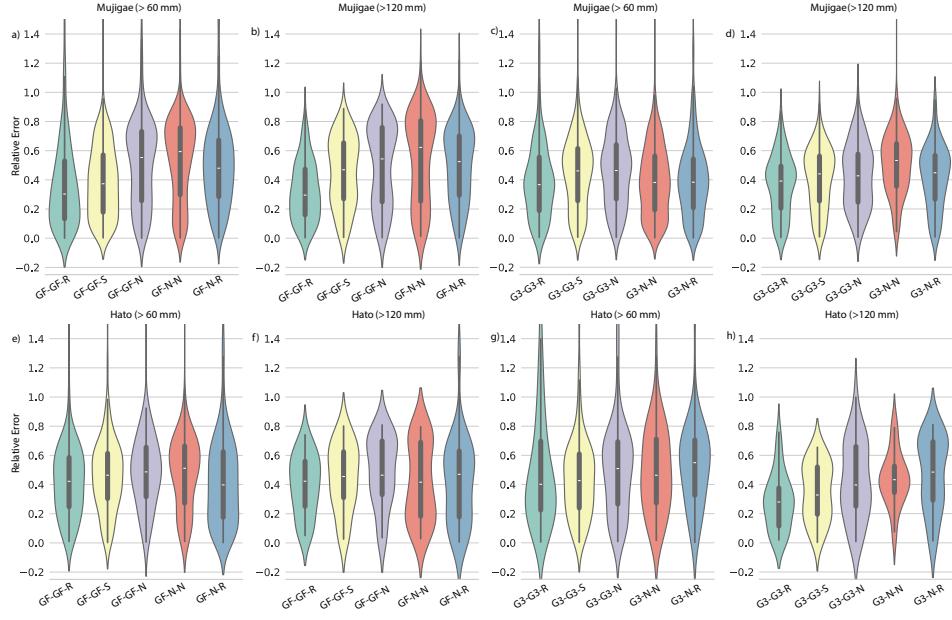


Figure 2. The relative error for the 24-hour accumulated precipitation for the different simulations over observation at 1303 stations for Typhoon Mujigae (a-d) and Typhoon Hato (e-h) at different thresholds (a)(c)(e)(g) 60mm, (b)(d)(f)(h) 120mm.

dictions of heavy rainfall, thereby improving the overall recall scores. However, it may also overestimate precipitation at some locations, leading to lower precision scores.

The relative error between the simulated accumulated precipitation and observed precipitation was calculated at thresholds of 60mm and 120mm to estimate the precipitation forecast performance for the different turbulence scheme configurations (Figure 2). Overall, the simulations using the RNA scheme for the horizontal turbulence show higher accuracy than others by decreasing the median relative error values at all the thresholds with both cumulus schemes. Simulations integrating the cumulus and RNA schemes outperform other simulations, especially for heavier hourly precipitation, which is consistent with the recall score in Figure 1. Specifically, the median value of the relative error of GF-GF-R (G3-G3-R) simulation is reduced by 52% (25%) than the GF-N-N (G3-N-N) simulation in Typhoon Mujigae at the 120mm threshold (Figure 2b (Figure 2d)). Additionally, the relative error also shows opposite results with different cumulus schemes. The GF-GF-R simulation shows lower relative errors than the G3-G3-R in the Mujigae case (Figure 2a,b), which is 25% lower than the G3-G3-R simulation at 120mm threshold. However, The results are opposite to the Hato case (Figure 2f,h). Furthermore, applying the Smagorinsky scheme for horizontal turbulence tends to weaken the precipitation precision, producing larger errors than the GF-GF-N (G3-G3-N) simulation. It is noteworthy that the distribution of relative errors in the simulation results exhibits different characteristics, with a higher proportion of smaller relative errors observed in the simulation results employing the RNA scheme, suggesting that the utilization of the RNA scheme in the simulations not only produces a smaller median value but also reduces errors at more stations.

The simulated reflectivity of the different experiments and the observed reflectivity for Typhoon Hato at 03:00 UTC, 23 August, is shown in Figure 3 as an example to determine the impact on the typhoon's structure and strength. The G3-N-N simulation generates intense rainfall over the Hainan island in Figure 3e which is spurious compared

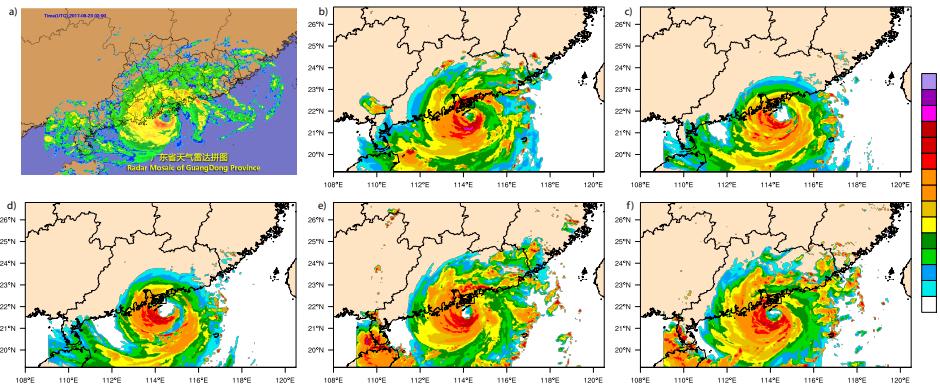


Figure 3. The observed and simulated reflectivity in different simulations for Typhoon Hato:

(a) Observation, (b) G3-G3-R, (c) G3-G3-S, (d) G3-G3-N, (e) G3-N-N, (f) G3-N-R.

with the observation, and simulations in Figure 3b-d show relatively weak reflectivity at around 20 dBZ in Hainan island due to the more compact structure by adopting the cumulus and RNA schemes. In addition, the grid-point convection on the east in the G3-N-N simulation tends to be relatively small in the G3-G3-R simulation due to the stronger convection to deplete convective instability, indicating that simultaneously employing the RNA scheme and cumulus parameterization can maintain the structure and intensity of the typhoon and further avoid causing spurious precipitation. The same features are found in the reflectivity simulation of typhoon Mangkhut (Supporting Figure S3), where applying the Grell-3 parameterization eliminates the false rainfall falling in the north of the Guangdong Province and the Guangxi Province.

The minimum sea level pressure and maximum wind are analyzed to evaluate the impact of the RNA scheme on the typhoon intensity and location (Supporting Figure S4). Applying the RNA scheme enhanced the typhoon intensity for Typhoon Hato and Typhoon Mujigae, showing lower sea level pressure during the prelanding period. For instance, the sea level pressure of typhoon Hato reaches 950hPa in the G3-G3-R and G3-N-R simulations, which is more intense than other simulations. The G3-G3-N and G3-G3-S simulations applying the Grell-3 cumulus scheme didn't show a significant difference in the sea level pressure compared with the G3-N-N (GF-N-N) simulation. On the other hand, the impact of the RNA and cumulus schemes on typhoon tracks is limited. Applying the cumulus and RNA schemes resulted in a higher maximum wind speed than other simulations. However, the G3-N-R simulation, which applies the RNA scheme alone, produces maximum wind speeds comparable to those of the G3-N-N (GF-N-N) simulation. It should be noted that the observations are based on best-track data and the comparison is not conducted at identical locations for both the observation and the simulation, satellite observations may provide further insights for evaluating wind speed over sea areas.

We further investigated the typhoon structure from the tangential and radial flow fields for the Typhoon Hato case (Supporting Figure S5). Applying the cumulus scheme yields a larger high wind speed radius in the G3-G3-R, G3-G3-N, and G3-G3-S simulations. The maximum tangential wind is also larger in the G3-G3-R case, which reaches 56 m/s, notably larger than the other simulations. Combining the RNA scheme and cumulus schemes produces stronger radial wind inflow, but simulations only applying RNA (G3-N-R) remain unchanged radial wind, which is consistent with the maximum wind. Furthermore, the depth of the radial inflow in the G3-G3-R simulation reaches 875 hPa, which is much larger than other simulations (975hPa). In conclusion, applying the cu-

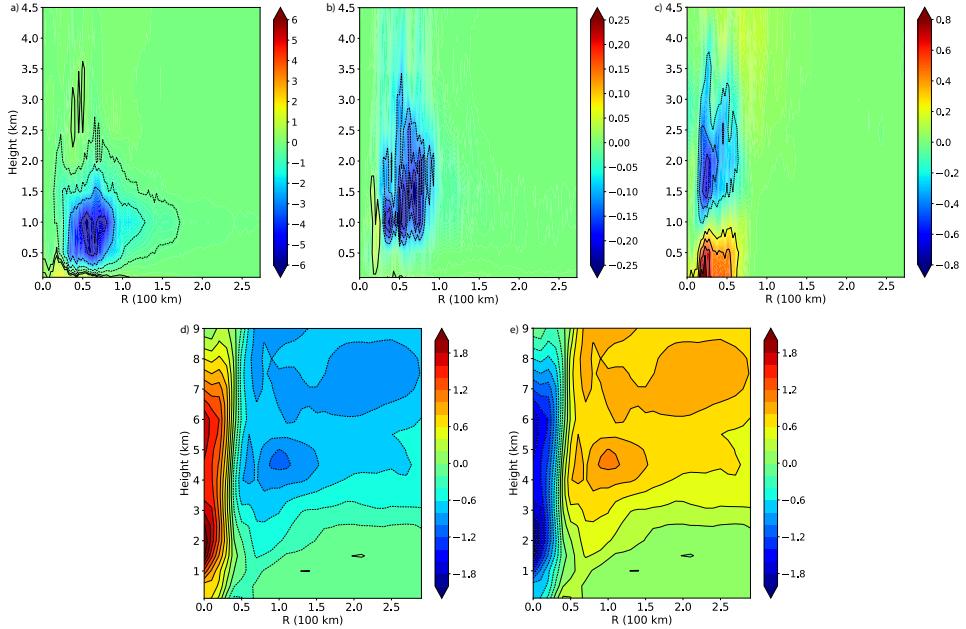


Figure 4. Time-averaged dynamics and difference in simulated potential temperature under different cumulus schemes for the 6 hours before landing. Panels (a-c) display the energy dissipation (positive) and backscatter (negative) within the G3-G3-R simulations for Typhoon Hato: (a) horizontal momentum, (b) vertical velocity, and (c) variance of potential temperature. Panels (d) and (e) show the difference in potential temperature for different cumulus schemes: (d) Grell-Freitas and (e) Grell-3.

271 mulus and RNA schemes simultaneously leads to larger intensity with a larger radius of
272 maximum wind and deeper radial inflow.

273 3.2 Dynamical Analysis

274 The difference in convection intensity is mainly due to the interactions between the
275 parameterized turbulence and the resolved flows. The product of parameterized flux
276 and gradients can be used to measure the downgradient or upgradient generated by the
277 horizontal turbulence parameterization (Shi et al., 2018). The parameterized horizontal mix-
278 ing of potential temperature (θ) is measured by

$$\Pi_\theta = -\tau_{\theta j} \frac{\partial \theta}{\partial \mathbf{x}_j} \quad (6)$$

279 where $\tau_{\theta j}$ is the parameterized horizontal turbulence flux of θ . This term can also
280 produce or destroy the turbulence potential energy in the governing equation of the sub-
281 filter scale θ variance (Shi, Enriquez, et al., 2019).

282 For the downgradient or upgradient mixing of the momentum can be measured by

$$\Pi = -\tau_{ij} \frac{\partial \mathbf{u}_i}{\partial \mathbf{x}_j} = -\tau_{ij} \mathbf{S}_{ij} \quad (7)$$

283 We can divide it into the horizontal and vertical momentum components. For the
284 Π_h , $i = 1, 2$ and $j = 1, 2$, for the vertical component, $i = 3$ and $j = 1, 2$. The positive and
285 negative values represent the downgradient and upgradient mixing of the momentum,
286 respectively. We show the momentum and potential temperature mixing in typhoon Hato

in Figure 4a-c. The mixing of the horizontal momentum shows larger negative values, meaning the RNA scheme generated upgradient mixing around 2 km. In contrast, the value above 2 km is positive with a much smaller magnitude, which implies weak downgradient (dissipation) transport happened. From the tangential wind analysis, we found that only applying the RNA scheme leads to a larger height of the maximum tangential wind than the G3-N-R simulation. By combining the cumulus and the RNA schemes, the maximum tangential wind height reaches 925 hPa. This demonstrates that the RNA scheme enhanced the low-level wind through backscattering. Moreover, the upgradient transport of the horizontal turbulence which enhanced the tangential wind can further enhance the convection in the secondary circulation by the dynamical adjustment. In Figure 4b, we also found the significant backscattering extended to 3 km, suggesting the upgradient transportation of the vertical velocity, which also favors the convection development in the typhoon eyewall. The flux shows the same configuration in other typhoon cases, although the effect is relatively weaker.

Figure 4c shows the calculated heat variance dissipation for the RNA scheme as a function of height and the radius from the typhoon center for Typhoon Hato. We can see that Π_θ displayed positive values meaning downgradient mixing at the low height level. The height of the downgradient mixing of potential temperature extends to 1km in typhoon cases. On the other hand, the heat flux is upgradient at high levels, indicated by the positive Π_θ values. As a result, the high-entropy air is transported from the eyewall to the outside, which further enhances the buoyancy of the updraft in the eyewall; in contrast, the backscattering at the upper level seems to be advantageous for deepening the convection, as it may potentially reduce the entrainment of environmental air, which will be investigated with further numerical experiments. In addition, different typhoon intensities may induce different magnitudes of heat and momentum fluxes, e.g., the heat flux of Typhoon Hato is stronger than Typhoon Mujigae. Nevertheless, both fluxes configurations contribute to the increased precipitation intensity, consistent with the enhanced typhoon precipitation forecast in the G3-G3-R simulations. The tangential wind is stronger near the eyewall in the G3-G3-R simulations than in the others. However, the G3-N-N simulation can produce stronger tangential wind in some situations, as in the Typhoon Hato Case.

We also examined the impact of different cumulus schemes by analyzing the difference in the potential temperature between the GF-GF-N and G3-G3-N simulations and their averaged field. We show the difference as a function of the radius from the center 6 hours before landing in Figure 4d,e. The Grell-3 scheme shows warmer air at the high level because the high entropy air from lower levels and the eye is transported to the environment and leads to more intense precipitation, resulting in higher recall scores and lower relative error in the heavy rainfall scale compared to the Grell-Freitas simulation. In addition, we found the moisture convergence in the G3-G3-N simulation is stronger than the GF-GF-N simulation for Typhoon Hato, especially before the landing stage, which means the Grell-3 scheme leads to intensified convection which is close to the observation. But for Typhoon Mujigae, allying the Grell-Freitas results in weaker moisture convergence which is consistent with the precipitation forecasts. However, as we mentioned before, the performance of the schemes can vary in different cases because of the various environments and typhoon structures, and the adaptation of the cumulus for the grey zone scheme still needs further investigation.

4 Conclusion

Tropical cyclones are significant weather systems, leading to extreme rainfall in coastal areas. Although convection-permitting-resolution numerical predictions of typhoons have become operational in many regions, forecasting precipitation remains challenging due to the controversial representation of convection and turbulence at grey zone resolutions. Traditional boundary layer turbulence schemes do not allow for horizontal turbulence,

339 which might hinder accurate typhoon precipitation predictions. Nevertheless, recent re-
 340 search has emphasized the importance of both vertical and horizontal subgrid-scale ef-
 341 fects in the simulation of typhoon development. This study evaluated the necessity and
 342 efficacy of the cumulus and RNA turbulence schemes on typhoon precipitation in kilometer-
 343 scale resolution simulations in border typhoon cases.

344 We found that applying the RNA turbulence scheme and integrating the cumulus
 345 scheme and turbulence scheme led to increased domain-averaged precipitation, higher
 346 recall scores, and reduced relative error compared to other simulations. Applying the cu-
 347 mulus and RNA turbulence schemes can enhance the typhoon intensity and generate more
 348 compact structures with lower minimum sea level pressure, and higher maximum wind
 349 speed. Combining the cumulus and RNA schemes also leads to a larger radius of max-
 350 imum wind and deeper radial inflow which benefit the intense convection. In addition,
 351 the two cumulus schemes exhibit varying impacts when integrated with the RNA scheme
 352 due to the specific characteristics of the schemes and typhoon cases. However, implement-
 353 ing the convection parameterization and RNA turbulence schemes does not necessarily
 354 enhance precipitation forecasting for weak precipitation events. The RNA scheme can
 355 generate horizontal downgradient mixing of potential temperature, increasing buoyancy
 356 flow towards the eyewall. Simultaneously, backscatter is observed in the upper level, re-
 357 ducing the convection core's depletion. The RNA scheme also promotes the upgradient
 358 transport of momentum in the lower troposphere, dynamically reinforcing typhoon cir-
 359 culation. We noticed that the magnitude of momentum and flux varies due to differing
 360 typhoon intensities, but the overall trend remains consistent.

361 Our study highlights the importance of considering cumulus and horizontal subgrid-
 362 scale turbulence impacts in typhoon precipitation forecasts at convection-permitting res-
 363 olutions, particularly for extreme precipitation events. They are useful to improve heavy
 364 rainfall warnings for typhoon cases. However, the specific impact of the RNA scheme and
 365 the advantage of the scale-aware convection scheme varies in different typhoon cases, prob-
 366 ably related to the distinct boundary layer environments, background fields, the sensi-
 367 tivity of combining the microphysical and cumulus scheme, the specific entrainment and
 368 typhoon structures of different cases. The results are also consistent with previous stud-
 369 ies, Liu et al. (2020) found that only the Grell-3 is superior for accumulated rainfall sim-
 370 ulation in the central Tianshan Mountains; Jeworrek et al. (2019) showed that GF per-
 371 formed better in the two case studies in the US Southern Great Plains. Ensemble nu-
 372 matical simulations will be conducted to investigate the cumulus and RNA turbulence
 373 parameterization schemes across different grid-resolution scales for typhoons exhibiting
 374 varying structures and intensities.

375 5 Open Research

376 The Weather Research and Forecast model is publicly available at <https://github.com/shixm-cloud/WRF-RNA>. We archived the namelist for our simulations at (WANG,
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