



DYNAMICS OF DIURNAL WARM LAYERS

Ken Hughes, Jim Moum, Emily Schroyer

Diurnal warm layers on a tabletop

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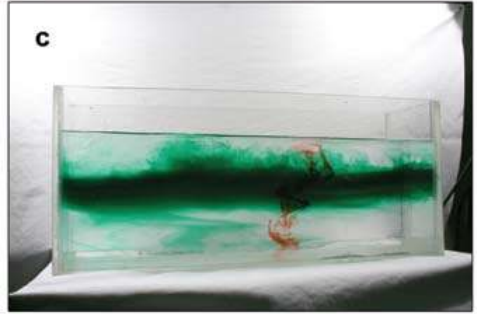
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Franks and Franks (2009)

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AUDIENCE

Designed for graduate oceanography courses, this simulation is suitable for students as young as elementary school age, provided the level of discussion is appropriately scaled.

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Designed for graduate oceanography courses, this simulation is suitable for a Zoom meeting of physical oceanographers, provided the level of discussion is appropriately scaled.

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Thin, stratified, and accelerated

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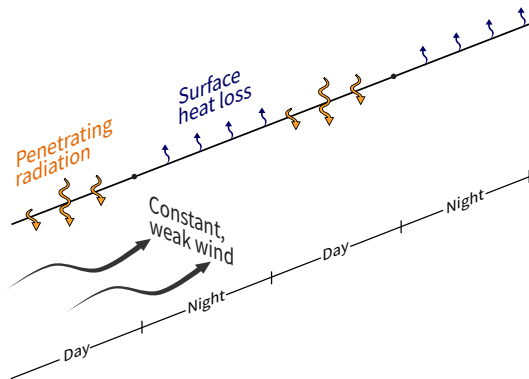
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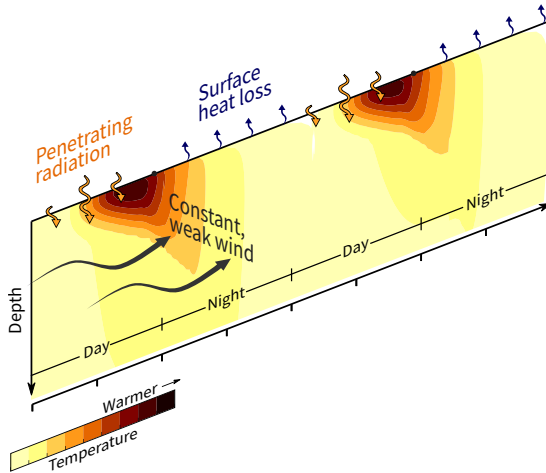
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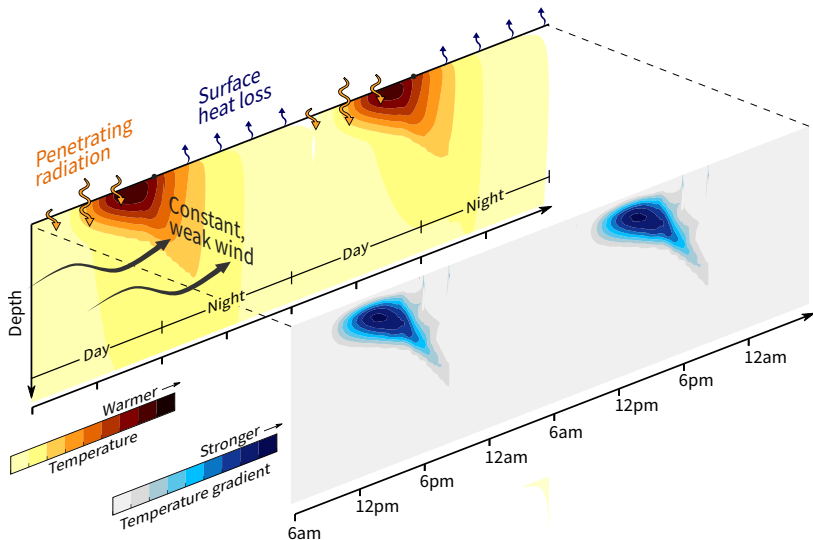
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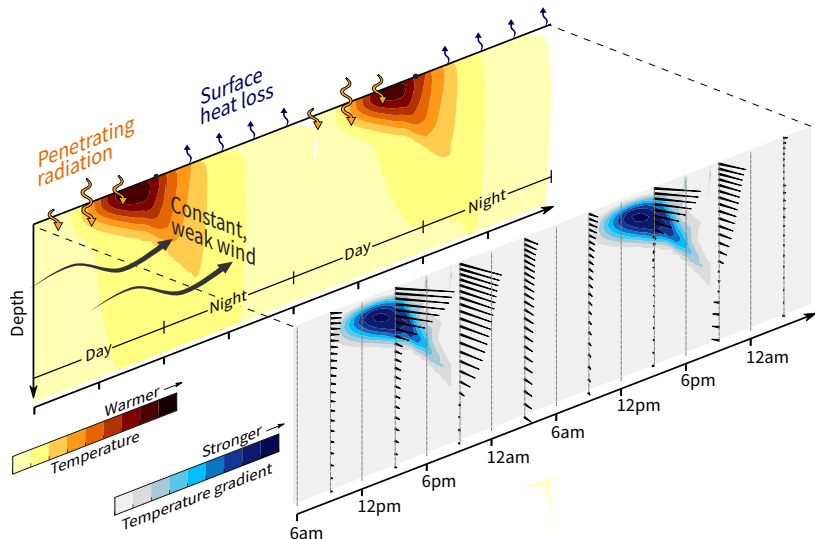
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Evolution of the Velocity Structure in the Diurnal Warm Layer

KENNETH G. HUGHES, JAMES N. MOUM, AND EMILY L. SHROYER

College of Earth, Ocean, and Atmospheric Sciences, Oregon State University, Corvallis, Oregon

(Manuscript received 23 August 2019, in final form 18 December 2019)

Heat transport through diurnal warm layers

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Something about shear instabilities?

Warm layers are ubiquitous

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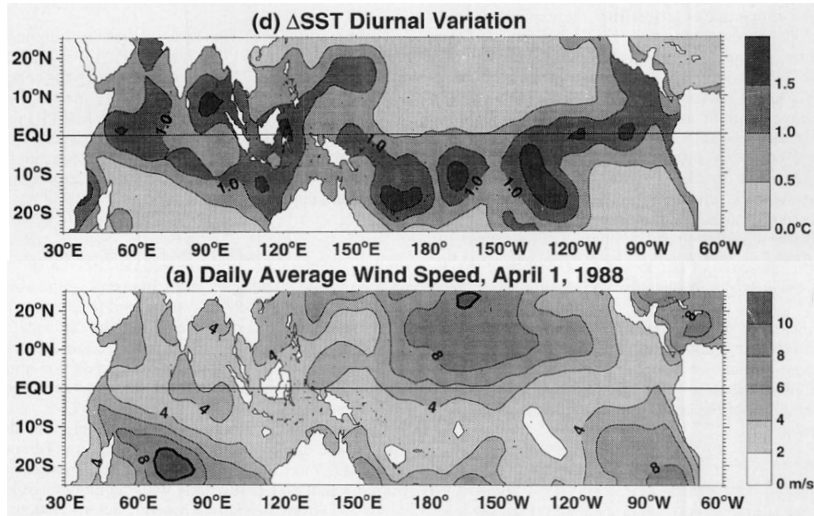
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Webster et al. (1996)

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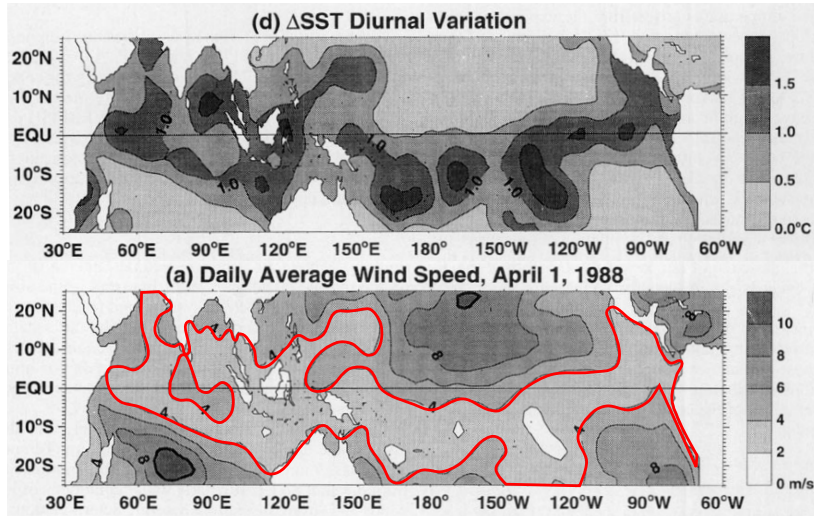
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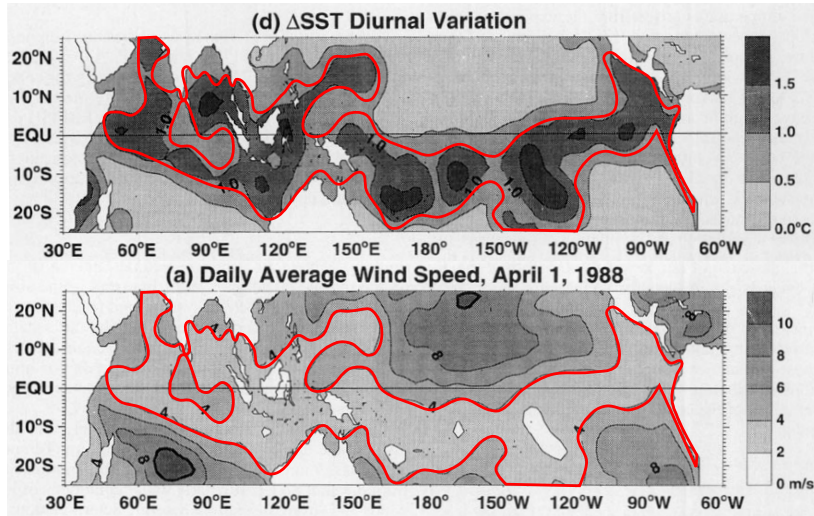
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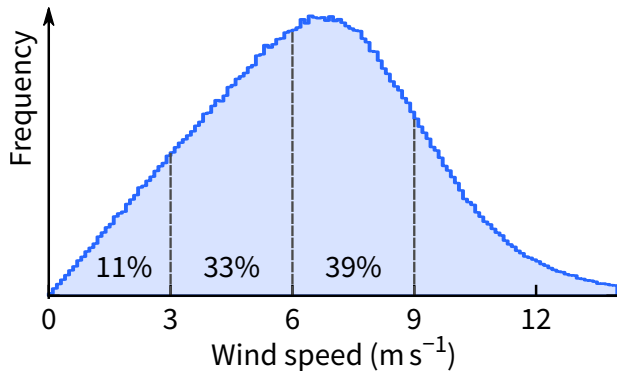
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Based on NCEP-DOE reanalysis

SST governs surface heat flux

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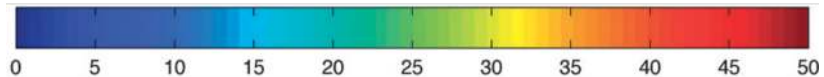
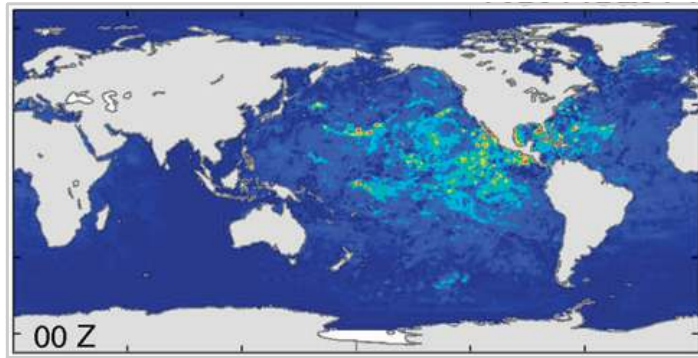
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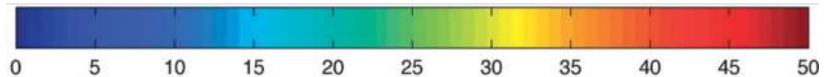
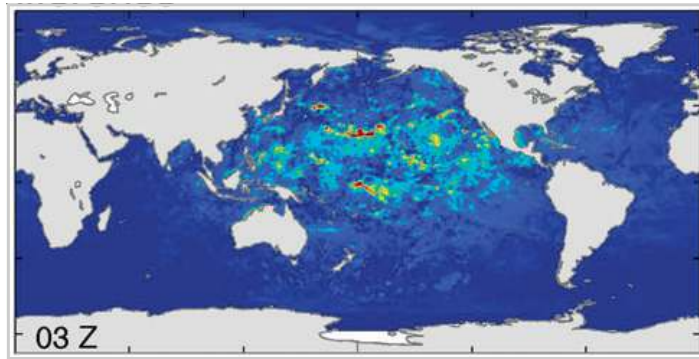
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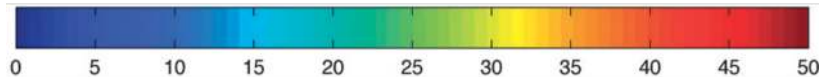
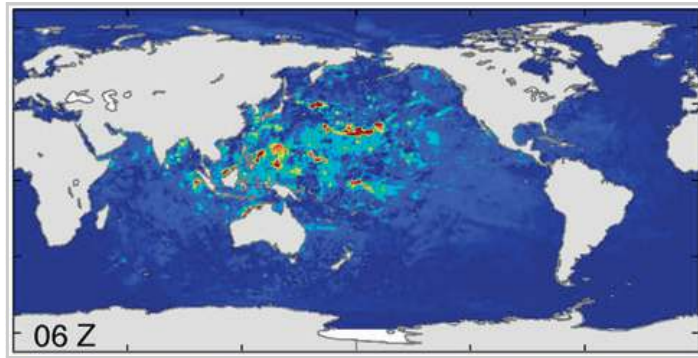
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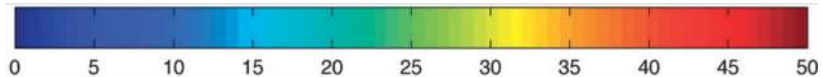
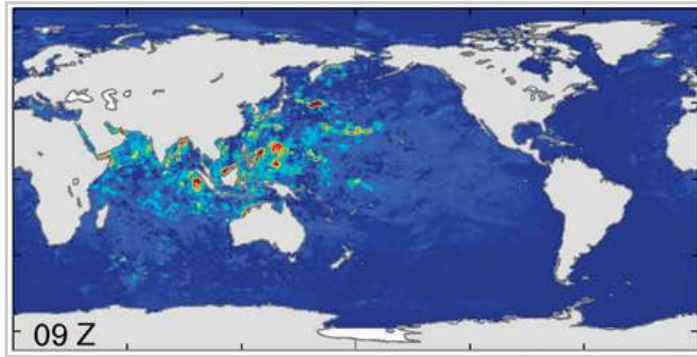
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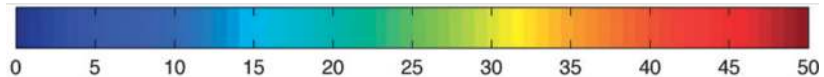
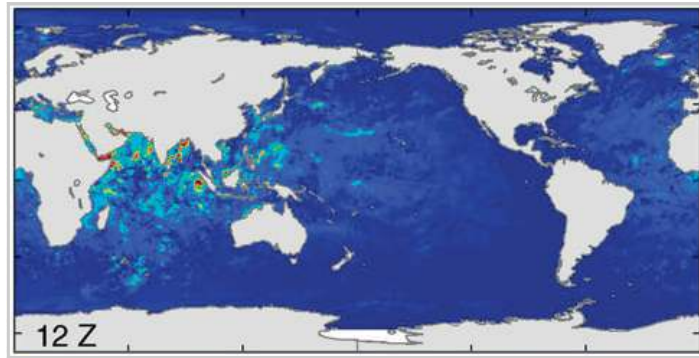
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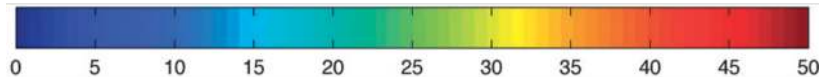
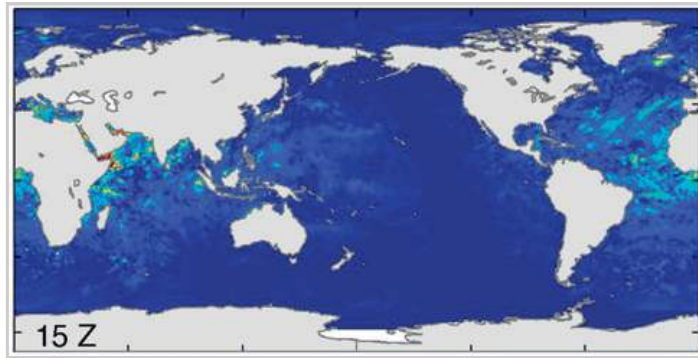
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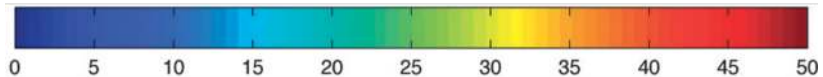
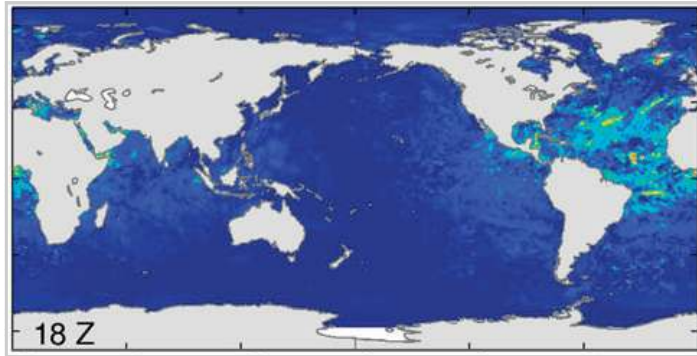
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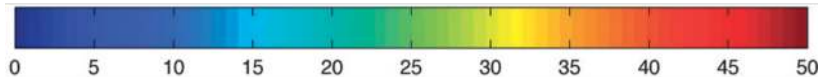
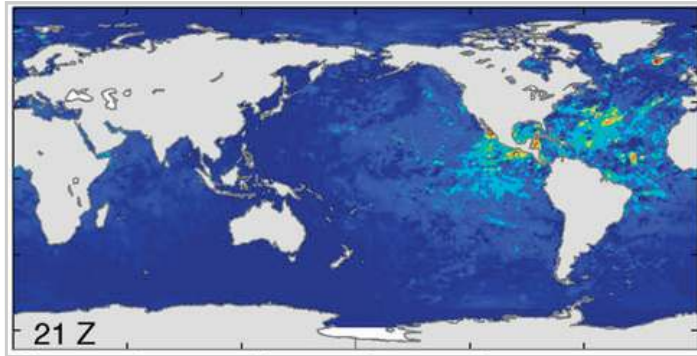
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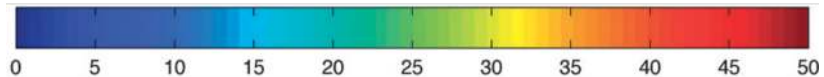
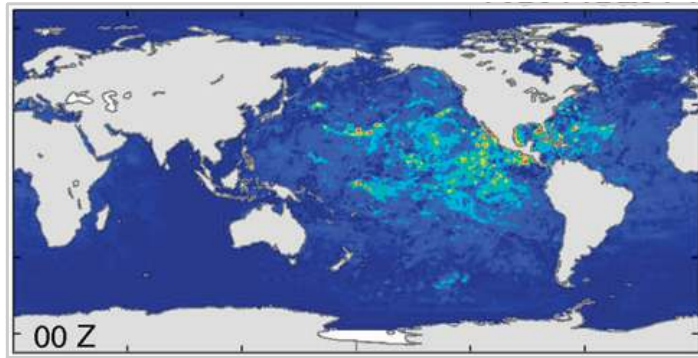
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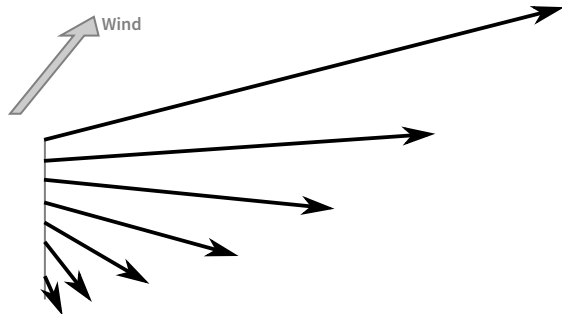
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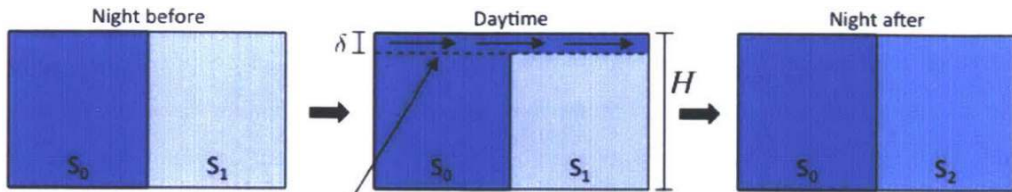
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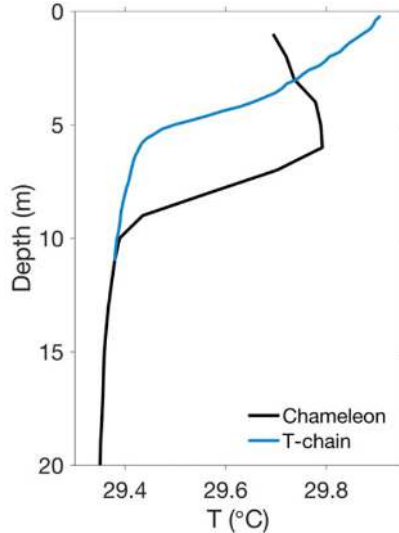
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Bogdanoff (2017)

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Surface measurements are difficult

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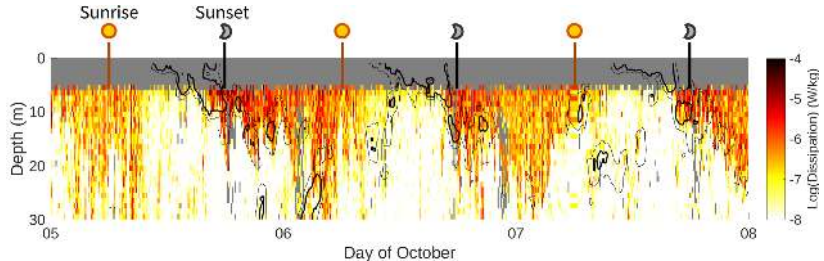
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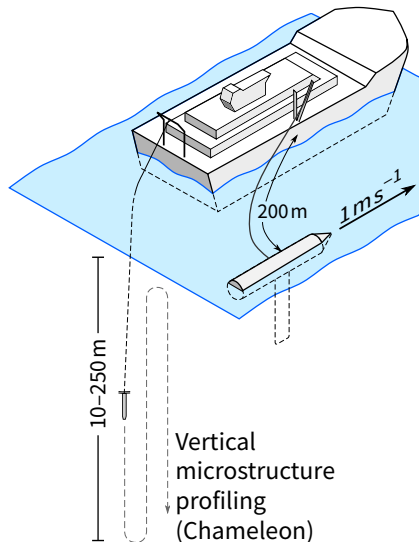
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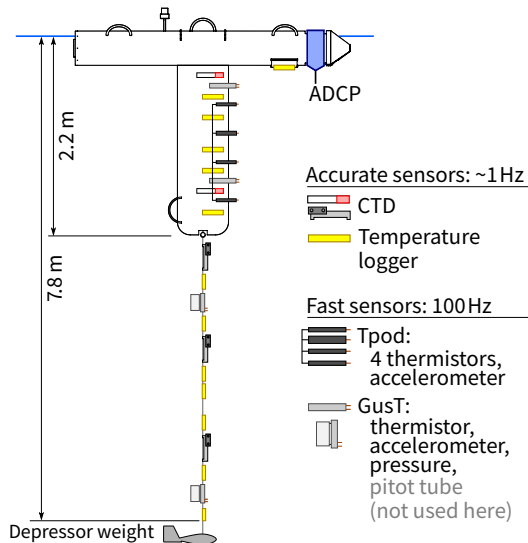
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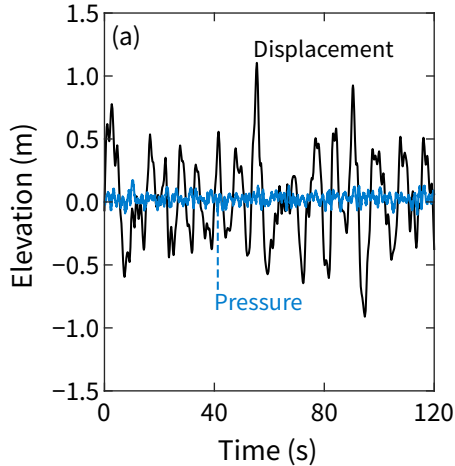
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SurfOtter follows the surface well



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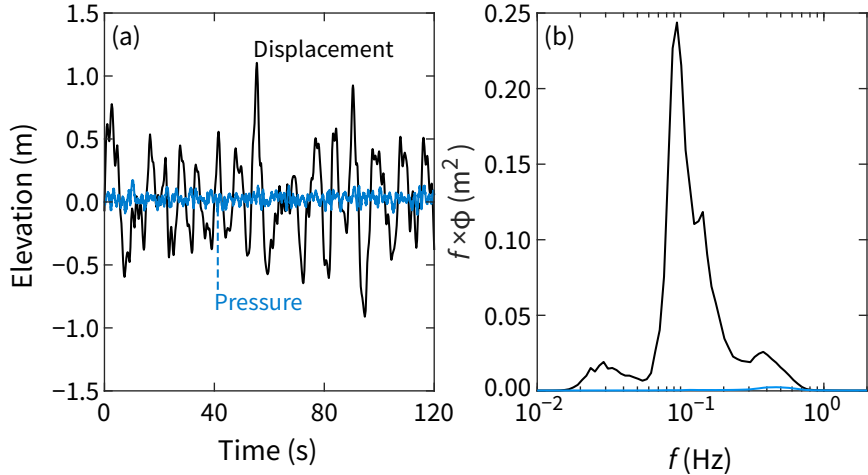
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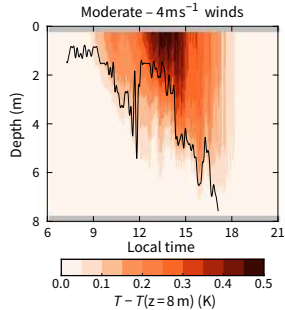
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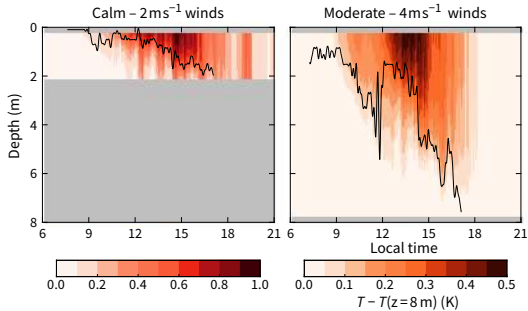
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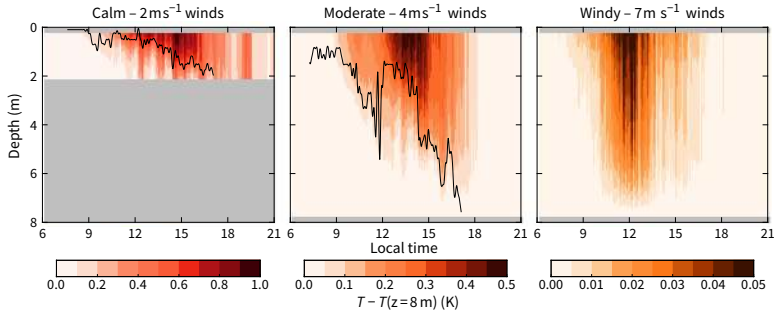
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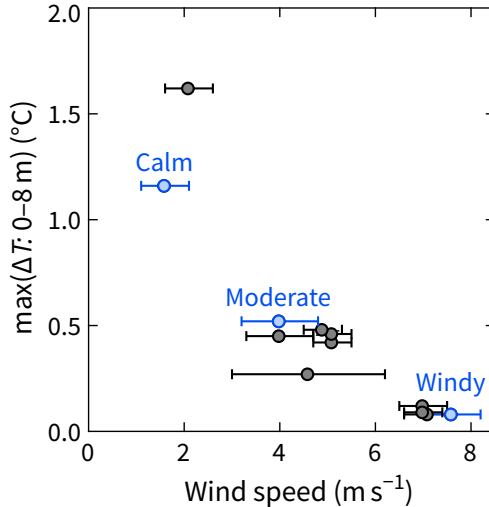
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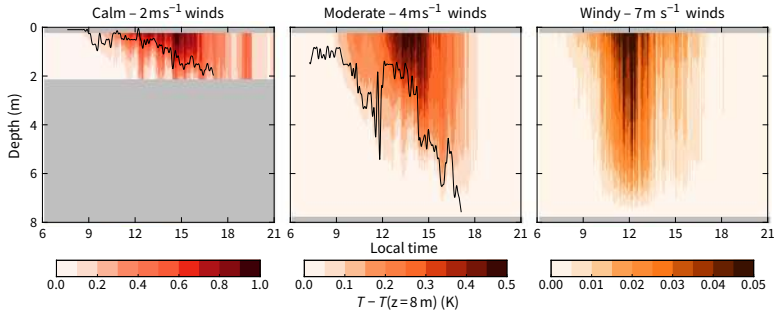
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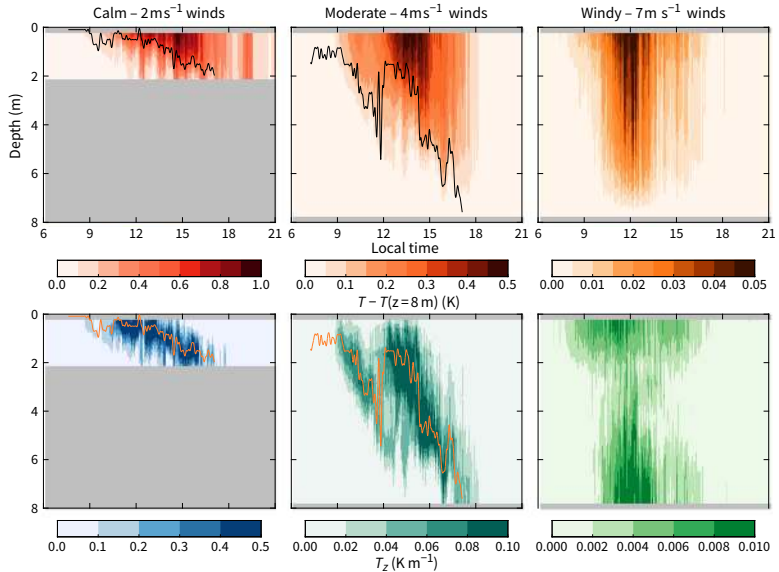
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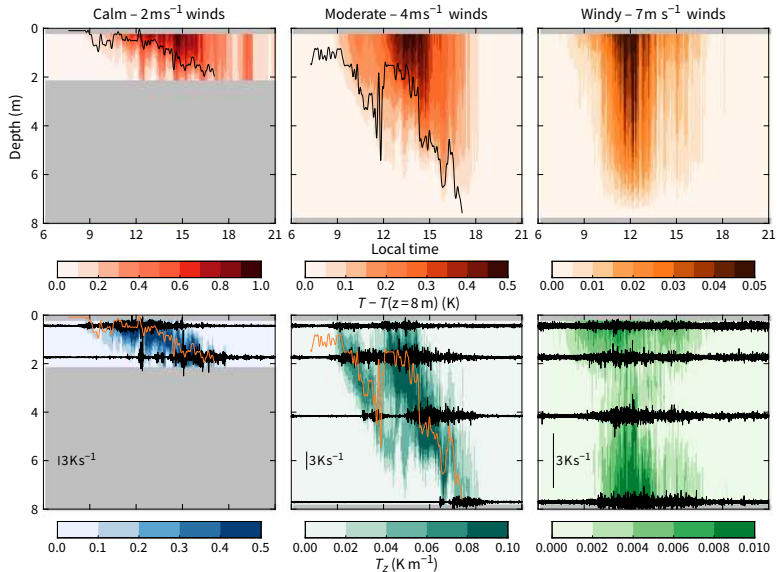
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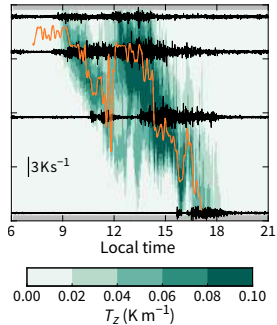
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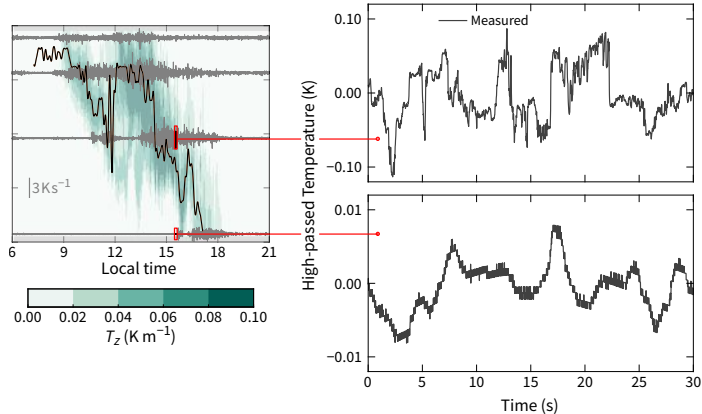
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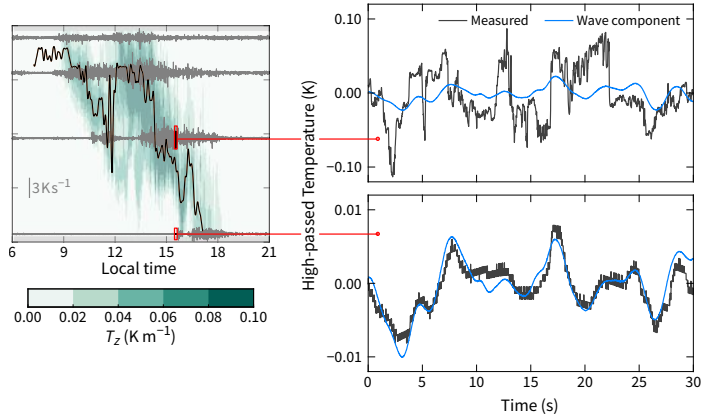
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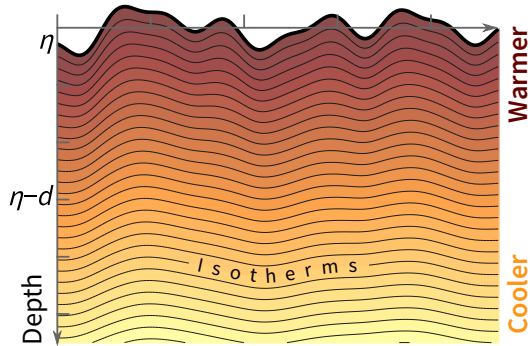
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SurfOtter temperatures always have a wave component



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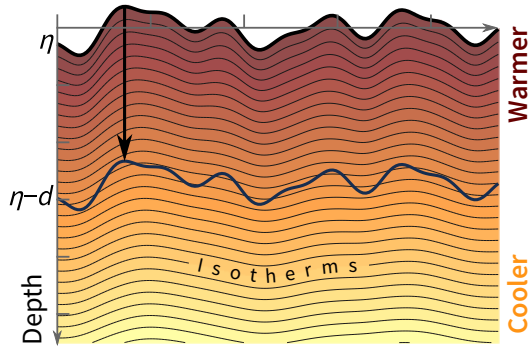
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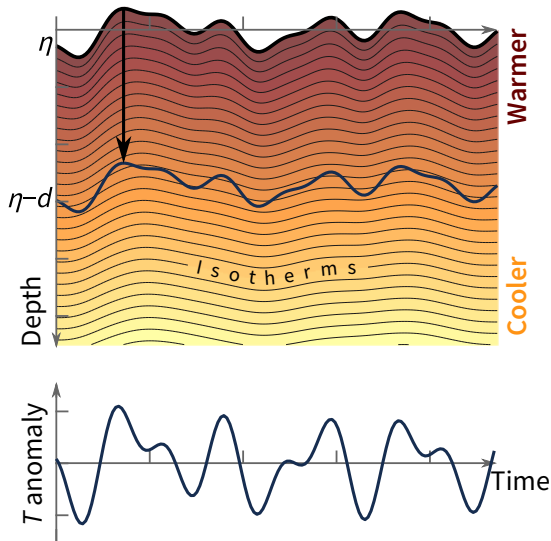
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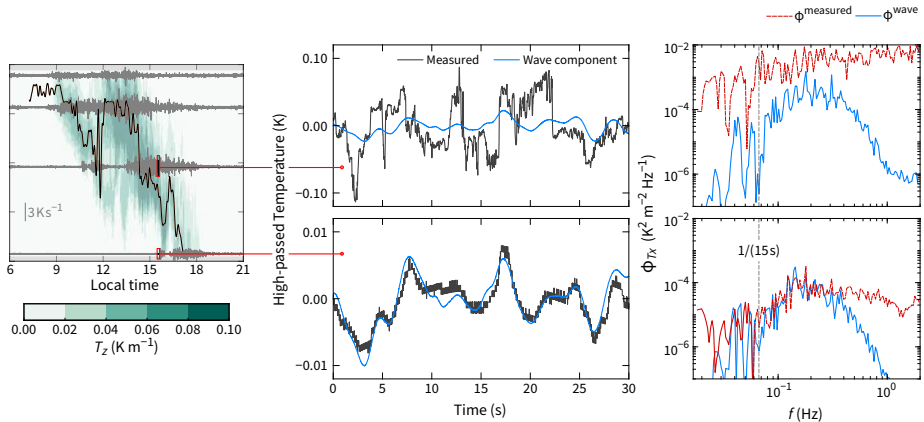
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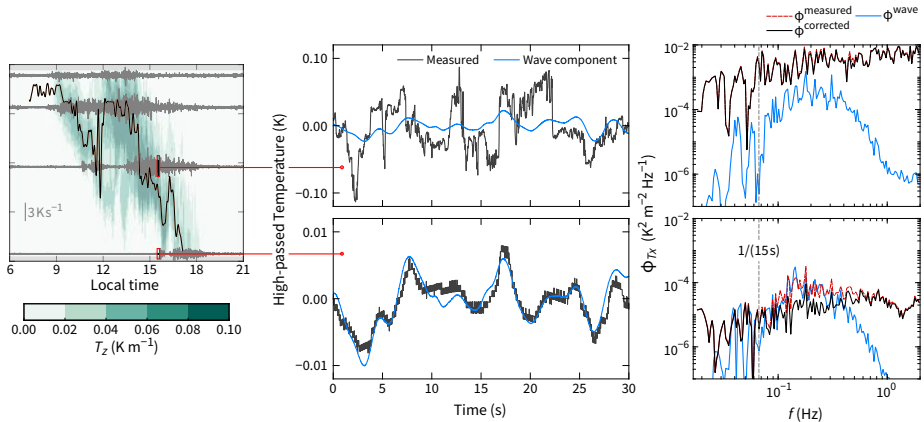
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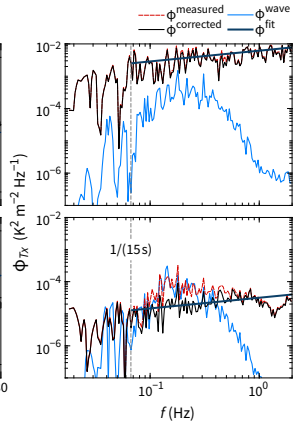
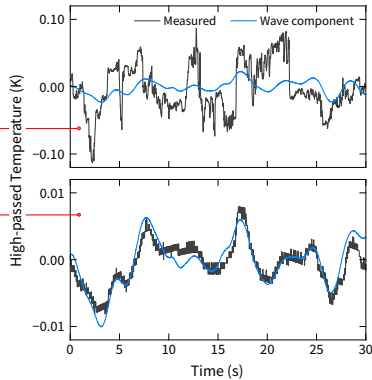
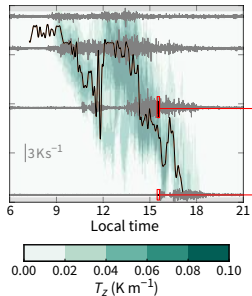
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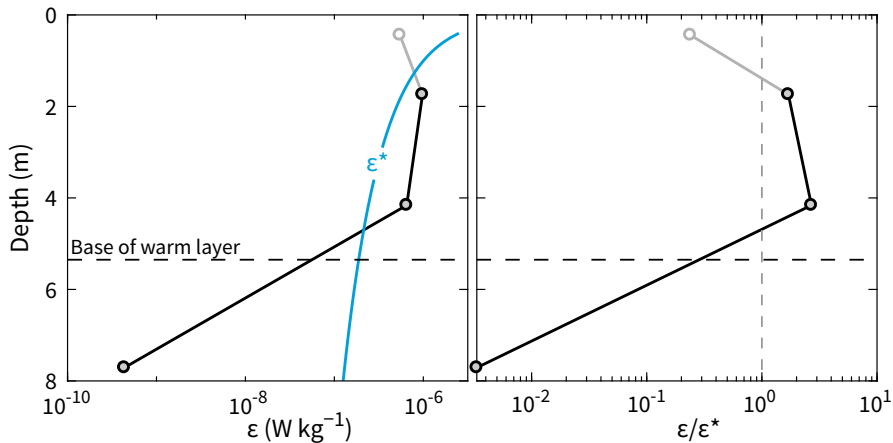
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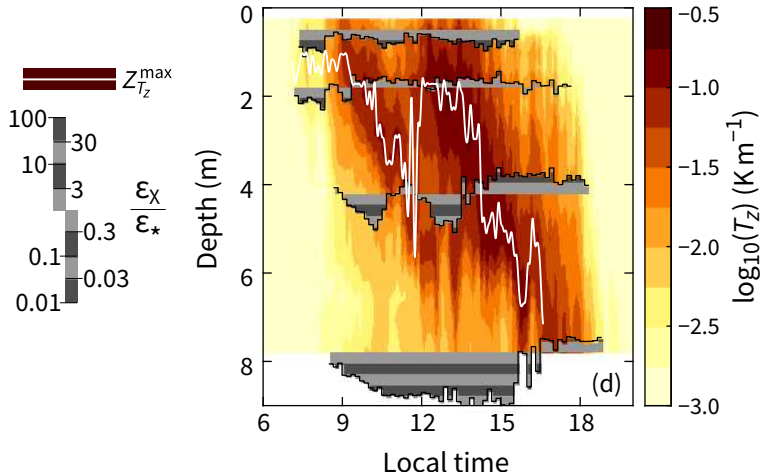
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Turbulent buoyancy fluxes completely shut off?

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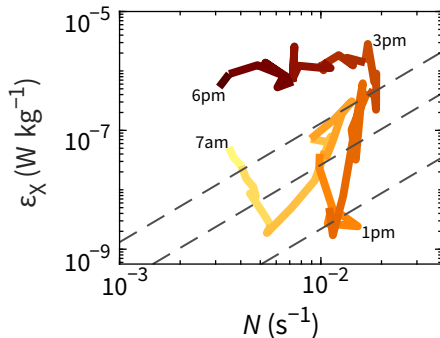
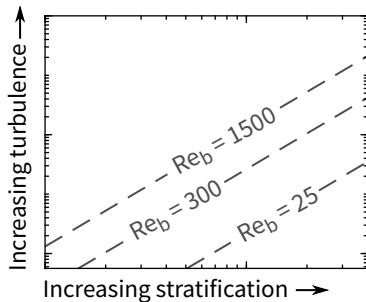
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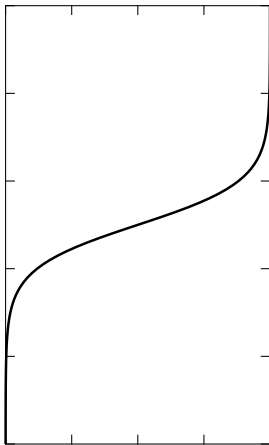
**Turbulent
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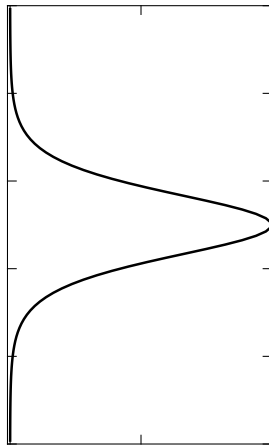
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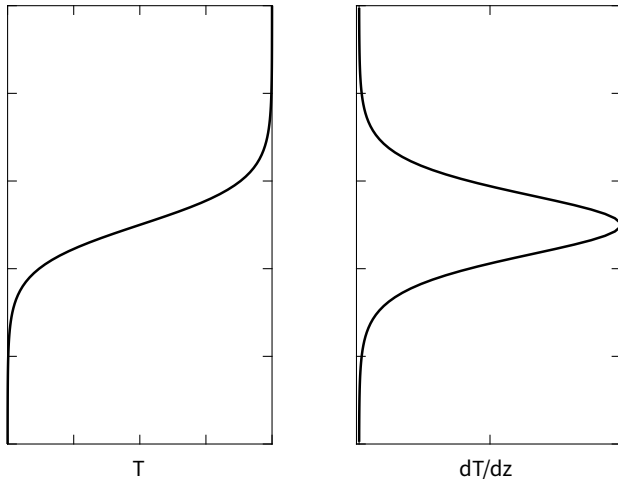
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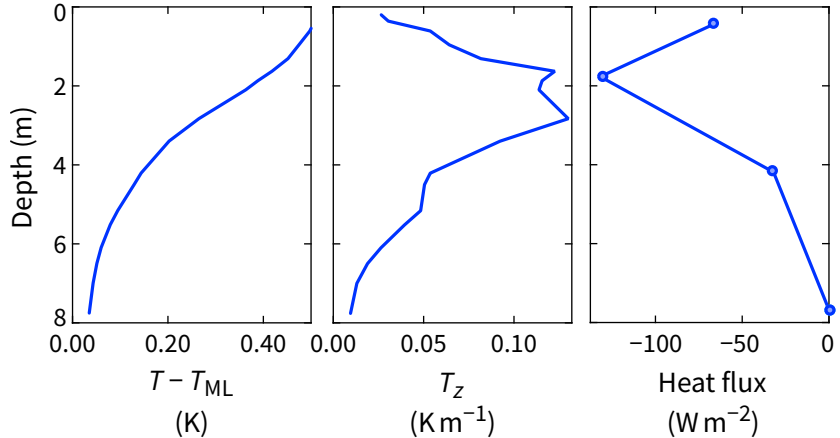
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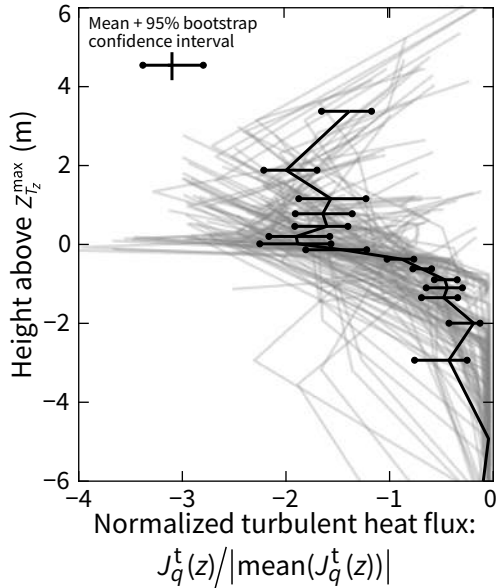
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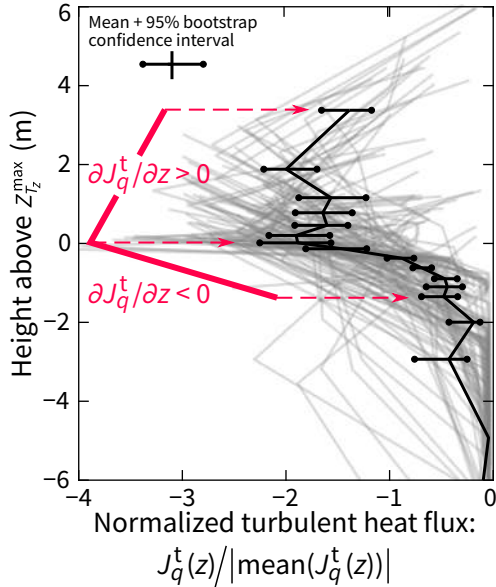
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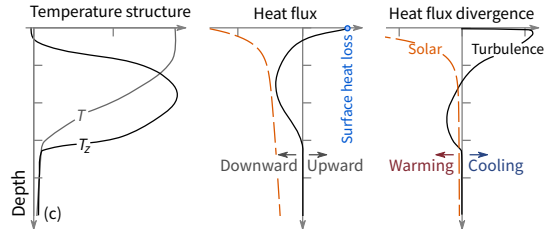
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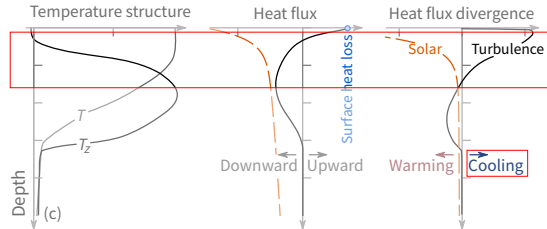
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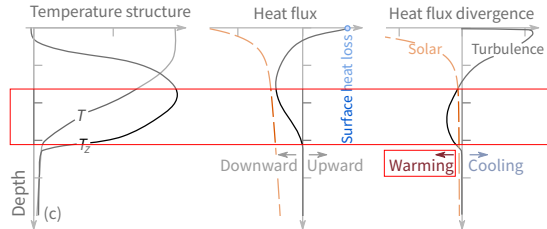
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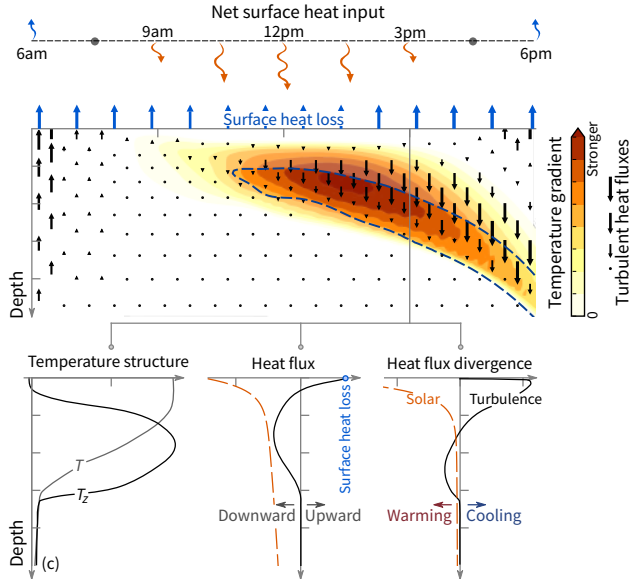
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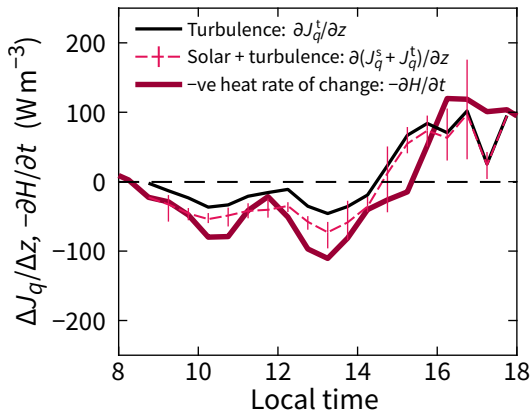
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$$\frac{-\partial H}{\partial t} = \frac{\partial J_q^t}{\partial z} + \frac{\partial J_q^s}{\partial z}$$

Independent estimates of heat transport agree



Warm layers exhibit marginal instability

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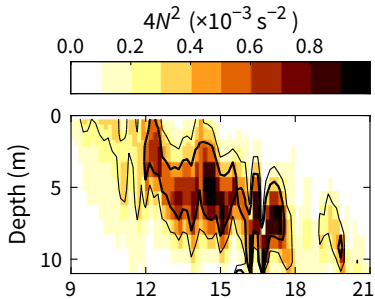
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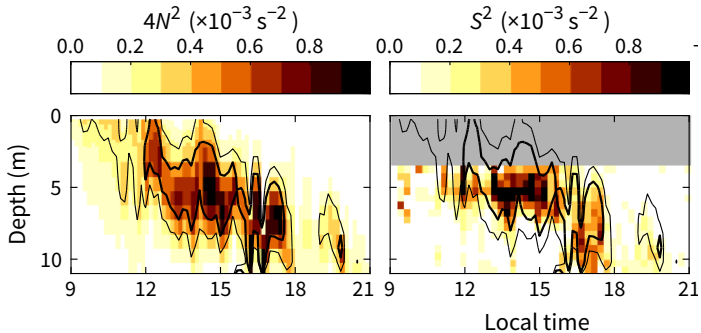
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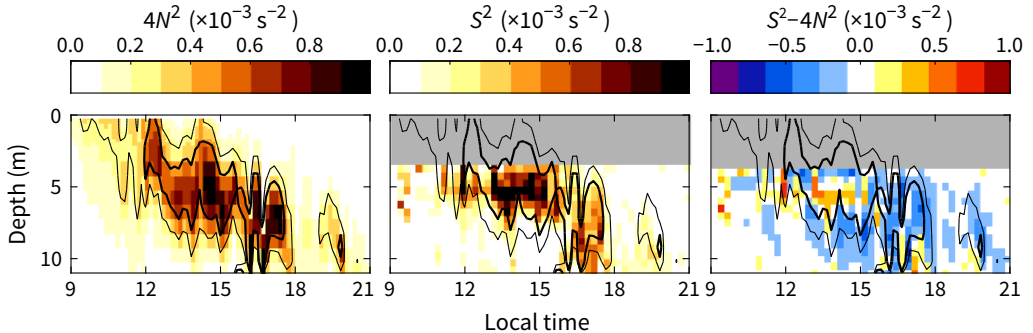
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Marginal instability only if wind $> 2 \text{ m s}^{-1}$

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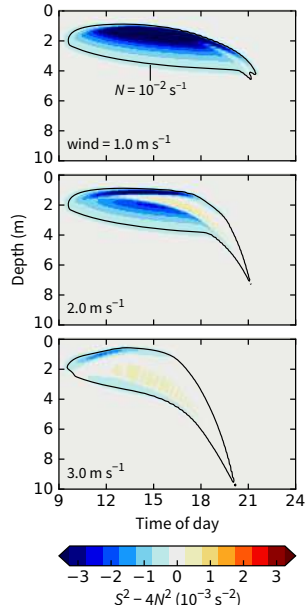
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Kinetic energy input:

$$\frac{u^2 + v^2}{2} = \frac{1}{2} \left(\frac{\tau}{h\rho_w f} \right)^2 (2 - 2\cos(ft))$$

Potential energy input:

$$J_bt = \frac{g\alpha}{\rho_w c_p} J_q t,$$

$$U_{\text{cr}} \approx 2 \text{ m s}^{-1}$$

Other evidence for 2 m s^{-1}

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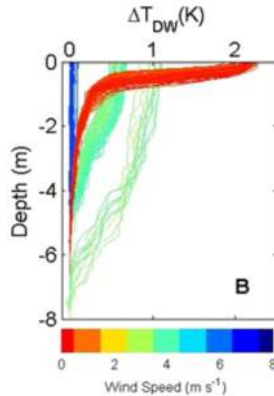
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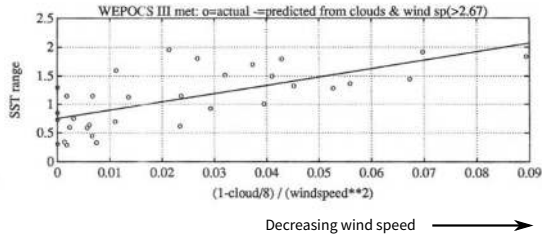
Gentemann et al. (2009)

$$\text{dSST} = f + a(\text{PS}) + b(P) + c \ln(U) \\ + d(\text{PS}) \ln(U) + e(U)$$

TABLE 5. Coefficients for determination of diurnal sea surface temperature amplitude (dSST) from (6).

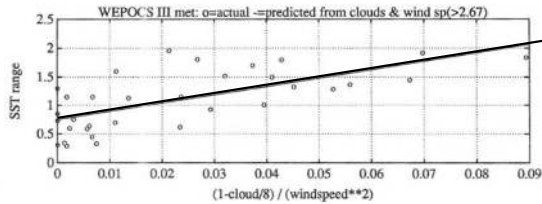
Coefficient	$U > 2 \text{ m s}^{-1}$ value	$U \leq 2 \text{ m s}^{-1}$ value
f	0.262	0.328
a	0.002 65	0.002
b	0.028	0.041
c	-0.838	0.212
d	-0.001 05	-0.000 185
e	0.158	-0.329

Other evidence for 2 m s^{-1}



Lukas (1991)

Other evidence for 2 m s^{-1}



Decreasing wind speed \longrightarrow

Lukas (1991)

Velocities almost to the surface

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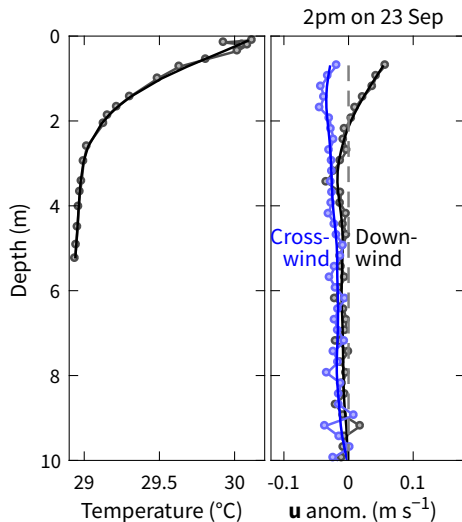
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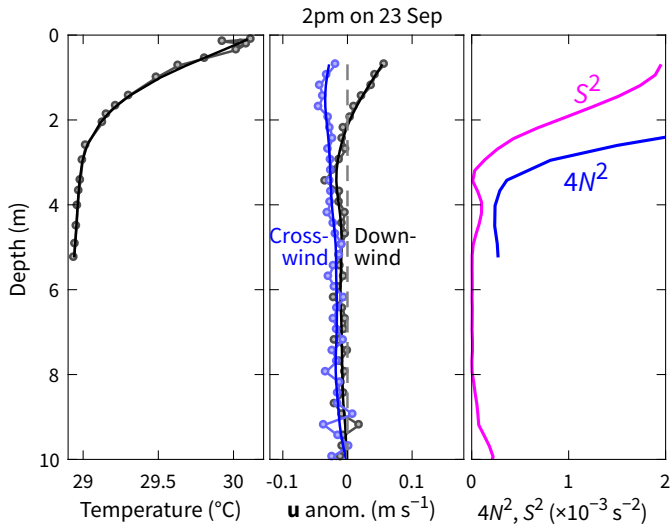
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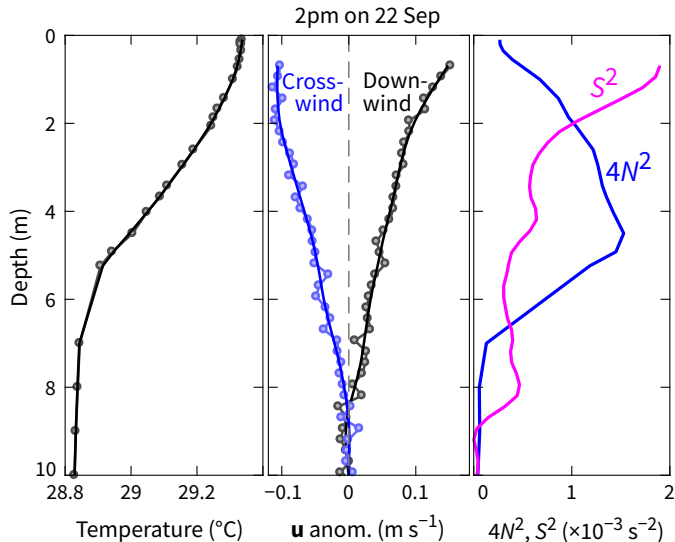
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Dynamics of the diurnal thermocline

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