

# 1 From Snowball Earth to the Cambrian Explosion: an Ediacaran 2 Subcommission field trip to Brazil

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23

## 24 Abstract

25 We report the findings from our International Subcommission on Ediacaran Stratigraphy  
26 (ISCS) trip to the Neoproterozoic to Cambrian Bambuí Group in Brazil. Important geochemical  
27 markers, glacial diamictites, and possible calcifying fossils have previously been reported from  
28 the Bambuí Group, which represent important criteria to be considered in subdividing the  
29 Ediacaran Period into meaningful Series and Stages.

30

31     **Introduction and background**

32       In late November 2023, the Ediacaran Subcommission, together with members of the  
33       Cryogenian and Cambrian subcommissions, held a field meeting to explore the Neoproterozoic to  
34       Cambrian Bambuí Group in Minas Gerais, Brazil. In our continued endeavour to subdivide the  
35       nearly 100 million years (635–539 Ma) of Ediacaran time into meaningful windows, Xiao et al.  
36       (2016) outlined the potential geological, geochemical, and paleontological observations that could  
37       be used as temporal markers. Chief amongst these are the identification of key geochemical  
38       events—especially the profound negative carbon isotope anomaly known as the Shuram Excursion  
39       as well as the first and last occurrence of Ediacara biotas and the stratigraphic position of a glacial  
40       diamictites like the one that corresponds to the ca. 579 Ma ‘Gaskiers’ ice age (Pu et al., 2016).  
41       This five-day field excursion included important stops to investigate a large-scale glacial unit of  
42       debatable Neoproterozoic age and to search for important Ediacaran biostratigraphic markers,  
43       especially the calcified funnel-in-funnel tubes of *Cloudina*—the first known animal to create a  
44       carbonate shell. As such, the Bambuí Group represents a key section that potentially provides  
45       insight into the selection of GSSPs (Global Boundary Stratotype Section and Point) to subdivide  
46       the Ediacaran System. In this report, we highlight the important localities visited by voting and  
47       corresponding members of both subcommissions and discuss some outstanding questions  
48       concerning the Bambuí Group of Brazil.

49

50       The Bambuí Group is a Neoproterozoic to Cambrian mixed carbonate-siliciclastic unit that  
51       outcrops over 350,000 km<sup>2</sup> of the Archean/Proterozoic São Francisco Craton in east central Brazil  
52       (Caxito et al., 2012; Reis et al., 2017; and references therein). The Bambuí Group comprises six  
53       major units (in ascending order): (1) The Jequitaí Formation, a diamictite that represents a  
54       Neoproterozoic glacial deposit of debatable age; (2) The Sete Lagoas Formation, which is divided  
55       into the basal Pedro Leopoldo Member (LSL) that contains originally aragonite crystal fans and a  
56       strong negative  $\delta^{13}\text{C}$  anomaly representing a post-glacial cap carbonate, and the Lagoa Santa  
57       Member (MSL and USL), comprising organic-rich limestone and dolomitic limestone with laminar  
58       and columnar stromatolites and thrombolites, in addition to the report of *Cloudina*, *Corumbella*,  
59       and putative trace fossils (Warren et al., 2014); (3) The overlying siltstone and mudstone of the  
60       Serra de Santa Helena Formation likely represents a basin-wide transgression. The occurrence of  
61       heterolithic facies associated with occasional salt pseudomorphs and mud cracks in this interval

62 suggests deposition in inter- to subtidal settings under episodic evaporitic conditions (Uhlein et al.,  
63 2019). Abundant microbial surface textures, pustular structures, and possible simple traces (Okubo  
64 et al., 2023) indicate extensive distribution of mat grounds in this shallow marine environment; (4)  
65 The overlying Lagoa do Jacaré Formation is composed of dark oolitic to muddy limestone and  
66 microbialites; (5) The Serra da Saudade Formation consisting of green siltstone, shale, and  
67 sandstone, and finally (6) The Três Marias Formation, a coarse siliciclastic molasse deposit,  
68 composed of sandstone, conglomerate, and shale, with potential trace fossils, as well as detrital  
69 zircons as young as ca. 527 Ma, indicating a Cambrian age (Tavares et al., 2020). Bambuí strata  
70 are divided by at least five distinct sequence boundaries, including a major regional unconformity  
71 between the Serra da Saudade and Três Marias formations (DaSilva et al., 2022).

72

### 73 **Out on the rocks**

74 Our first day began with a visit to the Brazilian Geological Survey's core facility in Caeté.  
75 There we investigated two cores (PSB13 and PSB14, Figs. 1, 2) drilled through the São Francisco  
76 Craton (henceforth SFC) near Januária. These rock libraries gave delegates unimpeded access to  
77 the Bambuí stratigraphy that formed the basis of our trip and allowed the trip leaders to guide us  
78 through the complex geochemical signals that have been recovered from the cores (Fig. 3). Both  
79 cores contain diamictite of glacial origin that are overlain by post-glacial cap carbonate with  
80 characteristically negative  $\delta^{13}\text{C}$  values. These are disconformably followed by limestone and  
81 dolomitic limestone in an 80+ meter interval of near 0‰ values in which the putative  
82 claudinomorphs were reported (Warren et al., 2014), and then an interval spanning three  
83 formations of >350 meter with profoundly positive values ranging from +5 up to +14‰.

84

85 The Ediacaran age assignment for the Bambuí Group relies on the presence of  
86 claudinomorphs in the Lagoa Santa Member, which was a primary target of the Subcommission  
87 field workshop. If the biomineralized fossils are verified at this level, they would represent the  
88 oldest known claudinomorphs given their stratigraphic position below the Middle Bambuí  
89 Excursion (MIBE), which based on the stratigraphic architecture of the Windermere Supergroup  
90 would underly the Shuram—constrained by Re-Os ages between 564 and 579 Ma (Rooney et al.,  
91 2020)—that is preserved in the Gametrail Formation (Macdonald et al., 2013). At present, the  
92 oldest known claudinomorph fossils are found near the end of the Shuram in the lower Nama

93 Group of Namibia (Kaufman et al., 2019). Unfortunately, the Shuram is missing in the Bambuí  
94 stratigraphy, which might be due to basin restriction (Uhlein et al., 2019; contra Moynihan et al.,  
95 2019 and DaSilva et al., 2022), in addition to the fossil impressions of the soft-bodied Ediacara  
96 biotas.

97

98 Our second day focused on examining outcrops and quarries exposing Sete Lagoas Formation  
99 carbonates as we travelled over 400 km north from Belo Horizonte to Montes Claros. Along the  
100 way we stopped at several localities, including those where tubular columnar stromatolites (Fig. 4  
101 left panel) and calcite (ex-aragonite) crystal fans were documented (Figs. 4 right panel, 5). On day  
102 three we continued our travels north towards Januária and observed spectacular outcrops of the  
103 Jequitaí Formation showcasing glacially striated pavements characterised by U-shaped grooves  
104 with internal striae (Fig. 6). These are interpreted as slumped plow ridges caused by saltation of  
105 grounded glacial ice. We visited additional sites where the Jequitaí diamictite included large  
106 dropstones and striated clasts. On day four we explored the Januária area in the northern Minas  
107 Gerais State. In addition to spectacular bladed barite ( $\text{BaSO}_4$ ) crystal fans (Fig. 7) from an outcrop  
108 at Riacho da Cruz, delegates searched the Barreiro Quarry from where thrombolites with chert  
109 nodules and the Ediacaran biomineralized macrofossil *Cloudina* have been reported (Warren et al.,  
110 2014). On day five, members explored several outcrops of the Serra de Santa Helena, Lagoa do  
111 Jacaré and Serra da Saudade formations in the Januária-Lontra section. Examples of microbially  
112 induced sedimentary structures (MISS) and putative Cambrian trace fossils were found in very  
113 fine-grained sandstone of the Serra de Santa Helena Formation. In the afternoon we moved to the  
114 Sapé section, where peritidal to subtidal carbonates of the Sete Lagoas Formation crop out. At this  
115 location, samples of possible fragments of skeletal organisms were collected. Our final field day  
116 was spent visiting the spectacular Janelões cave system (Fig. 8) of the Peruaçu National Park,  
117 which formed by karstification of the Sete Lagoas Formation carbonates.

118

## 119 Outstanding questions

120 As with all good field trips, exciting discussions took place on the outcrops and during  
121 evening meals. One of the topics that got a lot of us thinking was the age of the Bambuí Group,  
122 given the continuing debates in the literature. For example, the basal Pedro Leopoldo Member ex-

123 aragonite crystal fans at Sambra Quarry have been dated by the Pb-Pb carbonate technique at 720  
124  $\pm$  22 Ma (Babinski et al., 2007), which, at the time argued for a Sturtian age for the underlying  
125 diamictite. However, both high-resolution U-Pb and Re-Os acquired in the last decade indicate  
126 that the Sturtian glaciation lasted from ca. 717 to 660 Ma (Hoffman et al., 2017, and references  
127 therein), which would make the Pb-Pb age not compatible with a cap carbonate to that specific  
128 glaciation. Recently, in-situ LA-ICP-MS U-Pb dating of the same ex-aragonite fans at Sambra  
129 Quarry yielded lower intercept dates of  $615 \pm 6$  Ma and  $608 \pm 5$  Ma (Caxito et al., 2021) consistent  
130 with a Marinoan age for the Jequitaí diamictite. Using the same technique, crystal fans from the  
131 upper Lagoa Santa Member (separated by a sequence boundary from the Pedro Leopoldo Member  
132 according to DaSilva et al., 2022) at Tatiana Quarry yielded an age of  $573 \pm 11$  Ma, while  
133 stromatolites at the top of the unit provided an age of  $566 \pm 15$  Ma, which was interpreted as early  
134 diagenetic in origin (Caxito et al., 2021).

135

136 Similarly, detrital zircons from the Lagoa Santa Member yielded a weighted mean  
137  $^{206}\text{Pb}/^{238}\text{U}$  age of  $571 \pm 3$  Ma, interpreted as a maximum depositional age for this unit (recalculated  
138 by Caxito et al., 2021 from Paula-Santos and Babinski, 2018; Paula-Santos et al., 2015). The age  
139 estimate supports the Ediacaran designation for the entire succession (Paula-Santos et al., 2015).  
140 Additionally, both SHRIMP and LA-ICPMS detrital zircon data from the Bambuí Group support  
141 derivation of most of the stratigraphic package from the erosion of Ediacaran sources located on  
142 the surrounding Brasiliano mountain belts, including samples from the upper Sete Lagoas  
143 Formation (Pimentel et al., 2011; Dias et al., 2024), conglomerate wedges on the western portion  
144 of the basin (Uhlein et al., 2017), and samples from the Serra da Saudade Formation (Paula-Santos  
145 and Babinski, 2018; Kuchenbecker et al., 2020)—all indicating original provenance in the ca. 635–

146 560 Ma age range. Samples from the overlying Serra da Saudade and Três Marias Formations  
147 show even younger detrital zircons, as young as ca. 555–520 Ma (Tavares et al., 2020; Moreira et  
148 al., 2020; Rossi et al., 2020; Dias et al., 2024).

149

150 If these ages are correct, they suggest instead that the glacial units of the Jequitaí Formation  
151 were potentially co-eval with the Gaskiers glaciation at ~579 Ma (Pu et al., 2016), or perhaps even  
152 one of the less well characterized post-Gaskiers Ediacaran glaciations (e.g. Xiao et al., 2004;  
153 Hebert et al., 2010; Le Heron et al., 2019; Linnemann et al., 2021). However, it must be noted that  
154 no Ediacaran detrital zircons have been found in either the Pedro Leopoldo member cap carbonate  
155 or in the underlying Jequitaí Formation diamictites, with the main younger zircon peak at ca. 900  
156 Ma and sparse Cryogenian zircons (see compilation in Caxito et al., 2021). This supports the  
157 interpretation of deposition in distinct basinal settings with distinct provenances (cratonic versus  
158 mountain belts) for the glacial-related basal part of the Bambuí Group (diamictite + cap carbonate)  
159 in contrast to the upper part of the succession (Lagoa Santa member upwards) and leaves the  
160 question of the age of the glaciation open for debate. Further up section in a possible ash fall tuff  
161 breccia of the Serra da Saudade Formation, LA-ICPMS U-Pb ages for the youngest population (10  
162 out of 107) of detrital zircons yield an age of  $520.2 \pm 5.3$  Ma (Moreira et al., 2020; although see  
163 DaSilva et al., 2022 for an alternative view) suggesting that the upper Bambuí Group may be  
164 Cambrian in age. The full range of U-Pb ages span from ca. 520 to nearly 2800 Ma indicating a  
165 significant admixture of inherited grains. While this minimum age is likely to be robust, the  
166 igneous horizon lies over 100 meters above carbonates that are highly  $^{13}\text{C}$  enriched as discussed  
167 above and sit below carbonates of the Jaíba Member that have  $\delta^{13}\text{C}$  values between 1 and 3.5‰,  
168 as well as a tight distribution of  $^{87}\text{Sr}/^{86}\text{Sr}$  values near 0.7080. Strontium isotope age models are

169 consistent with these carbonates being deposited either before or after Marinoan aged glacial  
170 deposits in South China and elsewhere (e.g. Cui et al., 2015; Lau et al., 2017: see Fig. 1, left panel),  
171 or terminal Ediacaran strata in Arctic Siberia (Vishnevskaya et al., 2013; Kaufman et al., 2019),  
172 but not with those from the basal Cambrian interval (cf. Kaufman et al., 1996). Thus, at present,  
173 conflicting data make the age significance of the Serra da Saudade siltstone unclear; there is always  
174 the possibility of an unknown unconformity between the highly positive  $\delta^{13}\text{C}$  interval and the  
175 volcanioclastic horizon.

176 Considering that Ediacaran carbonates with such positive extremes are rare (see Moynihan  
177 et al., 2019 for the highest recorded Ediacaran values of ca. +12‰ in a 5 meter interval within  
178 carbonate deep marine turbidites of the Nadaleen Formation of the Windermere Supergroup: Fig.  
179 1, right panel), the extended Middle Bambuí Excursion might relate to basin restriction associated  
180 with the closing of the Goiás-Brasilides and Adamastor oceans (Uhlein et al., 2019; Caetano-Filho  
181 et al., 2020; Cui et al., 2020; Caxito et al., 2021). However, physical evidence of restriction is  
182 lacking, and a recent sequence stratigraphic analysis suggests continuous connection with the open  
183 ocean (DaSilva et al., 2022). The presumed Ediacaran equivalent in northern Canada would be the  
184 Nadaleen Formation where regional mapping reveals no evidence of basinal restriction.  
185 Alternatively, the positive carbon isotope anomaly coupled with non-radiogenic  $^{87}\text{Sr}/^{86}\text{Sr}$  values  
186 (<0.7075) might be Cryogenian in age (see Kaufman et al., 2009 for an extended discussion and  
187 Fig. 9, left panel for an example from Mongolia) and hence predate the Marinoan ice age (**M** in  
188 Fig. 9) of Snowball Earth renown. In this case, the basal Bambuí diamictite might be related to  
189 Sturtian (**S** in Fig. 9) glaciation instead. Notice, however, that this interpretation is not supported  
190 by the available geochronological (detrital zircon, U-Pb in carbonate), chemostratigraphic and  
191 paleontological data available as discussed above. Thus, more detailed work is necessary to better

192 constrain the age of both the glacial record and the extreme carbon isotope fluctuations recorded  
193 in the Bambuí Group.

194 The difficulties associated with the age dating of Bambui Group sedimentary rocks has  
195 important ramifications on our continued development of the Terminal Ediacaran Stage (Xiao et  
196 al., 2016). Despite a spirited search by the field trip members, except for possible fossil fragments,  
197 no new specimens of *Cloudina* were found at either Barreiros Quarry or the Sapé section where  
198 they have been previously reported (Warren et al., 2014). Further study of the fragmentary remains  
199 that have been described are warranted but lacking more robust fossil evidence and more clarity  
200 about age constraints and isotope age models, it remains unclear whether the Jequitaí glacial  
201 diamictite and the Bambuí Group containing the MIBE is Ediacaran or Cryogenian in age. As  
202 mentioned above, if claudinomorphs are confirmed in the Lagoa Santa Member and it is Ediacaran  
203 (but pre-Shuram) in age, then the first appearance of these early experiments in biomineralization  
204 would be 10s of millions of years older than previously known. This finding would thus play an  
205 important role in subdividing the Ediacaran Period.

206

### 207 **Moving Forward:**

208 Overall, the success of the Neoproterozoic Bambuí Group field trip marks an important  
209 return to Subcommission activities (Fig. 10) following the difficulties associated with the global  
210 pandemic. Considering that this trip was originally scheduled for Spring 2020, our division is  
211 excited to renew the process of defining the “Terminal Ediacaran Series” (Xiao et al., 2016) to  
212 which the Bambuí Group will feature prominently in the debate.

213

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218

219 **References:**

220 Babinski, M., Vieira, L.C., and Trindade, R.I. 2007. Direct dating of the Sete Lagoas cap carbonate  
221 (Bambuí Group, Brazil) and implications for the Neoproterozoic glacial events. *Terra Nova*,  
222 19(6), 401–406.

223

224 Caetano-Filho, S., Sansjofre, P., Ader, M., Paula-Santos, G., Babinski, M., Bedoya-Rueda, C., and  
225 Kuchenbecker, M., 2020, A large epeiric methanogenic Bambuí sea in the core of Gondwana  
226 supercontinent? *Geoscience Frontiers*, v. 12, pp. 203–218.

227

228 Caxito, F., Halverson, G.P., A., Uhlein, A., Stevenson, R., Dias, T.G., and Uhlein, G.J., 2012,  
229 Marinoan glaciation in east central Brazil: *Precambrian Research*, v. 200, pp.38–58.

230

231 Caxito, F., Lana, C., Frei, R., Uhlein, G.J., Sial, A.N., Dantas, E.L., Pinto, A.G., Campos, F.C.,  
232 Galvão, P., Warren, L. V., Okubo, J., and Ganade, C.E., 2021, Goldilocks at the dawn of  
233 complex life: mountains might have damaged Ediacaran–Cambrian ecosystems and prompted  
234 an early Cambrian greenhouse world. *Scientific Reports*, v. 11, pp. 1–15.

235

236 Cui, H., Kaufman, A.J., Xiao, S., Zhu, M., Zhou, C., and Liu, X.M. 2015. Redox architecture of  
237 an Ediacaran ocean margin: Integrated chemostratigraphic ( $\delta^{13}\text{C}$ – $\delta^{34}\text{S}$ – $^{87}\text{Sr}/^{86}\text{Sr}$ –Ce/Ce\*)  
238 correlation of the Doushantuo Formation, South China. *Chemical Geology*, 405, 48–62.

- 239 Cui, H., Warren, L.V., Uhlein, G.J., Okubo, J., Liu, X.M., Plummer, R.E., Baele, J.M., Goderis, S.,  
240 Claeys, P., and Li, F., 2020, Global or regional? Constraining the origins of the middle Bambuí  
241 carbon cycle anomaly in Brazil. *Precambrian Research*, v. 348, pp 105861.
- 242
- 243 DaSilva, L.G., Pufahl, P.K., James, N.P., Guimaraes, E.M. and Reis, C., 2022, Sequence  
244 stratigraphy and paleoenvironmental significance of the Neoproterozoic Bambui Group,  
245 Central Brazil. *Precambrian Research*, v. 379, pp. 106710.
- 246
- 247 Dias A.N.C., Martins-Ferreira M.A.C., Pereira V.Q., Sales A.S.W., Chemale Jr. F.; Insights into  
248 the Phanerozoic evolution of the São Francisco Craton based on detrital zircon  
249 thermochronology and U-Pb-Hf geochronology. *GSA Bulletin* 2024;  
250 doi: <https://doi.org/10.1130/B37281.1>
- 251
- 252 Hebert, C.L., Kaufman, A.J., Penniston-Dorland, S.C. and Martin, A.J., 2010, Radiometric and  
253 stratigraphic constraints on terminal Ediacaran (post-Gaskiers) glaciation and metazoan  
254 evolution. *Precambrian Research*, v. 182, pp. 402-412.
- 255
- 256 Hoffman, P.F., Abbot, D.S., Ashkenazy, Y., Benn, D.I., Brocks, J.J., Cohen, P.A., Cox, G.M.,  
257 Creveling, J.R., Donnadieu, Y., Erwin, D.H., and Fairchild, I.J. 2017. Snowball Earth climate  
258 dynamics and Cryogenian geology-geobiology. *Science Advances*, 3(11), e1600983.
- 259

- 260 Kaufman, A.J., Knoll, A.H., Semikhatov, M.A., Grotzinger, J.P., Jacobsen, S.B., and Adams, W.  
261 1996. Integrated chronostratigraphy of Proterozoic–Cambrian boundary beds in the western  
262 Anabar region, northern Siberia. *Geological Magazine*, 133(5), 509–533.
- 263
- 264 Kaufman et al., 2009
- 265
- 266 Kaufman et al., 2019
- 267
- 268 Kuchenbecker, M., Pedrosa-Soares, A. C., Babinski, M., Reis, H. L. S., Atman, D., & da Costa, R.  
269 D. (2020). Towards an integrated tectonic model for the interaction between the Bambuí basin  
270 and the adjoining orogenic belts: Evidences from the detrital zircon record of syn-orogenic  
271 units. *Journal of South American Earth Sciences*, 104, 102831.
- 272
- 273 Lau, K.V., Macdonald, F.A., Maher, K., and Payne, J.L. 2017. Uranium isotope evidence for  
274 temporary ocean oxygenation in the aftermath of the Sturtian Snowball Earth. *Earth and Planetary  
275 Science Letters*, 458, 282–292.
- 276
- 277 Le Heron, D.P., Vandyk, T.M., Kuang, H., Liu, Y., Chen, X., Wang, Y., Yang, Z., Scharfenberg, L.,  
278 Davies, B. and Shields, G., 2019, Bird's-eye view of an Ediacaran subglacial landscape.  
279 *Geology*, v. 47, pp.705-709.
- 280
- 281 Linnemann, U., Hofmann, M., Gärtner, A., Gärtner, J., Zieger, J., Krause, R., Haenel, R., Mende,  
282 K., Ovtcharova, M., Schaltegger, U. and Vickers-Rich, P., 2022, An Upper Ediacaran Glacial

- 283      Period in Cadomia: the Granville tillite (Armorican Massif)–sedimentology, geochronology and  
284      provenance. *Geological Magazine*, v. 159, pp.999–1013.
- 285
- 286      Macdonald, F.A., Strauss, J.V., Sperling, E.A., Halverson, G.P., Narbonne, G.M., Johnston, D.T.,  
287      Kunzmann, M., Schrag, D.P., and Higgins, J.A. 2013. The stratigraphic relationship between  
288      the Shuram carbon isotope excursion, the oxygenation of Neoproterozoic oceans, and the first  
289      appearance of the Ediacara biota and bilaterian trace fossils in northwestern Canada. *Chemical  
290      Geology*, 362, 250–272.
- 291
- 292      Moreira, D.S., Uhlein, A., Dussin, I.A., Uhlein, G.J., and Misuzaki, A.M.P., 2020, A Cambrian age  
293      for the upper Bambuí Group, Brazil, supported by the first U-Pb dating of volcanoclastic bed.  
294      *Journal of South American Earth Sciences*, v. 99, pp. 102503.
- 295
- 296      Moynihan, D.P., Strauss, J.V., Nelson, L.L., and Padgett, C.D. 2019. Upper Windermere  
297      Supergroup and the transition from rifting to continent-margin sedimentation, Nadaleen River  
298      area, northern Canadian Cordillera. *GSA Bulletin*, 131(9-10), 1673–1701.
- 299
- 300      Okubo, J., Inglez, L., Uhlein, G.J., Warren, L. V, Xiao, S., 2023, Simple structures and complex  
301      stories: potential microbially induced sedimentary structures in the Ediacaran Serra De Santa  
302      Helena Formation, Bambuí Group, eastern Brazil. *Palaios*, v. 38, pp. 188–209.
- 303
- 304      Paula-Santos, G.M., Babinski, M., Kuchenbecker, M., Caetano-Filho, S., Trindade, R.I., and  
305      Pedrosa-Soares, A.C., 2015, New evidence of an Ediacaran age for the Bambuí Group in

- 306 southern São Francisco craton (eastern Brazil) from zircon U-Pb data and isotope  
307 chemostratigraphy. *Gondwana Research*, v. 28, pp. 702–720.
- 308
- 309 Paula-Santos, G.M., and Babinski, M., 2018, Sedimentary provenance in the southern sector of  
310 the São Francisco Basin, SE Brazil. *Brazilian Journal of Geology*, v. 48, pp. 51–74.
- 311
- 312 Pimentel, M. M., Rodrigues, J. B., DellaGiustina, M. E. S., Junges, S., Matteini, M., & Armstrong,  
313 R. (2011). The tectonic evolution of the Neoproterozoic Brasília Belt, central Brazil, based on  
314 SHRIMP and LA-ICPMS U–Pb sedimentary provenance data: a review. *Journal of South  
315 American Earth Sciences*, 31(4), 345–357.
- 316
- 317 Pu, J.P., Bowring, S.A., Ramezani, J., Myrow, P., Raub, T.D., Landing, E., Mills, A., Hodgin, E.  
318 and Macdonald, F.A., 2016, Dodging snowballs: Geochronology of the Gaskiers glaciation and  
319 the first appearance of the Ediacaran biota. *Geology*, v. 44, pp. 955–958.
- 320
- 321 Reis, H.L.S., Alkmim, F.F., Fonseca, R.C.S., Nascimento, T.C., Suss, J.F., and Prevatti, L.D., 2017,  
322 The São Francisco Basin. *Regional Geology Reviews*. Springer, Cham., pp. 117–143.
- 323
- 324 Rooney, A.D., Cantine, M.D., Bergmann, K.D., Gómez-Pérez, I., Baloushi, B.A., Boag, T.H.,  
325 Busch, J.F., Sperling, E.A., and Strauss, J.V., 2020, Calibrating the coevolution of Ediacaran  
326 life and environment. *Earth, Atmospheric, And Planetary Sciences*, v. 117, pp. 16824–16830.
- 327

- 328 Rossi, A. V., Danderfer Filho, A., Bersan, S. M., Kelmer, L. R., Tavares, T. D., & de Carvalho  
329 Lana, C. (2020). Stratigraphic, isotopic, and geochronological record of a superposed pro-  
330 foreland basin in the eastern São Francisco craton during west Gondwana  
331 amalgamation. *Journal of South American Earth Sciences*, 97, 102406.
- 332
- 333 Tavares, T.D., Martins, M. De S., Alkmim, F.F., and Lana, C., 2020, Detrital zircons from the  
334 Upper Três Marias Formation, São Francisco basin, SE Brazil: record of foreland deposition  
335 during the Cambrian? *Journal of South American Earth Sciences*, v. 97, pp. 102395.
- 336
- 337 Uhlein, G. J., Uhlein, A., Stevenson, R., Halverson, G. P., Caxito, F. A., & Cox, G. M. (2017).  
338 Early to late Ediacaran conglomeratic wedges from a complete foreland basin cycle in the  
339 southwest São Francisco Craton, Bambuí Group, Brazil. *Precambrian Research*, 299, 101-116.
- 340
- 341 Uhlein, G.J., Uhlein, A., Pereira, E., Caxito, F.A., Okubo, J., Warren, L.V., and Sial, A.N., 2019,  
342 Ediacaran paleoenvironmental changes recorded in the mixed carbonate/siliciclastic Bambuí  
343 Basin, Brazil. *Palaeogeography, Palaeoclimatology, Palaeoecology*, v. 517 pp. 39–51.
- 344
- 345 Vishnevskaya et al., 2013
- 346
- 347 Warren, L.V., Quaglio, F., Riccomini, C., Simões, M.G., Poire, D.G., Strikis, N.M., Anelli, L.E.,  
348 Strikis, P.C., 2014, The puzzle assembled: Ediacaran guide fossil *Cloudina* reveals an old proto-  
349 Gondwana seaway. *Geology*, v. 42, pp. 391–394.
- 350

351 Xiao, S., Bao, H., Wang, H., Kaufman, A.J., Zhou, C., Li, G., Yuan, X. and Ling, H., 2004, The  
352 Neoproterozoic Quruqtagh Group in eastern Chinese Tianshan: evidence for a post-Marinoan  
353 glaciation. *Precambrian Research*, v. 130, pp. 1–26.

354

355 Xiao, S., Narbonne, G.M., Zhou, C., Laflamme, M., Grazhdankin, D.V., Moczydowska-Vidal, M.,  
356 and Cui, H., 2016, Towards an Ediacaran time scale: problems, protocols, and prospects.  
357 *Episodes*, v. 39, pp. 540–555.

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## Figure Captions

361   **Figure 1.** Members of the Cryogenian and Ediacaran subcommittees analyzing Bambuí Basin  
362   drill cores. From left clockwise: Huan Cui, Xiao-Dong Shang, Ying Zhou, Jay Kaufman, Graham  
363   Shields, Erik Sperling, Fabricio Caxito, Gabriel Correa Antunes, and Brandt Gibson. Photo from  
364   Huan Cui.

365

366   **Figure 2.** Members of the Ediacaran subcommission discussing the origin of the diamictite from  
367   the Jequitaí Fm. From left: Marc Laflamme, Shuhai Xiao, and Juliana Okubo. Photo from Huan  
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369

370   **Figure 3.** Simplified stratigraphic column and compilation of carbon isotopic data for the  
371   Bambuí Group (Caxito et al., 2021 and references therein)

372

373   **Figure 4.** Black columnar stromatolites (left) and aragonite fans (right) of the uppermost Sete  
374   Lagoas Formation (Lagoa Santa Member). Photo from Huan Cui.

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377   Tatiana Quarry. In this section there are several intervals of dolomite fans a few meters below the  
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381   glacial surface. Photo from Marc Laflamme.

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385

386   **Figure 8.** Sunlight inside the Janelões cave. Photo from Marc Laflamme.

387

388   **Figure 9.** Comparative Cryogenian and Ediacaran Earth history based on the presence of  
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390 Period claudinomorphs, and chemostratigraphic trends in both carbon and strontium isotopes.  
391 (Left panel) Cryogenian Taishir Formation of Mongolia (see Bold et al., 2015 and Lau et al.,  
392 2017), which lies between Sturtian and Marinoan glacial deposits and preserves a profound  
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405 missing in the Brazilian Bambuí Group.

406

407 **Figure 10.** Members of the Ediacaran and Cryogenian subcommissions at the end of a working  
408 day. From left to right, Chuan Yang, Ben Yang, Graham Shields-Zhou, Maoyan Zhu, Huan Cui,  
409 Tara Selly, Casey Bennett, Carolina Reis, Fabrício Caxito, Ying Zhou, Gabriel Uhlein, Juliana  
410 Okubo, Johannes Zieger, Gabriel Antunes, Mandy Zieger-Hofmann, Shuhai Xiao, Erik Sperling,  
411 Lucas Warren, Marc Laflamme, Brandt Gibson, Ulf Linneman, Xiao-Dong Shang.

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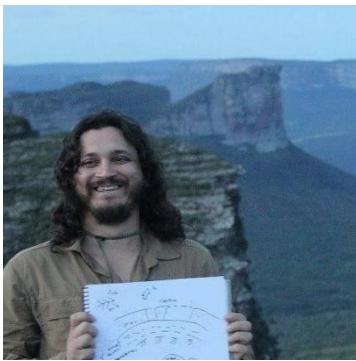
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Marc Laflamme is a professor of paleontology at the University of Toronto Mississauga. His research examines the origins, evolution, function, and preservation of the Ediacara biota. He is the current Chair of the International Subcommission on Ediacaran Stratigraphy of the International Commission of Stratigraphy in the International Union of Geological Sciences.



Lucas Veríssimo Warren Lucas Veríssimo Warren is an associate professor of sedimentology and biostratigraphy at the São Paulo State University, Brazil. Lucas focuses his research on basin analysis with an emphasis on sedimentology, sequence stratigraphy and Precambrian biostratigraphy mainly in Ediacaran successions in South America and Africa. He co-led the joint subcommission field trip to Brazil in 2023 and is currently secretary of the International Subcommission on Ediacaran Stratigraphy.



Fabrício de Andrade Caxito is a professor of geology at Universidade Federal de Minas Gerais, Brazil. His research is focused on applying field and geochemical tools to understand the interlinked evolution of the solid Earth and life that thrives on it. He is a voting member of the International Subcommission on Cryogenian Stratigraphy and co-led the joint subcommission field trip to Brazil in 2023.



Tara Selly's is a Research Assistant Professor and Assistant Director of the X-ray Microanalysis Lab at the University of Missouri. Her research utilizes modern microscopic techniques to answer questions about the origins of macroscopic life on Earth. Through this work, she explores how late Ediacaran (~550 million-year-old) organisms became incorporated into the fossil record and how their preservation influences paleontologists' ability to interpret their original biology.



Alan J. Kaufman is a Professor at the University of Maryland. His research has focused on the determination of changes in the isotopic composition of the oceans through time, by the analysis of stratigraphic suites of little-altered carbonate rocks. Thus far, most of these studies have centered around Neoproterozoic (ca. 1000-544 million-year-old) sedimentary successions in Svalbard/East Greenland, Namibia, arctic Canada and Alaska, India, and the western USA.



Graham A. Shields is a Professor from the University College London. His research utilizes geochemical and isotopic tracers to study the composition of past oceans and atmosphere, and the coevolution with life through crucial junctures in Earth history. His research group develops proxies to trace biogeochemical fluxes and related feedbacks that govern oxygen, carbon dioxide and nutrient budgets on Earth.



Shuhai Xiao is a professor of geobiology at Virginia Tech. He integrates paleobiological, sedimentological, and geochemical data to investigate the Precambrian Earth history, with a focus on the Ediacaran Period. In 2012–2020, he served as the Chair of the International Subcommission on Ediacaran Stratigraphy of the

International Commission of Stratigraphy in the International Union of Geological Sciences.



Maoyan Zhu is a research professor of geology and paleontology at the Nanjing Institute of Geology and Palaeontology, Chinese Academy of Sciences. His research focuses on early evolution of multicellular organisms and animals, Neoproterozoic-Cambrian stratigraphy and palaeo-environmental changes. He is the current Chair of the International Subcommission on Cryogenian Stratigraphy of the International Commission of Stratigraphy in the International Union of Geological Sciences.

Figure 1



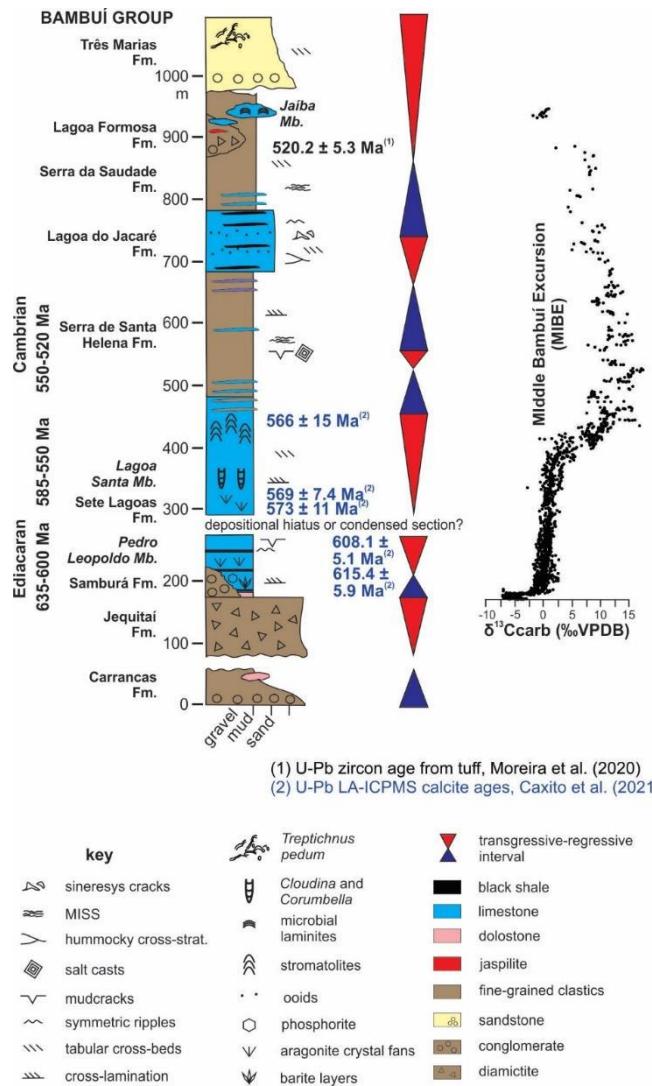
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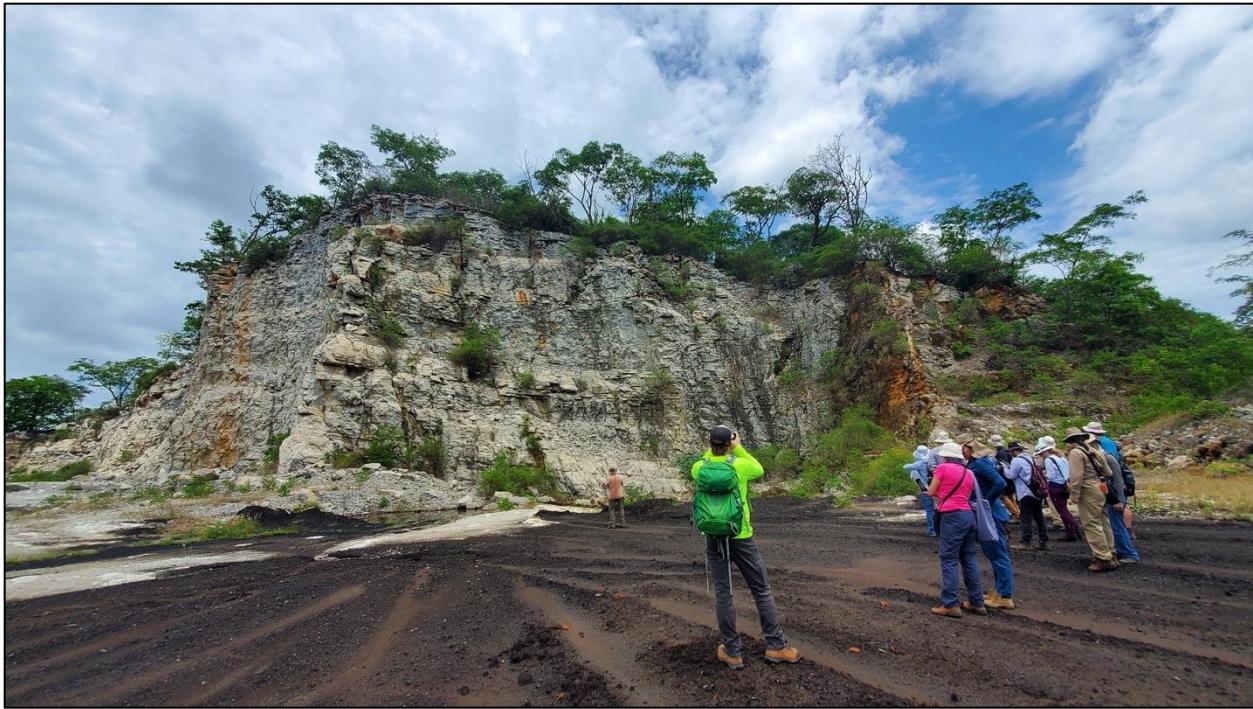
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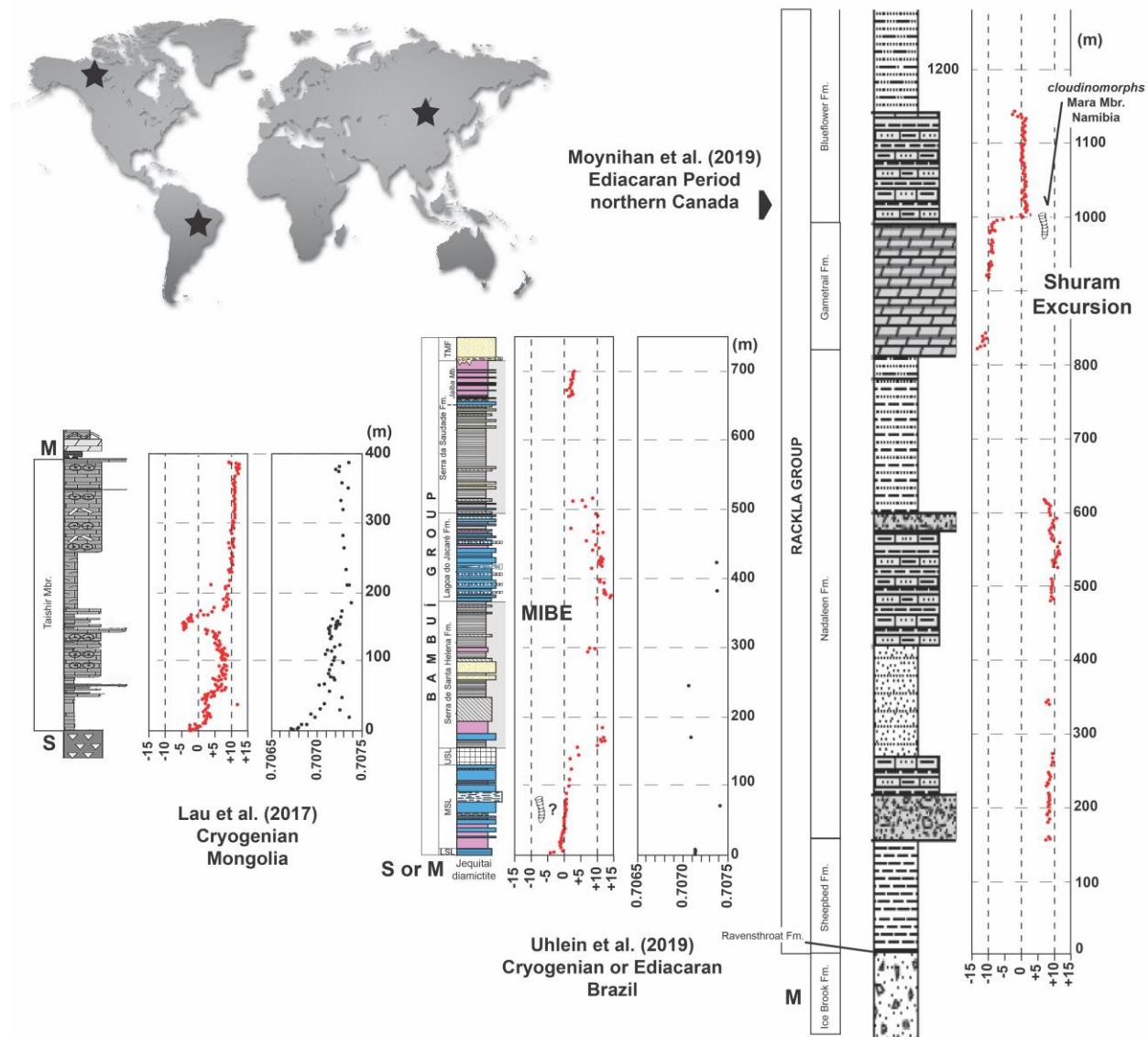


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