



# A decision-making framework for sedimentation analyses in dammed river corridor impoundments

**Ian M. Nesbitt<sup>1</sup>, Sean MC Smith<sup>1</sup>, Bess G. Koffman<sup>2</sup>, Seth W. Campbell<sup>1</sup>, Steven Arcone<sup>3</sup>**

<sup>1</sup> University of Maine, School of Earth and Climate Sciences

<sup>2</sup> Colby College, Department of Geology

<sup>3</sup> Cold Regions Research and Engineering Laboratory (CRREL), US Army Corps of Engineers



# A decision-making framework for sedimentation analyses in dammed river corridor impoundments

or

## How to pick a coring site for your study

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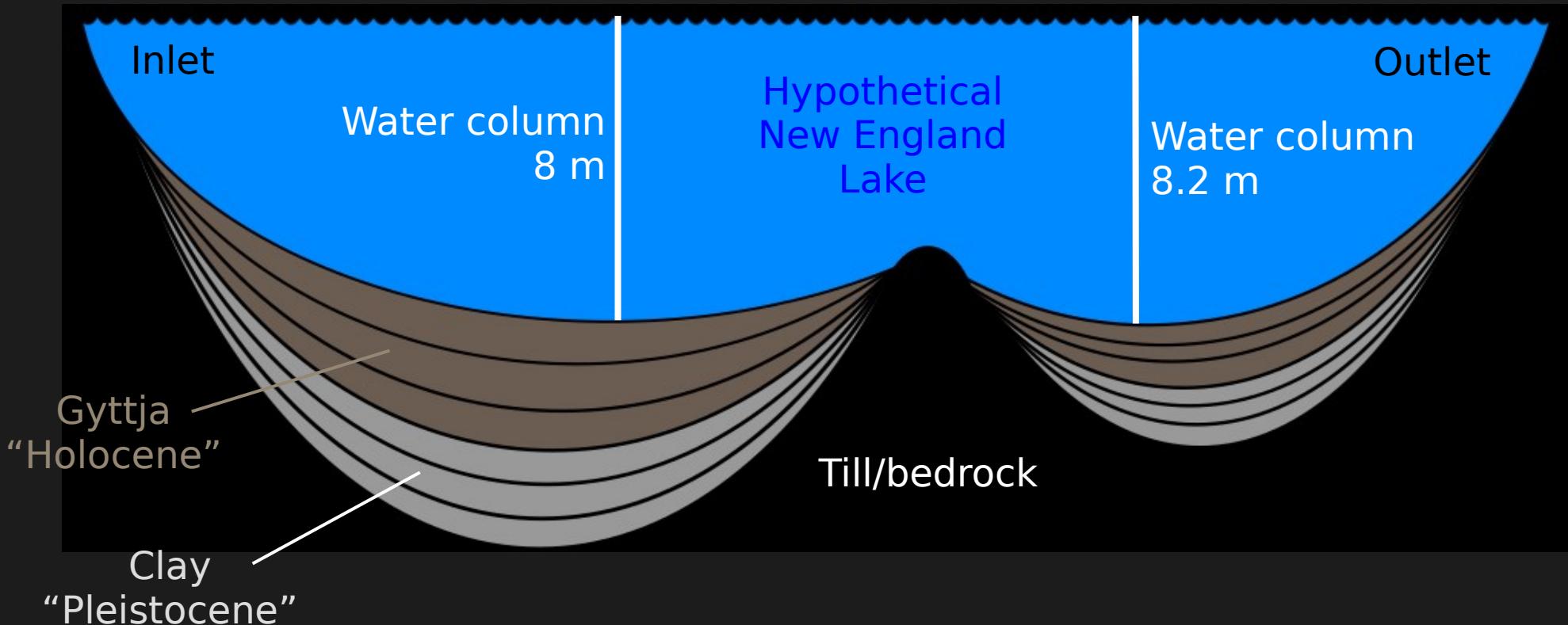
# How do we decide where to core?

- Why are we coring in the first place?
  - Sedimentation effects of damming a corridor
  - Micro- / macrofossils
  - Dating / sedimentation rates
  - Lake level change / paleovegetation / climate proxies
- How do we make sure we know where to core and what we'll be coring into for a given study?

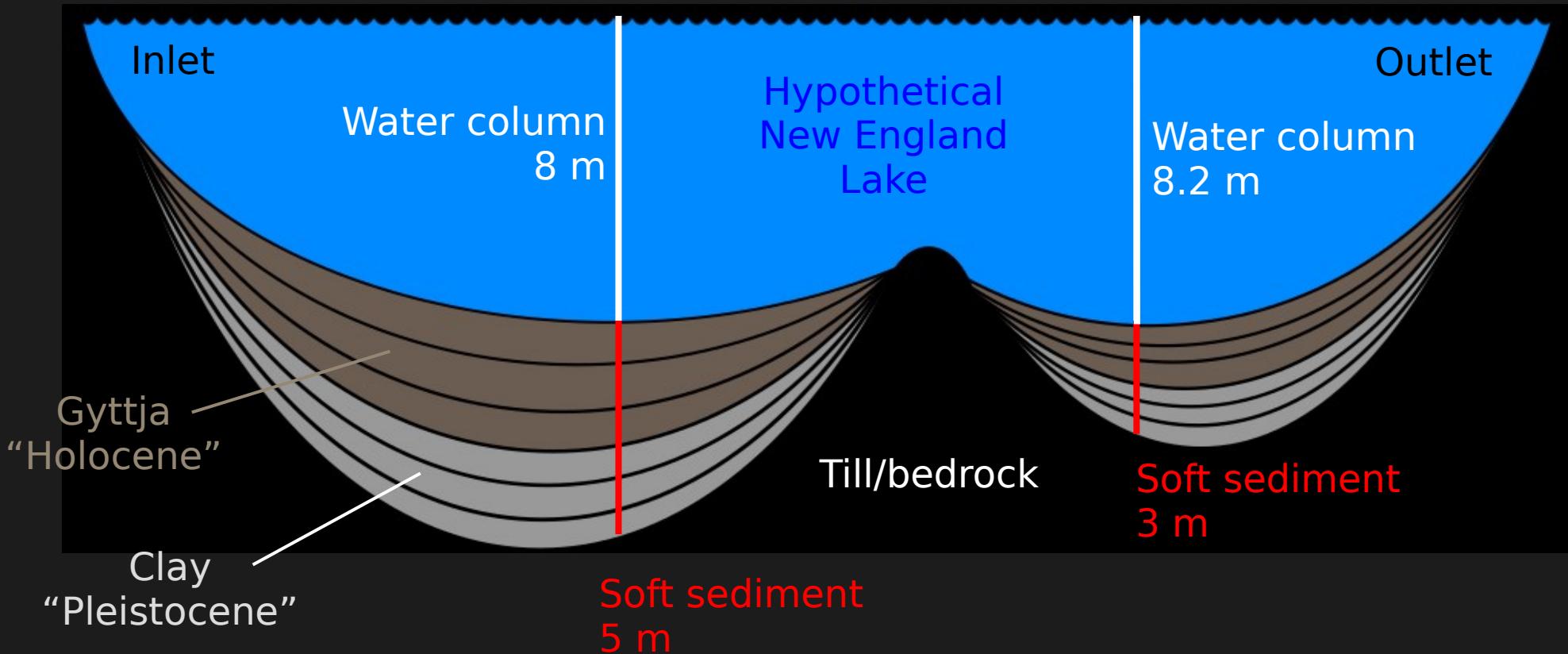
# Where would you core?



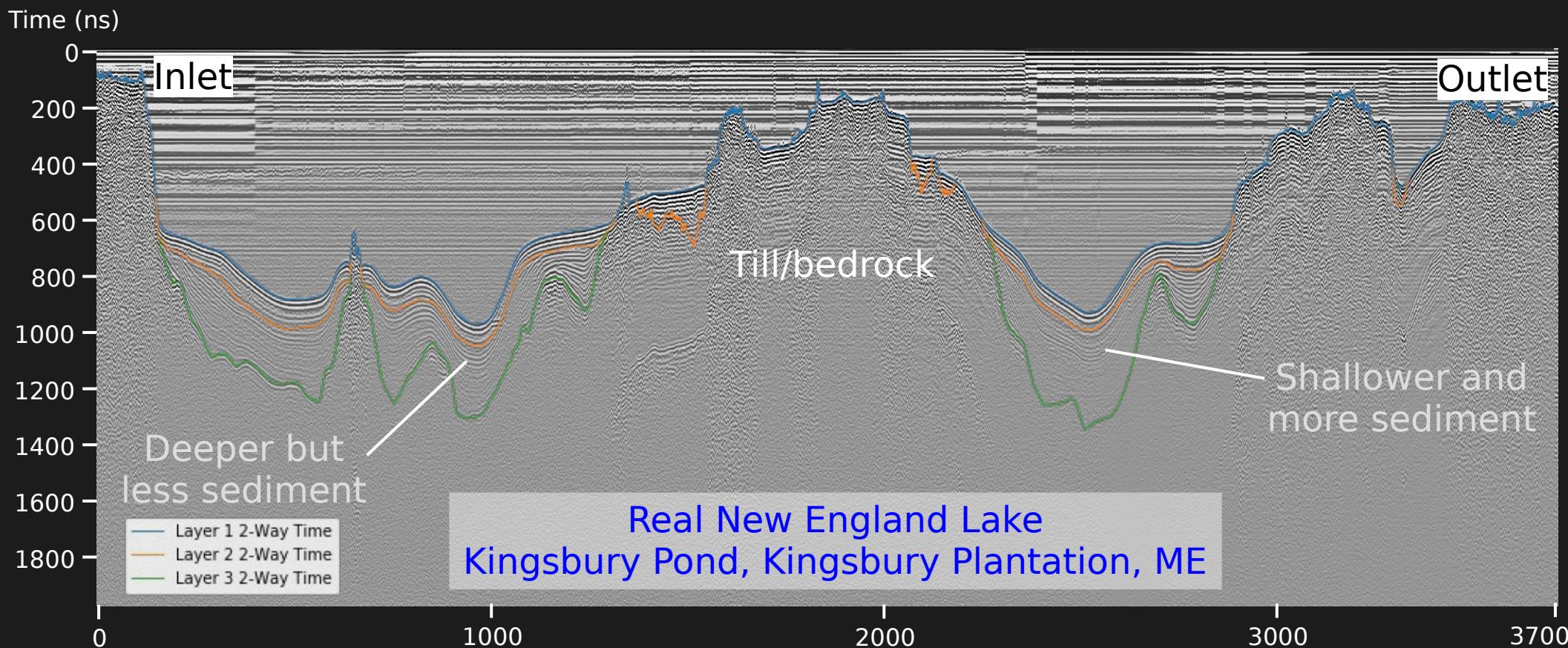
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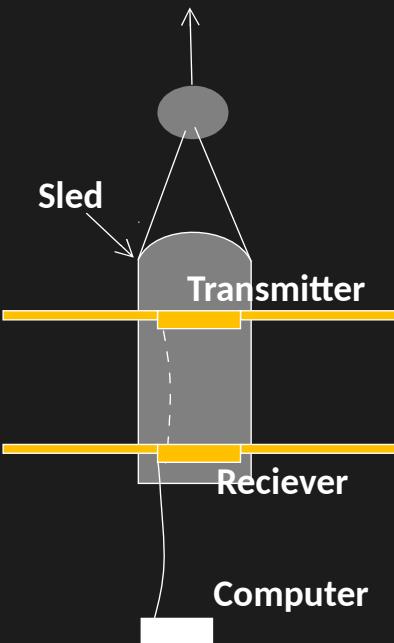
# Background

- Coring community has known about problems with spatial variability in New England for decades (Jacobson and Bradshaw, 1981; Davis and Ford, 1982)
- New England shallow lake GPR is not new (Arcone, 2018)
- GPR has been used to find core sites, but infrequently and in two dimensions (ex: Dieffenbacher-Krall and Nurse, 2005)

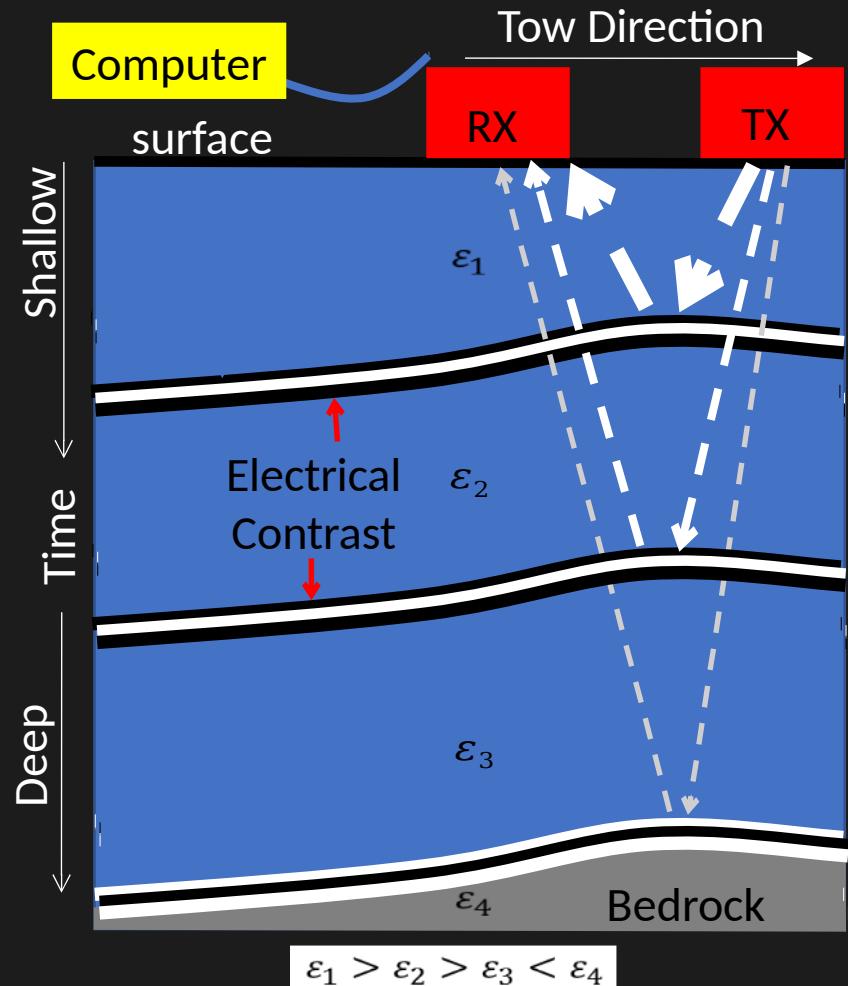
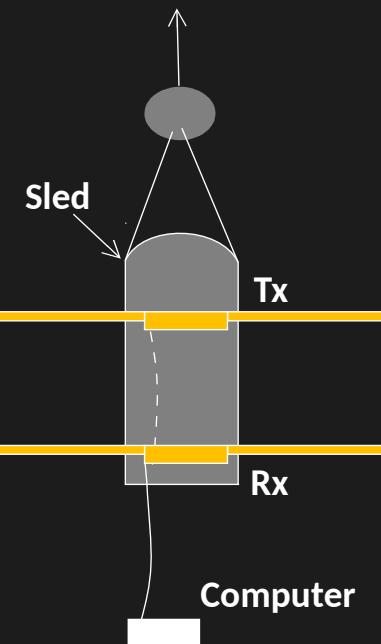
# Ground-Penetrating Radar (GPR) or Impulse Radar



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$$\varepsilon_1 > \varepsilon_2 > \varepsilon_3 < \varepsilon_4$$

# Ground-Penetrating Radar (GPR) or Impulse Radar

Permittivity ( $\epsilon$ ) - Material ability to store or release EM energy

$$\epsilon_r = \frac{\epsilon_{\text{material}}}{\epsilon_{\text{vacuum}}} \propto \frac{C_{\text{material}}}{C_{\text{vacuum}}}$$

$\epsilon_r$  values:

Air = 1

Dry snow = 1.4

Dry Firn = 2.2-2.6

Ice = 3.0-3.2

Wet snow = 4-6

Granite = 6-12

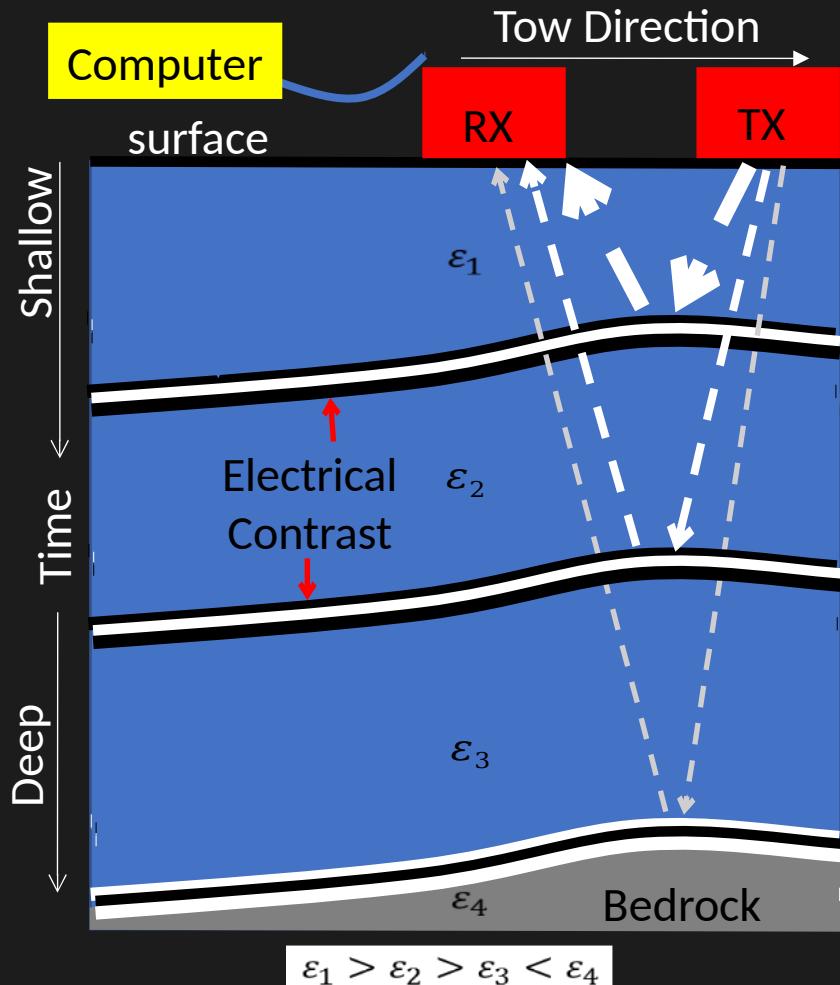
Permafrost = 5-6

Till = 12-32

Sands = 12-32

Water = 80-88

$$z = \frac{t_n - t_{n-1}}{2} * \frac{c}{\sqrt{\epsilon_r(n)}}$$



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- Cons:

- Water must be very fresh (typically  $<50 \mu\text{s}/\text{m}$ )
- Shallow bathymetry works best ( $<20 \text{ m}$ )
- Doesn't do well in hydrocarbon-rich layers (but better than sub-bottom acoustic)

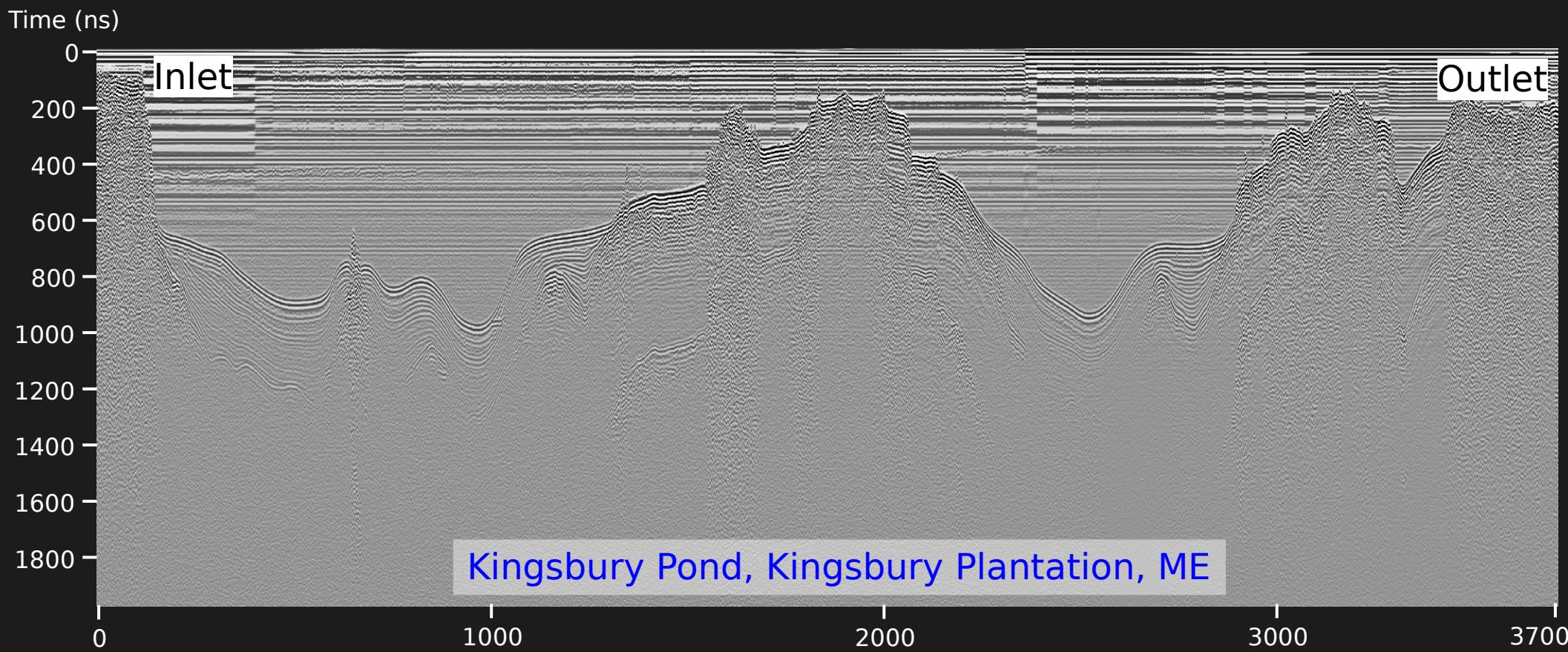
- Pros:

- Portable, unlike sub-bottom acoustic (walk, ski, paddle, motor)
- Can resolve stratigraphic detail
- "Relatively inexpensive"

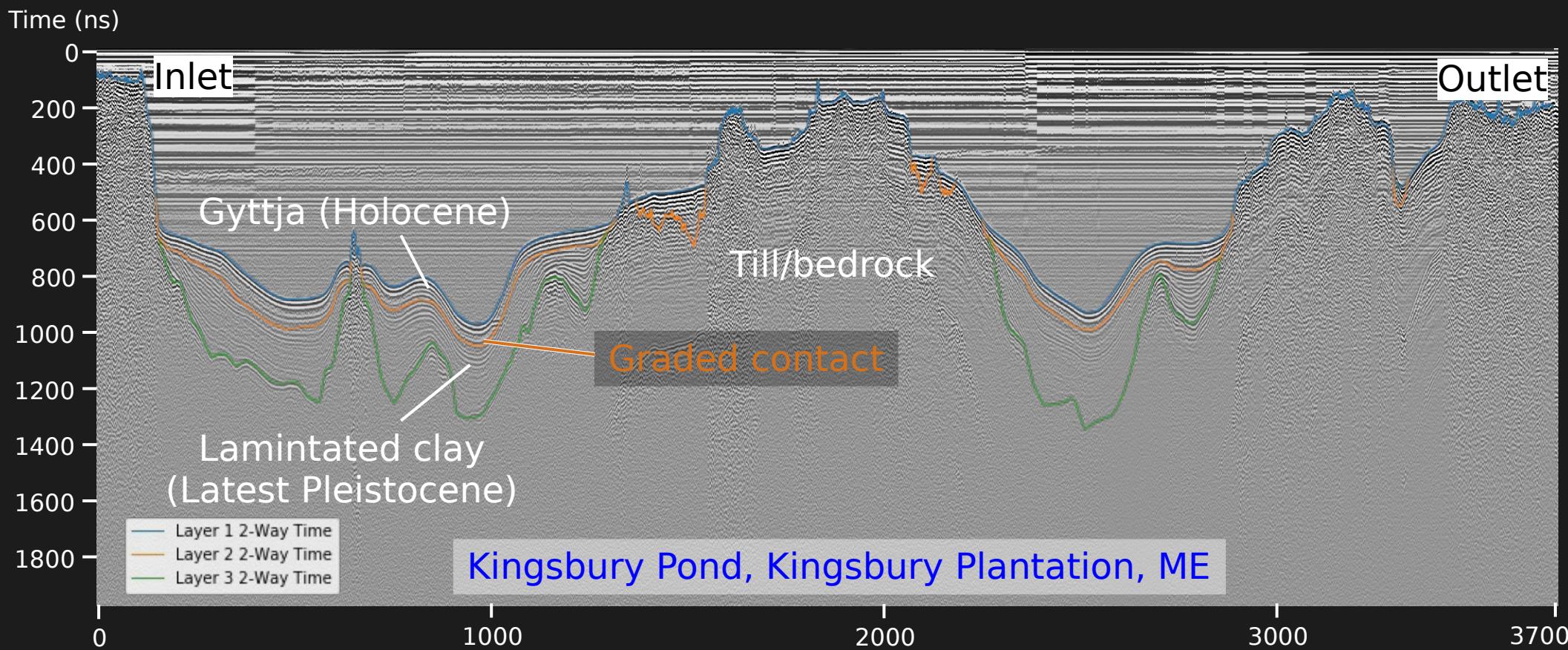
# Field scenarios



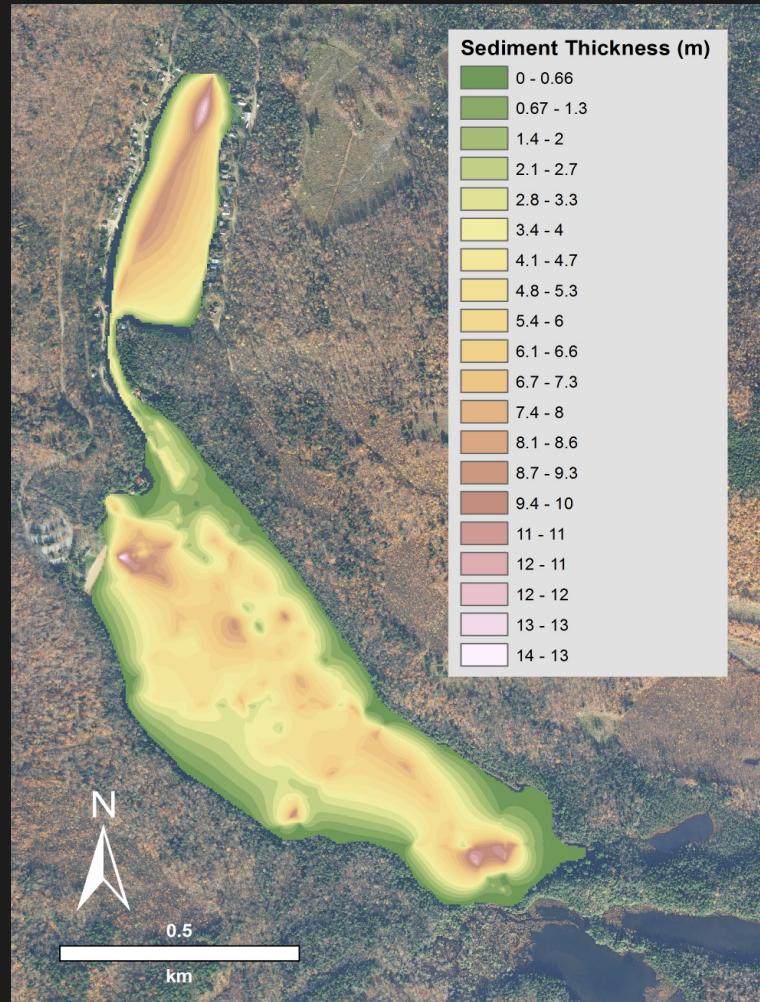
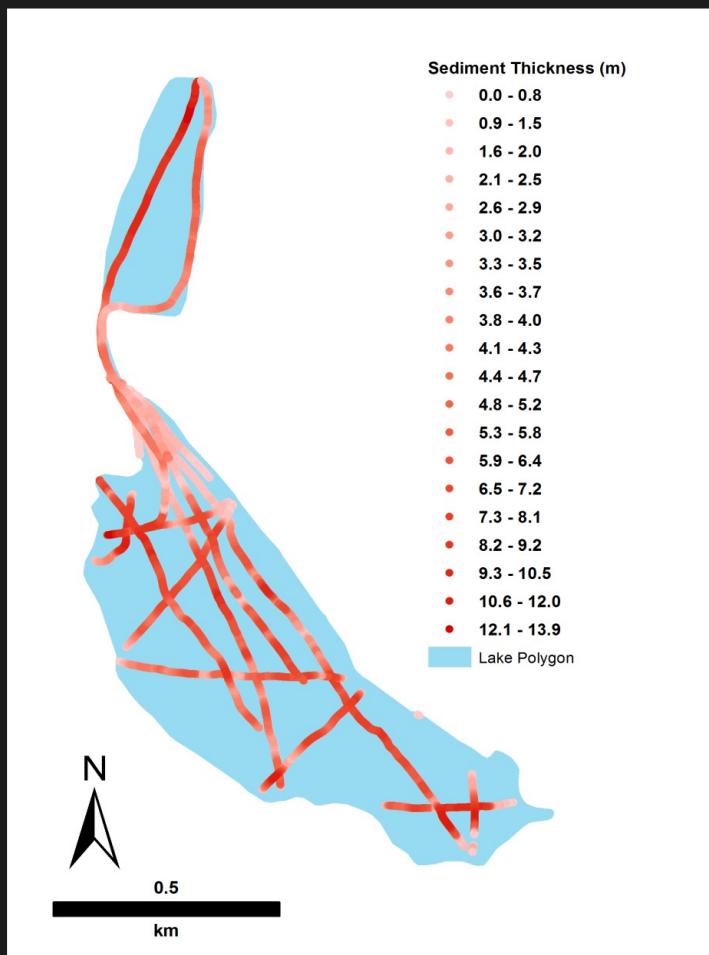
# Picking and depth calibration



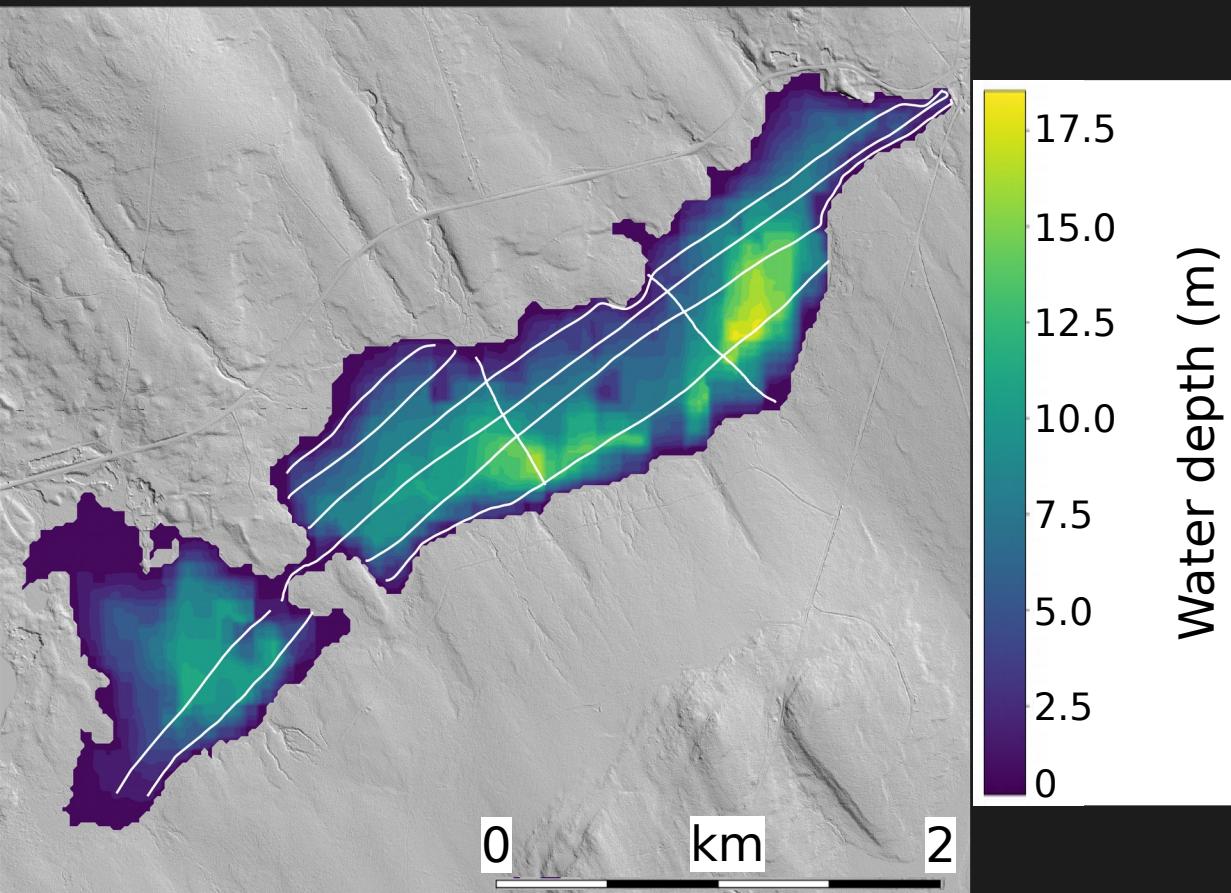
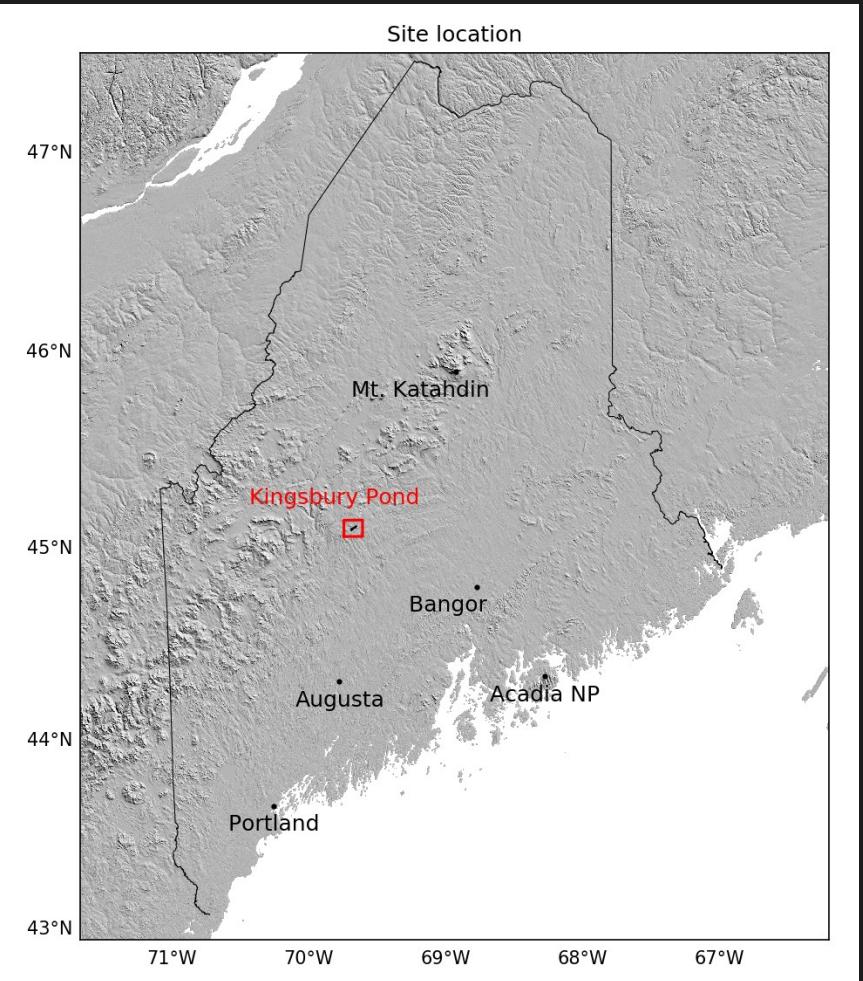
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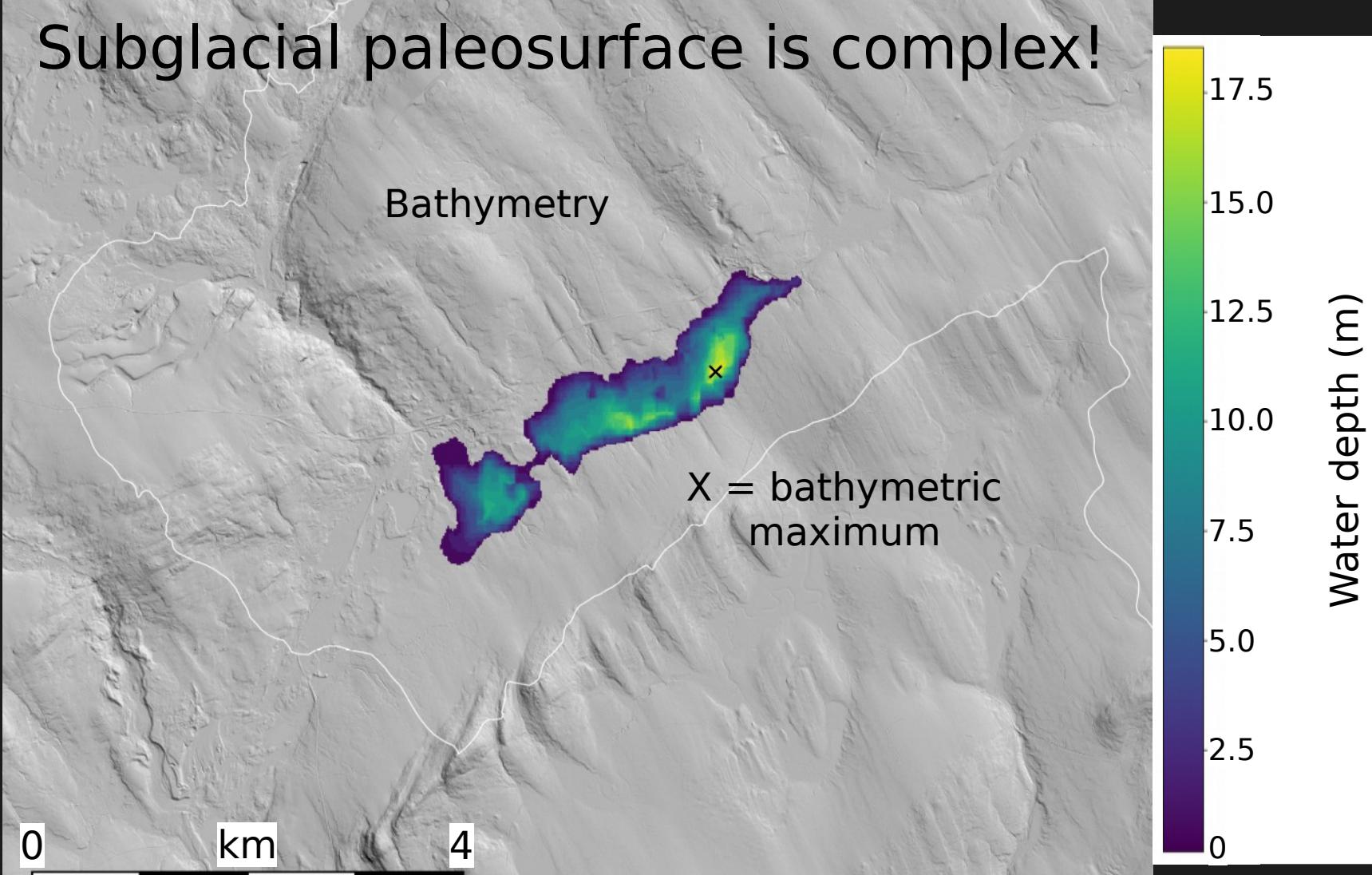
# Thickness map creation



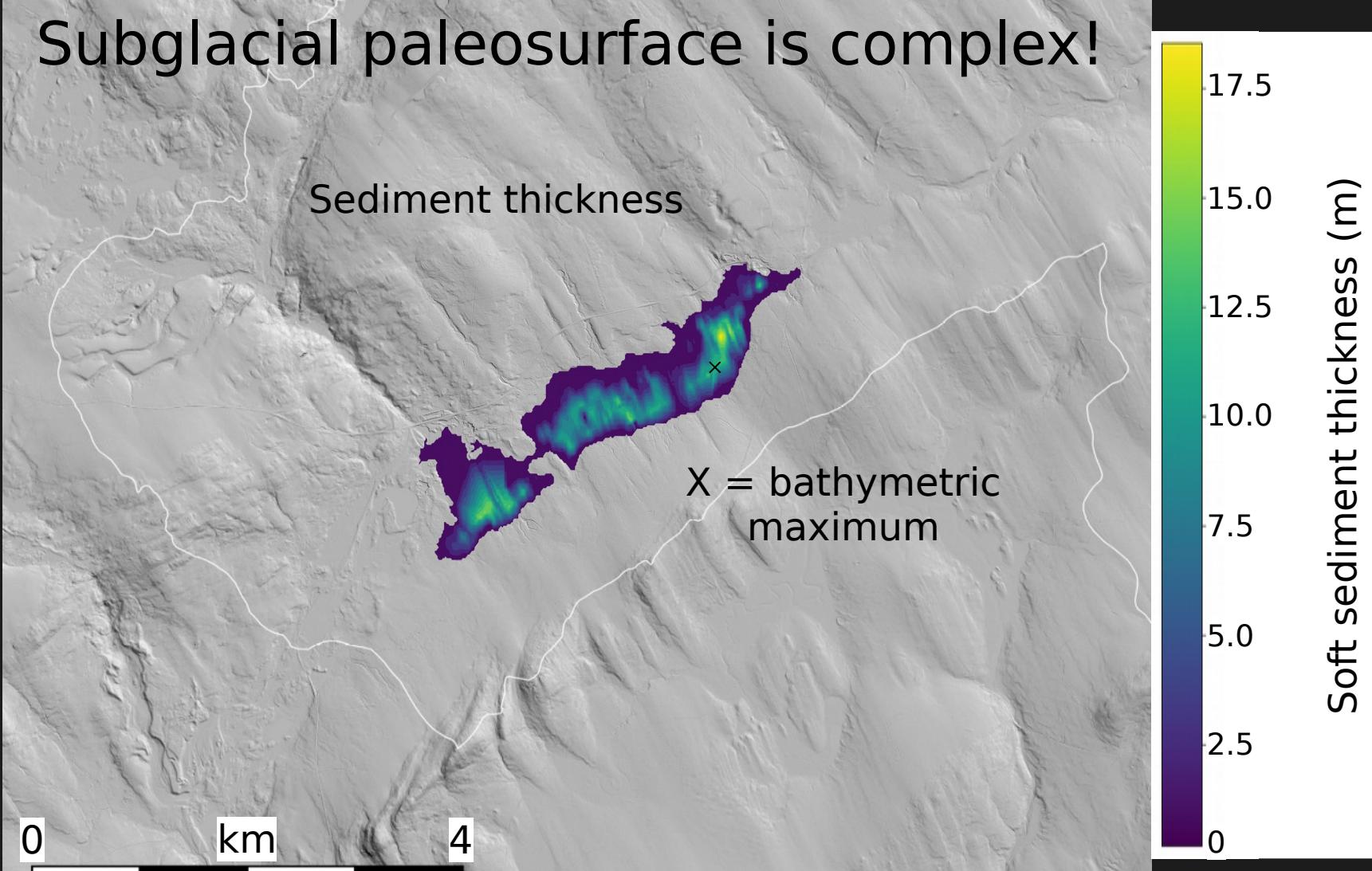
# Study site



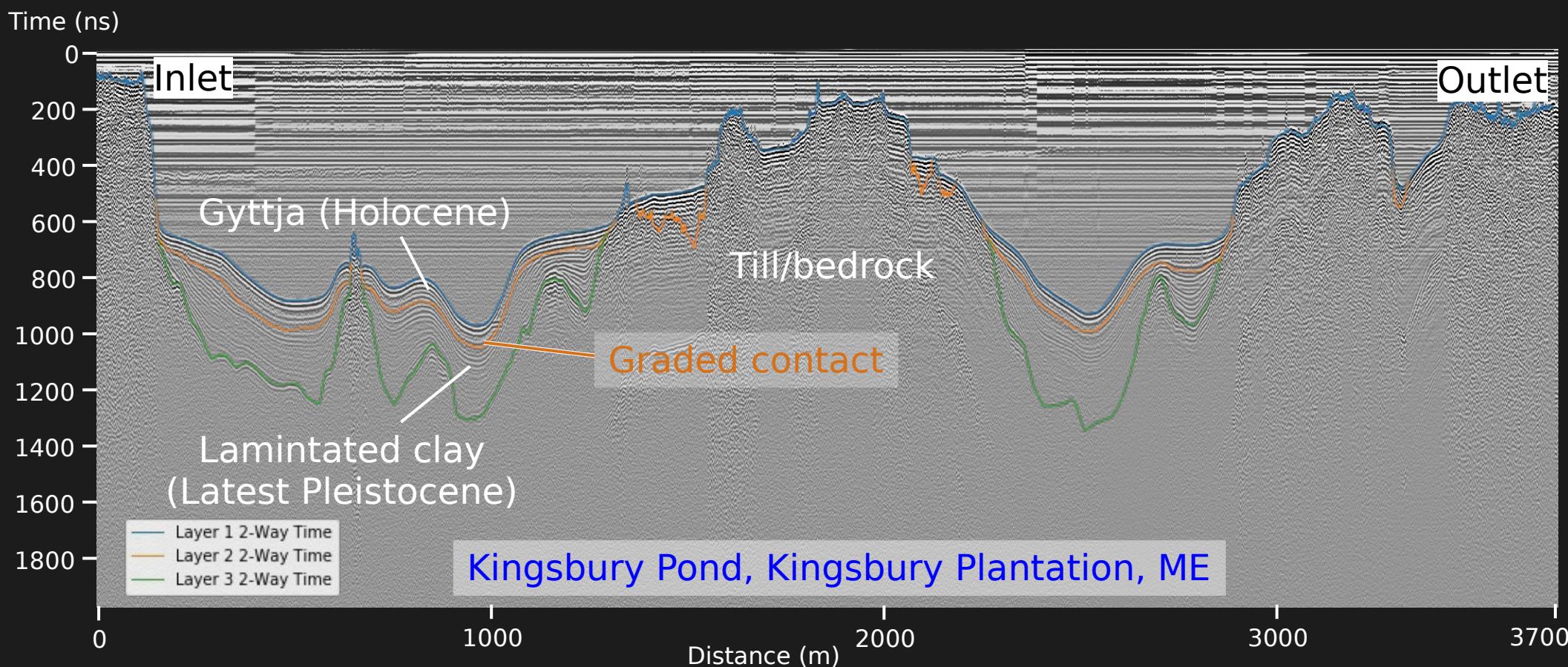
# Subglacial paleosurface is complex!

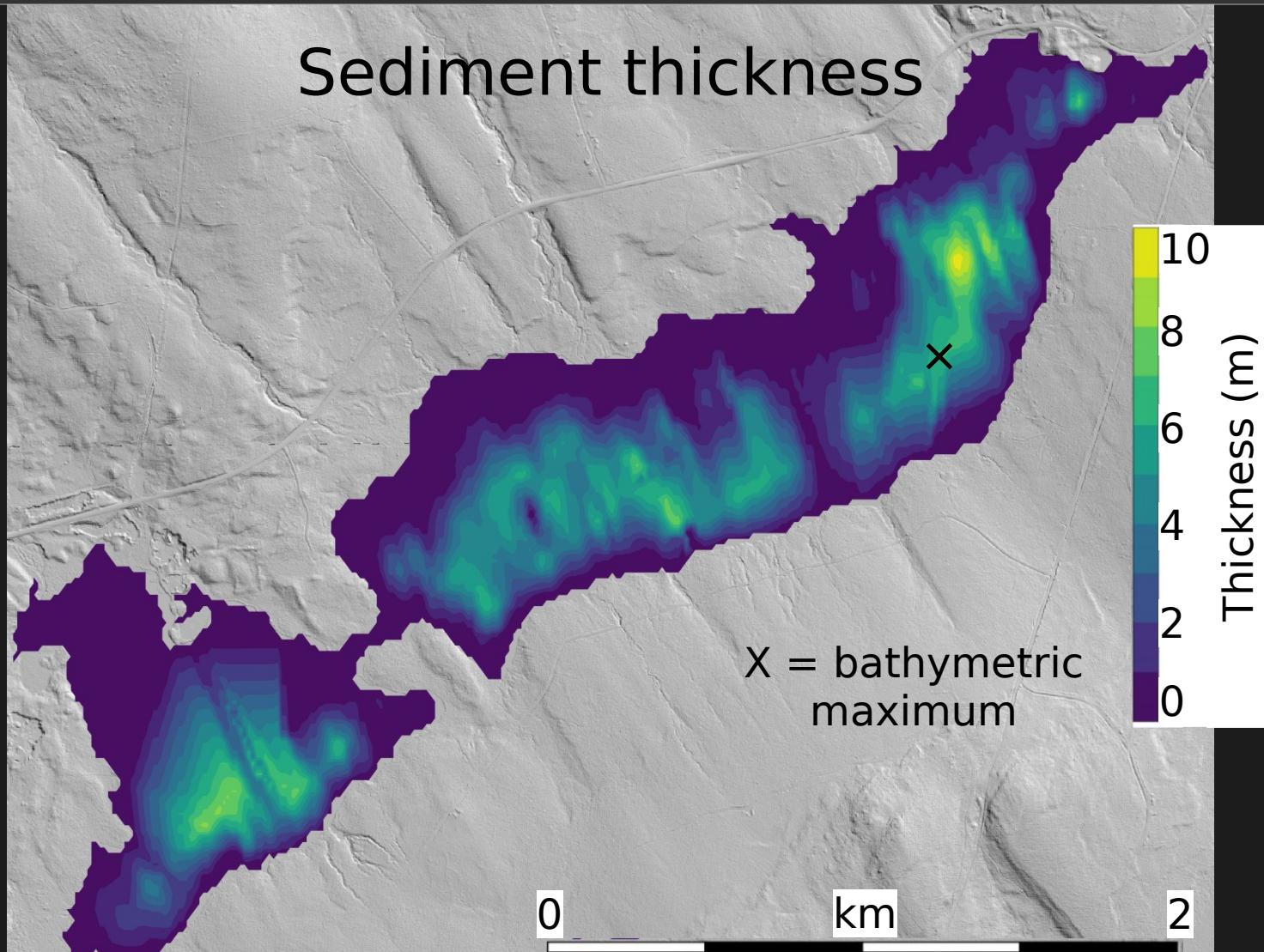


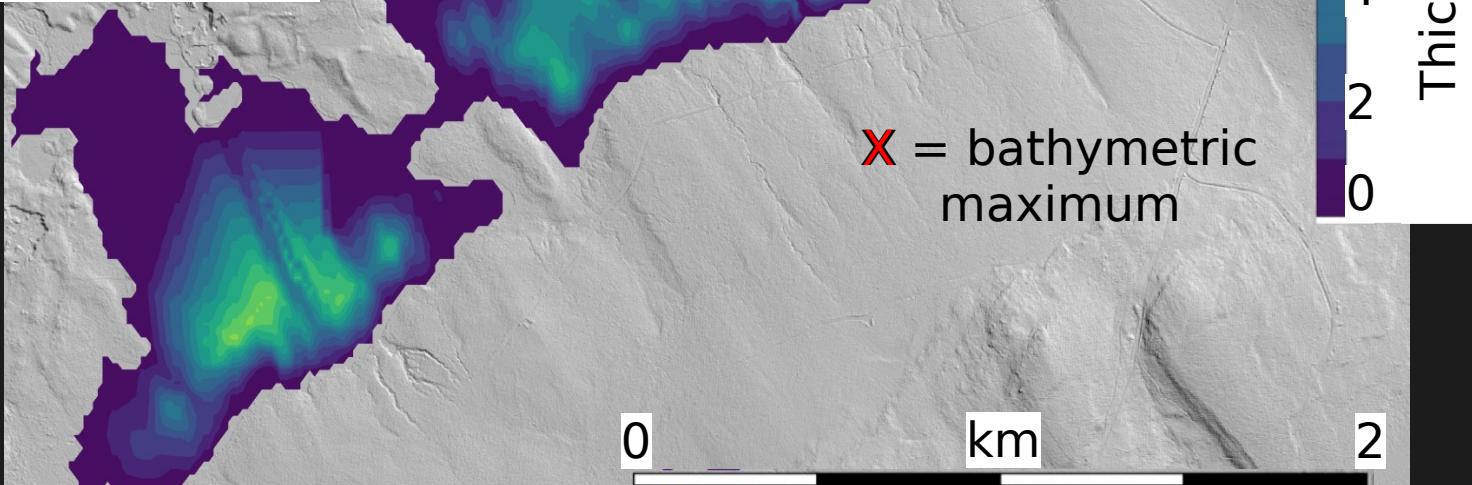
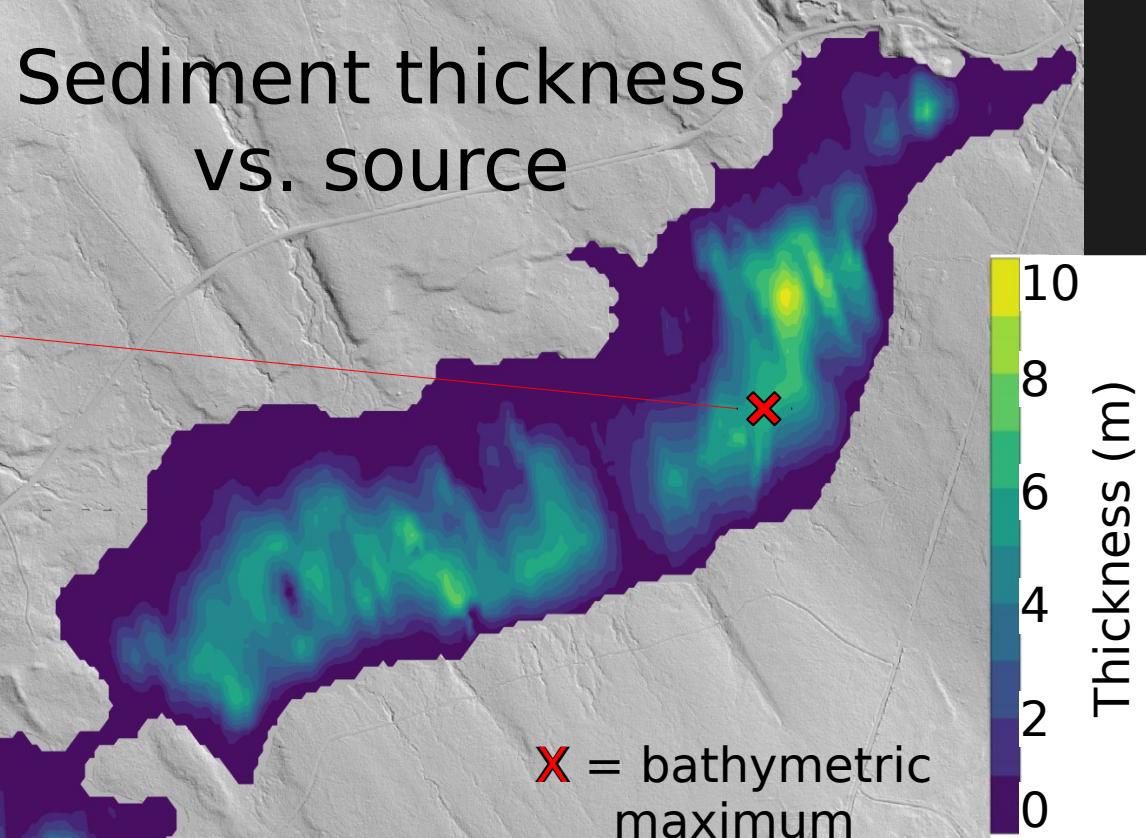
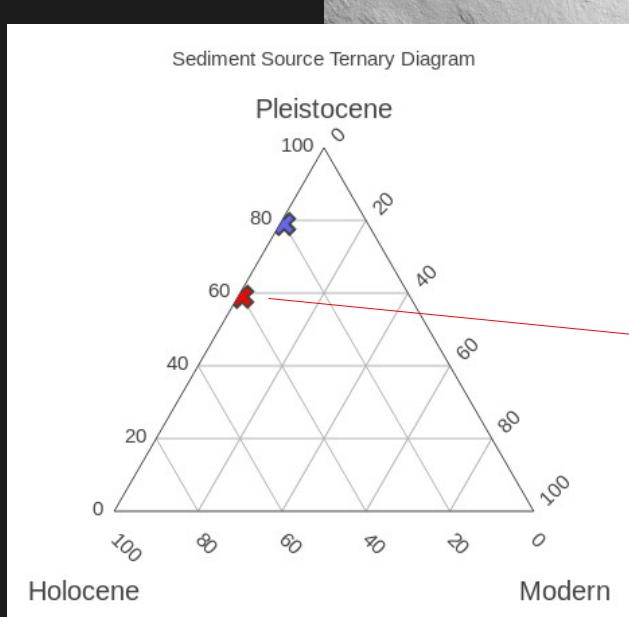
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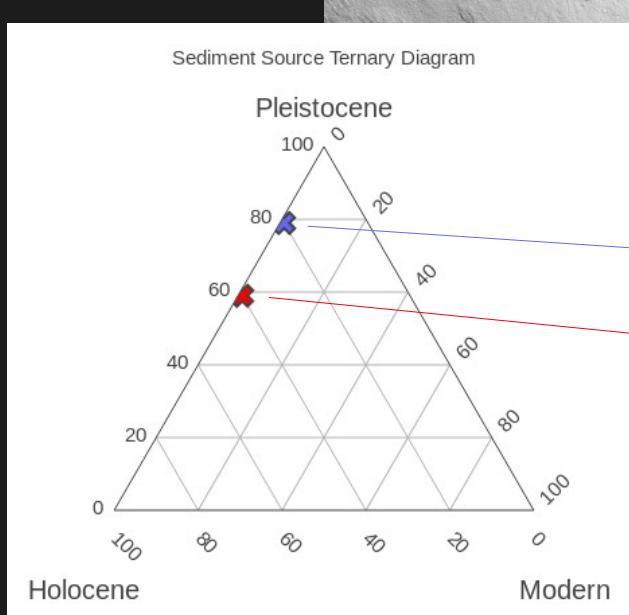


# Typical Kingsbury Pond radar transect

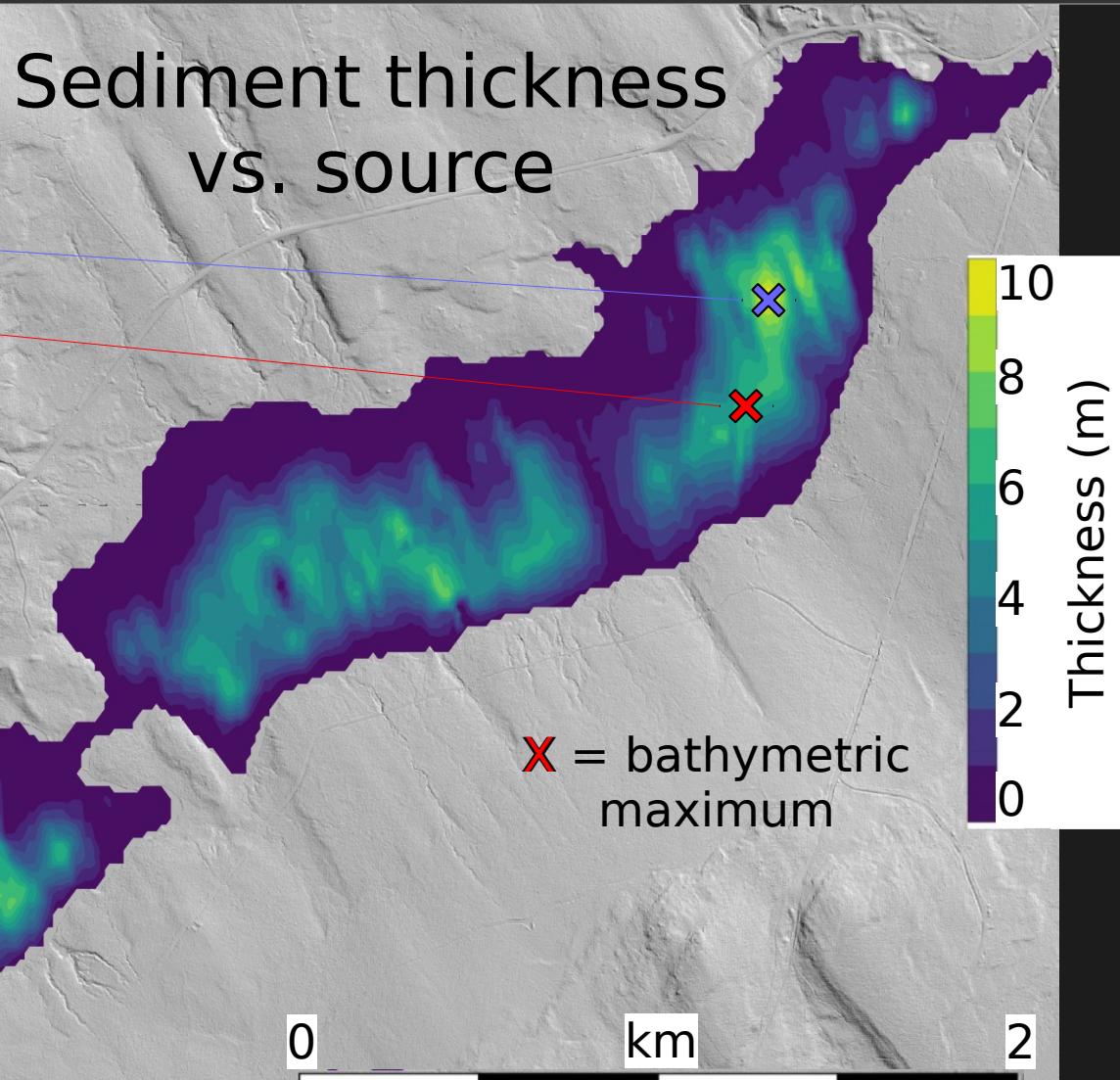


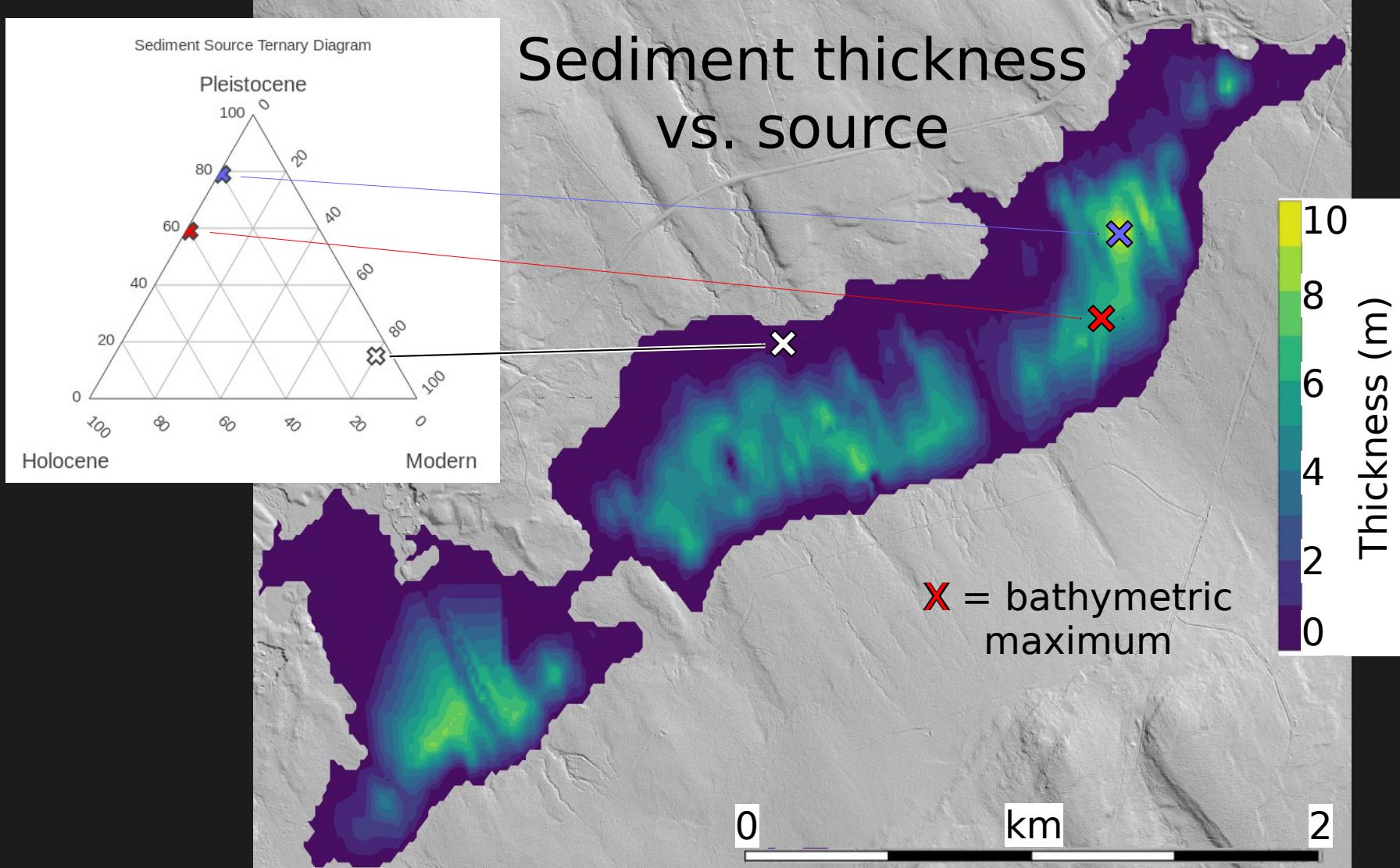


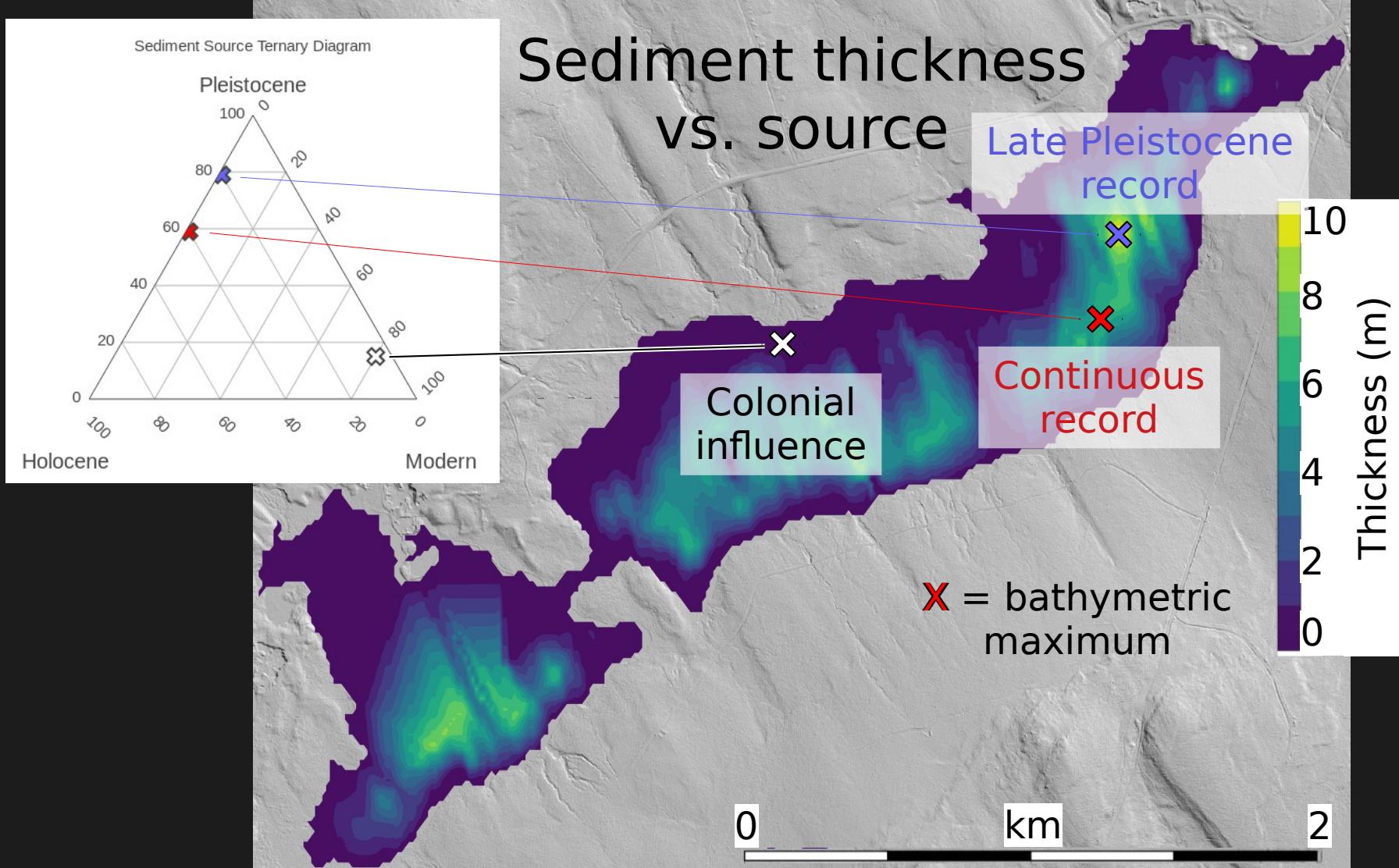




# Sediment thickness vs. source







# Summary

- Complexity has traditionally been a challenge in the lake coring community
- GPR can help establish:
  - Where to place core sites
  - Stratigraphic context for sites

# Coring



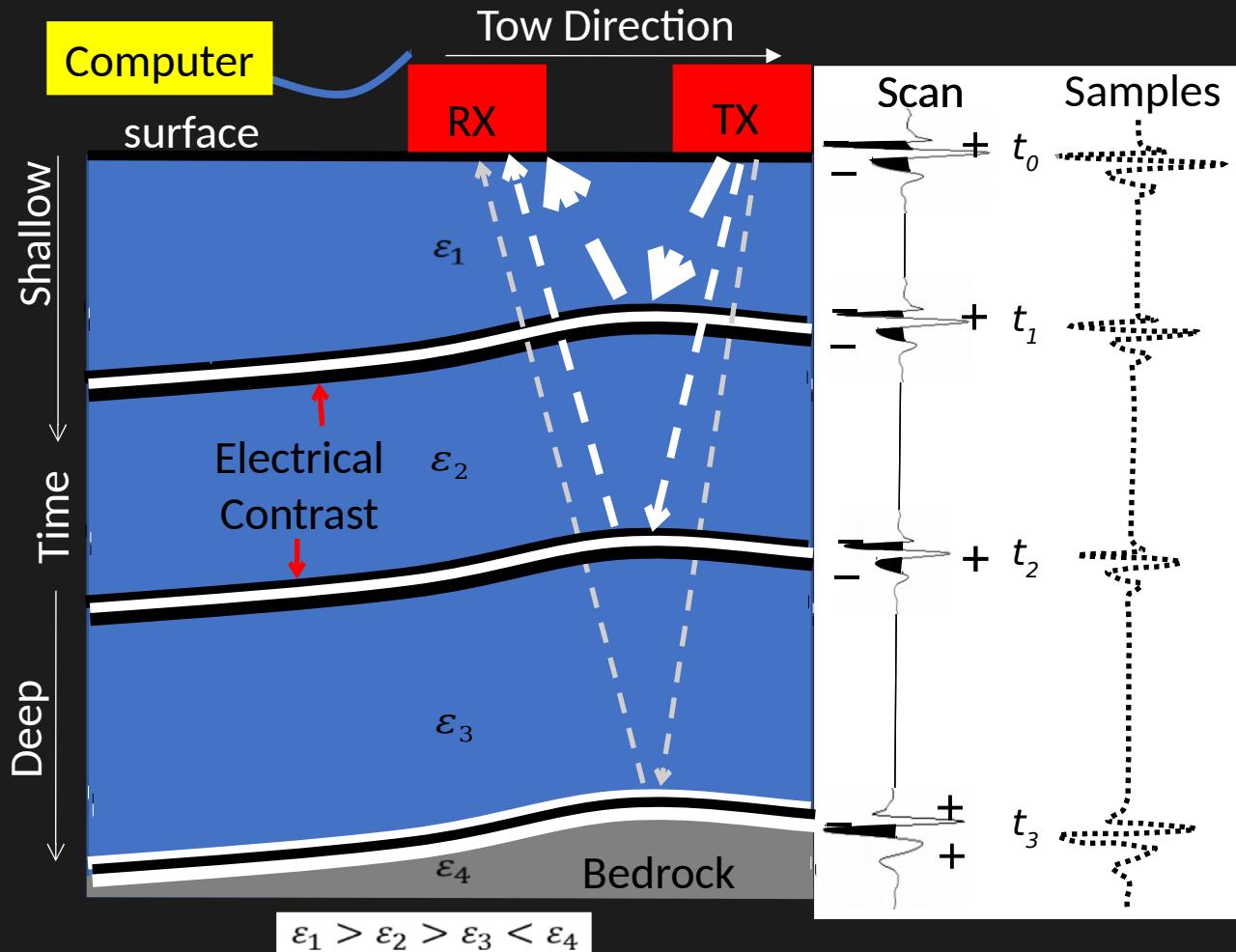
## Acknowledgments (Questions?)

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- Colby College Geology Department (coring & analysis)
- University of Maine Physics Department (analysis)
- George Jacobsen (advice)
- Mariama Dryak, Kate Hruby, Aaron Chesler (advice & moral support)

# References

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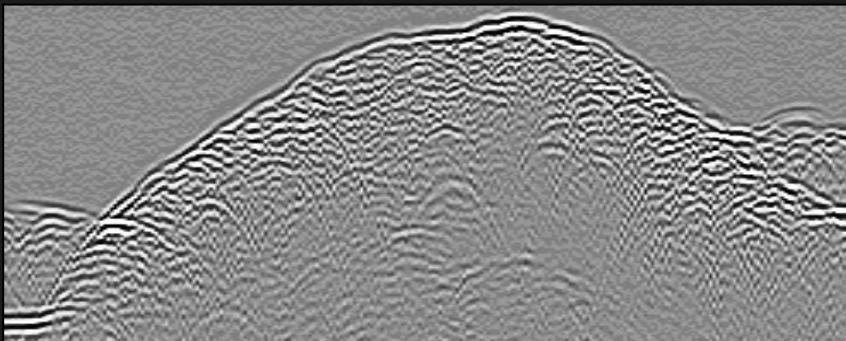
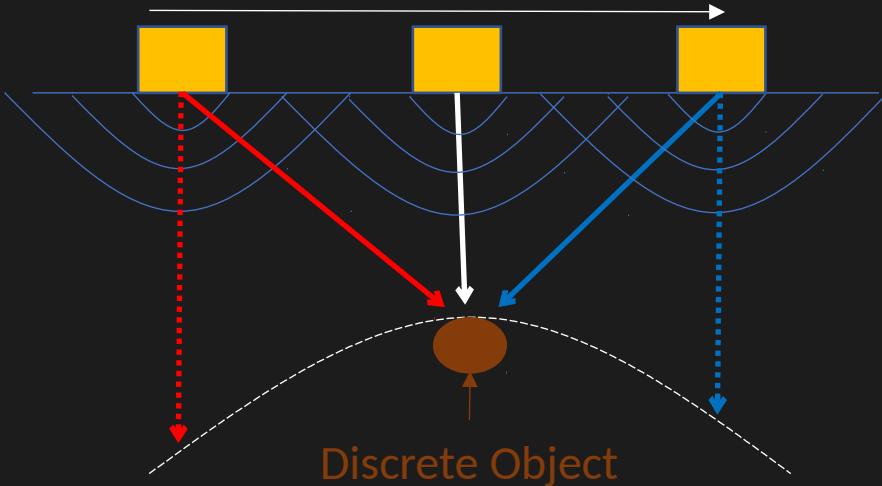
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# Ground-Penetrating Radar (GPR) or Impulse Radar

GPR Direction of Travel



## Key Points:

- Hyperbolas
- Off-Axis Reflections

