

# Waves

Ian Beckley  
University of Wisconsin-Madison

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## Descriptors

### What is a Wave?

- Quasi-periodic
- Transfer energy and momentum at large distances without transfer of fluid particles

### Free/Forced

- Free waves are resonant normal modes, excited by weak random motions in the atmosphere
- e.g., Hirota and Hirooka (1984), '5-Day Wave'
- Forced waves require continuous forcing at compatible space and time scales
- e.g., Tides

### Internal/External

- Internal: wave amplitude maximizes in the interior
- Sinusoidal form for phase variation  $\sim e^{im_r z}$  which indicates sinusoidal variation in the vertical
- e.g., synoptic Rossby waves, internal gravity waves
- External: wave amplitude maximizes at the edge ('evanescent waves'  $\sim$  amplitude decays in space)
- wave energy decreases away from the boundary  $\sim e^{-m_i z}$
- e.g., surface water waves (wave impact dampens away from the density gradient)

## Stationary/Travelling

- Stationary  $c_r = 0$
- e.g. 'standing wave' (fixed node seiche, guitar string, etc)
- Travelling  $c_r \neq 0$

## Steady/Transient

- Steady  $c_i = 0$ , fixed amplitude
- Transient  $c_i \neq 0$ , amplitude varies in time (growth or decay)

## Linear/Non-linear

- Linear (infinitesimal, theoretical)
- Nonlinear (finite amplitude, observable)

Take the Eulerian and advective form of the Lagrangian time rate of change of  $u$ . The flow is 'self-steepening' in that over faster flow will 'catch up to' slower flow and, subsequently, break. This process is fundamental to chaos theory. Note, however, that if  $\bar{u}$  is constant to first order, the total derivative is linearized ( $u = \bar{u} + u'$ ).

$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} = 0$  nonlinear advection, self-steepening with a tendency to break

If  $\frac{|u'|}{|\bar{u}|} \ll 1$  where  $u = \bar{u} + u'$  then

$\frac{\partial u'}{\partial t} + \bar{u} \frac{\partial u'}{\partial x} = 0$  advection linear in eddy amplitude  $u'$ , linearized operator

## Dispersive/Non-dispersive

- Dispersive: propagation speed depends on wavelength
- e.g., Rossby waves ( $c = -\frac{\beta}{k^2} \propto L_x^2$ )
- Non-dispersive: all waves travel at the same speed, regardless of wavelength
- e.g., linear shallow water waves,  $c = \sqrt{gh}$ ; linear sound waves,  $c = \sqrt{\gamma RT}$

## Dispersion relation

- A *dispersion relation* relates the time and space scales of a given wave type to physical parameters
- Expressed as the functional dependence of frequency  $\omega$  on size and physics ( $\omega = \omega(\vec{k}, \text{physics})$ )
- If  $\omega(k)$ , waves travel at different speeds
- Can also show how energy and wave crests move relative to one another
- A *critical surface* is defined where  $c = \bar{u}$ . This determines where wave energy emerges and cannot pass
- *phase speed* ( $c = \frac{\omega}{k}$ ) indicates how fast a wave crest or trough moves
- *group speed* ( $G = \frac{\partial \omega}{\partial k}$ ) indicates how fast energy is moving

## Wave Definitions

### Phase

*phase* is defined as  $\theta = kx + ly + mz - \omega t$ , which is the definition of a plane in 3D space with a temporal dependence (wave on a given phase can travel relative to the three axes). Phase increases in the direction of propagation...or phase decreases in time as a wave crest passes (hence negative sign on  $\omega t$  term).

### Disturbance

Take a generalized disturbance:  $\psi = \text{Re} [C e^{i\theta}] = \text{Re} [(C_r + iC_i)(\cos \theta + i \sin \theta)] = C_r \cos \theta - C_i \sin \theta = A \sin \theta + B \cos \theta$ .

### Wavenumber Vector

$$\vec{k} = \nabla \theta = \hat{i} \frac{\partial \theta}{\partial x} + \hat{j} \frac{\partial \theta}{\partial y} + \hat{k} \frac{\partial \theta}{\partial z} \quad (1)$$

Recall over one wavelength  $\partial \theta = 2\pi$ , thus,  $k = \frac{2\pi}{\lambda}$  (same for meridional and vertical wave numbers)

$$|\vec{k}| = \sqrt{k^2 + l^2 + m^2}$$

Since  $\vec{k} = \nabla\theta$ ,  $k \perp \theta$  pointing towards increasing  $\theta$ .

$\therefore$  phase decreases with time at a point

## Frequency

$\omega = -\frac{\partial\theta}{\partial t}$  'how fast  $\theta$  varies in time. When  $\omega > 0$  the wave is travelling to the east. When  $\omega < 0$  the wave is travelling to the west. Though, we don't really talk about the sign of wave periodicity.

In  $\tau$  (one period),  $\partial\theta = 2\pi$ , thus,  $|\omega| = \frac{2\pi}{\tau}$

## Conservation of Wave Crests

The local rate of change of the wavenumber vector and the frequency gradient sum to zero, or,

$$\begin{aligned} \frac{\partial \vec{k}}{\partial t} + \nabla\omega &= 0 \\ \frac{\partial}{\partial t}\left(\frac{\partial\theta}{\partial x}\right) + \frac{\partial}{\partial x}\left(-\frac{\partial\theta}{\partial t}\right) &= 0 \end{aligned}$$

This is easily interpreted through an example. If  $\omega$  is decreasing to the east ( $\nabla\omega > 0$ ) then  $\vec{k}$  must decrease in time (longer waves reaching the point).

## Phase Velocity

This is the rate that a point on a phase surface travels in the direction of  $\vec{k}$ .

$$\vec{c}_p = \frac{\omega}{|\vec{k}|} \frac{\vec{k}}{|\vec{k}|} = \frac{\omega}{k^2 + l^2 + m^2} (k\hat{i} + l\hat{j} + m\hat{k})$$

$$\vec{c}_p \perp \theta \text{ and } \vec{c}_p \parallel \vec{k}$$