

Stratosphere to troposphere ozone event characterisation and distribution over Melbourne, Macquarie Island, and Davis.

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Abstract

We develop a quantitative method for determining Stratosphere to Troposphere Transport events (STTs) and a minimum bound for this transported ozone quantity using ozonesondes over Melbourne, Macquarie Island, and Davis.

1 Introduction

1.1 Background

Tropospheric ozone is important for both air quality and climate change. Over the industrial period, tropospheric ozone, which is the third most potent greenhouse gas, has been estimated to exert a radiative forcing equivalent to a quarter of the CO₂ forcing. Ozone is present in the troposphere due to a variety of dynamical and photochemical processes, including downward transport from the ozone-rich stratosphere and anthropogenic pollution. The primary sources of tropospheric ozone are chemical creation and stratospheric input, estimated using a model ensemble to be 5100 ± 600 Tg/yr and 550 ± 170 Tg/yr, respectively [Stevenson et al., 2006]. The primary sinks are chemical destruction and dry deposition, estimated to be 4700 ± 700 Tg/yr and 1000 ± 200 Tg/yr, respectively [Stevenson et al., 2006].

Ozone-rich air mixes irreversibly down from the stratosphere during meteorologically conducive conditions [Sprenger et al., 2003, Mihalikova et al., 2012]X, these are referred to as Stratosphere - Troposphere Transport events (STTs). In

need some refs
in here

anything more recent?

the extra-tropics, STTs most commonly occur during synoptic-scale tropopause folds [Sprenger et al., 2003] and are characterised by tongues of high Potential Vorticity (PV) air descending to low altitudes. These tongues become elongated and filaments separate from the tongue which mix irreversibly into tropospheric air. Stratospheric ozone brought deeper (lower) into the troposphere is more likely to affect the surface ozone budget and tropospheric chemistry [Zanis et al., 2003, Zhang et al., 2014].

This belongs in your thesis
chapter but is too
basic for the audience
of your paper.

While the amount of tropospheric ozone is small compared with that found in the stratosphere, it is an important constituent. A high correlation is found between lower stratospheric and tropospheric ozone [Terao et al., 2008] with the highest STT associated with the jet-streams over the oceans in winter. Irreversible STT of ozone is important for explaining tropospheric ozone variability [Tang and Prather, 2011].

| some of this would fit better in previous paragraph

refs?

needs context! %
change? or 1965
levels?

In a future climate, a warmer, wetter troposphere will change the chemical processing of ozone. Dynamical processes such as STT, boundary layer ventilation and convection changes will alter tropospheric ozone distributions. Hegglin and Shepherd [2009] estimate that climate change will lead to increased STT of the order of 30 (121) Tg yr⁻¹ relative to 1965 in the southern (northern) hemisphere due to an acceleration in the Brewer Dobson circulation.

Using several years of ozonesonde flights from three locations spanning the latitudes of the Southern Ocean, we will characterise the seasonal cycle of STT events and determine their contribution to the total amount of tropospheric ozone. We will examine the depth of the intrusions and using case studies to relate these STT to meteorological events.

X2 Instruments and Data

Ozonesondes are weather balloons with an attached instruments which measure ozone concentrations roughly every 100m up to around 30km. These ozonesondes provide a high-vertical resolution profile of ozone. (TODO precision of ozone recordings?)

Ozonesondes are launched approximately weekly from Melbourne (145°E, 38°S), Macquarie Island (159°E, 55°S) and Davis (78°E, 69°S). For this study, we use the data collected from 2004-2013 for Melbourne and Macquarie, and 2006-2013 for Davis. While ozonesondes were launched prior to 2004 at the two northern sites, they were less frequent than after 2004, the analysis focuses on results from these more recent times. More frequent ozonesonde launches occur at Davis during the spring ozone hole season than at other times of the year [Alexander et al., 2013].

Ozonesondes provide much higher vertical resolution profiles of ozone than that available from reanalyses products. However, one data point per week from an ozonesonde flight is too low to be in itself useful to diagnose the evolution of STT exchange over time-scales associated with normal synoptic scale weather patterns present in the extra-tropics. The ozonesonde data are supplemented with the ERA-Interim reanalysis [Dee et al., 2011] to enable construction of an STT exchange climatology.

Paper introductions
don't typically have
sub-sections;
make this section
2.1?

This sentence is
irrelevant (and
misleading!) if you
aren't using O3
from the reanalysis.
I suggest you wait
until you use ERA
to introduce it.

Doesn't add anything,
irrelevant

2 Methods

2.1 Characterisation of STT events and associated fluxes

STT events are characterised in the ozonesonde vertical profiles of ozone as altitudes in the troposphere where the ozone mixing ratio exceeds a specified threshold. Usually stratospheric ozone ~~mixes~~ ^{intrudes} down into the troposphere in a synoptic-scale tongue of air. The vertical ozone profile observed by the ozonesonde depends upon the time in this cycle that it is observed [Sprenger et al., 2003]. As such, the altitude of the tropospheric ozone peak due to an STT event, and the amplitude of the event above the background tropospheric ozone profile, vary in space and time.

Two definitions of the tropopause height are calculated: the standard lapse rate tropopause [WMO, 1957] and the ozone tropopause [Bethan et al., 1996]. At Davis, the ozone tropopause definition is modified for polar sites, following Tomikawa et al. [2009], Alexander et al. [2013]. While the ozone tropopause can be less robust during stratosphere-troposphere exchange, it performs better than the lapse rate tropopause at polar latitudes in winter and near jet streams in the lower stratosphere [Bethan et al., 1996]. The lower of these two tropopause altitudes is referred to as the tropopause for this study. This choice of the lowest altitude of the tropopause avoids occasional unrealistically high tropopause heights due to perturbed ozone or temperature measurements. The monthly mean tropopause altitudes at each location are shown in Figure 1, along with the subset of altitudes from profiles for which an STT event was determined. The seasonal cycle in tropopause altitude at Melbourne is clearly apparent, as is the decreasing tropopause altitude poleward. Seasonally averaged ozone as recorded over the three stations (Figure 2) shows increased ozone extending down through the stratosphere during the peak STT months over Melbourne. It is worth noting that tropopause altitudes at Davis may exceed 11 km altitude under certain synoptic conditions [Alexander et al., 2013]: the relation of tropopause altitude with individual STT events will be investigated in detail below.

Figure 1 shows...
Why are you talking about tropopause height now?
Isn't there more to say from Fig 2?
To identify STT events, as
We exclude from our analysis perturbations
that...
in ozone profile new records?
backward phrasing? be explicit
for an event to qualify

→ this should come w/ Fig 1 not after Fig 2

The vertical profiles of ozone volume mixing ratio are linearly interpolated to a regular grid with 20m resolution up to 14km altitude and are then bandpass filtered so as to retain perturbations which have vertical scales between 0.5km and 5km. The choice of band limits is set empirically, but we note that to define an STT event, a clear increase above the background ozone level is needed, and a vertical limit of ~ 5 km removes seasonal-scale effects. The ozone perturbation profile is analysed at altitudes from 4 km above the surface to avoid surface pollution events, and 1 km below the tropopause to avoid the sharp transition to stratospheric air producing spurious false positives. Perturbations above the 99th percentile (locally) of all ozone levels are initially classified as STT events.

In order to remove unclear 'near tropopause' anomalies, we remove events where the gradient between the maximum ozone peak and the ozone at 1 km below the tropopause is greater than -20 ppbv km^{-1} and simultaneously require that the perturbation profile does not drop below zero between the event peak

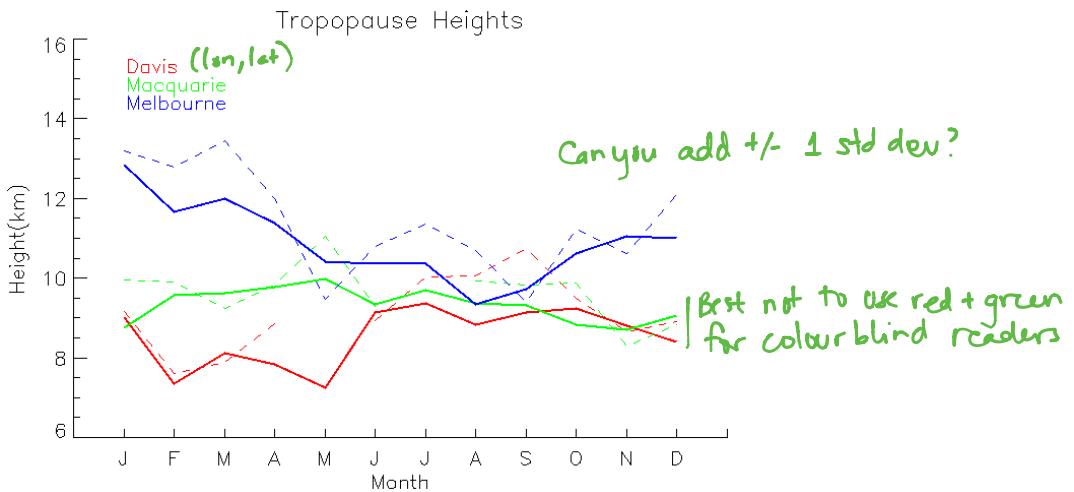


Figure 1: Monthly mean tropopause altitudes (minimum of lapse-rate and ozone defined tropopauses). Dashed lines show 'event only' seasonal tropopause altitudes.

↳ 3 sites

↳ define/explain in figure caption

and the tropopause. The addition of these filters removes several events, each with an ozone peak which could not be definitively said to be separated from the stratosphere.

filter

the

~what %?

We conservatively estimate To provide a conservative estimate of ozone flux into the troposphere for each event, the ozone concentration is integrated vertically over the interval for which an STT event is identified. An example of an ozone profile is illustrated in Figure 3 and indicates how the algorithm detects an STT event, defines the event boundaries, and calculates the ozone flux.

associated with altitude range

shows for an example
Ozone profile

from since?
emission?
point of measurement?
too basic here

rebs? | southern
Australian fires could
have equiv. influence
if from mid-lats -
rephrase

2.2 Removing STT events related to smoke plumes from biomass burning

Ozone production due to fire smoke plumes is complex and affected by photochemistry, fuel nitrogen load, and atmospheric plume interactions both during transport and at the plume's destination. Ozone precursors include nitrogen oxides ($NO_x = NO + NO_2$) and non methane volatile organic compounds (NMVOCs). Large biomass burning (BB) events emit substantial ozone precursors, some of which are capable of being transported far from their origins. Peroxyacetyl Nitrate (PAN) is a reservoir of NO_x which can lead to enhanced ozone far from the source of a fire [Jaffe and Wigder, 2012].

Biomass burning influence in the southern hemisphere comes mostly from South Africa and South America, however Australian and Indonesian fires can also influence the ozonesonde release sites. Transported BB plumes influence the southern mid latitudes generally between July and December [Pak et al., 2003]. Ozone production due to fire smoke plumes is complicated and dependent

Need a transition here. Why are we talking about fire? chemistry? dilution?
This paragraph is a bit rambling. Need to be clear that some BB (but not all) leads to O_3 prod which could affect your stats.

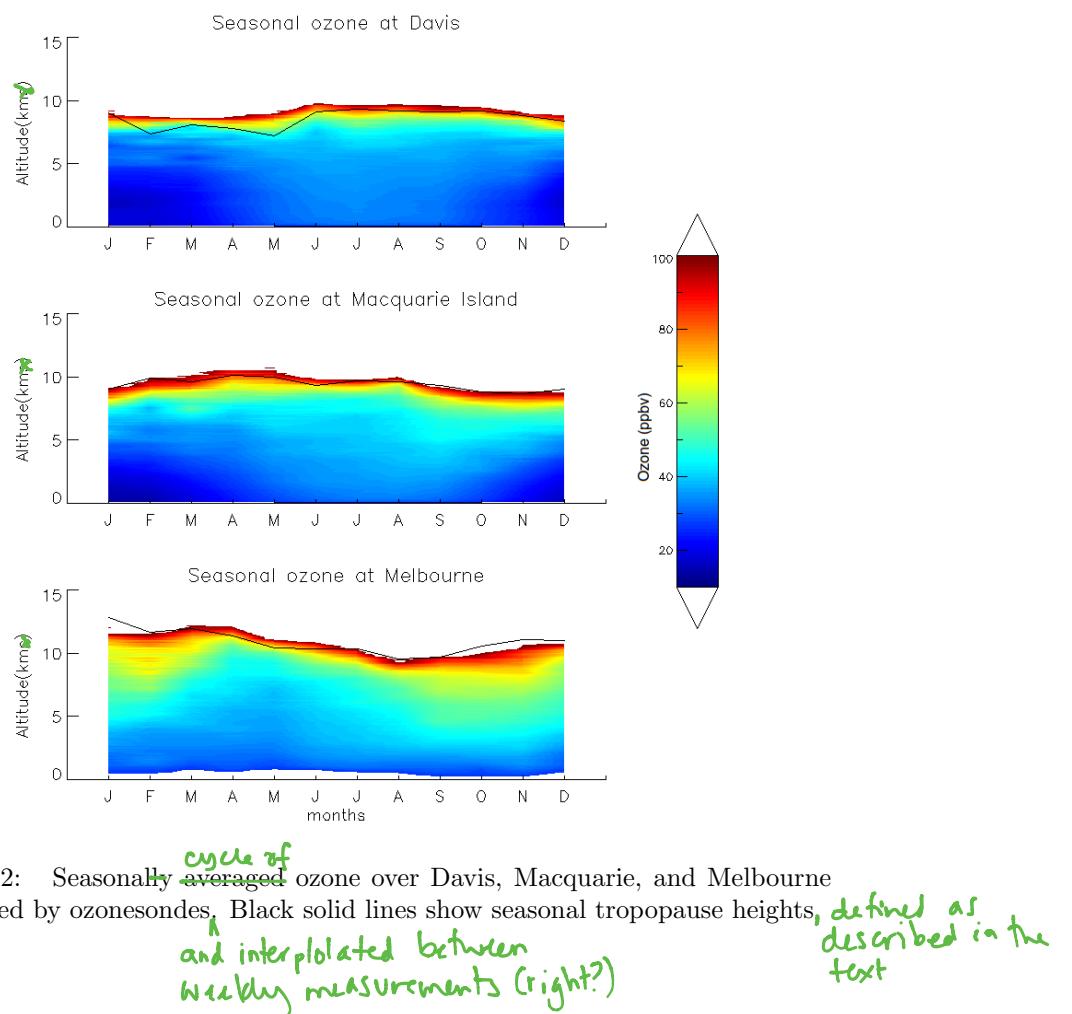
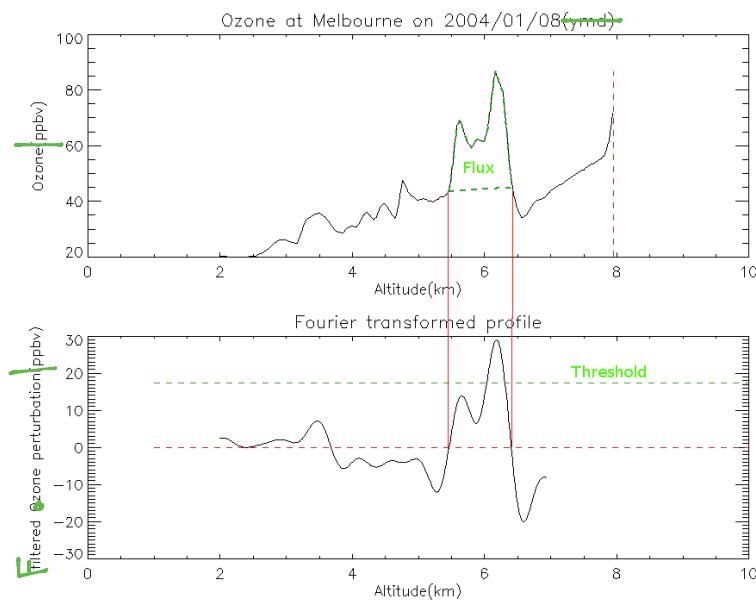


Figure 2: Seasonally averaged ozone over Davis, Macquarie, and Melbourne measured by ozonesondes. Black solid lines show seasonal tropopause heights, *cycle of and interpolated between weekly measurements (right?)*, defined as described in the text



Rotate to have alt as y-axis

change
Red/green

An example illustrating methods used for STI identification & flux estimation.

Figure 3: (a) An ozone profile between 2km altitude and the tropopause (indicated by the dashed vertical line). The flux area shows the estimate of stratospheric impact on tropospheric ozone. (b) The 99th percentile of filtered ozone perturbations (green dashed line) and the technique for determining the vertical extent of the event (red dashed and solid lines).

at Melbourne on
8 Jan 2004

D band pass filtered
O₃ profile from (a).
Coloured lines show

Transition! e.g.
"Here we identify
transported smoke
plumes using carbon
monoxide (CO)."
We use data
from

give w/e address
product / version
e.g. v6, level 3
data

Fig. 4a shows Africa

the example in
just work
into previous
sentence.

repetitive; combine
these two paragraphs.

on many chemical and meteorological factors. Due to this complication and the possible influence of smoke plumes on ozone concentrations, we exclude from analysis any dates where this influence is likely.

→ CO has a long enough lifetime to be an effective transport tracer. The primary source of atmospheric enhancement of CO is fires, making CO a good indicator of fire plumes. Using high CO levels as a proxy in order to determine where fire smoke plumes exist is a well established method (eg: Edwards [2003], Sinha et al. [2004], Edwards et al. [2006], Mari et al. [2008]) and is used here.

The AIRS (Atmospheric Infrared Sounder) instrument on board the AQUA satellite records column CO. In this work a visual inspection of AIRS' vertical columns of CO (provided by NASA [AIR, 2013]) over the southern hemisphere is used to exclude possible foreign smoke plume influence on the ozone profile at our three sites. Whenever high (approximately 2×10^{18} molecules cm^{-2}) CO concentrations coincide with sonde detected ozone events it's possible that the tropospheric ozone spike could be due to transported ozone or ozone precursors. All occasions where these coincidences are detected (through visual inspection) are removed.

I would say something like 'We diagnose smoke plumes when...
and remove these days from our analysis.' Also specify you are only looking at high CO at your sites.

Comment on timing of these relative to expected SH BB.

3 Results Just make these each their own section

3.1 Case Studies of Synoptic Conditions during STT events

We examine two STT case studies in detail to illustrate the synoptic scale conditions in which they can occur above Melbourne. Fig. 5 shows... for both events.

A cut-off low pressure system passed over Melbourne on 3 February 2005 (Figure 5b). The ozonesonde profile indicated low lapse-rate and ozonesonde tropopause (both > 450 hPa, see Figure 5a). An ozone intrusion into the troposphere is identified by our detection algorithm at ~ 520 hPa.

STT events also occur during frontal passages, an example of which is illustrated in Figure 5d over south-eastern Australia. The tropopause heights are much higher at this time and an ozone intrusion is identified centred around 200 hPa. Note the separation between this intrusion and the ozone tropopause (marked by the green dashed line), indicating the start of the stratosphere above Melbourne. During the frontal passage, stratospheric air descends and streamers of ozone-rich air likely break off and mix into the troposphere [Sprenger et al.,

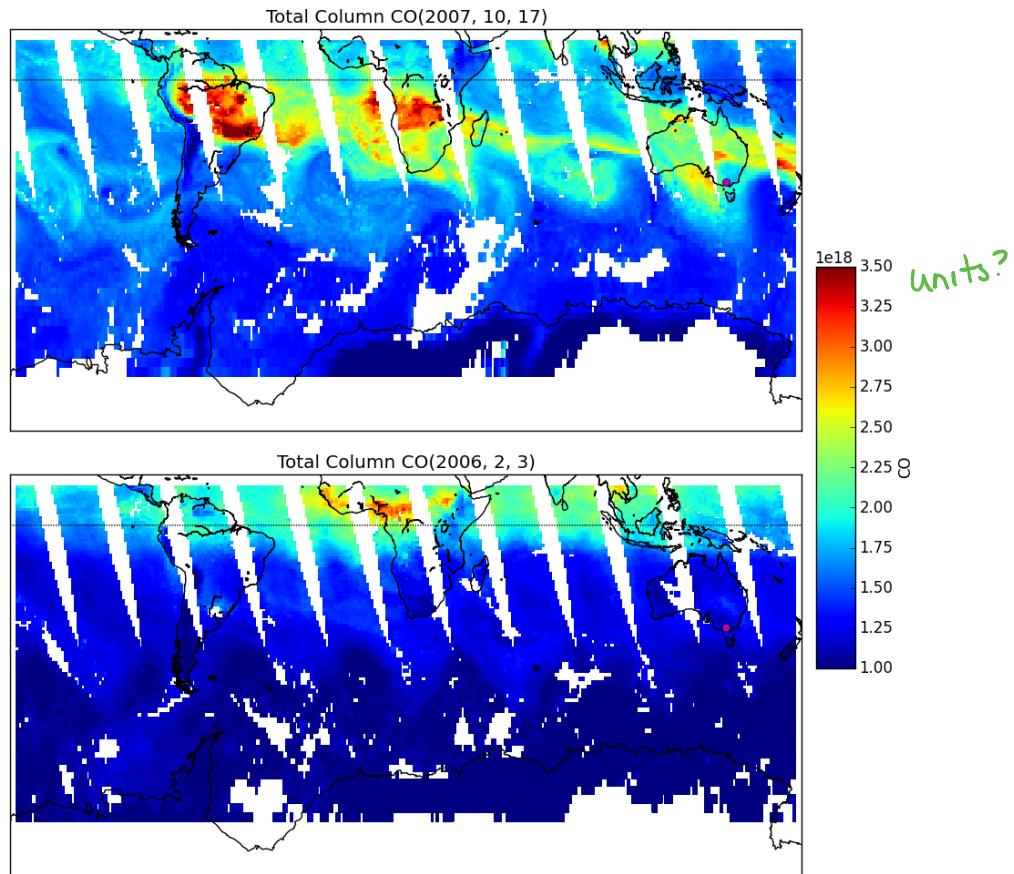


Figure 4: AIRS total column CO image showing two days separate days of swathes. The top panel (17 October 2007) shows an example of an excluded ozone event above Melbourne (marked by the red dot) which could have been caused by a transported biomass burning plume on October 17th, 2007. The bottom panel (3 February 2006) shows an example of a non-excluded ozone event. day when Melbourne was likely not influenced by transported ozone and was retained for our analysis

Need to fix resolution of these (unrealable)
I suggest post processing in inkscape
(save as postscript first, use !p.font=0 in (DL))

highlight your event on
these plots. Also calculate
flux associated with
them.

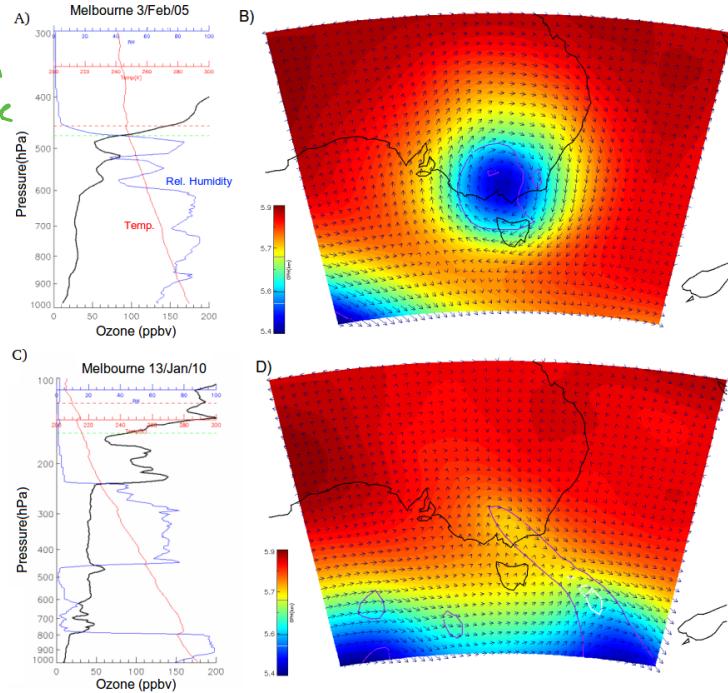


Figure 5: Vertical profiles show ozone ppbv (black line), relative humidity (blue line), and temperature (red line) for (a) 3 February 2005 and (c) 13 January 2010. Synoptic weather maps show the 500 hPa pressure level taken from the ERA-Interim reanalysis on (b) 3 February 2005 and (d) 13 January 2010. Vectors show wind direction and speed while the colour indicates the geopotential height. Also visible are the line contours of potential vorticity units, 1 PVU in purple and 2 PVU (often used to determine dynamical tropopause height) in white.

2003]. More detail again. Relate to more than one paper!

The relative humidity profiles are anticorrelated with ozone in the upper troposphere for these events, indicating again the stratospheric origin of the ozone-rich air mass. Some of this stratospheric air gets mixed into the troposphere, with one ozonesonde column showing an intrusion at around 200 hPa, with dry ozone rich air peaking below the tropopause (Figure 5c).

Does this say anything last paragraph didn't?

4 3.2 STT event Climatologies

The seasonal cycles and event climatologies of the STT events for each of the three locations are presented in Figure 6 to Figure 8. There is an annual cycle in the occurrence frequency of STT events (with a summertime peak) ↑

are shown in Fig 6
for Melbourne, Fig 7
for ..., Fig 8 for ...

Is there any difference in number of sondes launched per month @ each site? If so, needs to be commented on

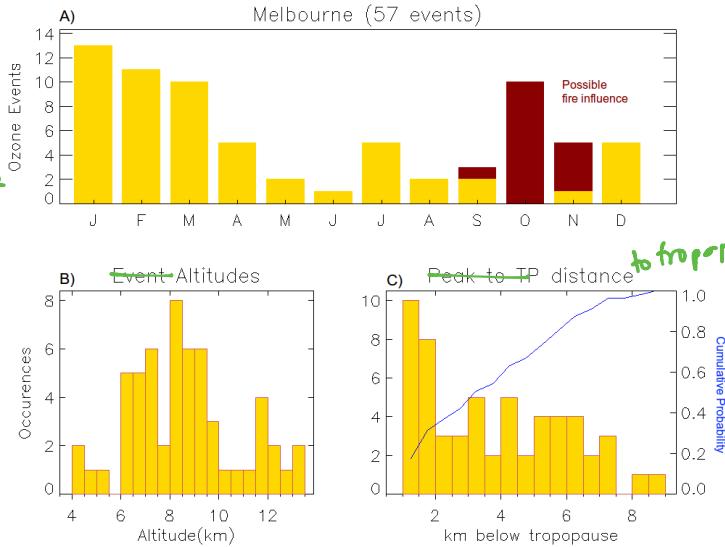


Figure 6: The climatology of STT events at Melbourne (A) Events sorted by month from the entire Melbourne ozonesonde dataset. The events filtered out as possibly smoke plume influenced are indicated in red. (B) The occurrence distribution of the ozone peak altitude, and (C) The distance between the ozone peak and the tropopause (bars) and the cumulative probability function of these distances (blue line).

In (b) and (c) altitude is at the event peak.

Melbourne and Macquarie Island. However, the occurrence frequency of STT events above Davis is relatively constant throughout the year.

The majority of events occur within 3 km of the tropopause at both Melbourne and Macquarie Island, and within 2 km of the tropopause at Davis. STT event altitudes most commonly occur at 6 – 10 km above Melbourne and below 8 km at Davis but are distributed more evenly in altitude at Macquarie Island.

Using 458 ozonesonde profiles over Melbourne, 72 ozone events are detected, of which 15 are discarded as possibly caused by transported fire smoke plumes. Over Macquarie island 48 events are detected from 380 ozonesondes of which 8 are discarded due to possible smoke influence. We also include on Figure 6 to Figure 8 the events which have possible fire influence. These events are concentrated in Spring at melbourne and Macquarie Island. For both of these sites, the STT events which are unlikely to be fire-related occur mostly in Summer and mostly during storms which can increase convection and upper tropospheric turbulence. For Davis, 36 events are detected from 240 ozonesonde profiles. There is no clear seasonal cycle over Davis.

Move to section or its fitting.

expand this! →

to tropopause
Shouldn't there be red fire influence in b/c, or have these already been removed?

binned by

Again, you need more discussion. What does it mean? How can you relate to literature?

Given the order of your discussion, I suggest you reshuffle subplots so Fig. 6 is seasonal cycle@ all 3 sites, Fig. 7 is altitudes, Fig 8. is distance to tropopause. Then discuss each in turn (as you sort of do now)

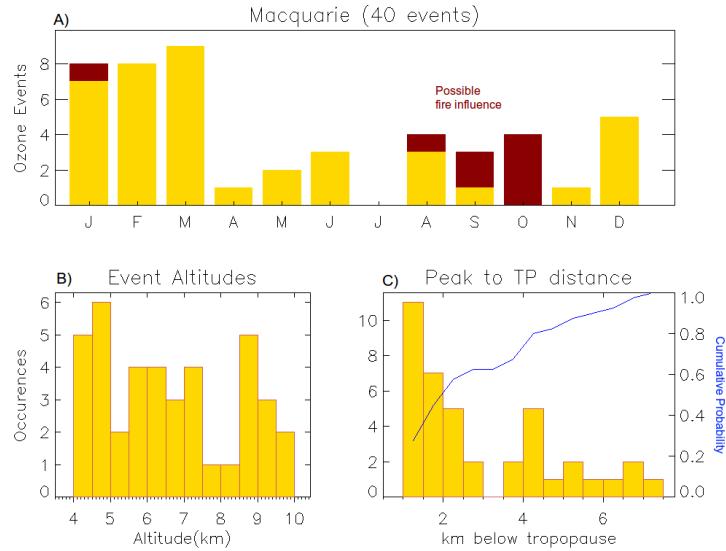


Figure 7: As for Figure 6 except showing the Macquarie Island STT events.

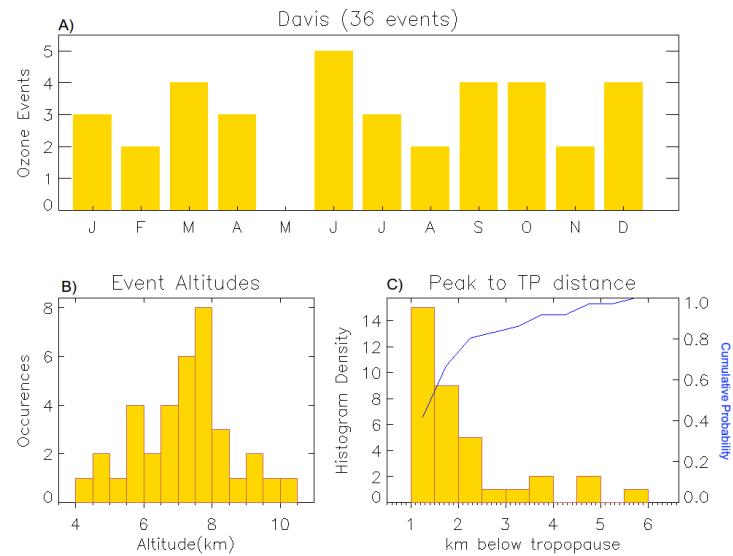


Figure 8: As for Figure 6 except showing the Davis STT events.

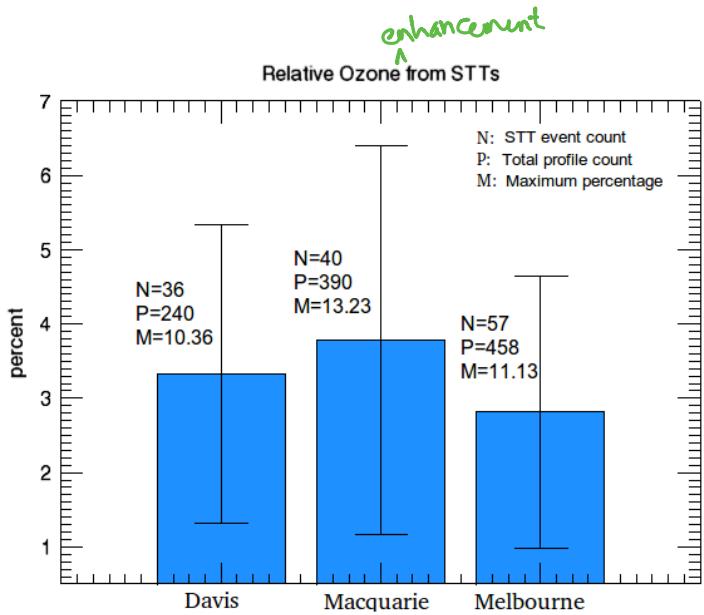


Figure 9: Fraction of total tropospheric column ozone attributed to stratospheric air intrusions during STT events. Error bars indicate one standard deviation.

4.3.3 Stratosphere to troposphere Ozone flux from STT events

Based on the integrated ozone amount associated with each STT event,

How do you calculate the total column? Is that from the stroke? Explain!

Would be nice to see a frequency distribution of flux associated w/ each event

Stratosphere to Troposphere ozone transport can potentially increase regional surface ozone levels above safe levels [Zhang et al., 2014]. Using our estimate of STT ozone flux (see section 2.1) we find a lower bound for the STT ozone flux over each of our three sites, (excluding possible fire influence) Figure 9 shows the climatological mean fraction of total tropospheric column ozone attributed to stratospheric ozone intrusions at each site, on days when an STT event occurs. These flux amounts are calculated after removal of the biomass burning events, although leaving the burning events in changes the means by less than 5%. The mean fractions of stratospheric ozone are 2–4%, although the largest fractional ozone in the tropospheric column attributed to stratospheric air exceeds 10% at all locations.

Our flux estimates are relatively insensitive to our biomass burning filter; including smoke influenced days changes the mean flux by <5%.

why is it a lower bound?

On individual days, this value can

4 Conclusion

Using ozonesonde data in the southern hemisphere can allow an overview of STT ozone transport which is independent of satellite data. Using a simple Fourier filter allows deterministic and quantitative analysis of STT ozone transport events. To BE CONTINUED! !!

To make this meaningful, I think you need at least a back of the envelope calculation of how much O₃ we are talking about if we extrapolate from your sites. Should be able to find papers w/ trop. O₃ columns, or we can grab some satellite data. If we combine w/ monthly likelihood + 3% should be able to estimate net O₃ inputs to trop. over southern ocean.

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