

**Joint stochastic theory of fluvial bedload transport and  
bed elevation changes deriving heavy-tailed sediment  
resting times**

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**Key Points:**

- We model fluvial bedload activity and local bed elevation as a two-species stochastic birth-death process.
- Resulting timescales of sediment storage by burial lie on heavy-tailed power-law distributions.
- These distributions have universal characteristics, offering a possibility of discriminating the signature of burial in tracer experiments.

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13      **Abstract**

14      A consensus has formed that fluvial bedload transports with heavy-tailed statistical dis-  
 15      tributions of resting times due to the effect of burial on its mobility, and this has key im-  
 16      plications for the diffusion of sediment and evacuation of contaminants from river chan-  
 17      nels. Owing to observational difficulties, only a handful of experiments have resolved these  
 18      heavy-tailed resting time distributions, and there have been few theoretical attempts to  
 19      build understanding, leaving open questions as to the form of the resting time distribu-  
 20      tion and its governing factors. We present a new theory which describes bedload trans-  
 21      port and bed elevation changes as a joint stochastic process and derives resting time dis-  
 22      tributions for sediment undergoing burial from these joint dynamics. Our theory pre-  
 23      dicted heavy-tailed power-law distributions of resting times for sediment undergoing burial,  
 24      with the longest resting times completely governed by the mean erosion rate across bed  
 25      aggradation/degradation cycles and its scaling with bed elevation changes. This implies  
 26      diffusion characteristics of bedload undergoing burial ultimately remain linked to the me-  
 27      chanics of sediment transport.

28      **1 Introduction**

29      The majority of the classic studies into fluvial sediment transport have attempted  
 30      to relate the bulk downstream flux of bedload to characteristics of the hydraulic forc-  
 31      ing (e.g. Meyer-Peter & Müller, 1948; Yalin, 1972), yet the relevance of this approach  
 32      to environmental problems is limited, as many contemporary issues are contingent on  
 33      the motion characteristics of individual grains rather than the average characteristics of  
 34      many grains. For example, the persistence of solid contaminants within river channels  
 35      ultimately links to the slowest moving grains, and not to the bulk flux (Malmon, Reneau,  
 36      Dunne, Katzman, & Drakos, 2005). Similar points can be made in relation to salmonid  
 37      habitat restoration (e.g. Gaeuman, Stewart, Schmandt, & Pryor, 2017) and morpholog-  
 38      ical adjustment to disturbances (Hassan & Bradley, 2017), highlighting the motions of  
 39      individual grains through river channels as an important topic for geophysics research.

40      A significant complication is that individual grains transport within a noisy envi-  
 41      ronment, with noise sources ranging from microscale fluid turbulence (Celik, Diplas, &  
 42      Dancey, 2014) and the irregular arrangement of bed surface grains (?), to channel mor-  
 43      phodynamics (Hassan & Bradley, 2017) and watershed hydrology (Phillips, Martin, &  
 44      Jerolmack, 2013). Owing to this noise, the transport characteristics of individual grains  
 45      are not deterministic (e.g. Einstein, 1937), even in the most controlled small-scale lab-  
 46      oratory experiments (e.g. Fathel, Furbish, & Schmeeckle, 2015; Heyman, Bohorquez, &  
 47      Ancey, 2016). This realization has led to random walk formulations of individual mo-  
 48      tions, where bedload grains are considered to move through an alternating series of steps  
 49      and rests, where step lengths and resting times are considered as random variables ly-  
 50      ing on statistical distributions (N. D. Bradley & Tucker, 2012; Einstein, 1937; Hassan,  
 51      Church, & Schick, 1991; Nakagawa & Tsujimoto, 1976; Yano, 1969).

52      These random-walk models describe bedload diffusion, or the spreading apart of  
 53      grains through time due to differences in their motions. The nature of bedload diffusion  
 54      is controlled by whether the probability distributions of step lengths and resting times  
 55      have light or heavy tails. A heavy-tailed step length or resting time distribution has an  
 56      exceedance distribution  $P(X > x) \sim x^{-\alpha}$  with tail parameter  $\alpha < 2$ , meaning large  
 57      values of  $x$  are relatively common, while a light-tailed distribution has  $\alpha \geq 2$ , mean-  
 58      ing large values of  $x$  are relatively rare. If both distributions have light tails, the dif-  
 59      fusion is said to be normal or Fickian, with a variance of particle positions  $\sigma_x^2$  which scales  
 60      with time  $t$  as  $\sigma_x^2 \propto t$ . However, if either of these distributions has a heavy-tail, the dif-  
 61      fusion is called anomalous, with the variance of particle position scaling as  $\sigma_x^2 \propto t^\gamma$  with  
 62       $\gamma \neq 1$ . In this expression,  $\gamma < 1$  is called sub-diffusion and  $\gamma > 1$  is called super-diffusion.  
 63      In a strongly assymmetric random walk such as bedload transport, heavy-tailed step length

distributions imply super-diffusion, while heavy-tailed resting time distributions imply either super or sub-diffusion depending on the value of the tail parameter  $\alpha$  (Weeks & Swinney, 1998; Weeks, Urbach, & Swinney, 1996).

Tracer experiments in gravel bed rivers show anomalous diffusion (D. N. Bradley, 2017; Phillips et al., 2013), light-tailed step length distributions (N. D. Bradley & Tucker, 2012; Hassan, Voepel, Schumer, Parker, & Fraccarollo, 2013), and heavy-tailed resting time distributions (D. N. Bradley, 2017; Olinde & Johnson, 2015; Pretzlav, 2016; Voepel, Schumer, & Hassan, 2013), forming a coherent experimental picture of super-diffusive bedload transport at long observation timescales (Martin, Jerolmack, & Schumer, 2012; Nikora, 2002). However, these field studies do not resolve the mechanism generating heavy-tailed resting times (e.g. D. N. Bradley, 2017), and the experimentally obtained resting time distributions are not entirely consistent with one another in form or characteristics, displaying different tail parameters  $\alpha$  and sometimes truncation (e.g. Voepel et al., 2013) or tempering back to a light-tailed distribution at large times (e.g. D. N. Bradley, 2017), which may be relics of necessarily limited observation periods (e.g. D. N. Bradley, 2017).

A predominant hypothesis is that sediment burial is a mechanism generating heavy-tailed resting times (Martin, Purohit, & Jerolmack, 2014; Voepel et al., 2013). Conceptually, when grains rest on the bed surface, material transported from upstream can deposit over top of them, burying them away from the flow and preventing their entrainment until the overlying material is removed, increasing sediment resting times and imparting a heavy tail to the distribution. Both Voepel et al. (2013) and Martin et al. (2014) created random-walk theories of local bed elevation and interpreted resting times as return periods from above in the bed elevation time-series, generating heavy-tailed distributions which are consistent with different experimental datasets, but inconsistent with one another. Voepel et al. (2013) explained the field data of Habersack (2001), deriving initially heavy-tailed distributions which temper to exponential decay at the largest resting times, while Martin et al. (2014) explained data from their own laboratory flume experiments, which are the first to directly resolve burial as the generating mechanism, deriving heavy-tailed power law distributions with no tempering and a tail parameter  $\alpha \approx 1$ . These theoretical developments are exciting, but their assumptions seem contingent on the datasets they strive to explain, meaning their generality can be called into question.

In this work, we approach the problem from a different angle. We link bed elevation changes to sediment transport in a joint stochastic model and compute resting time distributions as a consequence of the joint description. The key assumptions of our model are that (1) bedload erosion and deposition can be characterized by probabilities per unit time, or rates (e.g. Ancey, Davison, Böhm, Jodeau, & Frey, 2008; Einstein, 1950), and (2) that these rates are contingent on the local bed elevation, encoding the property that the erosion of sediment is emphasized from regions of exposure while deposition is emphasized in regions of shelter (e.g. Sawai, 1987; Wong, Parker, DeVries, Brown, & Burges, 2007). As we'll show, our theory supports heavy-tailed distributions with no tempering and a universal tail parameter  $\alpha \approx 1.1$  for a particular non-dimensionalization of the resting time, showing close correspondence to the theory of Martin et al. (2014), describing some imperfections in their results, and providing a signature of burial useful for interpreting experimental data.

## 2 Stochastic theory

We define a volume of downstream length  $L$  which contains some number  $n$  of moving particles in the water flow and some number  $m$  of stationary particles in the bed at an instant  $t$ . For simplicity, we consider all particles as approximately spherical with the same diameter  $2a$ , so that mobility and packing characteristics are similar from one par-



**Figure 1.** The conceptual picture of a control volume containing  $n$  moving particles and  $m$  resting particles. Migration in, entrainment, deposition, and migration out are represented by arrows, and the probability distribution of bed elevations is illustrated.

115 ticle to the next. We follow Ancey et al. (2008) to prescribe four events which can occur  
 116 at any instant to modify the populations  $n$  and  $m$ , and we characterize these events  
 117 using probabilities per unit time, or rates. The events are: (1) migration of a moving par-  
 118 ticle into the volume from upstream; (2) the entrainment of a stationary particle into  
 119 motion within the volume; (3) the deposition of a moving particle to rest within the vol-  
 120 ume; and (4) the migration of a moving particle out of the volume to downstream.

121 As the events occur at random intervals, they set up a joint stochastic evolution  
 122 of the populations  $n$  and  $m$ , leading to a joint probability mass function (pmf)  $P(n, m, t)$   
 123 having marginal pmfs  $P(n, t) = \sum_m P(n, m, t)$  and  $P(m, t) = \sum_n P(n, m, t)$  for the  
 124 number of particles in motion and at rest in the volume at  $t$ . These concepts are depicted  
 125 in figure 1.

The populations  $n$  and  $m$  link to the bedload transport rate  $q_s$  and the bed ele-  
 126 vation in the control volume. The mean bedload transport rate  $q_s$  is described by  $q_s \propto u_s \langle n \rangle$ , where  $u_s$  is the characteristic velocity of moving bedload and  $\langle n \rangle$  is the mean num-  
 127 ber of particles in motion (e.g. Ancey et al., 2008; Charru, Mouilleron, & Eiff, 2004; Fur-  
 128 bish, Haff, Roseberry, & Schmeeckle, 2012). To link the bed elevation  $z$  to the number  
 129 of resting particles  $m$ , we prescribe a mean number of particles at rest  $m_0$  and introduce  
 130 a packing fraction  $\phi$  of grains in the bed. Then from the bed geometry, considering a two-  
 131 dimensional bed (e.g. Einstein, 1950; Paintal, 1971), the deviation from the mean bed  
 132 elevation is

$$z(m) = \frac{\pi a^2}{\phi L} (m - m_0) = z_1(m - m_0). \quad (1)$$

126 The constant  $z_1 = \pi a^2 / (\phi L)$  is an important scale of the problem.  $z_1$  is the magnitude  
 127 of bed elevation change (in an average sense across the control volume) associated with  
 128 the addition or removal of a single particle.

We write the rates for the four possible transitions as (e.g. Ancey et al., 2008):

$$R_{MI}(n+1, m|n, m) = \nu \quad \text{migration in,} \quad (2)$$

$$R_E(n+1, m-1|n, m) = \lambda(m) + \mu(m)n \quad \text{entrainment,} \quad (3)$$

$$R_D(n-1, m+1|n, m) = \sigma(m)n \quad \text{deposition,} \quad (4)$$

$$R_{MO}(n-1, m+1|n, m) = \gamma n \quad \text{migration out.} \quad (5)$$

These rates are independent of the past evolution of the process, only depending on its current state  $n, m$ . This is the Markov hypothesis (e.g. Cox & Miller, 1965), which implies the time intervals between subsequent transitions are exponentially distributed (e.g. Gillespie, 2007).

In these equations  $\nu$  and  $\gamma$ , characterizing migration into and out of the volume are constants which do not depend on the populations  $n$  and  $m$ . In contrast,  $\lambda(m)$ ,  $\mu(m)$ , and  $\sigma(m)$ , characterizing entrainment, collective entrainment, and deposition, respectively, depend on  $m$ . Collective entrainment was introduced in Ancey et al. (2008) as a means to obtain bedload fluctuations of realistic magnitude and rectify some short-comings of an earlier work (Ancey, Böhm, Jodeau, & Frey, 2006). This quantity has been discussed carefully in follow-up studies (e.g. Ancey & Heyman, 2014; Heyman, Ma, Mettra, & Ancey, 2014; Heyman, Mettra, Ma, & Ancey, 2013; Ma et al., 2014), and we won't dwell on this topic.

The  $m$  dependence in these transition rates is through the local bed elevation  $z(m)$ . As is well-known, bed elevation changes modify the likelihood of entrainment and deposition (Sawai, 1987; Wong et al., 2007), meaning the entrainment and deposition rates depend on bed elevation. Wong et al. (2007) concluded from their experiments that bed elevation changes induce an exponential variation in entrainment and deposition probabilities, while Sawai (1987) concluded from his own experiments that the variation is linear. For simplicity, we incorporate the scaling of Sawai (1987) and note its equivalence to the Wong et al. (2007) scaling when bed elevation changes are small.

This scaling can be written  $\chi(m) = \chi_0(1 \pm z_1 z(m)/(2l)^2)$ , where  $\chi = \lambda, \mu, \sigma$ , and entrainment parameters take the plus sign while the deposition parameter takes the minus. With these substitutions, the entrainment and deposition rates become:

$$R_E(n+1, m-1|n, m) = (\lambda_0 + \mu_0 n)(1 + z_1 z(m)/(2l)^2) + O(dt), \quad (6)$$

$$R_D(n-1, m+1|n, m) = \sigma_0(1 - z_1 z(m)/(2l)^2)n + O(dt). \quad (7)$$

In these equations,  $l$  is a length scale of bed elevation change at which the entrainment and deposition rates are significantly affected, which we will clarify, and the ratio  $z_1/l$  controls the sensitivity of these effects to the addition or removal of a single particle. At  $z(m) = 0$ , these reduce to the transition rates of the Ancey et al. (2008) theory. Away from this elevation, entrainment and deposition are alternatively suppressed and accentuated, depending on the sign of  $z(m)$ , introducing a mean-reverting character to bed elevation changes.

In terms of the transition rates, we set up the Master equation describing the flow of probability through time as (e.g. Ancey et al., 2008; Cox & Miller, 1965; Gillespie, 1992):

$$\begin{aligned} \frac{\partial P}{\partial t}(n, m; t) = & \nu P(n-1, m; t) + \{\lambda(m+1) + [n-1]\mu(m+1)\}P(n-1, m+1; t) \\ & + [n+1]\sigma(m-1)P(n+1, m-1; t) + [n+1]\gamma P(n+1, m; t) \\ & - \{\nu + \lambda(m) + n\mu(m) + n\sigma(m) + n\gamma\}P(n, m; t). \end{aligned} \quad (8)$$

The solution  $P(n, m; t)$  of this equation provides the statistics of  $n$  and  $m$ , meaning it provides the statistics of bedload flux  $q_s$  and bed elevation  $z$  through equations ?? and 1. This equation represents a two-species stochastic birth-death model (e.g. Cox & Miller, 1965; Pielou, 1977) for the joint dynamics of bedload transport and bed elevations.

161 **3 Simulations**

162 Equation 8 does not admit an analytical solution, at least by the methods applied  
 163 to similar systems in the population ecology literature (e.g Swift, 2002). This difficulty  
 164 is a consequence of the products between  $n$  and  $m$  in equation 8. In response to this, we  
 165 proceed with numerical simulations. The simulation of these birth-death type master equa-  
 166 tions has been extensively studied for its relevance to chemical physics and population  
 167 ecology. Balancing conceptual simplicity against computational efficiency, we choose the  
 168 classic and foundational Gillespie algorithm (Gillespie, 1977, 1992, 2007).

169 The Gillespie algorithm leverages the defining property of Markov processes: due  
 170 to memorylessness in the rates 2-5, the time interval between subsequent transitions is  
 171 exponentially distributed (e.g. Cox & Miller, 1965). Once a transition occurs, its type  
 172 can be randomly selected using the relative rates of all possible transitions. Accordingly,  
 173 to step our birth-death process through a single transition, we can select the time to the  
 174 next transition by drawing a random value from an exponential distribution, then we  
 175 select the type of transition which occurs by selecting a random value from a uniform  
 176 distribution. After this transition is enacted, i.e., by stepping  $n$  and  $m$  by the shifts as-  
 177 sociated with the type of transition which occurred, this two-stage selection is repeated  
 178 indefinitely to form an exact realization of the stochastic process (Gillespie, 1977, 1992,  
 179 2007). Computing the joint statistics of  $n$  and  $m$  from these simulations provides a nu-  
 180 matical approximation of equation 8.

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 182  
 183 **Table 1.** Parameters from Ancey et al. (2008)  
 184 experiments describing the rates of migration in,  
 185 entrainment, deposition, and migration out of  
 186 the control volume when  $z(m) = 0$ . All units are  
 187  $s^{-1}$  (probability/time).

Flow	$\nu$	$\lambda_0$	$\mu_0$	$\sigma_0$	$\gamma$
(a)	5.45	6.59	3.74	4.67	0.77
(g)	7.74	8.42	4.34	4.95	0.56
(i)	15.56	22.07	3.56	4.52	0.68
(l)	15.52	14.64	4.32	4.77	0.48
(n)	15.45	24.49	3.64	4.21	0.36

201 0.3cm in accord with the Ancey et al. (2008) experiments.

202 One simulated realization of our joint stochastic process for bed elevations and bed-  
 203 load transport is depicted in figure 2 (a). These realizations determine the joint statis-  
 204 tics of  $n$  and  $m$ . The bedload transport statistics are depicted for a subset of all simu-  
 205 lation conditions in figure 2 (b). The Ancey et al. (2008) theory predicts negative bino-  
 206 mial distributions for the number of moving particles within the control volume, and the  
 207 mathematical form of these distributions is apparently not changed by our extension to  
 208 account for feedbacks between bed elevation changes and entrainment and deposition  
 209 probabilities. We obtain an excellent negative binomial fit to the marginal probability  
 210 distribution of  $n$ ,  $P(n) = \sum_m P(n, m, t)$ , for all 40 of our simulation results. However,  
 211 particle activity statistics, including the mean particle activity and the variance of par-  
 212 ticle activity, are definitely shifted by the inclusion of differential mobility with bed el-  
 213 evation changes. Bed elevation changes appear to buffer the magnitude of bedload fluc-

Using Gillespie's stochastic simulation algorithm, we simulated 5 flow conditions from the Ancey et al. (2008) experiments, prescribing 8 different values of the parameter  $l$ , ranging from half of the particle radius ( $l = a/2$ ) to 4 particle diameters ( $l = 8a$ ) for a total of 40 simulations. The parameters used for our simulations are taken from the experiments of Ancey et al. (2008) and are summarized in table 1. They are labeled (a), (g), and so on, in order of increasing bedload transport rate. Depending on the flow condition, exact stochastic trajectories consistent with equation 8 were simulated for either a 500 hr or 1,000 hr duration. In all simulations, we take the packing fraction  $\phi = 0.6$ , a typical value for a pile of spheres (Bennett, 1972), and set  $\Delta x = 22.5\text{cm}$  and  $a =$



**Figure 2.** Figure (a) depicts timeseries of particle activity and bed elevation over a 20 minute interval. Figures (b) and (c) display probability distribution functions of bed elevation (equation 1) and particle activity for a subset of the simulations. Colors represent flow conditions, while differing line styles represent different values of the differential mobility parameter  $l$ .

tuations by up to 30 percent, which is an expected effect of the model, since a rapid increase in the bedload rate induced by a series of many entrainments will lower the bed elevation and increase the probability of deposition, buffering the magnitude of the bedload rate increase.

Our bed elevation timeseries exhibit longer temporal correlations than related bedload activity series, evident in figure 2 (a). All 40 of our simulations develop clean unimodal distributions of bed elevations which are fit by Gaussian distributions with excellent correlation, and a subset of these marginal bed elevation pdfs with their Gaussian fits are displayed in figure 2 (c). The mean number of particles resting on the bed is  $m_0$ , corresponding to a relative elevation  $z(m_0) = 0$ . The variance of bed elevations is controlled by the differential mobility parameter  $l$ . Apparently, the simulations support the conclusion that  $\text{var}(m) = (l/z_1)^2$ . This conclusion is evident in figure 3, with generally excellent correspondence between this relationship and the simulation points, with some scatter we attribute to the finite duration of our simulations.

To compute the resting time distribution; at each elevation  $m$ , we extract the set of all departure times from this elevation, or times at which the bed was at this elevation and a deposition occurred; then we extract the set of all return times to this elevation, or times at which the bed was one increment above this elevation ( $m+1$ ) and an entrainment occurred. Taking differences between these two time-series returns the set of all return times from above marginal to the elevation  $m$ , which we binned across a 0.5s interval to compute the cumulative probabilities of return times at each elevation  $m$ ,  $P(T_r > t|m)$ . Following earlier investigators, we computed the unconditional or overall rest time distribution as the convolution of these conditional distributions over all bed elevations (e.g. Voepel et al., 2013; Yang & Sayre, 1971):

$$P(T_r > t) = \sum_m P(m)P(T_r > t|m), \quad (9)$$

where  $P(m)$  is the pdf of bed elevation like those depicted in figure 2 (b) and the sum is over all bed elevations attained during the simulation.

This analysis derives unconditional exceedance probabilities of resting times with heavy power-law tails. A subset of all of our simulation results are depicted in figure 4 (a)-(d). Apparently, for suitably long times, the tail parameter  $\alpha$  of these resting time distributions is independent of flow conditions or the differential mobility parameter  $l$ . However, the timescale at which particle resting transitions from exponential to power-law scaling shifts with flow conditions and  $l$ . Martin et al. (2014) obtained an approximate collapse at the tails of their experimental resting time distributions using the reciprocal of the rate of entrainment or deposition events occurring. They denoted this rate by  $a$ , so that their timescale was  $1/a$ . Scaling the resting times by  $1/a$  provides an incomplete collapse of the tails of our simulated resting time distributions, which may describe the incomplete collapse of the experimental data of Martin et al. (2014). It collapses the tails across flow conditions when  $l$  (the standard deviation of bed elevation) is fixed, i.e., it leads to the collapse seen between figures 4 (a) and 4 (b), but if  $l$  (which is the standard deviation of bed elevation) varies with flow, then the power-law scaling of resting times is no longer controlled by  $1/a$  alone. Instead, we must include a factor representing the differential entrainment and deposition characteristics of grains as the bed elevation changes. The timescale which provides universal collapse of the power-law tails of all our simulations is  $T_0 = l/(z_1 E)$ , where  $E$  is the entrainment rate.

We can understand  $T_0$  with a physical argument. According to figure 3, the typical length scale of bed elevation fluctuations is  $l$ , and as mentioned the length scale  $z_1$  is the magnitude of bed elevation change enacted by the entrainment or deposition of a single particle. In equilibrium bedload transport, Einstein (1950) tells us the condition  $E = D$  holds: this is a statement of mass conservation. Since  $E$  represents the mean number of particles removed from the bed in a unit of time, the product  $Ez_1$  can be interpreted as a representative velocity scale of bed return. It is the distance the bed lowers with the removal of a single particle divided by the mean time required to remove it. Hence we extract our timescale as a key distance scale over a key velocity scale:  $T_0 = l/(z_1 E)$ . Scaling the resting time as  $T_r/T_0$  exhibits a consistent power-law tail across all of our simulation results.



**Figure 3.** The standard deviation of bed elevation scales one-to-one with the differential mobility parameter ( $l/z_1$ ), indicating the ratio  $l/z_1$  controls the magnitude of bed elevation fluctuations.

## 4 Discussion

Our theory of bed elevations derives results similar to Martin et al. (2014) using bedload transport as a starting point, and it also provides a statistical characterization of bedload transport. Our assumptions derive a heavy-tailed power-law distribution of resting times with a tail parameter  $\alpha \approx 1$  which displays differences across flow conditions which partially collapse upon scaling by an activity timescale.

However, in extension of the Martin et al. (2014) theory, our model reveals another timescale which fully collapses the power-law tails of the resting time distributions, suggesting a universal power-law should characterize the asymptotic resting times of sed-



**Figure 4.** This figure summarizes the resting time exceedance distributions for a subset of all simulations. Part (a) displays resting time distributions for a range of flow conditions at fixed  $l$ , while part (b) displays the collapse obtained by scaling  $T_r$  by  $T_0$ . The collapse between (a) and (b) is analogous to Martin et al. (2014), and it is induced by the factor of  $1/E$  within  $T_0$ . Parts (c) and (d) display a similar collapse for a fixed flow condition at variable  $l$ . In this case, collapse is driven by the factor of  $z_1/l$  within  $T_0$ , and this influence of differential mobility in resting statistics has not to our knowledge been noticed up to now.

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iment undergoing burial if the assumptions of our model are correct. This timescale in-  
cludes an additional factor characterizing the dependence of entrainment and deposition  
probabilities on changes in local bed elevations. We hypothesize this new factor may ex-  
plain some of the differences between field (e.g. Olinde & Johnson, 2015) and laboratory  
(e.g. Martin et al., 2014) observations of resting time distributions, providing additional  
information to determine whether burial is the dominant mechanism of the heavy-tailed  
sediment resting times observed in natural channels.

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Pierce The model describes why Martin et al (2014) obtained incomplete collapse  
between their resting time distributions. They didn't include the  $(z_1/l)^2$  type factor in  
their scaling time-scale It is a first joint stochastic description of bed elevations and bed-  
load transport, mixing martin 2014 and ancey 2008. It provides a mechanism for heavy-  
tailed resting times and implies super-diffusion of bedload and a virtual velocity of sed-  
iment which decreases toward zero with time. what else? help

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Hassan 1. Discuss what is special in the model and how it works. (most of point  
number 2)

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2. Your point number 1  
3. The last paragraph from the introduction  
4. Your point number 3 with implications (What this means for the study of bed-  
load (and channel morphology in streams). The implications should be short.

301      **5 Conclusion**

302      **Acknowledgments**

303      The computer code used to simulate the presented model is available upon request from  
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306      **References**

- Ancey, C., Böhm, T., Jodeau, M., & Frey, P. (2006). Statistical description of sediment transport experiments. *Physical Review E - Statistical, Nonlinear, and Soft Matter Physics*, 74(1), 1–14. doi: 10.1103/PhysRevE.74.011302
- Ancey, C., Davison, A. C., Böhm, T., Jodeau, M., & Frey, P. (2008). Entrainment and motion of coarse particles in a shallow water stream down a steep slope. *Journal of Fluid Mechanics*, 595(2008), 83–114. Retrieved from [http://www.journals.cambridge.org/abstract{\\\_}S0022112007008774](http://www.journals.cambridge.org/abstract{\_}S0022112007008774) doi: 10.1017/S0022112007008774
- Ancey, C., & Heyman, J. (2014). A microstructural approach to bed load transport: mean behaviour and fluctuations of particle transport rates. *Journal of Fluid Mechanics*, 744(2014), 129–168. Retrieved from [http://www.journals.cambridge.org/abstract{\\\_}S0022112014000743](http://www.journals.cambridge.org/abstract{\_}S0022112014000743) doi: 10.1017/jfm.2014.74
- Bennett, C. H. (1972). Serially deposited amorphous aggregates of hard spheres. *Journal of Applied Physics*, 43(6), 2727–2734. doi: 10.1063/1.1661585
- Bradley, D. N. (2017). Direct Observation of Heavy-Tailed Storage Times of Bed Load Tracer Particles Causing Anomalous Superdiffusion. *Geophysical Research Letters*, 44(24), 12,227–12,235. doi: 10.1002/2017GL075045
- Bradley, N. D., & Tucker, G. E. (2012). Measuring gravel transport and dispersion in a mountain river using passive radio tracers. *Earth Surface Processes and Landforms*, 37, 1034–1045. doi: 10.1002/2017GL075045
- Celik, A. O., Diplas, P., & Dancey, C. L. (2014). Instantaneous pressure measurements on a spherical grain under threshold flow conditions. *Journal of Fluid Mechanics*, 741, 60–97. Retrieved from [http://www.journals.cambridge.org/abstract{\\\_}S0022112013006320](http://www.journals.cambridge.org/abstract{\_}S0022112013006320) doi: 10.1017/jfm.2013.632
- Charru, F., Mouilleron, H., & Eiff, O. (2004). Erosion and deposition of particles on a bed sheared by a viscous flow. *Journal of Fluid Mechanics*, 519(2004), 55–80. Retrieved from [http://www.journals.cambridge.org/abstract{\\\_}S0022112004001028](http://www.journals.cambridge.org/abstract{\_}S0022112004001028) doi: 10.1017/S0022112004001028
- Cox, D. R., & Miller, H. (1965). *The Theory of Stochastic Processes*. London: Chapman and Hall.
- Einstein, H. A. (1937). *Bed-load transport as a probability problem* (Unpublished doctoral dissertation). ETH Zurich.
- Einstein, H. A. (1950). *The Bed-Load Function for Sediment Transportation in Open Channel Flows* (Tech. Rep. No. 1026). Washington, DC: United States Department of Agriculture.
- Fathel, S. L., Furbish, D. J., & Schmeeckle, M. W. (2015). Experimental evidence of statistical ensemble behavior in bed load sediment transport. *Journal of Geophysical Research F: Earth Surface*, 120(11), 2298–2317. doi: 10.1002/2015JF003552
- Furbish, D. J., Haff, P. K., Roseberry, J. C., & Schmeeckle, M. W. (2012). A probabilistic description of the bed load sediment flux: 1. Theory. *Journal of Geophysical Research: Earth Surface*, 117(3). doi: 10.1029/2012JF002352
- Gaeuman, D., Stewart, R., Schmandt, B., & Pryor, C. (2017). Geomorphic response to gravel augmentation and high-flow dam release in the Trinity River,

- 352 California. *Earth Surface Processes and Landforms*, 42(15), 2523–2540. doi:  
 353 10.1002/esp.4191
- 354 Gillespie, D. T. (1977). Exact stochastic simulation of coupled chemical reactions.  
 355 *Journal of Physical Chemistry*, 81(25), 2340–2361. doi: 10.1021/j100540a008
- 356 Gillespie, D. T. (1992). *Markov Processes: An Introduction For Physical Sciences*.  
 357 Academic Press, Inc.
- 358 Gillespie, D. T. (2007). Stochastic Simulation of Chemical Kinetics. *Annual  
 359 Review of Physical Chemistry*, 58(1), 35–55. Retrieved from <http://www.annualreviews.org/doi/10.1146/annurev.physchem.58.032806.104637>  
 360 doi: 10.1146/annurev.physchem.58.032806.104637
- 361 Habersack, H. M. (2001). Radio-tracking gravel particles in a large braided river  
 362 in New Zealand: A field test of the stochastic theory of bed load trans-  
 363 port proposed by Einstein. *Hydrological Processes*, 15(3), 377–391. doi:  
 364 10.1002/hyp.147
- 365 Hassan, M. A., & Bradley, D. N. (2017). Geomorphic controls on tracer particle dis-  
 366 persions in gravel bed rivers. In *Gravel-bed rivers: Processes and disasters* (pp.  
 367 159–184). New York, NY: John Wiley & Sons Ltd. doi: 10.16719/j.cnki.1671  
 368 -6981.2015.03.007
- 369 Hassan, M. A., Church, M., & Schick, A. P. (1991). Distance of movement of coarse  
 370 particles in gravel bed streams. *Water Resources Research*, 27(4), 503–511.  
 371 doi: 10.1029/90WR02762
- 372 Hassan, M. A., Voepel, H., Schumer, R., Parker, G., & Fraccarollo, L. (2013). Dis-  
 373 placement characteristics of coarse fluvial bed sediment. *Journal of Geophysi-  
 374 cal Research: Earth Surface*, 118(1), 155–165. doi: 10.1029/2012JF002374
- 375 Heyman, J., Bohorquez, P., & Ancey, C. (2016). Entrainment, motion, and depo-  
 376 sition of coarse particles transported by water over a sloping mobile bed. *Jour-  
 377 nal of Geophysical Research: Earth Surface*, 121(10), 1931–1952. doi: 10.1002/  
 378 2015JF003672
- 379 Heyman, J., Ma, H. B., Mettra, F., & Ancey, C. (2014). Spatial correlations in  
 380 bed load transport: Evidence, importance, and modeling. *Journal of Geophys-  
 381 ical Research: Earth Surface*, 119(8), 1751–1767. Retrieved from <http://doi.wiley.com/10.1002/2013JF003003>.Received  
 382
- 383 Heyman, J., Mettra, F., Ma, H. B., & Ancey, C. (2013). Statistics of bedload trans-  
 384 port over steep slopes: Separation of time scales and collective motion. *Geo-  
 385 physical Research Letters*, 40(1), 128–133. doi: 10.1029/2012GL054280
- 386 Ma, H., Heyman, J., Fu, X., Mettra, F., Ancey, C., & Parker, G. (2014). Bed  
 387 load transport over a broad range of timescales: Determination of three  
 388 regimes of fluctuations. *Journal of Geophysical Research*, 119(12), 1–21.  
 389 doi: 10.1002/2014JF003308.Received  
 390
- 391 Malmon, D. V., Reneau, S. L., Dunne, T., Katzman, D., & Drakos, P. G. (2005). In-  
 392 fluence of sediment storage on downstream delivery of contaminated sediment.  
 393 *Water Resources Research*, 41(5), 1–17. doi: 10.1029/2004WR003288
- 394 Martin, R. L., Jerolmack, D. J., & Schumer, R. (2012). The physical basis for  
 395 anomalous diffusion in bed load transport. *Journal of Geophysical Research: Earth  
 396 Surface*, 117(1), 1–18. doi: 10.1029/2011JF002075
- 397 Martin, R. L., Purohit, P. K., & Jerolmack, D. J. (2014). Sedimentary bed evolution  
 398 as a mean-reverting random walk: Implications for tracer statistics. *Geophys-  
 399 ical Research Letters*, 41(17), 6152–6159. doi: 10.1002/2014GL060525
- 400 Meyer-Peter, E., & Müller, R. (1948). Formulas for Bed-Load Transport. *Proceed-  
 401 ings of the 2nd Meeting of the International Association of Hydraulic Research*,  
 402 39–64. doi: 1948-06-07
- 403 Nakagawa, H., & Tsujimoto, T. (1976). On Probabilistic Characteristics of Motion  
 404 of Individual Sediment Particles on Stream Beds. In *Hydraulic problems solved  
 405 by stochastic methods: Second international iahr symposium on stochastic  
 406 hydraulics*. Lund, Sweden.

- 407 Nikora, V. (2002). On bed particle diffusion in gravel bed flows under weak bed load  
 408 transport. *Water Resources Research*, 38(6), 1–9. Retrieved from <http://doi.wiley.com/10.1029/2001WR000513> doi: 10.1029/2001WR000513
- 410 Olinde, L., & Johnson, J. P. L. (2015). Using RFID and accelerometer-embedded  
 411 tracers to measure probabilities of bed load transport, step lengths, and rest  
 412 times in a mountain stream. *Water Resources Research*, 51, 7572–7589. doi:  
 413 10.1002/2014WR016259
- 414 Paintal, A. S. (1971). A Stochastic Model Of Bed Load Transport. *Journal of Hy-  
 415 draulic Research*, 9(4), 527–554. doi: 10.1080/00221687109500371
- 416 Phillips, C. B., Martin, R. L., & Jerolmack, D. J. (2013). Impulse framework for  
 417 unsteady flows reveals superdiffusive bed load transport. *Geophysical Research  
 418 Letters*, 40(7), 1328–1333. doi: 10.1002/grl.50323
- 419 Pielou, E. (1977). *Mathematical Ecology* (1st ed.). New York, NY: John Wiley &  
 420 Sons Ltd.
- 421 Pretzlav, K. L. (2016). *Armor development and bedload transport processes during  
 422 snowmelt and flash floods using laboratory experiments, numerical modeling,  
 423 and field-based motion-sensor tracers* (Unpublished doctoral dissertation).  
 424 University of Texas.
- 425 Sawai, K. (1987). Dispersion of bed load particles. *Bull. Disas. Prev. Res. Inst., Ky-  
 426 oto Univ.*, 37(Part 1, No. 323).
- 427 Swift, R. J. (2002). A Stochastic Predator-Prey Model. *Bulletin of the Irish Mathe-  
 428 matical Society*, 48, 57–63.
- 429 Voepel, H., Schumer, R., & Hassan, M. A. (2013). Sediment residence time dis-  
 430 tributions: Theory and application from bed elevation measurements. *Journal  
 431 of Geophysical Research: Earth Surface*, 118(4), 2557–2567. doi: 10.1002/jgrf  
 432 .20151
- 433 Weeks, E. R., & Swinney, H. L. (1998). Anomalous diffusion resulting from strongly  
 434 asymmetric random walks. *Physical Review E - Statistical Physics, Plasmas,  
 435 Fluids, and Related Interdisciplinary Topics*, 57(5), 4915–4920. doi: 10.1103/  
 436 PhysRevE.57.4915
- 437 Weeks, E. R., Urbach, J. S., & Swinney, H. L. (1996). Anomalous diffu-  
 438 sion in asymmetric random walks with a quasi-geostrophic flow exam-  
 439 ple. *Physica D: Nonlinear Phenomena*, 97(1-3), 291–310. Retrieved from  
 440 <http://linkinghub.elsevier.com/retrieve/pii/0167278996000826> doi:  
 441 10.1016/0167-2789(96)00082-6
- 442 Wong, M., Parker, G., DeVries, P., Brown, T. M., & Burges, S. J. (2007). Ex-  
 443 periments on dispersion of tracer stones under lower-regime plane-bed equi-  
 444 librium bed load transport. *Water Resources Research*, 43(3), 1–23. doi:  
 445 10.1029/2006WR005172
- 446 Yalin, M. S. (1972). *Mechanics of Sediment Transport*. Pergamon Press.
- 447 Yang, C. T., & Sayre, W. W. (1971). Stochastic model for sand dispersion. *Journal  
 448 of the Hydraulics Division, ASCE*, 97(HY2).
- 449 Yano, K. (1969). Tracer Studies on the Movement of Sand and Gravel. In *Proceed-  
 450 ings of the 12th congress iahr, vol 2.* (pp. 121–129). Kyoto, Japan.