



Geologic characterization of nonconformities using outcrop and core analogs: hydrologic implications for injection-induced seismicity

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Abstract. The occurrence of induced earthquakes in crystalline rocks kilometers from deep wastewater injection wells poses questions about the influence nonconformity contacts have on the downward and lateral transmission of pore-fluid pressure and poroelastic stresses. We hypothesize that structural and mineralogical heterogeneities at the sedimentary–crystalline rock nonconformity control the degree to which fluids, fluid pressure, and associated poroelastic stresses are transmitted over long distances across and along the nonconformity boundary. We examined the spatial distribution of physical and chemical heterogeneities in outcrops and core samples of the Great Unconformity in the midcontinent of the United States, capturing a range of tectonic settings and rock properties that we use to characterize the degree of past fluid communication and the potential for future communication. We identify three end-member nonconformity types that represent a range of properties that will influence direct fluid pressure transmission and poroelastic responses far from the injection site. These nonconformity types vary depending on whether the contact is sharp and minimally altered (Type 0), dominated by phyllosilicates (Type I), or secondary non-phyllosilicate mineralization (Type II). Our observations provide geologic constraints for modeling fluid migration and the associated pressure communication and poroelastic effects at large-scale disposal projects by providing relevant subsurface properties and much needed data regarding common alteration minerals that may interact readily with brines or reactive fluids.

1 Introduction

Deep wastewater injection near the nonconformity between the Phanerozoic sedimentary sequence and Proterozoic crystalline basement in the midcontinent United States (Sloss, 1963) is the primary means by which produced formation fluids are disposed of in Class II injection wells (Murray, 2015). Increased rates of seismicity in this region are associated with large volumes of wastewater injection (Ellsworth et al., 2015; Keranen et al., 2013; Nicholson and Wesson, 1990; Petersen et al., 2016; Zhang et al., 2013); the reduction of friction on preexisting faults and pressure diffusion away from the injection point are controlled by the permeability structure of the rocks in the subsurface (Goebel and Brodsky, 2018; Yehya et al., 2018). Recent midcontinent seismicity nucleates on faults in crystalline rocks kilometers from injection sites (Keranen et al., 2014; Weingarten et al., 2015; Zhang et al., 2016) and spans timescales of months to years post-injection, indicating that pore-fluid pressures and/or poroelastic loads are transmitted across or along the nonconformity or through connected fracture systems (including joints, faults, and veins) in the crystalline rocks (Ortiz et al., 2019). The depths of seismicity (up to 11 km) at some injection sites suggest that crystalline basement permeability is perhaps moderate to high (10^{-16} to 10^{-14} m^2) (Zhang et al., 2016) and is dynamically increased by elevated fluid pressures (Rojstaczer, 2008). The observations presented in this paper are also relevant to Class VI injec-

tion wells used for the geologic sequestration of CO₂; several of our analog sites include deep reservoirs being evaluated for CO₂ sequestration (Leetaru et al., 2009; Leetaru and McBride, 2009; Plains CO₂ Reduction (PCOR) Partnership, 2020; Thorleifson, 2008).

Numerical modeling of fluid flow and/or loading stresses associated with poroelastic effects across nonconformities indicate that (1) the presence of a high-storativity, low-permeability basal seal reduces the potential for basement-induced earthquakes; (2) poroelastic effects can trigger seismicity far away from the injection location; (3) the presence of conductive faults, including those that cut the nonconformity and those that are isolated in the basement, can provide direct fluid or fluid pressure pathways; and (4) permeable cross-nonconformity faults may exhibit high rates of seismicity (Chang and Segall, 2016; Goebel and Brodsky, 2018; Ortiz et al., 2019; Yehya et al., 2018; Zhang et al., 2013).

In this paper we summarize geologic observations made at the nonconformity zone, the altered rock volume surrounding the nonconformable contact. This zone varies in thickness and is defined by mineralogic and structural alteration of the protolith rocks surrounding the nonconformity. We characterize nonconformity zones associated with Precambrian granite, gabbro, gneiss, and schists that are overlain by porous sedimentary rocks including sandstone and mixed carbonate-clastic sequences. Study site locations were chosen based on their distribution within the midcontinent region and the suite of lithologies present (Fig. 1). These analogs represent the diversity of the nonconformity in the United States midcontinent region and are analogs for deep fluid injection from produced waters (Class II) and sequestration of CO₂ (Class VI). At each site we document the lithology and structural features of the rocks on either side of the nonconformity to characterize the range of rock types associated with the contact and identify any evidence of past cross-contact fluid flow. We present data on the mineralogic and structural heterogeneities observed in the outcrop and core, and these observations serve as proxies for variation in mineral alteration and deformation surrounding subsurface nonconformity zones, which may impact the future migration of fluids along and across the contact.

We find three end-member types of nonconformity zones. These zones range from diffuse to sharp (Type 0), can be phyllosilicate-rich (Type I), or can be dominated by non-phyllosilicate secondary minerals (Type II). Each contact type observed in this study has a range of mineralized textures and structural discontinuities. Due to weathering, deformation, diagenesis, and fluid–rock interactions, the nonconformity zone may be hydraulically heterogeneous at scales of millimeters to tens of meters and influence the migration of fluid and fluid pressures away from the injection well. Characterizing variations in rock properties at the nonconformity zone is critical for safe implementation of deep fluid injection, as the dimensions and hydraulic properties of the rocks in the nonconformity zones impact the subsurface

flow regimes (Ortiz et al., 2019). The lithologic character of the nonconformity zone has implications for hydraulically connected regions by allowing direct fluid communication, changes in pore-fluid pressure, and/or poroelastic loads. Because pressure diffusion and fluid migration depend on the permeability structure at a given location, our work can be used to improve hydrogeologic models that test the impact of lithologic changes and cross-nonconformity fractures on the transmission of pore fluids and/or poroelastic stress. We present results from hydrogeologic models based on observations of nonconformity zone characteristics, thereby testing the impact various nonconformity zone types have on the transmission of pore fluids.

2 Geologic setting

The North American craton, Laurentia, includes the Precambrian shields, the platforms and basins of the North American interior, and the reactivated Cordilleran foreland of the southwestern United States (Fig. 1). The craton includes Archean blocks, the Yavapai–Mazatzal and Grenville accretionary belts, and failed rifts (Hoffman, 1988; Marshak et al., 2017; Whitmeyer and Karlstrom, 2007). Precambrian exhumation produced erosional surfaces on the top of the crystalline basement, which were buried by Phanerozoic clastic and marine sedimentary rocks (Marshak et al., 2017; Sloss, 1988). The nonconformities studied in this paper are located within the Superior craton, an Archean basement complex of granite–greenstone or higher-grade equivalent overlain by erosional remnants of early Proterozoic platform facies (Hoffman, 1988), the Yavapai–Mazatzal province, 1.76–1.65 Ga juvenile arc terrane that includes the Central Plains Orogen (Karlstrom and Humphreys, 1998; Sims, 1985; Sims and Peterman, 1986), the Grenville province, 1.3–1.0 Ga imbricate thrust slices formed during continent–continent collision (Rivers, 1997), and the Midcontinent Rift, an approximately 1.1 Ga failed rift system dominated by volcanic rocks and basin-fill sedimentary rocks (Ojakangas et al., 2001) (Fig. 1).

2.1 Study areas

2.1.1 Outcrop locations

Exposed at Presque Isle and Hidden Beach along the southern shore of Lake Superior, Michigan, the nonconformity is defined by Proterozoic Jacobsville Sandstone overlying early Proterozoic altered peridotite crystalline basement (Fig. 2). The geologic history of the serpentinized peridotite is not well-constrained; it is thought that the peridotite was serpentinized between 1.80 and 1.1 Ga (Gair and Thaden, 1968). The overlying Jacobsville Sandstone is a dominantly fluvial sequence of feldspathic and quartzose sandstones (Malone et al., 2016), and at this study locality it consists of a variably indurated pebble to cobble conglomerate and a lentic-

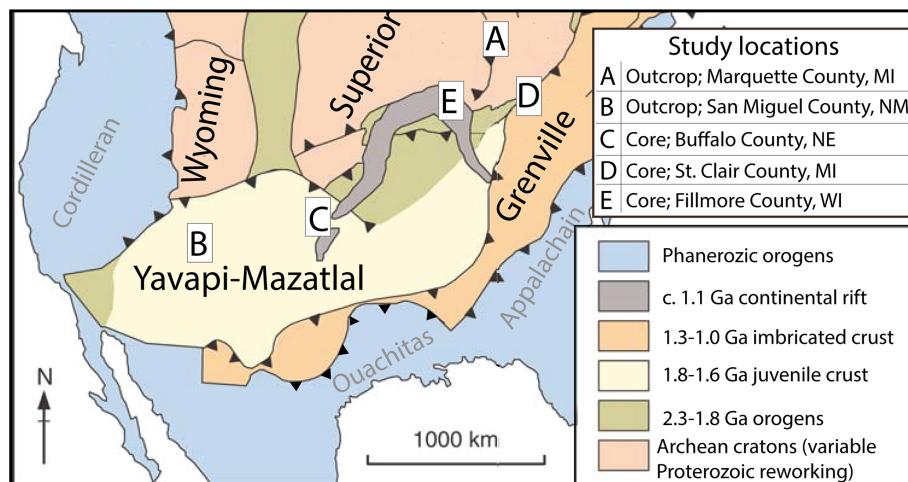


Figure 1. Precambrian tectonic elements map with the location of the nonconformity analog study sites (after St. Onge et al., 2009). **(a)** Lake Superior, Presque Isle, Michigan outcrop, **(b)** Gallinas Canyon, New Mexico outcrop, **(c)** R.C. Taylor 1 core, **(d)** CPC BD-139 core, and **(e)** BO-1 core.

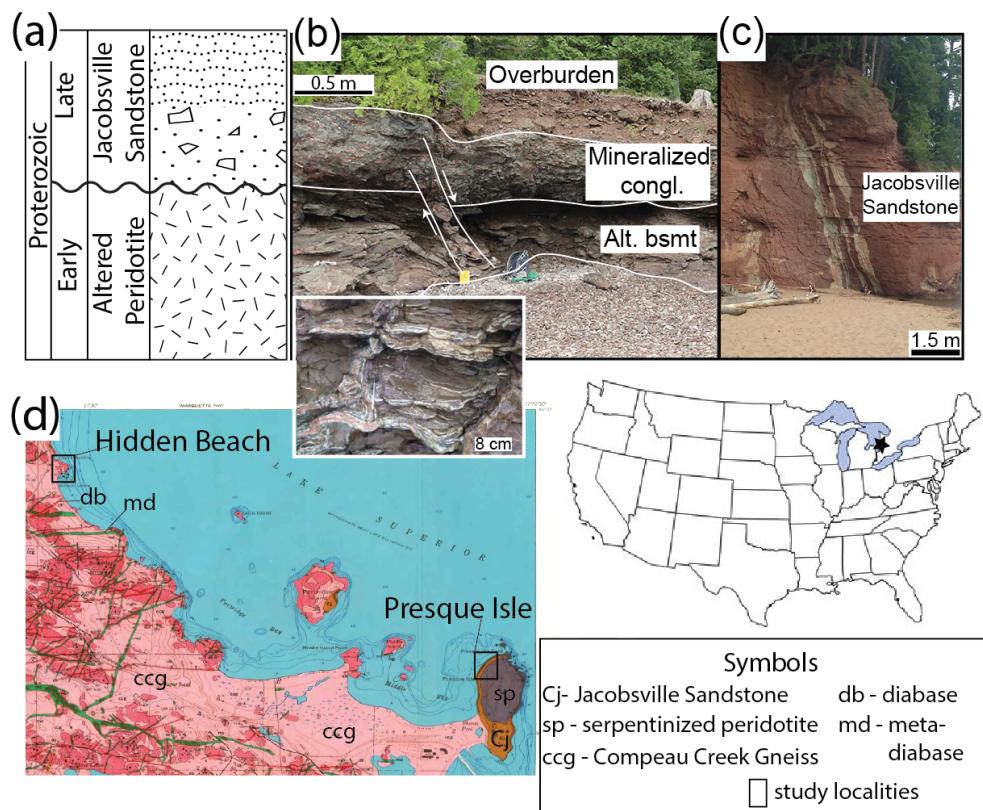


Figure 2. **(a)** Schematic lithologic log at Lake Superior, Michigan, where altered peridotite is overlain by the Jacobsburg Sandstone at Presque Isle and mineralized conglomerates of the Jacobsburg Sandstone overlie the Compeau gneiss at Hidden Beach. **(b)** Presque Isle – outcrop of small fault cutting the contact between the mineralized conglomerate of the Jacobsburg Sandstone and the underlying altered basement; the inset shows stockwork jasperoid veins in the underlying serpentized peridotite basement and **(c)** Hidden Beach outcrops. At this locality the Jacobsburg Sandstone overlies the Proterozoic altered peridotite basement rocks. **(d)** Geologic map of the Marquette, Michigan, field area showing the locations of Hidden Beach and Presque Isle; modified from Gair and Thaden (1968).

ular planar to cross-bedded light red quartz arenite (Fig. 2). The Presque Isle outcrops are analogs for geologic sequestration of CO₂ in the deep saline Jacobsville Sandstone reservoir (Leetaru and McBride, 2009).

The nonconformity at Gallinas Canyon is exposed along a 4 km long section in the southernmost Sangre de Cristo Mountains, New Mexico (Figs. 1 and 3). The outcrop exposure consists of the crystalline rocks of the Yavapai province, highly deformed compositionally layered quartz-feldspathic gneiss, amphibolitic gneiss, felsite, biotite schist, and granitic pegmatite (Lemen et al., 2015) overlain by the Devonian to Mississippian shallow marine transgressive sequence of carbonate and clastic rocks of the Espiritu Santo Formation. The Espiritu Santo Formation primarily consists of limestone and dolomitic limestone, with a basal conglomeratic sandstone known as the Del Padre Member (Baltz and Myers, 1999). The rocks are exposed within north-trending fault-bounded blocks uplifted during the Neogene Laramide Orogeny (Baltz and Myers, 1999; Lessard and Bejnari, 1976). This location provides an analog for the Raton Basin to the east where injection in Class II wells has been linked to basement earthquakes that began in 2001 (Nakai et al., 2017; Rubinstein et al., 2014).

2.1.2 Core samples

The R.C. Taylor 1 core samples the Cambrian La Motte Formation sandstone and sheared Proterozoic granitoids in the Central Plains Orogen of the 1.6 Ga Yavapai–Mazatzal province (Marshak et al., 2017; Sims, 1985; Whitmeyer and Karlstrom, 2007) (Fig. 1). The borehole was drilled adjacent to the Cambridge Arch and is associated with a series of northwest-trending transpressional faults of the Central Plains Orogen (Sims, 1985; Whitmeyer and Karlstrom, 2007) (Fig. 4). In this core, the arkosic La Motte Formation, regionally called the Reagan and Sawatch sandstones, is a fine-grained, well-sorted glauconitic sandstone deposited during a transgression and is an analog for Cambrian sandstones being evaluated for sequestration of CO₂ (Carr et al., 2005; Miller, 2012).

The CPC BD-139 core, recovered from the Michigan Basin, samples the contact between the Cambrian Mount Simon Sandstone and Precambrian altered granitoid gneiss of the Grenville Front Tectonic Zone (Figs. 1 and 5). The Precambrian crystalline rocks captured in this core are characterized as granitic to tonalitic gneiss (Easton and Carter, 1995) that form the basement of the Michigan Basin. The Michigan Basin is a thermally complex intracratonic basin situated over the lower peninsula of Michigan. Unexpectedly high levels of thermal maturity in the Paleozoic strata of the basin are thought to be attributed to elevated basal heat flow occurring up until Silurian time and the prior existence of ~2 km of Pennsylvanian and Permian strata that has since been eroded (Everham and Huntoon, 1999). The Mount Simon Sandstone reservoir is a unit of deep wastewater injec-

tion in Oklahoma, and it is also targeted for CO₂ storage (Barnes et al., 2009; Dewers et al., 2014; Leetaru et al., 2009; Liu et al., 2011).

The BO-1 core samples the lower Cambrian Mount Simon Sandstone overlying a Precambrian layered intrusive complex of altered gabbro and other mafic intrusions and felsic dikes (Fig. 6) (Smith et al., 2019). The crystalline basement rocks are part of the Northeast Iowa Intrusive Complex and are associated with the Midcontinent Rift system (Anderson, 2012). The Midcontinent Rift system extends from Kansas to Lake Superior and then southward through Michigan (Fig. 1). The geologic features associated with the Midcontinent Rift system, include axial basins filled with basalt and immature clastic rocks along with evidence of crustal extension (Ojakangas et al., 2001). The BO-1 core is analogous to several geologic settings anticipated in the subsurface of the midcontinent region where lower Cambrian rocks directly overlie Precambrian mafic igneous rocks of the Midcontinent Rift system along the Great Unconformity (Gilbert, 1962; Mossler, 1995). Northwest-trending fault systems near the borehole were identified by magnetic lineaments and are likely part of the regional NW–SE Belle Plaine Fault Zone (Drenth et al., 2015). The Midcontinent Rift system is being studied for deep injection of CO₂ (Abousif, 2015; Wickstrom et al., 2010).

3 Characterization of the nonconformity

Given the recognized importance of direct fluid transmission, variation in pressure, and poroelastic loads on induced seismicity (Chang and Segall, 2016; Ortiz et al., 2019; Yehya et al., 2018; Zhang et al., 2013), we provide an overview of rock properties observed at the nonconformity using integrated outcrop-based studies in Michigan and New Mexico, as well as analyses of cores from Michigan, Minnesota, and Nebraska (Fig. 1).

3.1 Methods

To describe the nonconformity zone in the core and outcrop and document structures and mineralogy across the boundary, we use a variety of microscale to mesoscale methods including lithological and structural logging of outcrop and core, optical thin-section petrography and X-ray diffraction (XRD) mineralogic studies, whole-rock X-ray fluorescence (XRF) elemental analysis, and gas or air permeability measurements, when possible. We evaluated fracture distribution in the outcrop and core, noting fracture types, the presence of fault and/or shear zones, and associated mineralogy. XRD and XRF were carried out at Utah State University (USU); XRD analysis was carried out at the Western Colorado University (WCU) petrography laboratory. At USU, XRD analyses were done using a Panalytical X’Pert Pro X-ray diffraction spectrometer (40 mA and 45 kV) with

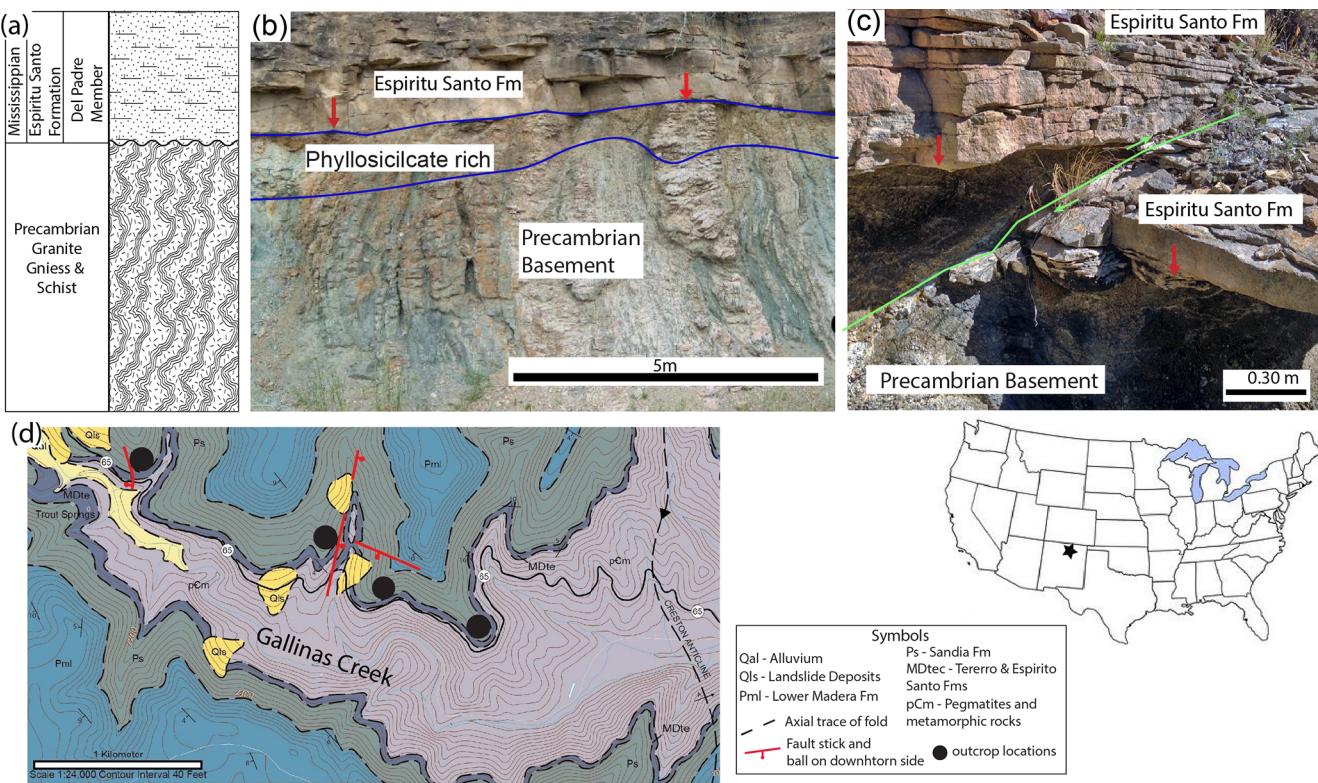


Figure 3. (a) Gallinas Canyon, New Mexico, outcrop lithology log. (b) Precambrian granitic gneiss and schist are overlain by the Mississippian Espiritu Santo Formation; red arrows mark the nonconformity, and blue lines mark the boundary of phyllosilicate alteration. (c) The 4 km long exposure in Gallinas Canyon; the nonconformity is cut by several centimeter to meter displacement faults. Red arrows mark the nonconformity, and the fault is shown by the green line. (d) Geologic map of the Gallinas Canyon study area; modified from Hesselton (2019). Faults that cut the nonconformity are shown in red with a ball on the downthrown side; at the map scale the nonconformity is relatively planar and parallels topographic contour lines.

monochromatic CuK α radiation utilizing X'Pert Highscore software for phase analysis. Whole-rock XRF analyses were conducted at the Washington State University Peter Hooper GeoAnalytical Lab using a Thermo-ARL automated X-ray fluorescence spectrometer. At WCU, XRD analysis was done using a Brucker D8 X-ray diffraction spectrometer (40 mA and 45 kV) with monochromatic CuK α radiation utilizing DIFFRAC.SUITE software for phase analysis.

A total of 25 samples from the BO-1 drill core were selected for gas permeability testing through Schlumberger Rock Mechanics and Core Analysis Services. Profile permeability measurements were made in steady-state conditions with a mini-permeameter whereby gas is injected directly onto the core slab surface. The profile permeameter has a measurable permeability range of 0.1 mD to 3 D ($\sim 9.9 \times 10^{-17}$ to $3.0 \times 10^{-12} \text{ m}^2$).

To illustrate the effects of reduced permeability above the nonconformity on fluid migration we compare three hydrogeologic models of basal reservoir injection that consider continuous and discontinuous zones of altered low-permeability rocks above the basement. We develop three-dimensional models to assess fluid migration along crys-

talline basement faults using MODFLOW, a public domain finite-difference groundwater flow code (Harbaugh and McDonald, 1996; Harbaugh et al., 2000) that solves the following groundwater equation:

$$\frac{\partial}{\partial x} \left(K_x \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left(K_y \frac{\partial h}{\partial y} \right) + \frac{\partial}{\partial z} \left(K_z \frac{\partial h}{\partial z} \right) = S_s \frac{\partial h}{\partial t} + Q(x, y, z, t), \quad (1)$$

where h is the hydraulic head (L), K is the hydraulic conductivity tensor (LT^{-1}), S_s is the specific storage (L^{-1}), Q is the fluid injection source term (i.e., injection well; T^{-1}), and t is time (T). Equation (1) represents single-phase, constant-density groundwater flow in a three-dimensional Cartesian coordinate system. Hydraulic conductivity is a lumped parameter that includes the influence of fluid and medium properties and is defined as $K = kr_f g$ (m^{-1}), where K is hydraulic conductivity (m s^{-1}), k is intrinsic permeability (m^2), r_f is fluid density (997 kg m^{-3} ; water), g is the acceleration due to gravity (9.81 m s^{-2}), and m is the dynamic viscosity of the fluid ($8.9 \times 10^{-4} \text{ kg (m s}^{-1})^{-1}$). Multiple researchers have implemented similar groundwater flow mod-

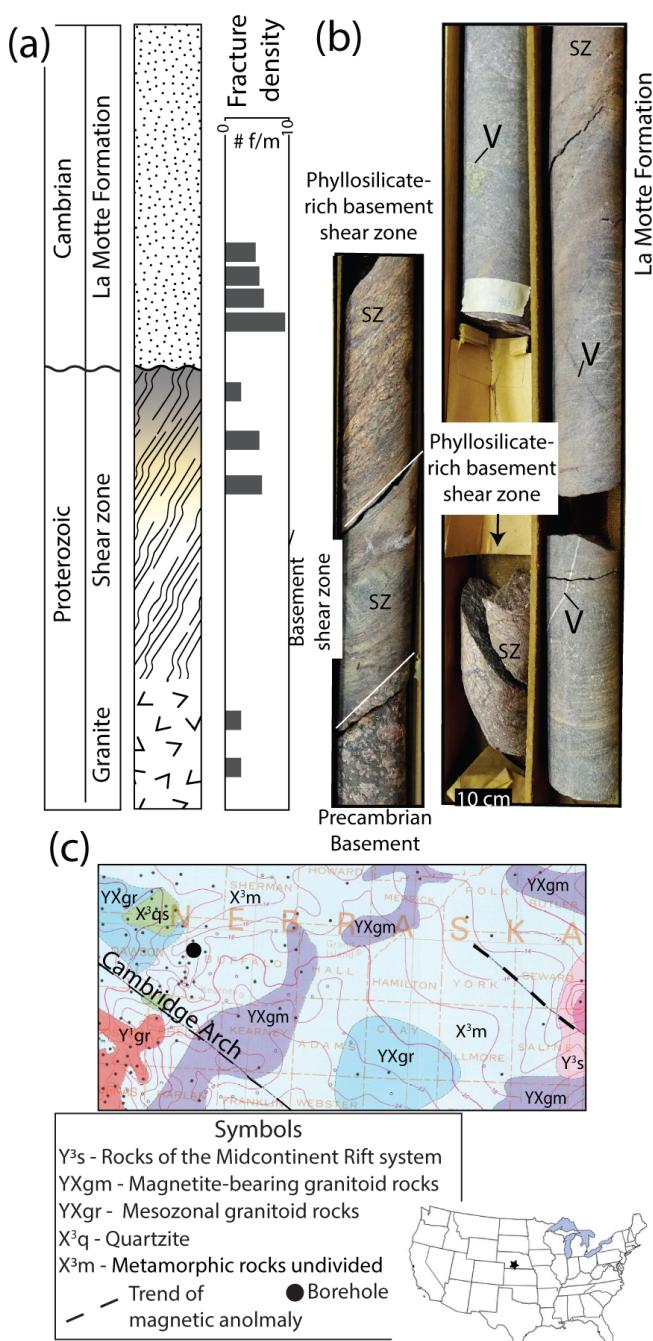


Figure 4. (a) Lithologic log of the R.C. Taylor 1 core, Nebraska, from 3984 to 4038 ft (1214–1231 m) of measured depth. The nonconformity occurs at 4018 ft (1225 m). Fracture density in the core is based on the number of fractures per meter of core. Four lithologic units are identified in the core, including sandstone, sedimentary-rock-hosted shear zone, altered basement shear zone, and minimally altered basement. The nonconformity occurs between the altered basement shear zone and overlying sandstone of the La Motte Formation. (b) Photographs of the R.C. Taylor 1 core. Both the basement shear zone (SZ) and overlying sandstone are cut by veins (V) of quartz, calcite, and Fe-oxides. (c) The borehole location is shown on the Precambrian basement map from Sims (1985).

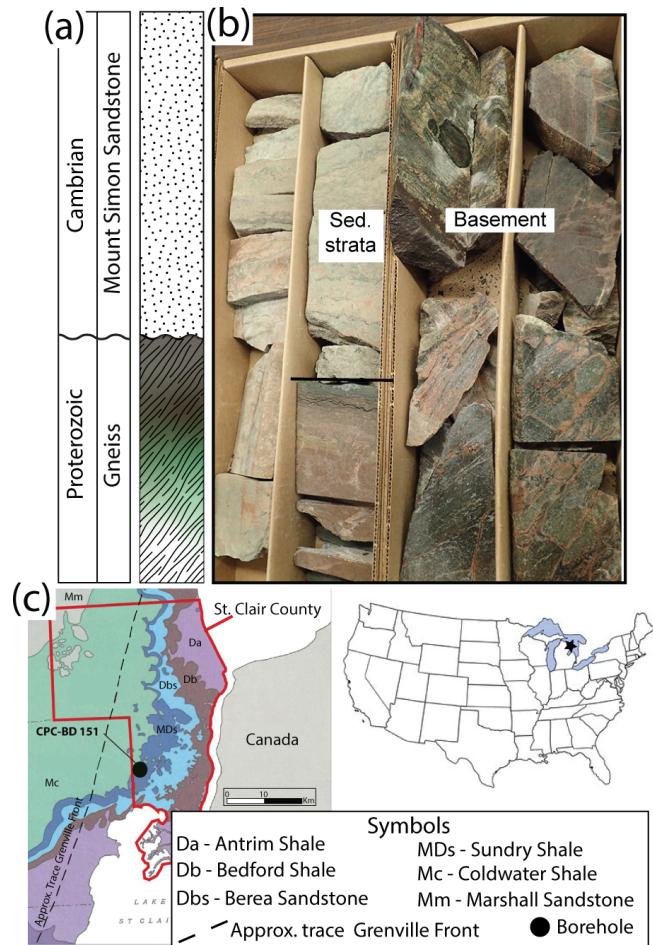


Figure 5. (a) Lithologic log of the CPC BD-139 core, Michigan, from 1404–1412.1 m of measured depth. There are five main lithologic units identified, including sandstone, dolomitized and undolomitized finely foliated gneiss, and dolomitized and undolomitized gneiss with subhorizontal white veins. (b) Photographs of the CPC BD-139 core showing the core between ~1404.5 and 1405.5 m as well as contact between the Cambrian Mount Simon Sandstone (light tan) and the underlying Precambrian gneiss. The gneiss directly at the contact is fine-grained, tan, and dolomitized. This is underlain by green altered gneiss with subvertical pink fractures. This lithology grades into a dark grey gneiss with subhorizontal white veins (core between 1411.5 and 1412.5 m), which extends through the bottom of the logged section. (c) Geologic map St. Clair County, Michigan; modified from Milstein (1987).

els in MODFLOW to investigate pore pressure propagation associated with basal reservoir injection (Zhang et al., 2013, 2016), and a more exhaustive description of cases and boundary conditions can be found in Ortiz et al. (2019). Each of our model simulations includes a 100 m thick basal reservoir ($3 \times 10^{-15} \text{ m}^2$) underlain by 9.9 km of relatively low-permeability ($k_x = k_z = 3 \times 10^{-17} \text{ m}^2$) crystalline basement rock. A 20 m wide conduit-barrier fault ($k_z/k_x = 105$; $k_z = 3 \times 10^{-10} \text{ m}^2$) is present in all simulations as is an injection

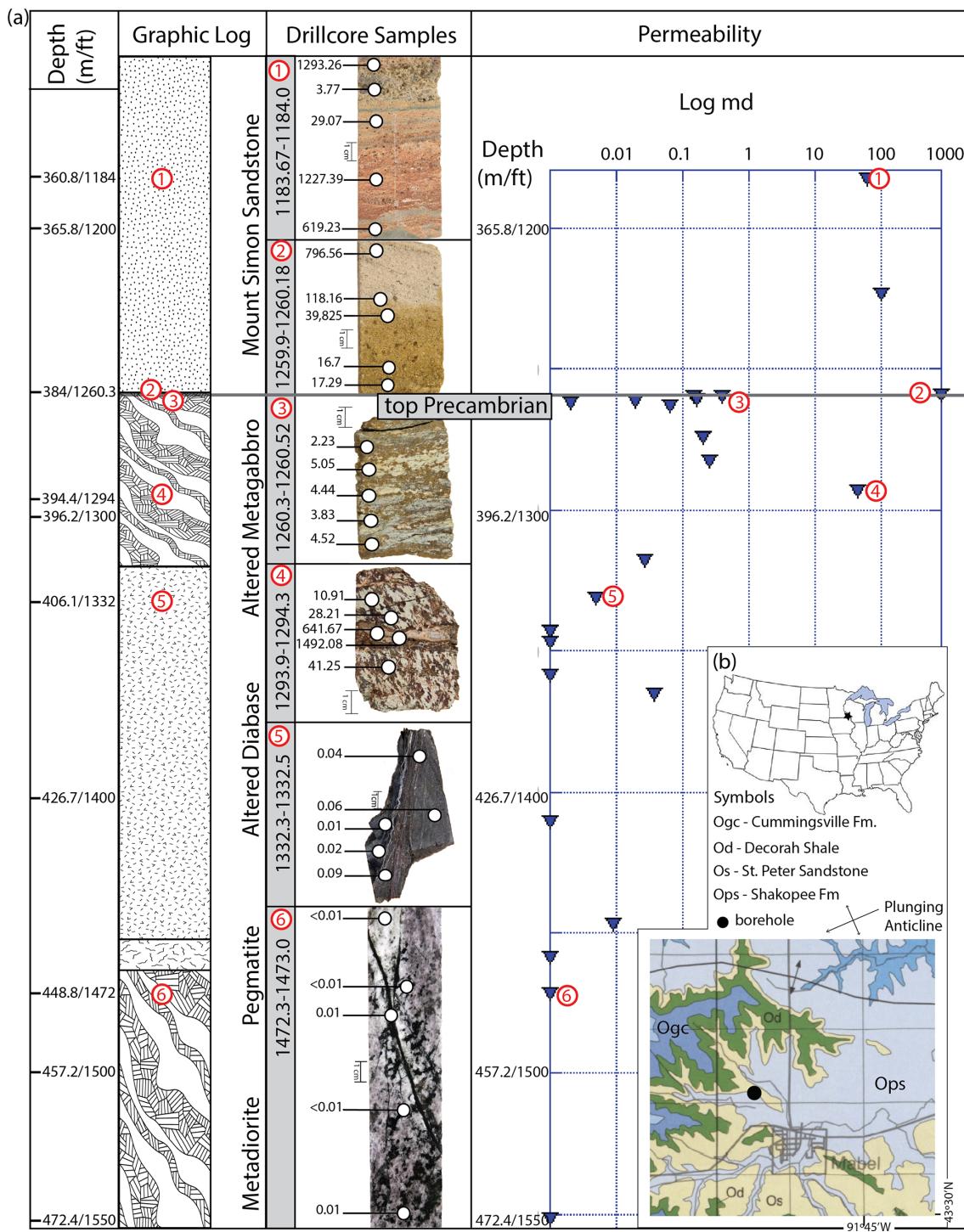


Figure 6. (a) BO-1 lithologic log with select representative core samples of each of the major lithologic units. Above the nonconformity, the analog reservoir, or the injection unit, the Cambrian Mount Simon Sandstone is porous with evidence for both dissolution and oxidation front. The crystalline basement rock consists of foliated, intensely altered, and altered metagabbro with localized faulting, variably altered and faulted diabase localized intrusions, pegmatite dikes, and, at greater depths, relatively unaltered and less-deformed metadiorite. Gas permeability measurements were made on 25 core samples spanning the nonconformity interface. For each core sample tested, five spot measurements were made (locations shown by white circles). For relative comparison across the contact and within the various lithologic units, permeability data (millidarcy, mD) are plotted using a log scale and are the averaged values for each sample. (b) Geologic map modified from Mossler et al. (1995).

well located 150 m from the fault zone. Wellhead pressures reached over 50 m excess hydraulic head after 4 d in response to 5000 m³ d⁻¹ of continuous injection.

3.2 Results

3.2.1 Lake Superior, Michigan

Outcrops of the nonconformity between late Proterozoic Jacobsville Sandstone and early Proterozoic altered peridotite crystalline basement are exposed at Presque Isle and Hidden Beach along the southern shore of Lake Superior, Michigan (Fig. 2) (Lewan, 1972). At Presque Isle the topographic relief of the basement nonconformity varies by 2.5 m over a distance of 1100 m of outcrop (Cuccio, 2017), is locally cut by small-offset faults (30 cm of throw), and is composed of mineralized conglomerate in direct contact with the underlying serpentized peridotite or is transitionally interbedded with the overlying sandstone (Fig. 2b). Where present, the conglomerate consists of subangular to rounded chalcedony, gneiss, and greenstone cobble clasts with fine-grained, poorly sorted, hematite-cemented angular quartz grains.

At Hidden Beach, poorly consolidated basal conglomerates of the Jacobsville Sandstone are in contact with the Precambrian Compeau Creek Gneiss. The quartz arenite consists of fine-grained, angular, moderately sorted quartz with some feldspar. Distinctive bleached open-mode fractures or reduction spots are associated with the lower Jacobsville Sandstone and range in orientation from near-vertical to bedding-parallel. The near-vertical reduction fractures (Fig. 2c) are not observed to extend into the basement.

Optical petrography across the transition from red sandstone protolith to a bleached fracture zone at Hidden Beach reveals a reduction in hematite grain coatings and cements. Whole-rock XRF analysis of the bleached areas of Jacobsville Sandstone indicates a minor depletion of K₂O and a minor enrichment of FeO and MgO relative to the unaltered Jacobsville Sandstone (Fig. 7). At Presque Isle, mineral alteration products in the conglomerate include nontronite, with trace zeolites and iron oxides (Fig. 7). The underlying serpentized peridotite is black to brown, with abundant white carbonate mesh veinlets and localized stockwork jasperoid veins up to 10 cm wide (Fig. 2). Jasperoid mineralization occurs along a few small faults that cross the nonconformity (Cuccio, 2017).

3.2.2 Gallinas Canyon, New Mexico

Devonian to Mississippian carbonate and clastic rocks of the Espiritu Santo Formation deposited on the Proterozoic quartzo-feldspathic and amphibolitic gneiss, biotite schist, and granitic pegmatite (Lemen et al., 2015) are exposed along a 4 km long section in Gallinas Canyon, eastern Sangre de Cristo Mountains, New Mexico. The nonconformity

is cut by centimeter to meter displacement faults; at this location we characterize both the faulted and un-faulted nonconformity zone (Hesseltine, 2019; Kerner, 2015). The top of the basement is defined by a phyllosilicate-rich zone with variable thickness, 0 to > 5 m, that is truncated by the Del Padre Sandstone. Locally the Del Padre Sandstone is laterally discontinuous (Hesseltine, 2019) but is reported to be up to 15 m thick, filling depressions in underlying crystalline rock elsewhere in New Mexico (Armstrong and Mamet, 1974). The carbonate and clastic rocks of the Espiritu Santo Formation include 1 m thick massive, fine-grained, rounded to sub-rounded sandstone with calcite nodules, ~ 1 m of microcrystalline dolomite that transitions upward into a chert nodule limestone, interbedded mudstone and limestone, and a massive microcrystalline limestone bed. A phyllosilicate-rich zone directly below the nonconformity is approximately 60 cm thick and is a poorly lithified zone that marks the transition from highly altered (weathering and hydrothermal alteration) to minimally altered crystalline rock (Fig. 8). The Precambrian crystalline rocks are cut by large thrust faults and smaller-scale normal faults (Baltz and Myers, 1999; Lessard and Bejnar, 1976) with some faults juxtaposing sedimentary and crystalline rock (Fig. 8d).

The predominant lithology of the crystalline basement is gneiss, with minor schist, pegmatitic granite, and basalt. Mineral alteration is greatest directly below the nonconformity. This zone is enriched in sericite within feldspars and clay minerals (mixed with hematite and associated with replacement of micas) (Fig. 8). Where cut by faults the nonconformity-associated phyllosilicates form a matrix that surrounds more rigid grains such as quartz, suggesting that deformation in this unit was accommodated by granular flow, a process associated with high pore-fluid pressure (Paterson, 2012). Microscopic fracturing has occurred within the crystalline basement; these fractures are mineralized with iron oxide, sericite, chert, and calcite. The majority of fractures within the crystalline basement occur along weak grains such as sericitized feldspar and altered mica or cut across quartz and feldspar grains. Authigenic calcite is rare within the crystalline basement, though it commonly occurs as coarsely crystalline calcite cement within grain fractures in feldspar and sericitized feldspar.

Where faults cut the altered crystalline basement locally, cataclasites are found throughout the fault core. Where faulted, the sedimentary rock damage zone includes large twinned calcite grains in fracture-filling cements and cataclasites that lie along the edges of the calcite veins. The cataclasites include pulverized quartz and feldspar grains, chert, pulverized protolith, and clay- and iron-oxide-rich minerals. Quantitative microprobe analyses of the carbonate and fine-grained matrix composition within the sedimentary and basement fault cores reveal that all calcite vein elemental values have a slightly more reduced level of Fe and Mg substitution for Ca than the calcite matrix. The fine-grained matrix within the sedimentary fault core is nearly pure silica, whereas the

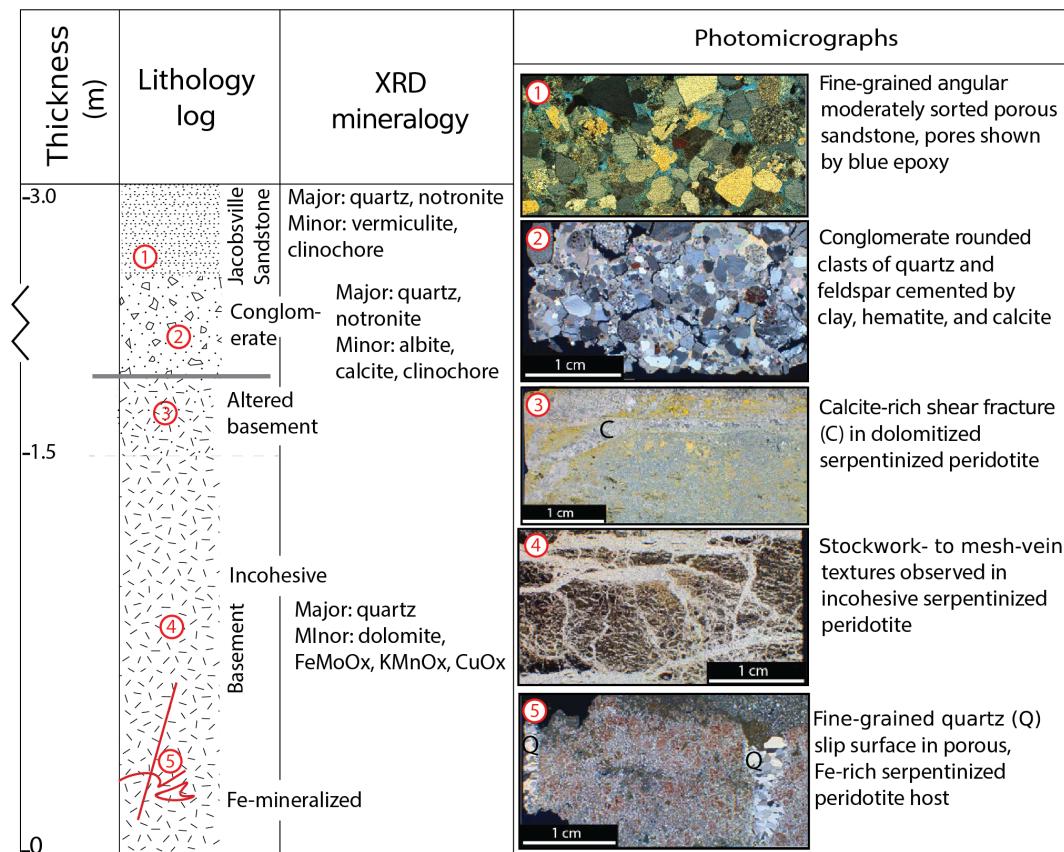


Figure 7. Petrographic summary figure, photomicrographs, and X-ray diffraction results of nonconformity units studied at Presque Isle, Michigan. (a) Jacobsburg Sandstone arenite ($100\times$, ppl), (b) Jacobsburg Sandstone altered conglomerate ($200\times$, ppl), (c) basement calcite-rich slip surface in dolomitized, serpentined peridotite ($200\times$, xpl), (d) basement serpentined peridotite ($100\times$, ppl), (e) basement slip surface within Fe-rich serpentined peridotite (red) ($100\times$, ppl).

fine-grained matrix within the crystalline basement fault core is aluminum-rich (Fig. 8).

3.2.3 R.C. Taylor 1 Core, Nebraska

Core from the R.C. Taylor 1 wildcat well was obtained in 1953 in south-central Nebraska (Table S1 in the Supplement). We examined 19.2 m of core recovered over the Cambrian La Motte Formation sandstone and sheared Proterozoic granitoids in the Central Plains Orogen of the Yavapai province (Marshak et al., 2017; Sims, 1985; Whitmeyer and Karlstrom, 2007). The arkosic La Motte Formation, regionally called the Reagan and Sawatch sandstones, is a fine-grained, well-sorted glauconitic sandstone (Fig. 6).

The basal La Motte Formation and uppermost basement are cut by quartz, calcite, dolomite, and iron-oxide veinlets (Fig. 9). Iron-oxide veins cut quartz veins, and both are cut by calcite veins, providing evidence for fracture reactivation (Fig. 9). Below the La Motte Formation is a phyllosilicate-rich zone composed of a 40 cm thick highly altered basement shear zone that overlies a minimally altered basement shear zone; both are comprised of fine crystalline sericitized

feldspar and chlorite-rich shear zones and overlie the coarse-crystalline, minimally altered granitic basement containing some sericitized feldspar (Fig. 9).

The altered basement shear zone is composed of quartz, feldspar, biotite, chlorite, and dolomite (Fig. 9). Quartz and feldspars are disintegrated; well-developed chlorite, hematite, and magnetite are altered from biotite, and granular disintegration has resulted in clay development. Open pore space occurs between host-rock grains and neo-formed clays. The basement shear zone is characterized by feldspar, quartz, mica, and the alteration minerals chlorite and dolomite (Fig. 9). The shear zones contain chlorite-lined slip surfaces and S-C fabrics within chloritized zones. The shear fabrics are cut by open-mode quartz, sparry calcite, iron-oxide, and dolomite veins. The basal, moderately altered basement unit is a coarse-crystalline granite composed of feldspar, quartz, biotite, and hornblende (Fig. 9). Chlorite is present and associated with minor shear fabrics. Open-mode calcite, dolomite, and quartz veins parallel and cross-cut the chlorite-rich shear fabrics and cut quartz and feldspar crystals (Fig. 9). In the coarse-crystalline granite, altered feldspars contain

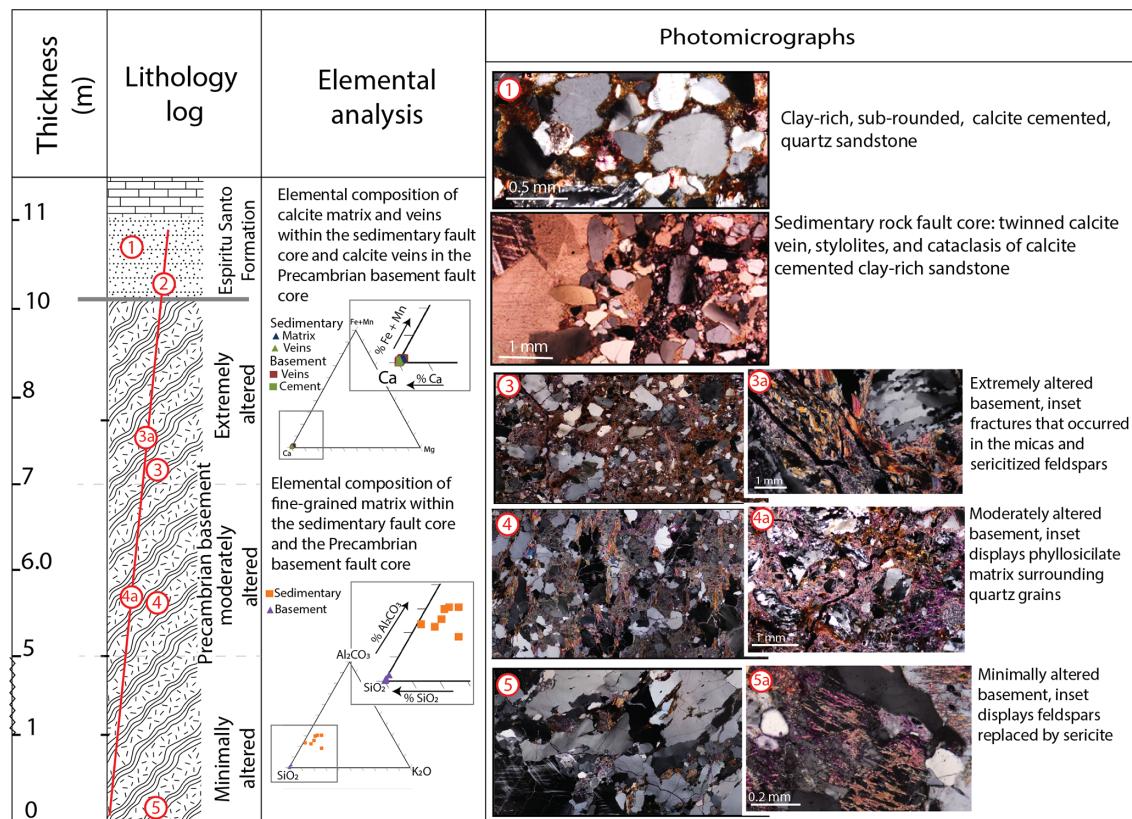


Figure 8. Petrographic and elemental analysis summary of nonconformity units at the Gallinas Canyon site. Thickness is measured in meters from base of outcrop section; the red line represents a fault. Elemental analysis shows a similar calcite composition of the veins within the sedimentary sequence and the Precambrian basement faults. (a, b) The Espiritu Santo sandstones are clay-rich calcite-cemented quartz sandstones; in the fault core, (b) the sandstones are cut by twinned calcite veins and stylolitic textures and contact cataclastic. (c) Adjacent to the nonconformity the granitic basement contain fractures in the micas and sericitized feldspars. (d, e) Basement alteration decreases away from the nonconformity, with a phyllosilicate matrix surrounding quartz grains and sericitization of feldspars occurring 10 m from the nonconformity.

sericite that has formed adjacent to twin planes. Open-mode fractures mineralized with dolomite, calcite, and hematite occur in the lower 7 m of the La Motte Formation and are observed through the underlying granitic shear zone covering 12.5 m of core.

3.2.4 CPC BD-139 Core, Michigan

The CPC BD-139 core, obtained in 1964 for the design of a brine disposal well, samples the contact between the Cambrian Mount Simon Sandstone and Precambrian altered granitoid gneiss of the Grenville Front Tectonic Zone (Table S1 in the Supplement). We divide the CPC BD-139 core into three lithologic units: a laminated sandstone, a finely foliated gneiss, and a gneiss with subhorizontal white veins.

Sandstone grains are rounded to sub-rounded and moderately to well-sorted. The Cambrian Mt. Simon Sandstone in the Michigan Basin is characterized as a porous (5%–15% pore space) arenite to sub-arkosic sandstone (Leeper, 2012). Permeability in the basal Mt. Simon Sandstone is reported

to be between 1×10^{-16} and $1 \times 10^{-12} \text{ m}^2$ (Frailey et al., 2011).

A discrete boundary separates the Mount Simon Sandstone from the underlying altered granitoid gneiss (Fig. 10). The uppermost 30 cm of the basement is composed of a tan, fine-grained dolomite horizon that grades into a dark green foliated gneiss cut by pink subvertical fractures over a span of $\sim 5 \text{ cm}$ (Fig. 10). The basal meter of the Mount Simon Sandstone is a tan, finely laminated arenite with minor amounts of iron-rich clay. The quartz-feldspathic granitoid gneiss near the contact contains the following alteration products: zeolites, vermiculite, Fe- and Mn-oxides, and carbonates including dolomite (Fig. 10). Dolomitization of the basement host rock reappears $\sim 2 \text{ m}$ below the nonconformity. The original basement foliation is preserved and is associated with micrometer-scale crystalline dolomite grains, radiating silica crystals, and subhorizontal calcite and dolomite open-mode veins (Fig. 10). Trace amounts of ankerite, clinochlore, and vermiculite are also present in the dolomitized basement rocks (Fig. 10).

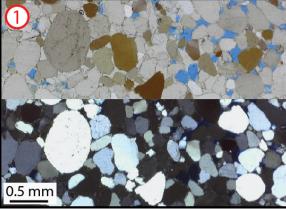
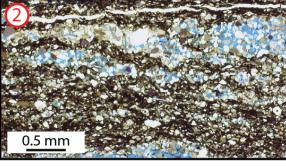
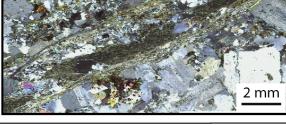
Depth (m/ft)	Graphic log	XRD mineralogy	Photomicrographs
-1201 / 3980	Shear zone	Major: quartz Sample 3984	 Coarse-grained rounded to sub-rounded, poorly sorted quartz-cemented quartz sandstone, 2% pore space PPL top, XPL bottom sample 3987 0.5 mm
-1213 / 3990	La Motte Sandstone	Major: microcline, orthoclase, sanidine Minor: chlorite-serpentine, bokite, glauconite, cameronite, vermiculite Sample 3995	 Medium-grained, poorly sorted, sub-angular to sub-rounded quartz sandstone, clay and oxide cement, calcite vein fill, 5% pore space Sample 3993 0.5 mm
-1225 / 4018	Finely foliated granitic shear zone	Top Precambrian	 Altered basement shear zone, quartz, Fe-oxide and calcite open-mode veins some porosity between ridge grains and neo-formed clays Sample 4031.6 2 mm
-1225.3 / 4020			 Altered basement shear zone, chlorite-lined shear planes, sericitization of k-spar grains Sample 4032 2 mm
-1228.3 / 4032			 Quartz, k-spar, plagioclase, with chlorite-lined shear fractures, sericitization of k-spar grains at twinning lamella Sample 4036 2 mm
-1231.4 / 4040			

Figure 9. Petrographic summary, photomicrographs, and X-ray diffraction results of nonconformity units studied in the R.C. Taylor 1 core; (a) La Motte Formation sandstone, rounded to sub-rounded, poorly sorted quartz sandstones (100×, ppl and xpl); (b) lower La Motte Formation, opaque Fe-oxide cements and vein fill, porosity shown by blue epoxy, (100× ppl); (c) top Precambrian crystalline altered basement shear zone. Syntaxial veins are mineralized with quartz, reactivated, and mineralized with Fe-oxide then sparry calcite. There is some porosity between rigid grains and neo-formed clays (50× PPL); (d) altered basement shear zone, chlorite-lined shear planes, sericitization of feldspars along twining lamella (100× ppl); (e) coarse-crystalline sericitization of feldspars adjacent to twin lamellae (150× XPL).

3.2.5 BO-1 Core, Minnesota

The BO-1 core was originally collected in 1962 as part of an exploratory mining project in Fillmore County, southeast Minnesota (Gilbert, 1962) (Table S1 in the Supplement). This core provides a continuous 300 m section of altered and mineralized rocks of lower Cambrian Mount Simon Sandstone overlying a Precambrian layered intrusive complex of altered and unaltered gabbro, other mafic intrusions, and felsic dikes (Smith et al., 2019; Fig. 11).

Sedimentary sequences in BO-1 extend to ~1.2 km where the nonconformity is marked by an approximately 12 cm zone of pervasive leaching and iron-hydroxide staining (goethite). Intense alteration extends into the basement rocks

for ~21 m, with ~50 m of argillitic and propylitic alteration and/or fracture mineralization observed to ~402 m of depth (Fig. 11). Localized faults as well as hybrid and open-mode fracture surfaces intersect the sampled basement core from within ~1 cm of the nonconformity contact and extend to 475.5 m; fracture density decreases with depth (Fig. 11). Slip surfaces exhibit oblique to dip-slip slickenlines, range from millimeters to centimeters thick, and are either coated in clay or contain mineral infillings (\pm carbonate, \pm silica, \pm chlorite, \pm iron oxides).

The Mount Simon Sandstone contains a ~0.5 m zone of intense iron-hydroxide (goethite) alteration at the nonconformity (Fig. 11). This iron-hydroxide oxidized zone extends for several meters into the slightly altered and meta-

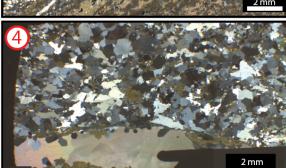
Depth (m)	Graphic log	XRD mineralogy	Photomicrographs
- 1404.5	Jacobsville Sandstone	Major: quartz Minor: k-spar	AT CONTACT: Argillaceous dolomitized basement host rock with dolomite veining 
- 1405.4		Major: quartz, dolomite	Sub-angular - sub-rounded qtz/feldspar/clinochlore fabric Fe-alt. of k-spar grains Microfractures perpendicular to fabric 
- 1405.7	Finely foliated gneiss	Major: quartz Minor: k-spar Trace: clinochlore, kaolinite (?), andalusite (?)	Banded qtz/feldspar gneiss Sericitization of k-spar grains 
- 1407.1		Major: quartz Minor: k-spar Trace: dolomite, clinochlore, zeolite, muscovite, dickite (?)	Angular - sub-angular quartz matrix 1-2mm wide carbonate vein w. colloform zeolite rim, euhedral qtz xtals 
- 1411.2	Gneiss with white veins		Fine-grained dolomitized gneiss, fabric intact, stringy quartz and carbonate veins, re-crystallization textures present (radiating, scalloped silica growth) 
1412.1			

Figure 10. Petrographic summary, depth in meters measured, photomicrographs, and X-ray diffraction results of nonconformity units studied in the CPC BD-139 core. (a) The basement sample at the contact is an argillaceous dolomitized gneiss with dolomite veins (XPL); (b) foliation defined by quartz–feldspar–clinochlore fabric with iron alteration of potassium feldspar grains (XPL); (c) banded quartz–feldspar gneiss with common sericitization of potassium feldspar grains (XPL); (d) carbonate vein with colloform zeolite rim and euhedral quartz crystals (XPL); (e) dolomitized gneiss with common quartz and carbonate veins (XPL).

morphosed crystalline basement rock. From petrographic and XRD analyses, we identify mineralogical assemblages (dolomite, siderite, iron oxides, iron hydroxides, illite, smectite, kaolinite–serpentinite, vermiculite) and textures that are indicative of weathering, diagenesis, and multiple episodes of fluid–rock interactions coupled with deformation within the broad ~ 50 m zone of intense alteration also marked by abundant structural discontinuities (fractures, faults, veins) across the nonconformity zone (Fig. 11).

Measured gas permeability values are highest above the nonconformity within the porous Mount Simon sedimentary reservoir (up to 1000 mD) and vary significantly from 0 to 500 mD below the nonconformity contact. Locally permeability increases in direct correlation with the presence of structural discontinuities (Fig. 6).

3.3 Hydrogeologic models

The first model (Fig. 12a) is a Type 0 nonconformity, represented by a sharp contact between the basement and overlying injection reservoir. The second model simula-

tion, a Type I nonconformity, includes a 20 m thick, low-permeability ($k_x = k_z = 3 \times 10^{-18} \text{ m}^2$) zone (Fig. 12b); this layer is 1 order of magnitude less permeable than the basement host rock and a further 1 order of magnitude less permeable in the fault core. The continuous low-permeability zone reduces the permeability of the basement fault damage zone by 4 orders of magnitude, making the fault damage zone nonconductive. Pressure does not propagate into the crystalline basement although there was some diffusion of the 2 m excess hydraulic head front to depths ≤ 500 m. In the third simulation, a discontinuous low-permeability zone is present (Fig. 12c). Where this zone is absent, the pressure front propagates into the basement along the fault damage zone to a depth of 2.5 km. The fault zone was not blocked by the low-permeability zone, and elevated pore pressures propagated downward to depths of 2.5 km via the fault zone (Fig. 12c). Elevated fluid pressures likewise appeared to be forced down in other areas where the low-permeability zone pinches out, such as towards the right-hand side of Fig. 12c.

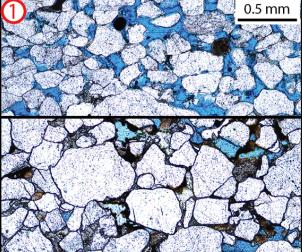
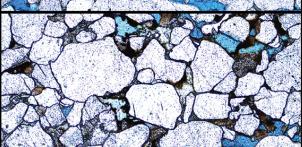
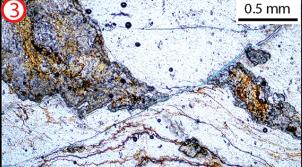
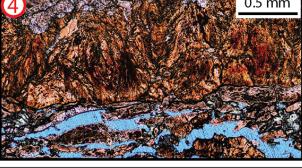
Depth (m/ft)	Graphic log	XRD mineralogy	Photomicrographs
-365.8/1200		Mount Simon Sandstone Major: quartz Minor: Fe-oxides Major: quartz, Fe-oxides Minor: Fe-hydroxide (goethite)	 Fine- to medium-grained, subangular to subrounded, well-sorted porous sandstone  Iron-coated and irregularly cemented, siderite crystals partially fill pore space Sample: 1259.9-1260.18
-384/1260.3	① ② ③	Nonconformity contact Major: Fe-hydroxide (goethite), Mn-oxides, Fe-oxides, magnesio-chloritoid (ferroan) Minor: ankerite, dolomite	 Altered meta-gabbro-norite with dolomitization Sample: 1260.3-1260.52
-394.4/1294	④	HOST Major: dolomite, Ti-oxide (anatase) Minor: ramsbeckite	 Argillization along discrete discontinuities Multiple centimeter- and micrometer-scale slip surfaces Sample: 1260.52-1261.25
-396.2/1300	⑤	SLIP Major: dolomite, birnessite, Mo-oxide, quartz, nacrite clay (kaolinite-serpentine)	 Argillization extends outward from fracture surfaces Open fractures Sample: 1293.9-1294.3
-398/1305.9	⑥	Major: Fe-oxides, Cu-hydroxides (clinoclase), muscovite, illite	 Multi-layered vein, slip surface dilatational pull-apart microstructures crack-seal textures serpentinization, dolomitization Sample: 1305.85-1305.95
-406.1/1332	⑦	HOST Major: siderite, dolomite, adularia minor: Fe-rich smectite (nontronite)	 Multi-layered shear zone of multi-phase composition Anastomosing and cross-cutting vein textures Sample: 1332.3-1332.5
-426.7/1400		HOST Major: diopside, clinochlore (Fe-rich) Minor: Mn-oxides, andesine, ramsbeckite	 Moderately altered feldspars Ultracataclasite along mm-scale shear zones Sample: 1472.3-1473
-448.8/1472		SLIP Major: quartz, Fe-oxides Minor: fluorapatite, zeolites, Mn-oxides	
-457.2/1500		Pegmatite Major: albite (Ca), anorthite, Minor: calcite, ferro-actinolite, vermiculite, Na-Ca plagioclase (andesine)	
		Metadiorite	

Figure 11. X-ray diffraction mineralogy and photomicrographs of BO-1 drill core samples showing representative compositions and textures across the nonconformity interface contact. Samples within centimeters of the contact (1–3) are strongly weathered, altered, and slightly metamorphosed gabbro-norite. Alteration and diagenesis assemblages include iron-oxide hydroxides with chlorite, ankeriteite, and dolomite. Alteration extends for ~50 m into the basement. Sample 3 illustrates millimeter-scale offset across the argillite layer. Note the fracture permeability (blue epoxy) parallel to the slip surface. Fracture surfaces within Sample 4 at 34 m below the contact are several millimeters of mixed chlorite-clay alteration and fine-scale permeability (blue epoxy). Sample 5 at 1304.9 m shows multiple phases of fluid-rock interactions coupled with dilation, serpentinization, and dolomitization. There is a multi-layered clay-rich fault core gouge within Sample 6 at 1332 m or ~70 m below the contact. Note the open fractures within the central portion of the fault core gouge (blue epoxy). At 1472 or 272 m below the contact within the meta-grano-diorite unit, clay alteration is observed within feldspar grains at the microscale.

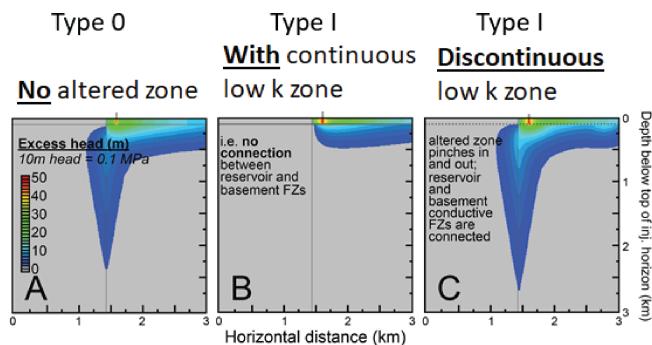


Figure 12. Cross-sectional views of pore pressure envelope propagation resulting from injection into a reservoir underlain by a low-permeability altered zone. Excess hydraulic heads after 4 years of constant-rate injection are presented for a Paleozoic conduit-barrier fault scenario. **(a)** Faulted Type 0 nonconformity with no low-permeability altered zone; **(b)** Type I nonconformity with an altered zone present as a 20 m thick confining layer (represented by two horizontal grey lines) that is continuous such that the reservoir and basement fault zones do not connect; **(c)** Type I nonconformity with a discontinuous altered zone that pinches in and out in 20 m horizontal intervals (i.e., undulating) but with the reservoir and basement fault zones fully connected. Results are zoomed in to the top 3 km × 3 km of the model domain. Vertical grey lines indicate the location of the fault zone. The injection well location is indicated on top of each panel. Transition from grey to dark blue (and all subsequent contour lines) denotes a 2 m increase in hydraulic head (0.02 MPa). Adapted from Ortiz (2017; for details of the modeling approach, see Ortiz et al., 2019).

4 Discussion

The nonconformities examined in this study range from sharp contacts to zones several meters thick and exhibit a range of mineralized textures and structural discontinuities (Table S2 in the Supplement). We observe mineralogic alterations across the nonconformity that are expected to impact diffusivity and storativity, and the sites evaluated provide geological and hydrogeological analogs that aid in understanding the impact circulating fluids may have on altering rock properties at depth (Oliver et al., 2006). Based on observations made in this work, we divide nonconformities into three end-member types (Table 1): Type 0 – a sharp contact between sedimentary strata and basement rocks; Type I – an interface dominated by phyllosilicates; and Type II – an interface dominated by non-phyllosilicate secondary mineralization (Table 1). All the nonconformity types observed in this study are cut by structural discontinuities; therefore, several possible contact subtypes exist within these three proposed end-members (Fig. 13). Based on our observations, structural and mineralogical heterogeneities at the sedimentary–crystalline rock nonconformity are thought to control the degree to which fluids, fluid pressure, and associated poroelastic stresses are transmitted over long distances across and along the nonconformity boundary. The structural elements

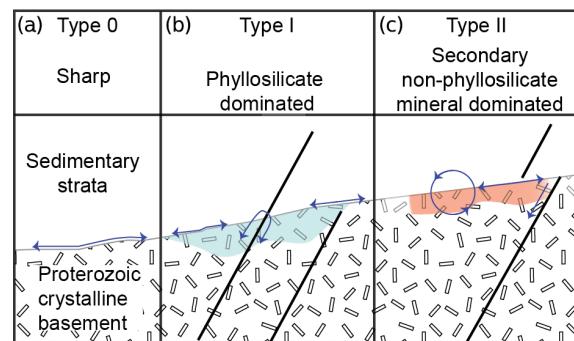


Figure 13. Proposed geologic schematics of the nonconformity contact region. **(a)** Type 0 – sharp contact expected to prevent direct fluid pressure communication across the contact while promoting migration parallel to the contact distributing fluids laterally. **(b)** Type I – phyllosilicate-dominated zone above the crystalline basement is expected to inhibit fracture propagation across the nonconformity, prevent fluid migration due to permeability contrast, and promote lateral migration; downward fluid migration can occur at a permeable fault zone. **(c)** Type II – secondary-mineralization-dominated zone with lateral migration due to permeability contrast; mineralization due to fluid–rock interactions suggests that deep fluid circulation occurs even without enhanced permeability from fractures. All nonconformity types may be cut by structural discontinuities. Blue arrows indicate potential flow paths of injected fluids.

and fluid-related alteration patterns observed in these analog sites support the hypothesis that the nonconformity interface zone influences or controls the potential for cross-contact fluid flow and distribution of fluids within the crust.

Our collective field and core observations in various basement tectonic settings document the occurrence of significant variations in altered or mineralized zones that lead to contrasts in permeability across the nonconformity. Where present, the structural discontinuities include small-offset faults, shear fractures, and veins. In thin section we note evidence for dissolution, recrystallization, new mineral growth, and veins that reflect mineralization or deformation at depth and are not the result of alteration due to weathering alone. Crack–seal textures and calcite twinning lamella suggest vein mineralization at depth (Burkhard, 1993), and the reactivation of preexisting fractures documents episodic fracture growth (Davatzes and Hickman, 2005; Laubach et al., 2004).

At a Type 0 nonconformity, the nonconformity zone is expected to prevent direct fluid pressure communication across the contact due to a significant contrast in rock permeabilities that would hinder cross-contact fluid migration while promoting migration parallel to the contact, distributing fluids laterally away from the injection site (Fig. 13a). At a Type I nonconformity, a phyllosilicate-dominated contact is expected to inhibit fracture propagation across the nonconformity (Ferrill et al., 2012; Larsen et al., 2010; Schöpfer et al., 2006) and therefore maintain a significant permeability contrast, preventing direct fluid migration. In such cases,

Table 1. Summary of nonconformity types and associated features.

Nonconformity type	Site location	Summary features
Type O	Hidden Beach; Marquette, Michigan	Sharp, discrete nonconformity contact ± topographic variations between porous sedimentary sequences and nonporous crystalline basement
Type I	Gallinas Canyon, New Mexico R.C. Taylor core	Altered contact with phyllosilicate mineralization; faults within basement only and/or cross-cutting nonconformity contact into overlying sedimentary rocks, open-mode veins in basement, and overlying sedimentary rocks
Type II	Presque Isle; Marquette, Michigan CPC BD-139 core BO-1 core	Altered contact with non-phyllosilicate secondary mineralization and hydrothermal mineralization; faults within basement only and/or cross-cutting nonconformity contact into overlying sedimentary rocks; open-mode and hybrid veins in basement and overlying sedimentary rocks

nonconformities result in a poor hydrologic connection between the sedimentary section and deeper basement rocks (Fig. 13b).

However, repeated brittle failure and mineralization, observed in Type I nonconformities, suggest that phyllosilicate-dominated shear zones can act as a zone of mechanical weakness that can be reactivated, allowing for the development of fracture permeability. In this fractured nonconformity we observed alteration as deep as 5 m below the nonconformity in the crystalline rocks examined; however, previous work highlights the potential for fractures and connectivity to basement fault zones at much greater depths (Duffin et al., 1989). Preexisting basement shear zones that are reactivated may allow future fluid circulation during injection scenarios.

Type II nonconformities (Table 1, Fig. 13c, Table S2) are mineralized contacts that include secondary alteration minerals found within 10 cm to several meters below the nonconformity. The mineralization due to fluid–rock interactions at Type II nonconformities suggests that deep fluid circulation occurs even without enhanced permeability from fractures (Cuccio, 2017) (Fig. 12c). This nonconformity type could prevent brittle deformation but may be more influenced by poroelastic loads. The impact of these contacts on hydrogeologic properties is not yet well-understood or modeled.

The impact the morphology of the nonconformity has on the downward propagation of fluid pressures into the crystalline basement has been shown by several numerical hydrogeologic studies (Ortiz et al., 2019; Segall and Lu, 2015; Yehya et al., 2018; Zhang et al., 2016). Models suggest that direct pore-fluid pressure communication (Ortiz et al., 2019; Segall and Lu, 2015; Yehya et al., 2018) and significant changes in poroelastic stress (Goebel and Brodsky, 2018; Zhang et al., 2016) can occur well way from the injection zones. Numerical models predict that nonconformities with throughgoing fractures distribute fluid deeper into the base-

ment rocks and that direct pore pressure communication can destabilize faults at depth (Ortiz et al., 2019; Segall and Lu, 2015; Yehya et al., 2018). All the nonconformity types observed here are cut by structural discontinuities, and several possible contact subtypes exist within these three proposed end-member scenarios (Fig. 12). Fractures, especially fault zones, are expected to distribute fluids and propagate fluid pressures to a greater depth regardless of nonconformity type (Yehya et al., 2018). Because nonconformity interface zones with preexisting deformation fabrics may be preferential flow pathways that distribute fluid pressure away from the injection zone, high-permeability damage zones transmit fluid pressure to greater depths than non-conduit fault zones (Yehya et al., 2018).

To illustrate the effects of reduced permeability above the nonconformity and the impact of permeable fault zones on fluid migration, we compare three models of basal reservoir injection that consider continuous and discontinuous zones of altered low-permeability rocks above the basement (Fig. 12). The Type 0 nonconformity, represented by a sharp contact between the basement and overlying injection reservoir, results in lateral migration away from the injection well and downward migration whereby it encounters a fault zone (Fig. 13a). In the second model simulation, a Type I nonconformity, the presence of a 20 m thick, low-permeability zone and no connection between basement and sedimentary fault zones result in lateral migration and pressure does not propagate into the crystalline basement (Fig. 12b). The third simulation models a Type I nonconformity with a discontinuous low-permeability zone; where this zone is absent, the pressure front propagates into the basement along the connected fault damage zone to a depth of 2.5 km and elevated fluid pressures appear to be forced downward where the low-permeability zone pinches out (Fig. 13c).

Our collective field and core observations document the occurrence of significant variations in altered or mineralized zones, which would impact permeability values associated with the nonconformity zone, and indicate that alteration coupled with abundant structural discontinuities can result in relatively higher permeability that extends for tens of meters into the crystalline basement rock below the nonconformity.

5 Conclusions

We define key rock types and structural elements of the nonconformity zone and split the analog nonconformities into three end-member types. The three nonconformity end-member types provide a broad hierarchy of nonconformities in the midcontinent (Table S1) and are observed at nonconformity sites elsewhere. We expect these nonconformity types to either distribute fluid pressure away from the injection point or provide direct communication with basement rocks, moving fluids to a greater depth across the nonconformity. We observe that fractures cut all nonconformity types and expect in these cases that changes in fluid pressure or poroelastic loads could result in triggered earthquakes within basement rocks (Chang and Segall, 2016; Zhang et al., 2013). Numerical modeling of Type 0 and Type I end-members that include fault zones predicts downward propagation of fluid pressure and changes to poroelastic loads. The data presented here can be used to improve model inputs for evaluation of cross-contact fluid and pressure communication, whether through the creation or modification of existing permeability, poroelastic pathways, or rheological changes associated with fluid–rock interactions. We show that conditions along the nonconformity zone vary, and the data from outcrop and core observations also suggest that injection of brines at depth may drive mineralogical alteration and potential fault zone weakening; these data can also be used to understand the impact that long-term storage of chemically reactive fluids has on rock properties (Callahan et al., 2020). Once fluids penetrate the basement, flow is likely controlled by fracture and fault systems, and the reactivation of preexisting structures is possible. However, microporosity within basement rocks may enhance mineralogical changes over the long term and transmit fluids deeper in the basement while promoting short-term lateral migration along the nonconformity.

Our observations illustrate that the contact should not be treated as an impermeable barrier to fluid flow nor as one cut by faults of various permeabilities but should instead be evaluated on a site-by-site basis prior to injection of large fluid volumes.

Data availability. Underlying and ongoing research data related to this paper are listed in the Supplement in Table S1. Access to underlying data associated with student theses can be found through university websites: <https://digitalcommons.usu.edu/etd/>

6889/ (last access: 17 September 2020, Cuccio, 2017) and <https://digitalcommons.usu.edu/etd/7497/> (last access: 17 September 2020, Hesseltine, 2019).

Supplement. The supplement related to this article is available online at: <https://doi.org/10.5194/se-11-1803-2020-supplement>.

Author contributions. This work was collaborative, and author contributions include but are not limited to the following. ESP collected data and characterized the R.C. Taylor 1 core (Nebraska), conceived and designed the analysis, provided funding support for sample collection and analysis, and wrote the paper. KKB collected field data at the Presque Isle outcrop, conducted data analysis, provided funding support for sample collection and analysis, and wrote the paper. LC collected field data at the Presque Isle outcrop and the CPC BD-139 core (Michigan), conducted data analysis, and provided funding support for sample collection. KS collected data from the BO-1 core (Minnesota), conducted data analysis, and provided funding support for sample collection. JPE conceived and designed the analysis, contributed to data analysis and data collection for the CPC BD-139 core (Michigan), provided funding support for sample collection and analysis, and wrote the paper. JPO conceived and designed the analysis as well as hydrogeologic modeling based on outcrop and core observations. KK collected and characterized data at the Gallinas Canyon, New Mexico, outcrop. MP contributed analysis tools, conceived and designed the analysis, and conducted hydrogeologic modeling. PM conceived and designed the analysis for Gallinas Canyon, conducted analysis of mineral assemblages and data, and provided funding support for sample collection and analysis.

Competing interests. The authors declare that they have no conflict of interest.

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