

Supporting Information for “VESIcal Part I: An open-source thermodynamic model engine for mixed volatile ($\text{H}_2\text{O}-\text{CO}_2$) solubility in silicate melts”

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Introduction This supplement contains text, figures, tables, and jupyter notebooks that serve to both enhance the user experience in VESIcal (e.g., with additional worked ex-

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amples of how the code can be used) and give a complete and thorough background as to the thermodynamic calculations and methodologies employed both in VESIcal and in the models contained therein. Captions for the datasets and jupyter notebooks (all of which are uploaded as separate files) are given below.

Captions for Datasets S1 to S7

Seven Microsoft Excel files containing compositional data on various silicate melts are included. These files are as follows:

1. example_data.xlsx: Silicate melts from basalts to rhyolites. These data are used as the example data in the main manuscript.
2. alkaline.xlsx: Alkaline experimental samples from Iacovino et al. (2016), ranging from basanite to phonotephrite
3. basalts.xlsx: Basalts from Tucker et al. (2019) and Bennett et al. (2019)
4. cerro_negro.xlsx: Basaltic melt inclusions from Cerro Negro volcano from Roggensack (2001)
5. rhyolites.xlsx: Rhyolites from Myers et al. (2019) and Mercer et al. (2015)
6. liu_2005.xlsx: Calibration dataset of Liu et al. (2005); bulk composition data only
7. moore_1998.xlsx: Calibration dataset of Moore et al. (1998)

Captions for Jupyter Notebooks S1 to S10

Ten Jupyter Notebooks are included with this manuscript. These notebooks are given along with all necessary files to run them plus a PDF copy of the evaluated notebooks such that the code therein can be accessed without needing to run the notebook itself.

Each notebook and accessory files are located within a folder following the naming scheme described here:

1. Allison: For testing VESIcal's reproduction of the Allison et al. (2019) model

- S1_Testing_Allison_et_al_2019.ipynb
- S1_Testing_Allison_et_al_2019.pdf
- S1_Testing_Allison_et_al_2019.xlsx

2. Dixon: For testing VESIcal's reproduction of the Dixon (1997) model (VolatileCalc)

- S2_Testing_Dixon_1997_VolatileCalc.ipynb
- S2_Testing_Dixon_1997_VolatileCalc.pdf
- S2_Testing_Dixon_1997_VolatileCalc.xlsx
- VolatileCalc_Test1.png
- VolatileCalc_Test2.png

3. Iacono-Marziano: For testing VESIcal's reproduction of the Iacono-Marziano (2012) model

- S3_Testing_Iacono-Marziano_et_al_2012.ipynb
- S3_Testing_Iacono-Marziano_et_al_2012.pdf
- S3_Testing_Iacono-Marziano_et_al_2012.xlsx

4. Liu: For testing VESIcal's reproduction of the Liu et al. (2005) model

- S4_Testing_Liu_et_al_2005.ipynb
- S4_Testing_Liu_et_al_2005.pdf
- S4_Testing_Liu_et_al_2005.xlsx

- Testing_Liu_img1.png

- Testing_Liu_img2.png

- Testing_Liu_img3.png

5. MagmaSat: For testing VESIcal's reproduction of MagmaSat (Ghiorso and Gualda, 2015) model.

- S5_Testing_Magmasat.ipynb

- S5_Testing_Magmasat.pdf

- S5_Testing_Magmasat.xlsx

- Testing_MagmaSat_img1.png

6. Moore: For testing VESIcal's reproduction of the Moore et al. (1998) model

- S6_Testing_Moore_et_al_1998.ipynb

- S6_Testing_Moore_et_al_1998.pdf

- S6_Testing_Moore_et_al_1998.xlsx

7. Shishkina: For testing VESIcal's reproduction of the Shishkina et al. (2014) model

- S7_Testing_Shishkina_et_al_2014.ipynb

- S7_Testing_Shishkina_et_al_2014.pdf

- S7_Testing_Shishkina_et_al_2014.xlsx

- Testing_Shishkina_img1.png

8. Cerro_Negro_Calculations: For running calculations discussed in Section 4.1 of the manuscript

- S8_CerroNegro_isobar_comparison.ipynb

- S8_CerroNegro_isobar_comparison.pdf
- cerro_negro_satP_compare.xlsx

9. Isobar_Comparison_Plotter: For running calculations discussed in Section 4.2 of the manuscript

- S9_Isobar_Comparison_Plotter.ipynb
- S9_Isobar_Comparison_Plotter.pdf

10. Model_Hybridization: Notebook illustrating how to hybridize models, as discussed in Section 3.9 of the manuscript

- S10_Model_Hybridization.ipynb
- S10_Model_Hybridization.pdf

Text S1. VESIcal's verbose installation instructions

Below are instructions for installing the necessary prerequisite libraries and VESIcal on your computer in order to run the VESIcal code locally.

Step 1: Install dependencies. Dependencies are prerequisite files or libraries that you must have installed on your computer for VESIcal to run properly. VESIcal's dependencies are:

1. Python3 (<https://www.python.org/downloads/>)
 - (i) Python pandas (<https://pandas.pydata.org/>)
 - (ii) Numpy (<https://numpy.org/>)
 - (iii) SciPy (<https://www.scipy.org/>)
 - (iv) Matplotlib (<https://matplotlib.org/>)
2. ENKI thermoengine (<https://gitlab.com/ENKI-portal/ThermoEngine>)

Install Python3. First, determine if Python3 (or some other version of Python) is installed on your computer. On Mac or Linux systems, open the application called Terminal and type the following, which will start running python if you have it installed: `python`

If python is installed, you will see a message telling you which python is installed. If this message says “Python 3.x.x” where x can be any number, then you have Python3 already installed. If the message also says “Anaconda, Inc.” then that’s even better – you have Anaconda’s version of Python3, which means you have all of the basic dependencies already. To quit python, type: `exit()`

If python is not installed, or if you see a different python version, we suggest installing Python3 via Anaconda3 (<https://www.anaconda.com/products/individual>). Anaconda is Python plus all of the most commonly used libraries, many of which are required to run VESIcal. Pandas, Numpy, SciPy, and Matplotlib all come prepackaged in Anaconda.

Install ENKI thermoengine. Next, ENKI’s thermoengine (the python implementation of MELTS) must be installed. Because installation instructions for thermoengine may change, we will not replicate them here. Instead we refer you to the ENKI portal at <https://gitlab.com/ENKI-portal/ThermoEngine>. From there, download the thermo-engine files and follow the installation instructions.

Step 2: Install VESIcal. Now that all dependencies are installed, you can install VESIcal in your terminal by typing: `pip install VESIcal`

Step 3: Use VESIcal. You are now ready to begin coding with VESIcal. At the beginning of your python script, type “`import VESIcal as v`” to use VESIcal in the same way shown in this manuscript.

Text S2. Text S2. Generic methods for calculating mixed-fluid properties

VESIcal provides a set of methods for calculating the properties of mixed CO₂-H₂O fluids, which can be used with any combination of H₂O and CO₂ solubility model. Using generic methods allows additional models to be added to VESIcal by defining only the (simpler) expressions describing pure fluid solubility. Non-ideality of mixing in the fluid or magma phases can be incorporated by specifying activity and fugacity models. We provide a summary here of how we approach calculating the properties of mixed-fluid systems.

S2.1 Finding the saturation pressures of pure fluids

Whilst the variables required for each pure-fluid solubility model are different, we take a very similar approach to calculating fluid properties for each. Most pure-fluid models provide analytical expressions for the concentration of a volatile species dissolved in a magma as a function of either its partial-pressure or fugacity. However, in most cases, an analytical expression does not exist for the inverse calculation (calculating the saturation pressure given the concentration of dissolved volatile species), it must be solved numerically.

Numerical methods obtain the answer by making guesses for the saturation pressure, checking whether the predicted concentration of the dissolved volatile species matches the measured value, and refining the guesses as appropriate. When the saturation pressure for a dissolved volatile concentration $[V]^L$ is found, this equality holds:

$$[V]_{measured}^L - [V]_{predicted}^L(P, \dots) = 0 \quad (1)$$

An advantage of using python is the ease with which other libraries can be incorporated, for example the `scipy` library (Virtanen et al., 2020), which contains optimized algorithms for finding solutions to problems like this. To solve for pure-fluid saturation pressures we employ the `scipy.optimize.root_scalar()` method.

S2.2 Finding saturation pressure of mixed fluids Magmas in equilibrium with mixed-fluids can be modelled using a combination of pure-fluid CO₂ and H₂O models, provided the mole-fraction of each species ($X_{\text{CO}_2}^F$ and $X_{\text{H}_2\text{O}}^F$) in the vapour phase is specified. In an ideal gas the partial pressure of the vapour phase is given by:

$$P_{\text{CO}_2} = X_{\text{CO}_2}^F \times P \quad (2)$$

For non-ideal fluids, the H₂O and CO₂ fugacities are a function of $X_{\text{CO}_2}^F$ and $X_{\text{H}_2\text{O}}^F$, but will tend to zero as $X_{\text{CO}_2}^F$ and $X_{\text{H}_2\text{O}}^F$ tend to zero.

When calculating the saturation pressure of a mixed CO₂-H₂O fluid, we have an additional unknown: $X_{\text{H}_2\text{O}}^F$ ($X_{\text{CO}_2}^F$ is given by $[1 - X_{\text{H}_2\text{O}}^F]$). Like for the pure-fluid case described above, the problem must be solved numerically, but for two parameters simultaneously. The solution can be expressed as the vector equality:

$$\begin{pmatrix} [\text{H}_2\text{O}]_{\text{measured}}^L \\ [\text{CO}_2]_{\text{measured}}^L \end{pmatrix} - \begin{pmatrix} [\text{H}_2\text{O}]_{\text{predicted}}^L(P, X_{\text{H}_2\text{O}}, \dots) \\ [\text{CO}_2]_{\text{predicted}}^L(P, X_{\text{H}_2\text{O}}, \dots) \end{pmatrix} = \begin{pmatrix} 0 \\ 0 \end{pmatrix} \quad (3)$$

VESIcal then employs the `scipy.optimize.root()` method to obtain the value of the vector $\begin{pmatrix} P \\ X_{\text{H}_2\text{O}} \end{pmatrix}$ for which the above equality holds.

S2.3 Finding the equilibrium fluid composition

When calculating the dissolved volatile concentrations in a progressively degassing magma, we break the calculation into steps, each with a small decrement in pressure. As pressure decreases, the mass of fluid in equilibrium with the magma increases, and its com-

position will change. It follows that at each stage we must solve for the mass and composition of the fluid in equilibrium with the magma, given the total mass of volatiles in the system. This calculation is implemented in MixedFluid.calculate_equilibrium_fluid_comp(), which in turn is called by MixedFluid.calculate_degassing_path() for each increment of degassing magma. Here we summarize the methodology for the calculation.

Mass balance in the system can be expressed in terms of the mole fractions of H₂O (X_{H_2O}) and CO₂ (X_{CO_2}) in the silicate-liquid (L), the H₂O-CO₂O fluid (F), and the system total (T), and the fraction of fluid present (f).

$$X_{CO_2}^F + X_{H_2O}^F = 1 \quad (4)$$

$$(1 - f)X_{CO_2}^L + fX_{CO_2}^F = X_{CO_2}^T \quad (5)$$

$$(1 - f)X_{H_2O}^L + fX_{H_2O}^F = X_{H_2O}^T \quad (6)$$

The fraction of fluid (f) can be eliminated by substitution, obtaining Eqn (8) of Papale et al. (1999):

$$\frac{X_{CO_2}^T - X_{CO_2}^L}{X_{CO_2}^F - X_{CO_2}^L} - \frac{X_{H_2O}^T - X_{H_2O}^L}{X_{H_2O}^F - X_{H_2O}^L} = 0 \quad (7)$$

In principle, this equation allows us to numerically solve for mass balance; however, the equation is valid only for both $X_{CO_2}^F \neq X_{CO_2}^L$ and $X_{H_2O}^F \neq X_{H_2O}^L$. When $X_{CO_2}^F \rightarrow X_{CO_2}^L$ or $X_{H_2O}^F \rightarrow X_{H_2O}^L$ the equation tends to $\pm\infty$, which the `scipy.optimize.root_scalar()` routine will mistake for a root of the equation. The regions over which the expression is not infinite can be found; however we find a more reliable approach when using the `scipy.optimize` library is to use the equation in a different form.

Rearranging the CO₂ mass-balance expression allows the fraction of fluid (f) to be calculated:

$$f = \frac{X_{CO_2}^T - X_{CO_2}^L}{X_{CO_2}^F - X_{CO_2}^L} \quad (8)$$

This is valid for $X_{CO_2}^F \neq X_{CO_2}^L$. It is much more unlikely that a solution will exist near $X_{CO_2}^F = X_{CO_2}^L$ than near $X_{H_2O}^F = X_{H_2O}^L$. By not directly eliminating f by dividing through by both $(X_{CO_2}^F - X_{CO_2}^L)$ and $(X_{H_2O}^F - X_{H_2O}^L)$ we significantly increase the reliability of using this expression with `scipy.optimize.root_scalar()`.

In this expression, the only known value is $X_{CO_2}^T$. However, if a guess is provided for $X_{CO_2}^F$, the partial pressure of CO₂ can be calculated and the corresponding value for $X_{CO_2}^L$ obtained. Since $X_{H_2O}^F = 1 - X_{CO_2}^F$, when we provide a guess for $X_{CO_2}^F$ we also fix $X_{H_2O}^F$. This in turn fixes the partial pressure of H₂O, allowing us to calculate $X_{H_2O}^L$ consistent with that guess. We can then test whether our guess for $X_{CO_2}^F$ satisfies mass balance by checking for the equality:

$$(1 - f)X_{H_2O}^L + fX_{H_2O}^F - X_{H_2O}^T = 0 \quad (9)$$

The value of $X_{CO_2}^F$ that uniquely satisfies mass balance can then be found numerically by solving this equation using `scipy.optimize.root_scalar()`.

Text S3 - Code Benchmarking and Testing

A series of tests have been performed to ensure that models included in VESIcal are implemented correctly. For each model there is a folder containing a Jupyter notebook which runs through these tests, named “Testing_Model.pynb”. This reads from an excel spreadsheet named “Testing_Model.xlsx”. We detail these tests below:

Moore et al., 1998

- Benchmark 1 compares the solubility of H₂O in the melt calculated using the excel spreadsheet of Moore et al. (1998) to the outputs of VESIcal for a synthetic array of inputs. The outputs match to within +- 0.5% (Fig. 1).
- Benchmark 2 compares saturation pressures calculated with the Moore et al. (1998) spreadsheet to the outputs of VESIcal. The outputs match to within 0.15% for pressures within the calibration range (Fig. 2).

Dixon (1997), as implemented in VolatileCalc (Newman and Lowenstern, 2002)

- Benchmark 1 compares saturation pressures from VESIcal to the VolatileCalc spreadsheet, as well as an Excel Macro which has been used by a number of studies (e.g., Tucker et al., 2019; Fig. 3).
- Benchmark 2 compares X_{H₂O} in the fluid phase at volatile saturation to that predicted by VolatileCalc and the Excel Macro (Fig. 5).
- Benchmark 3 compares isobars to those produced by VolatileCalc for SiO₂=49 wt% (Fig. 6).

- Benchmark 4 compares an open system degassing path kindly provided by J. Dixon to VESIcal, as well as comparing VESIcal to VolatileCalc-Basalt (Fig. 7)

Liu et al., 2005

- Benchmark 1 recreates the model lines shown on Fig. 5 of Liu et al. (2005) by applying calculate_saturation_pressure to a synthetic array of inputs with variable H₂O contents (at three different temperatures, Fig. 8).
- Benchmark 2 recreates the model lines shown on Fig. 7 of Liu et al. (2005) by applying calculate_saturation_pressure to a synthetic array of inputs with variable CO₂ contents (at 1050°C; Fig. 10).
- Benchmark 3 recreates the isobars shown in Fig. 6b of Liu et al. (2005) using calculate_isobars_and_isopleths.

Iacono-Marziano et al., 2012

- Benchmark 1 compares saturation pressures obtained from the web calculator hosted at <http://calcul-isto.cnrs-orleans.fr/> to VESIcal outputs for compositions in the calibration dataset (Fig. 21a).
- Benchmark 2 compares saturation pressures obtained from the web calculator to VESIcal outputs for a synthetic dataset (Fig. 21b).
- Benchmark 3 compares dissolved volatile contents calculated from the web app to those from VESIcal (Fig. 11).

Shishkina et al. 2014

- Benchmark 1 uses the Π values in the supporting spreadsheet of Shishkina et al. to calculate CO₂ solubility, and compares this to the results from VESIcal (Fig. 12).

MagmaSat (Ghiorso and Gualda, 2015)

- Benchmark 1 compares saturation pressures from Bennett et al. (2019), who used the Mac App, to those calculated using VESIcal (Fig. 13).
- Benchmark 2 compares the isobars shown in Fig. 14 of Ghiorso and Gualda (2015) to those calculated with VESIcal. We note that although the figure caption says that the composition of the Late Bishop Tuff was used, VESIcal matches their isobars exactly if the composition of the Early Bishop Tuff is used and so this is the comparison shown (Fig. 14).
- Benchmark 3 compares X_{H_2O} calculated using the 'Fluid+magma from bulk composition' option of the web app with the calculate_equilibrium_fluid_comp function of VESIcal for a set of synthetic inputs (Fig. 15).

Allison et al. (2019)

- Benchmark 1 compares saturation pressures calculated in the spreadsheet provided by Allison et al. (2019) for the Sunset Crater model, using a synthetic array of inputs with variable CO₂ contents (H₂O=0 wt%; Fig. 16).
- Benchmark 2 compares saturation pressures from the spreadsheet of Allison et al. (2019) to those calculated by VESIcal for all 6 composition for 100, 5000 and 10,000 ppm CO₂ (and H₂O=0 wt%; Fig. 17).

Text S4 - Model comparisons with original calibration datasets

In addition to a benchmark comparison between VESIcal results and results from a published model calculator (e.g., web app or excel spreadsheet), or where it was not possible to compare VESIcal results directly to those of a published tool, we have compared VESIcal

results with measured or known values from experiments that make up the original calibration dataset. As with benchmarking tests shown above, a Jupyter notebook is provided for each model shown here, which runs through these tests, named “Testing_Model.pynb”, which reads from an excel spreadsheet named “Testing_Model.xlsx”.

Moore et al. (1998)

- Test 1 compares the experimental pressures for the H₂O-only experiments in the calibration dataset of Moore et al. (1998) with the saturation pressures obtained from VESIcal using the “MooreWater” model (Fig. 18a).

Liu et al. (2005)

- Test 1 compares the experimental pressures for the H₂O-only experiments in the calibration dataset of Liu et al. (2005) with the saturation pressures obtained from VESIcal using the “LiuWater” model (Fig. 19).

Iacono-Marziano et al. (2012)

- Test 1 compares the experimental pressures of the H₂O-only experiments in their calibration dataset to the saturation pressures obtained from VESIcal using the “Iacono-MarzianoWater” model (Fig. 20a).
- Test 2 compares the experimental pressures of the H₂O-CO₂ experiments in their calibration dataset to the saturation pressures obtained from VESIcal for the “Iacono-Marziano” model (Fig. 20b).

Shishkina et al., 2014

- Test 1 compares the experimental pressures of the H₂O-only experiments in the calibration dataset of Shishkina et al. (2014) to the saturation pressures obtained from VESIcal using the “ShishkinaWater” model (Fig. 22a).
- Test 2 compares the experimental pressures of the CO₂-only experiments in the calibration dataset to the saturation pressures obtained from VESIcal using the “ShishkinaCarbon” model (Fig. 22b).
- Test 3 compares the experimental pressures for mixed H₂O-CO₂ experiments presented in their to the saturation pressures obtained from VESIcal using the “Shishkina” model (Fig. 23).

Text S5 - Normalization

As mentioned in the main text, MagmaSat, Iacono-Marziano and Dixon are sensitive to normalization. To demonstrate this, first we consider the basalt composition with a volatile-free analytical total of 96.77 wt% described in the main text. In Figs. 24 and 25, we show saturation pressures calculated using these three different normalization routines, as well as raw data which has not been normalized. All three of VESIcal’s normalization routines change the absolute concentration of major elements, so change the saturation pressure calculated using the Dixon model (which is sensitive to SiO₂ content). The effect of melt SiO₂ on pure H₂O solubility is relatively small, so saturation pressures change by only 5% for different normalization routines (Fig. 24c). It is apparent from comparing Fig. 24c to 24f that saturation pressures calculated for pure H₂O are more sensitive to the changes in major element contents than the small change in volatile contents (i.e., biggest change in saturation pressure is between the circle and pentagon, which have the same

H_2O content, but SiO_2 contents that differ by >1 wt%. The change in saturation pressure for the pure CO_2 example is significantly greater; normalization using any of the three routines causes the saturation pressure to increase by a factor of 1.46–1.47× compared to the raw input. This reflects the very high sensitivity of the Dixon CO_2 expression to absolute SiO_2 content.

To assess whether data should be normalized for use with Dixon, we plot the North Arch glasses of Dixon (1997) in SiO_2 vs. Π space, and compare this to Figure 2C in Dixon (1997) (Fig. 27). Raw data, with analytical totals ranging from 97.382 to 99.99 wt%, provides a better match than normalized data (which is offset to higher SiO_2 contents). We also contacted Jackie Dixon, who confirmed that she had used unnormalized data. Thus, we recommend that the Dixon expression is best applied to unnormalized data, to be consistent with the calibration dataset. However, users should be aware of the extreme sensitivity of this model to normalization. If their totals are significantly lower than those in Dixon (1997), perhaps due to excess charge build up or high melt H_2O contents (see Huges et al., 2019), or literature data has been scaled up so major elements sum to 100 wt%, not doing any form of normalization may yield spurious results.

Using Iacono-Marziano (Hydrous and Anhydrous), the most significant change in saturation pressure is encountered using the standard normalization routine. Tests using the web app of Iacono-Marziano yield the same saturation pressure for raw, additional-volatile normalized, and fixed volatile normalized. This is because input major elements are normalized to 100% before any calculations are performed (identical to the additionalvolatiles routine in VESIcal). Only changing the input volatile contents changes saturation pres-

sures on the web app. Based on the implementation in the web app, in VESIcal, all inputs are automatically normalized using the additionalvolatiles routine for consistency.

In VESIcal, the Raw and standard normalization routine for MagmaSat yield the same saturation pressure as one another, which is higher than the saturation pressures from additional- and fixedvolatiles (which are also identical). However, the 10 bar resolution of MagmaSat must be considered for such comparisons. To test this further, we consider the same composition, but instead with 10 wt% H₂O (causing much larger differences between normalization routines, that can be distinguished considering the model resolution; Fig. 26). As with Iacono-Marziano, the absolute concentrations of majors and volatiles are not important to the calculated pressure in MagmaSat. Instead, it is the relative abundances of different elements in the system. Thus, raw data (no normalization) and standard normalization will always give the same saturation pressures, because simply scaling up or down major and volatile elements by the same amount to give a total of 100 does not affect relative abundances.

The 10 wt% H₂O example is the easiest way to visualize the differences between additionalvolatiles and fixedvolatiles. Additionalvolatiles effectively scales up the major elements in the system to 100 wt%, keeping H₂O=10 wt%. Considering the system as having 110 g of mass, this means that 100 g of the system is major elements, and 10 g is H₂O. Thus the proportion of H₂O in the system is 10/110=0.09 (pentagon, Fig. 26). In contrast, fixedvolatiles rescales major elements to 90 wt%, and H₂O at 10 wt%. For a 100 g system, 90 g is major elements, and 10 g is H₂O, so the proportion of H₂O is 10/100=0.1. This proportion is higher than for fixedvolatiles, which explains why the saturation pressure is higher (diamond vs. pentagon in Fig. 26). These differences are

too small to recognize for the 1 wt% H₂O example. Using the same line of reasoning, the raw composition contains 97.38 g of major elements, and 10 g of H₂O. The proportion of H₂O in this system is $10/(97.38+10)=0.093$. Similarly, the standard normalized example has scaled up volatile and major elements by the same factor (to H₂O=9.31 wt% and major element total=90.68 wt%), so the proportion $(9.31/(9.31+90.68))$ is also 0.093. Thus, these two normalization routines predict saturation pressures between the fixed and additional volatiles example. In the MagmaSat app, users enter major element and volatile concentrations. The app then scales the major element concentrations while keeping the volatile element concentrations constant, such that the hydrous total is 100 (Fig. 28). Thus, saturation pressures obtained using the Fixed Volatile normalization routine in VESIcal gives answers comparable to the web app.

References

- Dixon, J. (1997). Degassing of alkalic basalts. *American Mineralogist*, 82, 368–378.
- Huges, E., Buse, B., Kearns, S., Blundy, J., Kilgour, G., & Mader, H. (2019). Low analytical totals in EPMA of hydrous silicate glass due to subsurface charging: Obtaining accurate volatiles by different. *Chemical Geology*, 505, 48–56.

Moore et al., 1998 - Benchmark 1

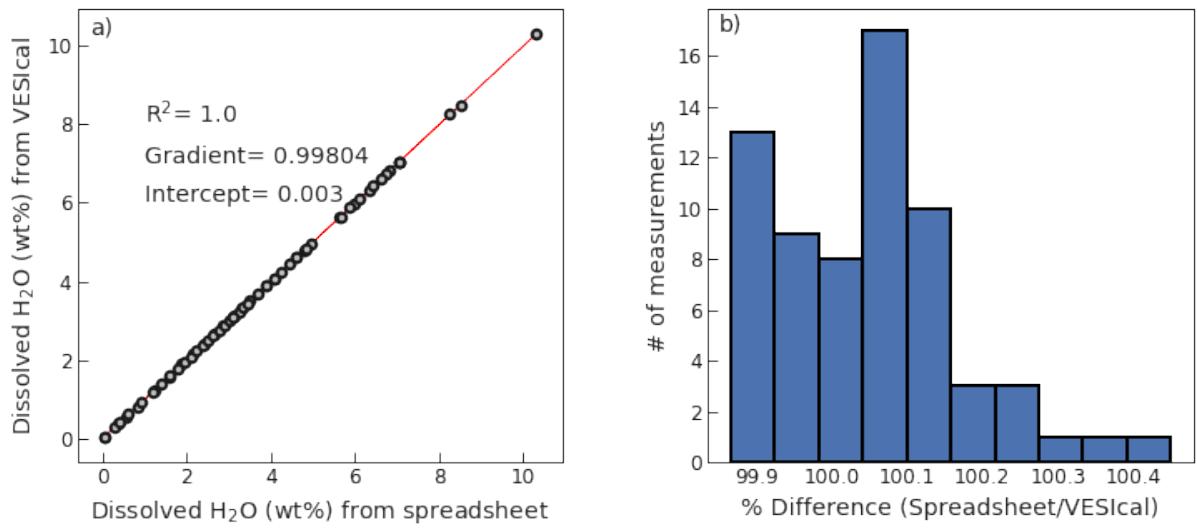


Figure 1. a) Comparison of the wt% H_2O dissolved in the melt calculated using the excel spreadsheet of Moore et al. (1998) with that calculated in VESICAL. The very small discrepancies vary as a function of pressure (see Fig. 2.)

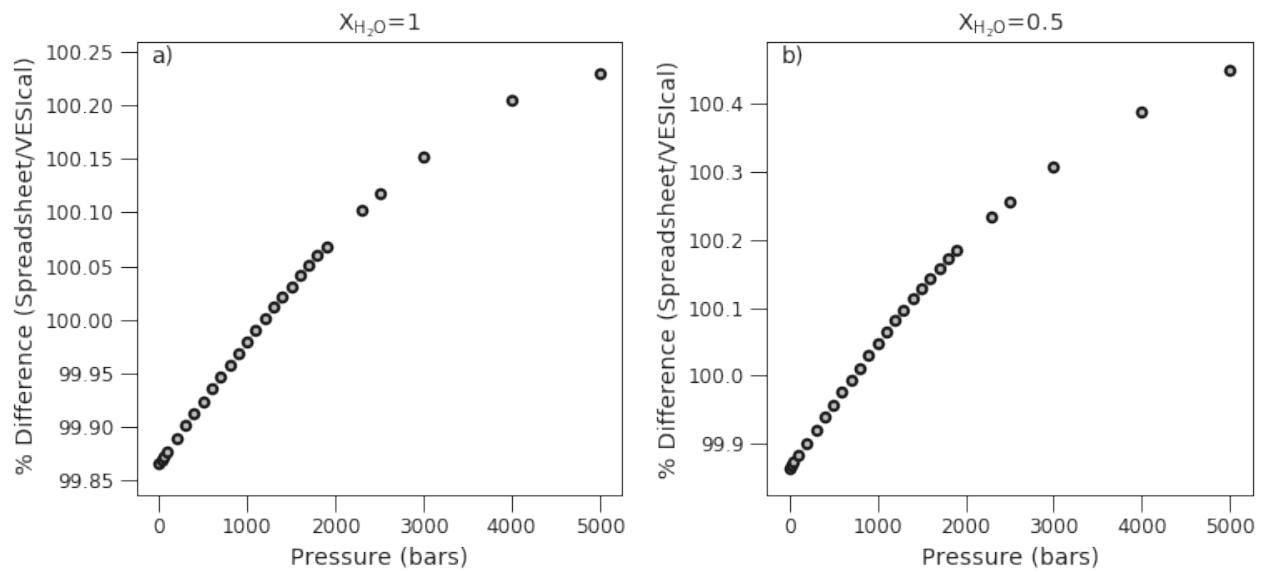
Moore et al., 1998 - Benchmark 2

Figure 2. The discrepancy between the wt% H₂O in the melt calculated using the Moore spreadsheet and VESICAL varies with pressure, but is <0.15% for pressures within the calibration range (0-3000 bars).

VolatileCalc implementation of Dixon (1997) - Benchmark 1

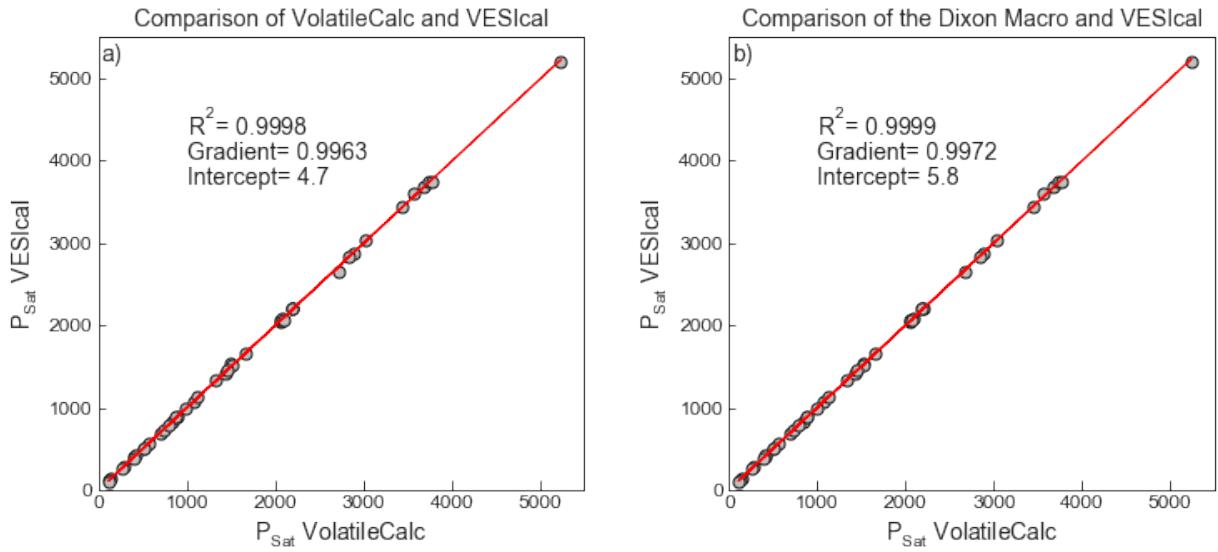


Figure 3. Comparison of the saturation pressures calculated using the VolatileCalc excel spreadsheet to those calculated in VESIcal, and a widely-used excel macro (e.g., Tucker et al., 2019).

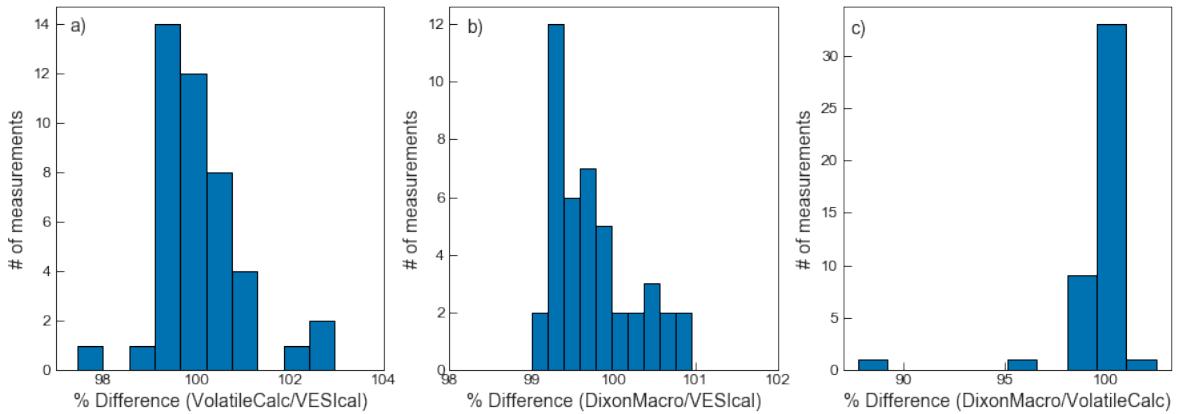


Figure 4. The percentage difference between saturation pressures from VESIcal vs. VolatileCalc (a) and VESIcal vs. the excel macro (b) are similar in magnitude to those from VolatileCalc vs. the excel macro (c).

VolatileCalc implementation of Dixon (1997) - Benchmark 2

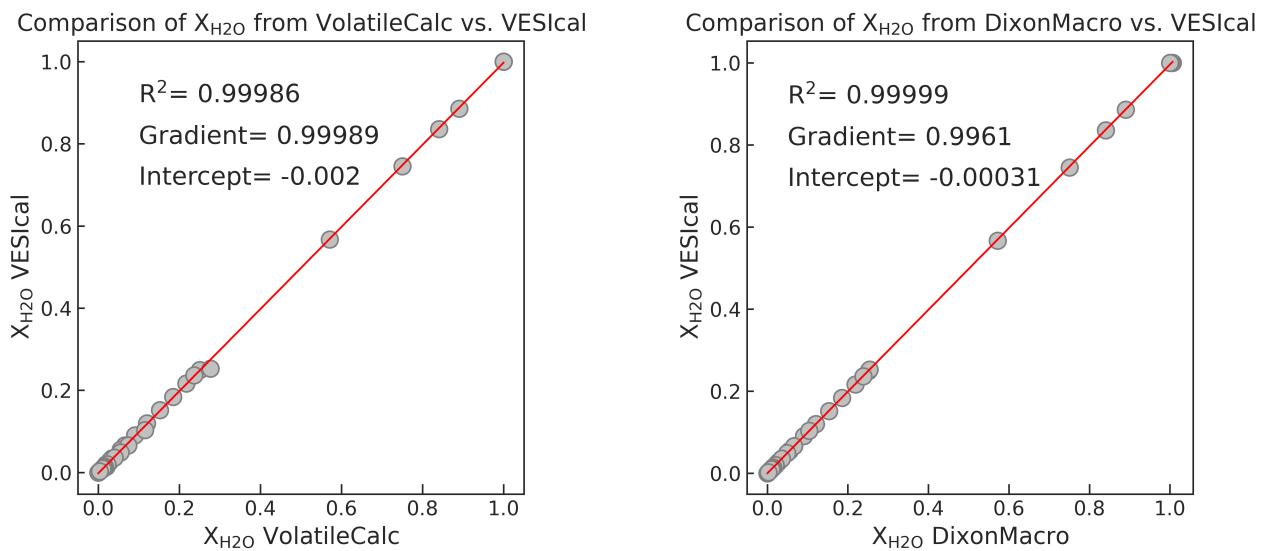


Figure 5. Comparison of X_{H_2O} in the fluid phase at volatile saturation calculated using the VolatileCalc excel spreadsheet to those calculated in VESIcal, and a widely-used excel macro (e.g., Tucker et al., 2019).

VolatileCalc implementation of Dixon (1997) - Benchmark 3

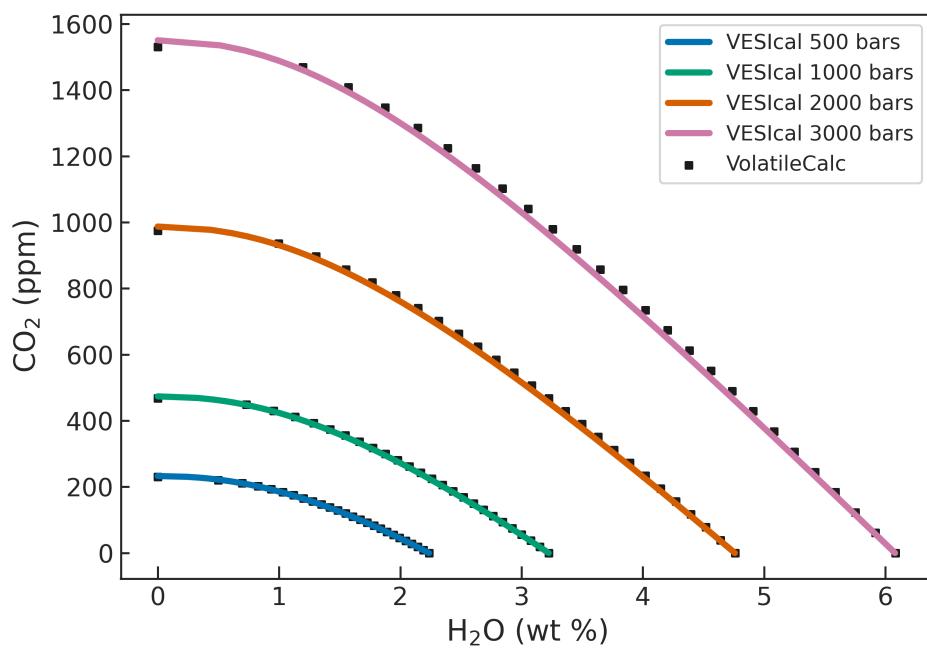


Figure 6. Comparison of isobars generated in VolatileCalc to those in VESIcal.

VolatileCalc implementation of Dixon (1997) - Benchmark 4

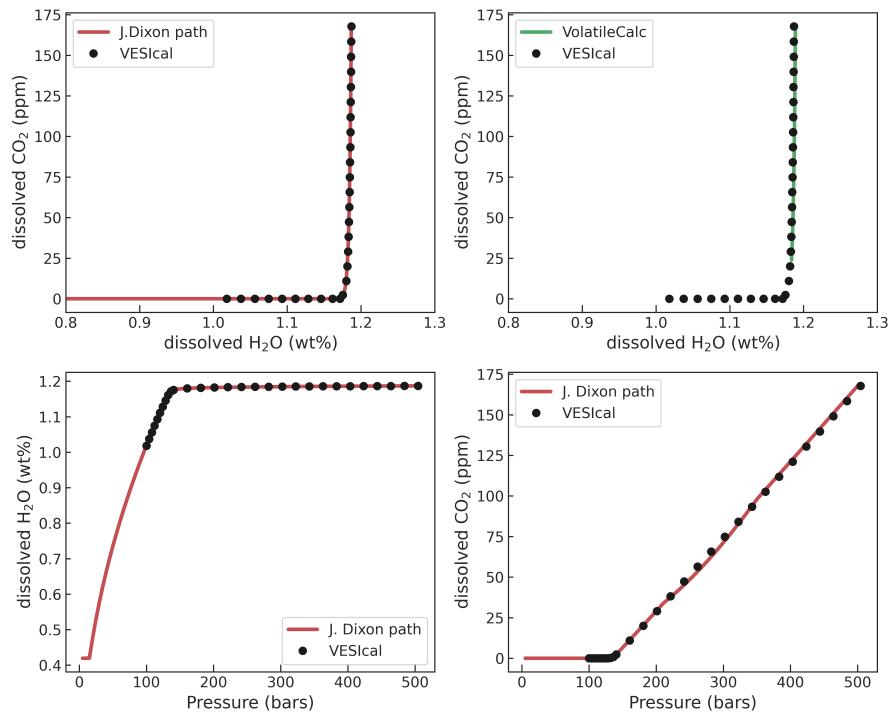


Figure 7. Comparison of an open system degassing path for a tholeiite calculated by VESIcal to that supplied by J. Dixon and that calculated by VolatileCalc. We limited calculations to 100 steps, so VESIcal does not calculate the very long tail to extremely low CO_2 contents calculated by J. Dixon.

Liu et al., 2005 - Benchmark 1

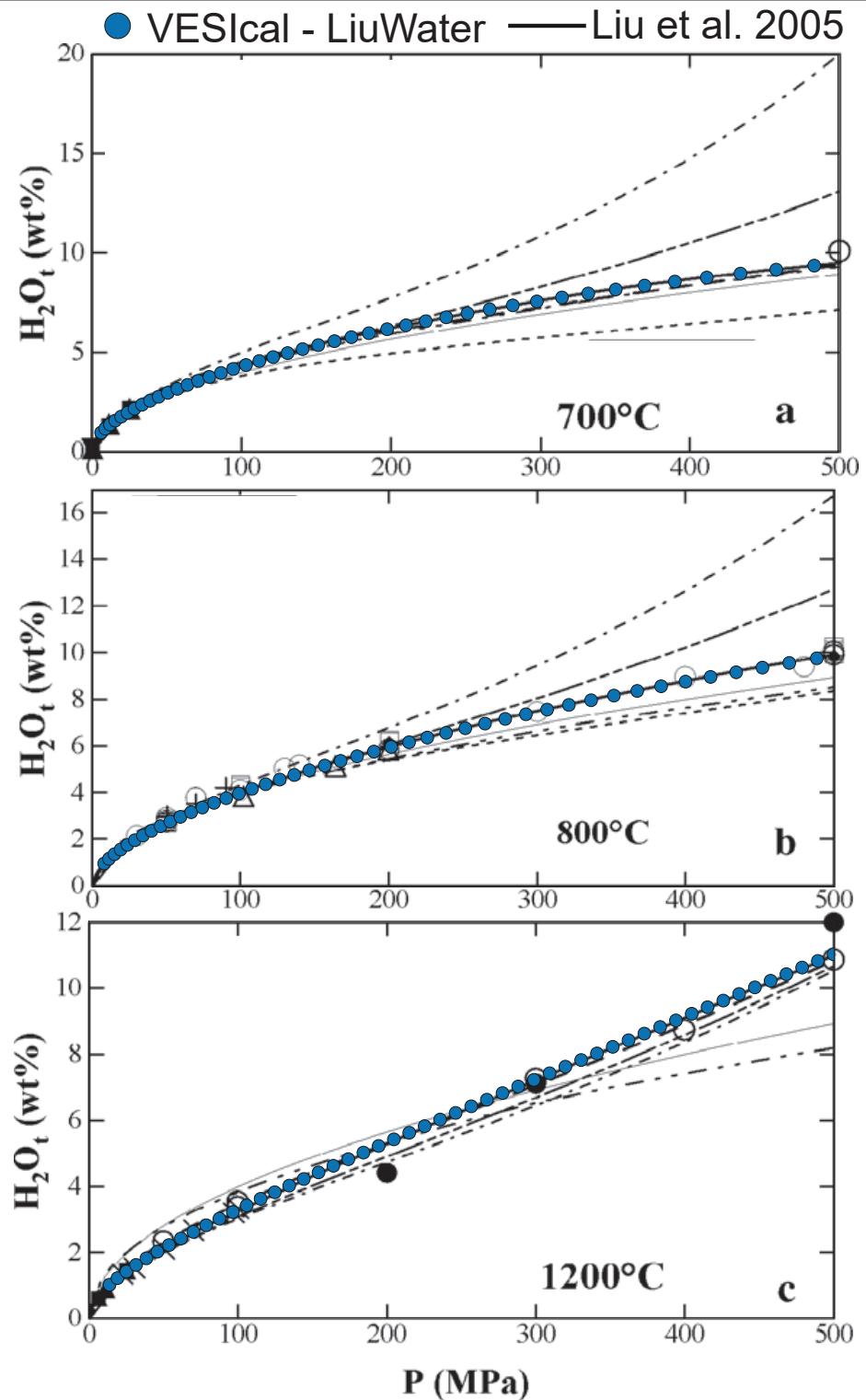


Figure 8. Relationship between pressure and dissolved H_2O content calculated in VESIcal for three different temperatures overlain on Fig. 5 of Liu et al. (2005).

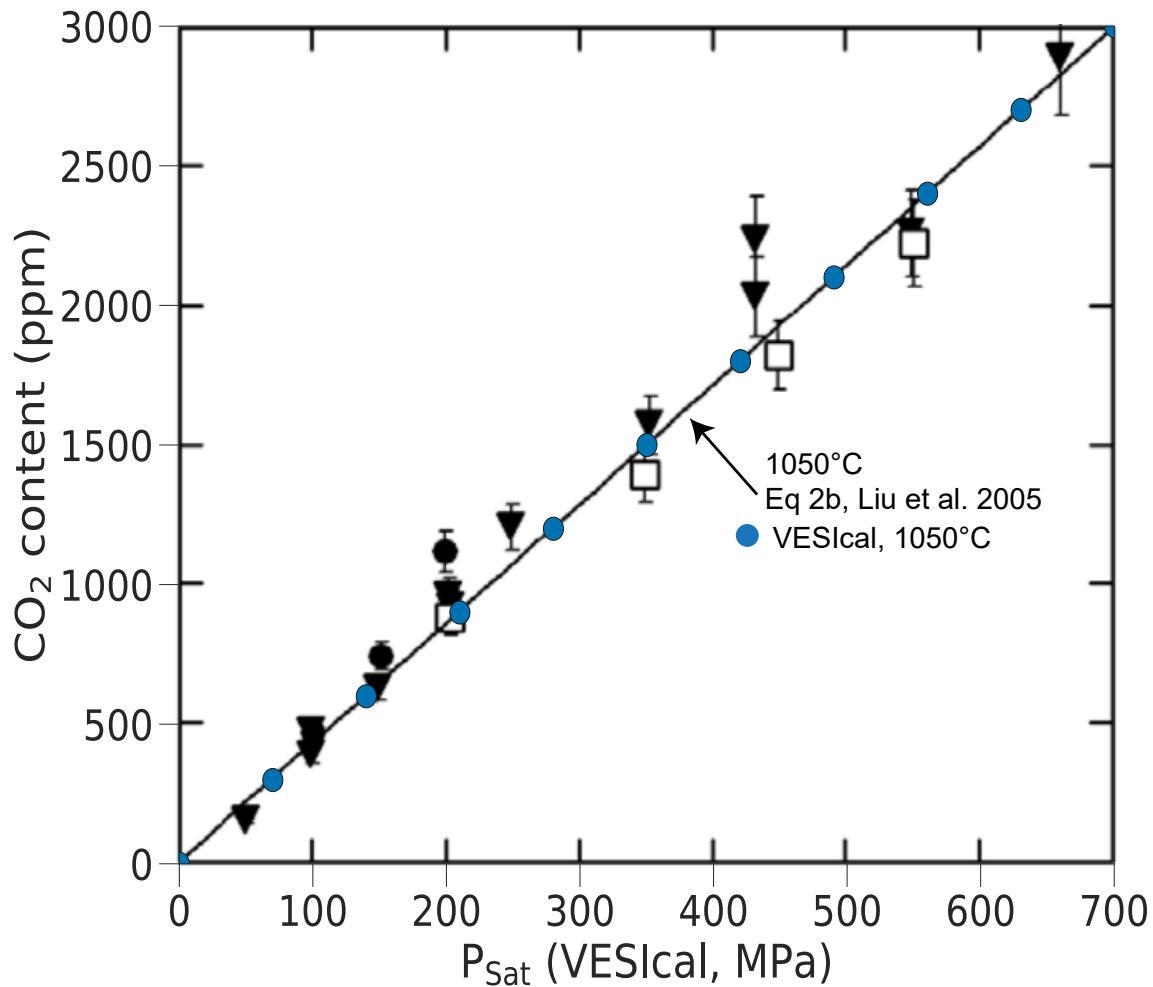
Liu et al., 2005 - Benchmark 2

Figure 9. Relationship between pressure and dissolved CO_2 content calculated using VESIcal overlain on Fig. 7 of Liu et al. (2005)

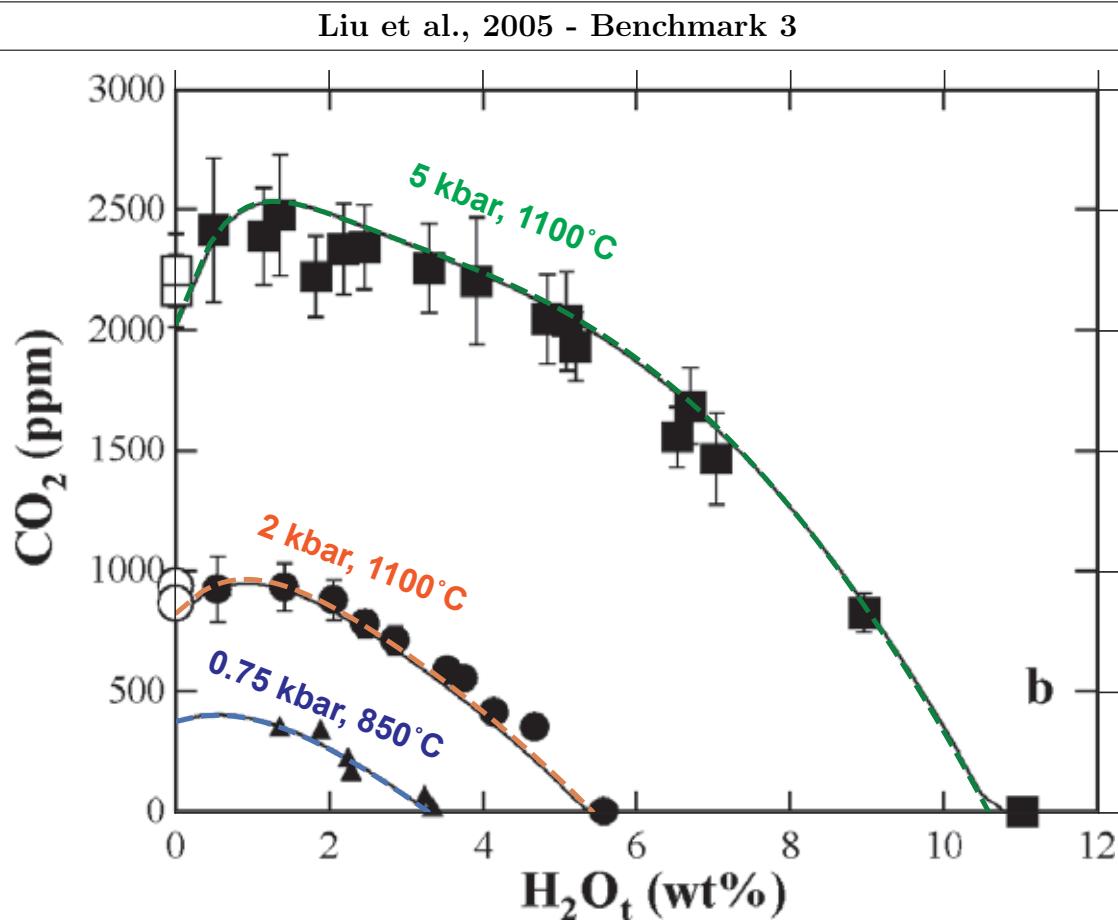


Figure 10. Isobars calculated using VESIcal overlain on Fig. 6b of Liu et al. (2005). For the 5 kbar isobar, H_2O and CO_2 contents are multiplied by 0.98 following their description in the figure caption.

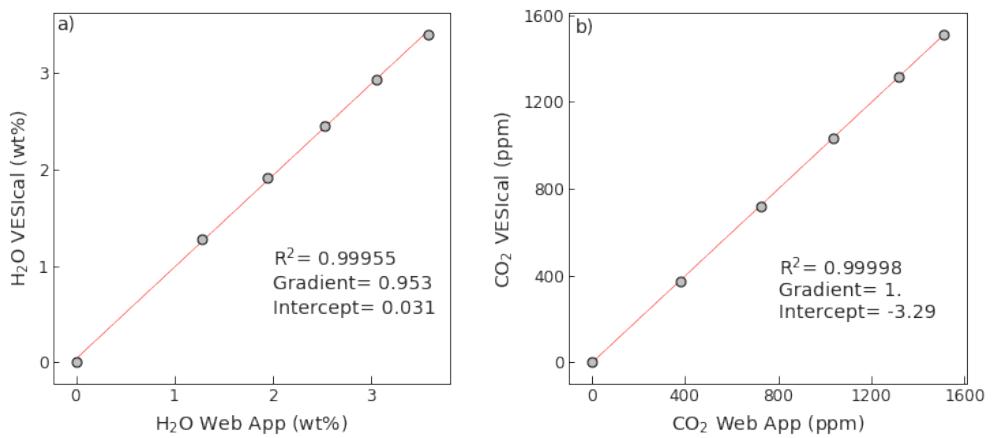
Iacono-Marziano et al., 2012 - Benchmark 1

Figure 11. Comparison of dissolved volatiles calculated using the web app and VESICAL for a synthetic dataset.

Shishkina et al., 2014 - Benchmark 1

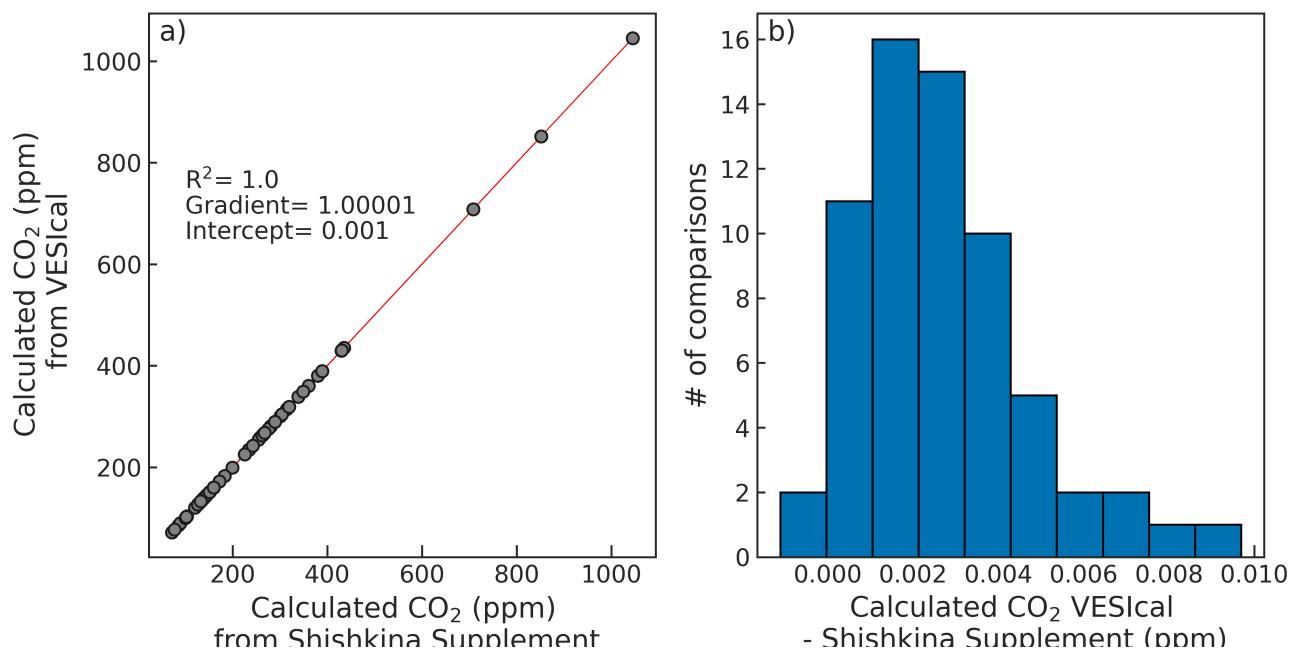


Figure 12. Comparison of CO_2 solubility calculated using the Π values in the supporting information of Shishkina with those calculated using VESIcal.

MagmaSat (Ghiorso and Gualda, 2015) - Benchmark 1

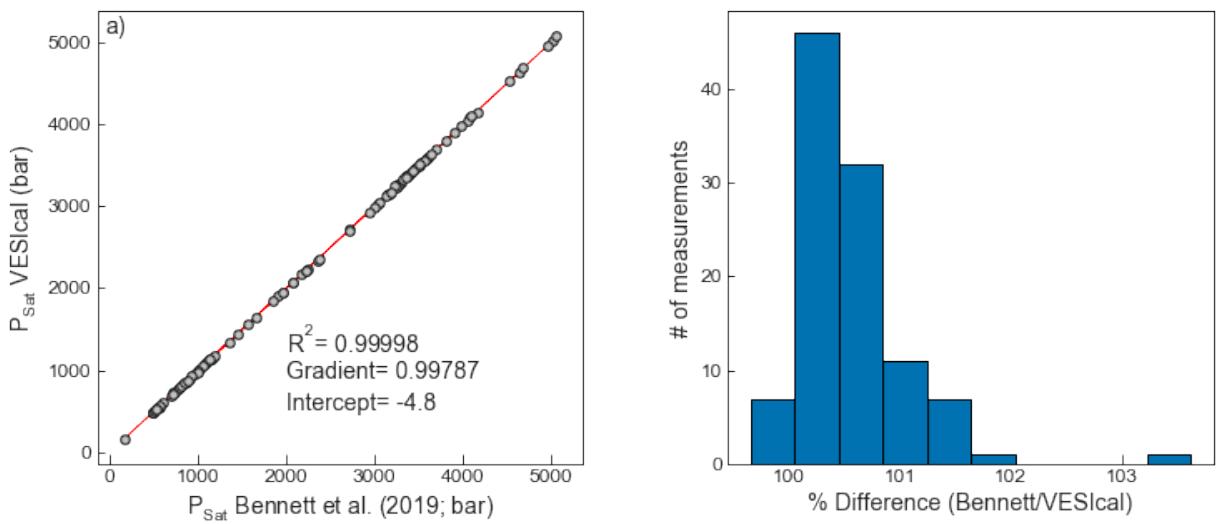


Figure 13. Comparison of saturation pressures calculated in VESIcal to those published by Bennett et al. (2019, who used the Mac App). These results agree to within 1% (within error of the rounding within VESIcal to the nearest MPa).

MagmaSat (Ghiorso and Gualda, 2015) - Benchmark 2

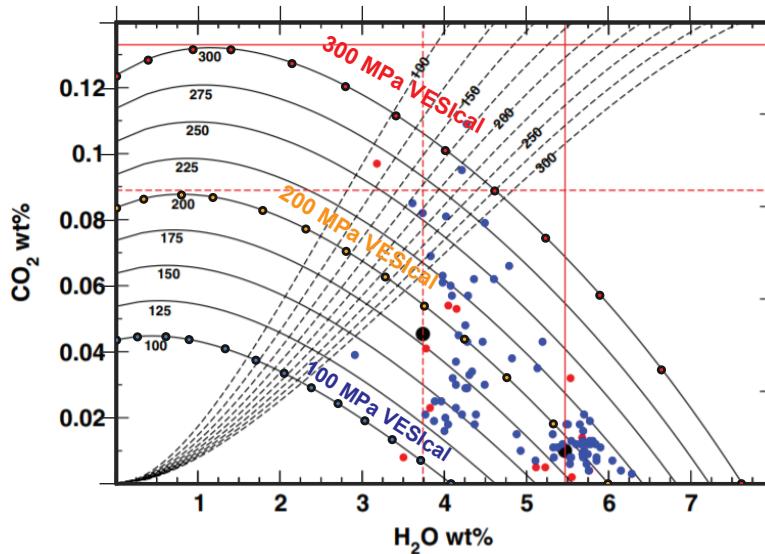


Figure 14. Comparison of isobars produced in VESICal for the Bishop Tuff (colored dots) to those shown in Fig. 14 of Ghiorso and Gualda (lines). Although the published figure states that these isobars are drawn for the Late Bishop Tuff, the best match uses the Early Bishop Tuff composition from the table in their paper, and so that is the comparison we show here.

MagmaSat (Ghiorso and Gualda, 2015) - Benchmark 3

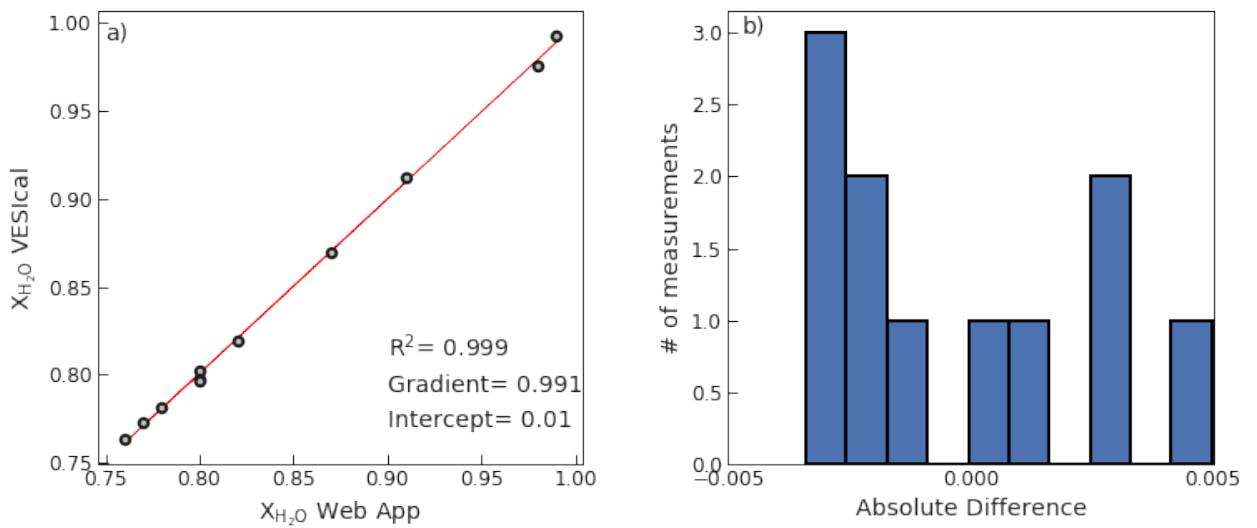


Figure 15. Comparison of X_{H_2O} from VESIcal and the web app (<http://209.180.202.74:8080/H2O-CO2BxApp/>) for a synthetic array of inputs. The web app only returns values to the nearest 0.01. The discrepancies between VESIcal and the web app lie within this rounding error.

Allison et al. (2019) - Benchmark 1

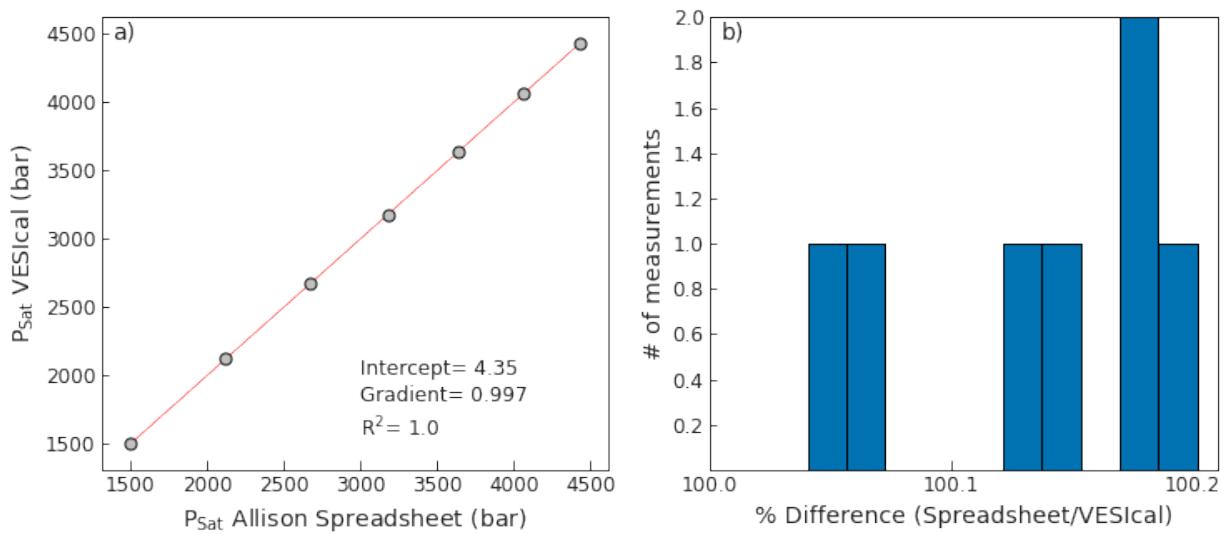


Figure 16. a) Comparison of saturation pressures from the Excel spreadsheet provided by Allison et al. (2019) for the Sunset Crater Model for a synthetic array of inputs.

Allison et al. (2019) - Benchmark 2

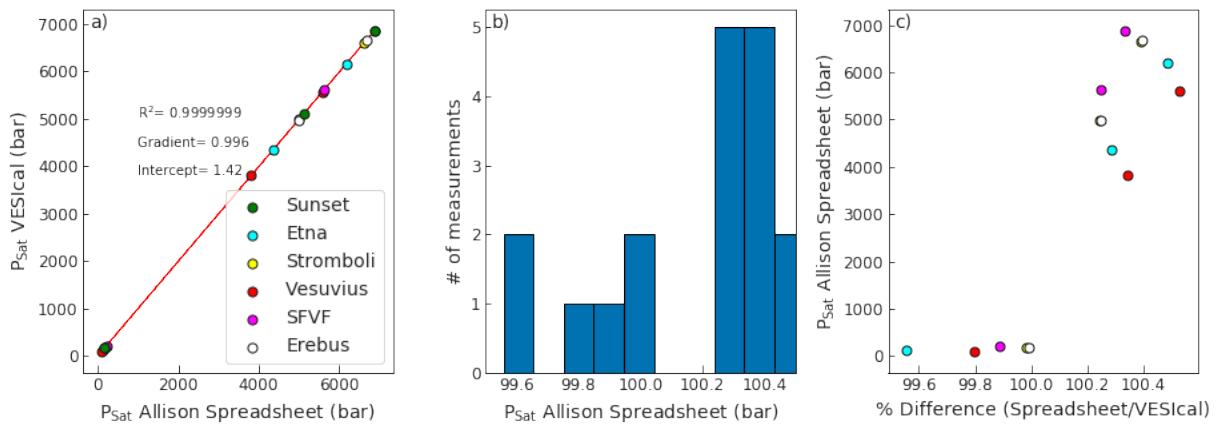


Figure 17. a) Comparison of saturation pressures from the Excel spreadsheet provided by Allison et al. (2019) for all 6 model compositions for a synthetic array of inputs. b) VESIcal matches the spreadsheet to within $\pm 0.5\%$. c) The percentage difference increases with increasing pressure.

Moore et al., 1998 - Test 1

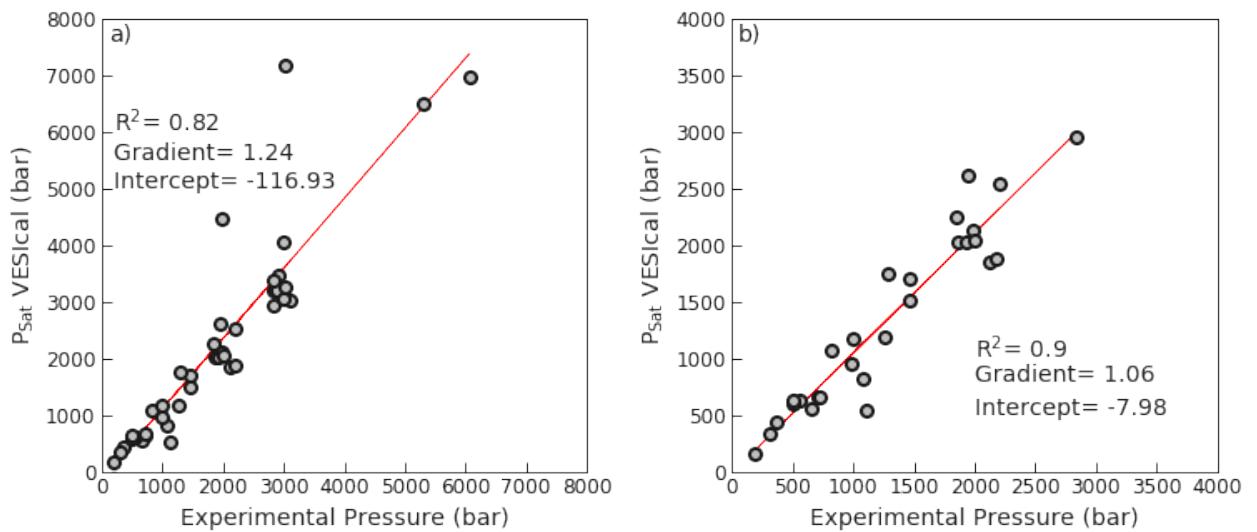


Figure 18. a) Comparison of the experimental pressures of the H_2O -only experiments in the calibration dataset of Moore et al. (1998) with saturation pressures calculated by VESIcal using ‘MooreWater’. b) The correspondence is improved if experimental and saturation pressures lying outside the calibration range of 0–3000 bars are removed.

Liu et al., 2005 - Test 1

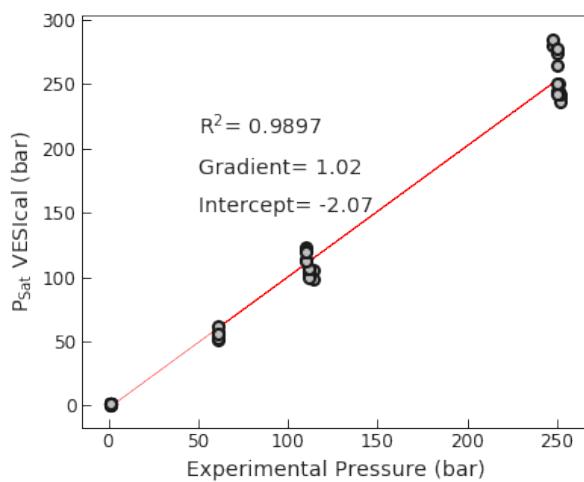


Figure 19. Comparison of the experimental pressures in the H₂O-only experiments in the calibration dataset of Liu et al. (2005) with saturation pressures calculated in VESIcal.

Iacono-Marziano et al., 2012 - Test 1

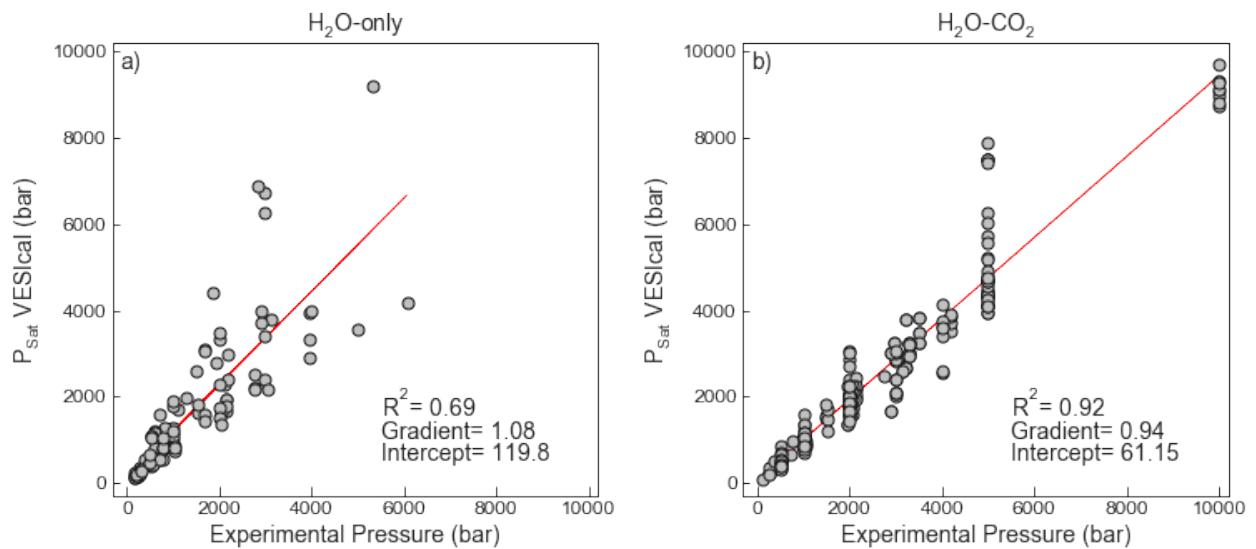


Figure 20. Comparison of the experimental pressures in H₂O-only experiments (a) and the H₂O-CO₂ experiments (b) in the calibration dataset of Iacono-Marziano with the saturation pressures calculated in VESIcal.

Iacono-Marziano et al., 2012 - Test 2

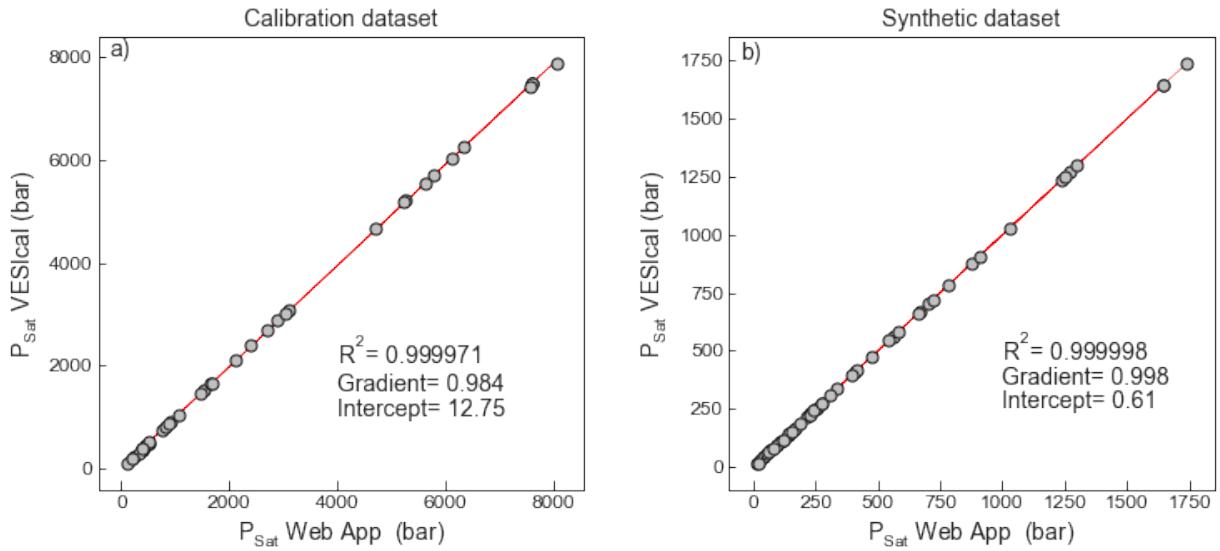


Figure 21. a) Comparison of saturation pressures from the web app of Iacono-Marziano (<http://calcul-isto.cnrs-orleans.fr/apps/h2o-co2-systems/>) and VESical for the major element and volatile contents of the calibration dataset (a) and a synthetic array of inputs (b).

Shishkina et al., 2014 - Test 1 and 2

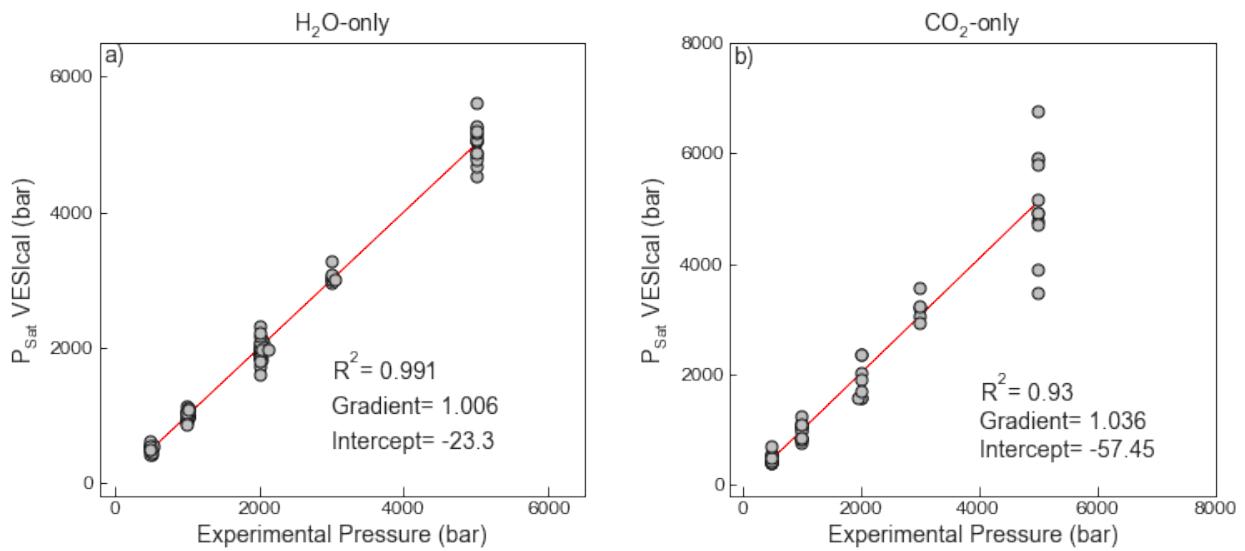


Figure 22. Comparison of the experimental pressures in H₂O-only experiments (a) and the CO₂ experiments (b) in the calibration dataset with the saturation pressures calculated by VESIcal.

Shishkina et al., 2014 - Test 3

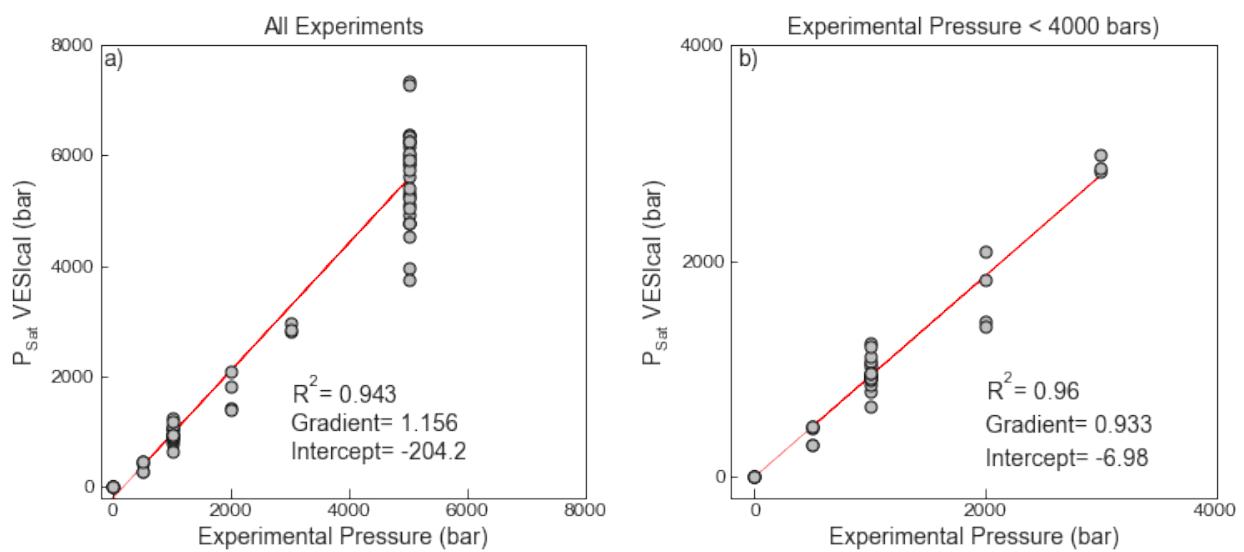


Figure 23. a) Comparison of the experimental pressures for the H₂O-CO₂ experiments presented in Table 2 of the main text with the saturation pressures calculated by VESIcal. b) Regression excluding the experiments run at >4000 bars, which show considerable scatter.

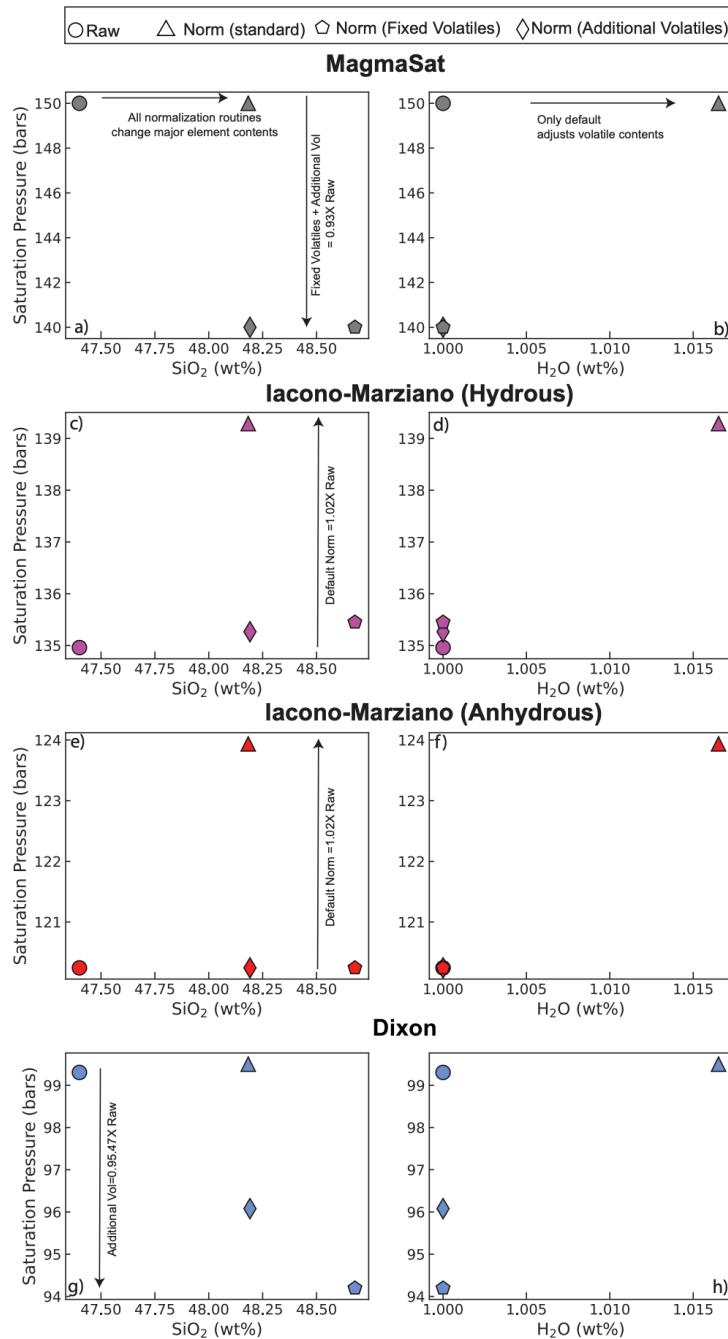


Figure 24. Saturation pressures calculated for a basalt with 1 wt% H_2O , 0 ppm CO_2 using different normalization routines.

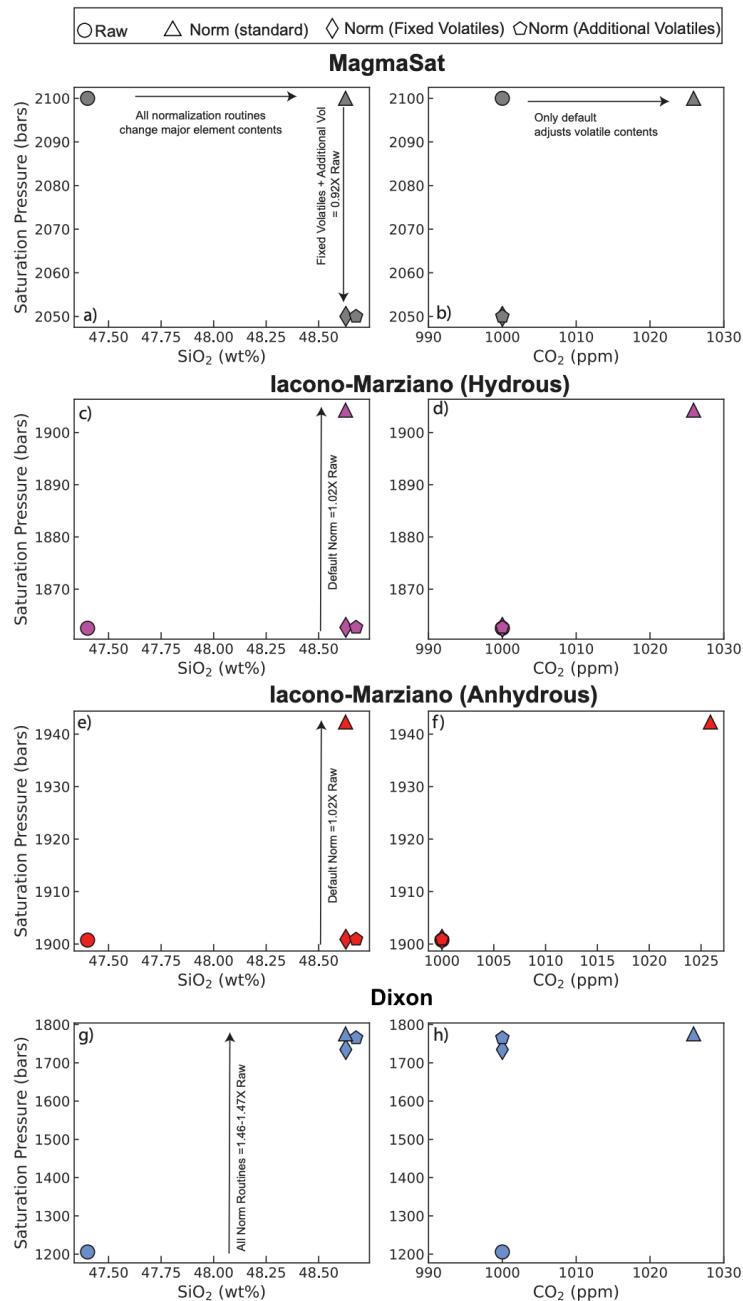


Figure 25. Saturation pressures calculated for a basalt with 0 wt% H_2O , 1000 ppm CO_2 using different normalization routines.

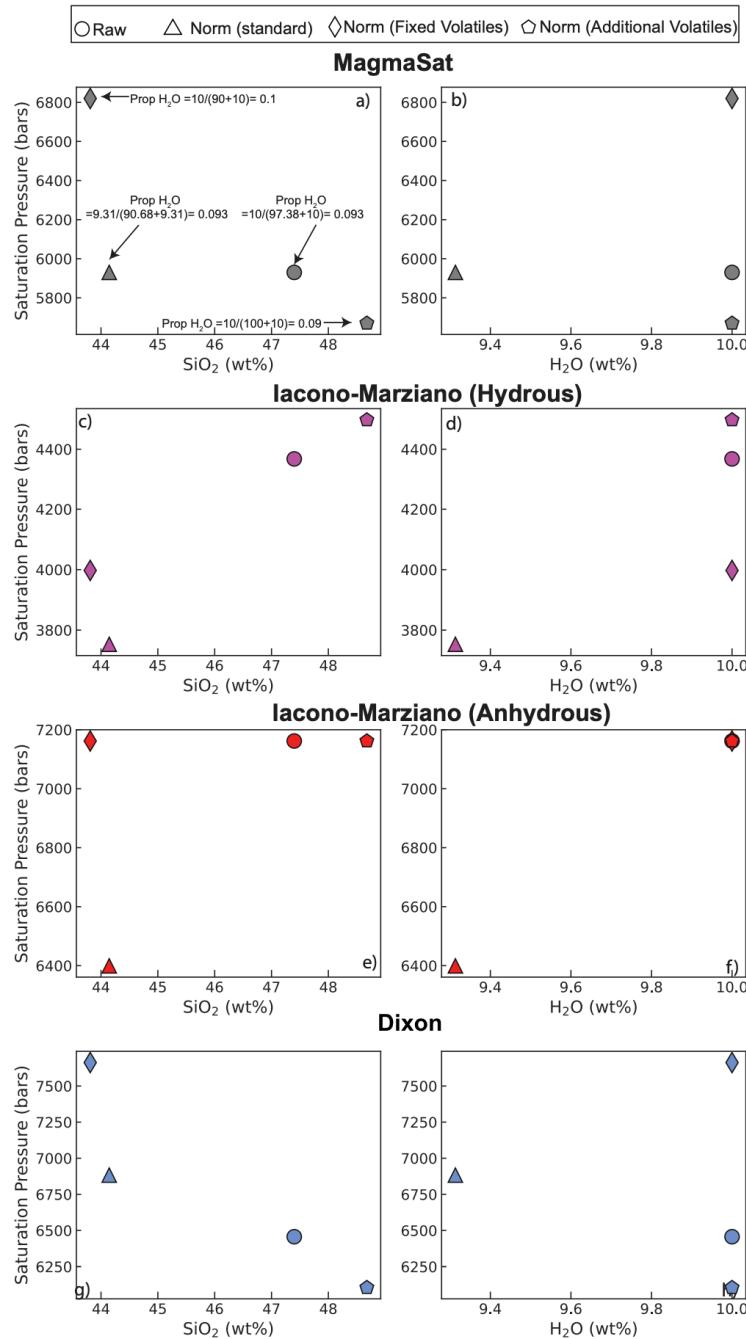


Figure 26. Saturation pressures calculated for a basalt with 10 wt% H_2O , 0 ppm CO_2 using different normalization routines.

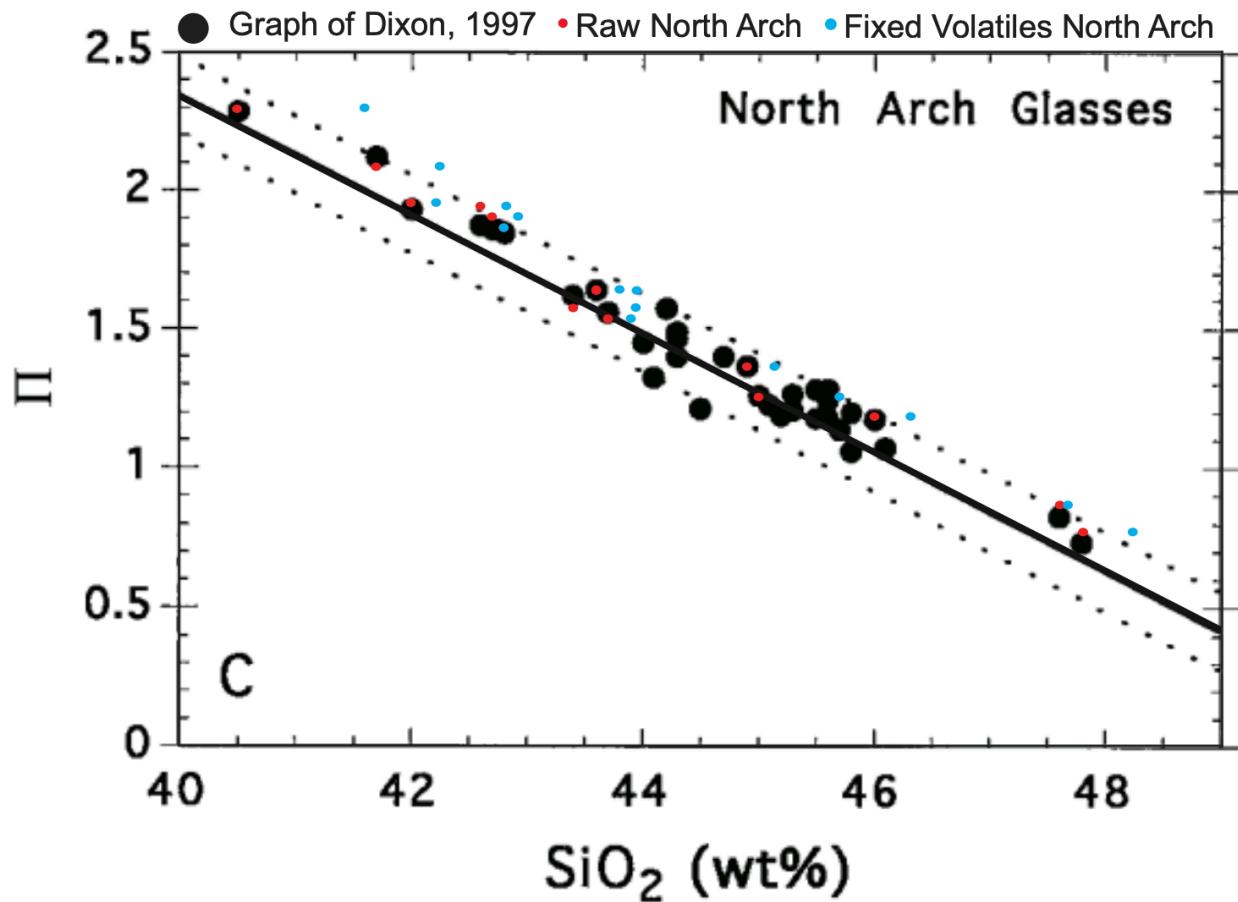


Figure 27. Π vs. SiO_2 graph shown in Fig. 2C of Dixon (1997). Red dots show the raw data presented in Table 1 of Dixon (1997), with analytical totals ranging between 97.38 to 99.995 wt%, while cyan dots show this data following the VESIcal fixedvolatiles normalization routine. It is apparent that raw data is a significantly better fit; all normalization routines shift SiO_2 contents to higher values.

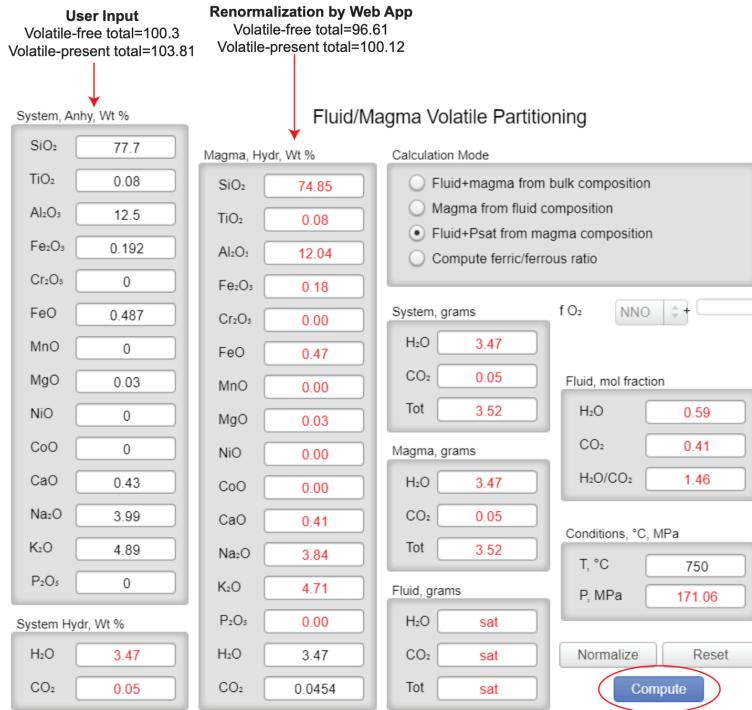


Figure 28. Screenshot of the MagmaSat app. Users type in data into the column on the left labelled system, anhyd wt%. After pressing the compute button (circled red), the app recalculates major element concentrations, and fill them into the second column (Magma, hyd wt%). Effectively, the app is performing a normalization routine identical to fixedvolatiles in VESIcal, leaving volatile concentrations the same, and scaling major elements so that the hydrous total is 100.