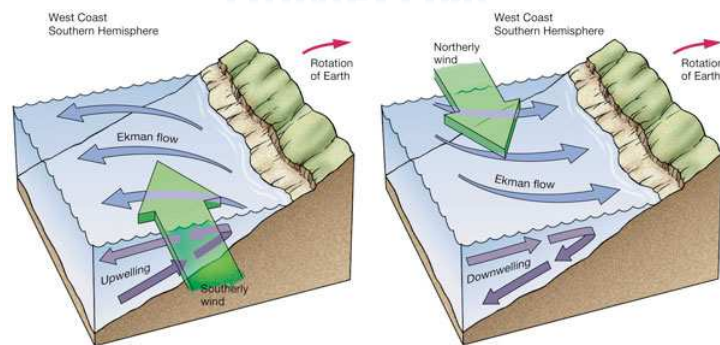


# Ekman transport, ENSO & IOD

## EKMAN TRANSPORT & UPWELLING

- Upwelling occurs in the open ocean and along coastlines when wind causes surface water to subside along a coastline.
- Subsurface water rises to the surface as a result and is typically colder, nutrient-rich, and biologically productive supporting rich fishing grounds (e.g. west coasts of Africa and South America).
- How does this happen?

## EKMAN TRANSPORT & UPWELLING



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## ENSO

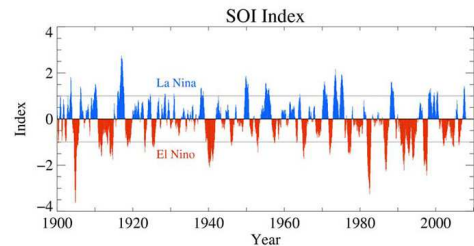
- El Niño and La Niña are opposite phases of the *El Niño-Southern Oscillation (ENSO)* cycle;
- ENSO describes the fluctuations in temperature between the ocean and atmosphere in the east-central Equatorial Pacific; **La Niña is the cold phase** and **El Niño the warm phase**.
- These deviations from normal surface temperatures can have large-scale impacts not only on ocean processes, but also on global weather and climate.

<http://oceanservice.noaa.gov>

## ENSO

- El Niño and La Niña episodes typically last 9 to 12 months, but some events may last years. They often begin to form between June and August, reach peak strength between December and April, and then decay between May and July of the following year.

<http://oceanservice.noaa.gov>

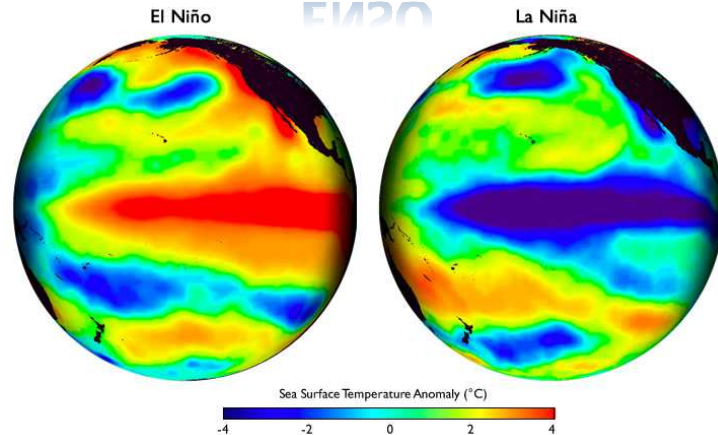


## ENSO

- While irregular, El Niño and La Niña events occur about every 3 to 5 years. Typically, El Niño occurs more frequently than La Niña.
- El Niño:** Large-scale ocean-atmosphere climate interaction linked to periodic warming in sea surface temperatures across central and east-central Equatorial Pacific.
- La Niña:** Periods of below-average sea surface temperatures across the east-central Equatorial Pacific.

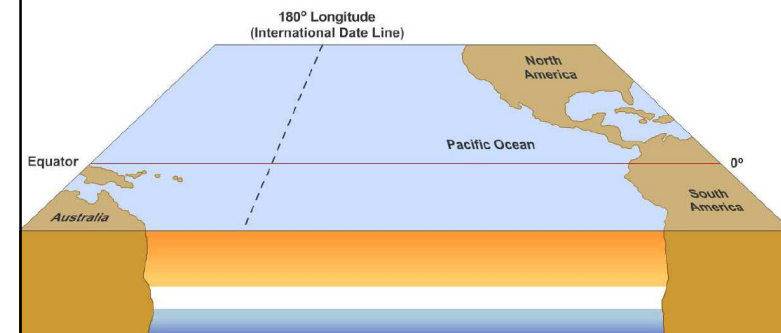
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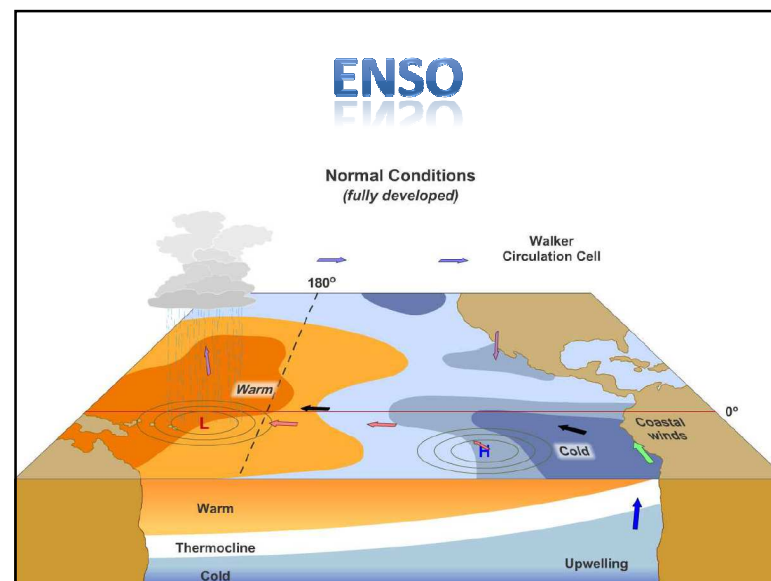
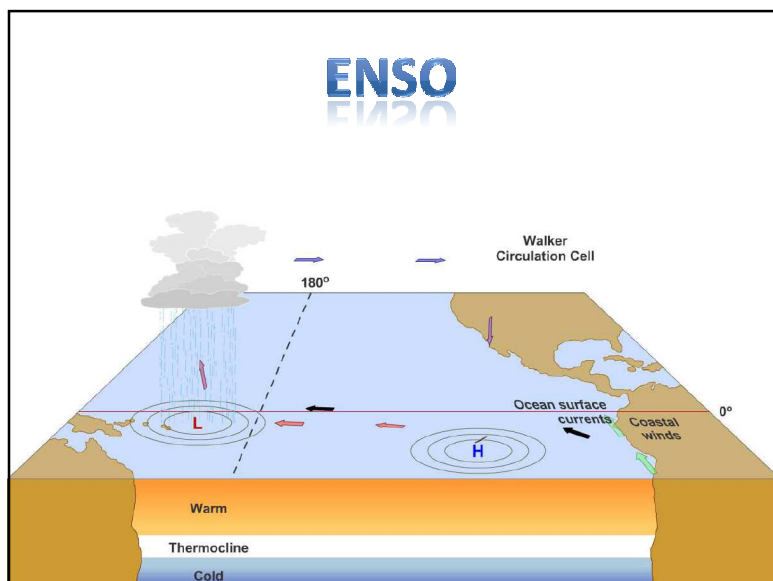
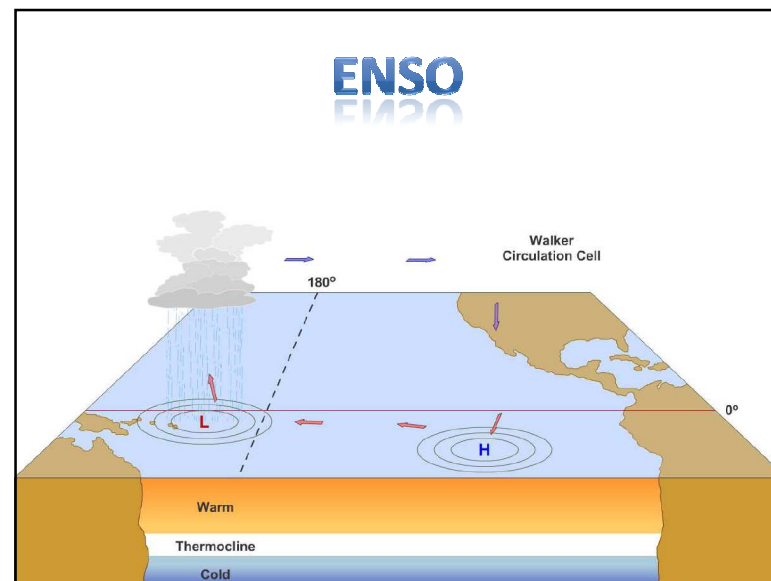
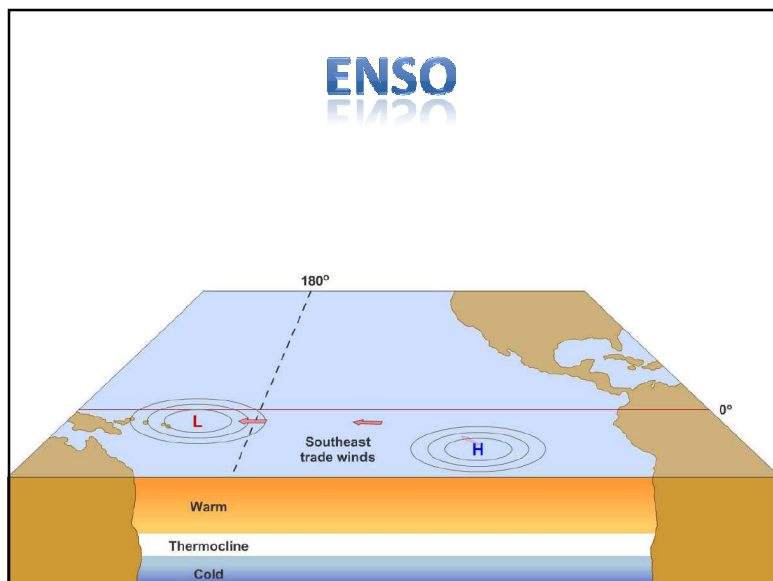
## ENSO

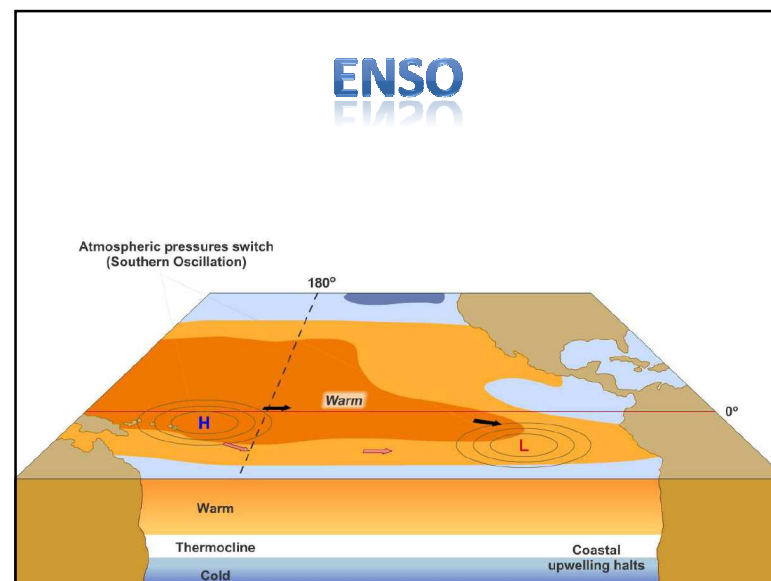
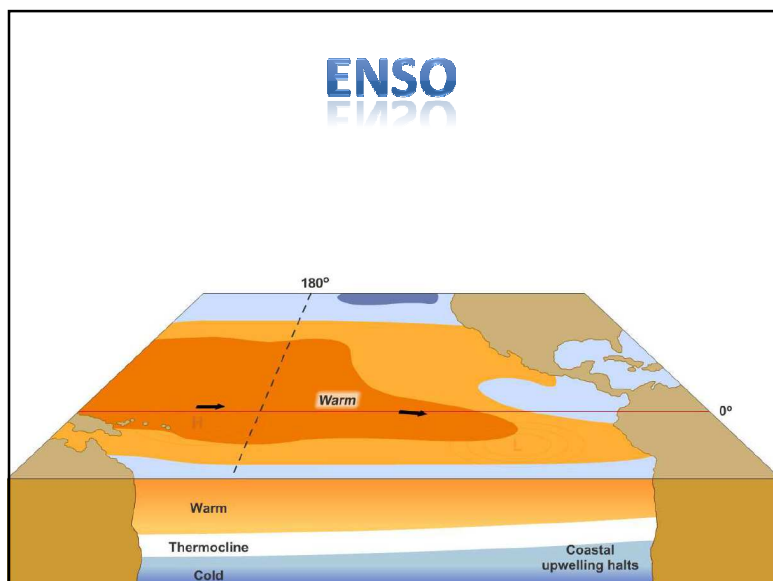
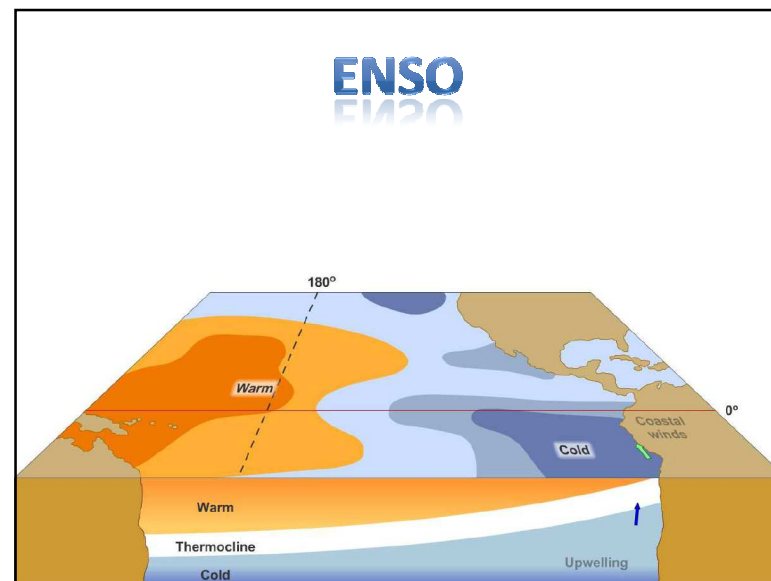
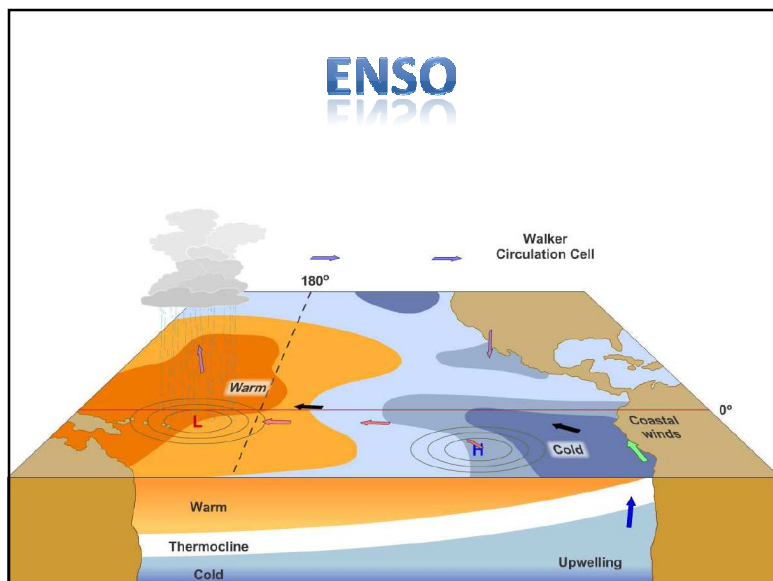


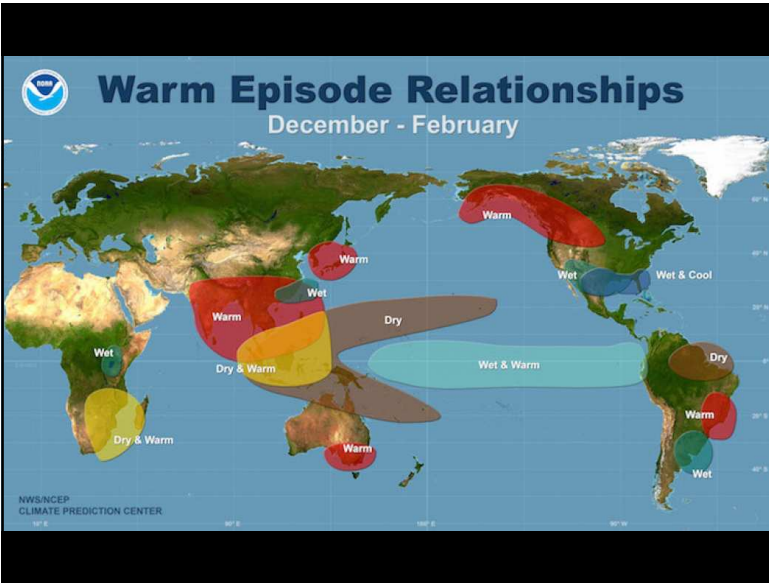
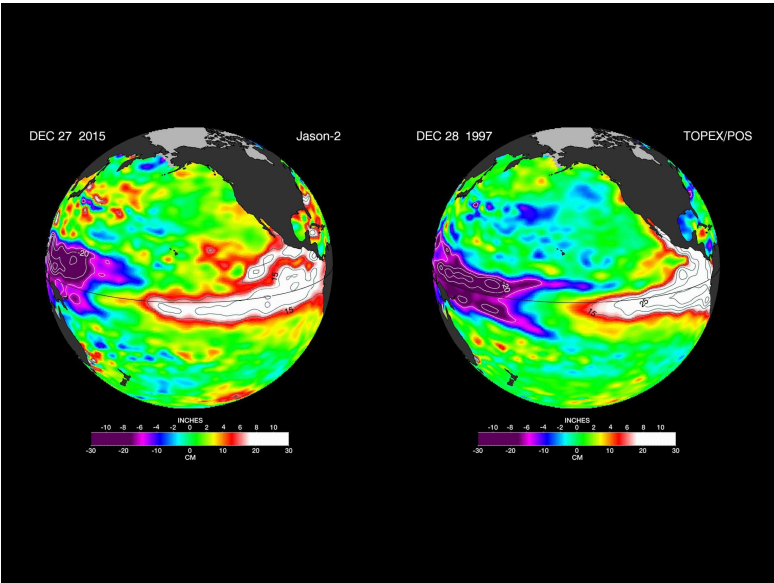
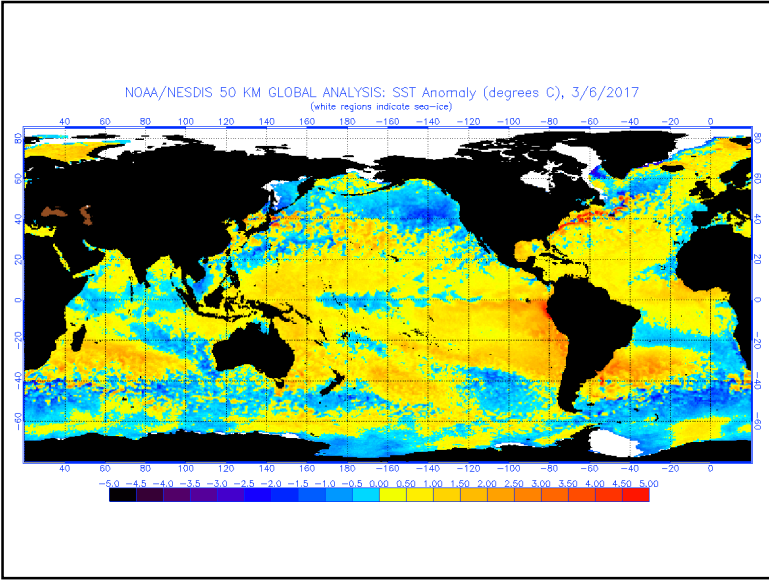
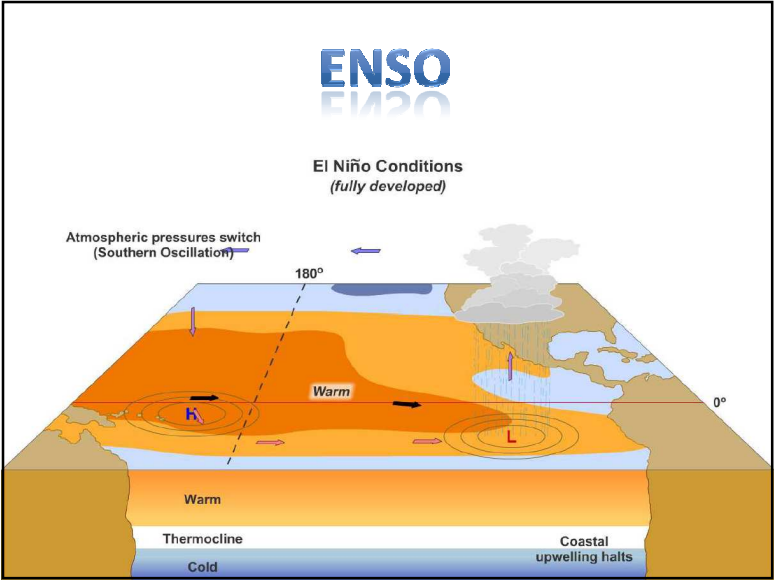
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## ENSO

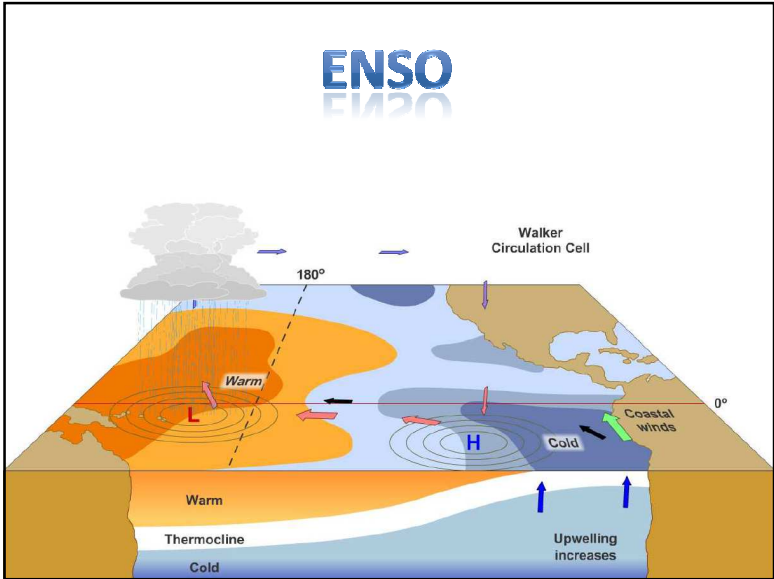
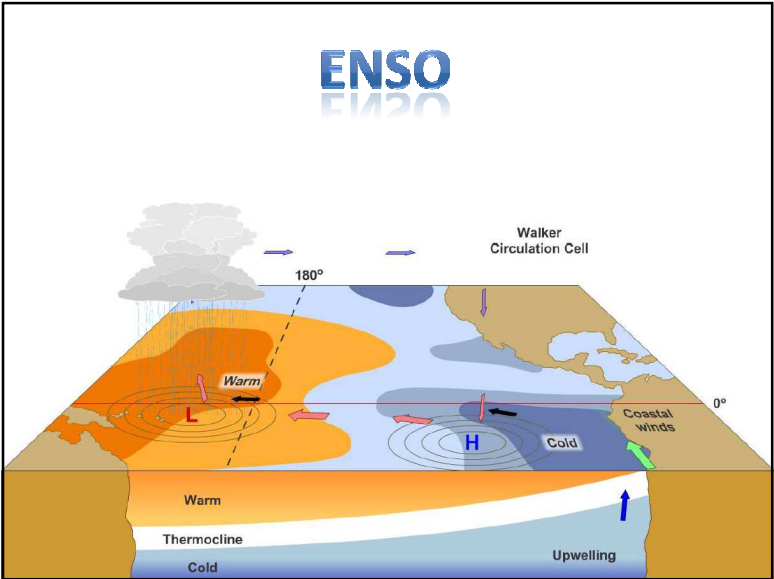
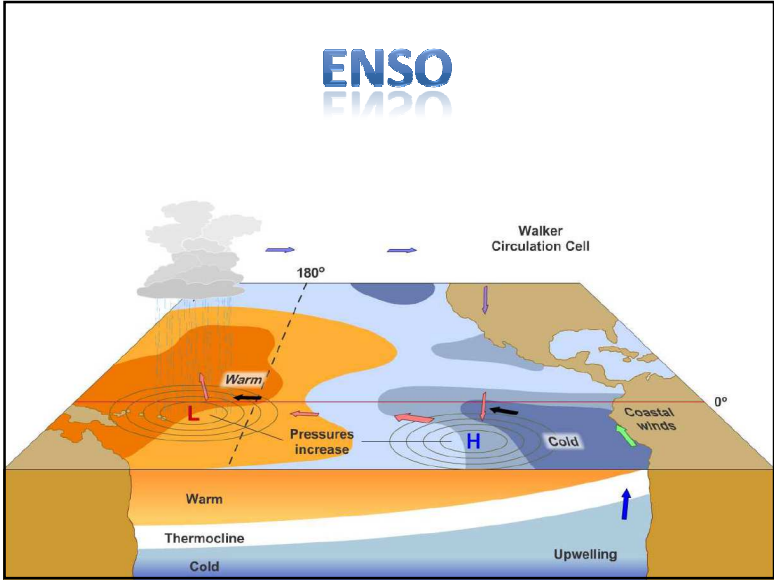
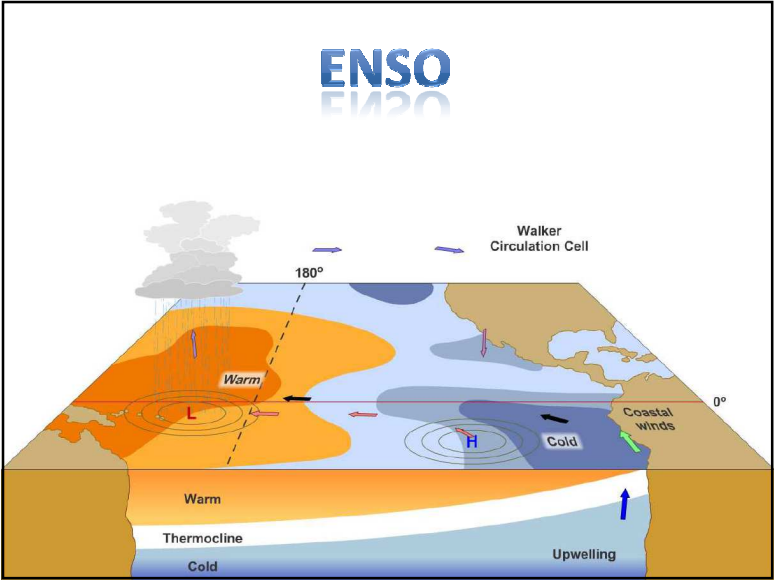


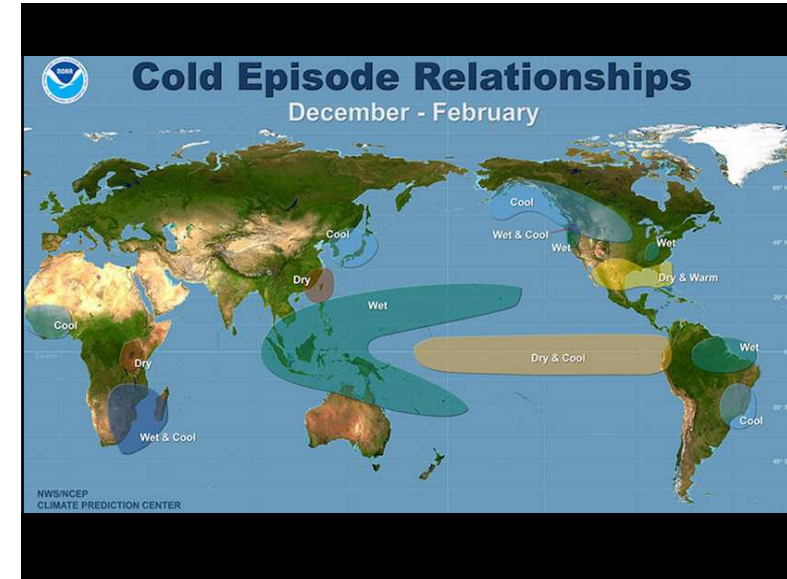
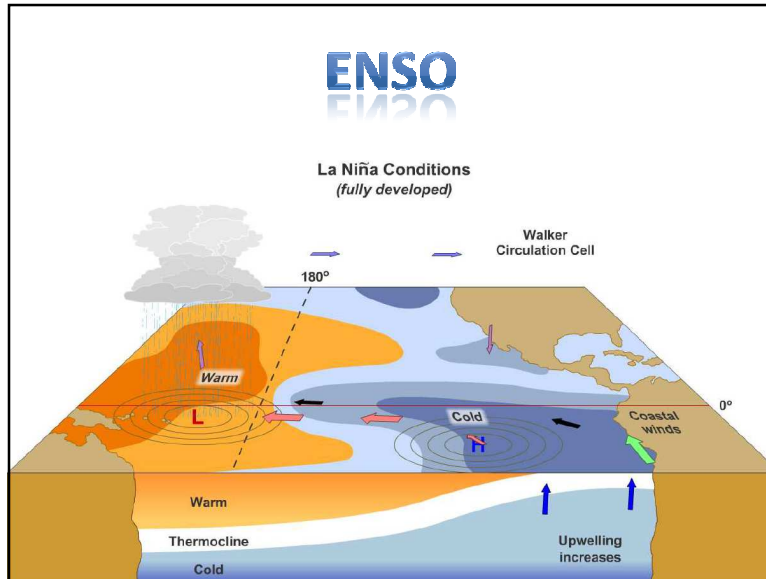










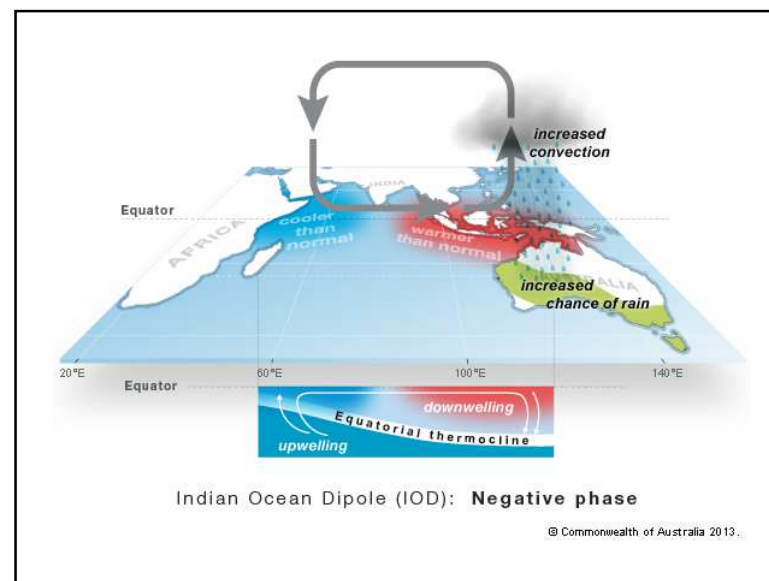
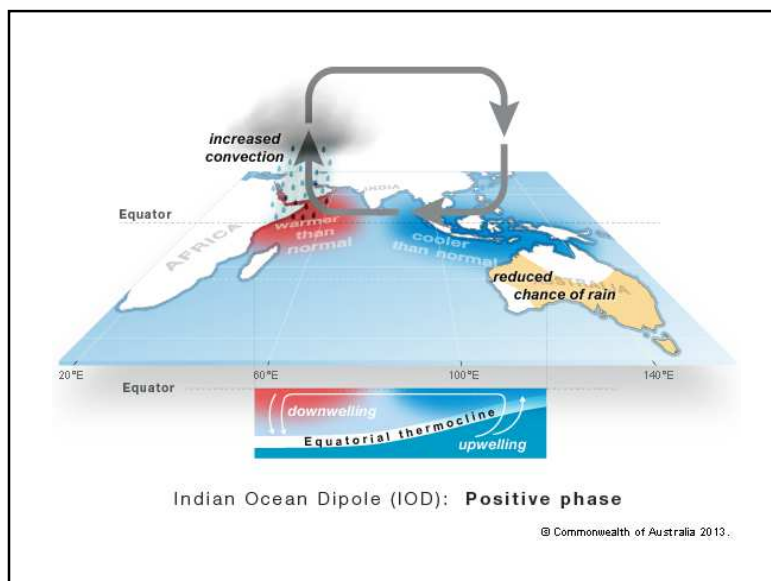
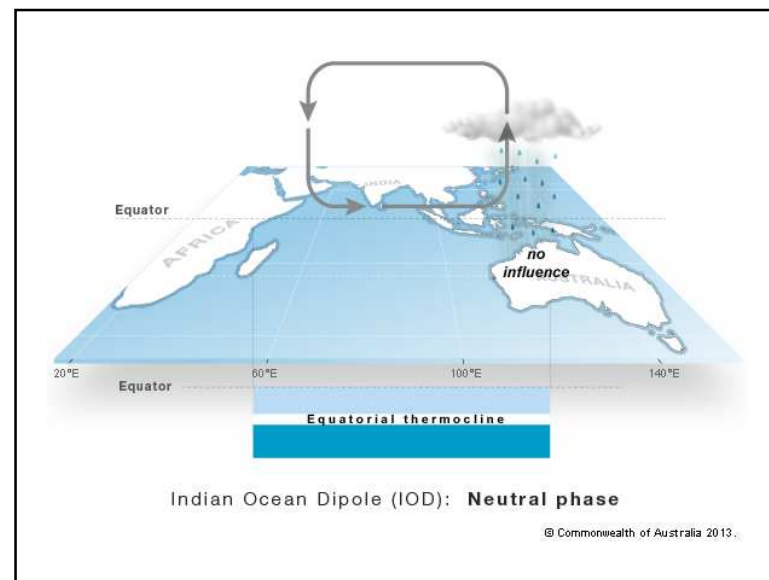
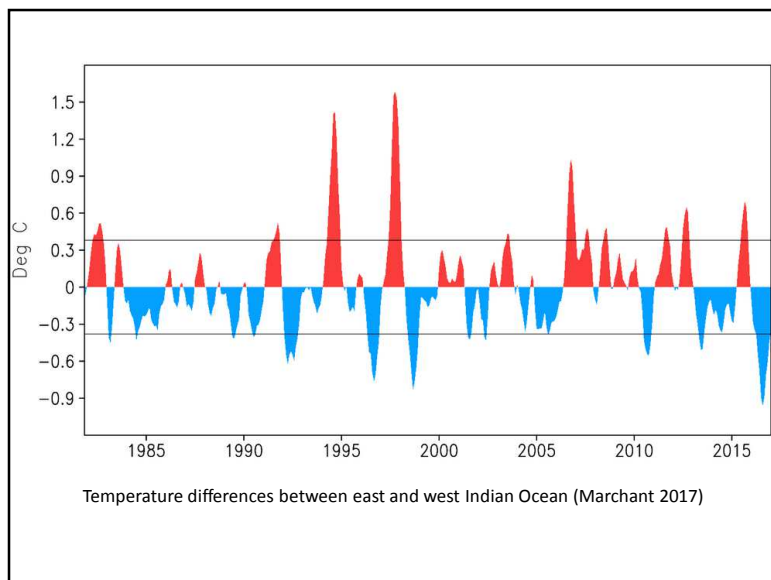


## INDIAN OCEAN DIPOLE

- East African droughts can be partly explained by the IOD; sustained changes in difference between sea surface temperatures of the tropical western and eastern Indian Ocean.
- Three phases: neutral, positive and negative.
- Usually starts early winter (May/June), peaks in spring (August/October) and decays at end of spring.

<http://www.bom.gov.au/climate/iod/>

- Since 1960 to 2013 there have been nine negative and nine positive IOD events (each phase occurs every 3-5 years).
- Farmers in the Horn of Africa rely on moisture from the Indian Ocean to generate the “short rains” that run from October to December and the “long rains” from March to June.
- During the negative phase, winds blowing out to sea makes the air drier than usual; exacerbated drought events.

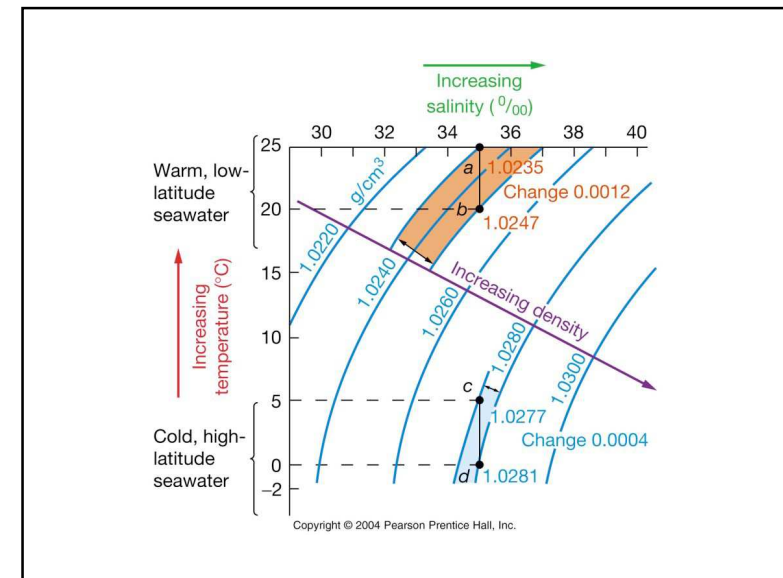
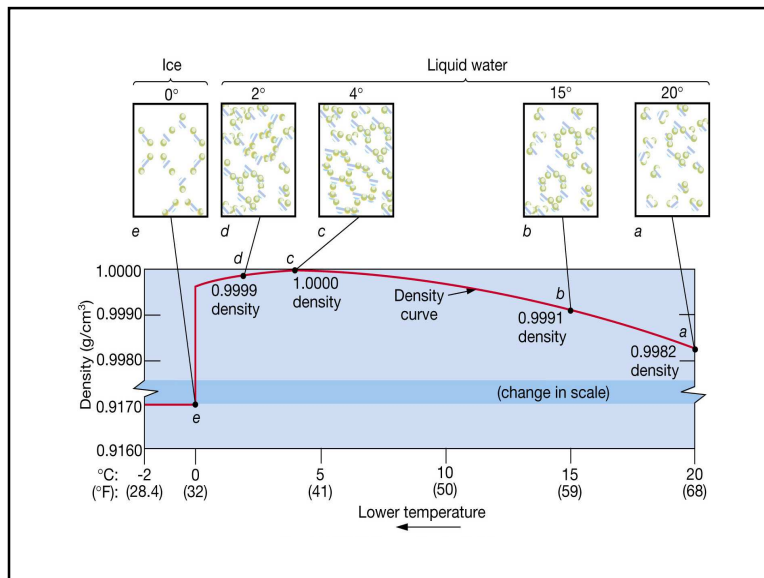


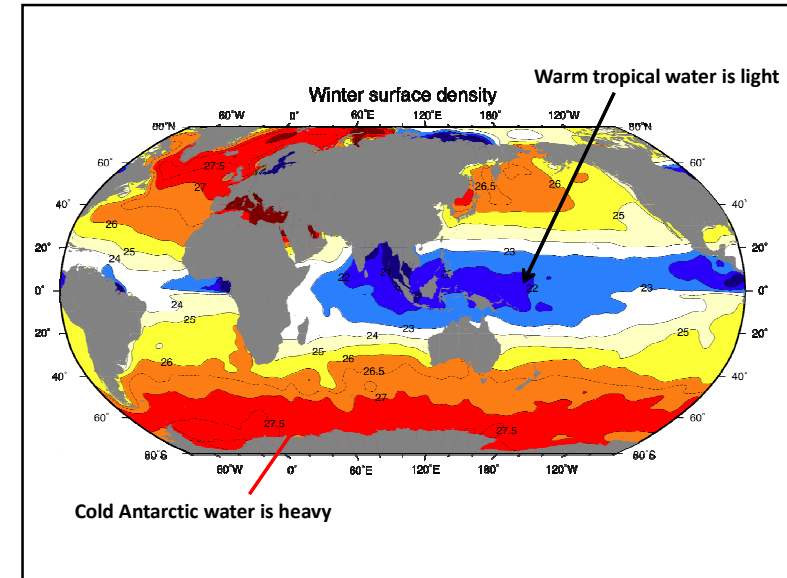
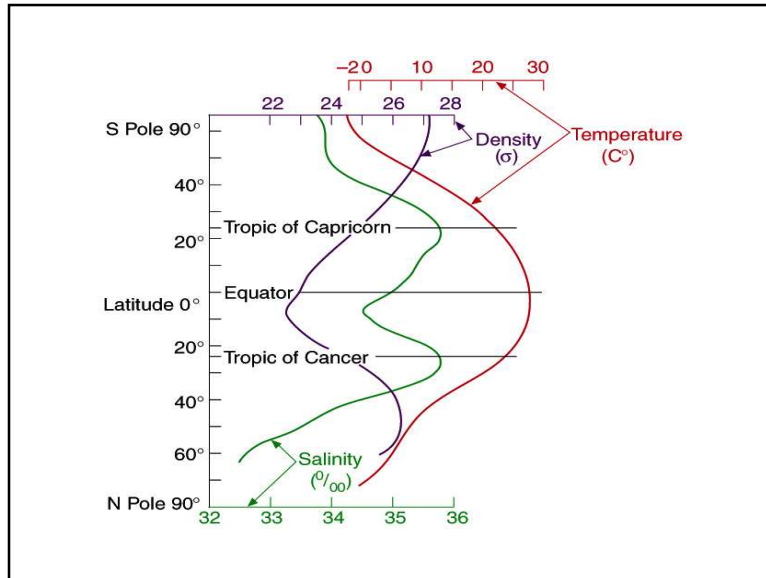


# Thermohaline Circulation & Great Ocean Conveyor

## Water density ( $\text{kg m}^{-3}$ )

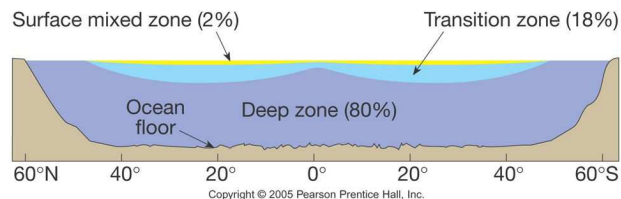
- Dependent on temperature, salinity and pressure
- Differences in fluid density → horizontal layering (stratification); less dense fluid floats on top of denser fluid. E.g. oil floating on water.
  - **Pure water =  $1000 \text{ kg m}^{-3}$**
  - **Sea water =  $1025 \text{ kg m}^{-3}$**





## Three-layered ocean

- 1) Surface (mixed) layer:** ~100-200 m deep. Generally mixed by wind, waves and currents but calm summer conditions may warm this layer creating temporary thermoclines.



- 2) Intermediate (pycnocline) layer:** Below surface layer at ~1000-1500 m. Main thermocline lies within this layer; this is generally permanent and a key feature of open ocean (continental shelf too shallow).
- 3) Deep & bottom layers:** >1500 m. These are similar in being uniformly cold (<4°C).

**NB:** Thermoclines and pycnoclines seldom develop at high latitudes. **WHY?**

