**EAGE** 

TECHNOLOGY AND TALENT SECURE AND SUSTAINABLE ENERGY FUTURE



## Multigeophysical Inversion and Play-Fairway Analysis in the Hengill **Volcanic System**











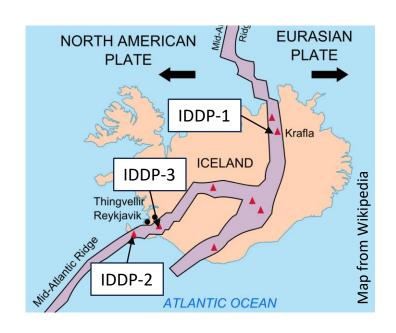


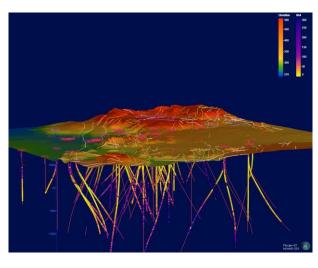
# Multigeophysical Inversion and Play-Fairway Analysis in the Hengill Volcanic System

Ketil Hokstad\*<sup>1</sup>, Jan Inge Tollefsrud<sup>1</sup>, Ásdís Benediktsdóttir<sup>2</sup>, Vala Hjörleifsdóttir<sup>3</sup>, Anne Obermann<sup>4</sup>, Bettina Goertz-Allmann<sup>5</sup>, Nadege Langet<sup>5</sup>

<sup>1</sup>Equinor, <sup>2</sup>Reykjavik Energy (OR), <sup>3</sup>Reykjavik University, <sup>4</sup>ETHZ, <sup>5</sup>NORSAR

### The Iceland Deep Drilling Project (IDDP)





Hengill (figure from OR/ISOR)

#### Objectives of our study

- Contribute to siting IDDP-3
- Multigepophysical inversion (MGI)
- Geothermal PFA
- Part of Geothermica
   DEEPEN project

#### Purpose of IDDP:

- Drill SC geothermal systems (T> 374 oC, P>22 MPa)
- Assess economic potential of SC resources

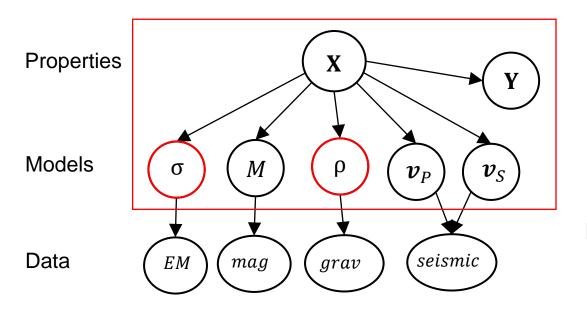


## Multigeophysical inversion (MGI)

IDDP-2 at TD (4659 mMD):

• Inversion: 535 ± 50 °C

Horner plot: ~530 °C

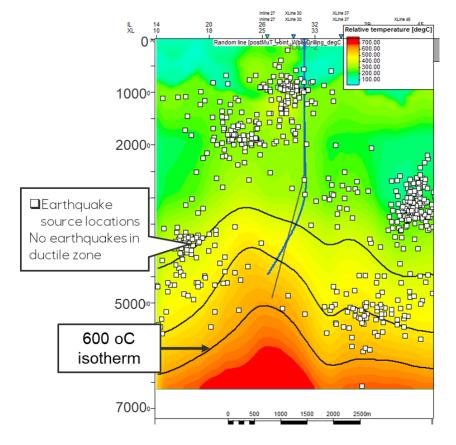


Rock physics

Laws of physics

Bayesian network for MGI

$$p(\mathbf{X}|m_1,\cdots,m_n) = C \prod_{i=1}^n p(m_i|\mathbf{X}) p(\mathbf{X};\lambda)$$

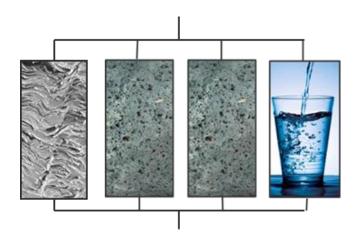


Subsurface temperature from MT and gravity

Hokstad, K. and Tänavsuu-Milkeviciene, K., 2017. Temperature Prediction by Multigeophysical Inversion: Application to the IDDP-2 Well at Reykjanes, Iceland, GRC, Transactions, **42**, 1141-1152.

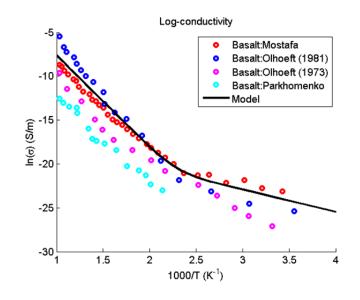


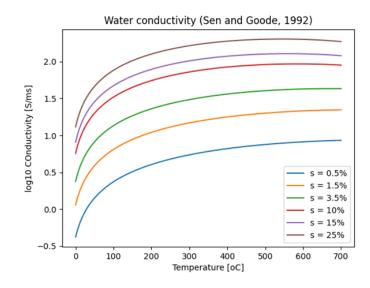
## Electric conductivity (resistivity)





- Igneous rock (basalt)
- Pore-water (salts)
- Clay (cation exchange in smectites)





- Cation exchange capacity
  - Waxman-Smits equation

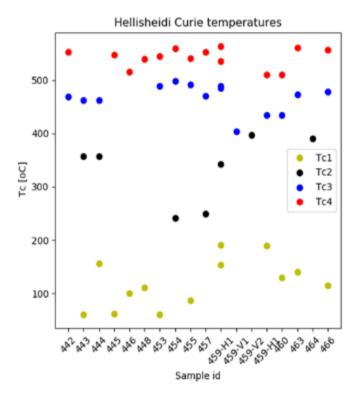




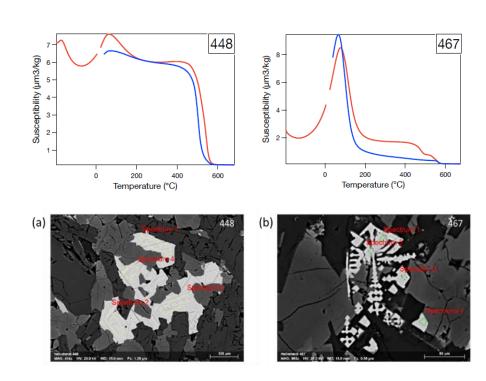
## Magnetization – Samples from Hellisheiði







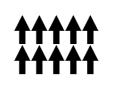
MagLab measurements by Nathan Church and Geertje ter Maat, NTNU



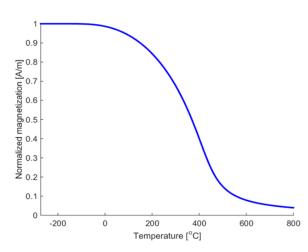
EDS measurements by Claudia Kruber, Equinor



#### Magnetization forward model







Ising (1925) model

Spin-spin and spin-field interactions

$$\varepsilon = -J \sum_{\langle ij \rangle} \sigma_i \, \sigma_j - \mu_B H \sum_j \sigma_j$$

Macroscopic magnetization

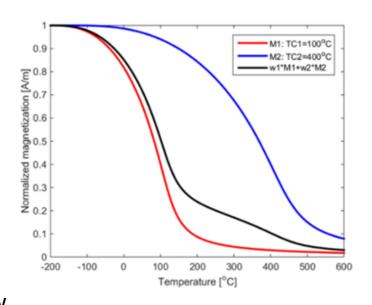
$$\frac{M}{M_0} = \tanh \left[ \frac{T_C M}{T M_0} + \frac{CH}{T M_0} \right]$$

Susceptibility

$$\chi = \frac{M(T,H) - M(T,0)}{H}$$

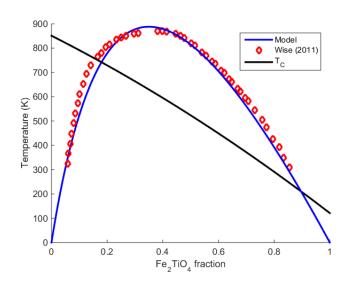
High-temp approximation: Curie-Weiss law

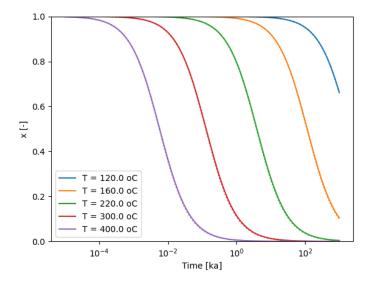
$$\chi \simeq \frac{C}{T - T_C}$$

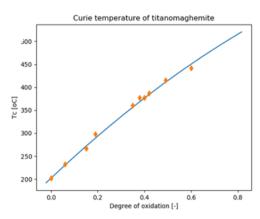




#### Exsolution and oxidation







Curie temperature (Moskowitz and Banerjee;1981)

#### Spinodal exsolution

$$\frac{\partial^2 G}{\partial u^2} = -2W + k_B T \left[ \frac{(1+h)}{u} + \frac{(1-h)}{v} \right] = 0$$

- Curie temperature from Lattard et al. (2006)
- Model calibration using data from Wise et al. (2011)

#### Oxidation: Mixed 1st/2nd order Arrhenius

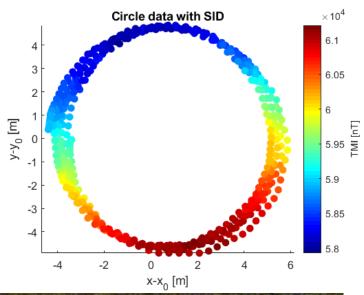
$$\frac{dx}{dt} = -\left(\frac{k_1 - k_2}{x_0}\right)x^2 - k_2x$$
$$k_i = Ae^{-\frac{E_i}{RT}}$$







#### Hellisheiði Hi-Res magnetics





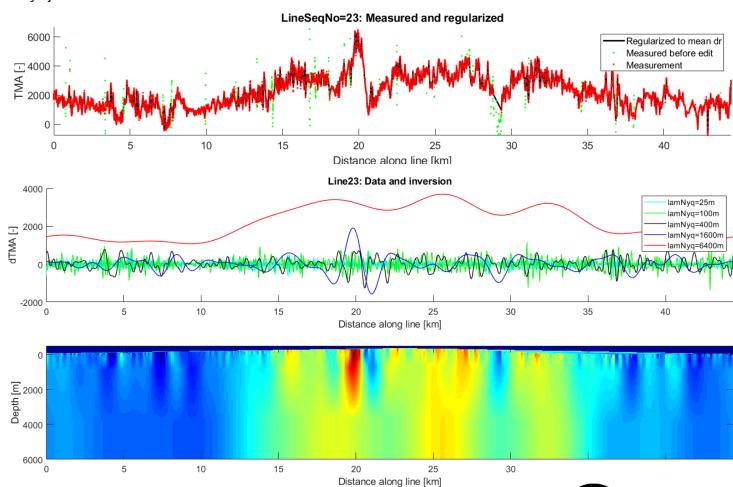






Reykjavik

Selfoss



## P-wave velocity, S-wave veocity, density

Thermal expansion coefficient:

$$\Phi = \int_{T_0}^T \alpha(T')dT' = a_0 [(T - T_0) - 20(\sqrt{T} - \sqrt{T_0})]$$

Hacker et al. (2003)

Porosity, effective stress and clay:

$$v_P(T, \phi, V_{cl}, \sigma') = v_{P_m}(T) - A_P \phi - B_P V_{cl} + C_P (\sigma' - D_P e^{-\beta \sigma'})$$

- Han et al. (1986)
- Eberhardt-Phillips et al. (1989)
- Effective stress is first order
- Temperature is 2nd order

$$\rho_m(T) = \rho_0 e^{-\Phi}$$

$$K_m(T) = K_0 e^{-\delta_T \Phi}$$

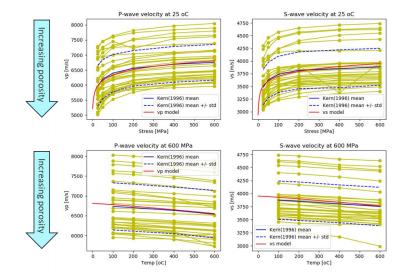
$$\rho_m(T) = \rho_0 e^{-\Phi}$$

$$K_m(T) = K_0 e^{-\delta_T \Phi}$$

$$\mu_m(T) = \mu_0 e^{-\Gamma_T \Phi}$$

$$v_{P_m}(T) = \sqrt{\frac{K_m(T) + (4/3)\mu_m(T)}{\rho_m(T)}}$$

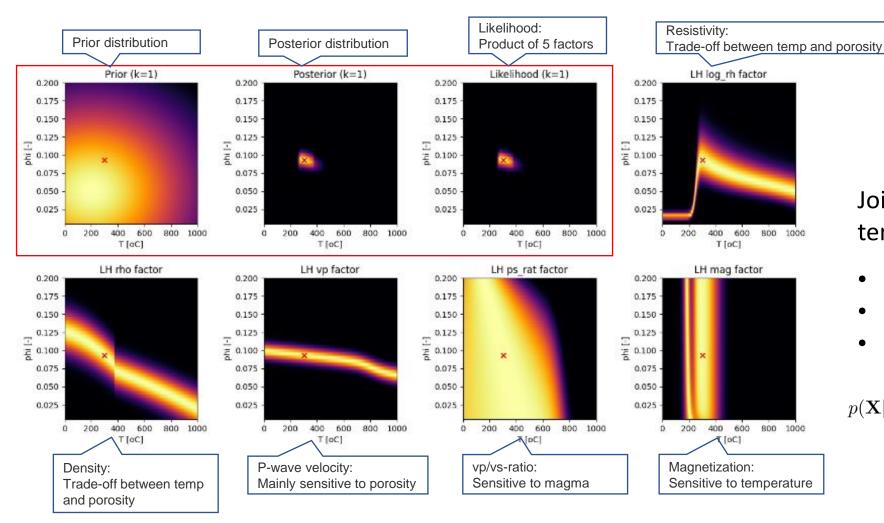
$$v_{S_m}(T) = \sqrt{\frac{\mu_m(T)}{\rho_m(T)}}$$

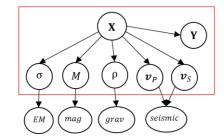


Lab measurements from Kern et al. (1996)



#### Bi-variate multigeophysical inversion





Joint distributions for temperature and porosity

- Posterior
- Likelihoods
- Prior

$$p(\mathbf{X}|m_1,\cdots,m_n) = C \prod_{i=1}^n p(m_i|\mathbf{X}) p(\mathbf{X};\lambda)$$



# Multigeophysical inversion (Hengill)

500

450

400

350

300

250

200

150

405.0

402.5

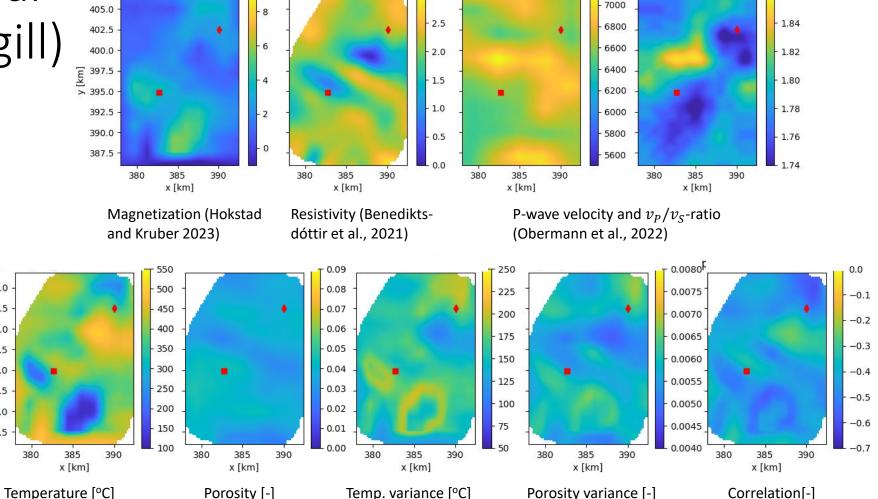
400.0

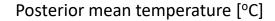
392.5

390.0

387.5

<u>y</u> 397.5 ≤ 395.0





x [km]

385.0 387.5 390.0

post mu T (z=-4000.0m)

405.0

402.5

400.0

397.5 [kw]

395.0

392.5

390.0

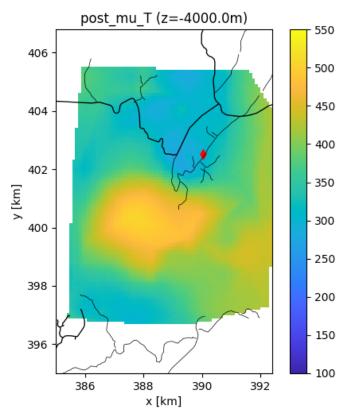
387.5

380.0 382.5

Horizontal cross-sections at z=4000m (MSL)



Multigeophysical inversion (Nesjavellir)



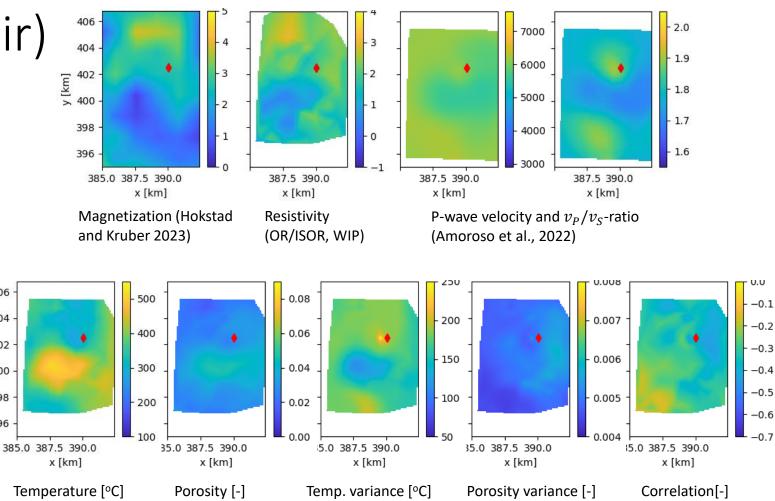
406

[w] 402 5 400

398

396

Posterior mean temperature [°C]



Horizontal cross-sections at z=4000m (MSL)

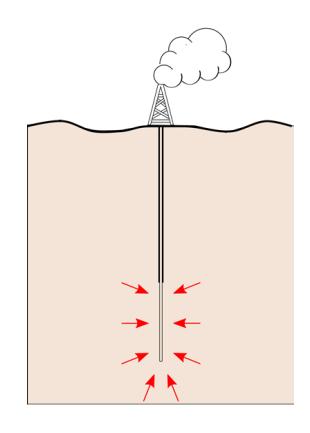


## Play-fairway analysis

#### **Geothermal PFA**

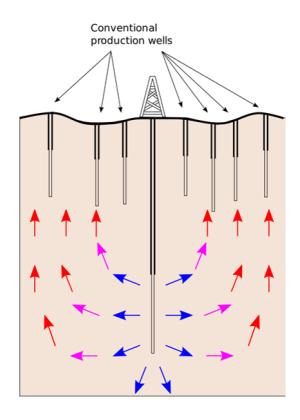
- Heat source
- Permeability
- Fluid
- Seal (SC/SH)







- Extended to SC/SH systems (Kolker et al., GRC, 2022)
- Poux and O'Brien (2020)

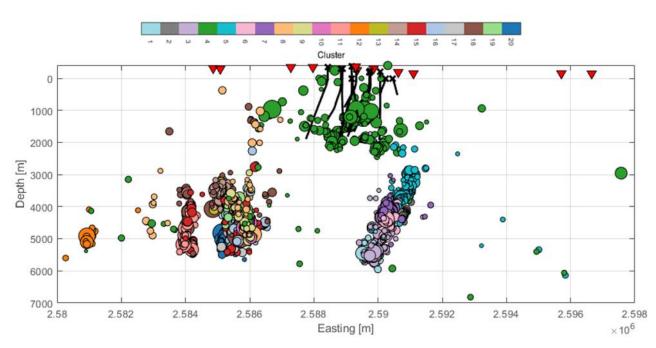


Hybrid strategy for SC/SH systems (from Gunnarsson et al., 2022)

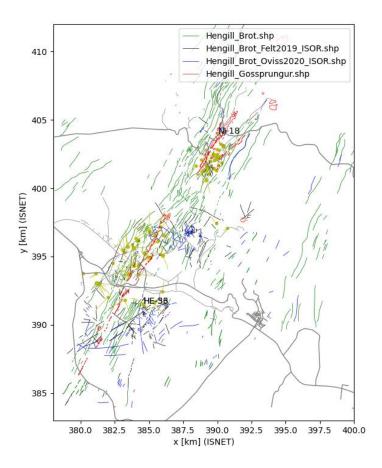


#### Faults and fissures

- Fault density and distance to faults
- Fault intersections and fault-fissure intersections



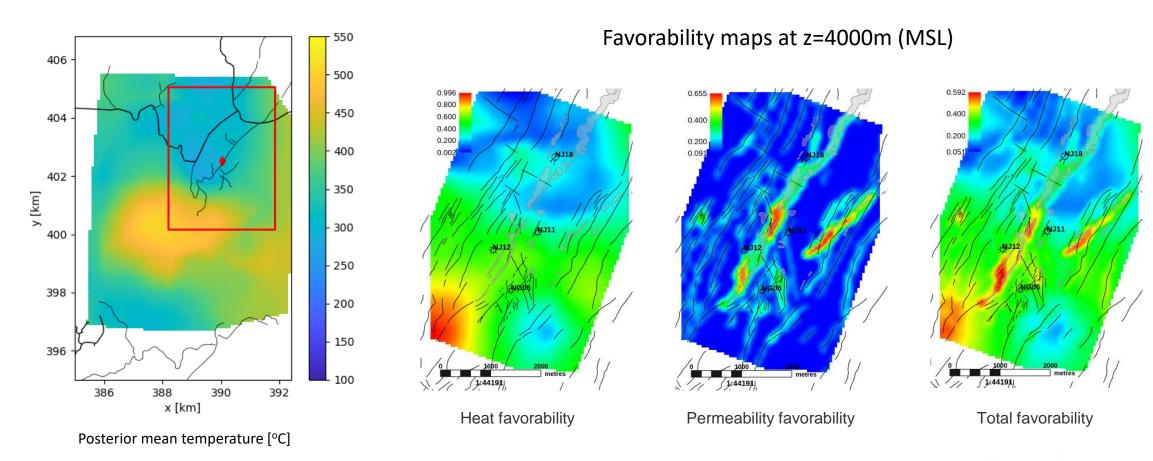
Microseismicity (figure from Goertz-Allmann et al., GRC, 2023)



Faults and fissures (data from OR/ISOR)



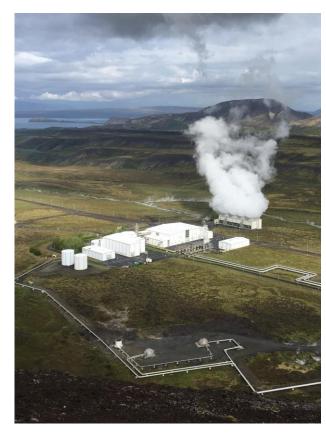
## Play-fairway analysis (Nesjavellir)





#### Conclusions

- Multigeophysical inversion (MGI) for geothermal exploration
  - Utilizing MT, seismic, (gravity), magnetics
- Bivariate inversion for temperature and/or porosity
- Demonstrated on field data and models from Hengill
- MGI used as part of PFA



Nesjavellit power plant





#### Acknowledgements

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- Kristján Óttar Klausen (ISOR)
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- Carsten F. Sørlie, Claudia Kruber (Equinor)

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- Jo Eidsvik, Nathan Church, Geertje ter Maat (NTNU)

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- We thank DEEPEN partners for sharing field data and models from the Hengill geothermal area
- We thank Equinor and OR for permission to publish this work
- Thank you for your attention!























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