Submesoscale processes at Depth in Gulf of Mexico

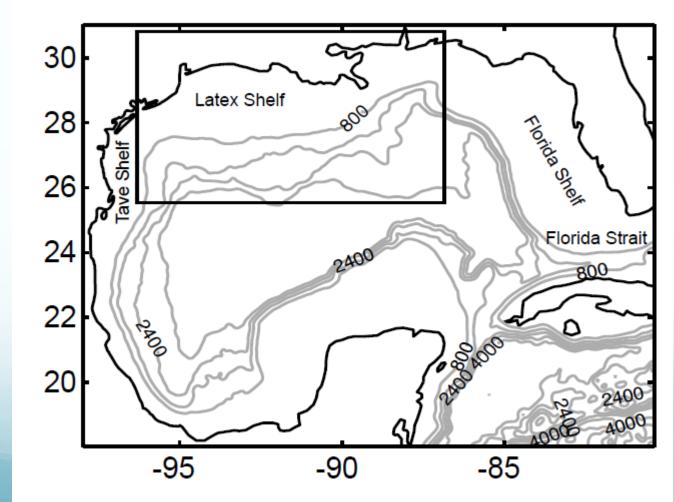
Keshav Joshi, Annalisa Bracco, Hao Luo, Jim McWilliams





Model Setup: ROMS-AGRIF 2.1

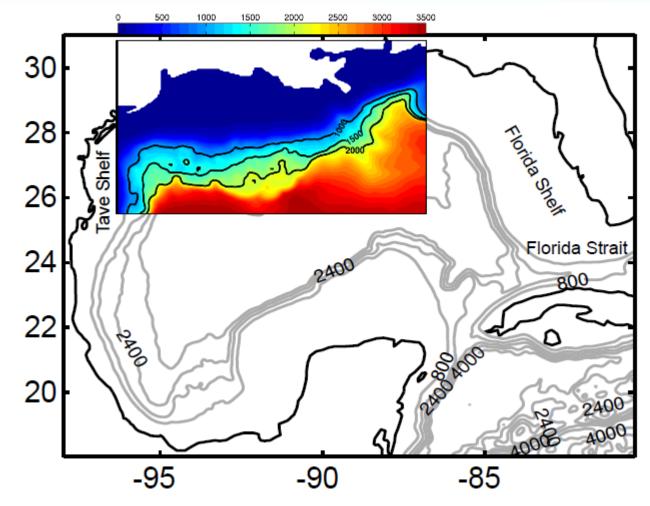
- Domain:
 - LR -> 5km (98.0W~80.4W) & (18.0N~31.0N)
 - **HR** -> 1.6km (96.3W~86.9W) & (25.4N~30.7N)
- 70 vertical terrain-following layers
- HYCOM GoM 31.0: Boundary and Initial Condition
- ECMWF ERA-interim reanalysis (6-hour wind stress and daily heat fluxes)
- Run dates
 Jan 2010 Dec 2012





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Model Setup: ROMS-AGRIF 2.1 Vertical Layers along 26N

-3500 -96

-95

-94

Domain: LR -> 5km (98.0W~80.4W) & (18.0N~31.0N) -500 • **HR** -> 1.6km (96.3W~86.9W) & (25.4N~30.7N) -100070 vertical terrain-following layers HYCOM GoM 31.0: Boundary and Initial -1500 Condition **ECMWF ERA-interim** reanalysis (6-hour wind -2000 stress and daily heat fluxes) Run dates Jan 2010 - Dec 2012 -2500-3000

-93

-92

-91

-90

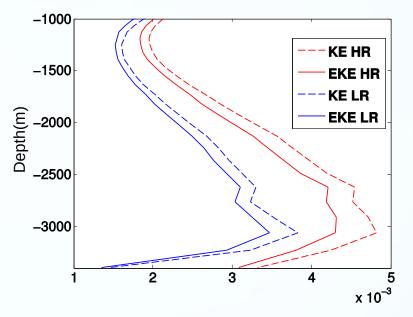
Mean Flow

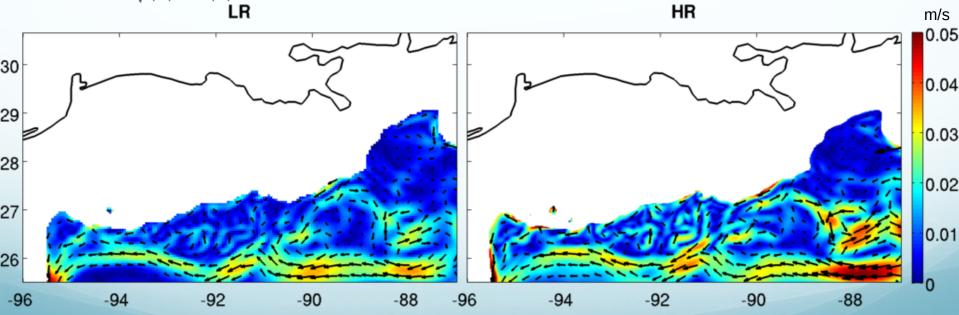
Mean velocity field vectors over 3 year run

contours from mean speed

$$= \sqrt{\langle u \rangle^2 + \langle v \rangle^2}$$
LR

Energy Profile below 1000m (right)





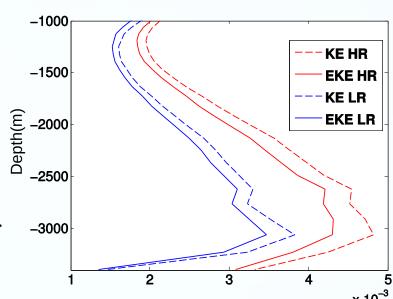
EKE variability

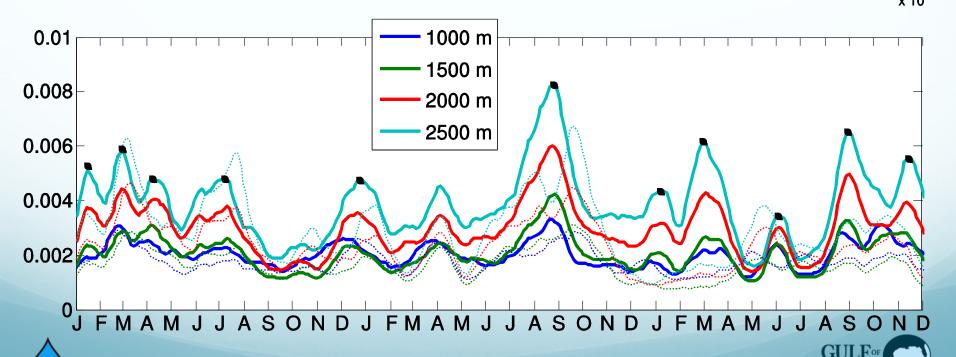
- greater EKE variability in HR
- black dots: consistent separation cyclone off Mississippi Fan

$$KE = \frac{1}{2}(u^2 + v^2)$$

$$u' = u - \langle u \rangle; \ v' = v - \langle v \rangle$$

$$EKE = \frac{1}{2}(u'^2 + v'^2)$$





Vorticity generation

Seasonal Vorticity PDF @1000m LR – lighter dashed line HR – darker line

Vorticity:
$$\frac{\zeta}{f} = \left(\frac{dv}{dx} - \frac{du}{dy}\right)/f$$

HR: $\zeta/f \sim 1$ generated in

- Separation cyclone
- · Boundary layer instability
- Sigsbee escarpment

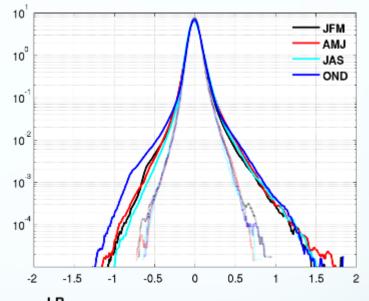
Skewness

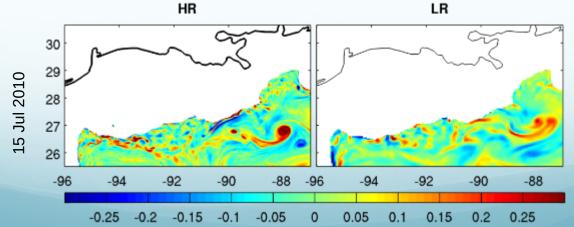
@1000m

HR - 0.95; LR - 0.55

@2000m

HR - 1.58; LR - 1.10

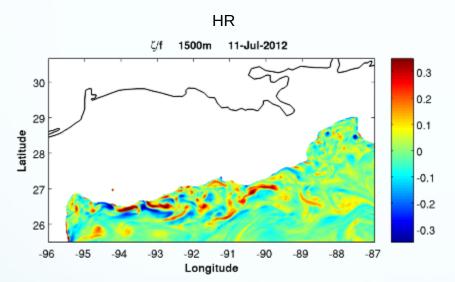


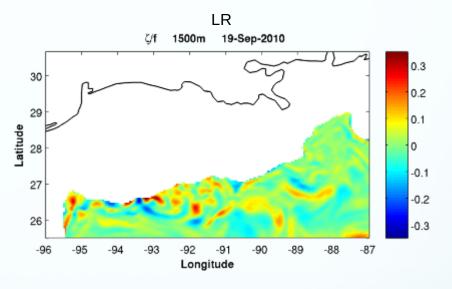


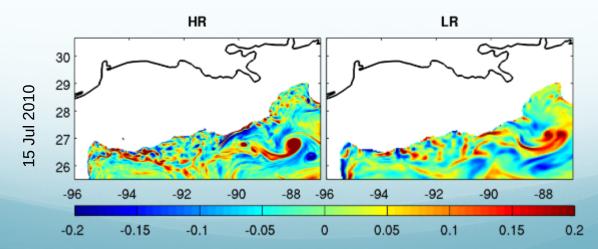




Vorticity generation



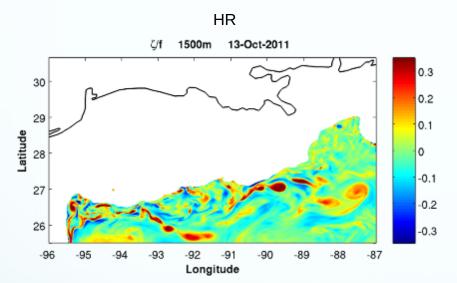


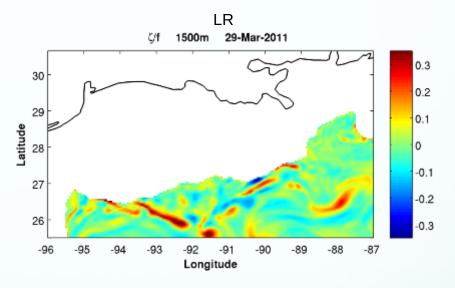


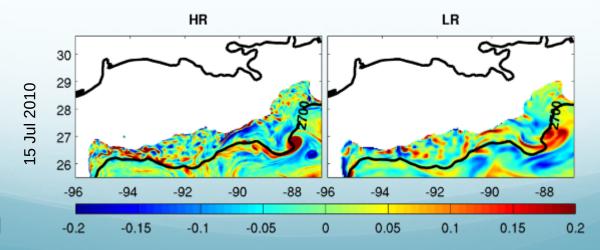




Vorticity generation







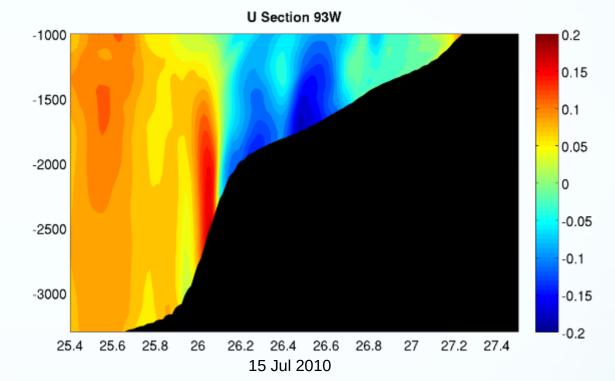


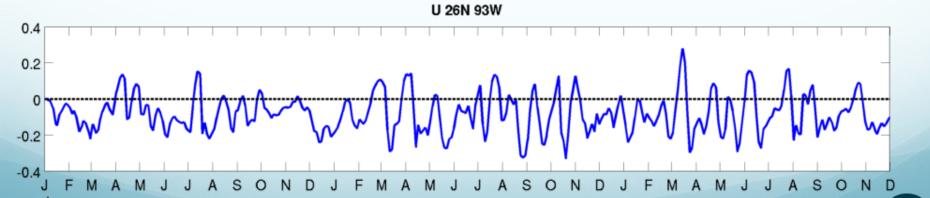


Boundary Layer

- generates cyclonic vorticity filament consistently
- better resolved boundary layer at 1.6km, higher population of eddies

$$\zeta = \frac{dv}{dx} - \frac{du}{dy} \sim -\frac{du}{dy} \sim -\frac{du}{dy}$$

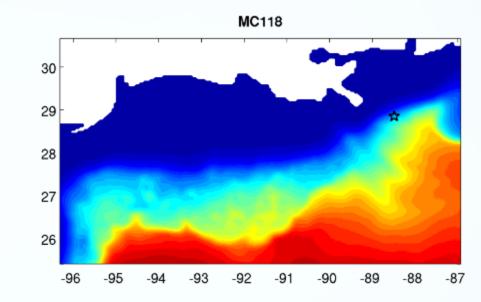


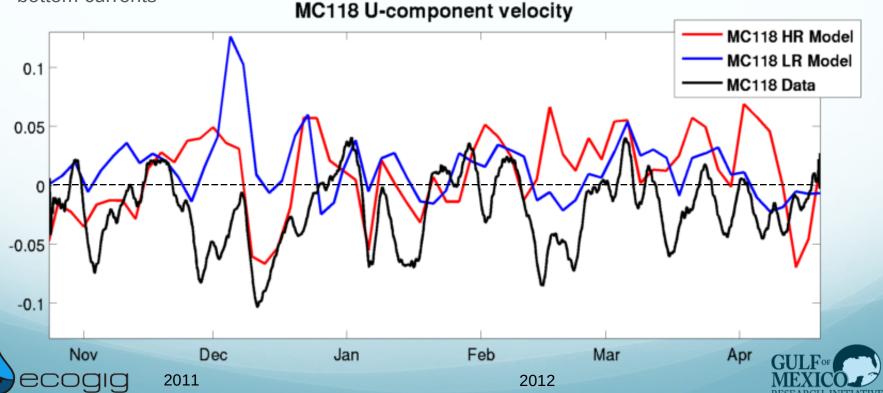


U-variability: data

Data gathered every 2 minutes by lander at MC118 location, natural seep site

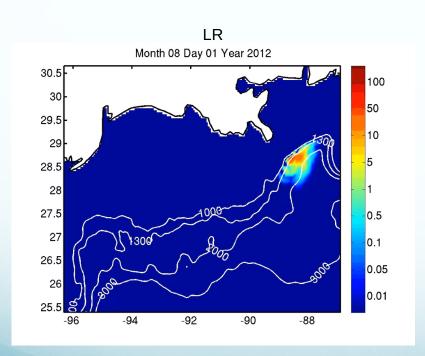
3-day lowpass applied to all data More data needed for validation of bottom currents

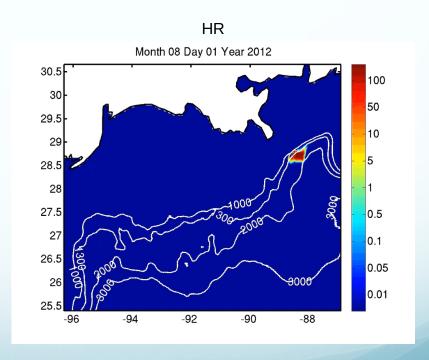




Tracer Evolution

Tracers released on Jul. 28 2012, close to Macondo Blowout site, advected for 4 months

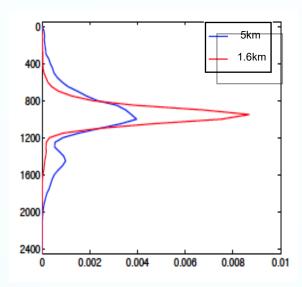






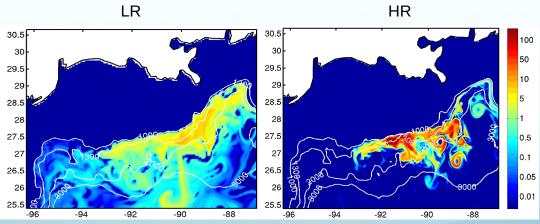


Tracer vertical distribution



LR tracer diffuses more both horizontally and vertically

HR traps tracer in submesoscale structures along the continental slope

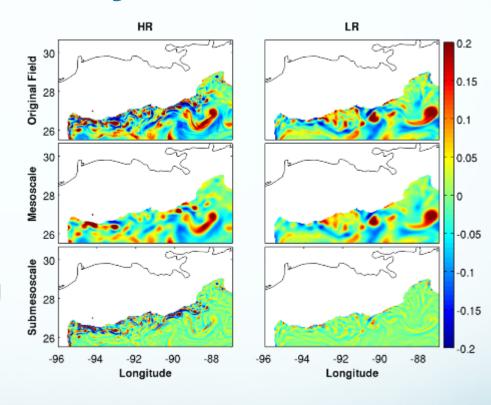






Summary

- 1.6 km run
 - better resolves boundary layer and consequent instabilities
 - produces high vorticity populations: $\zeta/f \sim 1$, at depth as a result
- Submesoscale resolving models important for predicting transport and mixing of tracer or particles suspended in the ocean, even at depth



15-Sep-2011



