# Run 200

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require(MASS)	
require(rstan)	
<pre>options(mc.cores = parallel::detectCores())</pre>	
rstan_options(auto_write = TRUE)	

# Set up simulation

Like Run101 except:

• von Karman topography on main fault (and smaller elements on main fault 50 m instead of 100 m)

#### Build fault model

Make a simple fault model consisting of a single, vertical, strikeslip fault with von Karman topography, and then add random small faults around it.

```
L = 5e3 # fault length

W = 3e3 # fault vertical extent

y0 = -L/2 # center fault on origin

zmax = -2e3 # top edge of fault

1 = w = 50 # element length and width

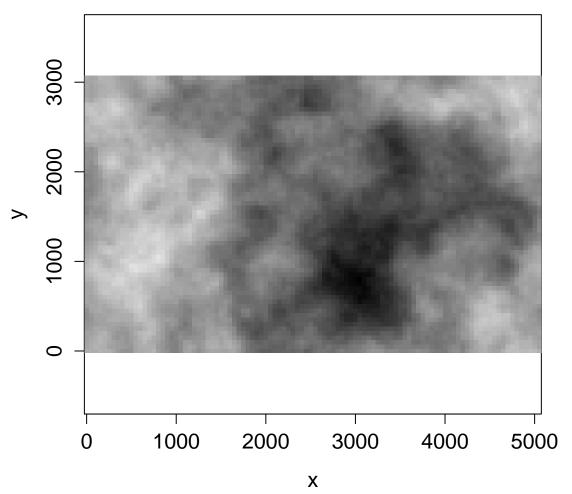
strike = 0

dip = 90

rake = 0

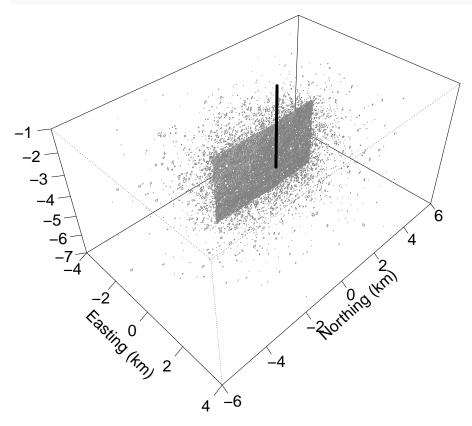
ddot = 0
```

# dx = 50; ax = 500; ay = 500; H = 0.9; ag = 0



#### Plot of resulting fault system

```
# injection point that we will use below
xi = matrix(c(500, 0, -3.5e3),1, 3) # well location: 500 m to east of center pt. of main fault
p = plotFault3d(flt2, theta=45, phi=40, xlim=c(-4,4), ylim=c(-6,6), zlim=c(-7,-1), lwd=0.5)
lines(trans3d(c(xi[1,1], xi[1,1])/1e3, c(xi[1,2], xi[1,2])/1e3, c(0, xi[1,3])/1e3, p), lwd=3)
```



#### Set up injection history and resulting stressing file

Make a simple injection history and RSQSim external stressing file from it. Note tha maxErr parameter of 100 (Pa), default is 1, which is much more precise than necessary. This helps both the execution time and the file size of test.externalStress.txt (ends up being 106 MB compared to 455 MB in previous runs, despite having twice as many total fault elements here).

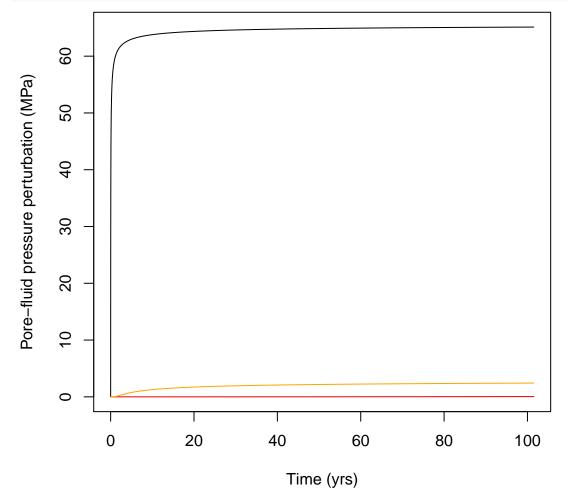
```
yr = 365.25*86400
injHist = makeInjectionHistory(xi=xi) # use defaults for all hydrologic params

t = findSampleTimes(flt2, injHist, tmax=100*yr, eps=0.05, minh=1000, maxErr=100)
writePoreFluidStressingHistoryFile(flt2, injHist, t, "test.externalStress.txt")

s = readExternalStressingHistoryFile("test.externalStress.txt", flt2)
```

#### Plots of pressure histories

The closest element, the farthest element, and the median distance element.



#### Initial stresses

```
set.seed(234)
tau00.main = 63.2 + 1
tau00.random = 63
sdtau.main = 0
sdtau.random = 1
main = which(flt2$segName == "NA")
random = which(flt2$segName != "NA")
tau0 = rep(NA, flt2$np)
tau0[main] = rnorm(length(main), tau00.main, sdtau.main)
```

```
tau0[random] = rnorm(length(random), tau00.random, sdtau.random)
write(tau0, "initTau.txt", ncol=1)
```

#### Run simulation

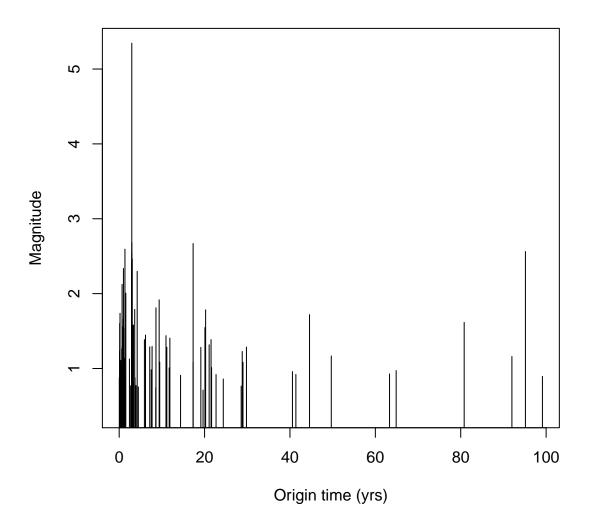
Copy Run3's parameter file and make adjustments.

#### Results

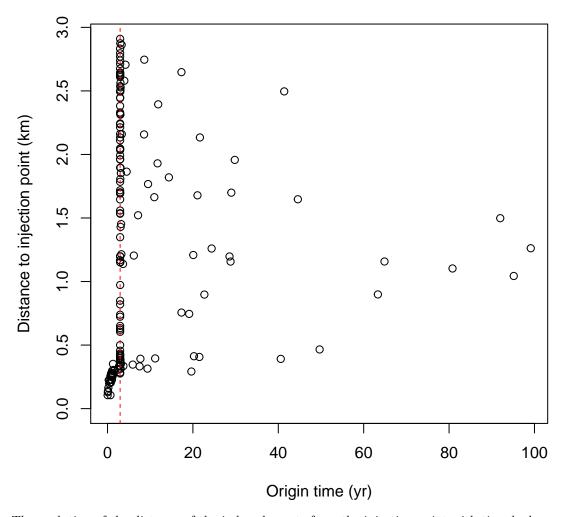
```
eqs = readEqs(paste("eqs.", params$outFnameInfix, ".out", sep=""))
if (max(eqs$t0) > max(s$t)) {
   eqMax = min(which(eqs$t0 > max(s$t))) - 1
   eqs = readEqs(paste("eqs.", params$outFnameInfix, ".out", sep=""), eqMax = eqMax)
}
```

#### Magnitude-time plot

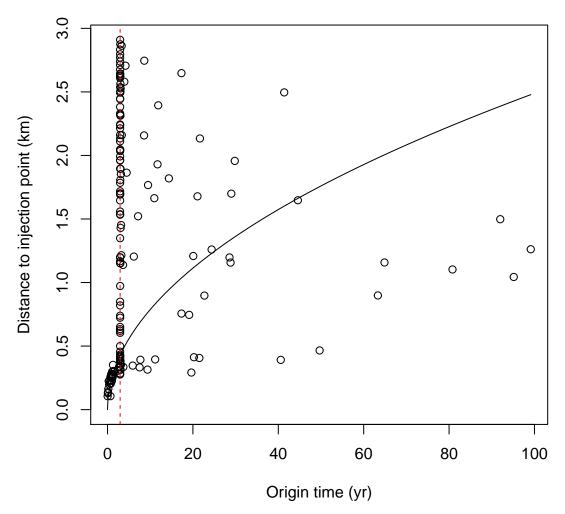
```
plot(eqs$t0yr, eqs$M, typ='h', xlab="Origin time (yrs)", ylab="Magnitude")
```



# Hypocenter-well distance



The evolution of the distance of the induced events from the injection point with time looks roughly like the expected  $\sqrt{t}$  (until the M5+). It is not clear what the exact relationship between the coefficient of  $\sqrt{t}$  and the hydraulic parameters should be. Shapiro et al. [1997] advocate for  $x = \sqrt{4\pi\kappa t}$ , where  $\kappa$  is the diffusivity. The argument of the erfc in our pore-fluid pressure expression is equal to one at  $x = \sqrt{4\kappa t}$ , but there is also a  $|x|^{-1}$  factor in front of the erfc. Adding in the least-squares fit to the pre-M5 events:



The  $\sqrt{t}$  is not a great fit: event times are systematically earlier at near distances and later at farther distances than the best fit  $\sqrt{t}$  curve. I would attribute this to the  $|x|^{-1}$  factor in the pore-fluid pressure expression. It could also be due to RSF evolution effects. If we ignore that and ask what value of  $\kappa$  is implied by the Shapiro et al. [1997] expression or just setting the argument of the erfc to 1, we get  $1.6 \times 10^{-4}$  and  $4.9 \times 10^{-4}$  m<sup>2</sup>/s, respectively. These are both about an order of magnitude lower than the actual diffusivity used in generating the pore-fluid pressures  $(0.008 \text{ m}^2/\text{s})$ . Note that this difference here between the actual diffusivity and that estimated by assuming  $x = \sqrt{4\pi\kappa t}$  is not necessarily relevant to the data in Shapiro et al. [1997]. In that paper the diffusivity is much higher and the time scales are much shorter (a few days), so that RSF evolution effects may be less important.

Elizabeth mentions that some papers of Talwani estimates the diffusivity by fitting curves of the form  $x = \sqrt{\kappa t}$ . Also said that might have been for fluid pressures diffusing in 2-d (e.g. if a fault is a high permeability conduit) instead of in the 3-d bulk. If we ignore the 2-d/3-d issue and just estimate  $\kappa$  from curves of that form, we get  $0.002 \text{ m}^2/\text{s}$ , just a factor of two less than the simulation value.

### Pore-fluid pressure at failure

As an indication of how important RSF evolution effects may be, let's look at the pore-fluid pressure one each hypocentral element just before it failed. For friction with a fixed failure stress (e.g. slip-weakening) and uniform initial stresses as we have here, this would be expected to be the same for all events as long as the stress transfer between the elements is neglible (which it will be in this fault system everywhere except the main fault). However, with RSF, the more distant, later-failing elements will fail at a lower pore-fluid

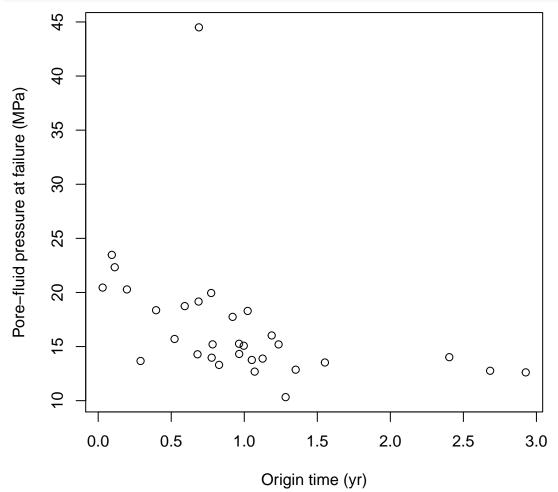
pressure because (once they are above steady-state) they will have been weakening.

Note that the instantaneous pore-fluid pressure increment that would cause immediate failure is given by that P which solves

$$V^* \exp\left(\frac{\tau - \mu_0(\sigma_0 - P)}{a(\sigma_0 - P)}\right) \left(\theta_0 \left(\frac{\sigma_0 - P}{\sigma_0}\right)^{\alpha/b} \frac{V^*}{D_c}\right)^{-b/a} = V^{\text{EQ}}.$$

For this simulation, this comes out to about 26.9 MPa.

Below I plot the pore-fluid pressure perturbation at failure for all events before the M5 (after which stress transfer dominates).



These are indeed all less than 26.9 MPa and are smaller at later times / farther distances, but very scattered here because of the variable pre-stress.

So, for now, it is an open question how much of the departure from  $\sqrt{t}$  behavior is due to the  $|x|^{-1}$  and how much due to RSF evolution effects.

## Estimation of Omori p-value and aftershock fraction n

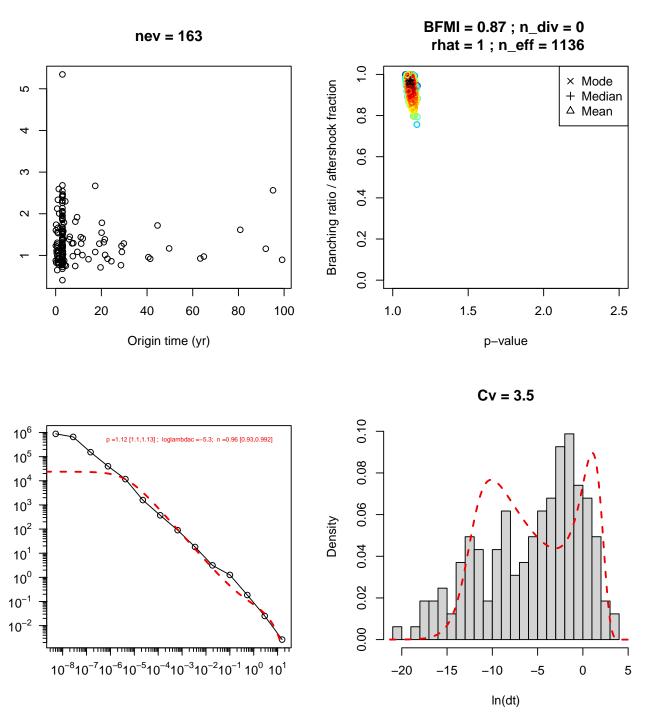
See templateFamilies.[Rmd|pdf]

```
use = 1:length(eqs$t0)
alpha = 0.32 # so 68% CIs
c = 90 # seconds, needs to match dSS2007.krd.fixedloglambdac.stan
st = do.stan(eqs, use=use, alpha=alpha, c=c)
```

#### Plotting function definition

The functions to make plots like in templateFamilies.pdf; modified from there. Not printing them out because they take a lot of page space.

#### Actually make the plots

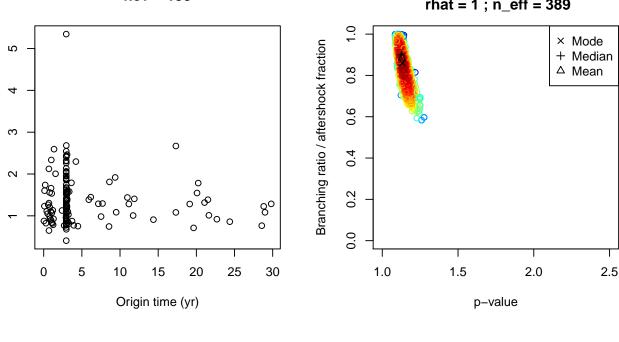


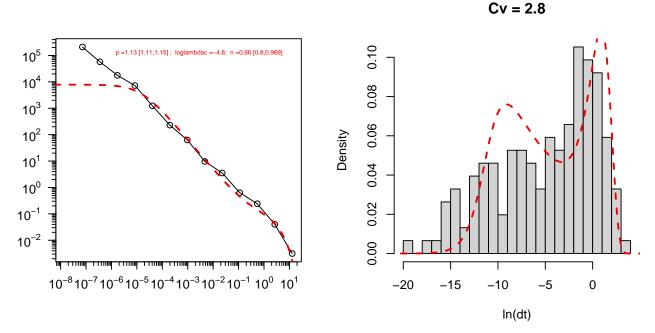
This estimates the aftershock fraction to be 0.96. Just given the data asked for my opinion, I probably would have counted the 97 events that occurred in the couple of years after the M5+ as aftershocks and said the aftershock fraction was 97/163 = 0.6.

#### Leaving out very late events

Perhaps those last 10 events way off by themselves are biasing things? What happens if we leave them out?

```
use2 = 1:(length(eqs$t0)-nout)
st2 = do.stan(eqs, use=use2, alpha=alpha, c=c)
```





Now estimated aftershock fraction is 0.86, much closer.

Note that the estimate of *Hainzl et al.* [2006] would give 0.92 and 0.87, when using all events and all the but last 10 events, respectively.

Both the estimated aftershock fraction from fitting the Saichev and Sornette [2007] functional form and

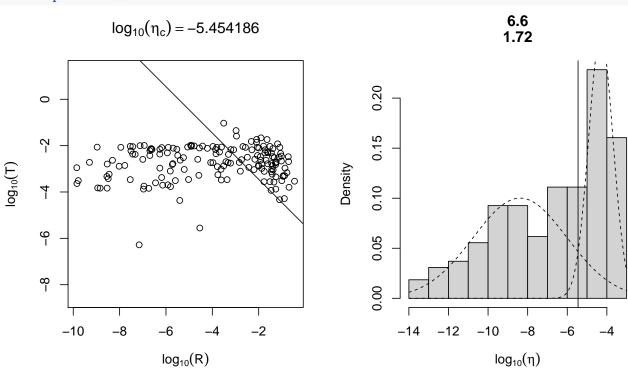
the  $C_v$  are very affected by these outliers. If we used the 1-norm version of the  $C_v$ , call it  $C_{v1}$ ,  $C_{v1} = \text{smad}(dt)/\text{median}(dt)$ , it would be essentially completely unaffected: including the 10 events gives 1.4825964 and excluding them gives 1.4825888.

# Aftershock fraction ala Zaliapin and Ben-Zion [2016]

```
sderrh = 10
z = zaliapin(eqs, use=use, sderrh=sderrh, est.se=TRUE)
z2 = zaliapin(eqs, use=use2, sderrh=sderrh)
```

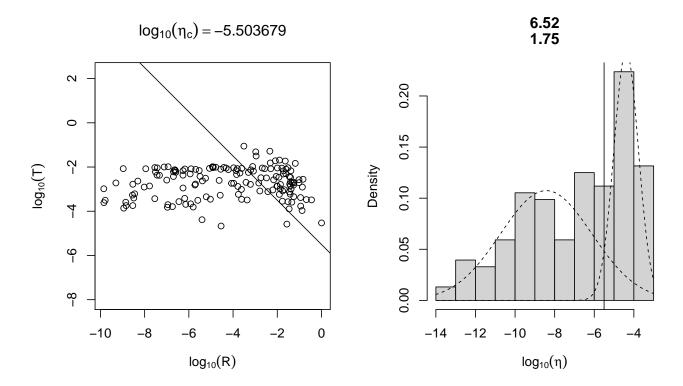
#### Using all events

```
source("../mkZaliapinPlots.R")
mkZaliapinPlots(z)
```



#### Excluding late events

mkZaliapinPlots(z2)



## Summary

- Full-main-fault event at 3 yrs
- Number of events
  - total: 163
  - before full-main-fault event: 31
  - after full-main-fault event: 131
- Estimated aftershock fractions
  - naive count: 0.6
  - SS2007
    - \* using all events: 0.96
    - \* excluding last 10 events: 0.86
  - H2006
    - \* using all events: 0.92
    - \* excluding last 10 events: 0.87
  - ZB2016
    - \* using all events: 0.54
    - \* excluding last 10 events: 0.59

# To do

- how to get more realistic time-scales and pore-pressures?
- different uniform stresses on main fault

- variable stresses on main fault
- different forms / amplitudes of variable stresses on random faults
- denser random faults
- well closer and farther from main fault
- higher and lower injection
- how to get aftershocks other than whole-fault event on main fault?
  - heterogeneous stresses on main fault?
  - more realistic fault connectivity(?) for random faults?

## References

Hainzl, S., F. Scherbaum, and C. Beauval (2006), Estimating background activity based on interevent-time distribution, *Bulletin of the Seismological Society of America*, 96(1), 313–320.

Saichev, A., and D. Sornette (2007), Theory of earthquake recurrence times, *Journal of Geophysical Research:* Solid Earth, 112(B4).

Shapiro, S. A., E. Huenges, and G. Borm (1997), Estimating the crust permeability from fluid-injection-induced seismic emission at the ktb site, *Geophysical Journal International*, 131(2), F15–F18.

Zaliapin, I., and Y. Ben-Zion (2016), Discriminating characteristics of tectonic and human-induced seismicity, Bulletin of the Seismological Society of America, 106(3), 846–859.