# Run 7

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require(MASS)	
require(rstan)	
<pre>options(mc.cores = parallel::detectCores())</pre>	
rstan_options(auto_write = TRUE)	

# Set up simulation

Difference from Run3 is higher initial stress on the random faults. Same fault system and injection history going back to Run2, so can reuse fault system and external stressing file.

### Copy fault model

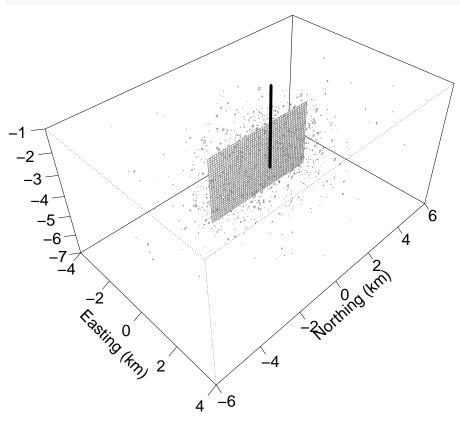
Make a simple fault model consisting of a single, vertical, planar strikeslip fault, and then add random small faults around it.

```
file.copy(file.path("..", "Run2", "test.addRandom.10.flt"), ".")
flt2 = readFault("test.addRandom.10.flt")
```

### Plot of resulting fault system

```
# injection point that we will use below xi = matrix(c(500, 0, -3.5e3), 1, 3) # well location: 500 m to east of center pt. of main fault
```

p = plotFault3d(flt2, theta=45, phi=40, xlim=c(-4,4), ylim=c(-6,6), zlim=c(-7,-1), lwd=0.5)lines(trans3d(c(xi[1,1], xi[1,1])/1e3, c(xi[1,2], xi[1,2])/1e3, c(0, xi[1,3])/1e3, p), lwd=3)



### Set up injection history and resulting stressing file

Make a simple injection history and RSQSim external stressing file from it. Can just copy this over from Run2. Or, actually, to save disk space (current test.externalStress.txt is 450 MB), should just point to Run2's copy.

```
yr = 365.25*86400
injHist = makeInjectionHistory(xi=xi) # use defaults for all hydrologic params

#t = findSampleTimes(flt2, injHist, tmax=100*yr, eps=0.05, minh=1000)
#writePoreFluidStressingHistoryFile(flt2, injHist, t, "test.externalStress.txt")
#file.copy(file.path("..", "Run2", "test.externalStress.txt"), ".")

s = readExternalStressingHistoryFile(file.path("..", "Run2", "test.externalStress.txt"), flt2)
```

#### Plots of pressure histories

The closest element, the farthest element, and the median distance element.

```
lines(s$t/yr, -s$sigma[ord[flt2$np],], col="red2")
lines(s$t/yr, -s$sigma[ord[round(flt2$np/2)],], col="orange")
Pore-fluid pressure perturbation (MPa)
       30
       20
       10
       0
                0
                              20
                                                                                        100
                                            40
                                                           60
                                                                          80
                                                Time (yrs)
```

### Initial stresses

```
set.seed(234)
tau00.main = 63.2
tau00.random = 65
sdtau.main = 0
sdtau.random = 1
main = which(flt2$segName == "NA")
random = which(flt2$segName != "NA")
tau0 = rep(NA, flt2$np)
tau0[main] = rnorm(length(main), tau00.main, sdtau.main)
tau0[random] = rnorm(length(random), tau00.random, sdtau.random)
write(tau0, "initTau.txt", ncol=1)
```

# Run simulation

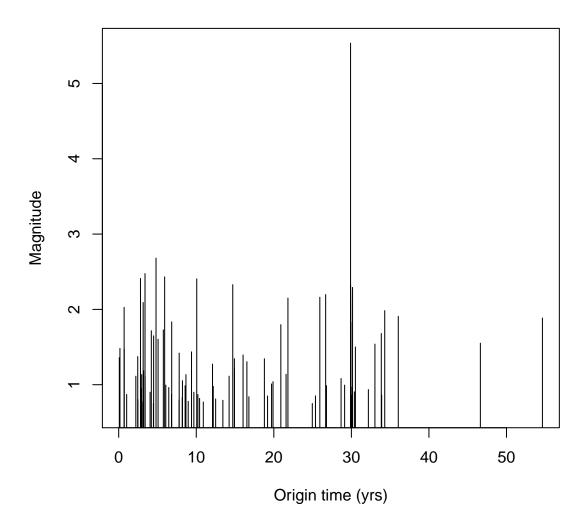
Copy Run3's parameter file and make adjustments.

## Results

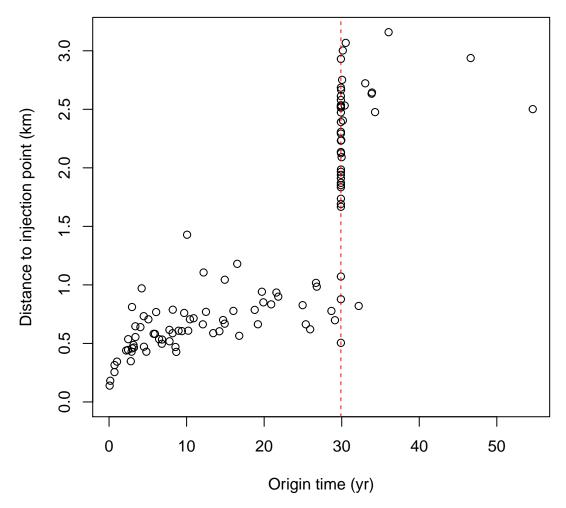
```
eqs = readEqs(paste("eqs.", params$outFnameInfix, ".out", sep=""))
if (max(eqs$t0) > max(s$t)) {
   eqMax = min(which(eqs$t0 > max(s$t))) - 1
   eqs = readEqs(paste("eqs.", params$outFnameInfix, ".out", sep=""), eqMax = eqMax)
}
```

#### Magnitude-time plot

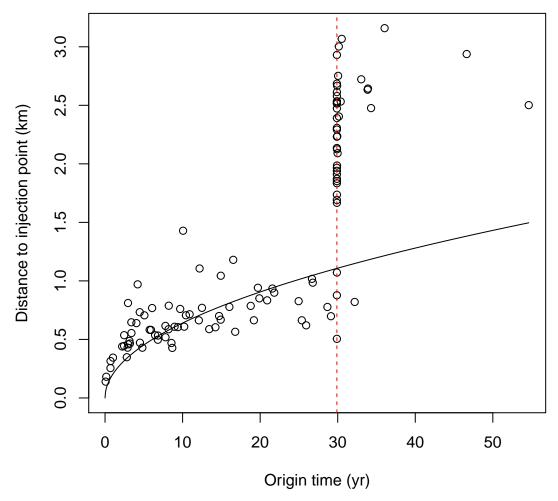
```
plot(eqs$t0yr, eqs$M, typ='h', xlab="Origin time (yrs)", ylab="Magnitude")
```



# Hypocenter-well distance



The evolution of the distance of the induced events from the injection point with time looks roughly like the expected  $\sqrt{t}$  (until the M5+). It is not clear what the exact relationship between the coefficient of  $\sqrt{t}$  and the hydraulic parameters should be. Shapiro et al. [1997] advocate for  $x = \sqrt{4\pi\kappa t}$ , where  $\kappa$  is the diffusivity. The argument of the erfc in our pore-fluid pressure expression is equal to one at  $x = \sqrt{4\kappa t}$ , but there is also a  $|x|^{-1}$  factor in front of the erfc. Adding in the least-squares fit to the pre-M5 events:



The  $\sqrt{t}$  is not a great fit: event times are systematically earlier at near distances and later at farther distances than the best fit  $\sqrt{t}$  curve. I would attribute this to the  $|x|^{-1}$  factor in the pore-fluid pressure expression. It could also be due to RSF evolution effects. If we ignore that and ask what value of  $\kappa$  is implied by the Shapiro et al. [1997] expression or just setting the argument of the erfc to 1, we get  $10^{-4}$  and  $3.2 \times 10^{-4}$  m<sup>2</sup>/s, respectively. These are both about two orders of magnitude lower than the actual diffusivity used in generating the pore-fluid pressures  $(0.008 \text{ m}^2/\text{s})$ . Note that this difference here between the actual diffusivity and that estimated by assuming  $x = \sqrt{4\pi\kappa t}$  is not necessarily relevant to the data in Shapiro et al. [1997]. In that paper the diffusivity is much higher and the time scales are much shorter (a few days), so that RSF evolution effects may be less important.

Elizabeth mentions that some papers of Talwani estimates the diffusivity by fitting curves of the form  $x = \sqrt{\kappa t}$ . Also said that might have been for fluid pressures diffusing in 2-d (e.g. if a fault is a high permeability conduit) instead of in the 3-d bulk. If we ignore the 2-d/3-d issue and just estimate  $\kappa$  from curves of that form, we get 0.0013 m<sup>2</sup>/s, still more than an order of magnitude less than the actual value.

Why are these estimates so much lower here than in the uniform pre-stress cases?

#### Pore-fluid pressure at failure

As an indication of how important RSF evolution effects may be, let's look at the pore-fluid pressure one each hypocentral element just before it failed. For friction with a fixed failure stress (e.g. slip-weakening) and uniform initial stresses as we have here, this would be expected to be the same for all events as long as the stress transfer between the elements is neglible (which it will be in this fault system everywhere except

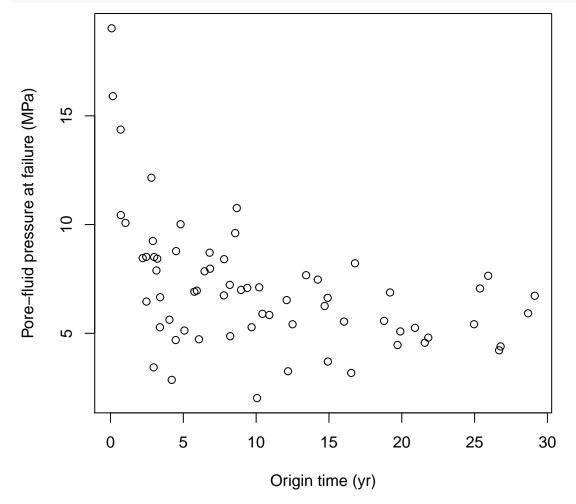
the main fault). However, with RSF, the more distant, later-failing elements will fail at a lower pore-fluid pressure because (once they are above steady-state) they will have been weakening.

Note that the instantaneous pore-fluid pressure increment that would cause immediate failure is given by that P which solves

$$V^* \exp\left(\frac{\tau - \mu_0(\sigma_0 - P)}{a(\sigma_0 - P)}\right) \left(\theta_0 \left(\frac{\sigma_0 - P}{\sigma_0}\right)^{\alpha/b} \frac{V^*}{D_c}\right)^{-b/a} = V^{\text{EQ}}.$$

For this simulation, this comes out to about 26.9 MPa.

Below I plot the pore-fluid pressure perturbation at failure for all events before the M5 (after which stress transfer dominates).



These are indeed all less than 26.9 MPa and are smaller at later times / farther distances, but very scattered

here because of the variable pre-stress.

So, for now, it is an open question how much of the departure from  $\sqrt{t}$  behavior is due to the  $|x|^{-1}$  and how much due to RSF evolution effects.

### Estimation of Omori p-value and aftershock fraction n

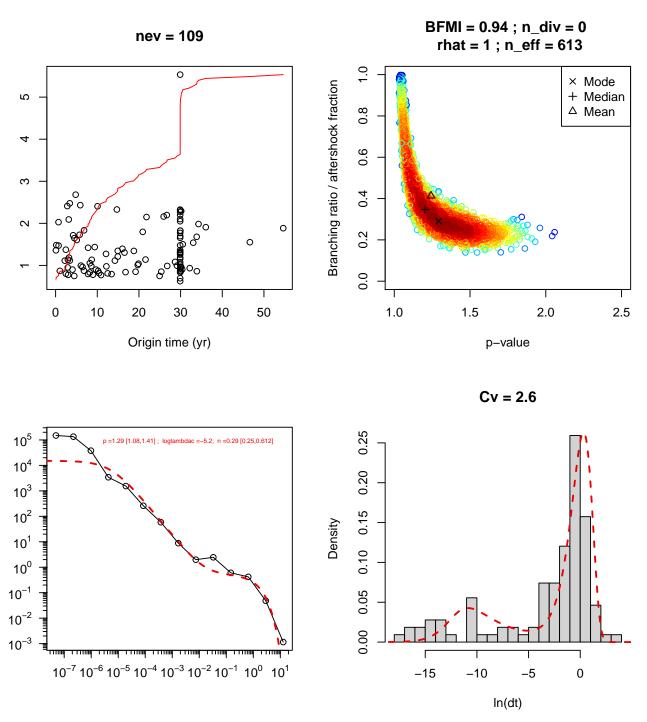
See templateFamilies.[Rmd|pdf]

```
use = 1:length(eqs$t0)
alpha = 0.32 # so 68% CIs
c = 90 # seconds, needs to match dSS2007.krd.fixedloglambdac.stan
st = do.stan(eqs, use=use, alpha=alpha, c=c)
```

#### Plotting function definition

The functions to make plots like in templateFamilies.pdf; modified from there. Not printing them out because they take a lot of page space.

### Actually make the plots

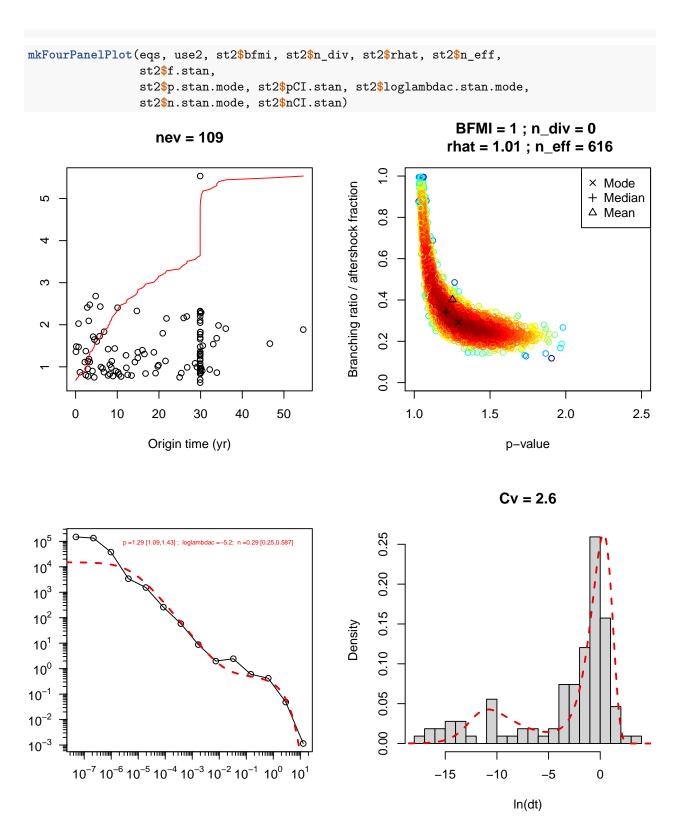


This estimates the aftershock fraction to be 0.29. Just given the data asked for my opinion, I probably would have counted the 39 events that occurred in the couple of years after the M5+ as aftershocks and said the aftershock fraction was 39/109 = 0.36.

#### Leaving out very late events

Perhaps those last 0 events way off by themselves are biasing things? What happens if we leave them out?

```
use2 = 1:(length(eqs$t0)-nout)
st2 = do.stan(eqs, use=use2, alpha=alpha, c=c)
```



Aftershock fraction didn't change much: 0.29.

Note that the estimate of  $Hainzl\ et\ al.\ [2006]$  would give 0.85 and 0.85, when using all events and all the but the last 0 events, respectively.

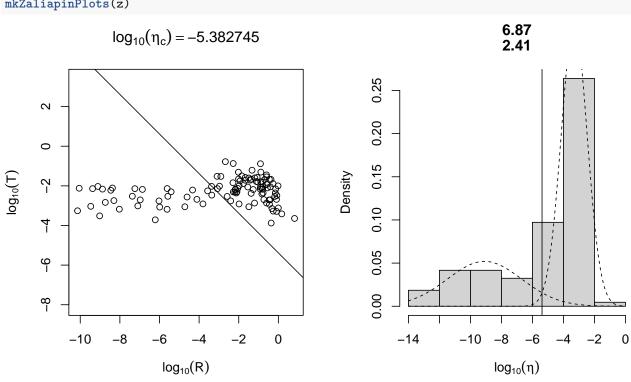
Both the estimated aftershock fraction from fitting the Saichev and Sornette [2007] functional form and the  $C_v$  are very affected by these outliers. If we used the 1-norm version of the  $C_v$ , call it  $C_{v1}$ ,  $C_{v1} = \text{smad}(dt)/\text{median}(dt)$ , it would be essentially completely unaffected: including the 0 events gives 1.4824502 and excluding them gives 1.4824502.

# Aftershock fraction ala Zaliapin and Ben-Zion [2016]

```
sderrh = 10
z = zaliapin(eqs, use=use, sderrh=sderrh, est.se=TRUE)
z2 = zaliapin(eqs, use=use2, sderrh=sderrh)
```

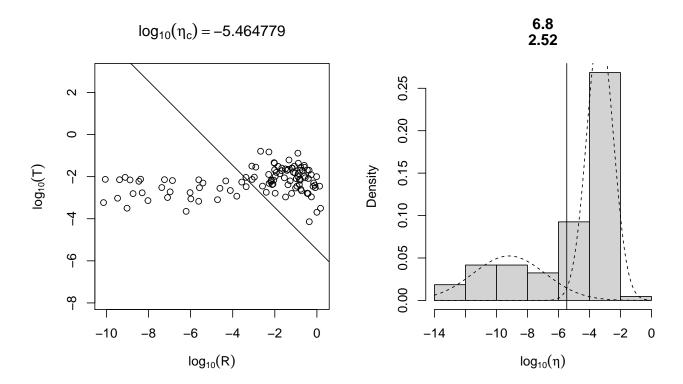
#### Using all events

```
source("../mkZaliapinPlots.R")
mkZaliapinPlots(z)
```



#### Excluding late events

mkZaliapinPlots(z2)



## Summary

- $\bullet\,$  Full-main-fault event at 29.9 yrs
- Number of events
  - total: 109
  - before full-main-fault event: 66
  - after full-main-fault event: 42
- Estimated aftershock fractions
  - naive count: 0.36
  - SS2007
    - \* using all events: 0.29
    - \* excluding last 0 events: 0.29
  - H2006
    - \* using all events: 0.85
    - \* excluding last 0 events: 0.85
  - ZB2016
    - \* using all events: 0.3
    - \* excluding last 0 events: 0.29

# To do

- how to get more realistic time-scales and pore-pressures?
- different uniform stresses on main fault

- variable stresses on main fault
- different forms / amplitudes of variable stresses on random faults
- denser random faults

# References

Hainzl, S., F. Scherbaum, and C. Beauval (2006), Estimating background activity based on interevent-time distribution, *Bulletin of the Seismological Society of America*, 96(1), 313–320.

Saichev, A., and D. Sornette (2007), Theory of earthquake recurrence times, *Journal of Geophysical Research:* Solid Earth, 112(B4).

Shapiro, S. A., E. Huenges, and G. Borm (1997), Estimating the crust permeability from fluid-injection-induced seismic emission at the ktb site, *Geophysical Journal International*, 131(2), F15–F18.

Zaliapin, I., and Y. Ben-Zion (2016), Discriminating characteristics of tectonic and human-induced seismicity, Bulletin of the Seismological Society of America, 106(3), 846–859.