

ASLO

LIMNOLOGY AND OCEANOGRAPHY

# e-Lectures

## Non-invasive Flux Measurements at the Benthic Interface: The Aquatic Eddy Covariance Technique

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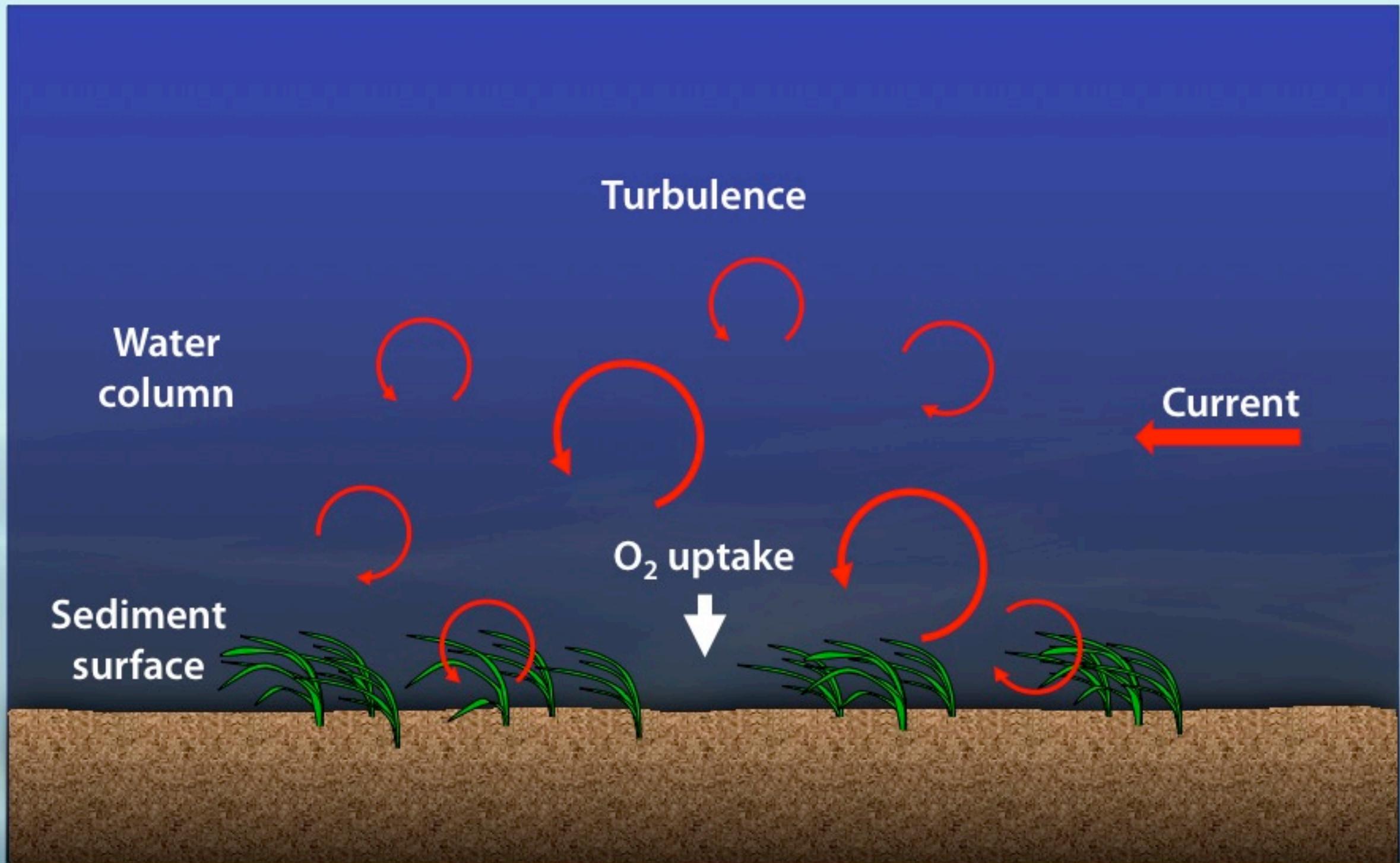
## A versatile approach



Photo: AWI Sea Ice Physics group

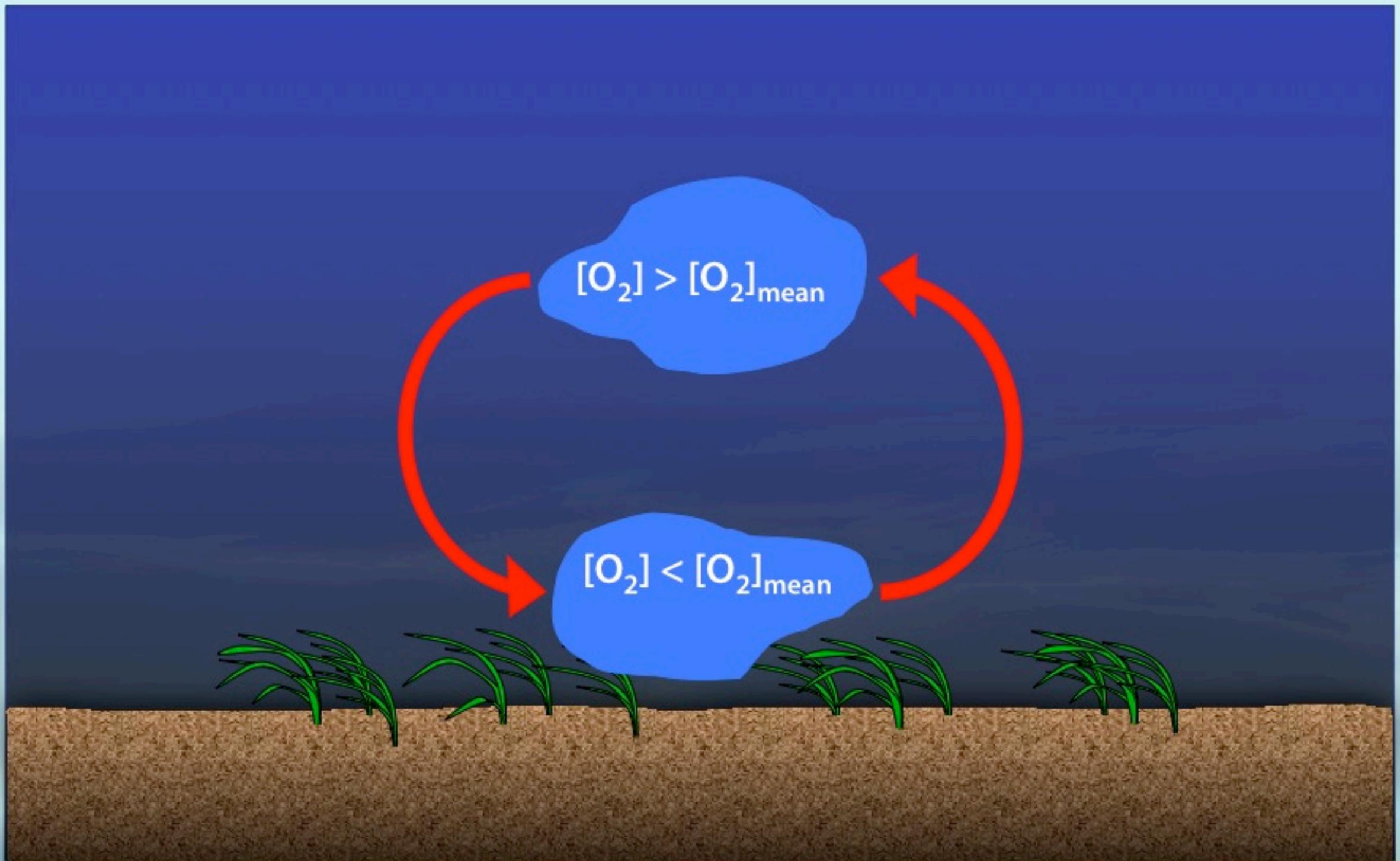


# Basic principles of eddy covariance





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# Basic principles of eddy covariance

- Key variables: vertical velocity ( $w$ ) and oxygen concentration ( $C$ )
- Instantaneous vertical advective flux =  $w C$



Unit:  $[\text{cm s}^{-1}] [\text{nmol cm}^{-3}] = [\text{nmol s}^{-1} \text{ cm}^{-2}]$

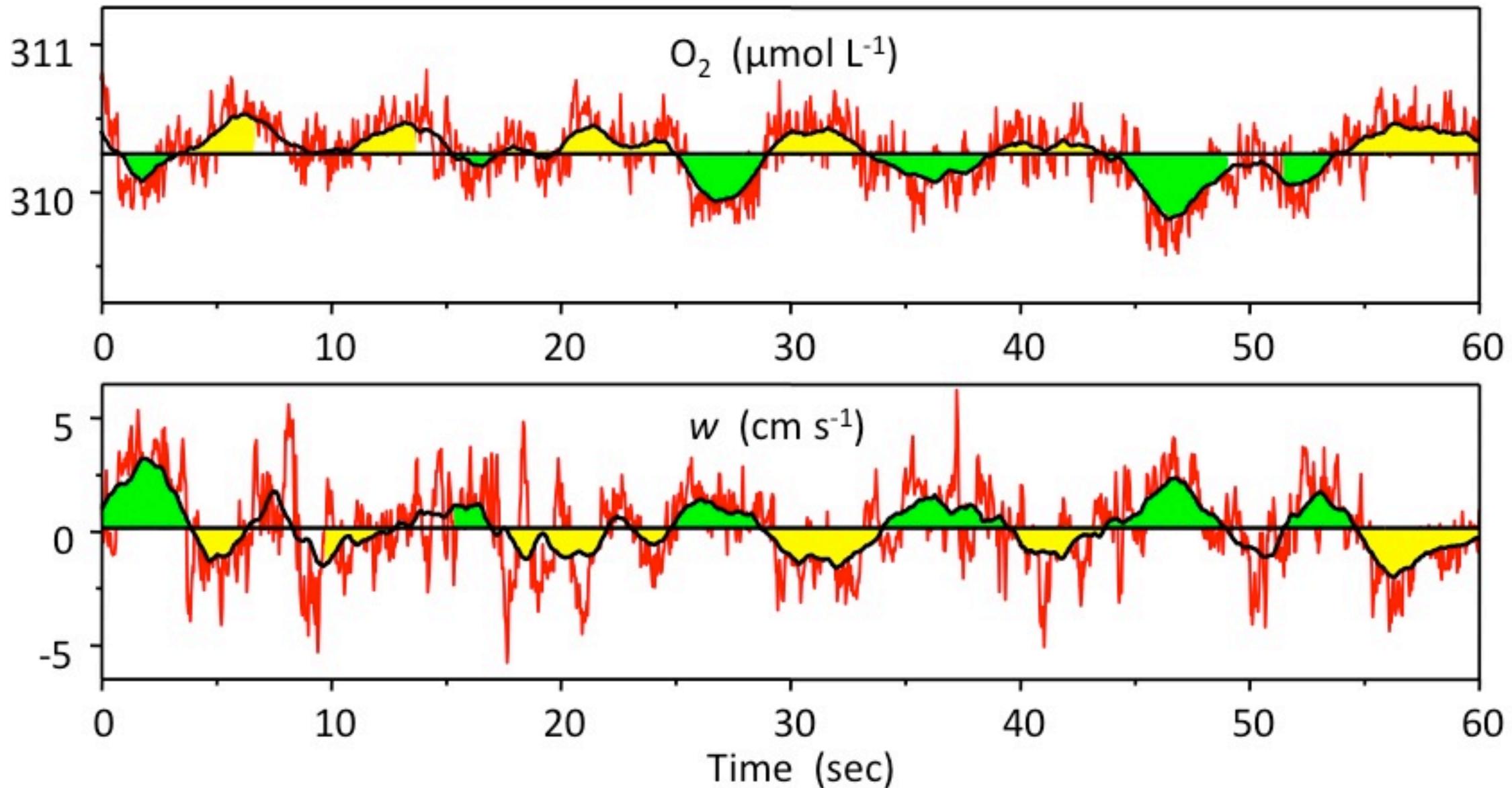


← Measurements at a point





# Basic principles of eddy covariance

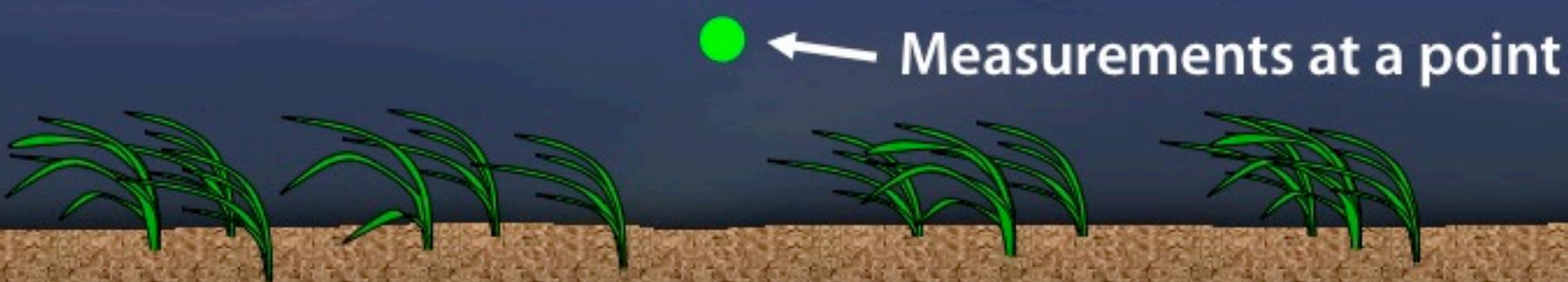


- One min of real data measured over an oxygen-consuming substrate



## Basic principles of eddy covariance

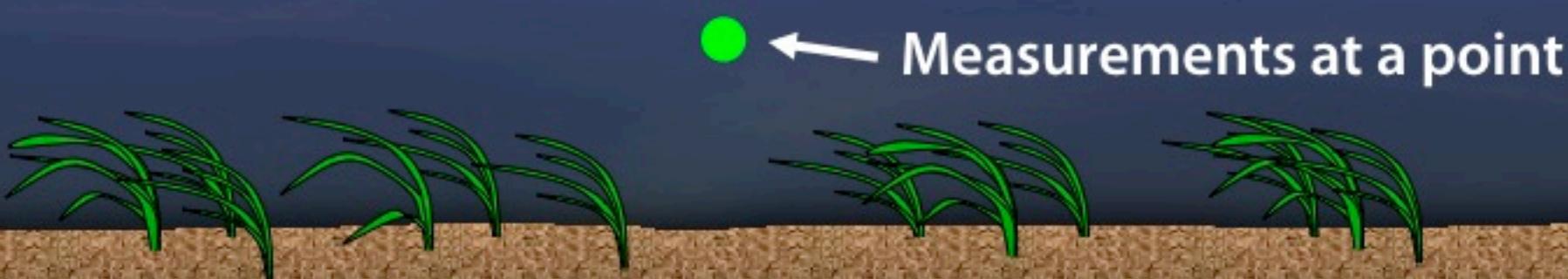
- Reynolds decomposition:  $w = \bar{w} + w'$  and  $C = \bar{C} + C'$
- Averaged flux =  $\bar{wC}$





## Basic principles of eddy covariance

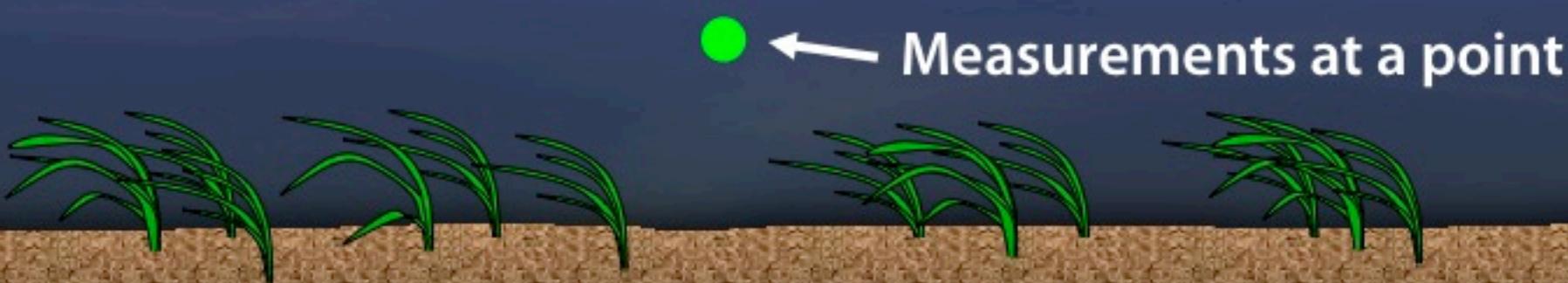
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## Basic principles of eddy covariance

- Reynolds decomposition:  $w = \bar{w} + w'$  and  $C = \bar{C} + C'$
- Averaged flux =  $\overline{wC} = \bar{w}\bar{C} + \cancel{\bar{w}\bar{C}'} + \cancel{\bar{w'C}} + \bar{w'}\bar{C}'$

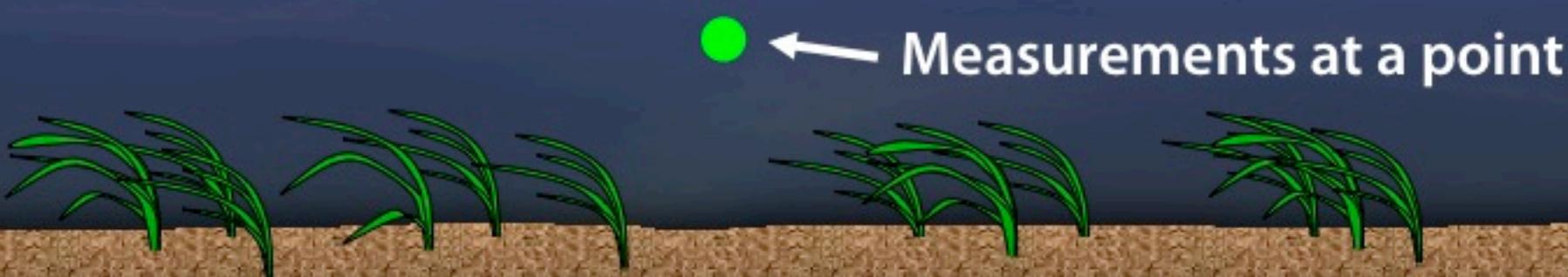




# Basic principles of eddy covariance

- Reynolds decomposition:  $w = \bar{w} + w'$  and  $C = \bar{C} + C'$
- Averaged flux =  $\bar{wC} = \cancel{\bar{w}\bar{C}} + \cancel{\bar{w}\bar{C}'} + \cancel{\bar{w'}\bar{C}} + \bar{w' C'}$
- Assumption:  $\bar{w} = 0 \rightarrow$

$$\text{Averaged flux} = \bar{w' C'}$$

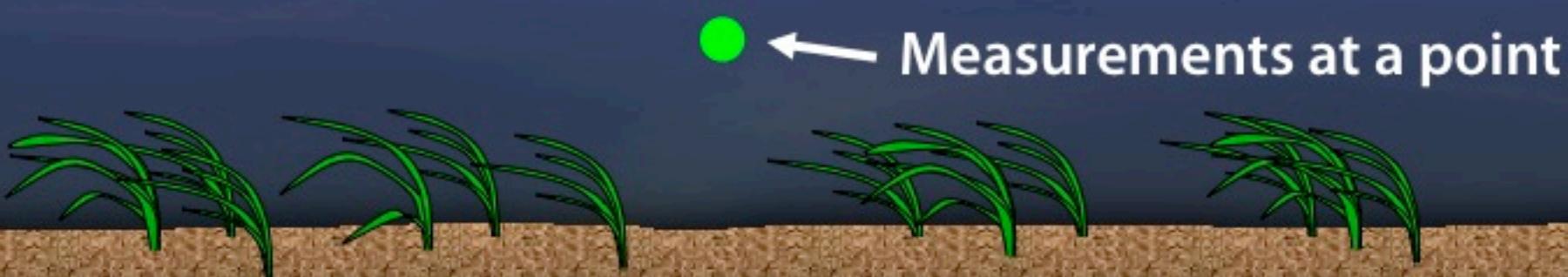




# Basic principles of eddy covariance

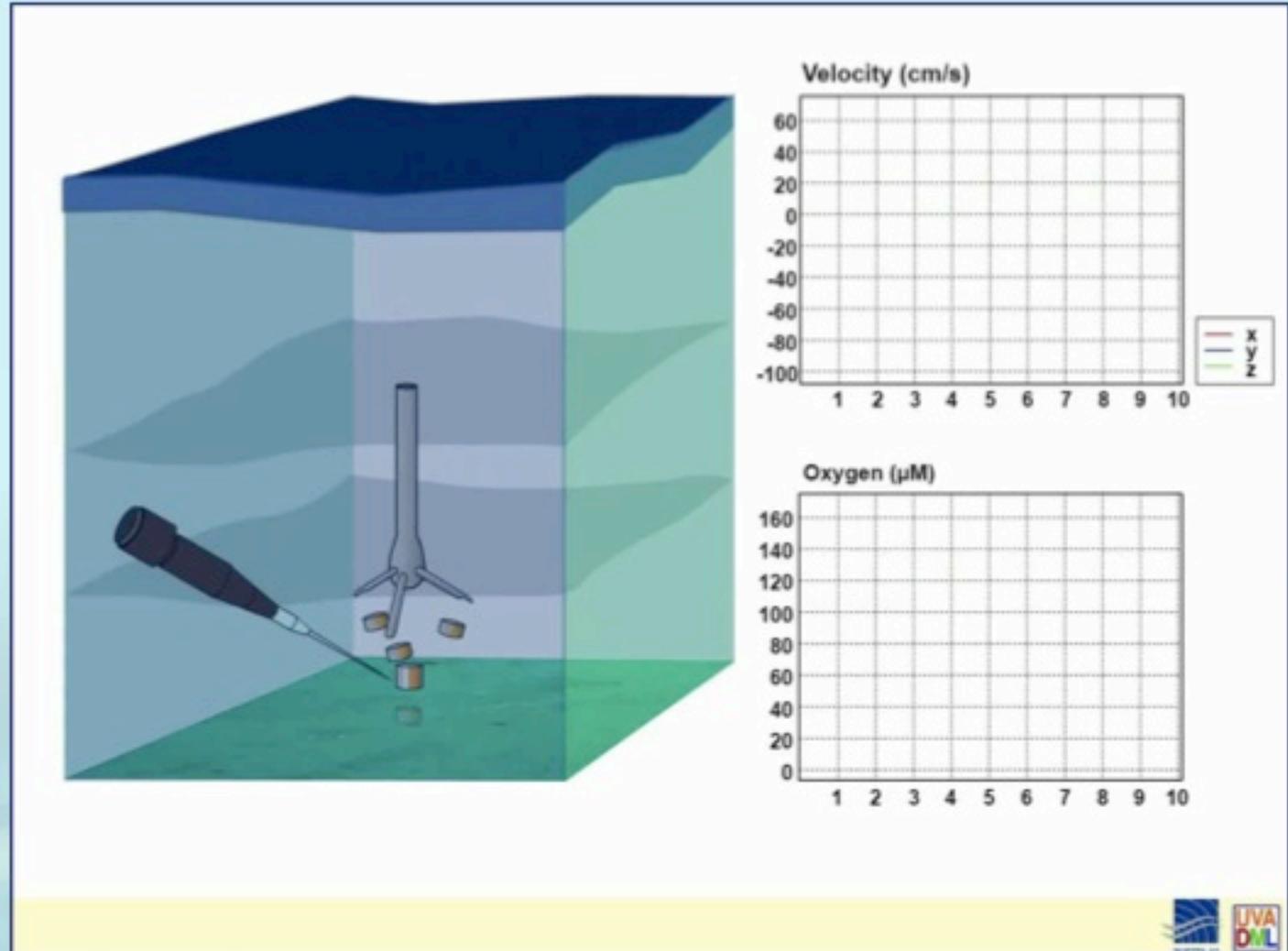
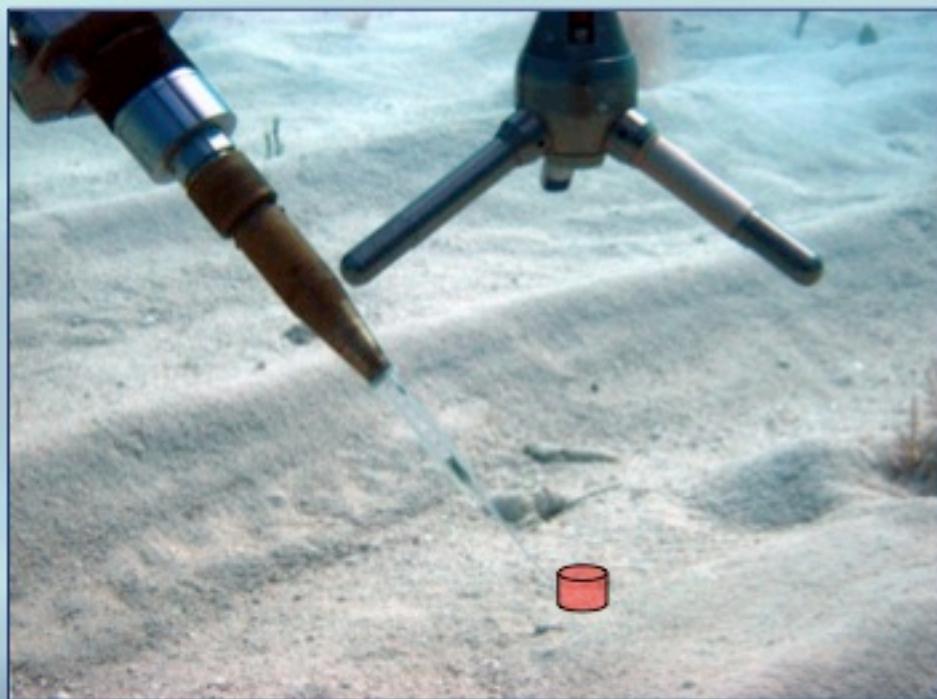
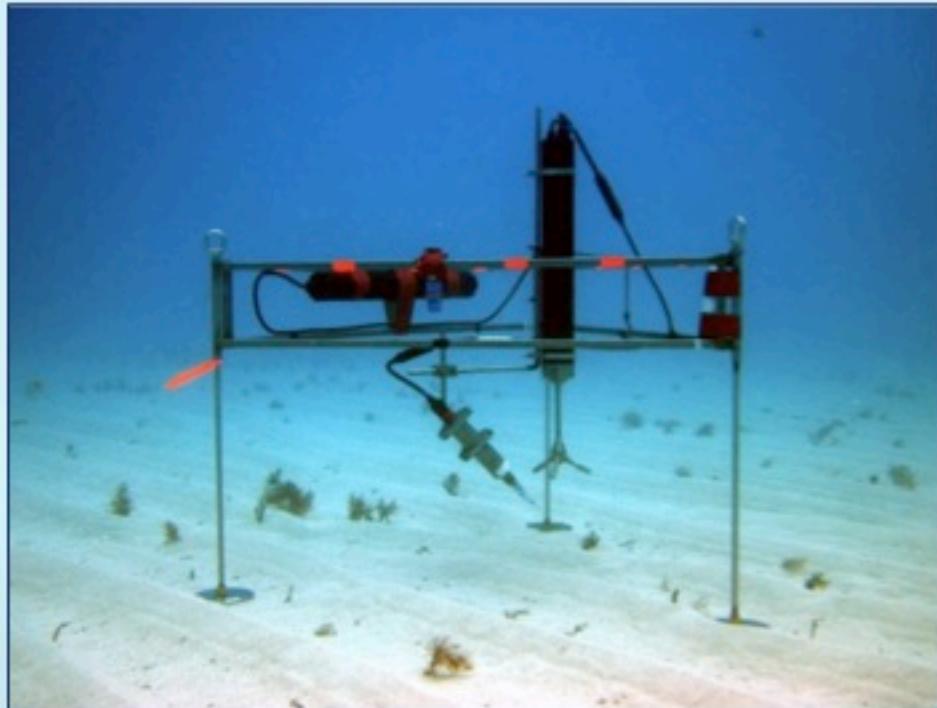
- Reynolds decomposition:  $w = \bar{w} + w'$  and  $C = \bar{C} + C'$
- Averaged flux =  $\bar{w}\bar{C}$  =  ~~$\bar{w}\bar{C} + \bar{w}\bar{C}' + \bar{w}'\bar{C} + \bar{w}'\bar{C}'$~~
- Assumption:  $\bar{w} = 0 \longrightarrow$

$$\text{Averaged flux} = \frac{1}{N} \sum_{i=1}^N w'_i C'_i$$





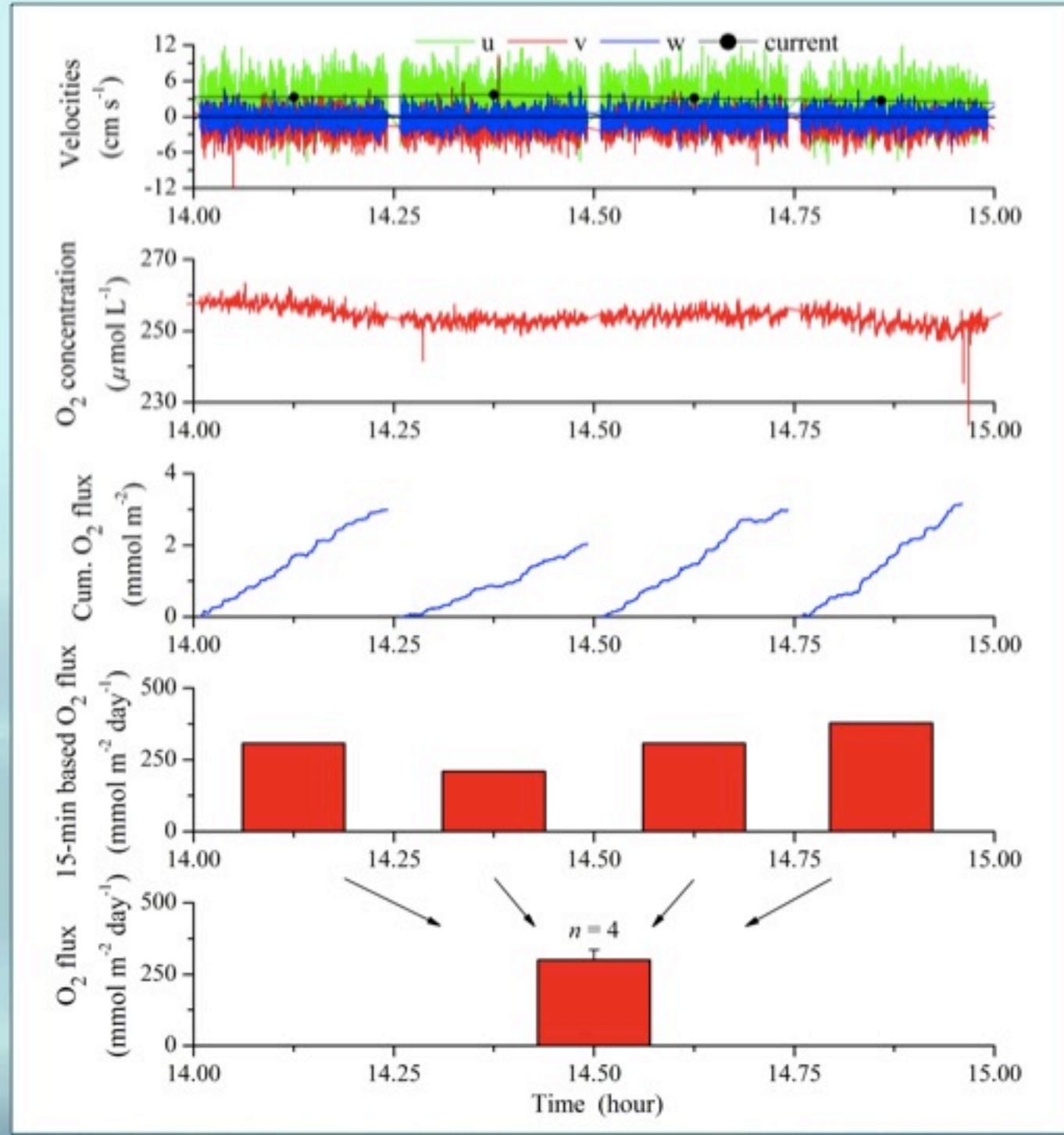
# Basic principles of eddy covariance



- Data typically recorded at 32 - 64 Hz and 5 - 30 cm above benthic surface



# Basic principles of eddy covariance

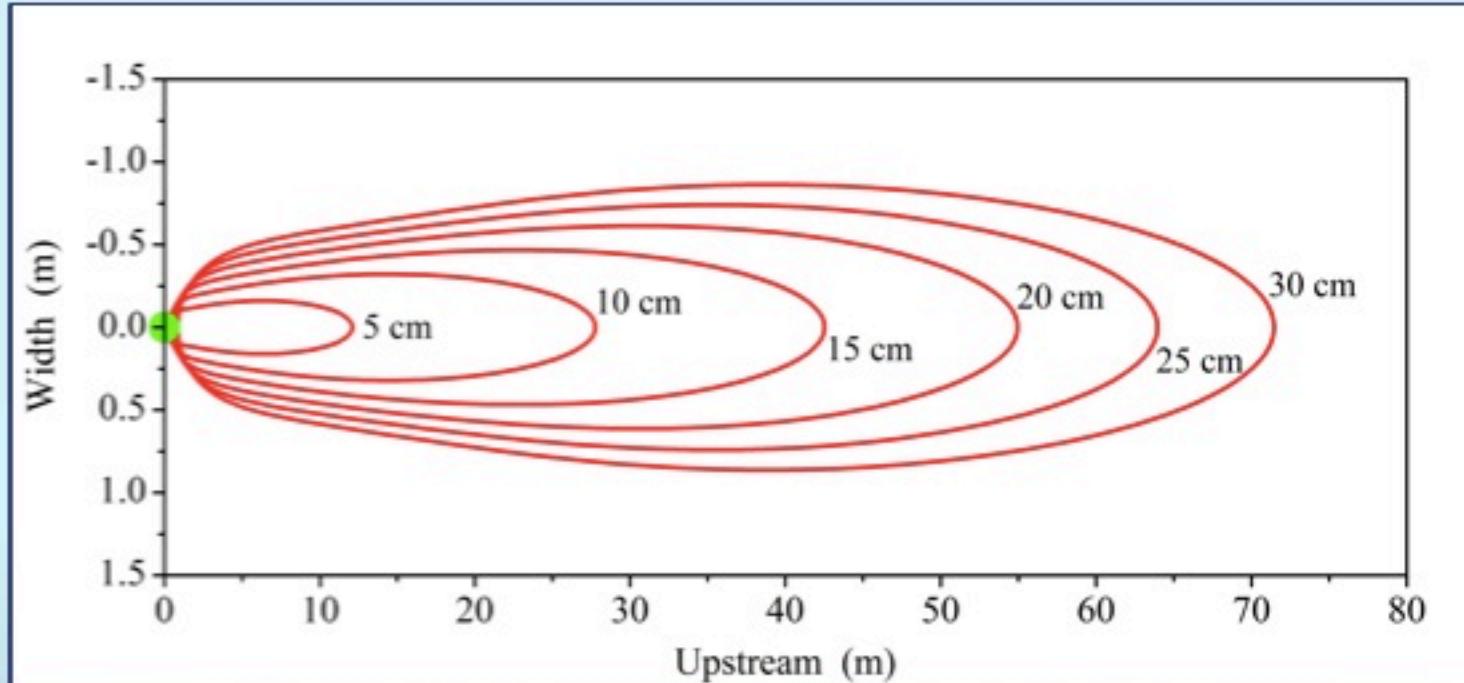
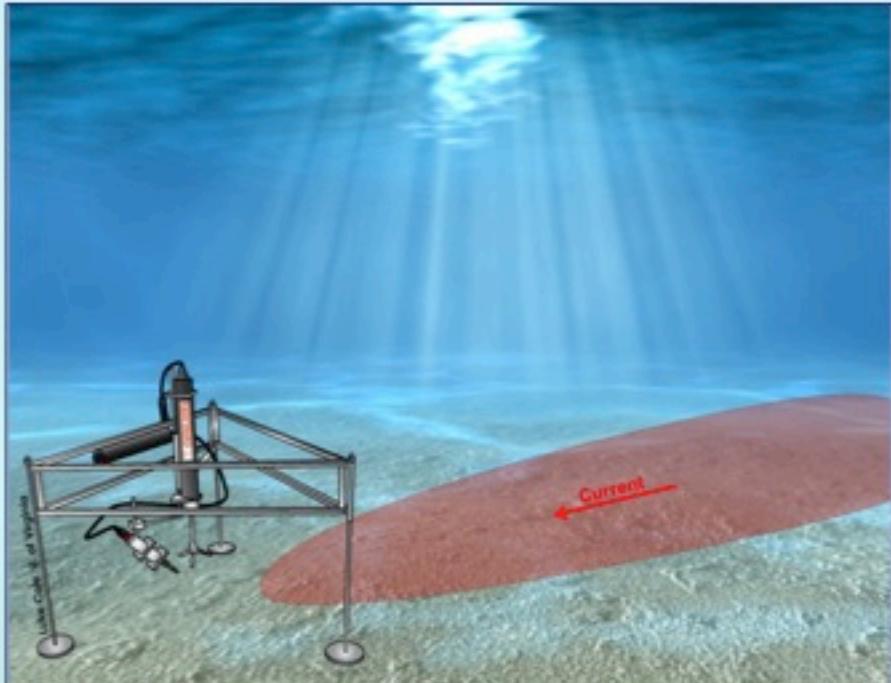


Hume et al. (2011) L&O





# Where does the eddy flux come from?

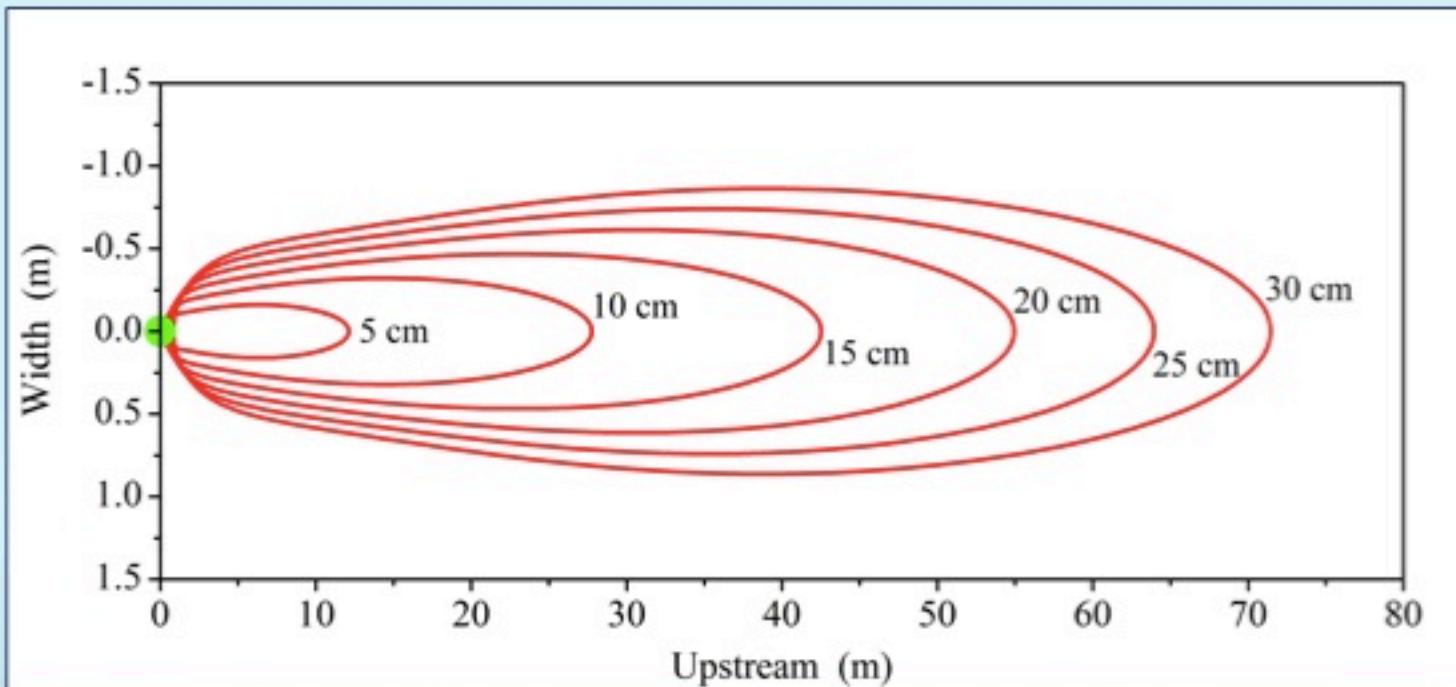
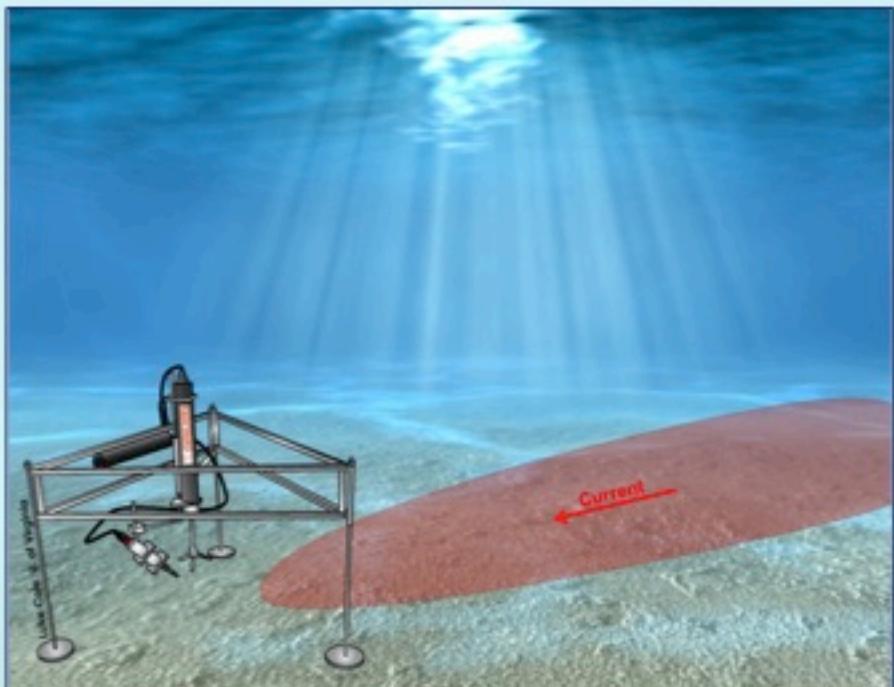


Berg et al. (2007) L&O

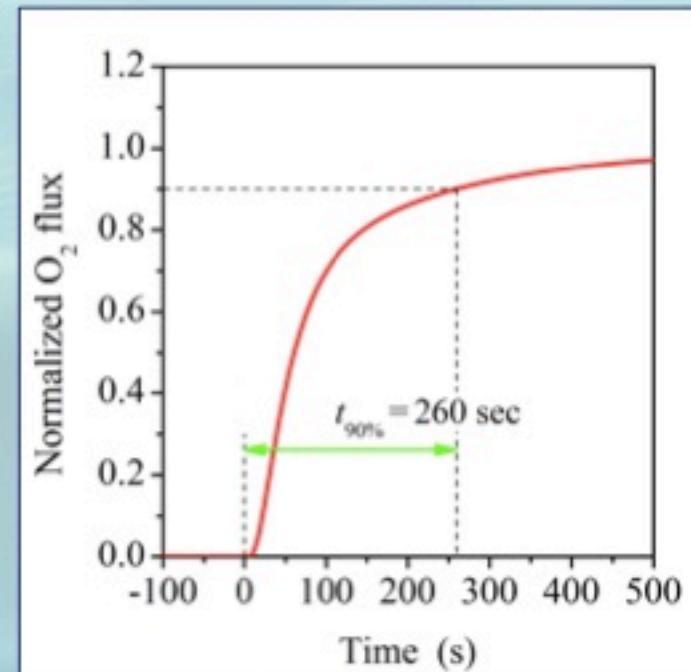
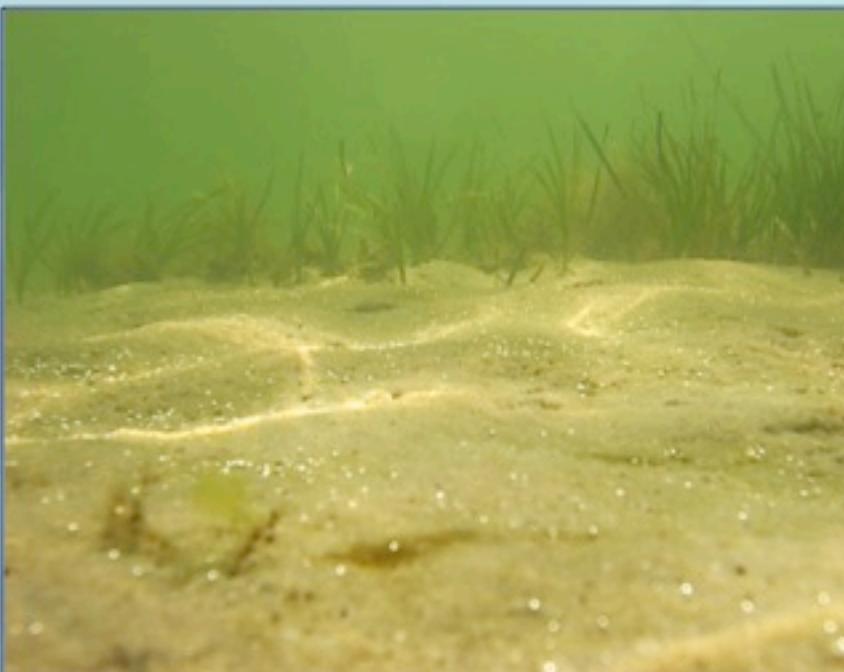
- Footprint definition: smallest area on the bottom that contributes 90% of the measured flux
- Typical size: 5 to 100 m<sup>2</sup>
- Footprint different (and as yet undetermined) at shallow-water sites with waves because of oscillatory wave motions



# Where does the eddy flux come from?



Berg et al. (2007) L&O



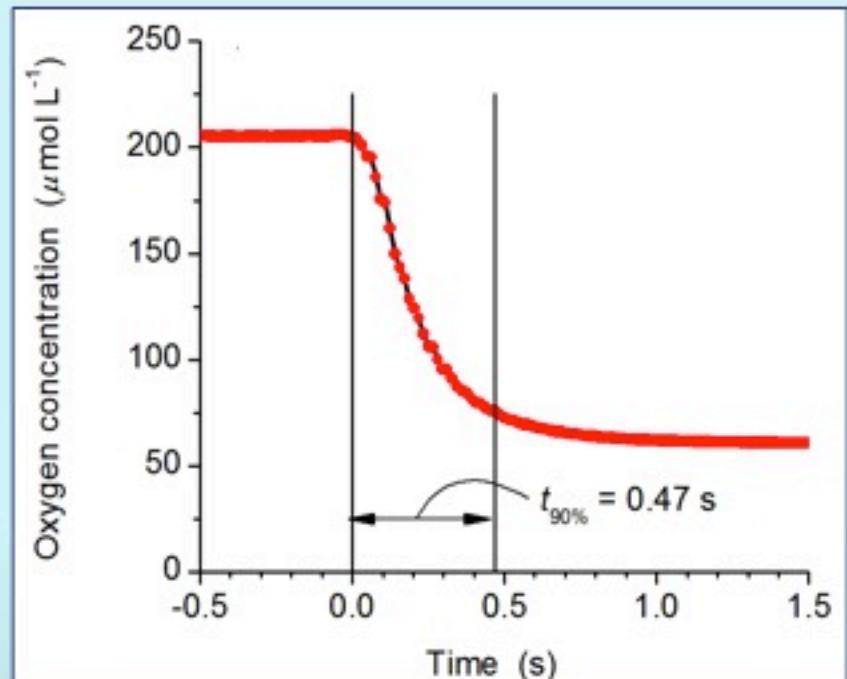
Rheuban and Berg (2013) L&O Methods



# Requirements of approach

- Record velocity and oxygen data at  $> \sim 8 \text{ Hz}$
- Record velocity and oxygen data no more than 2 – 3 cm apart
- Sensor response time  $t_{90\%} < \sim 0.5 \text{ s}$
- More complex to apply than other methods
- Full system price: \$30,000 - \$60,000

Why do all this?

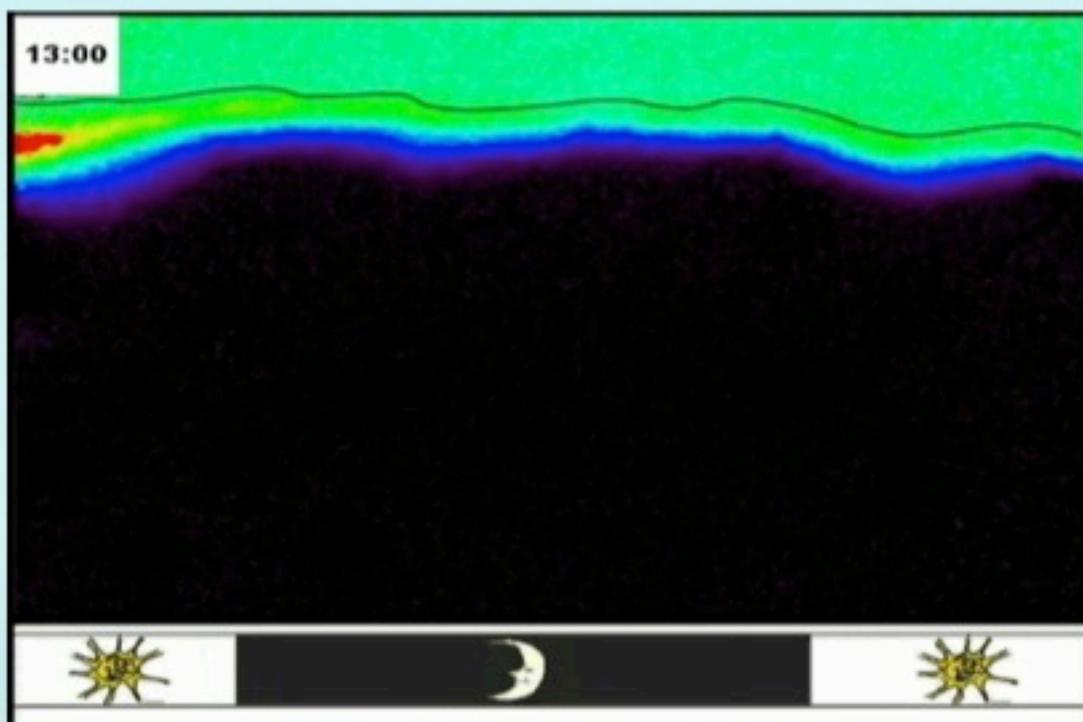


Berg et al. (2016). L&O Methods

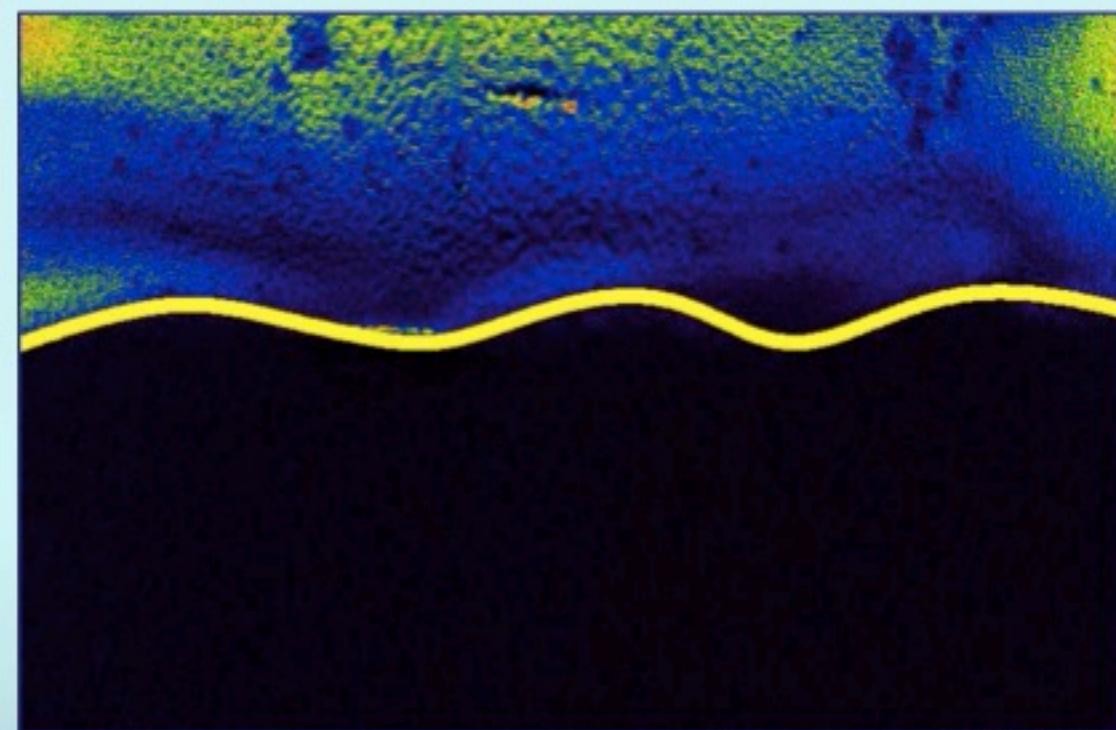




# Natural environment complex and dynamic



Wenzhofer and Glud (2004)



Kessler et al. (Unpubl.)



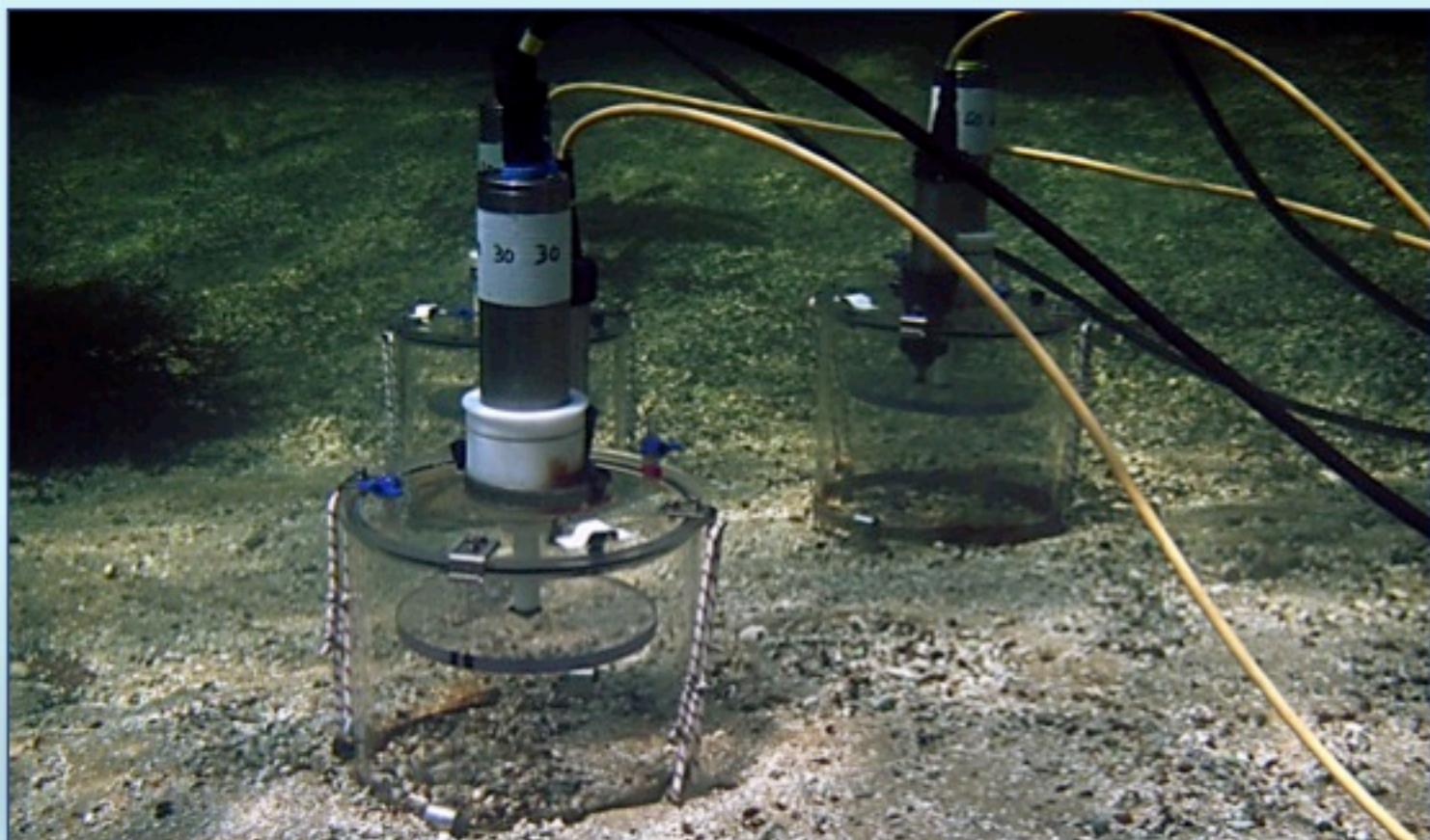
Reimers et al. (2004)



Reidenbach et al. (2007)



## Traditional flux methods have limitations

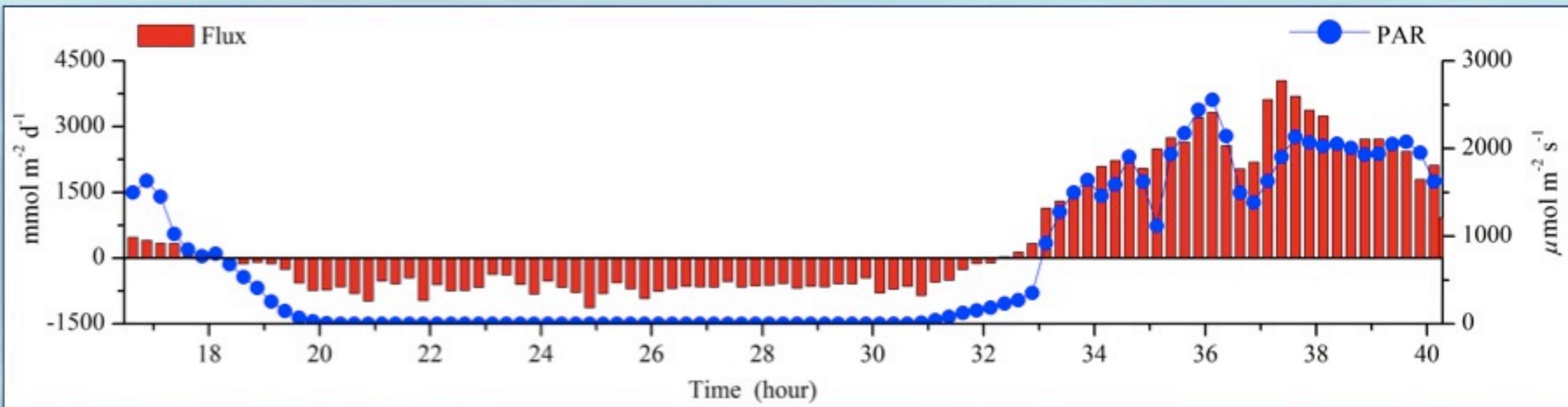


Sediment core incubations and in situ chamber deployments:

- Exclude or seriously alter the most important drivers of flux (current flow, wave action, light, etc.)
- Disturb the sediment
- Cut off all exchange with the natural dynamic water column



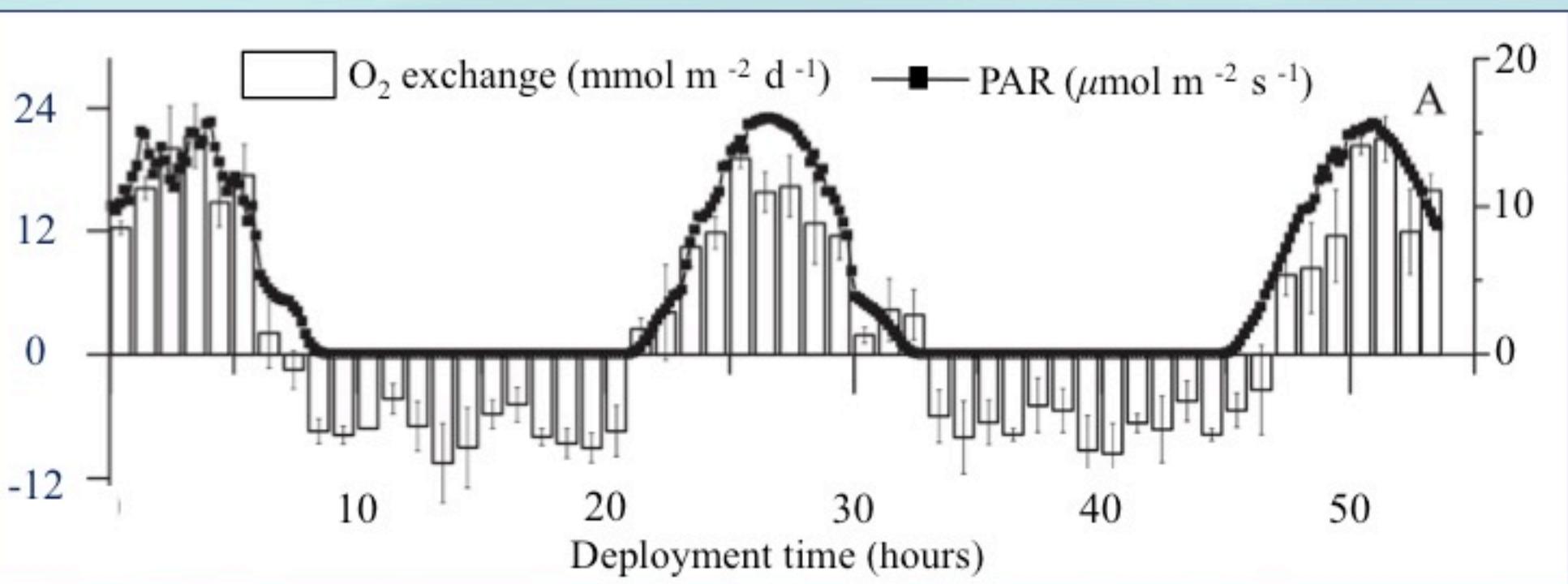
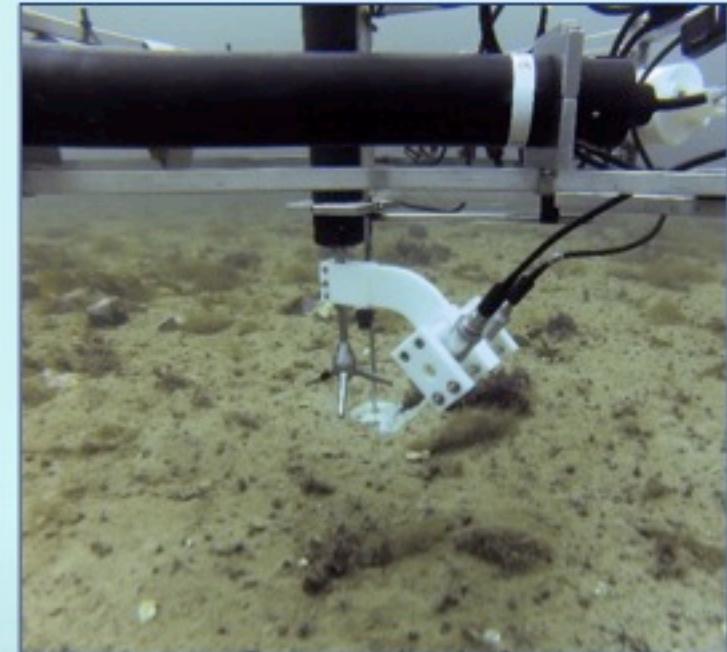
# Effects of natural light



Long et al. (2013) PLoS ONE



# Effects of natural light

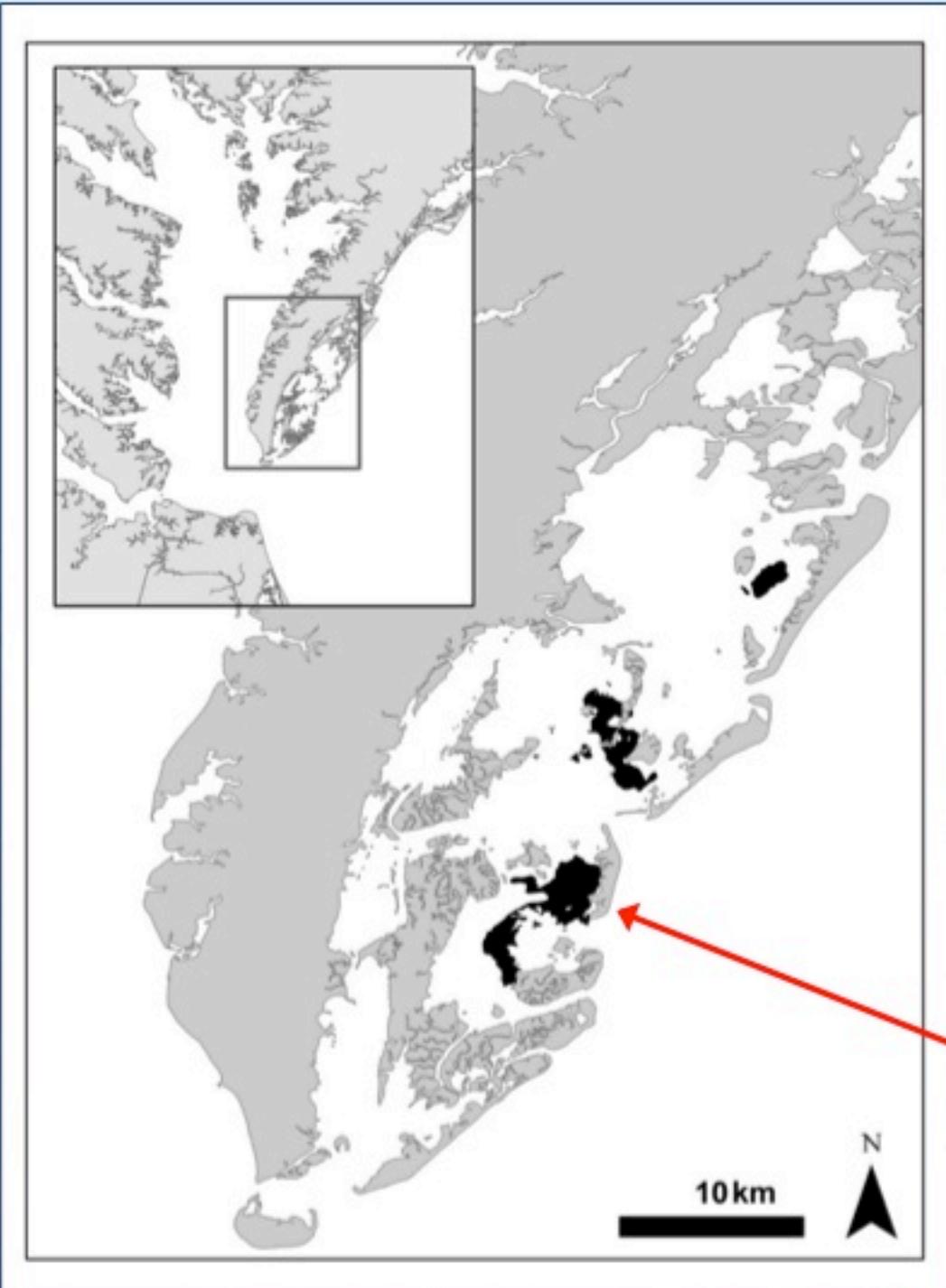


Attard et al. (2014) L&O

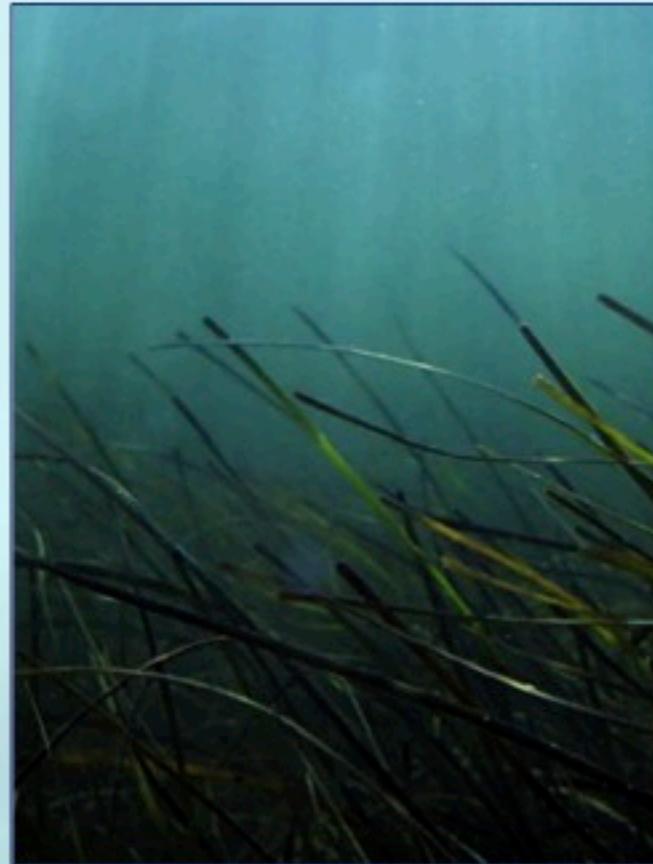




# Seagrass at Virginia Coast Reserve LTER



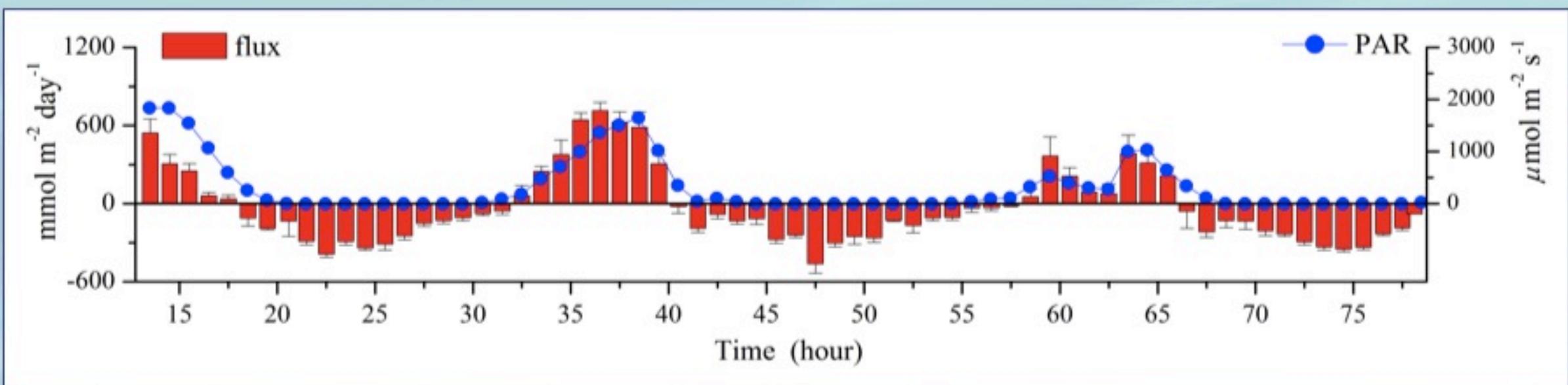
Map: M. Oreska & J.J. Orth



- Restored seagrass meadows – larger than 2500 ha
- Site of extensive aquatic eddy covariance work



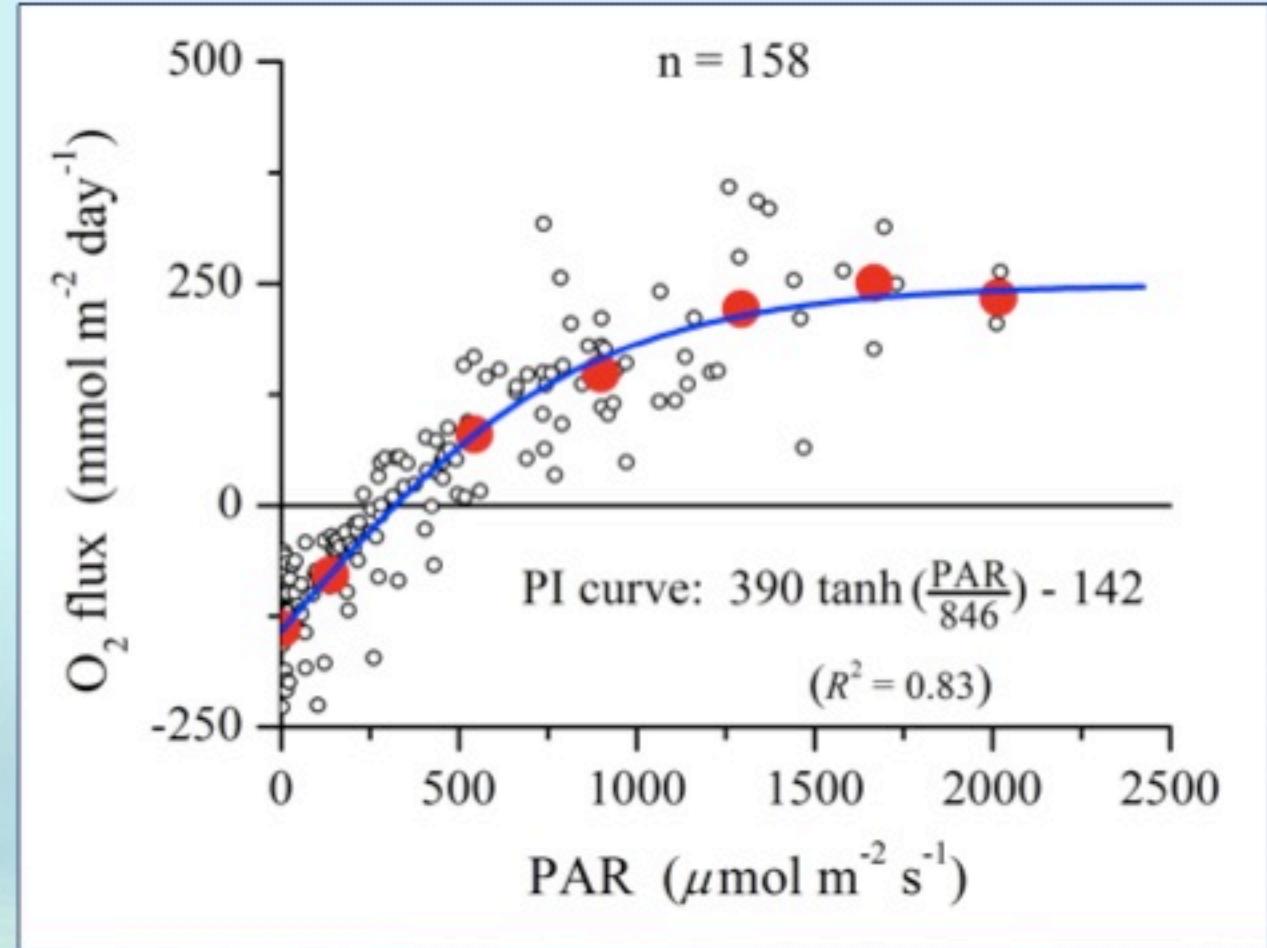
# Effects of natural light



Delgard et al. (In prep A)



# Effects of natural light

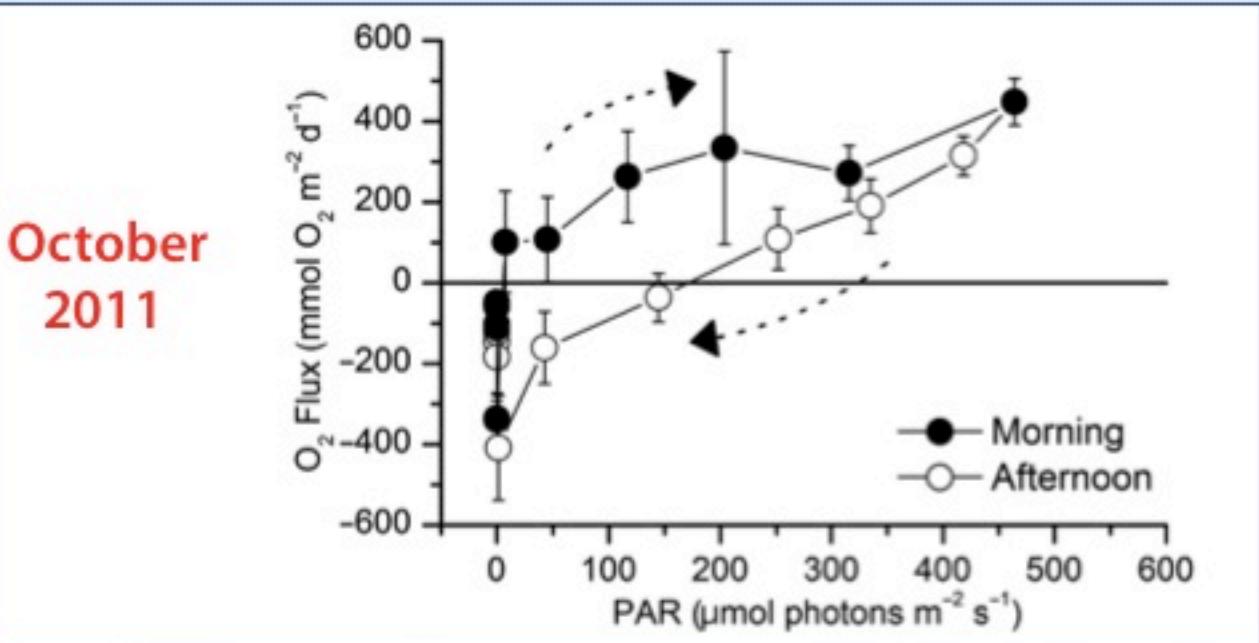


Delgard et al. (In prep A)

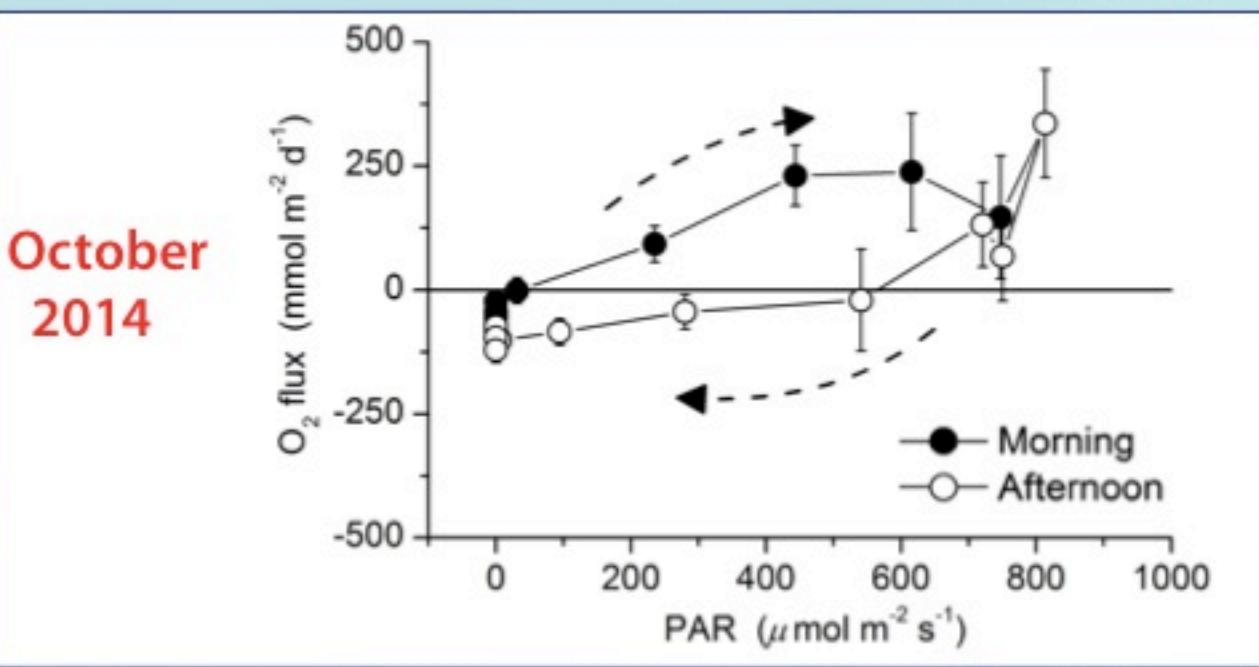
- Averaged system-scale response to light
- Integrated horizontally and vertically (through canopy)



# Effects of natural light



Rheuban et al. (2014) MEPS



Delgard et al. (In prep A)

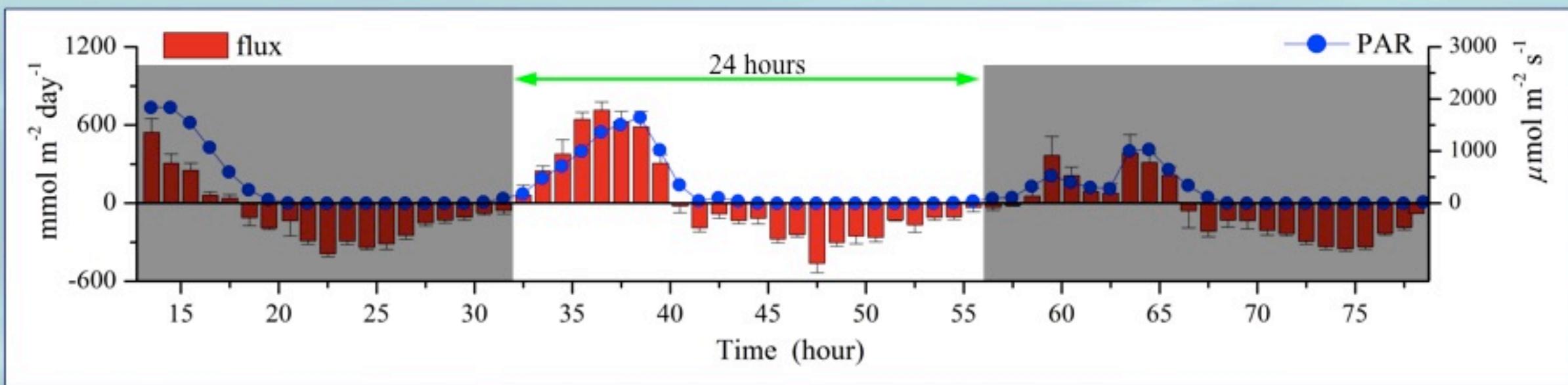


# Seagrass metabolism



- $R = \frac{1}{24} \left( \sum \text{flux}_{\text{dark}} + \frac{\sum \text{flux}_{\text{dark}}}{h_{\text{dark}}} h_{\text{light}} \right)$
- $GPP = \frac{1}{24} \left( \sum \text{flux}_{\text{light}} + \frac{|\sum \text{flux}_{\text{dark}}|}{h_{\text{dark}}} h_{\text{light}} \right)$

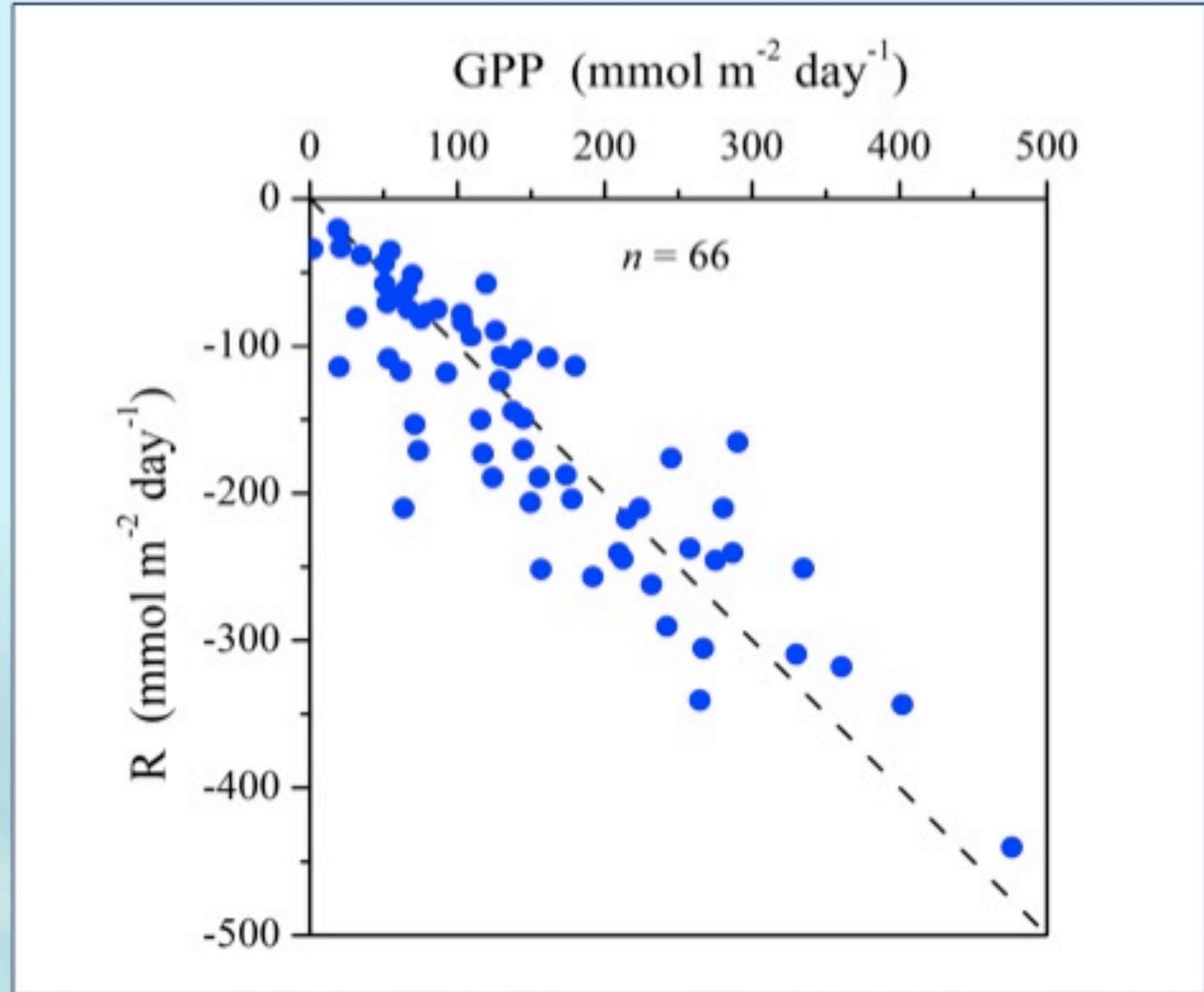
Hume et al. (2011) L&O



Delgard et al. (In prep A)



# Seagrass metabolism

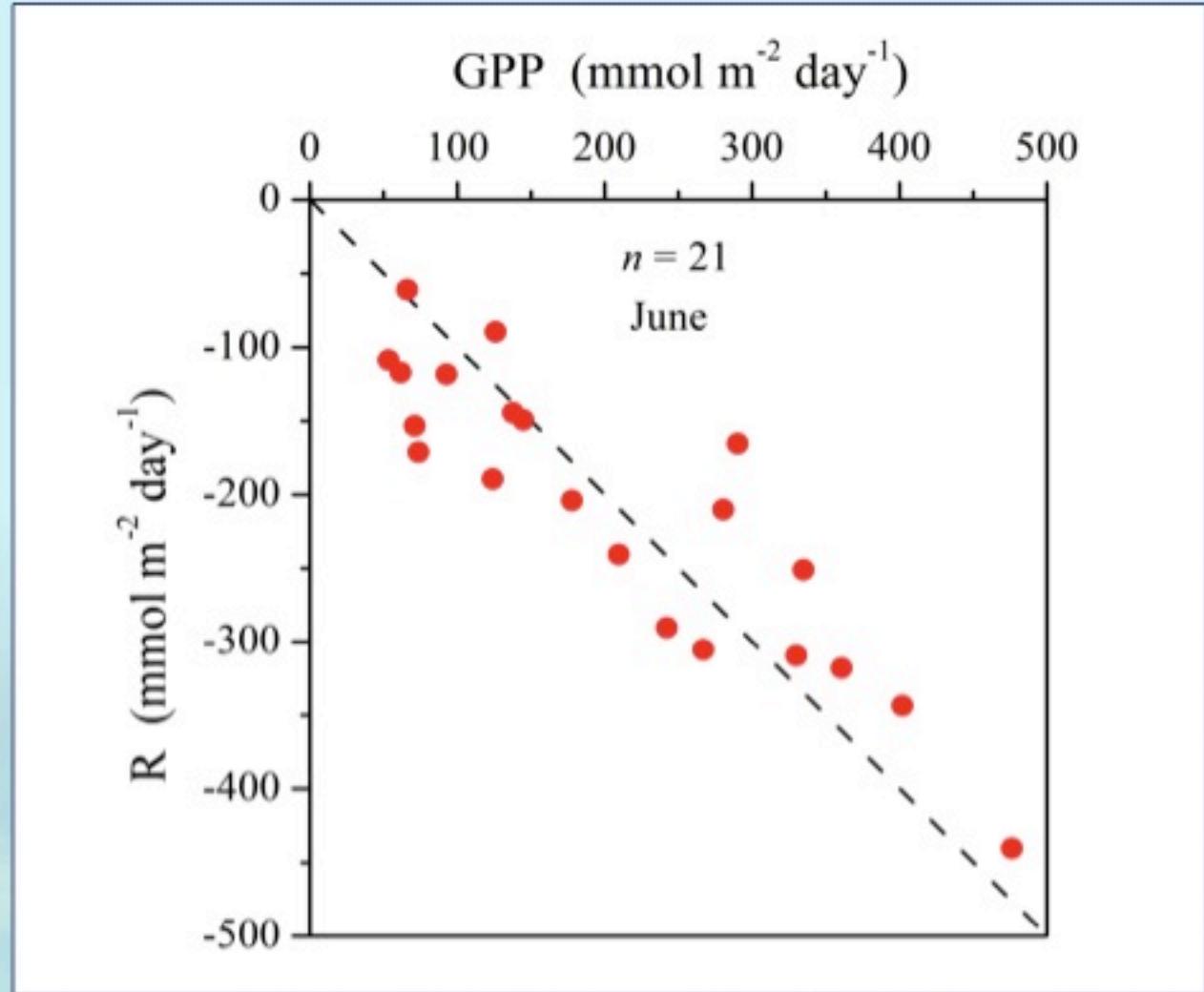


Hume et al. (2011) L&O; Rheuban et al. (2014) MEPS; Delgard et al. (In prep B)

- Tight coupling between R and GPP
- System in metabolic balance



# Seagrass metabolism

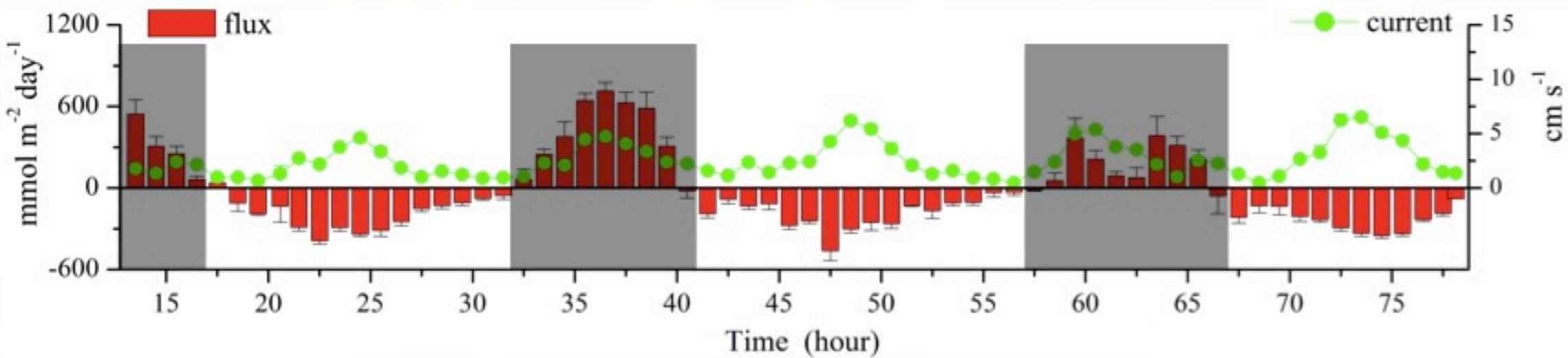
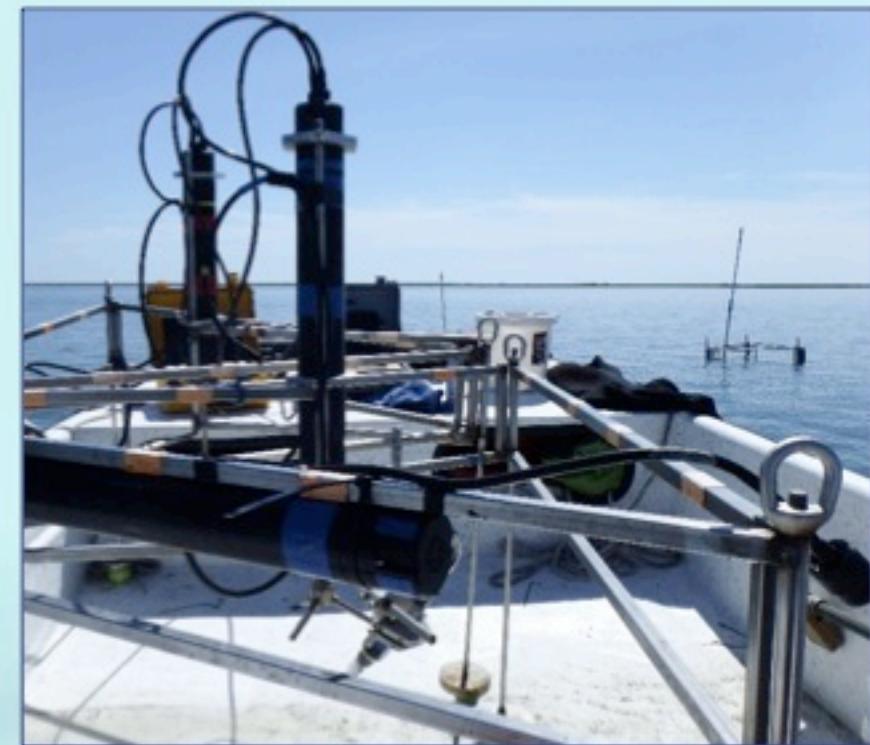


Hume et al. (2011) L&O; Rheuban et al. (2014) MEPS; Delgard et al. (In prep B)

- Variation in R and GPP in June is almost as large as over entire year
- Challenging to find “good” mean values



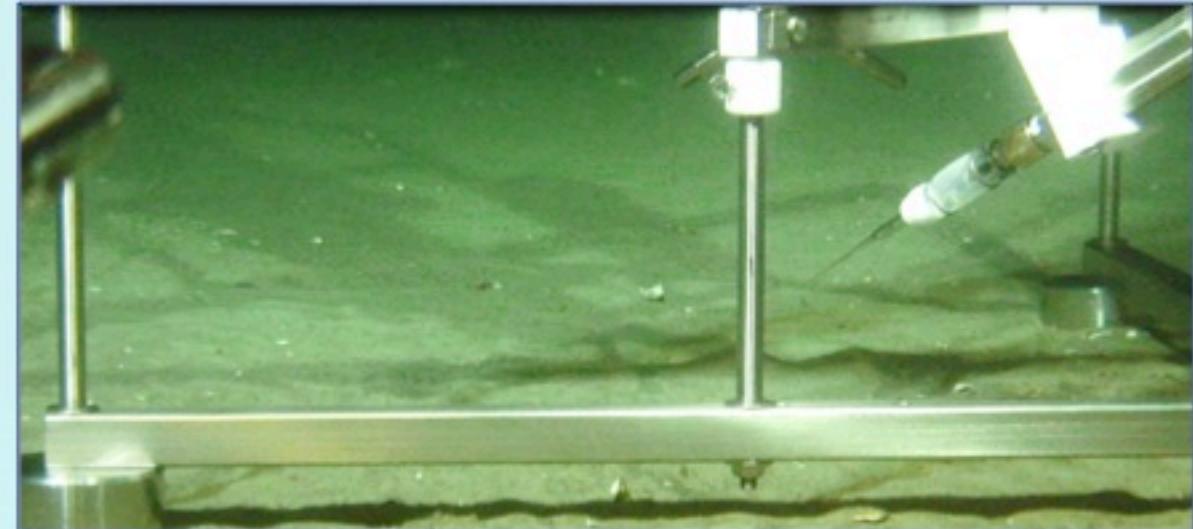
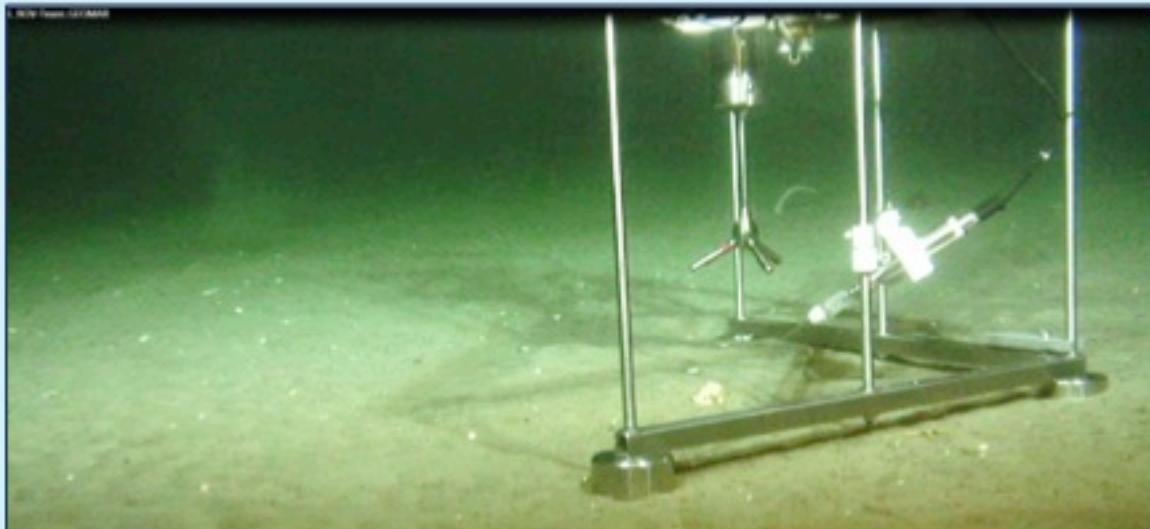
# Effects of natural flow



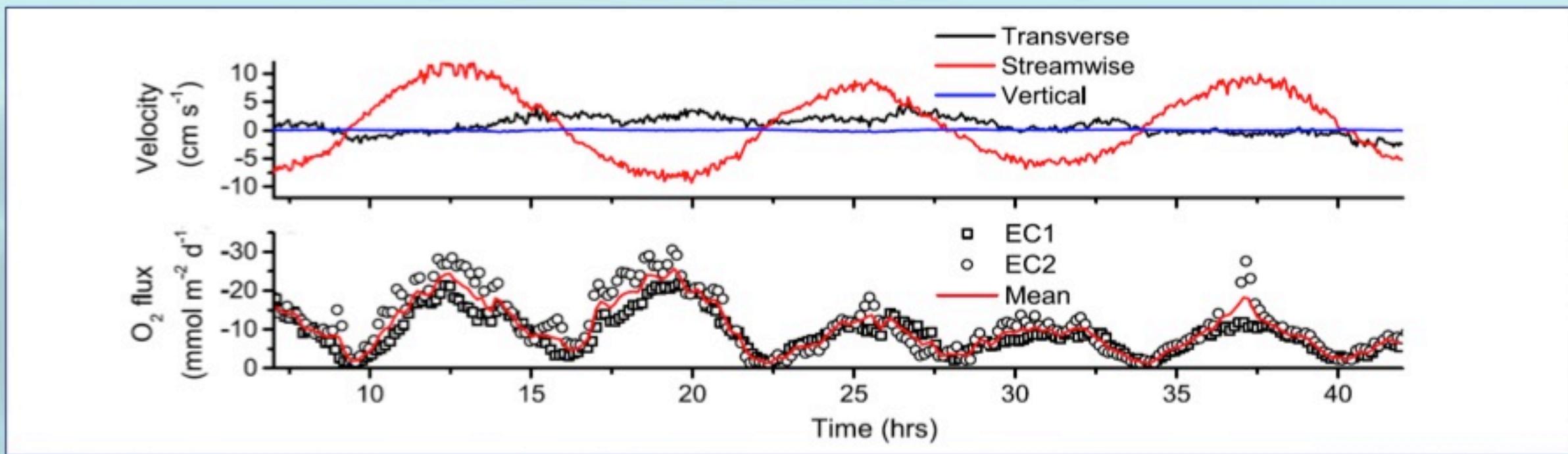
Delgard et al. (In prep A)



# Effects of natural flow



Photos: GEOMAR ROV Team

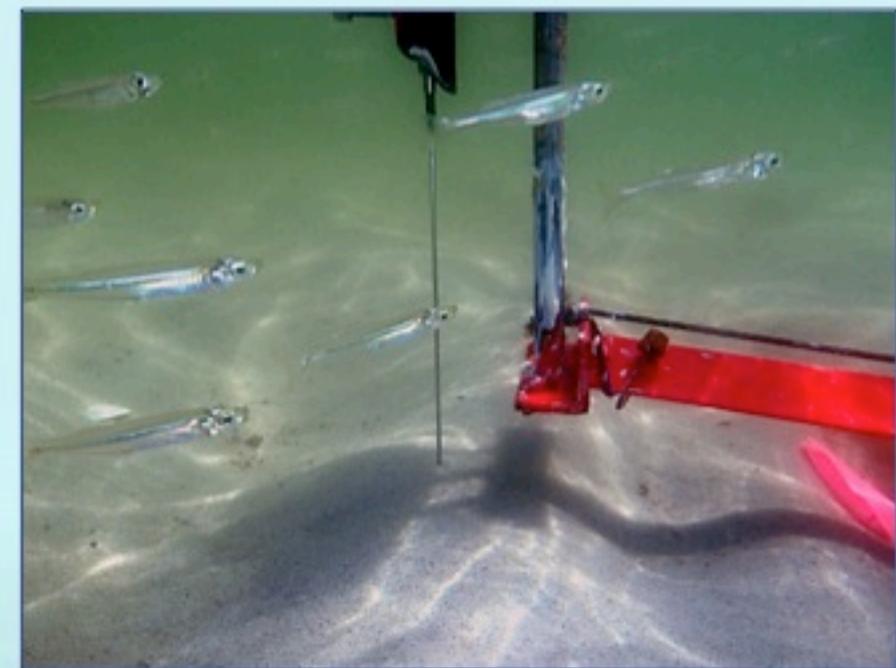


McGinnis et al. (2014)

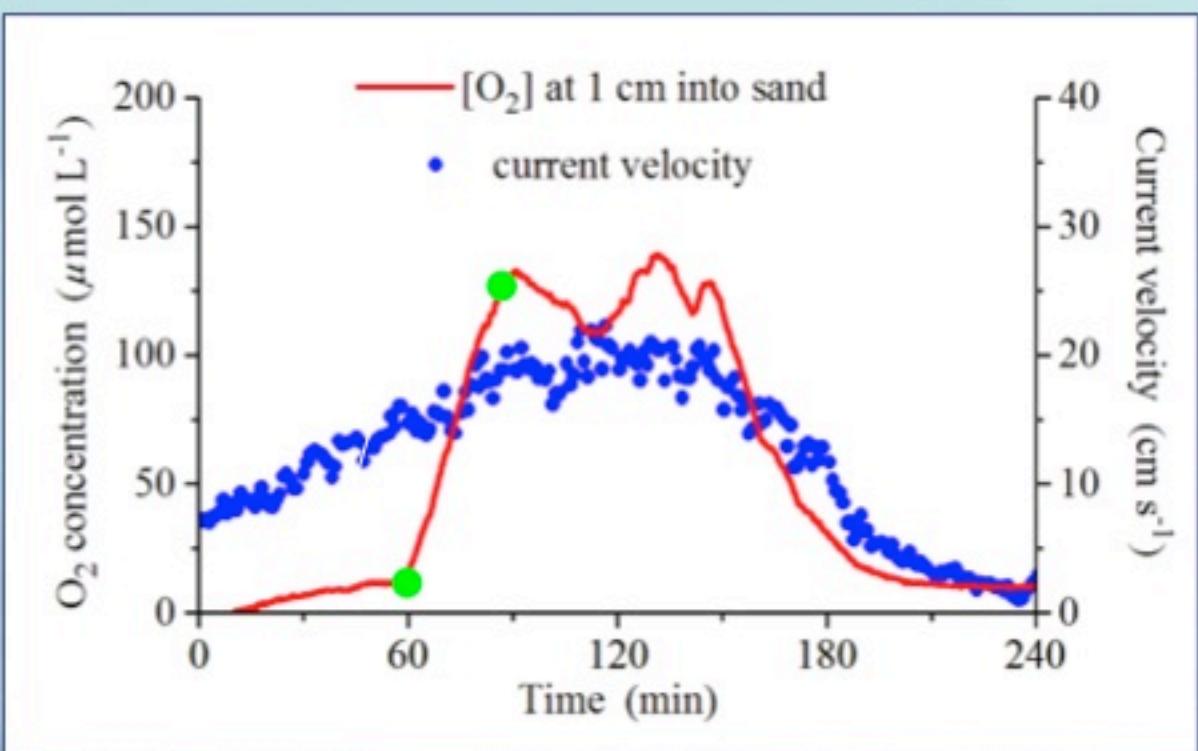
- Flow over permeable sands can stimulate O<sub>2</sub> flux, here up to 25-fold



# Effects of natural flow



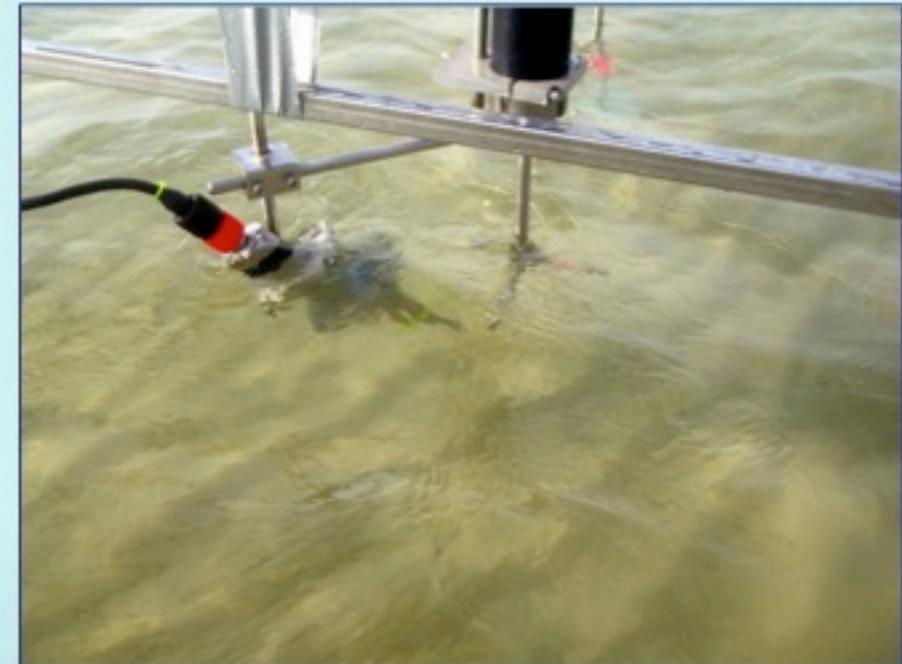
- Flushing of anoxic pore water out of sediment equals flux of  $\sim -50 \text{ mmol m}^{-2} \text{ day}^{-1}$
- Flushing of reduced compounds (e.g.,  $\text{H}_2\text{S}$ ,  $\text{Fe}^{2+}$ ,  $\text{Mn}^{2+}$ ) reacting with oxygen adds to this flux



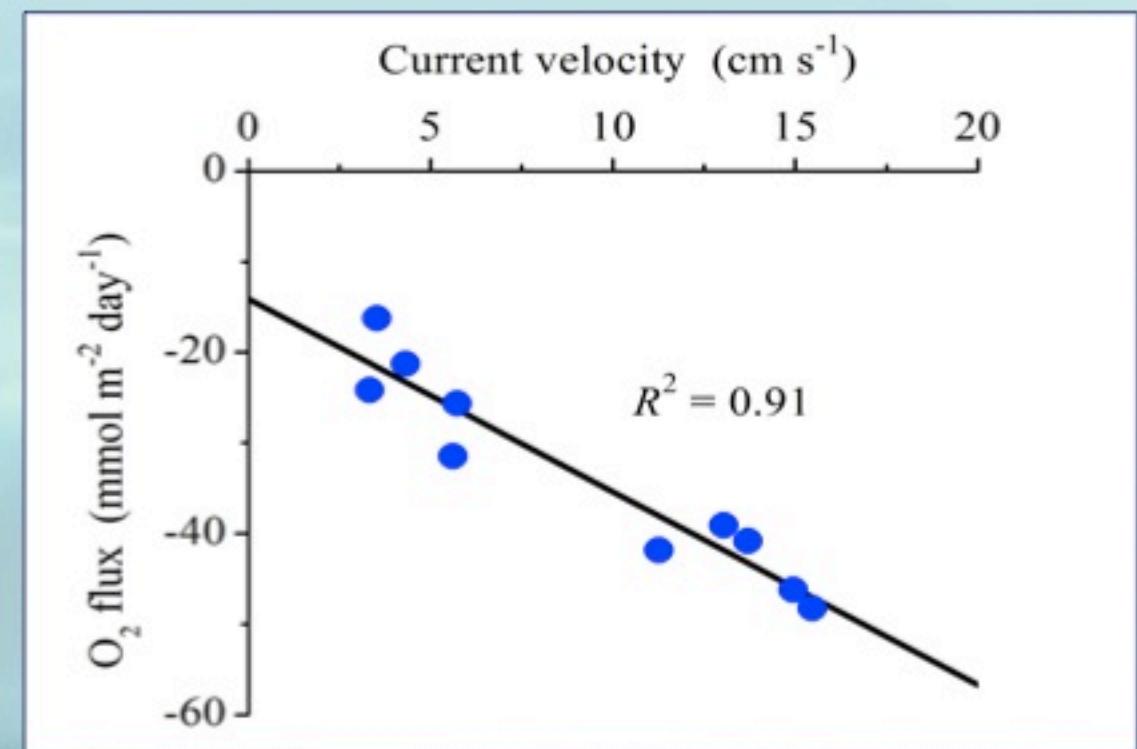
Berg et al. (2013) L&O



# Effects of natural flow



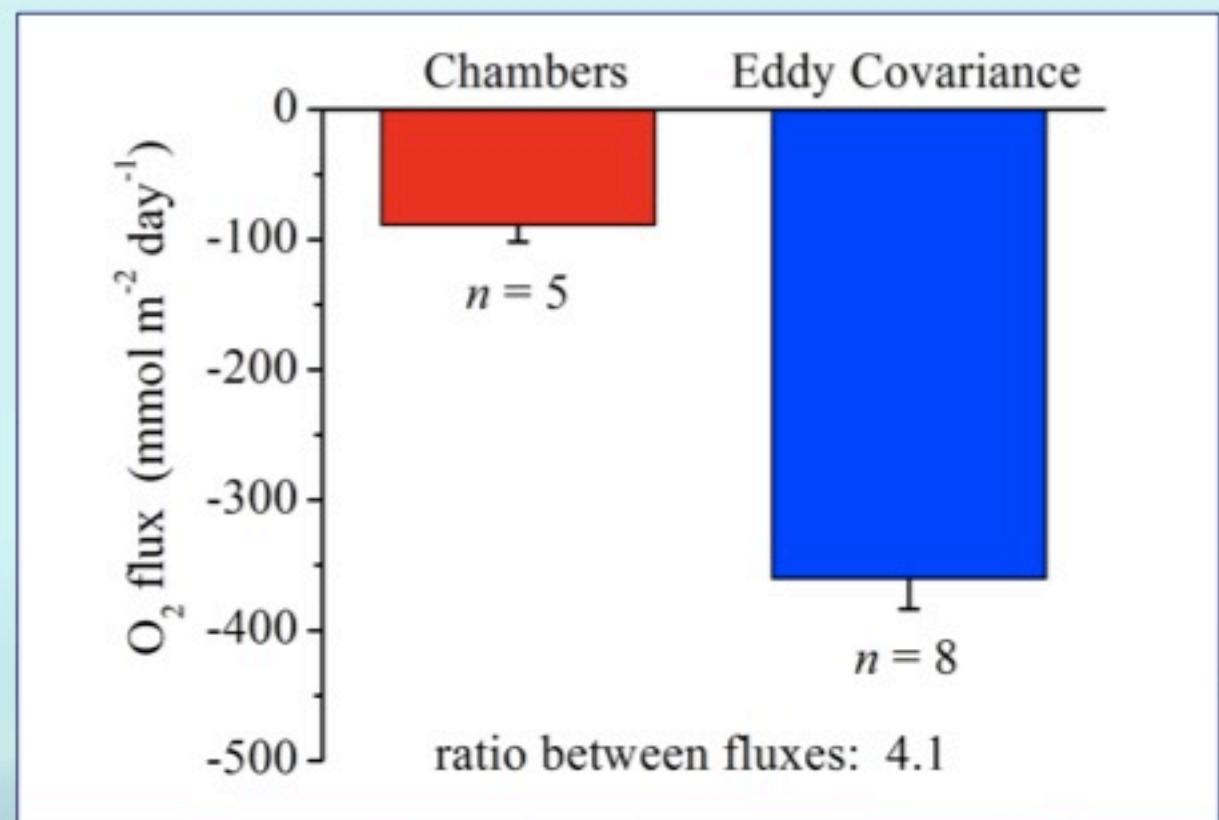
- Fluxes averaged over several hours still contain strong current flow signal
- Fluxes increase ~4-fold with flow increase from 1 to 20 cm s<sup>-1</sup>



Berg et al. (2013) L&O

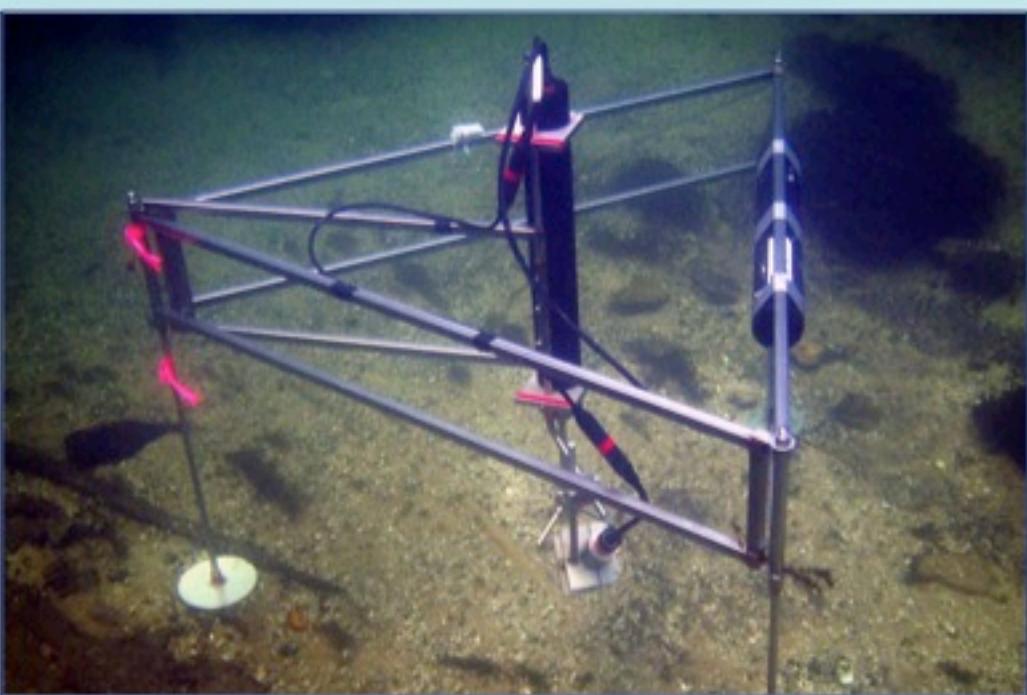


# Effects of natural flow



Berg et al. (2013) L&O

- Permeability of sand:  $2.3 \times 10^{-10} \text{ m}^2$
- Current velocity:  $31 \text{ cm s}^{-1}$
- Use chambers with caution in highly permeable sands at high current flows





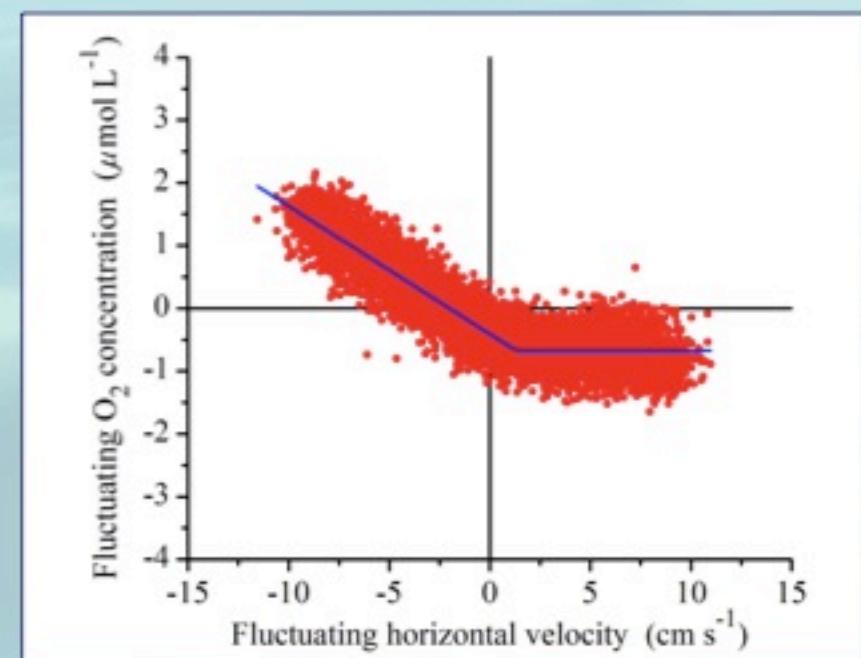
# Sensors for eddy covariance

## Clark-type glass microelectrode:

- Great, well-tested sensor
- Used in almost all studies published so far
- But fragile and difficult to work with
- Expensive (\$700)
- Often a microelectrode is broken or the tip fouled within a 24-hour deployment
- Stirring sensitive



Photo: Sarah Larsen, Dan McGinnis



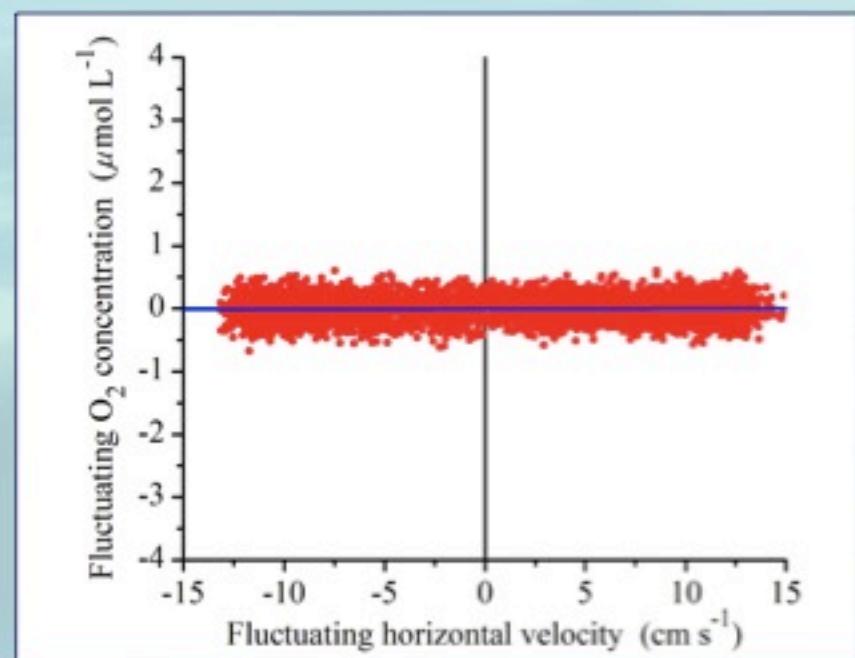
Berg et al. (2016). L&O Methods



# Sensors for eddy covariance

## RINKO micro sensor:

- New dual sensor (oxygen and temp.)
- Developed specifically for eddy covariance (Berg et al. 2015. L&O methods)
- 10mm tip size, robust
- Disturbs flow for some flow directions
- Has not shown signs of stirring sensitivity in tests



Berg et al. (2016). L&O Methods



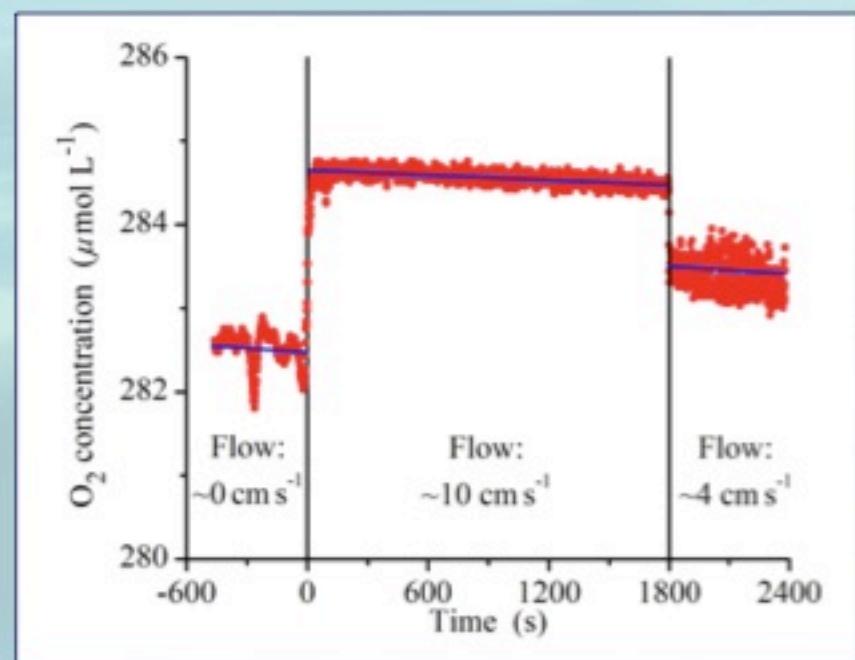
# Sensors for eddy covariance

## Fiber optode:

- Used in a few new studies
- More robust than Clark-type electrodes
- Still only lasts a few days due to breakage or bleaching
- Somewhat expensive (\$360)
- Can apparently be stirring-sensitive. More work is needed to understand why.



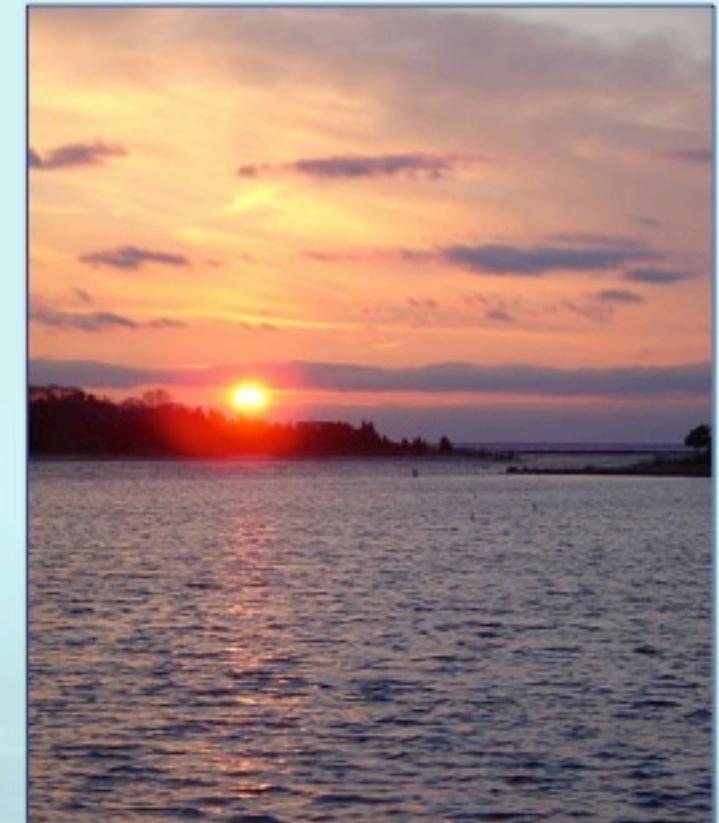
Photo: Dirk Koopmans



Glud, et al. (unpubl)



## New applications – air-water exchange



Air-water exchange of gases ( $O_2$ ,  $CO_2$ , and  $CH_4$ ) for lakes, rivers, and estuaries is critical for whole-system estimates of:

- Primary production and respiration
- Carbon budgets
- Emission of greenhouse gases

$$\text{Flux} = k (C_{\text{air}} - C_{\text{water}})$$



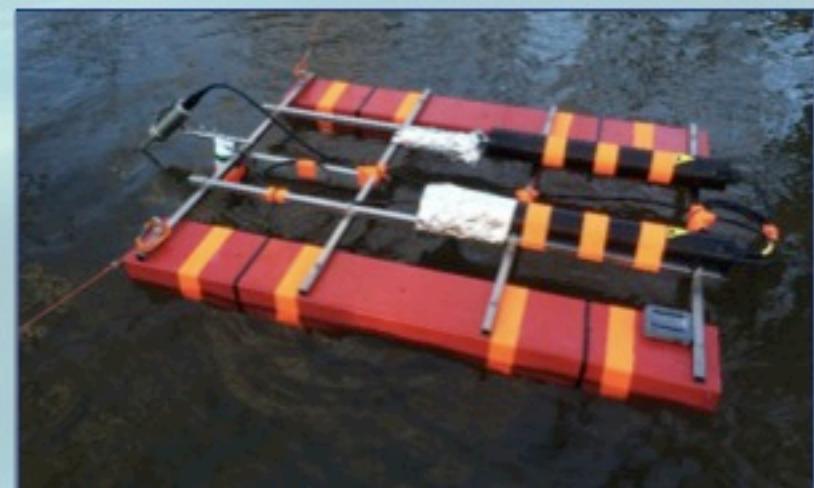
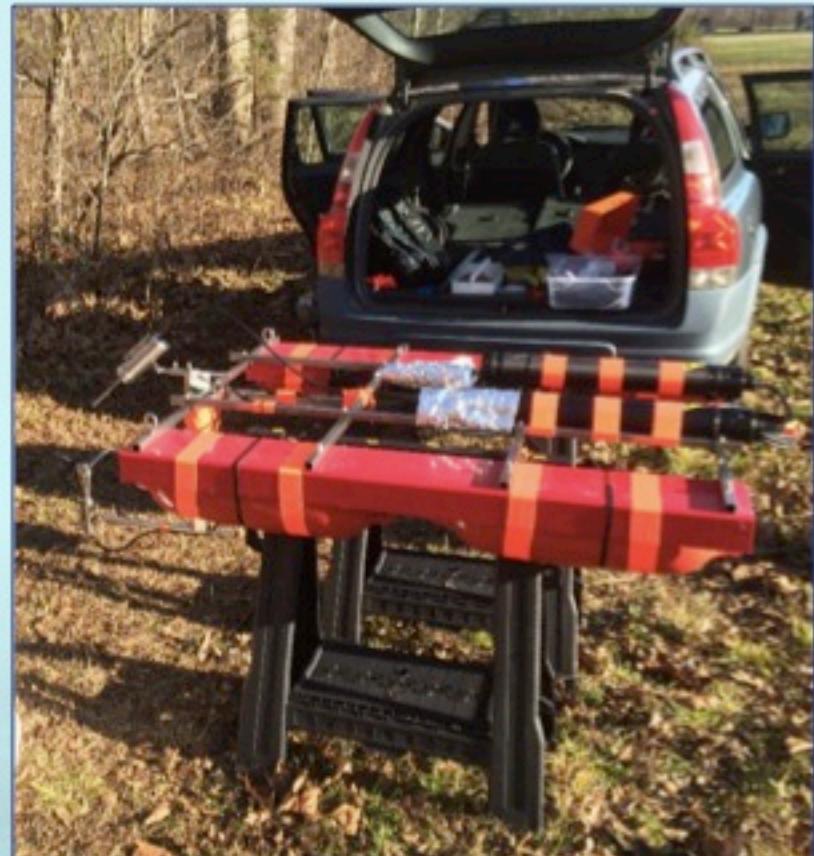
# New applications – air-water exchange

$$\text{Flux} = k (C_{\text{air}} - C_{\text{water}})$$

- Means of  $C_{\text{air}}$  and  $C_{\text{water}}$  are easy to measure (or calculate) for gases of interest
- Values for  $k$  are difficult to assess
- Uncertainty in  $k$  is often huge

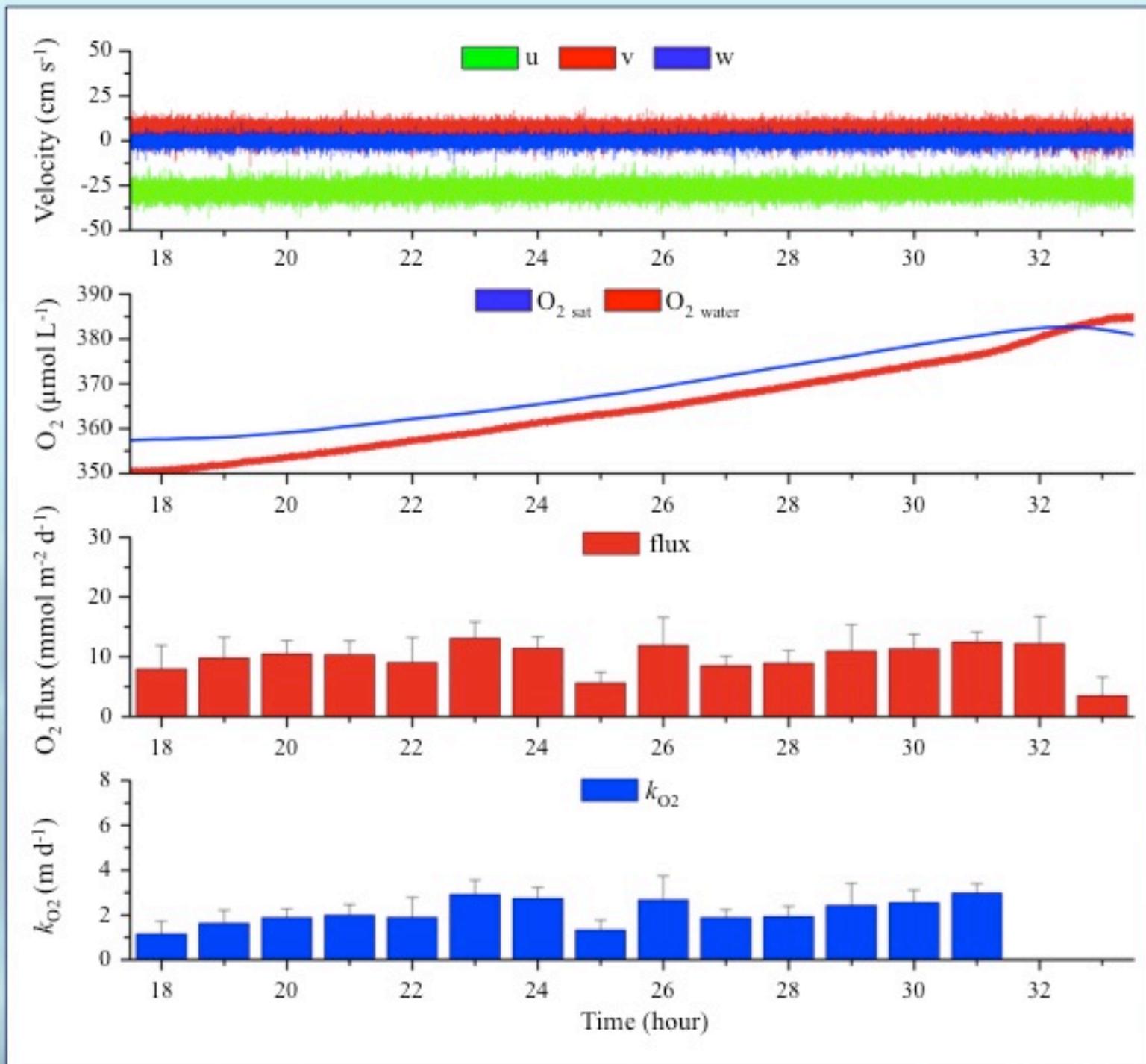
## New idea:

- Measure eddy fluxes of oxygen under air-water interface along with means of  $C_{\text{air}}$  and  $C_{\text{water}}$
- Determine  $k$  for oxygen ( $k_{\text{O}_2}$ )
- Convert  $k_{\text{O}_2}$  to gases of interest ( $\text{CO}_2$ ,  $\text{CH}_4$ , etc.)

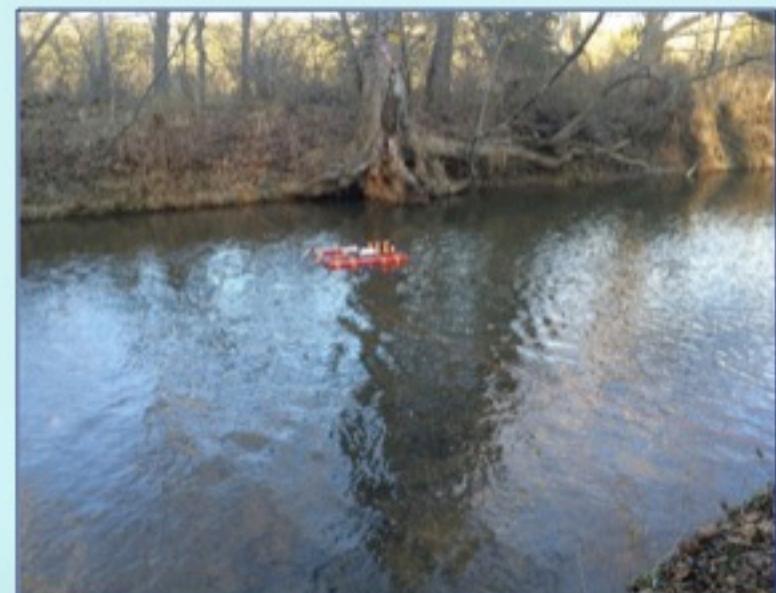




# New applications – air-water exchange

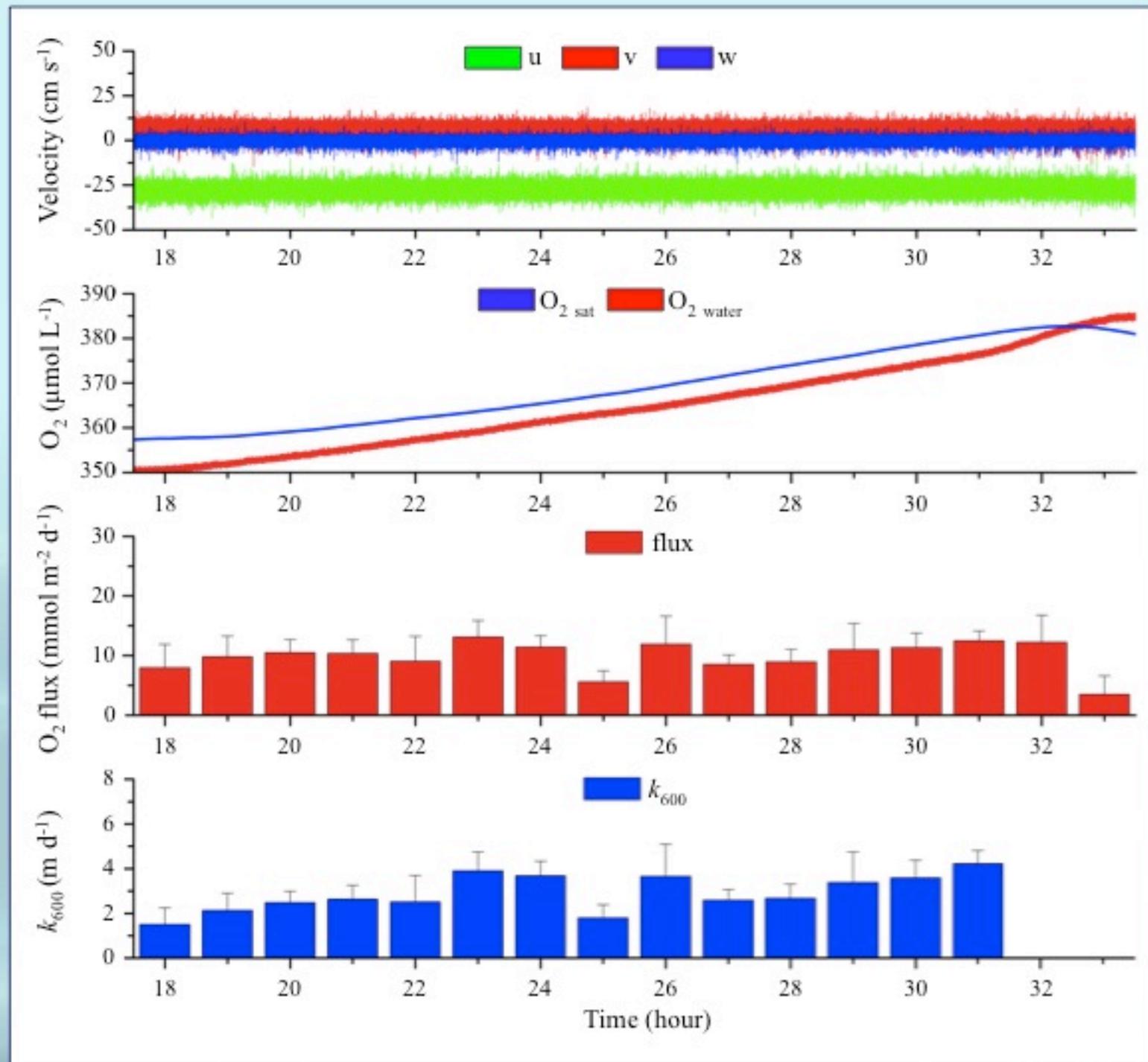


Berg and Pace (In prep)

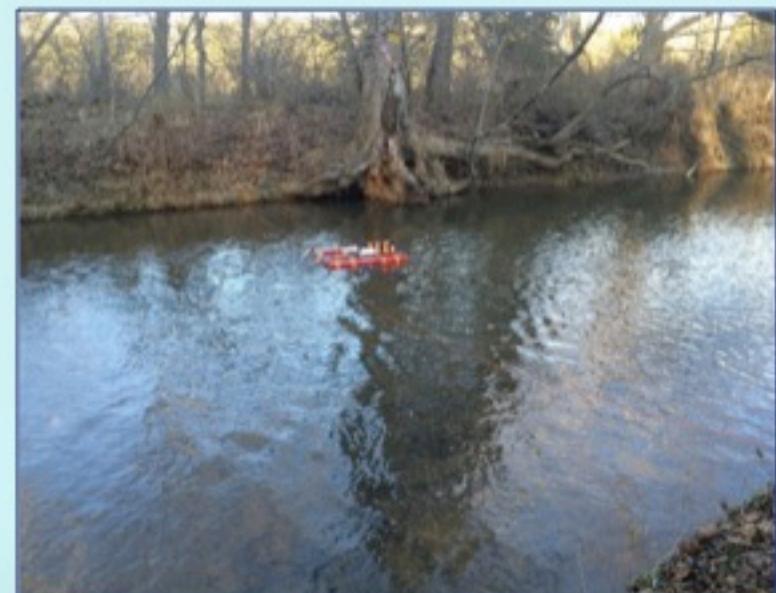




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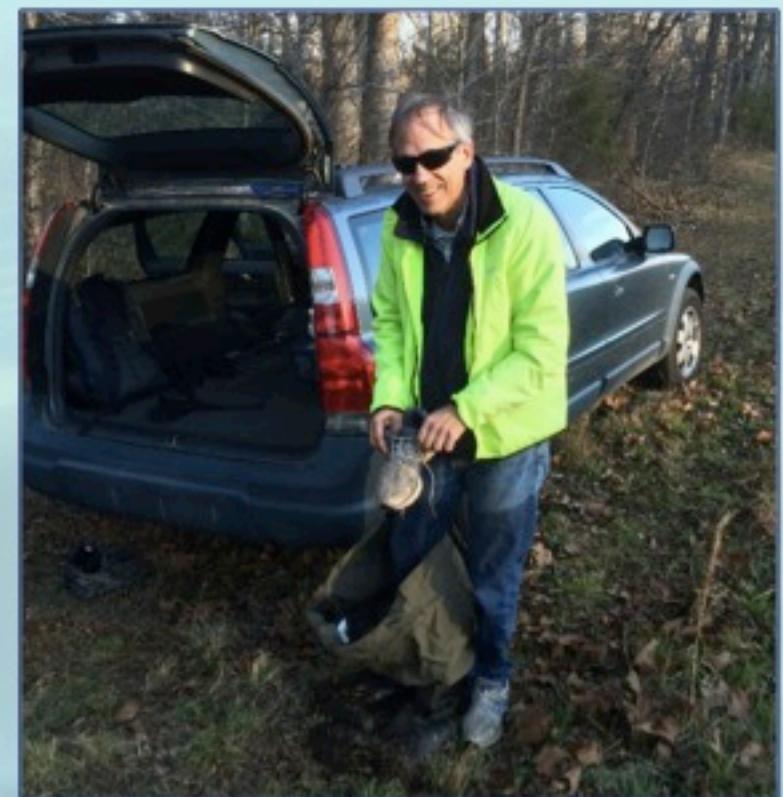
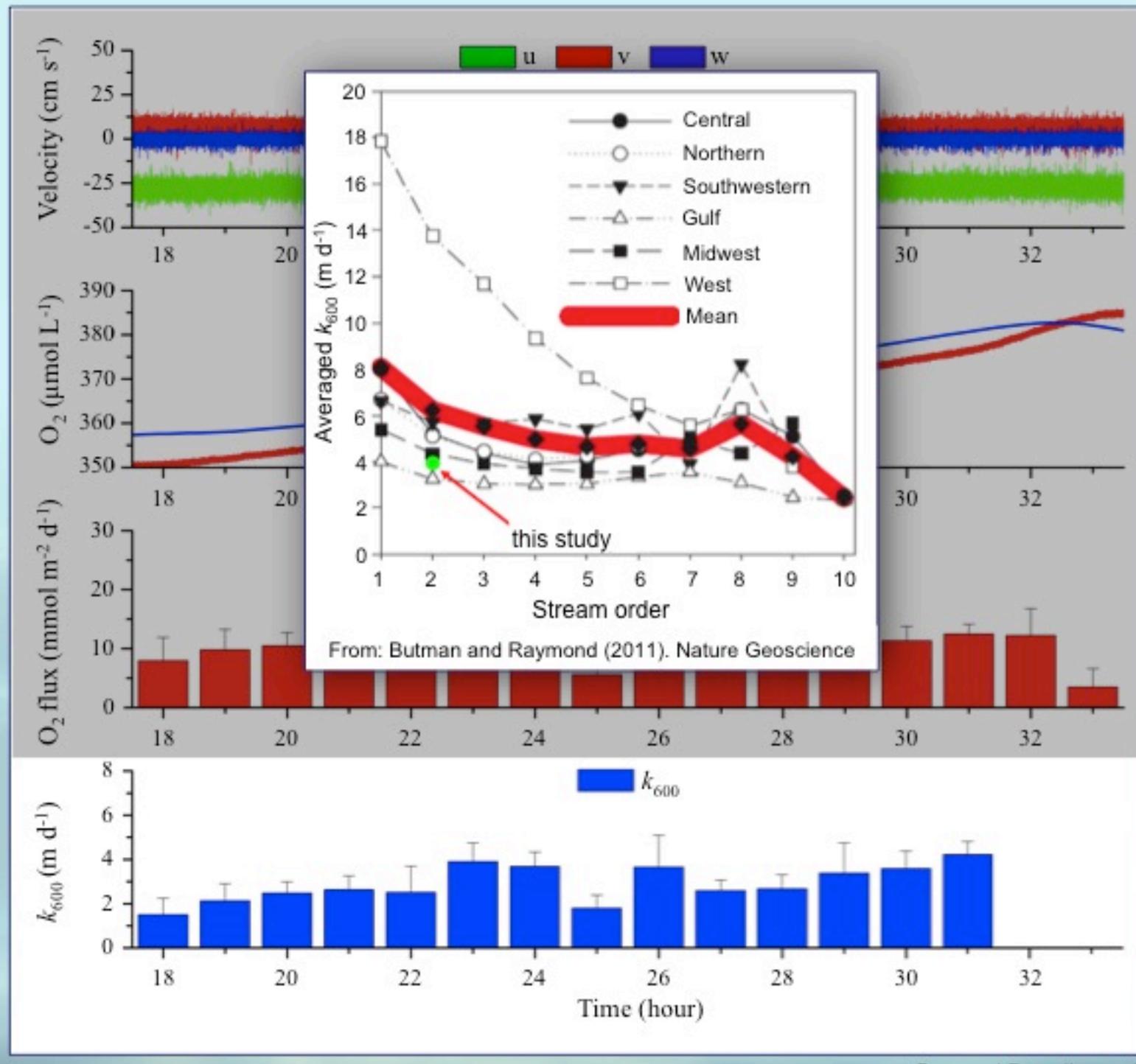


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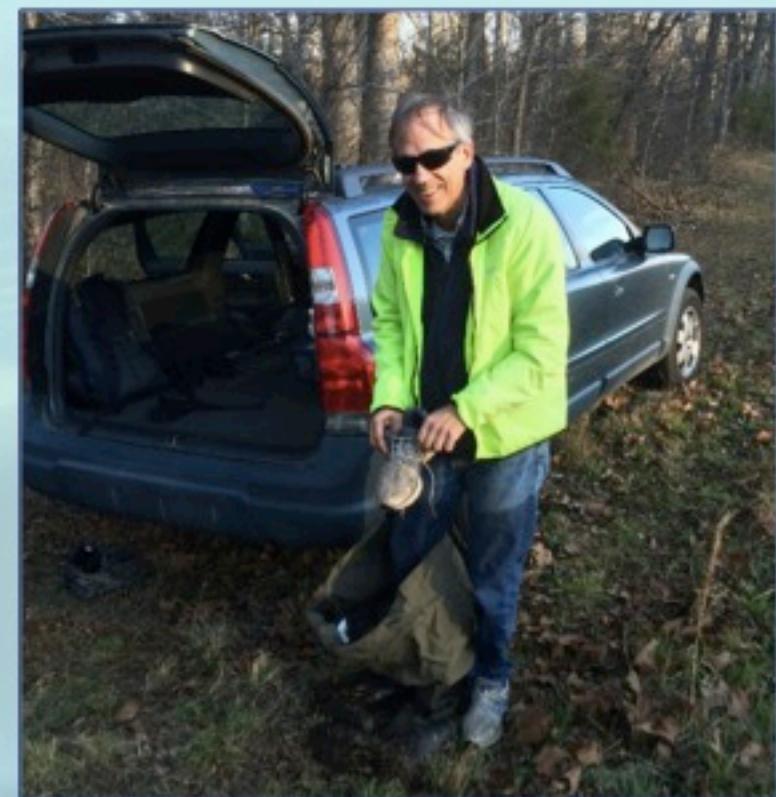
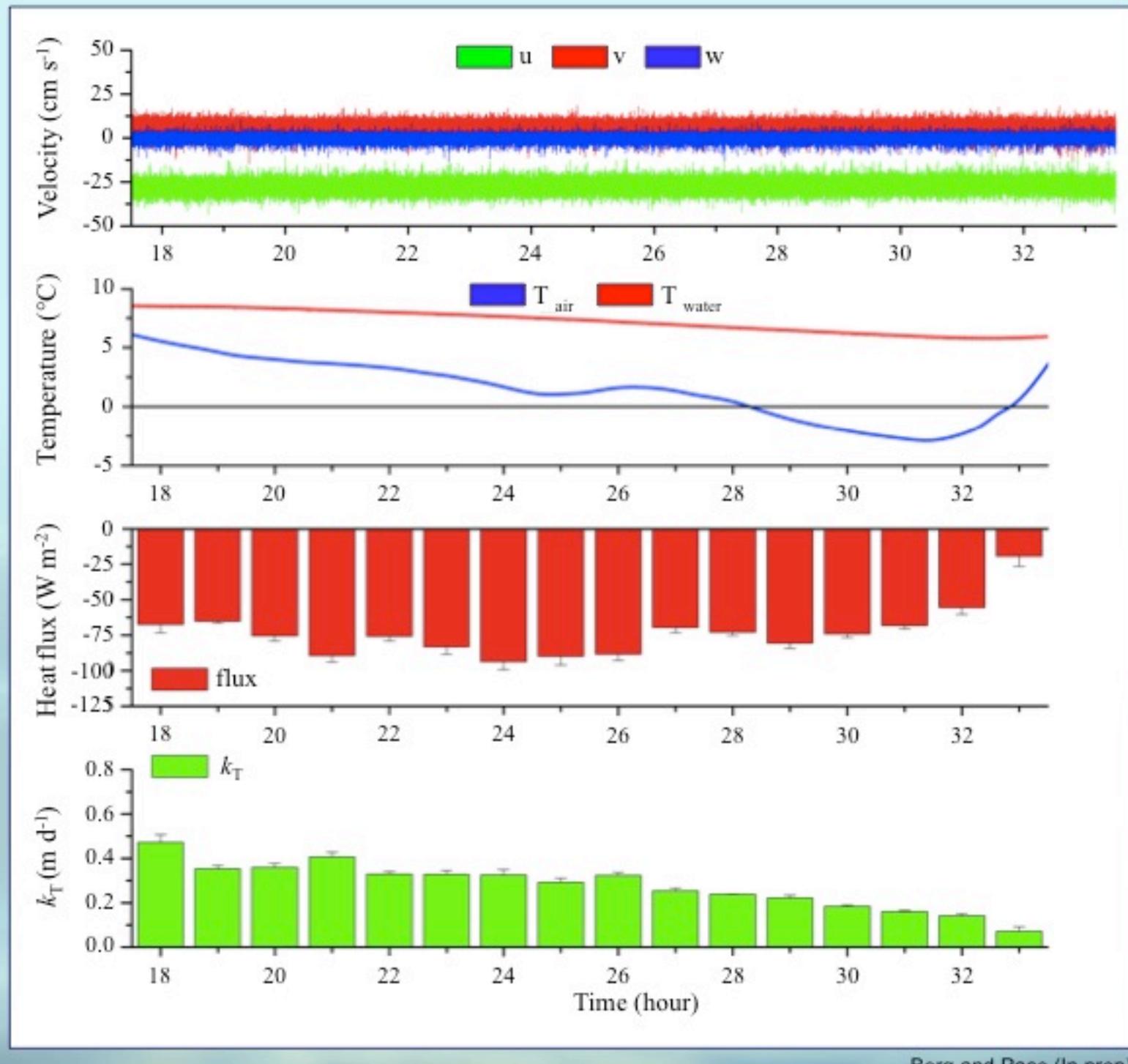


# New applications – air-water exchange



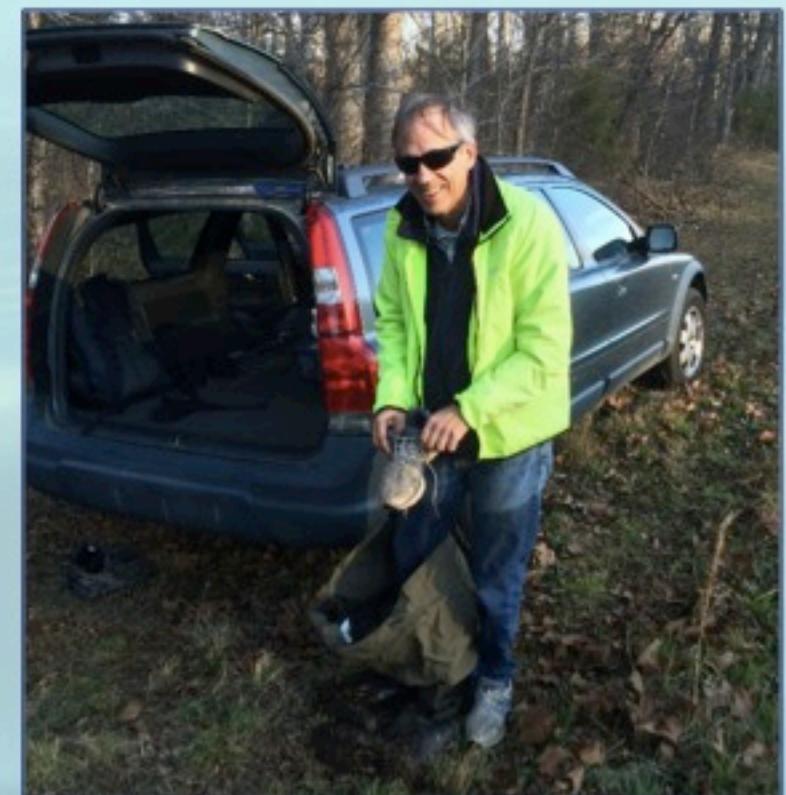
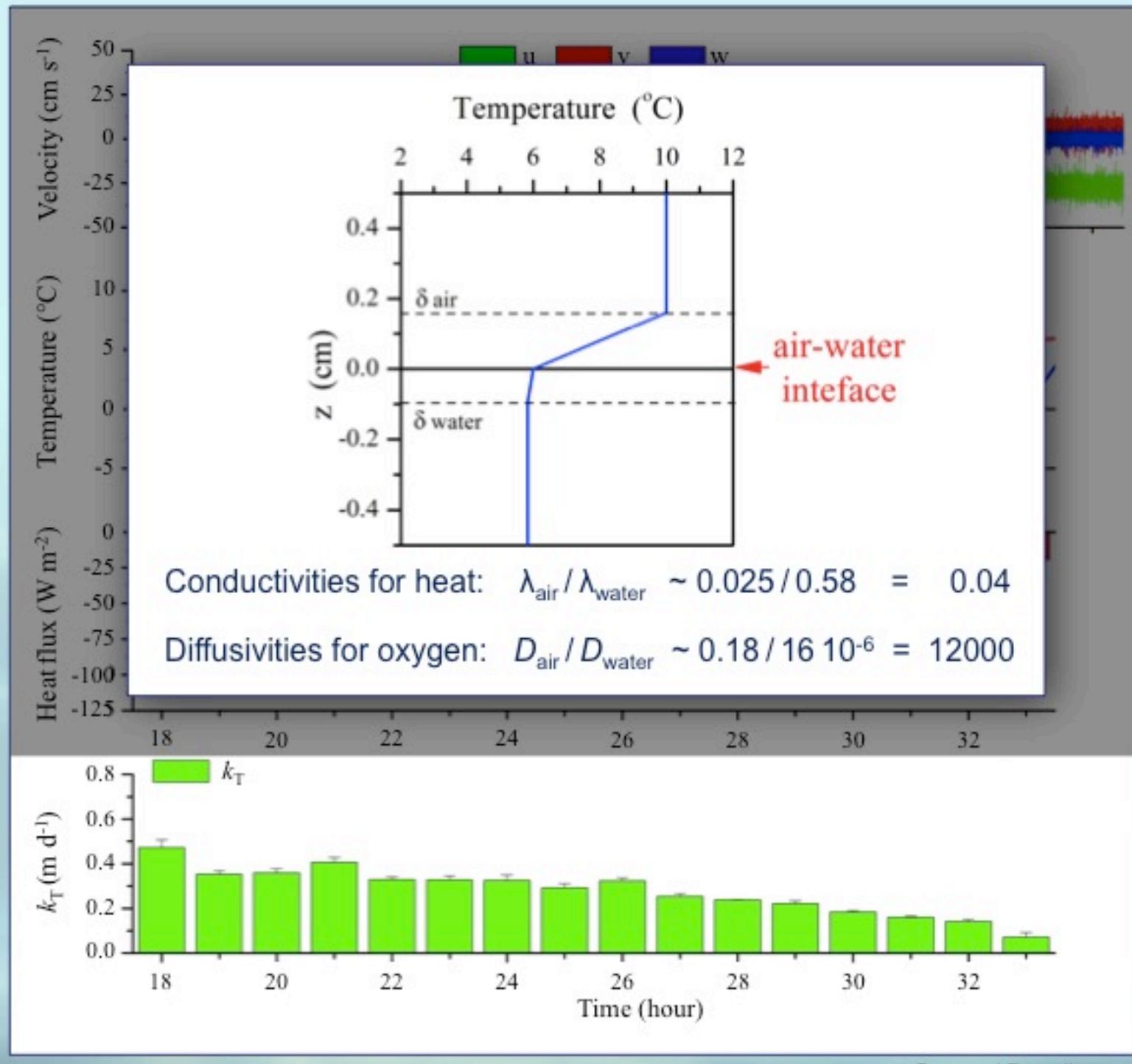


# New applications – air-water exchange



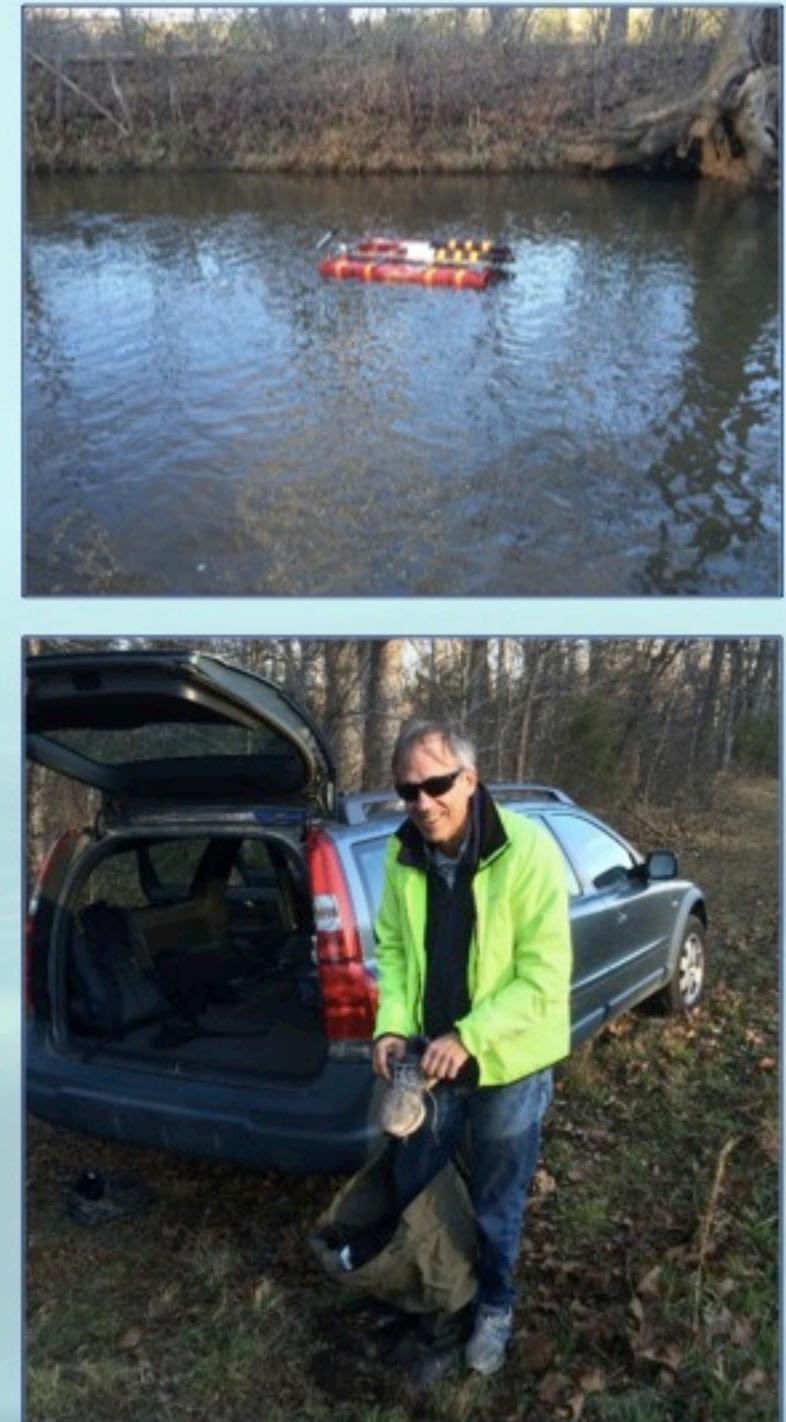
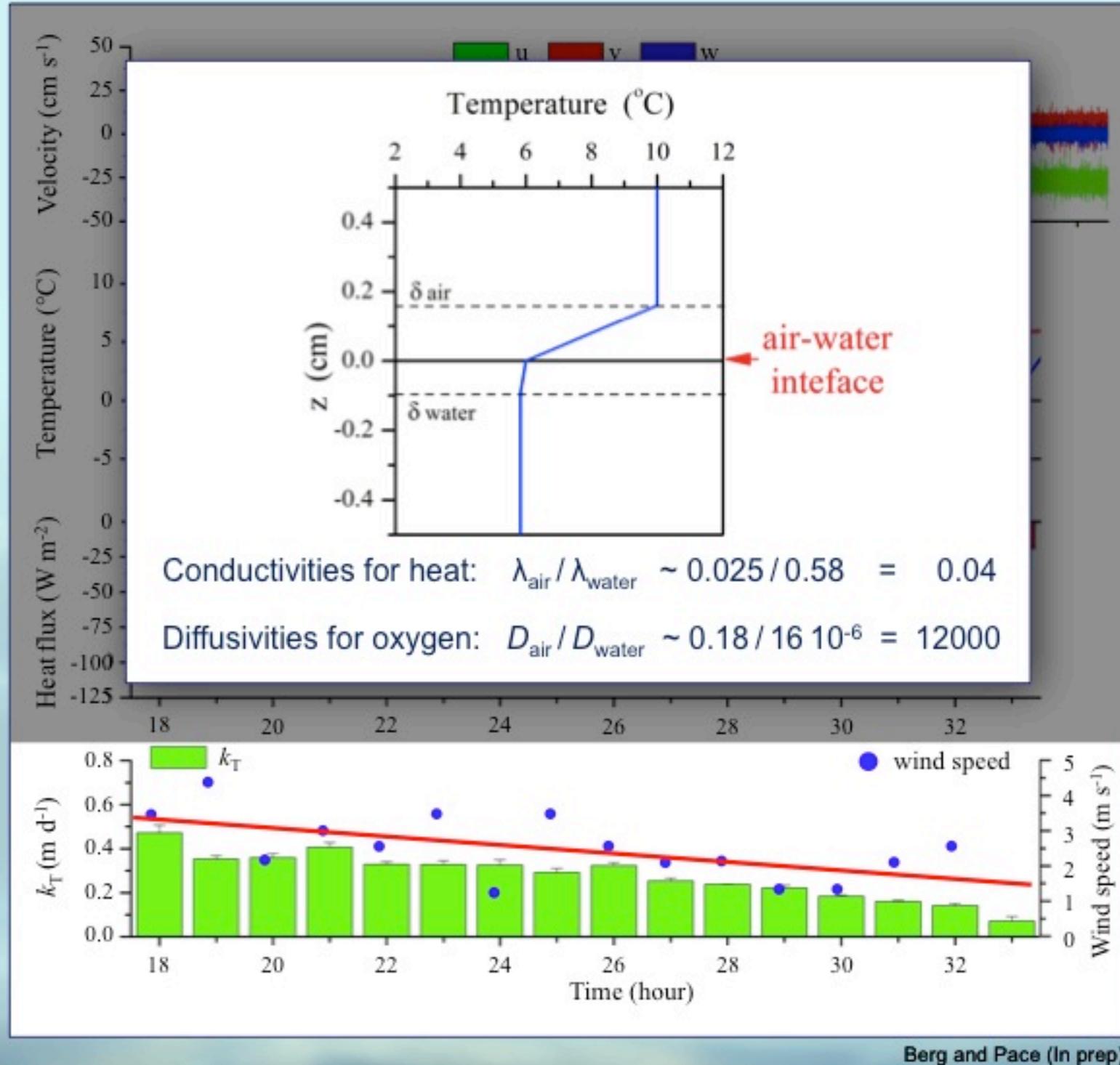


# New applications – air-water exchange





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## New applications – “fly-over”



LI-COR Biosciences



Bluefin Robotics

- Fly-over eddy covariance measurements in air are common today
- A possibility under water? “Fly” over the seafloor or under the air-water interface to get fluxes and exchange coefficients? Maybe...



## Summary and next steps

**Main point:** Aquatic eddy covariance measures fluxes at high temporal resolution under true in situ conditions. Allows studies of whole-system in situ fluxes at a level of detail not possible before.

**Specific foci - develop:**

- New more robust sensors – including for other constituents than oxygen
- Plug-and-play instruments
- Flux extraction procedures for imperfect field conditions
- New applications, mobile platforms, and observatory systems
- Courses and workshops to discuss issues and developments, and to train new users
- Open access databases



# Acknowledgments/funding



National Science Foundation



University of  
Virginia



University of  
Southern Denmark



Florida State  
University



Oregon State  
University



# Teamwork





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*Non-invasive Flux Measurements at the Benthic Interface:  
The Aquatic Eddy Covariance Technique*

doi:10.1002/loe2.10005

e-Lecture received June 2016; accepted May 2017

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