# Development of RANS boundary condition for AMR-Wind

Document all notes here.

# 1 RANS ABL grid stretching

The physical domain coordinates x and the computational domain  $\xi$  are related by the mapping function  $x(\xi)$ :

$$x(\xi) = x_{geom}(\xi)(1 - h(\xi)) + h(\xi)x_{const}(\xi)$$

where the blend function  $h(\xi)$  is defined as

$$h(\xi, \xi_T, W) = \frac{1}{2} \left( 1 + \tanh\left(\frac{\xi - \xi_T}{W}\right) \right)$$

and depends on the transition location  $\xi_T$  and width W.

$$\frac{dx}{d\xi} = (1 - h)x'_{geom} - h'x_{geom} + hx'_{const} + h'x_{const}$$

### 1.1 Geometry growth mapping function

The geometric growth function  $x_{geom}$  is defined as :

$$x_{geom}(\xi) = \frac{\Delta_0}{r} \frac{1-r^{\xi+1}}{1-r} - \frac{\Delta_0}{r}$$

With the first cell height  $\Delta_0$  and stretch rate r > 1 provided by the user. The inverse of this function is

$$\xi_{inv}(x) = \frac{\ln\left(1 - (x - \Delta_0/r)\frac{r(1-r)}{\Delta_0}\right)}{\ln(r)}$$

The derivative of this function is:

$$\frac{dx_{geom}}{d\xi} = \frac{\Delta_0}{r(r-1)} r^{\xi-1} \ln(r)$$

#### 1.2 Constant cell size mapping function

$$x_{const} = s(\xi - \xi_0) + L_{match}$$

slope s, offset  $\xi_0$ , and match height  $L_{match}$ .

To set up the constants in  $x_{const}$ , there are two options:

1. We can compute the  $\xi$  that would correspond to L input by the user:

$$\xi_L = \inf\{x_{inv}(L, \delta_0, r)\}\$$

2. We can compute the  $\xi$  that would correspond to  $\Delta_{max}$  input by the user:

$$\xi_L = \text{floor}\left[\frac{\ln(\Delta_{max}r^2/\Delta_0)}{\ln r}\right] - 1$$

We'll set the offset  $\xi_0 = \xi_L$ . Because  $\xi_L$  has been rounded to the nearest integer, we'll compute  $L_{match}$  at  $\xi_L$ 

$$L_{match} = x_{geom}(\xi_L, \delta_0, r)$$

And the slope will be the cell size at  $\xi_L$ :

$$s = x_{geom}(\xi_L, \delta_0, r) - x_{geom}(\xi_L - 1, \delta_0, r)$$

The derivative of  $x_{const}$  is

$$\frac{dx_{const}}{d\xi} = s$$

## 2 RANS k- $\epsilon$ ABL boundary conditions

Implemented in MOData.cpp as calc\_phi\_m\_alinot():

$$\phi_m(\zeta) = \begin{cases} (1 - \beta_m \zeta)^{-1/4}, & \zeta < 0\\ 1 + \gamma_m \zeta & \zeta > 0 \end{cases}$$
 (1)

where  $\zeta = z/L$  (see [1])

Implemented in MOData.cpp as calc\_phi\_eps\_alinot():

$$\phi_{\epsilon}(\zeta) = \begin{cases} 1 - \zeta & \zeta < 0\\ \phi_{m}(\zeta) - \zeta & \zeta > 0 \end{cases}$$
 (2)

The following calculations are implemented in ShearStress.H. As ShearStressAlinot.calc\_mu()

$$\mu_{t0}(z) = \frac{\rho \kappa u_* z}{\phi_m(\zeta)} \tag{3}$$

As ShearStressAlinot.calc\_eps()

$$\epsilon_0(z) = \frac{u_*^3}{\kappa z} \phi_{\epsilon}(\zeta) \tag{4}$$

As ShearStressAlinot.calc\_tke()

$$k_0(z) = \sqrt{\frac{\kappa u_* z \epsilon_0(z)}{C_\mu \phi_m}} \tag{5}$$

$$k_0(z) = \sqrt{\frac{\mu_{t0}(z)\epsilon_0(z)}{\rho C_\mu}} \tag{6}$$

 $As \verb| ShearStressAlinot.calc_omega()|$ 

$$\omega_0(z) = \frac{\epsilon_0(z)}{C_\mu k_0(z)} \tag{7}$$

# References

[1] ALINOT, C., AND MASSON, C. k- $\epsilon$  model for the atmospheric boundary layer under various thermal stratifications. *Journal of Solar Energy Engineering* 127, 4 (06 2005), 438–443.