



Model Evaluations I: How to think about and what to expect from dynamical core and GCM tests

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Overview & Questions

- What are the sources of uncertainty in General Circulation Models (GCMs) and dynamical core experiments?
- Review of the typical test hierarchy for GCMs
- What is the particular impact of the dynamical core on the flow, illustrated via a tropical cyclone test
- Review of dry dynamical core tests with examples from the 2008 Dynamical Core Summer Colloquium
- Overview of the DCMIP test cases and example results



Sources of Uncertainty in GCMs

Structural:
choice of physical parameterizations
and dynamical cores, model
resolution, ...

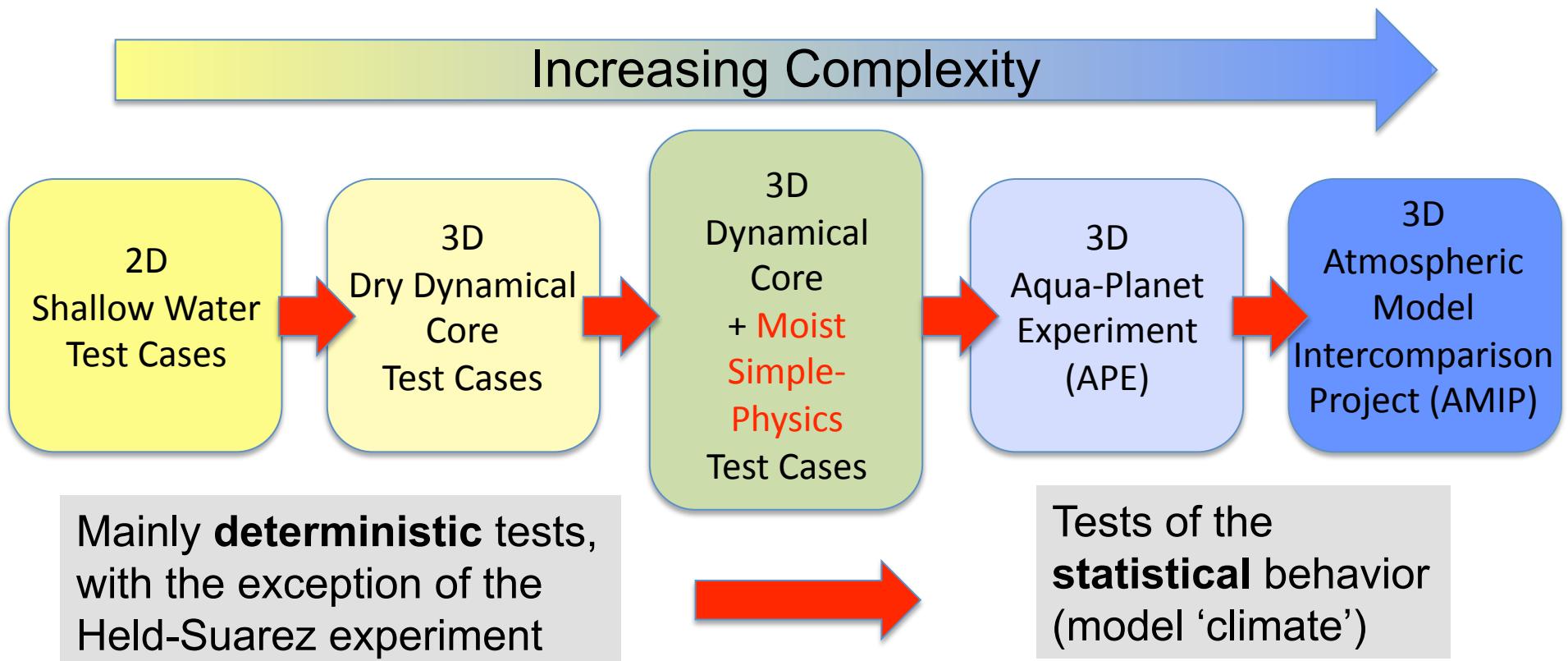
Data:
Initial data,
boundary data
(e.g. SSTs, ...)

Parameter: physics
tuning, physical
constants, diffusion
coefficients, ...



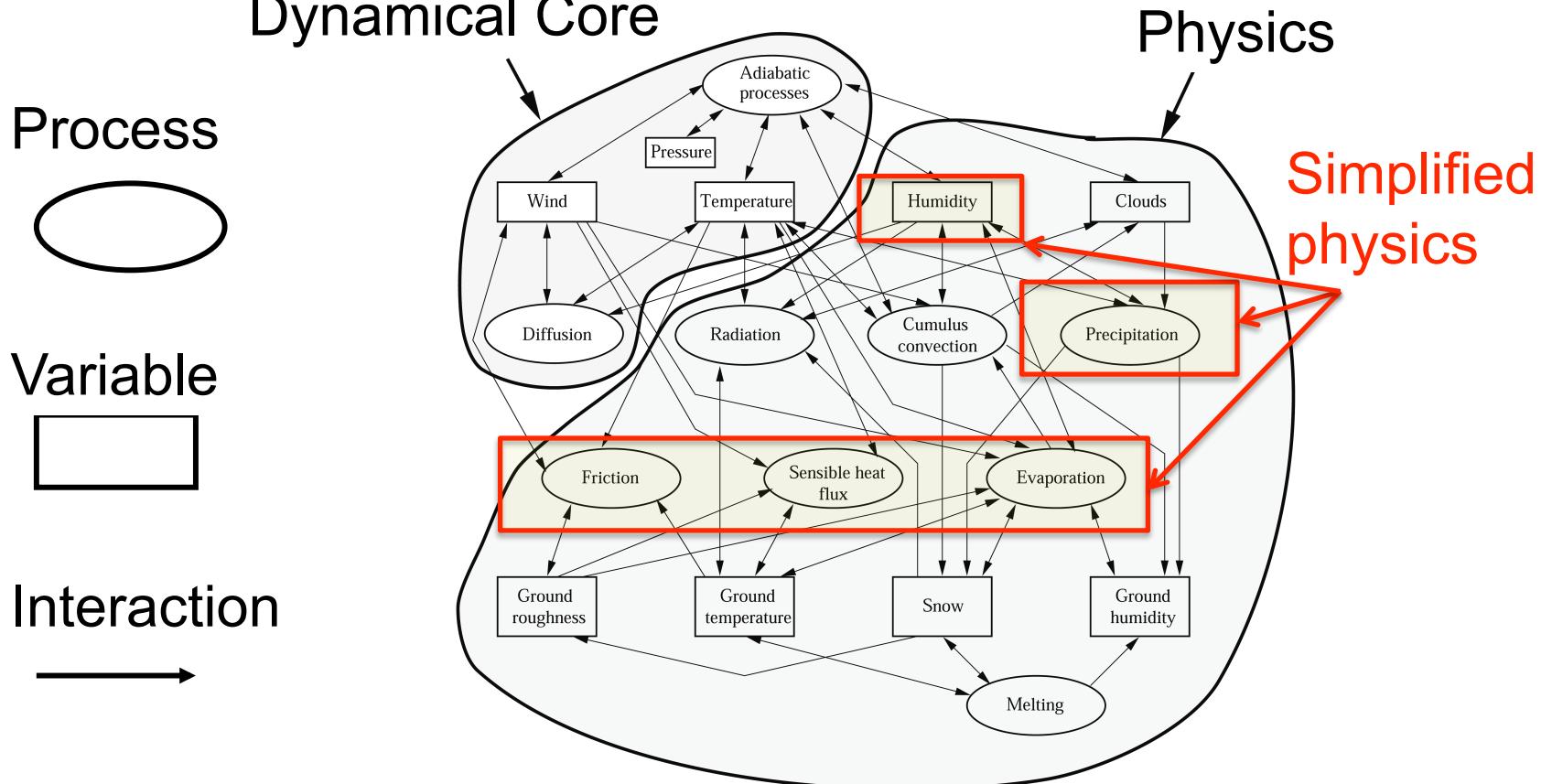
GCM Test Hierarchy

- Typical evaluation hierarchy for Dynamical Core and GCMs assessments





Dynamics & Physics: How to think about simplified physics

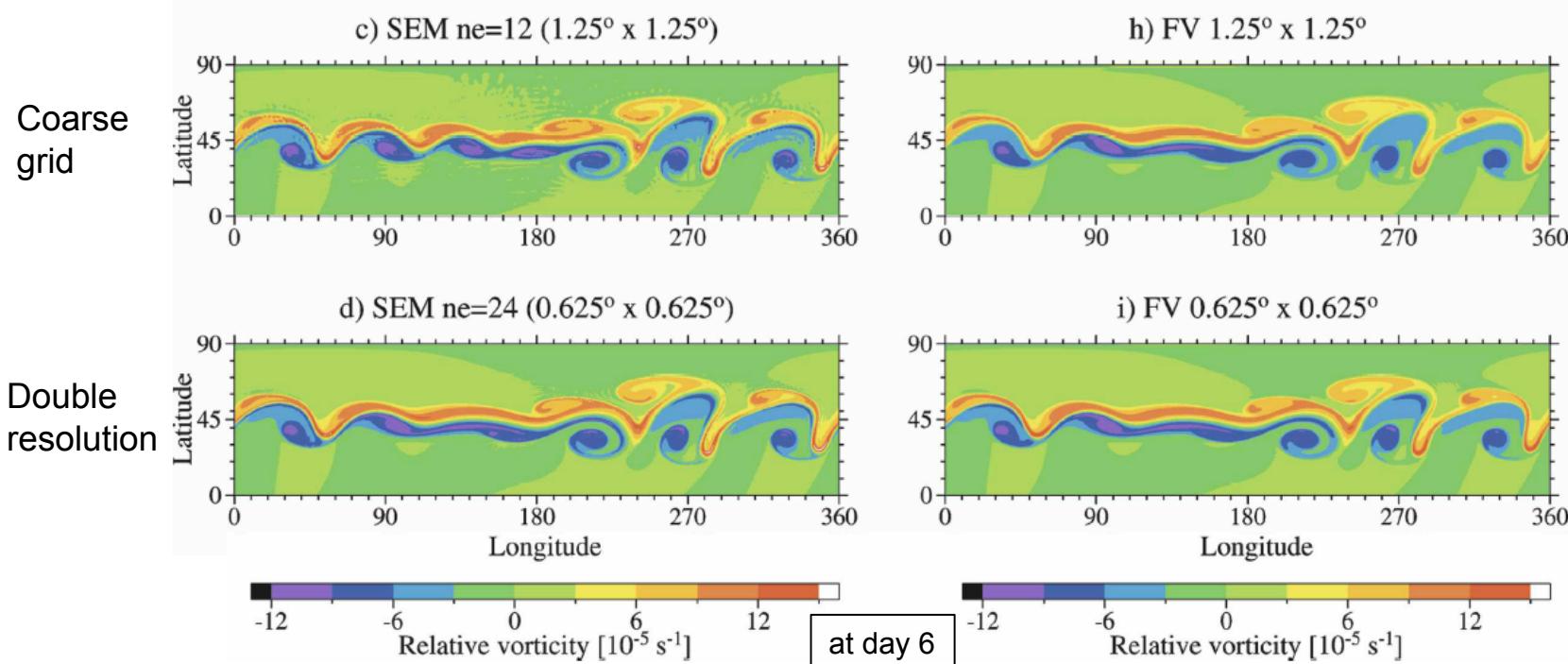


Replace the physics package: use only simple surface fluxes, large-scale condensation and vertical diffusion in BL



GCM Test Hierarchy: 2D SW

- 2D shallow water (SW) tests evaluate the characteristics of the horizontal and temporal discretizations
- An example is the barotropic wave test by Galewsky et al., Tellus (2004), here shown for two SW models SEM and FV

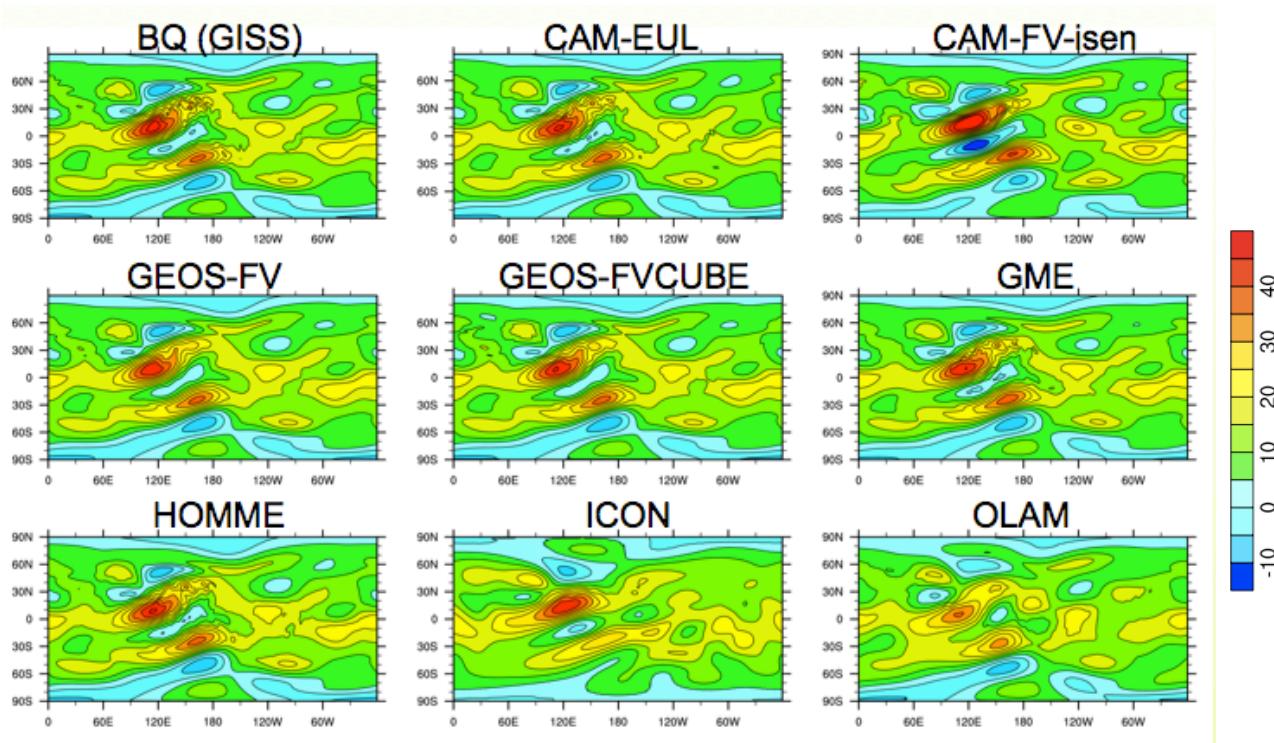


from St-Cyr, Jablonowski, et al., MWR, 2008



GCM Test Hierarchy: 3D dry

- 3D dry dynamical core tests evaluate the characteristics of the horizontal, **vertical** and temporal discretizations



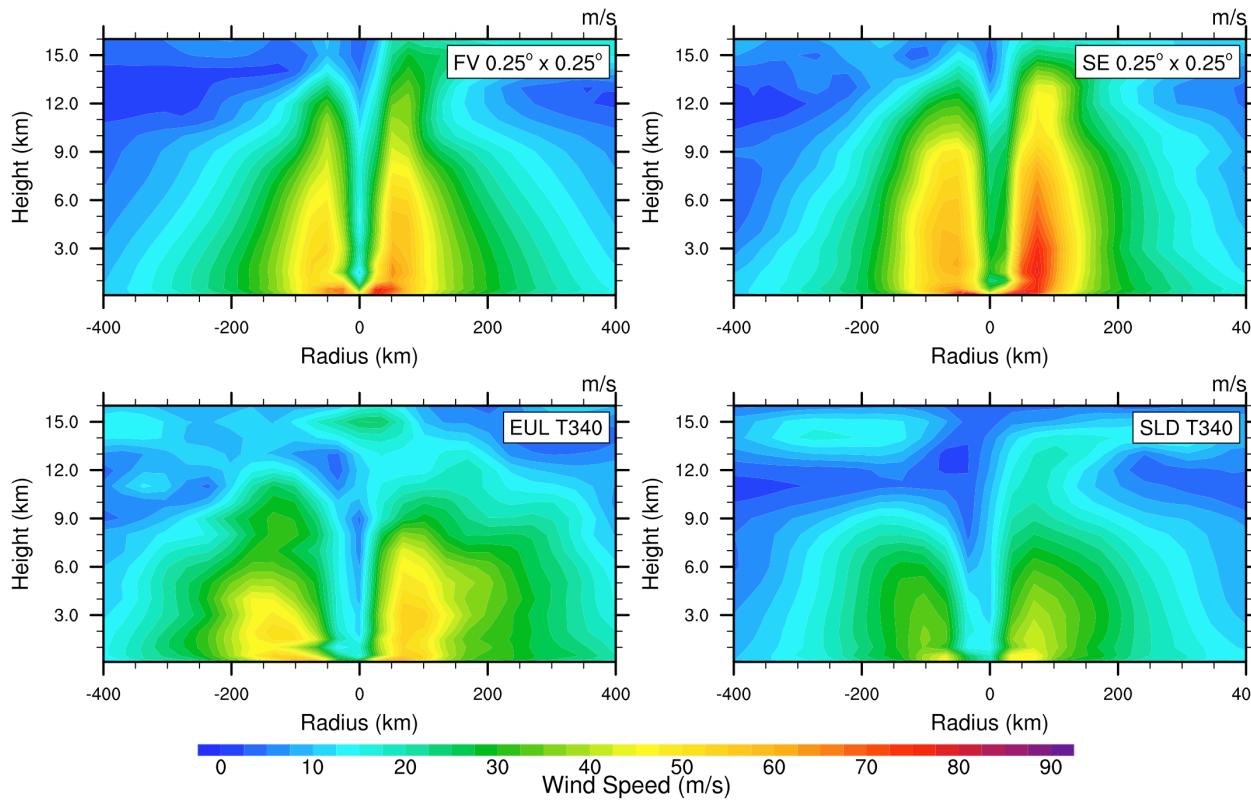
9 different
dynamical
cores

Example from the 2008 Dynamical Core Intercomparison during the NCAR ASP Summer Colloquium: Mountain-generated Rossby waves, 700 hPa zonal wind (m/s) at day 15



GCM Test Hierarchy: 3D moist

- 3D (simplified)-moist dynamical core tests evaluate the characteristics of the horizontal, vertical and temporal discretizations, and **their interaction with simple physical parameterizations**: tropical cyclone example with CAM



4 dynamical cores of the
NCAR Community
Atmosphere Model CAM:

FV & SE at 0.25°
(≈ 28 km)

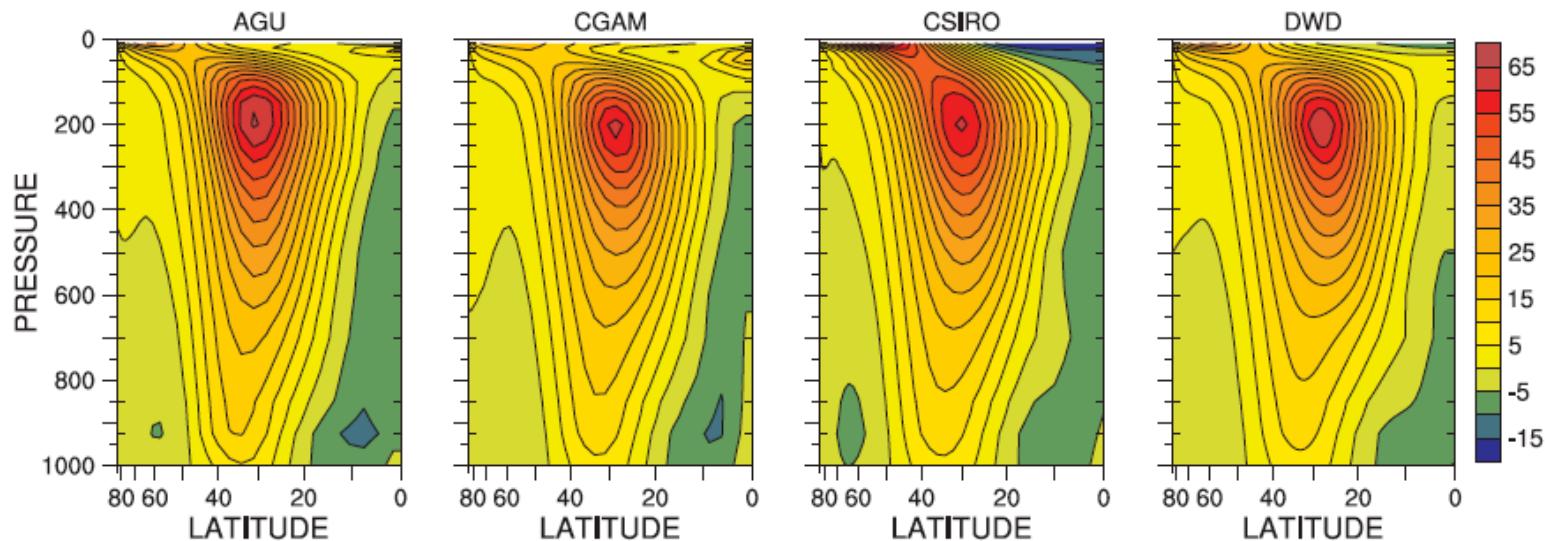
EUL & SLD at T340
(≈ 39 km)

Reed and Jablonowski, JAMES (2012)



GCM Test Hierarchy: Aqua-planet

- 3D aqua-planet tests evaluate the interaction between the dynamical core with **complex physical parameterizations** using a **simplified lower boundary** (flat ocean-covered Earth with analytically prescribed sea surface temperatures (SSTs))



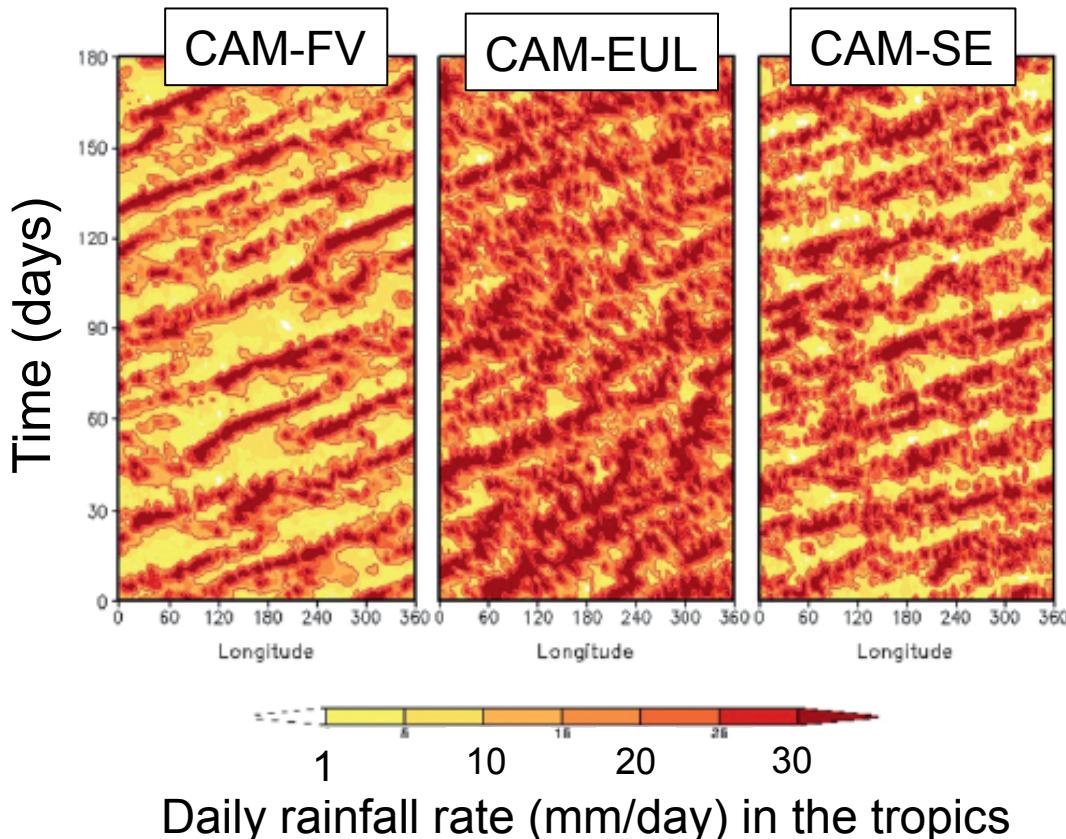
Zonal-mean **3-year-mean** zonal wind: Snapshots of 4 GCMs that participated in the Aqua-Planet Experiment (APE)

Williamson et al., NCAR Technical Note TN-484+STR (2012)

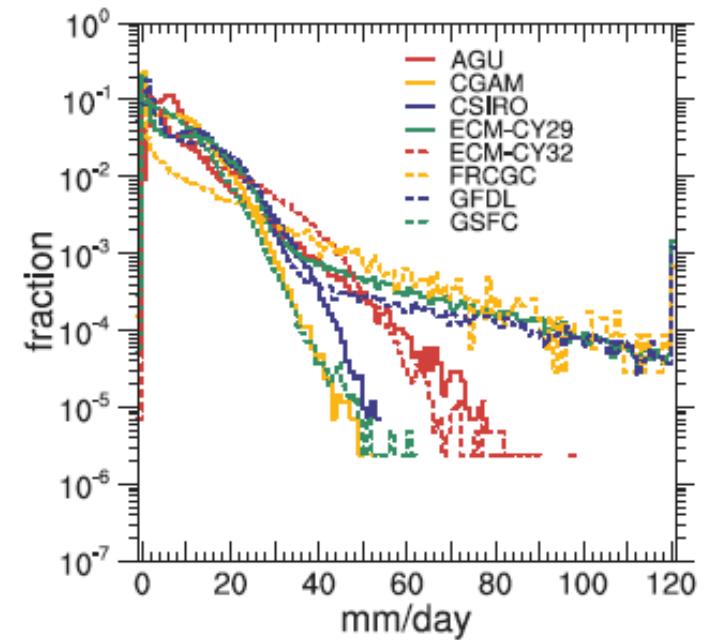


GCM Test Hierarchy: Aqua-planet

- 3D aqua-planet tests give insights into the characteristics of moisture processes, what drives these differences?



Mishra et al., J. Climate (2011)



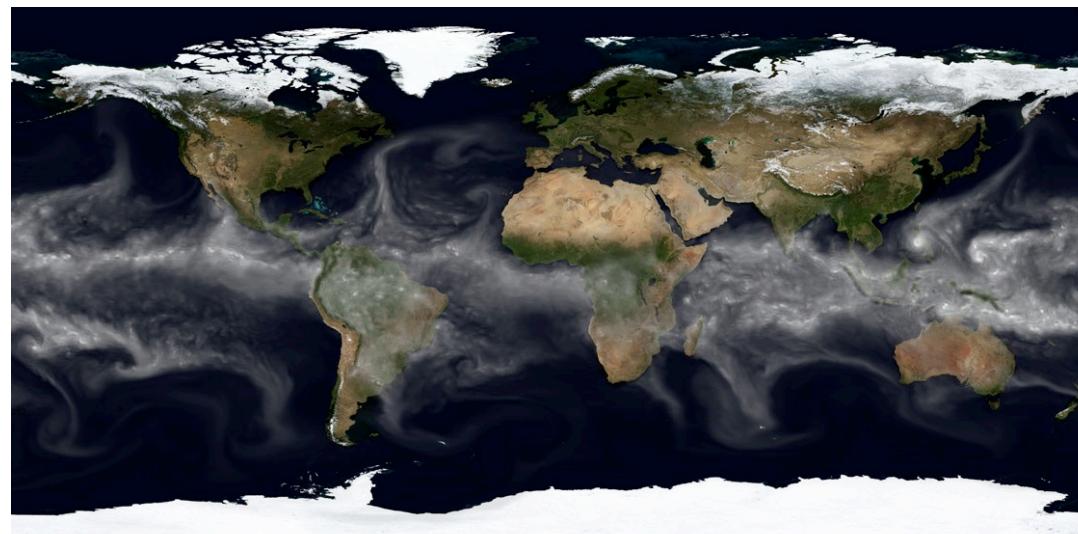
Fraction of time precipitation in the tropics is between 0-120 mm/day

Williamson et al., NCAR Technical Note TN-484+STR (2012)



GCM Test Hierarchy: AMIP

- 3D AMIP tests evaluate the interaction between the dynamical core with complex physical parameterizations (maybe even including chemistry packages) using a **complex but prescribed lower boundary** (orography, prescribed observation-based SSTs and sea ice) over **25-year** time frames



Snapshot of the precipitable water in a CAM-FV (25 km) AMIP simulation conducted by Kevin Reed (University of Michigan) and Michael Wehner (LBNL)



GCM Test Hierarchy: AMIP

- Time-loop of the precipitable water in the CAM-FV (25 km) AMIP simulation (produced by the Lawrence Berkeley Lab)

Preliminary CAM5 hi-resolution simulations (0.25°, prescribed aerosols)

Michael Wehner, Prabhat, Chris Algieri, Fuyu Li, Bill Collins
Lawrence Berkeley National Laboratory

Kevin Reed, University of Michigan

Andrew Gettelman, Julio Bacmeister, Richard Neale
National Center for Atmospheric Research

June 1, 2011





GCM Test Hierarchy: Fully coupled

- The most complex GCM evaluations utilize fully coupled atmosphere–ocean–ice–land–chemistry–carbon-cycle Earth System Models, sometimes prescribed greenhouse gas concentrations are used
- Fully coupled simulations of past time periods are typically compared against observations, sometimes in form of re-analysis data
- Differences between simulations are very hard to understand due to the complexity and nonlinear interactions
- Fully coupled GCMs are used for the assessment of future climate scenarios, e.g. for the Intergovernmental Panel on Climate Change (IPCC) assessments



Measures of ‘Truth’

- What is truth? How can we judge whether GCM model simulations are robust, reliable and accurate?
- The higher we go up in the test hierarchy the more difficult it is to understand the causes and effects, and to determine the accuracy of a simulation.
- Only very idealized test cases have analytic solutions.
- In non-linear dry dynamical core test cases we rely on **ensembles of high-resolution references** solutions to determine the perceived ‘truth’ and its uncertainty.
- Dry dynamical core tests converge within some uncertainty with increasing resolution.
- Moist dynamical cores and GCMs do not converge when resolution is increased.



Ensembles

- Ensembles are one way to assess the robustness of the simulations, and to gain insight into the uncertainty of the model simulations

There are three types of ensembles

- Perturbed parameter ensembles, e.g. variations of
 - empirical tuning factors in the physical parametrizations
 - diffusion coefficients or physical constants in the dynamical core
- Initial data and boundary value ensembles, e.g.
 - slight variations of the initial data
 - different topography data set, different SSTs
- Multi-model ensembles, e.g. different
 - GCMs or different versions of the same GCM



Ensembles: Some Examples

- Use the idealized tropical cyclone (TC) test case from DCMIP (test 52 setup)
- Utilize 4 dynamical cores from NCAR's CAM 5 model:
 - Finite-Volume (FV)
 - Spectral Element (SE)
 - Eulerian Spectral Transform (EUL)
 - Semi-Lagrangian Spectral Transform (SLD)
- Dycores use the same hydrostatic equation set, but there are many differences in the numerical methods, computational grids, staggering, choice of prognostic variables, and the dissipation mechanisms

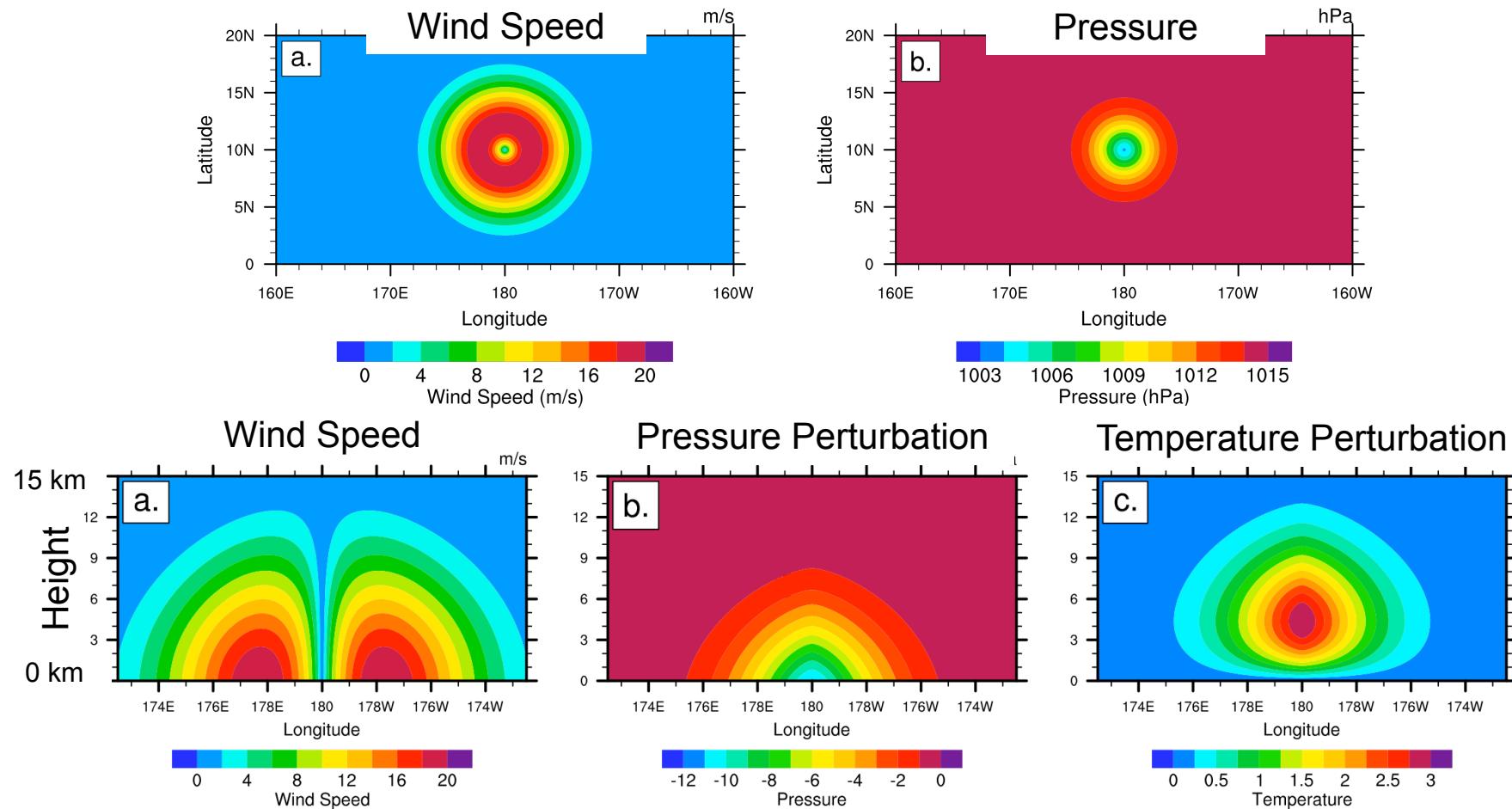


Ensembles: Experimental Design

- Use idealized initial conditions to spin up an idealized TC over 10 simulation days in an aqua-planet configuration (SST = 29 °C) of CAM
- Quick look at initial data and structural uncertainty in **CAM-FV** due to
 - slightly (2%) perturbed initial conditions, modified physical constants and longitude-position of vortex
 - CAM 4 physics versus CAM 5 physics
- Quick look at structural uncertainty due to **different dynamical cores** with the same CAM 5 physics package



Tropical Cyclone (TC) Test Case: Analytic Initial Conditions



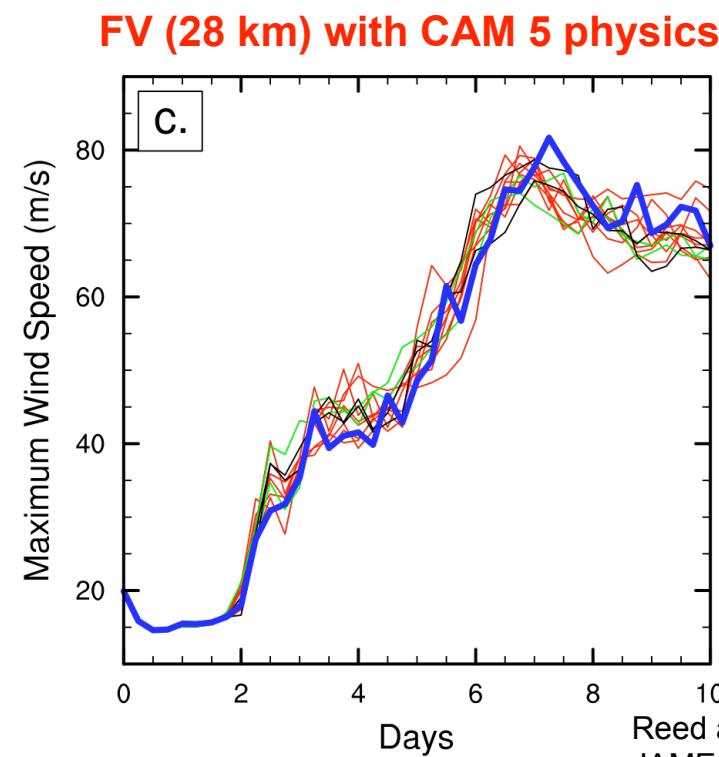
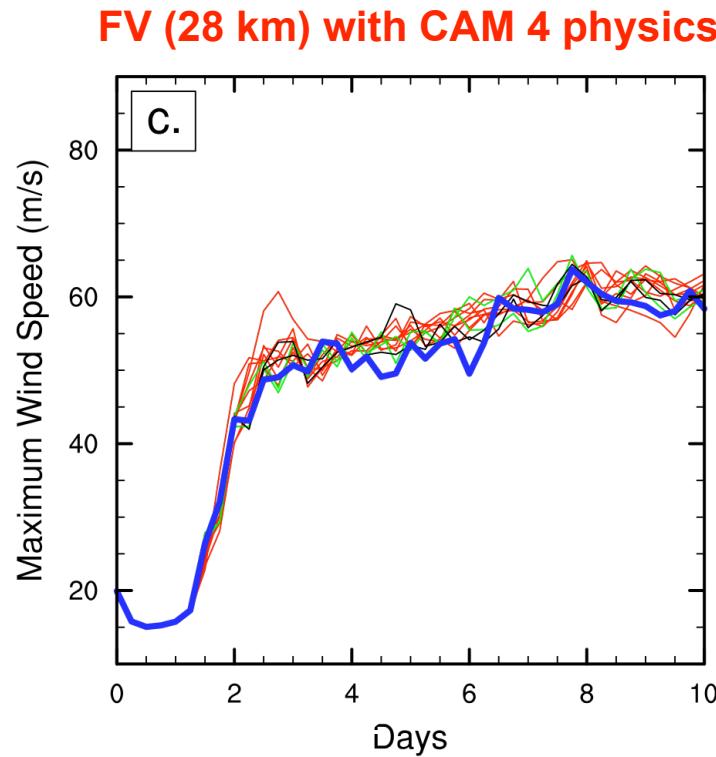
The vortex is centered at 10°N and 180°E .

Reed and Jablonowski, MWR (2011a)



TC Test: CAM-FV 4 / CAM-FV 5

- Initial data (colored lines, blue is unperturbed) and structural (physical parameterization, left/right) uncertainties: Structural uncertainty is large



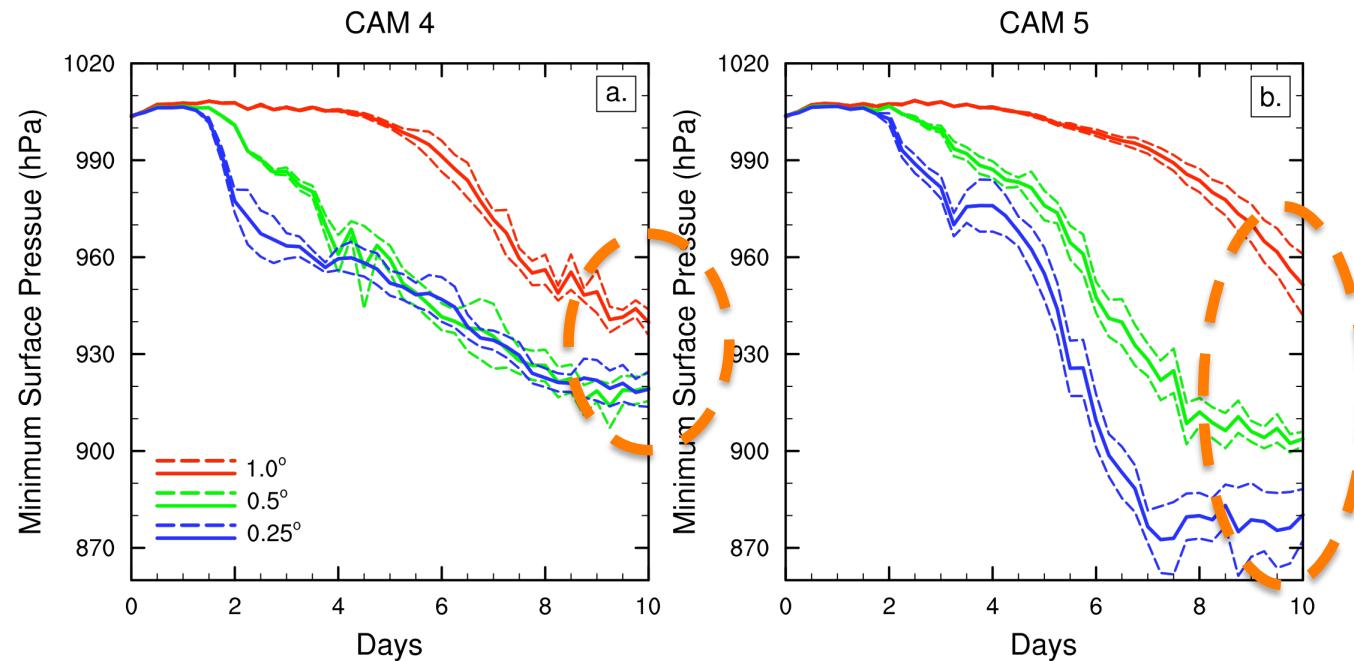
Reed and Jablonowski,
JAMES (2011c)



TC Test: CAM-FV 4 / CAM-FV 5

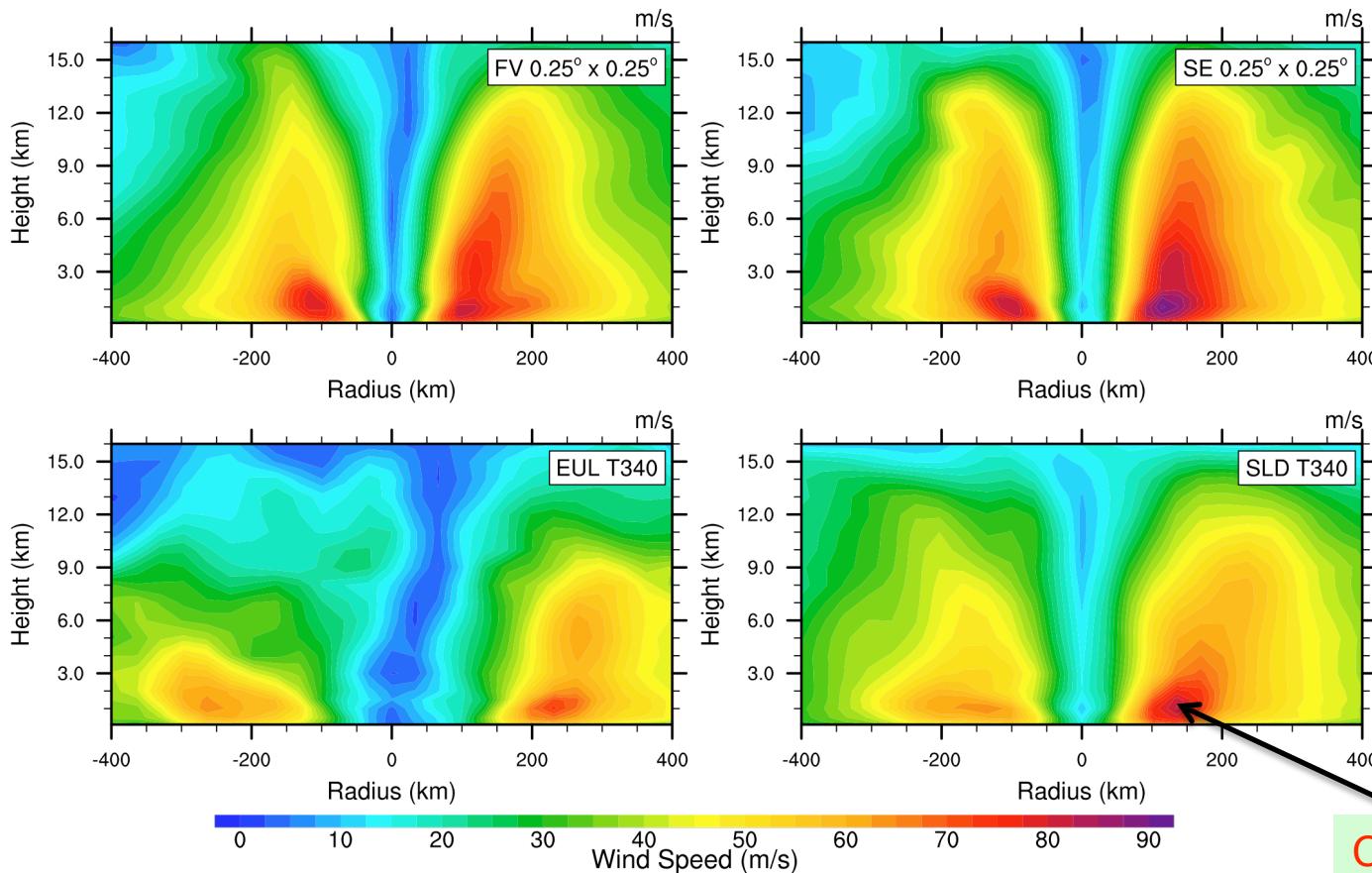
- Initial data (solid, dashed lines are uncertainty estimates) and structural (horizontal resolution, colors) uncertainties: Structural uncertainty is large

FV (28 km) with CAM 4 physics FV (28 km) with CAM 5 physics





Impact of the Dynamical Core: CAM 5 Full Physics Simulations



Wind Speed (m/s)
At Day 10

Differing strengths
and shapes of the
tropical cyclone:

FV & SE
at 0.25°
(≈ 28 km)

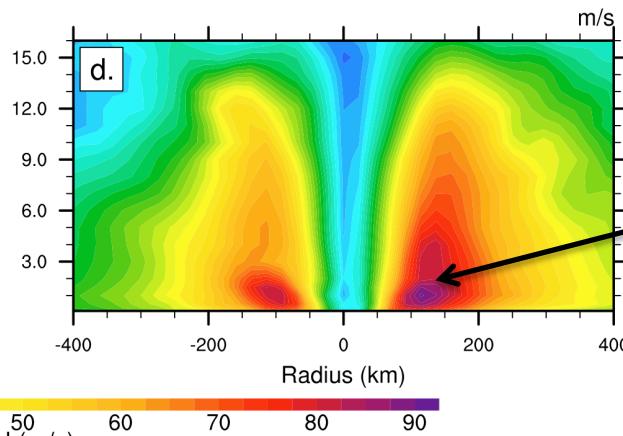
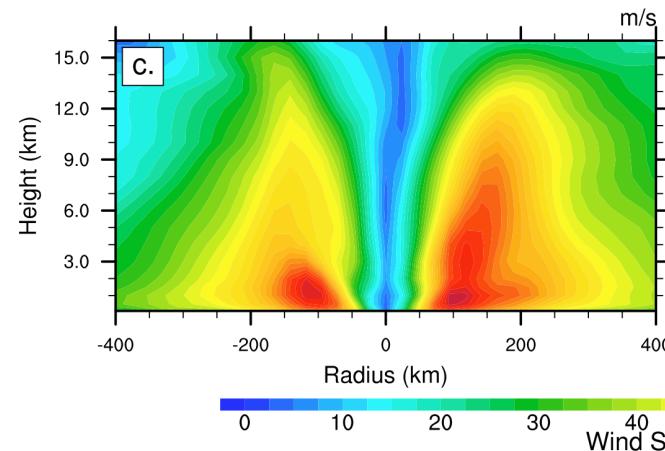
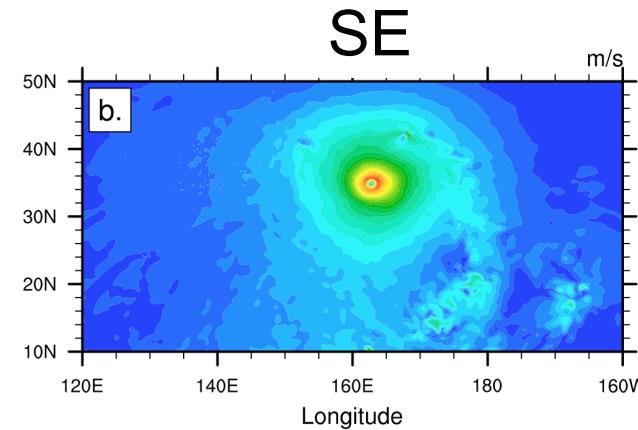
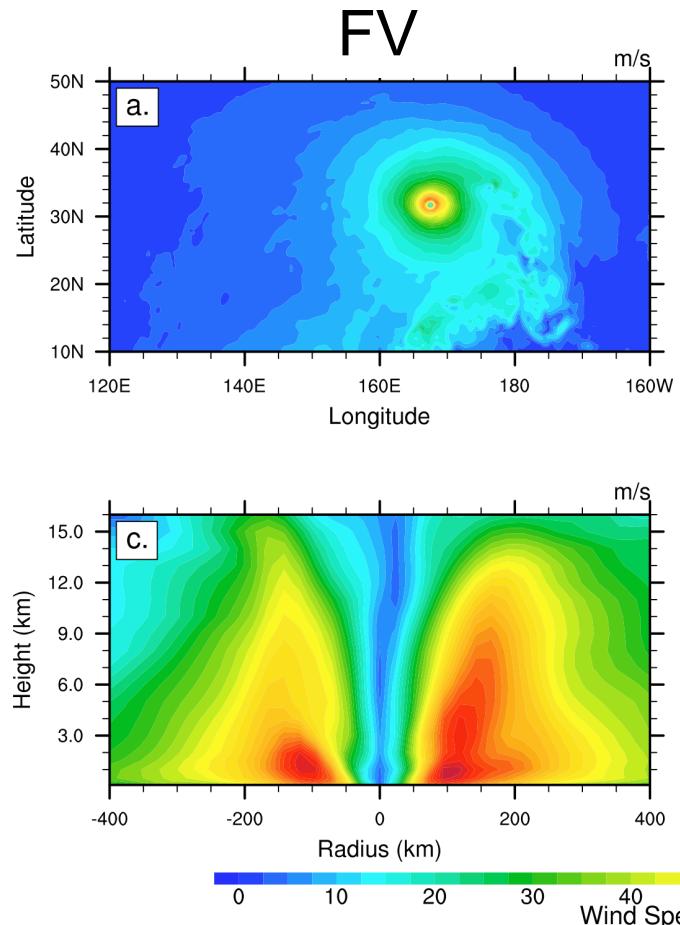
EUL & SLD
at T340
(9 km)

EUL and SLD: weaker and broader storms

Reed and Jablonowski, JAMES (2012)



Impact of the Dynamical Core: CAM 5 Full Physics Simulations



FV and SE: similar in strength and TC characteristics

Wind Speed
(m/s)
At Day 10

$0.25^\circ \approx 28 \text{ km}$

Category-5
cyclone

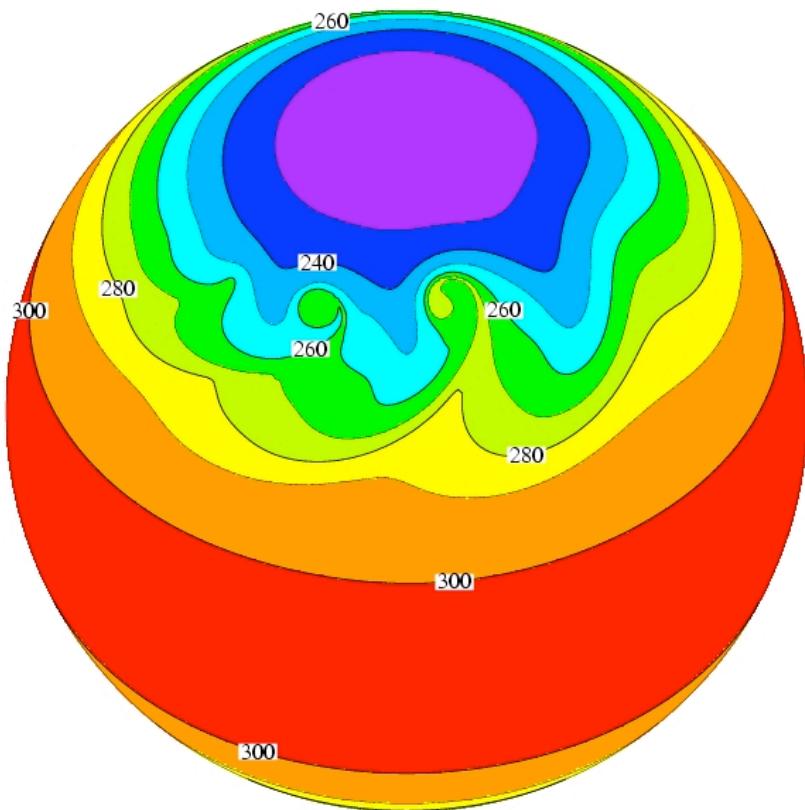


Review: Examples of Dry Dynamical Core Tests

- In 2008 we organized a 2-week NCAR Summer Colloquium on Dynamical Cores and collected & designed a suite of dry dynamical core test cases
- All had analytic initial conditions
 1. Steady-state test case (various rotations angles α)
 2. Evolution of a baroclinic wave (various rotations angles)
 3. 3D advection experiments (various rotations angles α)
 4. 3D Rossby-Haurwitz wave with wavenumber 4
 5. Mountain-induced Rossby wave train
 6. Pure gravity waves and inertial gravity waves



Review: Dry Baroclinic Wave Test (DCMIP test 410)



- 850 hPa temperature field (in K) of an idealized baroclinic wave at model day 9
- Initially smooth temperature field develops strong gradients
- Explosive cyclogenesis after day 7
- Baroclinic wave breaks after day 9
- Models start converging at 1°

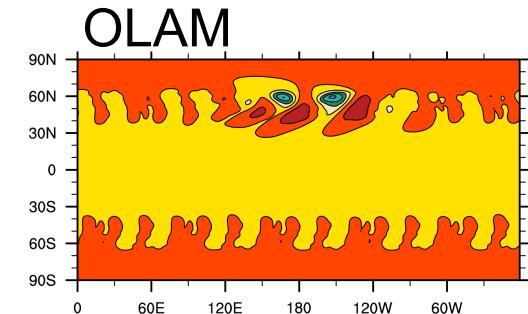
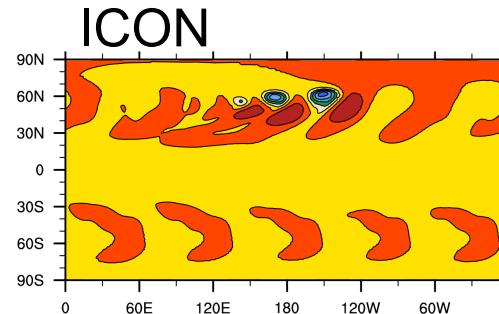
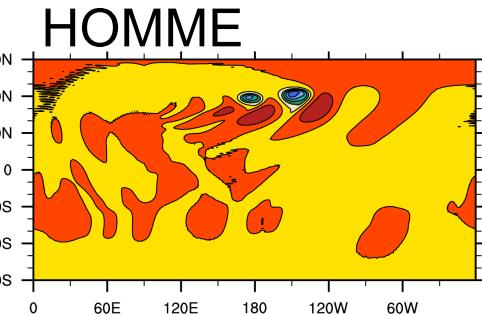
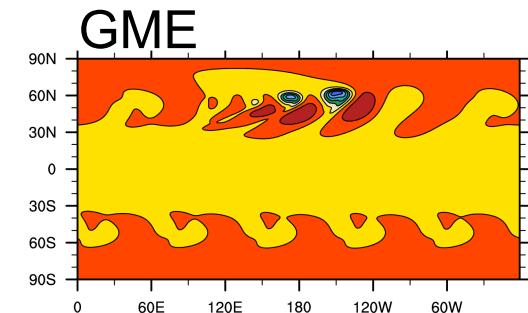
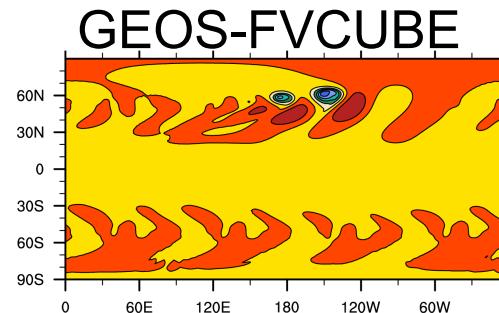
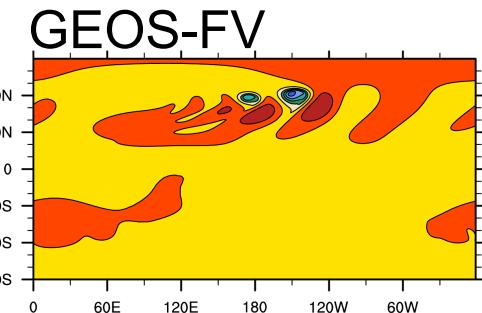
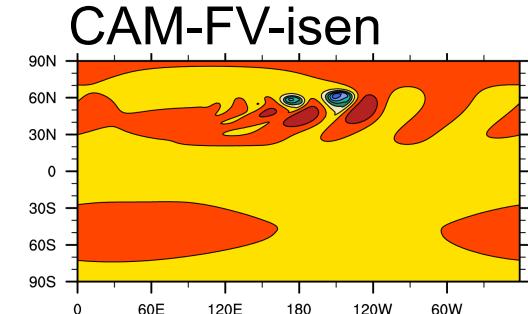
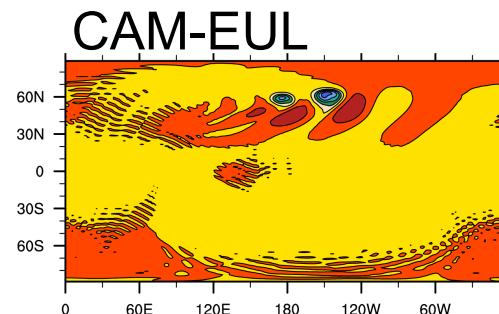
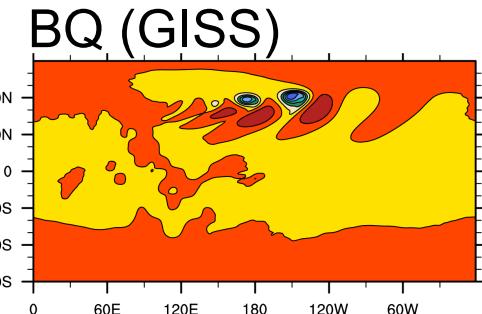
Jablonowski and Williamson, QJ (2006)



Dry Baroclinic Wave Test: p_s

Look
for grid
imprinting,
numerical
noise like
spectral
ringing,
strength of
the system

Results from 9
dynamical cores
during the 2008
NCAR
Colloquium



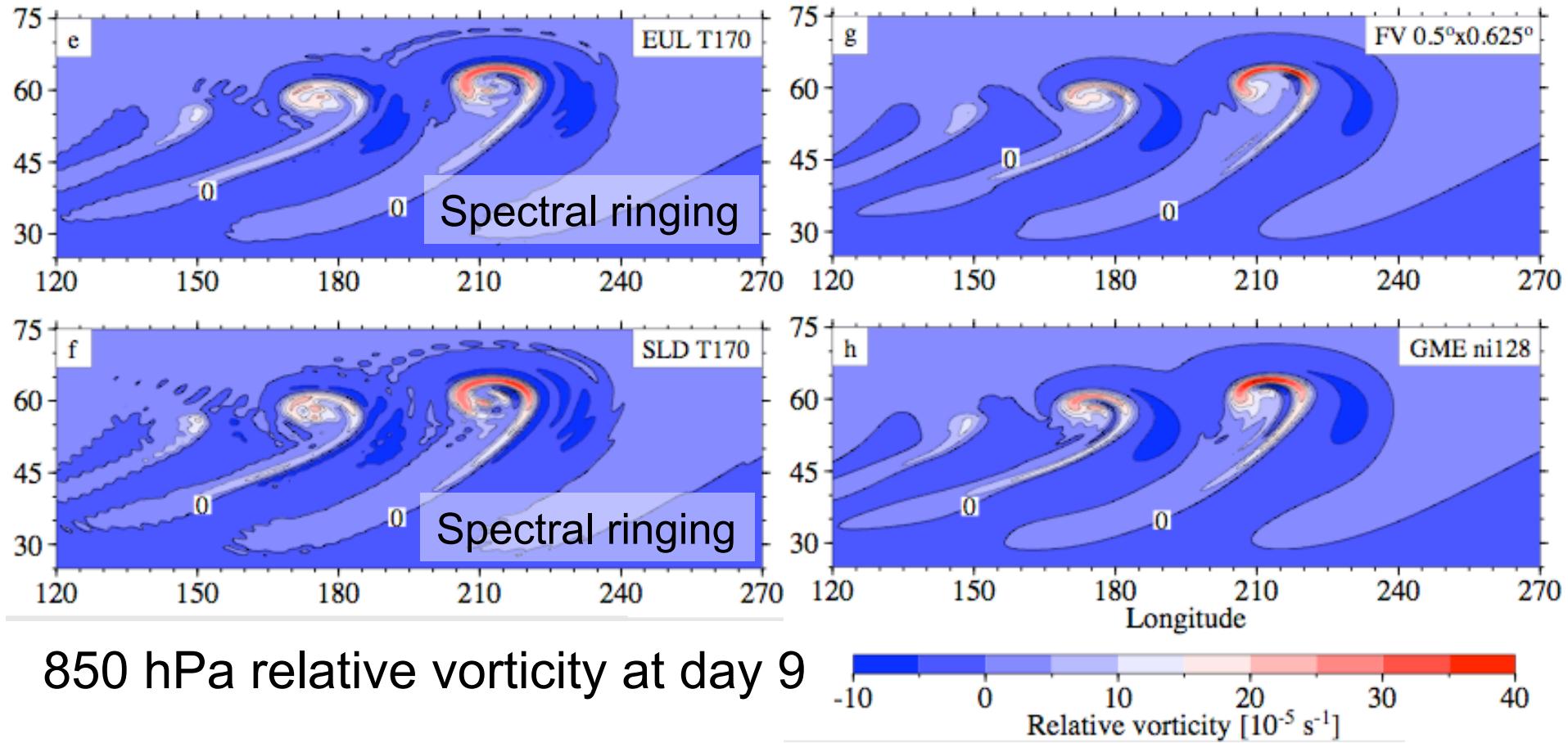
940 960 980 1000 1020 hPa
surface pressure

with $\alpha=0^\circ$, resolution $\approx 1^\circ \times 1^\circ$ L26



Dry Baroclinic Wave Test: $\zeta_{850\text{hPa}}$

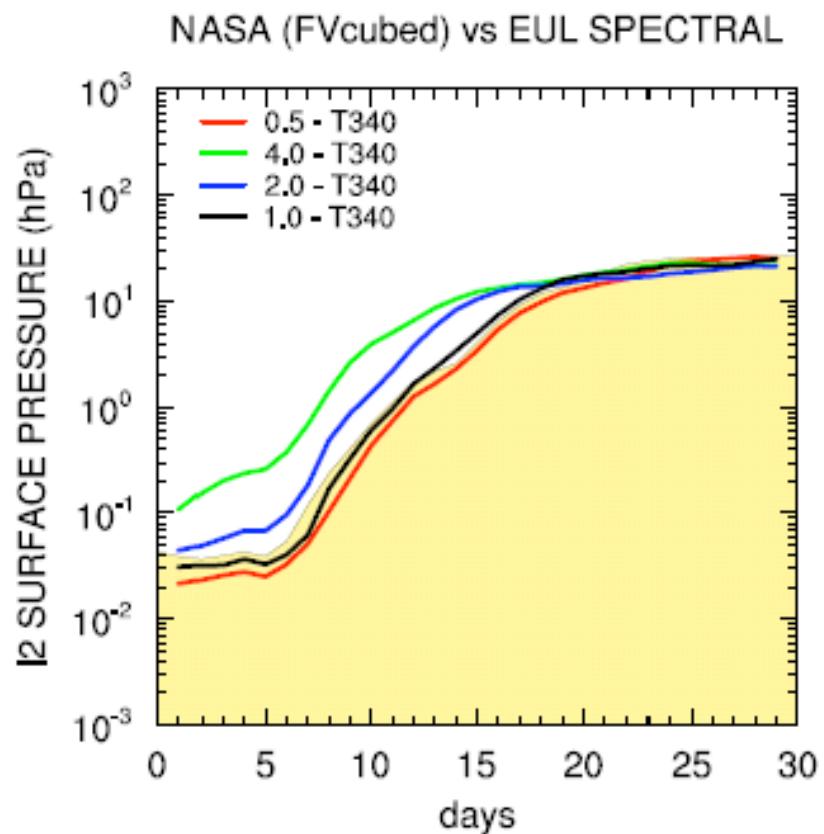
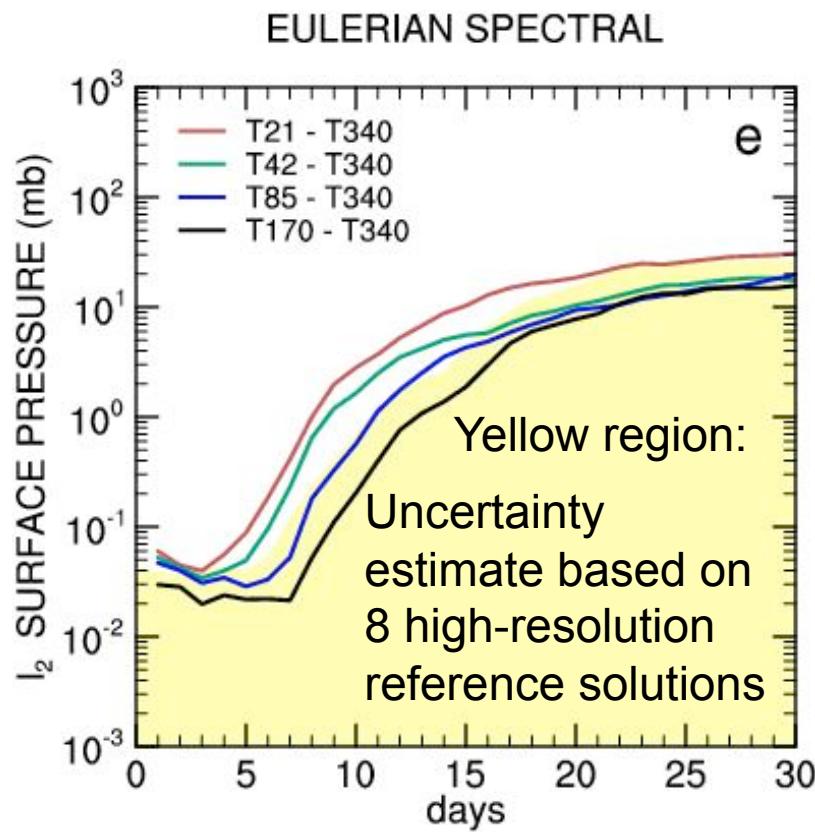
- Differences in the vorticity fields grow faster than p_s diff.
- Small-scale differences easily influenced by diffusion





Convergence & Uncertainty

- $\| \cdot \|_2$ error norms: Dry baroclinic wave simulations converge within the uncertainty estimate for the resolutions T85 (EUL & SLD), around 1° (FV), GME (55km / $ni=128$)

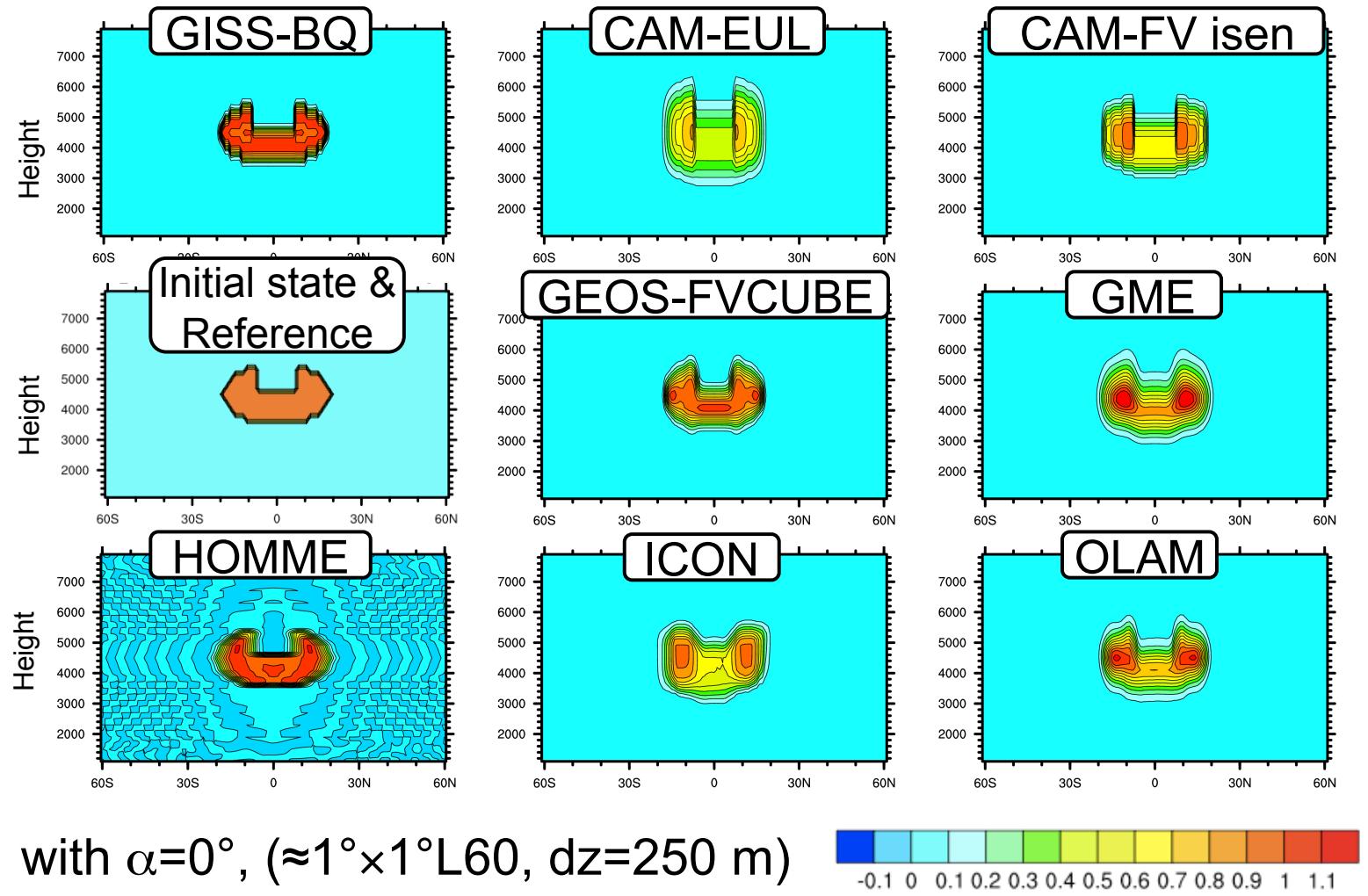




3D Tracer Transport with Prescribed Winds

Results from 8 dynamical cores during the 2008 NCAR Colloquium

Test case description in Jablonowski et al. (2008), similar to DCMIP 11

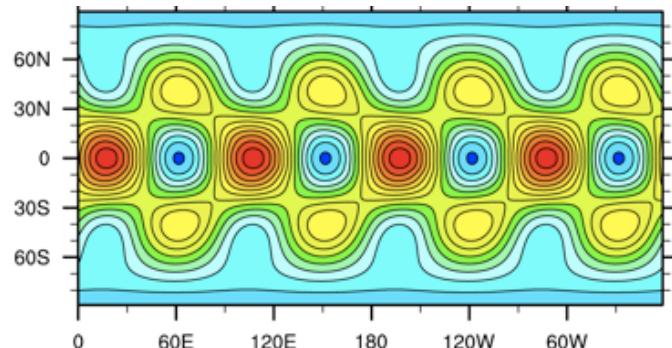




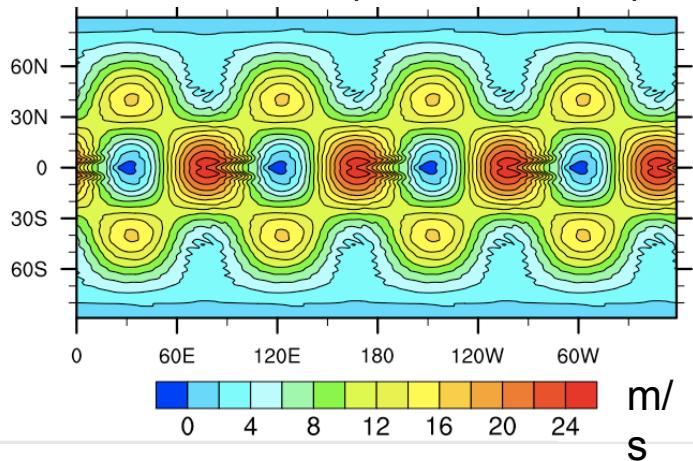
3D Rossby-Haurwitz Wave

Zonal wind at day 15 ($\approx 1^\circ \times 1^\circ$ L26)

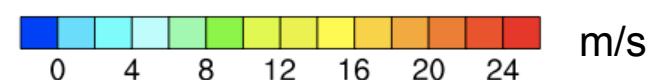
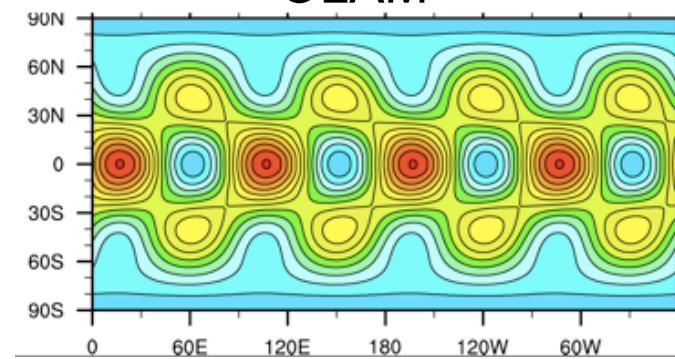
CAM-EUL



CAM-EUL (no diffusion)



OLAM



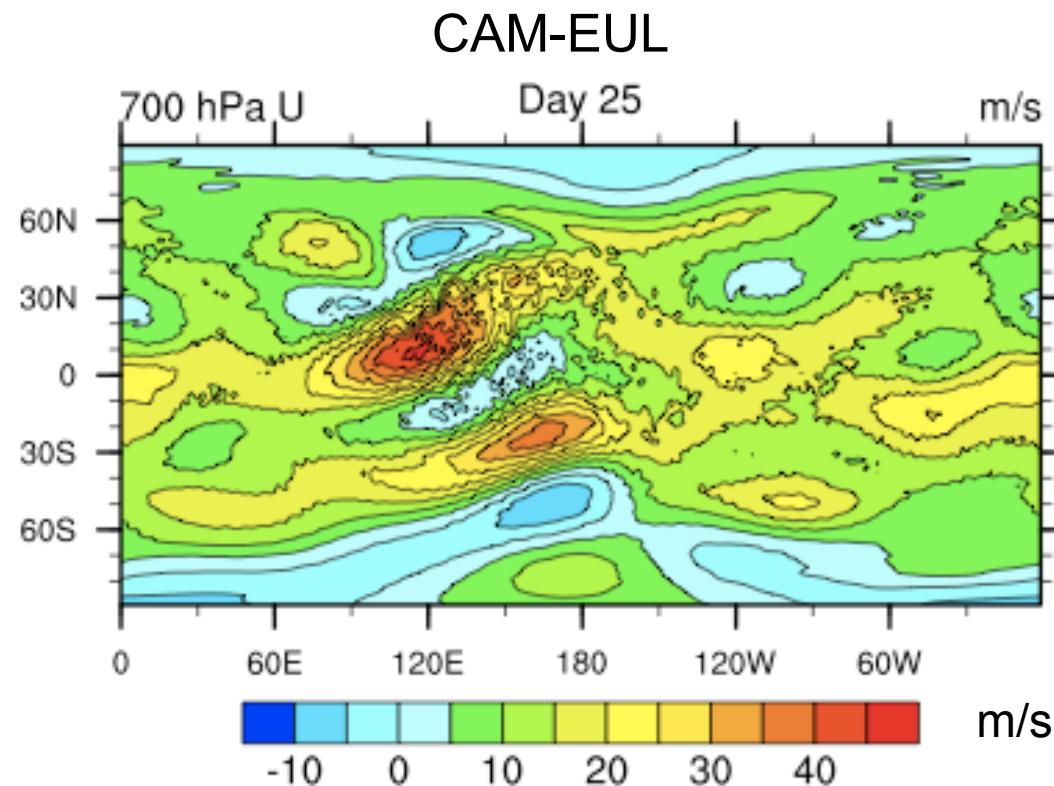
- Diffusion needed for stability in EUL
- OLAM shows reduced amplitudes

Test case description in Jablonowski et al. (2008)



Mountain-Induced Rossby Wave

700 hPa zonal wind at day 25 ($\approx 1^\circ \times 1^\circ$ L26)

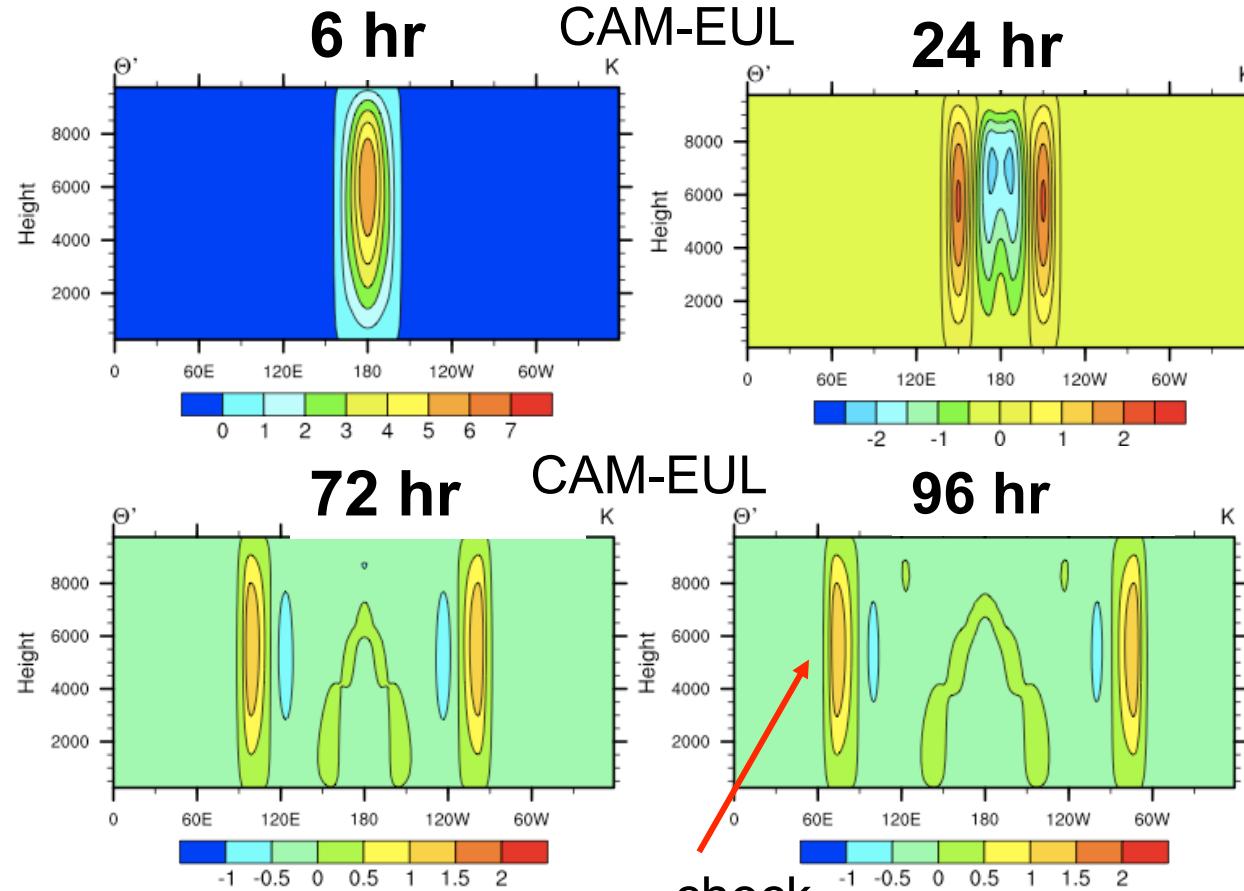


As modeler evaluators, we need to judge what we see: Numerical noise or physical nonlinear effects?



Irrational Gravity Waves

CAM-EUL T106 L20 with standard diffusion, Θ' cross section along equator



Test case description in Jablonowski et al. (2008)
similar to DCMIP 31

check
sharpness, propagation speed

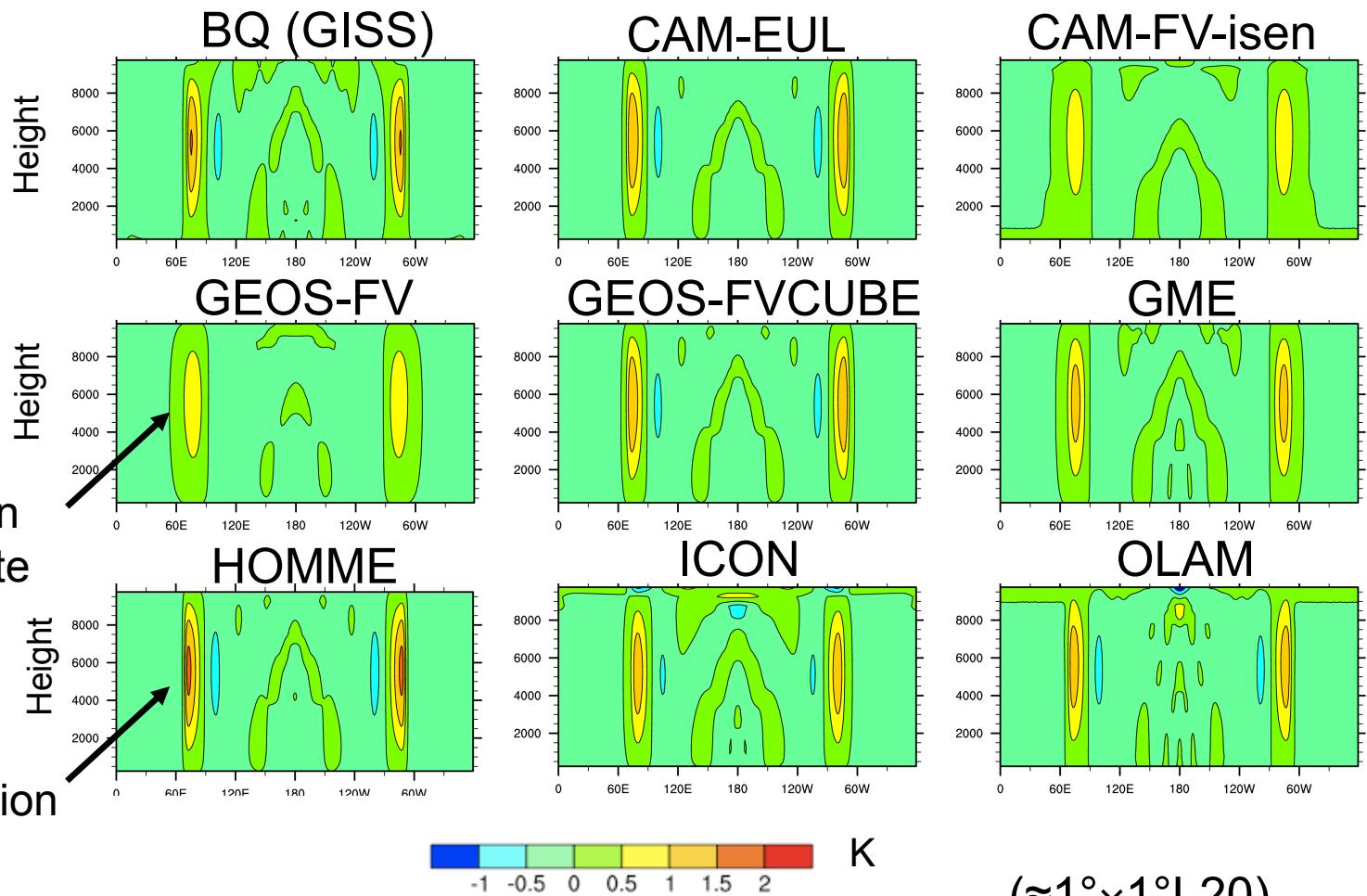


Irrational Gravity Waves

Θ' cross section along the equator at day 4

This extreme diffusion was due to explicit diffusion with an inadequate coefficient.

sharp,
low diffusion



Results from 9 dynamical cores during the 2008 NCAR Summer Colloquium.

($\approx 1^\circ \times 1^\circ L20$)



The DCMIP Test Case Hierarchy

- Advection
 - Pure 3D advection without orography
 - Pure 3D advection in the presence of orography
- Dry dynamical core without rotation
 - Stability of a steady-state at rest in presence of a mountain
 - Mountain-induced gravity waves on small planets
 - Thermally induced gravity waves on small planets
- Dry dynamical core with the Earth's rotation
 - From large (hydrostatic) to small (nonhydrostatic) scales, nonlinear baroclinic waves on a shrinking planet
- Simple moisture feedbacks
 - Moist baroclinic waves with large-scale condensation
 - Moist baroclinic waves with simplified physics (simple-physics)
 - Idealized tropical cyclones with simple- and full-physics

For descriptions see http://earthsystemcog.org/projects/dcmip-2012/test_cases



The DCMIP Test Case Hierarchy

Table 1: Overview of all DCMIP test cases

- | | |
|-----|---|
| 11 | 3D deformational flow |
| 12 | 3D Hadley-like meridional circulation |
| 13 | 2D transport of thin cloud-like tracers in the presence of orography |
| 200 | optional: Steady-state at rest in the presence of moderately-steep orography |
| 201 | optional: Steady-state at rest in the presence of steep orography on a small planet ($X=500$) |
| 21 | Mountain waves over a Schaeer-type mountain on a small planet without shear ($X=500$) |
| 22 | Mountain waves over a Schaeer-type mountain on a small planet with shear ($X=500$) |
| 31 | Gravity wave on a small planet, along the equator ($X=125$) |
| 410 | Dry baroclinic instability with dynamic tracers EPV and Θ and $X=1$ |
| 411 | Dry baroclinic instability with dynamic tracers EPV and Θ and $X=10$ (scaled small planet) |
| 412 | Dry baroclinic instability with dynamic tracers EPV and Θ and $X=100$ (scaled small planet) |
| 413 | Dry baroclinic instability with dynamic tracers EPV and Θ and $X=1000$ (scaled small planet) |
| 42 | Moist baroclinic instability (with large-scale condensation) |
| 43 | optional: Moist baroclinic instability (with simplified physics forcing) |
| 51 | Idealized tropical cyclone (with simplified physics forcing) |
| 52 | optional: Idealized tropical cyclone (with full physics forcing) |

[Check the
DCMIP-2012
web page](#)

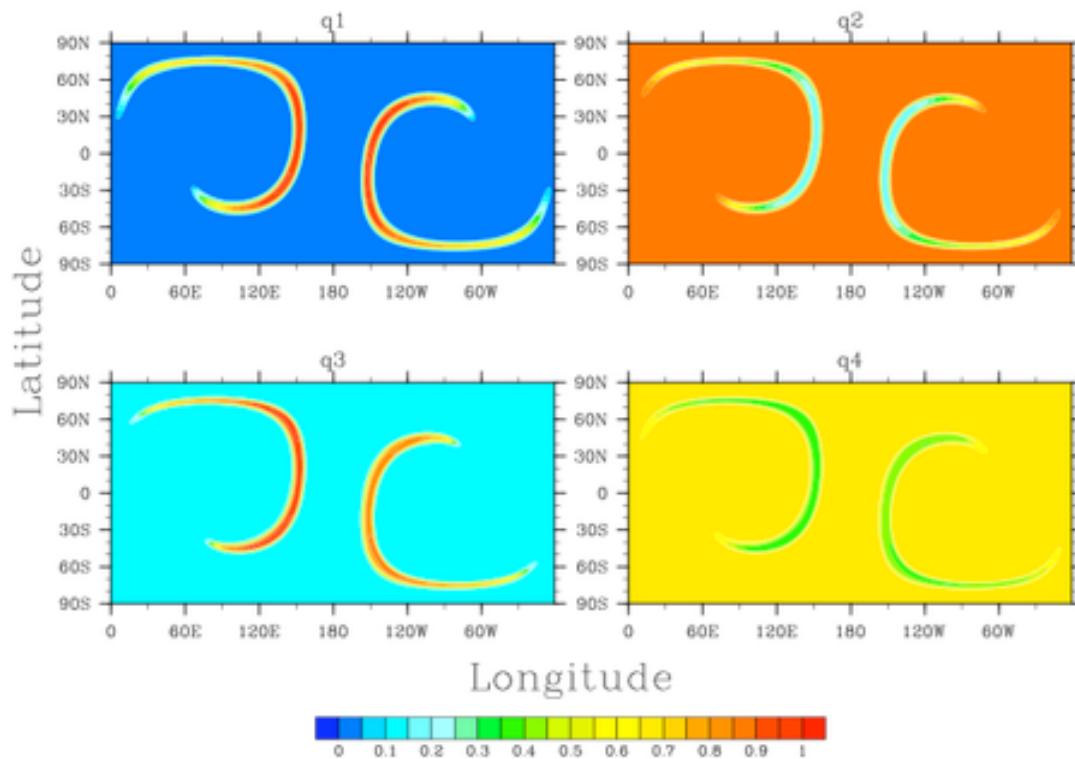


DCMIP 3D Advection Tests

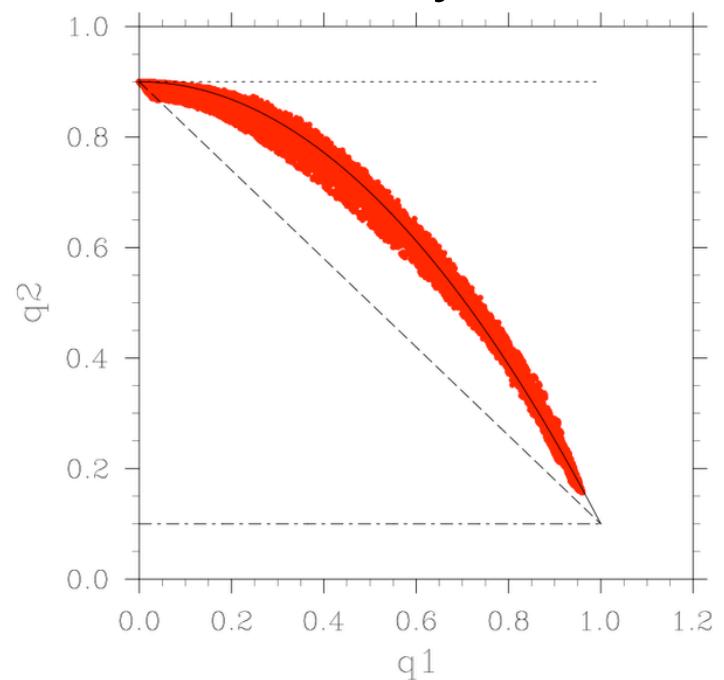
Test 11: Correlated tracers in a reversing sheared flow

4 tracers at time of flow reversal

4900 m, $t = 6$ days



Mixing diagnostics
at day 6

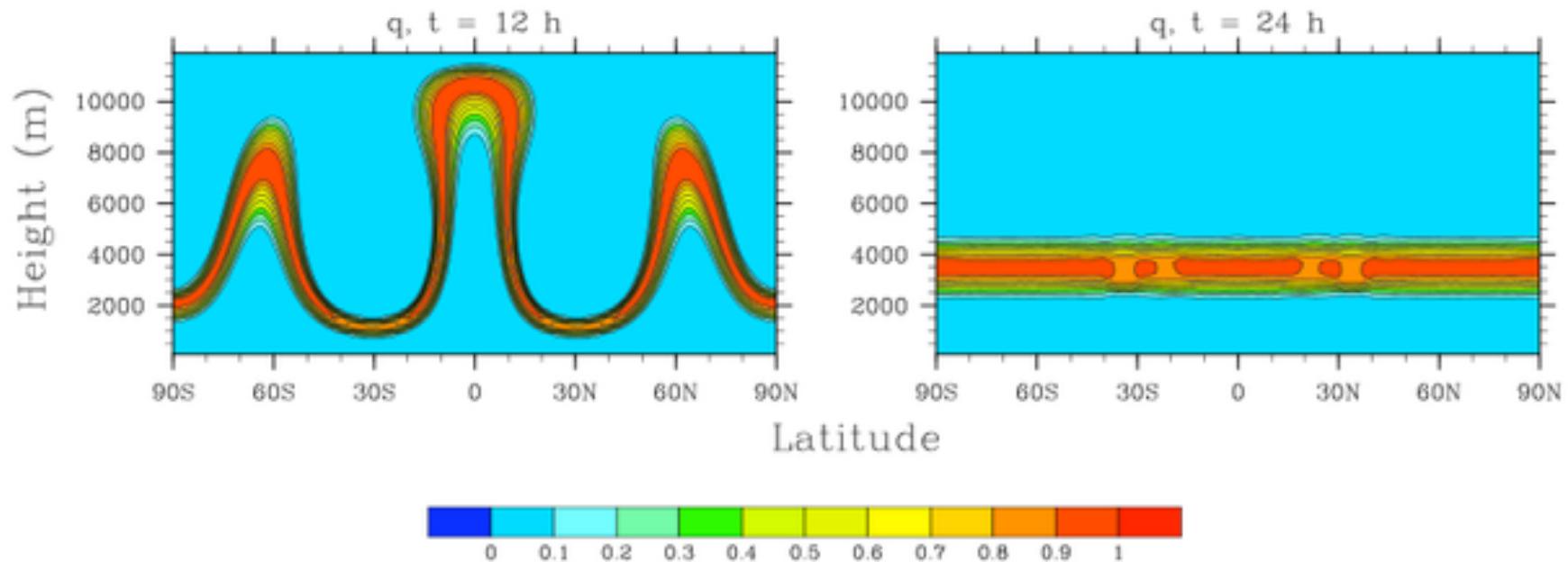




DCMIP 3D Advection Tests

Test 12: Tracers in a Hadley-cell like meridional circulation

Tracer q is stretched thin, but is still resolved on a $1^\circ \times 1^\circ$ grid, comes back to its original position after 1 day



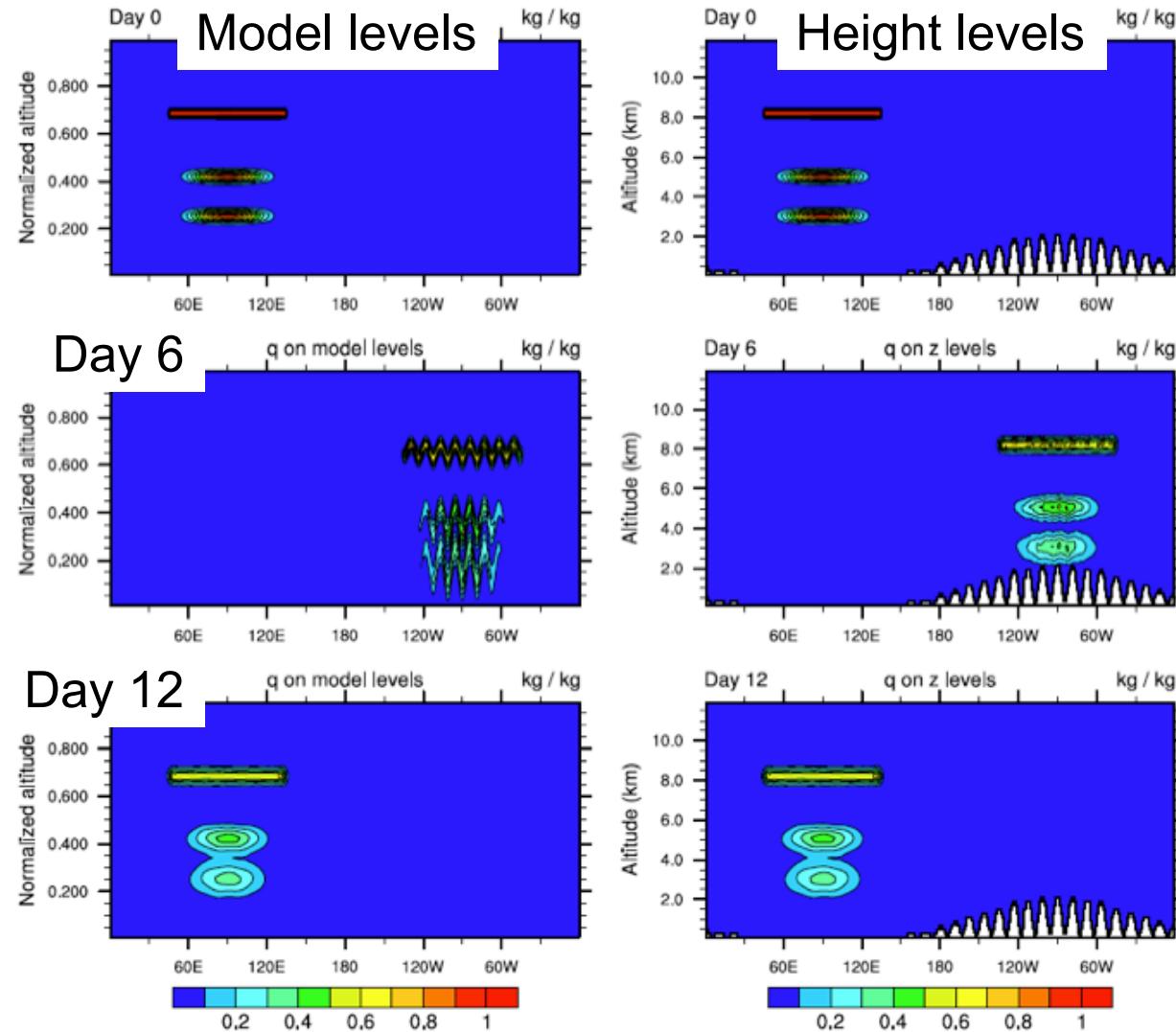
Initial state is the reference solution: check error norms



DCMIP 3D Advection Tests

Test 13:
Horizontal
advection of thin
cloud-like tracers
(at a 30° angle to
the equator)
in the presence
of orography

Model: MCORE

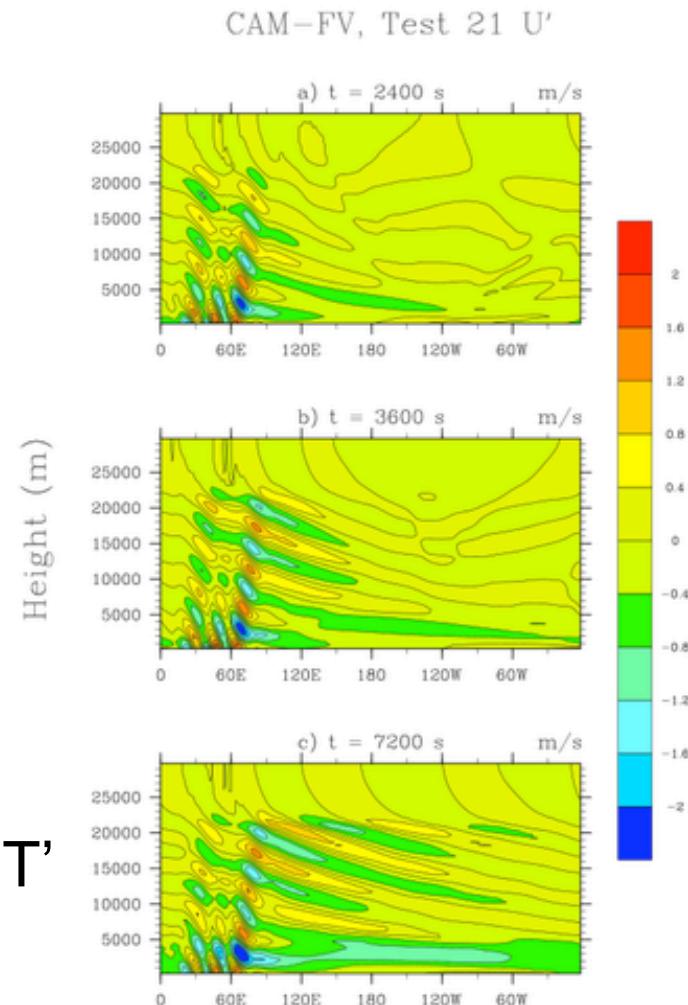




DCMIP: Flow over Topography and Gravity Waves

Test 21 on a reduced-size Earth
with a circumference at the equator of 80 km distinguishes between **hydrostatic** and **nonhydrostatic** responses

Test 21 without vertical wind shear:
Hydrostatic response in the zonal wind perturbation u' in CAM-FV at the scaled times 2400 s, 3600 s, 7200 s, the temperature perturbation T' looks similar



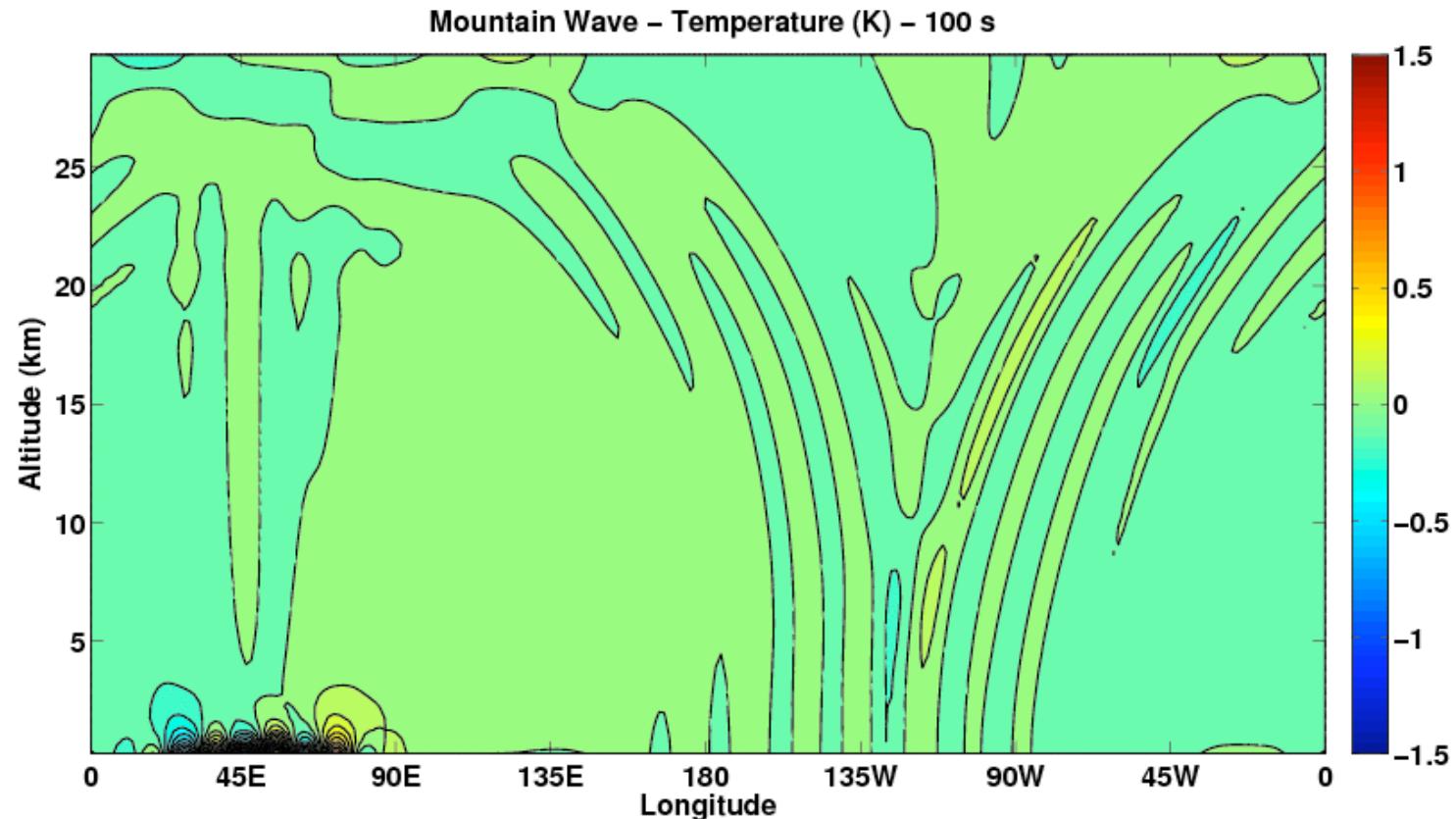
The nonhydrostatic solution is very different.



DCMIP: Flow over Topography and Gravity Waves

Test 22 on a reduced-size Earth:

with vertical wind shear: Example of nonhydrostatic response in the temperature perturbation field T' (model MCORE, animation)

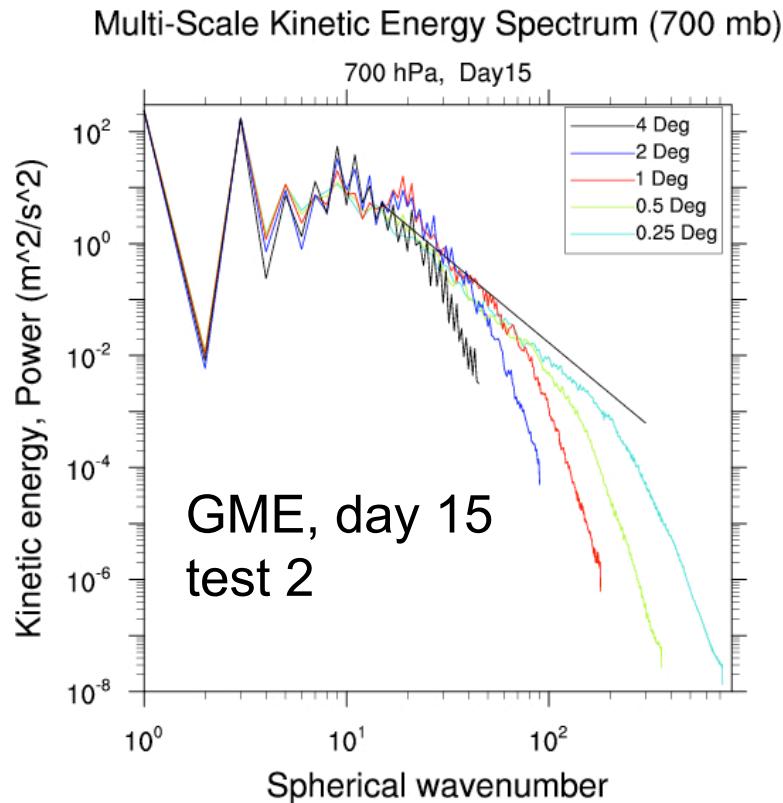




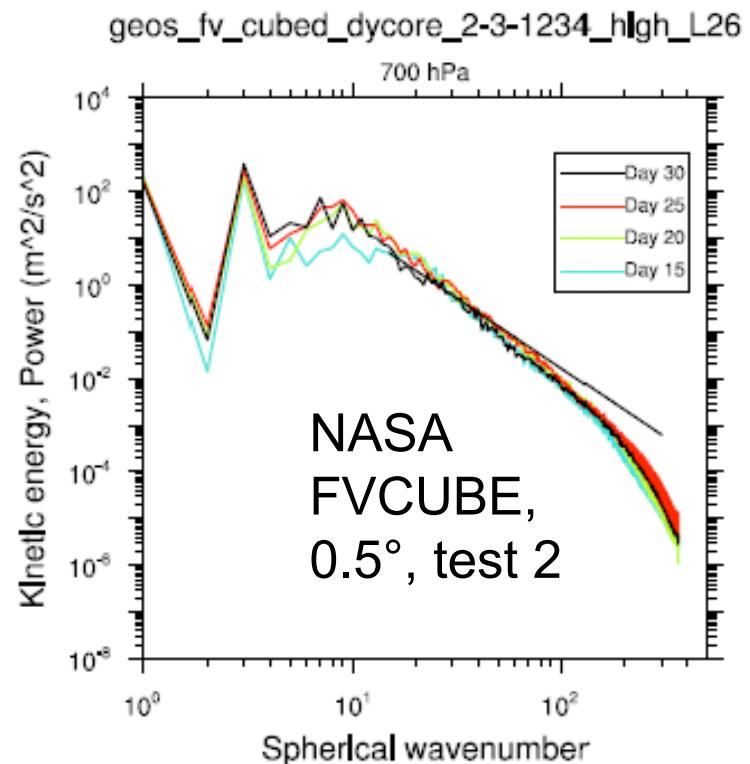
DCMIP Test 41X: KE spectra

Test 41X: Baroclinic wave KE spectra as diagnostic tools,
Diffusion characteristics are expressed by steepness of tail

Optional: Variation with resolution



Variation with time

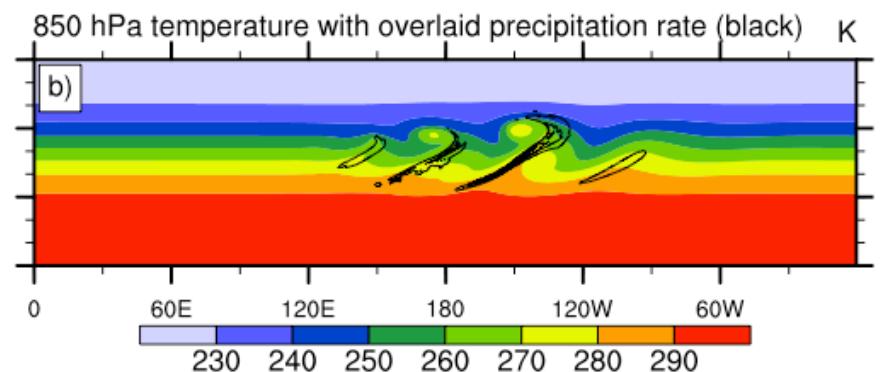
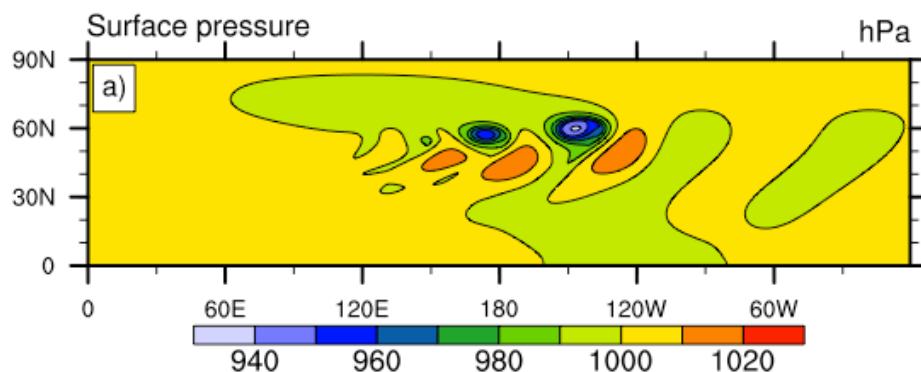




DCMIP Test 42: Moisture and Large-Scale Condensation

Test 42:

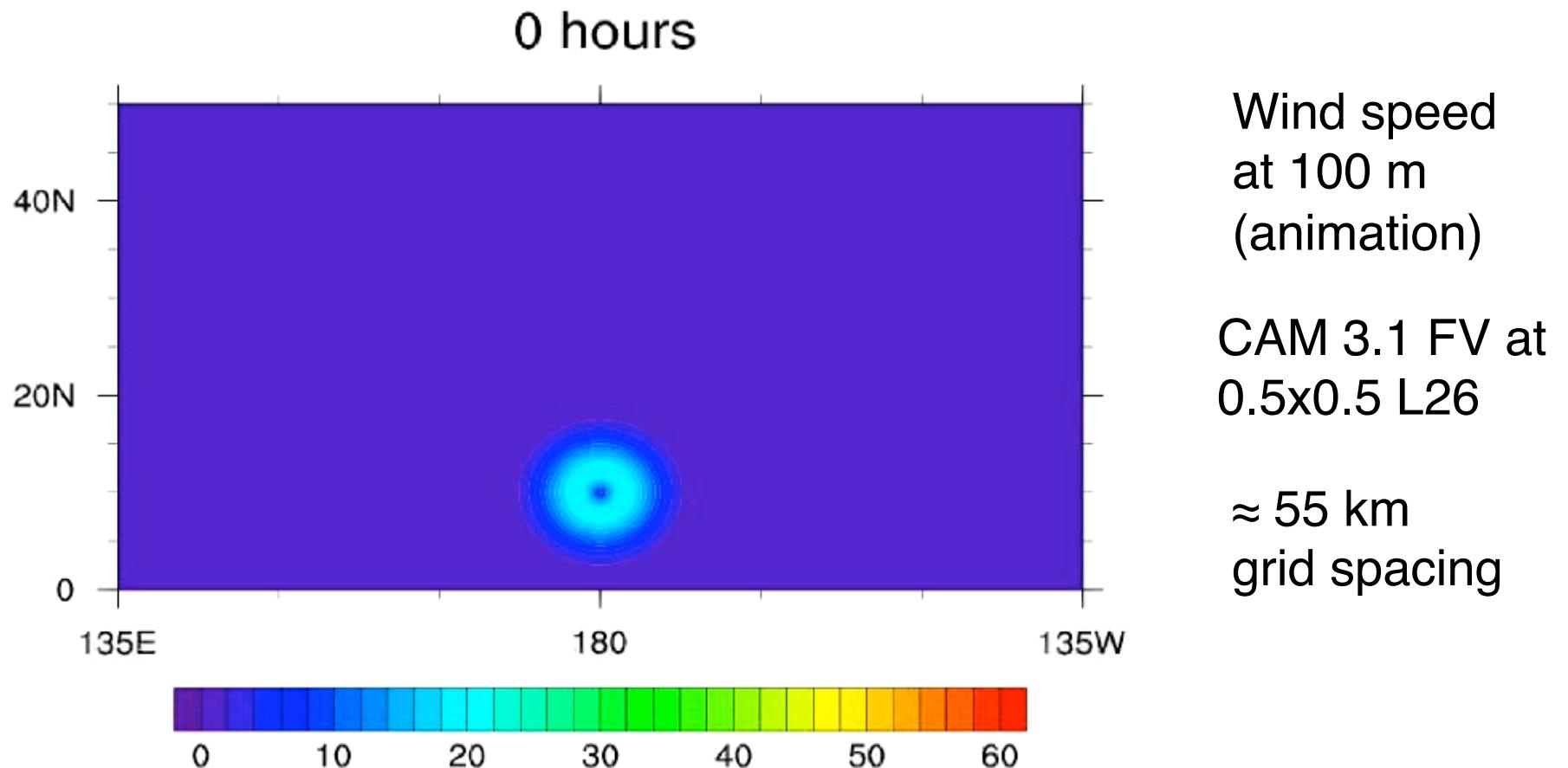
Large-scale condensation in the moist baroclinic wave leads to an intensification of the baroclinic wave in CAM-FV ($1^\circ \times 1^\circ$ L30), here at day 9





DCMIP Test 51: Tropical Cyclone with Simple-Physics

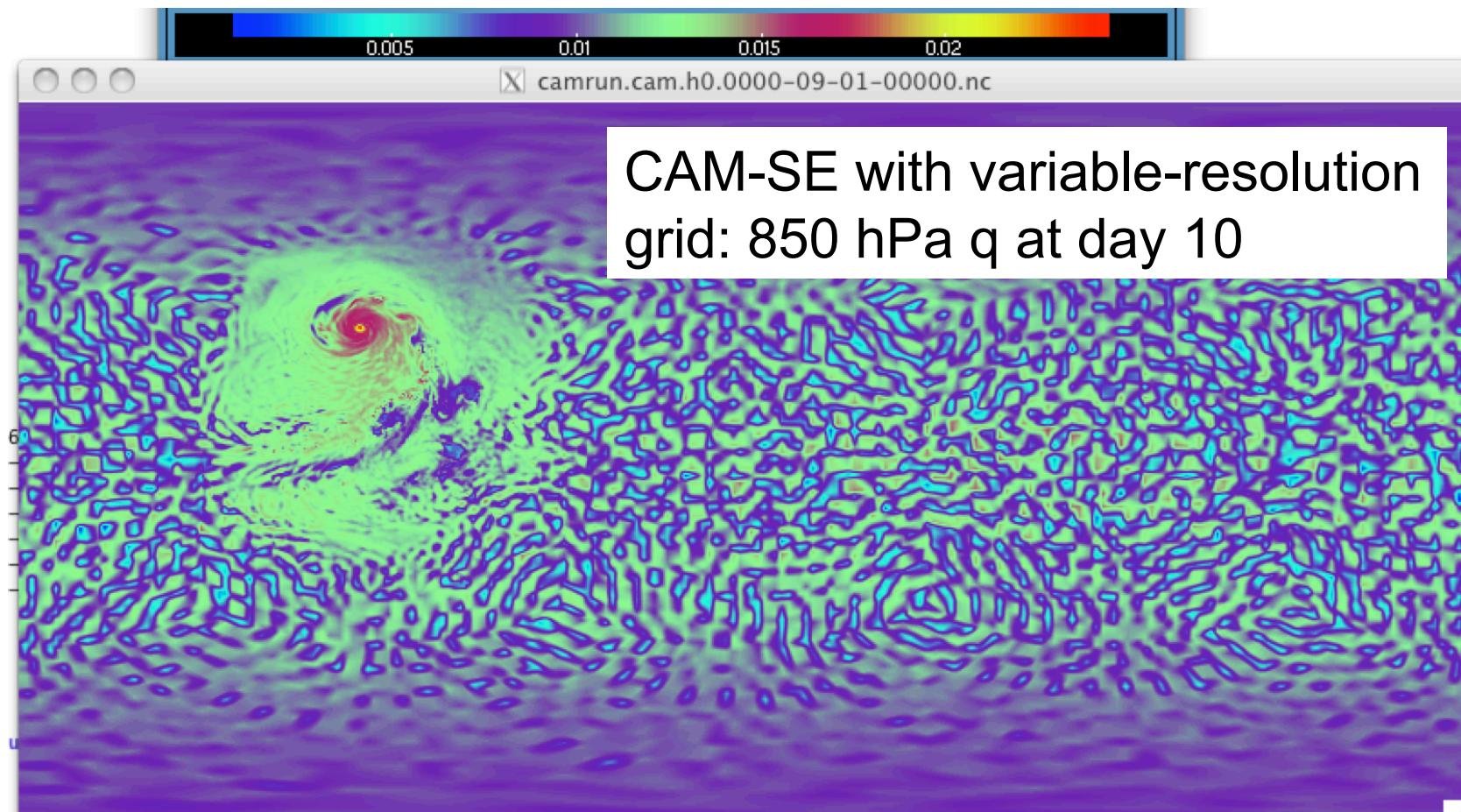
Idealized TC: Aqua-planet Simulations with Simple-Physics
(here including a Betts-Miller-type convective adjustment)





DCMIP Test 52: Tropical Cyclone with Full Physics

Test 52: Check grid imprinting signatures in the moisture field q





Some Final Thoughts

- Component testing like the test of the dynamical core is a crucial stepping stone in the GCM evaluation hierarchy.
- We (the dynamical core community) just started testing our dynamical cores in a systematic way, assessing **model ensembles and uncertainties** will become standard.
- Let's close the gap between dry dynamical cores and full-physics simulations in the evaluation hierarchy by assessing **simple moisture feedbacks**.
- We hope that DCMIP contributes valuable ideas towards a **standardized test case suite** for dry and moist dycores. Feedback is appreciated.



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