

Tracer Advection II: Advanced Numerical Methods for Transport Problems

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Atmospheric Numerical Modeling : [Desirable Properties]

Numerical algorithms for the next generation atmospheric models should be based on the following criteria:

- Inherent local and global conservation
 - High-order accuracy
 - Computational efficiency
 - Geometric flexibility (complex domain boundaries, AMR capability)
 - Non-oscillatory advection (monotonic or positivity preservation)
 - High parallel efficiency (local method, petascale capability aiming $O(100K)$ processors)
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- Examples of numerical methods which can address the above requirements:-
Continuous Galerkin or Spectral Element (SE) method, Multimoment Finite-Volume (FV) Method and Discontinuous Galerkin (DG) Method etc..
 - The DG method (DGM) is a hybrid approach which combines nice features of SE and FV methods

Part-I

- How to solve the basic building block of a complex model – the advection problem – with DGM?

Flux-Form Atmospheric Equations (Conservation Laws)

- A large class of atmospheric equations of motion for compressible and incompressible flows can be written in **flux (conservation) form**.
- Conservation laws are systems of nonlinear partial differential equations (PDEs) in flux form and can be written:

$$\frac{\partial}{\partial t} U(\mathbf{x}, t) + \sum_{j=1}^3 \frac{\partial}{\partial x_j} F_j(U, \mathbf{x}, t) = S(U),$$

where

- \mathbf{x} is the 3D space coordinate and time $t > 0$. $U(\mathbf{x}, t)$ is the state vector represents mass, momentum and energy etc.
- $F_j(U)$ are given flux vectors and include diffusive and convective effects
- $S(U)$ is the source term
- Linear transport problem is a simple example of conservation law:

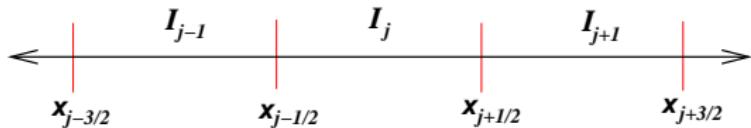
$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{V}) = 0, \quad \text{or } \rho_t + \operatorname{div}(\rho \mathbf{V}) = 0$$

Discontinuous Galerkin Method (DGM) in 1D

- 1D scalar conservation law:

$$\begin{aligned}\frac{\partial U}{\partial t} + \frac{\partial F(U)}{\partial x} &= 0 \quad \text{in } \Omega \times (0, T), \\ U_0(x) &= U(x, t = 0), \quad \forall x \in \Omega\end{aligned}$$

- E.g., $F(U) = c U$ (Linear advection), $F(U) = U^2/2$ (Burgers' Equation)
- The domain Ω (periodic) is partitioned into N_x non-overlapping elements (intervals) $I_j = [x_{j-1/2}, x_{j+1/2}], j = 1, \dots, N_x$, and $\Delta x_j = (x_{j+1/2} - x_{j-1/2})$



DGM-1D: Weak Formulation

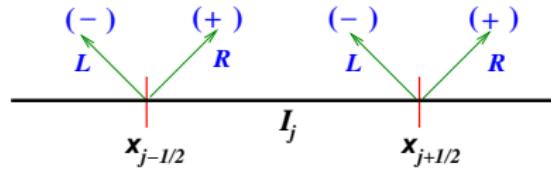
A **weak formulation** of the problem for the approximate solution U_h is obtained by multiplying the PDE by a *test function* $\varphi_h(x)$ and integrating over an element I_j :

$$\int_{I_j} \left[\frac{\partial U_h}{\partial t} + \frac{\partial F(U_h)}{\partial x} \right] \varphi_h(x) dx = 0, \quad U_h, \varphi_h \in \mathcal{V}_h$$

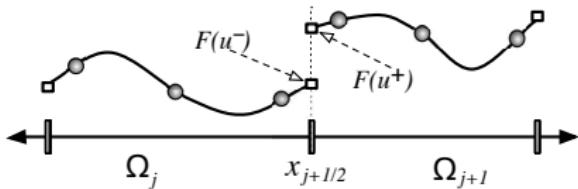
Integrating the second term by parts \implies

$$\int_{I_j} \frac{\partial U_h(x, t)}{\partial t} \varphi_h(x) dx - \int_{I_j} F(U_h(x, t)) \frac{\partial \varphi_h}{\partial x} dx + \\ F(U_h(x_{j+1/2}, t)) \varphi_h(x_{j+1/2}^-) - F(U_h(x_{j-1/2}, t)) \varphi_h(x_{j-1/2}^+) = 0,$$

where $\varphi(x^-)$ and $\varphi(x^+)$ denote "left" and "right" limits.



DGM-1D: Flux term (“Gluing” the discontinuous element edges)



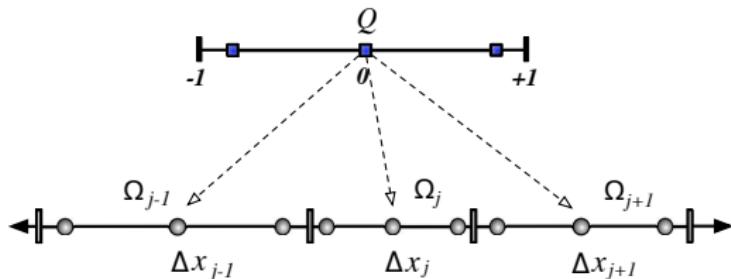
- Flux function $F(U_h)$ is **discontinuous** at the interfaces $x_{j\pm 1/2}$
- $F(U_h)$ is replaced by a **numerical flux** function $\hat{F}(U_h)$, dependent on the left and right limits of the discontinuous function U . At the interface $x_{j+1/2}$,

$$\hat{F}(U_h)_{j+1/2}(t) = \hat{F}(U_h(x_{j+1/2}^-, t), U_h(x_{j+1/2}^+, t))$$

- Typical flux formulae ([Approx. Riemann Solvers](#)): Gudunov, Lax-Friedrichs, Roe, HLLC, etc.
- Lax-Friedrichs numerical flux formula:-

$$\hat{F}(U_h) = \frac{1}{2} \left[(F(U_h^-) + F(U_h^+)) - \alpha(U_h^+ - U_h^-) \right].$$

DGM-1D: Space Discretization (Evaluation of the Integrals)



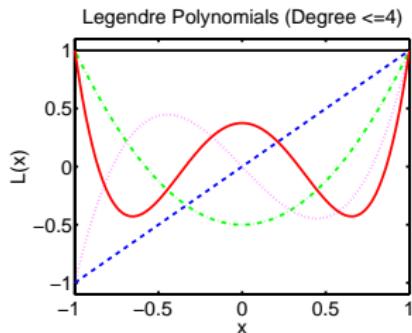
- Map every element Ω_j onto the reference element $[-1, +1]$ by introducing a local coordinate $\xi \in [-1, +1]$ s.t.,

$$\xi = \frac{2(x - x_j)}{\Delta x_j}, \quad x_j = (x_{j-1/2} + x_{j+1/2})/2 \Rightarrow \frac{\partial}{\partial x} = \frac{2}{\Delta x_j} \frac{\partial}{\partial \xi}.$$

- Use a high-order Gaussian quadrature such as the Gauss-Legendre (GL) or Gauss-Lobatto-Legendre (GLL) quadrature rule. The GLL quadrature is 'exact' for polynomials of degree up to $2N - 1$.

$$\int_{-1}^1 f(\xi) d\xi \approx \sum_{n=0}^N w_n f(\xi_n); \quad \text{for GLL, } \xi_n \Leftarrow (1 - \xi^2) P'_\ell(\xi) = 0$$

DGM-1D: Representation of Test function & Approximate Solution



- The model basis set for the \mathcal{P}^k DG method consists of Legendre polynomials, $\mathcal{B} = \{P_\ell(\xi), \ell = 0, 1, \dots, k\}$.
- Test function $\varphi_h(x)$ and approximate solution $U_h(x)$ belong to \mathcal{B}

$$U_h(\xi, t) = \sum_{\ell=0} U_h^\ell(t) P_\ell(\xi) \quad \text{for } -1 \leq \xi \leq 1, \quad \text{where}$$

$$U_h^\ell(t) = \frac{2\ell+1}{2} \int_{-1}^1 U_h(\xi, t) P_\ell(\xi) d\xi \quad \ell = 0, 1, \dots, k.$$

$$\int_{-1}^1 P_m(x) P_n(x) dx = \frac{2}{2m+1} \delta_{m,n} \Leftarrow \text{Orthogonality}$$

- $U_h^\ell(t)$ is the degrees of freedom (dof) evolves w.r.t time.

DGM-1D: Modal Basis Set for a “ \mathcal{P}^2 ” Method

- For the \mathcal{P}^2 method, $\mathcal{B} = \{P_0, P_1, P_2\} = \{1, \xi, (3\xi^2 - 1)/2\}$.
- Approximate solution:

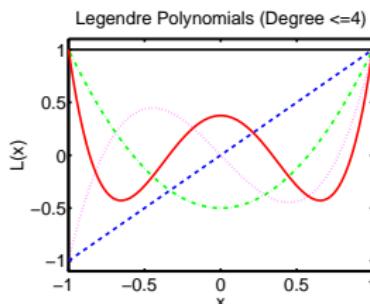
$$U_h(\xi, t) = U_h^0(t) + U_h^1(t) \xi + U_h^2(t) [3\xi^2 - 1]$$

- The degrees of freedom to evolve in t are:

$$U_h^0(t) = \frac{1}{2} \int_{-1}^1 U_h(\xi, t) d\xi \quad \Leftarrow \text{Average}$$

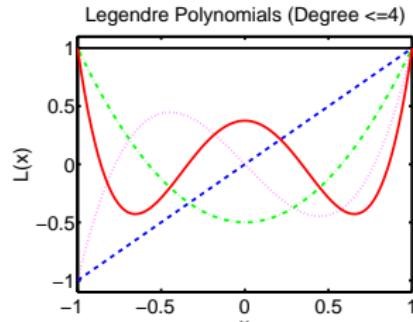
$$U_h^1(t) = \frac{3}{2} \int_{-1}^1 U_h(\xi, t) \xi d\xi$$

$$U_h^2(t) = \frac{5}{2} \int_{-1}^1 U_h(\xi, t) [3\xi^2 - 1] d\xi$$

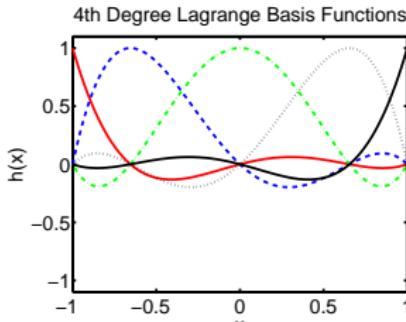


DGM-1D: Orthogonal Basis Set (Modal Vs Nodal)

Modal basis functions



Nodal basis functions



- The nodal basis set \mathcal{B} is constructed using Lagrange-Legendre polynomials $h_i(\xi)$ with roots at Gauss-Lobatto quadrature points (physical space).

$$U_j(\xi) = \sum_{j=0}^k U_j h_j(\xi) \quad \text{for } -1 \leq \xi \leq 1,$$

$$h_j(\xi) = \frac{(\xi^2 - 1) P'_k(\xi)}{k(k+1) P_k(\xi_j) (\xi - \xi_j)}, \quad \int_{-1}^1 h_i(\xi) h_j(\xi) = w_i \delta_{ij}.$$

- Nodal version was shown to be more computationally efficient than the Modal version (see, Levy, Nair & Tufo, *Comput. & Geos.* 2007)
- Modal version is more “friendly” with monotonic limiting

DG-1D: Semi-Discretized Form

- Finally, the **weak formulation** leads the PDE to the time dependent ODE

$$\int_{I_j} \left[\frac{\partial U_h}{\partial t} + \frac{\partial F(U_h)}{\partial x} \right] \varphi_h(x) dx = 0 \Rightarrow \frac{d}{dt} U_h^\ell(t) = \mathcal{L}(U_h) \quad \text{in } (0, T) \times \Omega$$

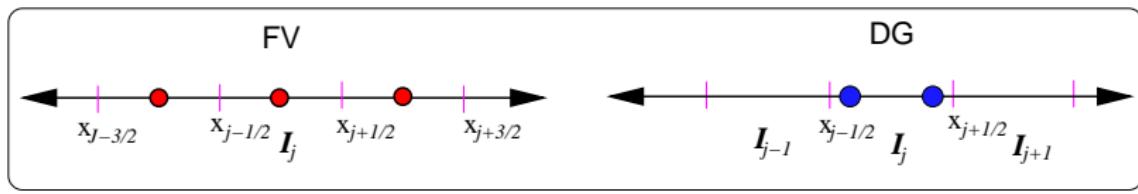
Example: For the \mathcal{P}^1 case on an element I_j , we need to solve:

$$\frac{d}{dt} U_h^0(t) = \frac{-1}{\Delta x_j} [F(\xi = 1, t) - F(\xi = -1, t)]$$

$$\frac{d}{dt} U_h^1(t) = \frac{-3}{\Delta x_j} \left([F(\xi = 1, t) + F(\xi = -1, t)] - \int_{-1}^1 U_h(\xi, t) d\xi \right)$$

Solve the ODEs for the modes at new time level $U_h^\ell(t + \Delta t)$ For the \mathcal{P}^1 case,

$$U_h(\xi, t + \Delta t) = U_h^0(t + \Delta t) + U_h^1(\xi, t + \Delta t) \xi$$



Time Integration

- For the ODE of the form,

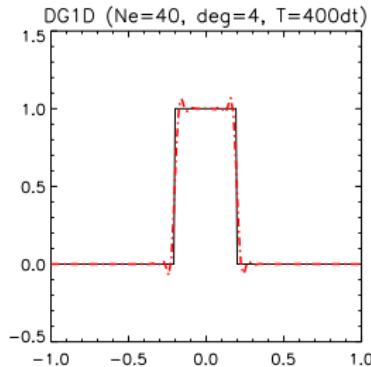
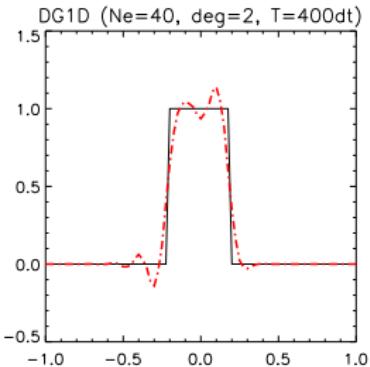
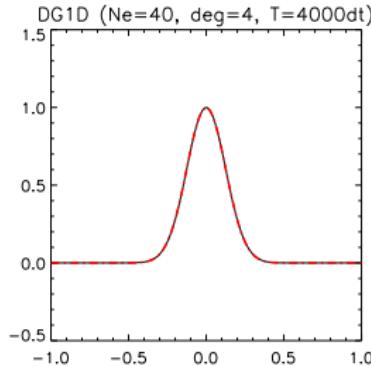
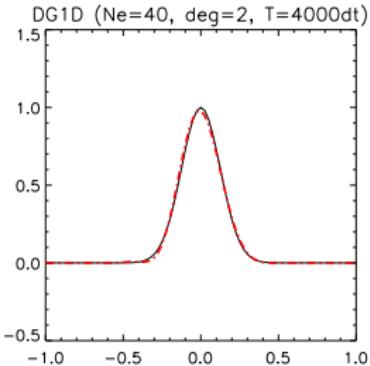
$$\frac{d}{dt} U(t) = \mathcal{L}(U) \quad \text{in } (0, T) \times \Omega$$

- Strong Stability Preserving third-order Runge-Kutta (SSP-RK) scheme (*Gottlieb et al., SIAM Review, 2001*)

$$\begin{aligned} U^{(1)} &= U^n + \Delta t \mathcal{L}(U^n) \\ U^{(2)} &= \frac{3}{4} U^n + \frac{1}{4} U^{(1)} + \frac{1}{4} \Delta t \mathcal{L}(U^{(1)}) \\ U^{n+1} &= \frac{1}{3} U^n + \frac{2}{3} U^{(2)} + \frac{2}{3} \Delta t \mathcal{L}(U^{(2)}). \end{aligned}$$

- CFL for the DG scheme is **estimated** to be $1/(2k + 1)$, where k is the degree of the polynomial (*Cockburn and Shu, 1989*).
- Remedy: Use low-order polynomials ($k \leq 3$) or efficient semi-implicit / implicit time integrators or high-order multi-stage R-K method.

DGM-1D: Results (Simple Linear Advection Test)

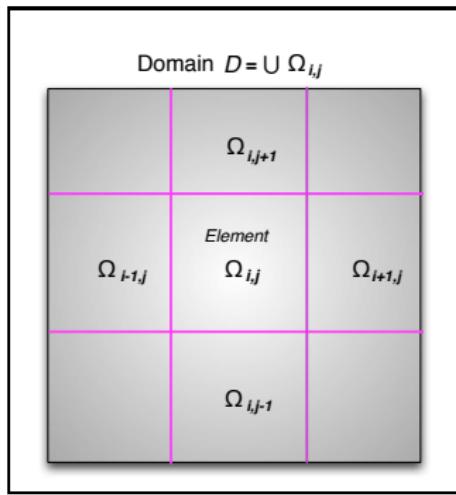


Discontinuous Galerkin (DG) Methods in 2D Cartesian Geometry

2D Scalar conservation law:

$$\frac{\partial U}{\partial t} + \nabla \cdot \mathbf{F}(U) = S(U), \quad \text{in } (0, T) \times \mathcal{D}; \quad \forall (x^1, x^2) \in \mathcal{D},$$

where $U = U(x^1, x^2, t)$, $\nabla \equiv (\partial/\partial x^1, \partial/\partial x^2)$, $\mathbf{F} = (F, G)$ is the flux function, and S is the source term.



- The domain \mathcal{D} is partitioned into non-overlapping elements Ω_{ij}
- Element edges are discontinuous
- Problem is locally solved on each element Ω_{ij}

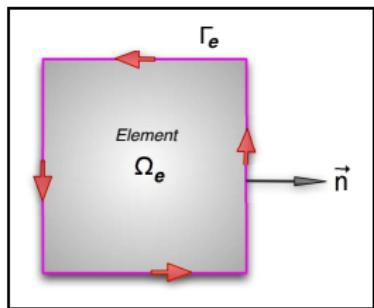
DG-2D Spatial Discretization for an Element Ω_e in \mathcal{D}

- Approximate solution U_h belongs to a vector space \mathcal{V}_h of polynomials $\mathcal{P}_N(\Omega_e)$.
- The **Galerkin formulation**: Multiplication of the basic equation by a *test function* $\varphi_h \in \mathcal{V}_h$ and integration over an element Ω_e with boundary Γ_e ,

$$\int_{\Omega_e} \left[\frac{\partial U_h}{\partial t} + \nabla \cdot \mathbf{F}(U_h) - S(U_h) \right] \varphi_h d\Omega = 0$$

- **Weak Galerkin formulation** : Integration by parts (Green's theorem) yields:

$$\frac{\partial}{\partial t} \int_{\Omega_e} U_h \varphi_h d\Omega - \int_{\Omega_e} \mathbf{F}(U_h) \cdot \nabla \varphi_h d\Omega + \int_{\Gamma_e} \mathbf{F}(U_h) \cdot \vec{n} \varphi_h d\Gamma = \int_{\Omega_e} S(U_h) \varphi_h d\Omega$$



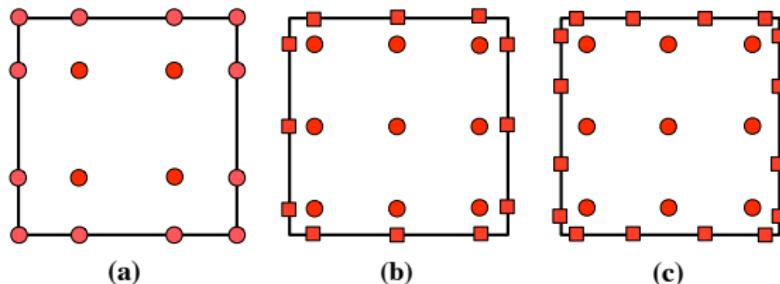
- Orthogonal polynomials (basis functions) are employed for approximating U_h and φ_h on Ω_e .
- Surface and line integrals are evaluated with high-order Gaussian quadrature rule
- Exact Integration: The flux (line) integral should be an order higher than the surface integral (Cockburn & Shu, 1989).

DG-2D: High-Order Nodal Spatial Discretization

- The nodal basis set is constructed using a tensor-product of Lagrange polynomials $h_i(\xi)$, with roots at Gauss-Lobatto-Legendre (GLL) or Gauss-Legendre (GL) quadrature points $\{\xi_i\}$.

$$[h_i(\xi)]_{GLL} = \frac{(\xi^2 - 1) P'_N(\xi)}{N(N+1) P_N(\xi_i) (\xi - \xi_i)}; \quad \int_{-1}^1 h_i(\xi) h_j(\xi) \simeq w_i \delta_{ij}.$$

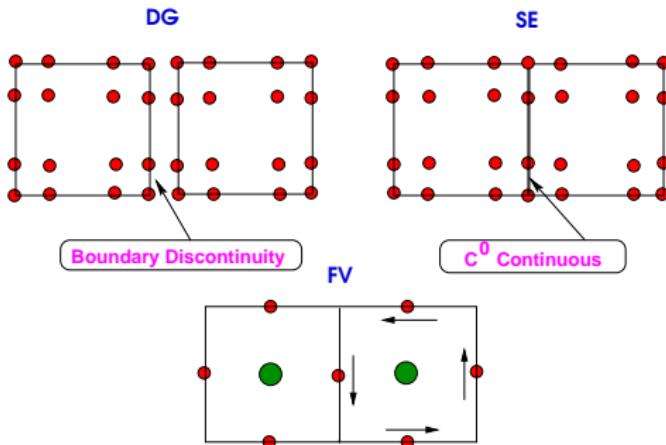
- $P_N(\xi)$ is the N^{th} degree Legendre polynomial; and w_i are Gauss quadrature weights



The approximate solution U_h and test function are represented in terms of nodal basis set.

$$U_{ij}(\xi, \eta) = \sum_{i=0}^N \sum_{j=0}^N U_{ij} h_i(\xi) h_j(\eta) \quad \text{for } -1 \leq \xi, \eta \leq 1,$$

The DG, SE & FV Methods



- For DGM degrees of freedom (*d.o.f*) to evolve per element is N^2 , where N is the order of accuracy.
- For FV method the *d.o.f* is 1 (cell-average), irrespective of order of accuracy.
- DGM is based on **conservation laws** but exploits the spectral expansion of SE method and treats the element boundaries using FV “tricks.”

Monotonic Limiter for DG transport

- Importance:
 - In atmospheric models, mixing ratios of the advecting chemical species and humidity should be non-negative and free from spurious oscillation.
 - The model should avoid creating unphysical negative mass
- Challenges:
 - Godunov theorem (1959): “Monotone scheme can be at most first-order accurate”
 - There is a “conflict of interest” between the high-order methods and monotonicity preservation!
 - In principle, a limiter should eliminate spurious oscillation and preserve high-order nature of the solution to a maximum possible extent
- Existing Limiters for DGM:
 - Minmod limiter (Cockburn & Shu, 1989): Based on van Leer's slope limiting, but too diffusive
 - Limiters based on WENO or H-WENO (Qui & Shu 2005), Expensive and no positivity preservation
 - New bound-preserving limiter: Positivity-preserving and local (Zhang & Shu, 2010)

Local Bound-Preserving Limiter for DGM

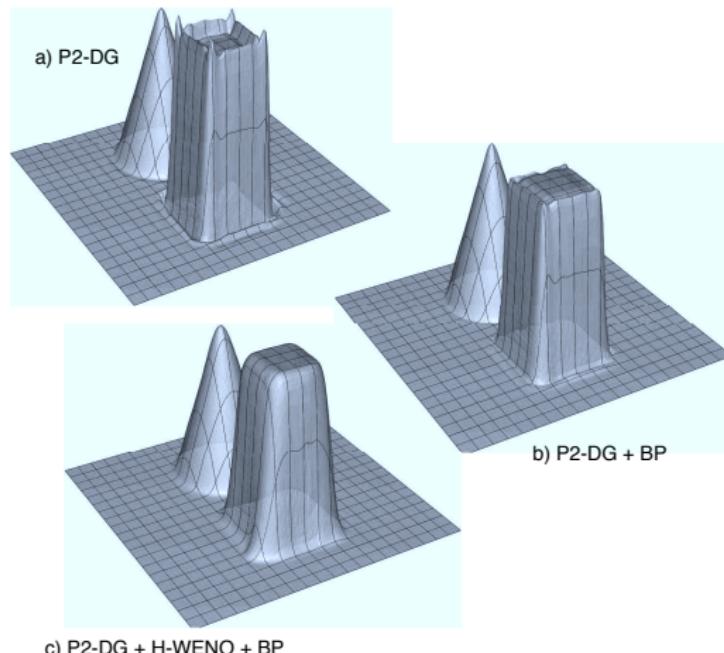
- If the global maximum M and minimum m values of the solution $\rho_{i,j}(x, y)$ is known, then the limited solution $\tilde{\rho}_{i,j}(x, y)$:

$$\tilde{\rho}_{i,j}(x, y) = \hat{\theta} \rho_{i,j}(x, y) + (1 - \hat{\theta}) \bar{u}_{i,j}, \quad \hat{\theta} = \min\left\{ \left| \frac{M - \bar{u}_{i,j}}{M_{i,j} - \bar{u}_{i,j}} \right|, \left| \frac{m^* - \bar{u}_{i,j}}{m_{i,j}^* - \bar{u}_{i,j}} \right|, 1 \right\},$$

- $\bar{u}_{i,j}$ is the average solution in the element $\Omega_{i,j}$, $M_{i,j} = \max_{(x,y) \in \Omega_{i,j}} \rho_{i,j}(x, y)$ and $m_{i,j}^* = \min_{(x,y) \in \Omega_{i,j}} \rho_{i,j}(x, y)$.
- $\hat{\theta} \in [0, 1]$. The positivity preserving option is a special case of BP filter, and can be achieved my setting $m^* = 0$.
- This limiter is conservative and local to the element (Zhang & Shu, JCP, 2010)

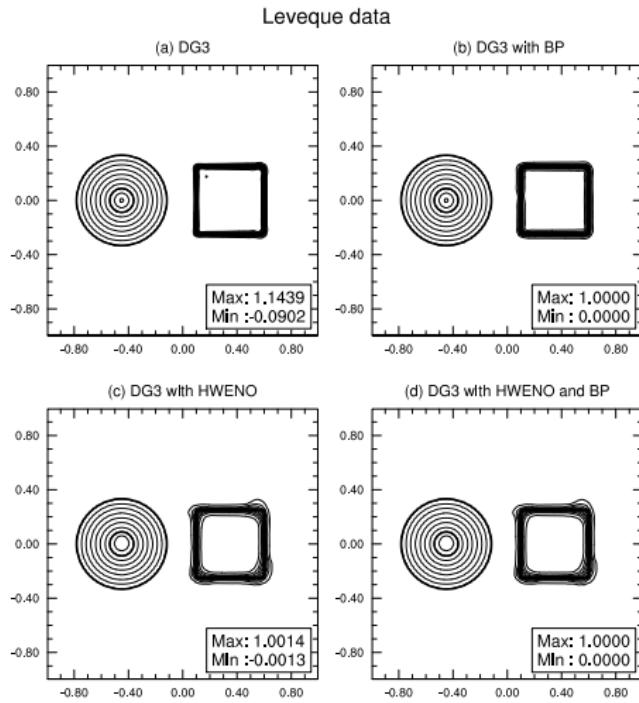
DG Advection on 2D Cartesian Grid (Solid-Body Rotation)

- A DG P^2 (third-order) Model version with 6 DOFs on 3×3 G-L grid (*Zhang & Nair, MWR, 2012*)
- Solid-body rotation (*Leveque, 2002*), 80×80 elements.



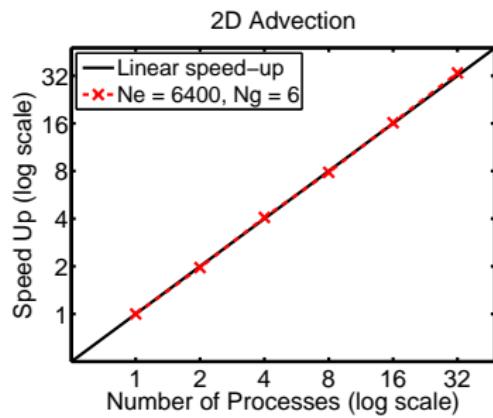
DG Advection on 2D Cartesian Grid

- HWENO uses 3×3 cells and completely removes oscillation, but more diffusive.

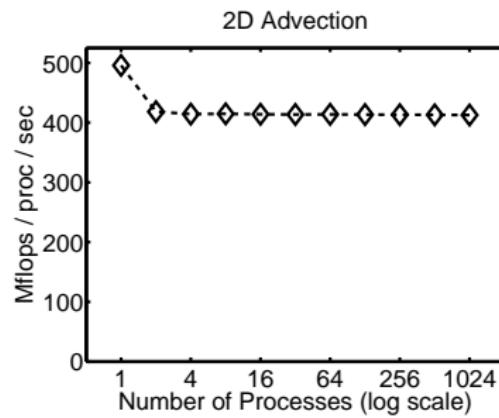


DG-2D: Scaling Results (Levy, Nair & Tufo, 2007)

- Problem: Advection of a Gaussian-hill, 80×80 elements with 6×6 GLL grid
- Strong scaling** is measured by increase the number processes running while keeping the problem size constant
- Weak scaling** is measured by scaling the problem along with the number of processors, so that work per process is constant



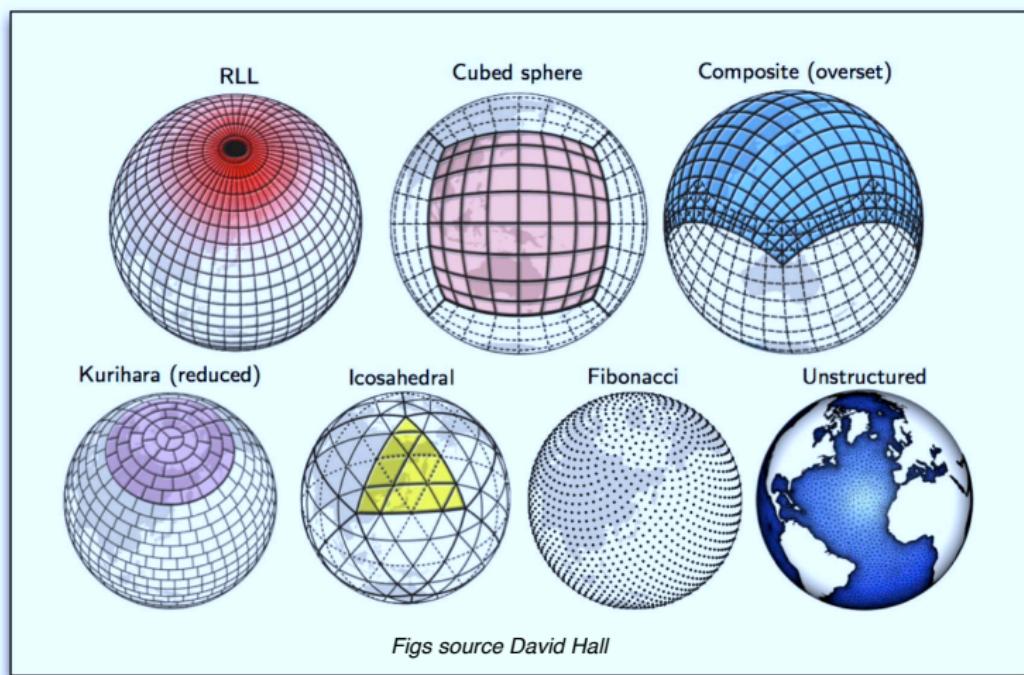
Strong scaling



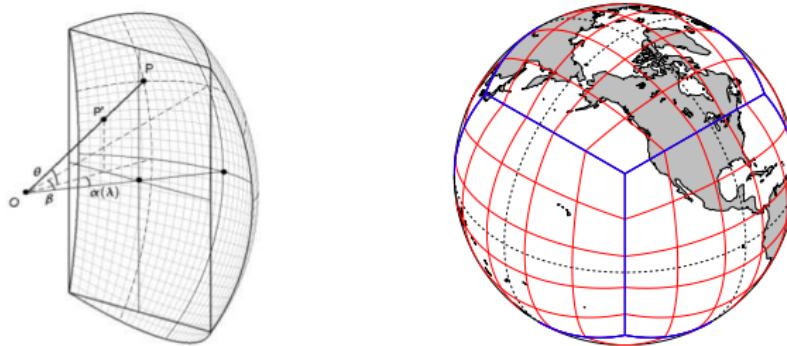
Weak scaling

Extending DG Methods to Spherical Geometry: Various Grid Options

- DG method can be potentially used on various spherical mesh with triangular or quadrilateral (or both) elements



Cubed-Sphere: Central Equiangular (Gnomonic) Projection



- The sphere is decomposed into 6 identical regions, and free of polar singularities (*Sadourny, MWR, 1972*).
 - Equiangular projection using central angles (x^1, x^2).
 - Non-orthogonal grid lines and discontinuous edges
 - All the grid lines are great-circle arcs
 - Quasi-uniform rectangular mesh, well suited for the element-based methods such as DG or SE methods (CAM-HOMME)

Non-Orthogonal Cubed-Sphere Grid System

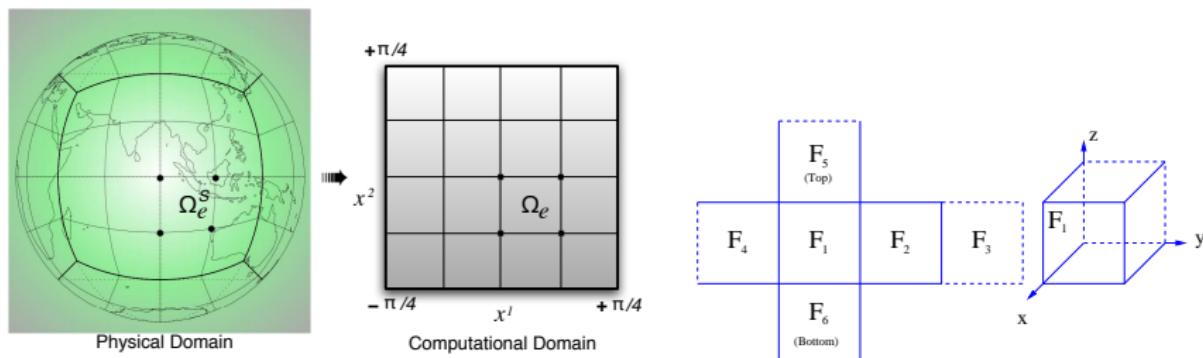
Metric term (Jacobian) of [Cubed-Sphere \rightleftharpoons Sphere] Transform on the cubed-sphere: \sqrt{G}

Central angles $(x^1, x^2) \in [-\pi/4, \pi/4]$, ($\Delta x^1 = \Delta x^2$) are the independent variables.

Transport equation (Nair et al. MWR, 2005):

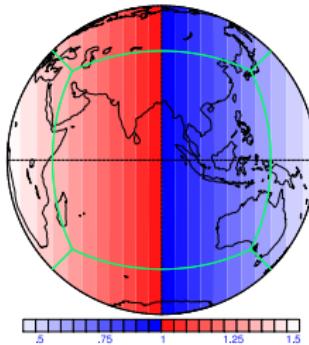
$$\frac{\partial}{\partial t}(\sqrt{G} h) + \frac{\partial}{\partial x^1}(\sqrt{G} u^1 h) + \frac{\partial}{\partial x^2}(\sqrt{G} u^2 h) = 0$$

Computational domain is the surface of cube $[-\pi/4, +\pi/4]^3$

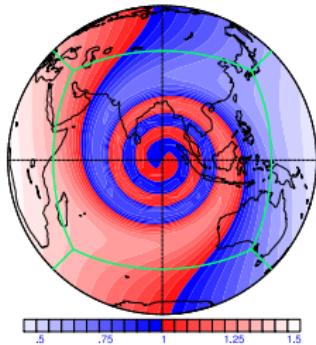


Advection: Deformational Flow (Moving Vortices on the Sphere)

DG: Moving Vortex on the Sphere (HOMME/Nair)



DG: Moving Vortex on the Sphere (HOMME/Nair)



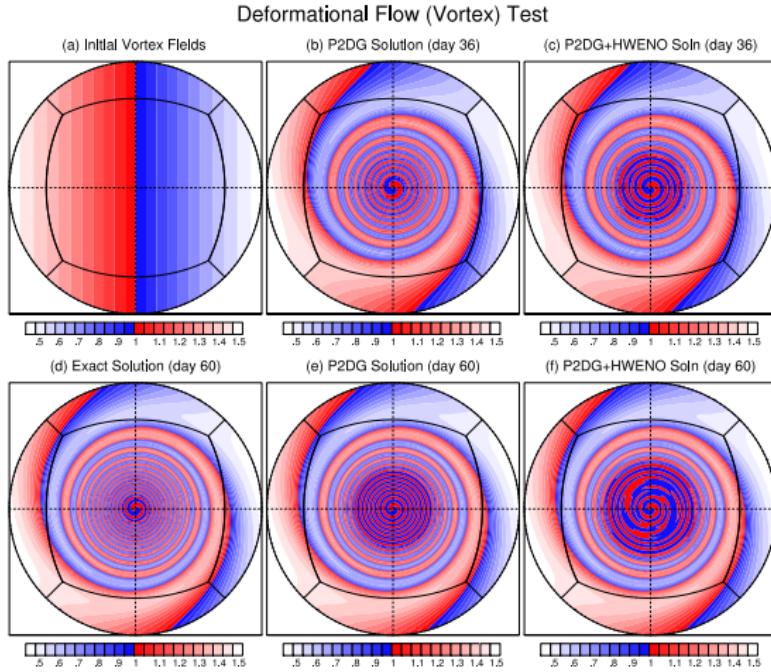
Initial field and DG solution after 12 days. Max error is $\mathcal{O}(10^{-5})$

A Smooth Deformational Flow Test [Nair & Jablonowski (MWR, 2008)]

- The vortices are located at diametrically opposite sides of the sphere, the vortices deform as they move along a prescribed trajectory.
- Analytical solution is known and the trajectory is chosen to be a great circle along the NE direction ($\alpha = \pi/4$).

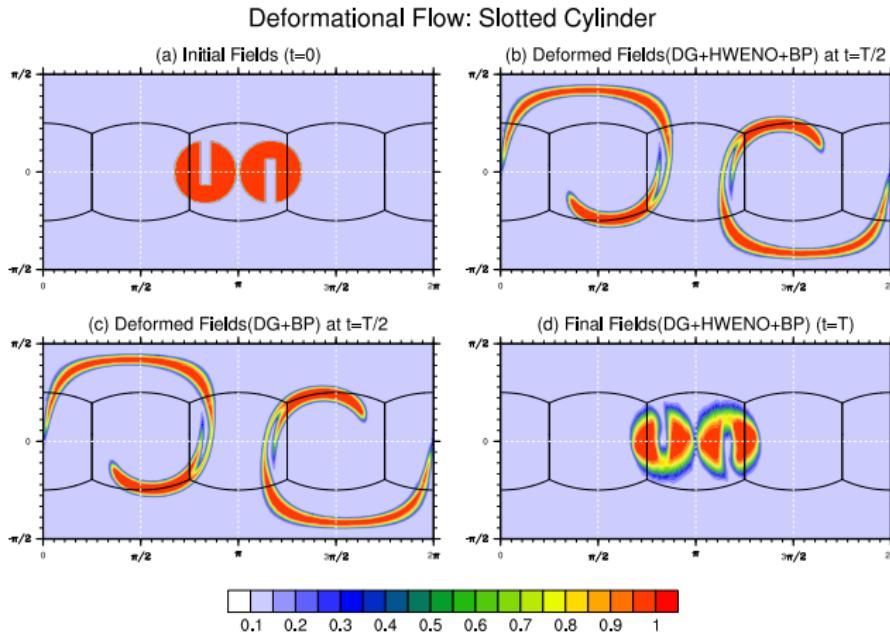
DGM Advection: Extreme deformation

- Deformational flow: Fine filament preservation (*Zhang & Nair, MWR, 2012*)
- Modal P^2 -DG with $100 \times 100 \times 6$ cells, $\Delta t = 600$ s, 60-day simulation



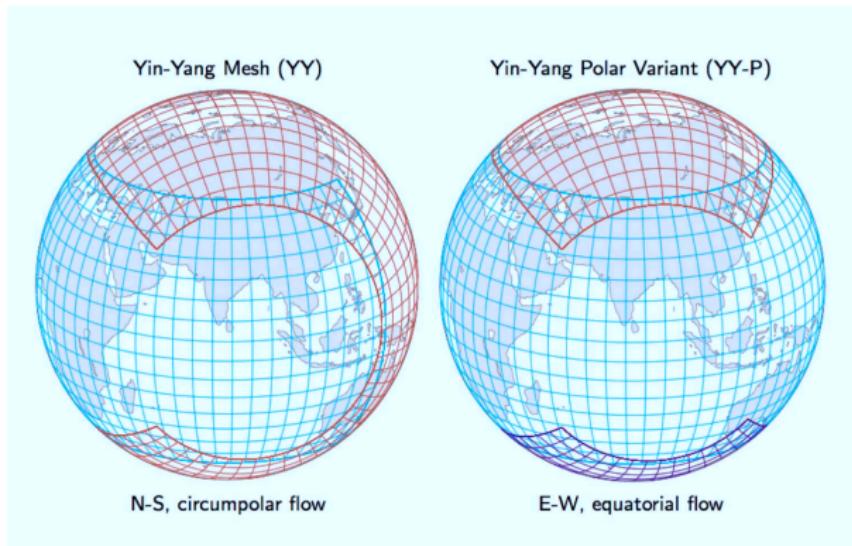
DGM Advection: Deformational flow (Slotted-Cylinder)

- Deformational flow (non-smooth deformation) (*Nair & Lauritzen, JCP, 2010*)
- Modal P^2 -DG with $45 \times 45 \times 6$ cells, $\Delta t = 0.00125\text{s}$, $T = 5$.



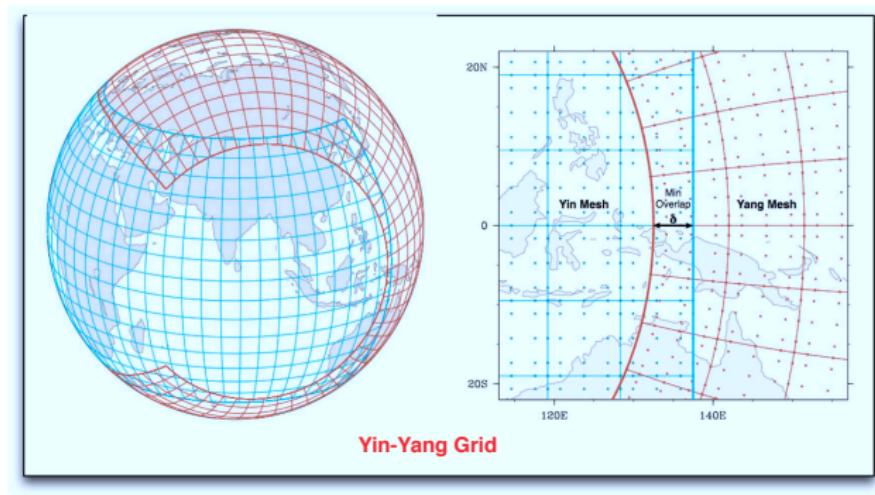
DG on Yin-Yang Overset Spherical Grid [Kageyama and Sato (2004)]

- It avoids the pole problem of the RLL grid, and there are no singular points.
- The grid spacing is quasi-uniform with a largest to smallest grid-length ratio $\sqrt{2}$
- Each grid component is orthogonal, producing a simple analytical form for PDEs.
- Overlap regions provide two set of solution.
- Numerical schemes require special treatment for conservation



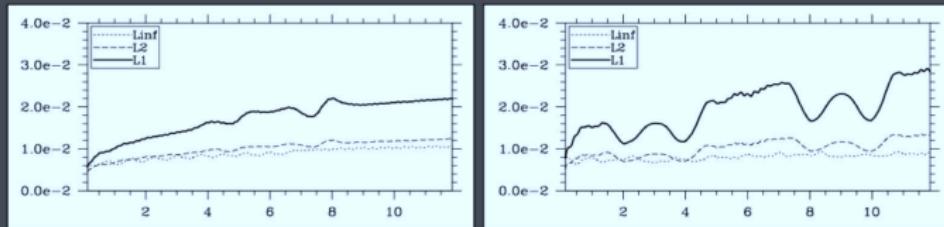
DG on Yin-Yang Grid: Advection [Hall & Nair, MWR, 2012]

- Sphere $S = Y \cup Y'$ where Y : Yin region, and Y' : Yang region. $Y \perp Y'$
- Y is a rectangular region in lat/ion (θ, λ) -space, $\lambda \in [-3\pi/2 - \delta, 3\pi/2 + \delta]$, $\theta \in [-\pi/4 - \delta, \pi/4 + \delta]$ where δ is the overlap region.
- There are total $6 \times N_e^2$ elements (DOF) for the DG spatial discretization.

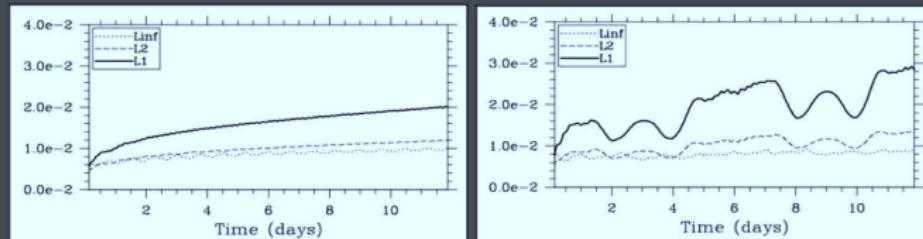


DG Advection on YY Grid: Cosine-Bell Test (Time-traces of $\ell_1, \ell_2, \ell_\infty$ errors)

Yin-Yang grid at $\alpha = 0^\circ$ and $\alpha = 45^\circ$, $N_e = 4$, $N_g = 8$



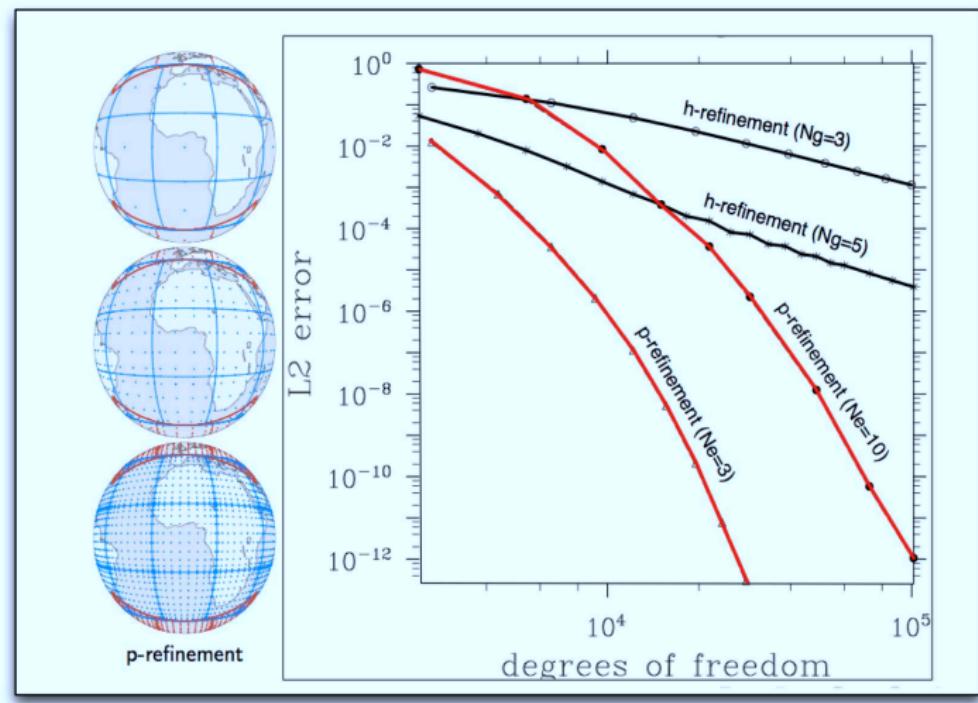
YY-Polar grid at 0° and 45° , $N_e = 4$, $N_g = 8$



- Cosine bell transport results with $N_e = 4$ and $N_g = 8$ nodes per element (approximately 3.2° resolution, and 6144 DOF).
- Figs from *Hall & Nair, MWR, 2012*
- Note: Exact mass conservation can be enforced by additional integral constraints (*Baba et al. (2010), Peng et al. (2006)*)

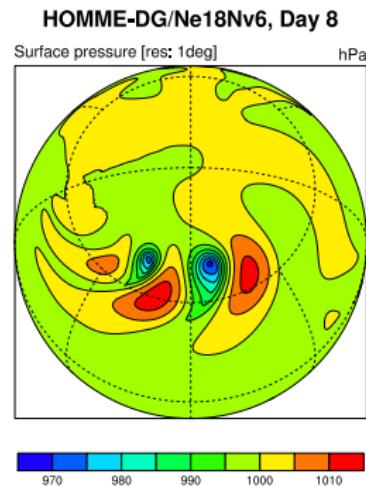
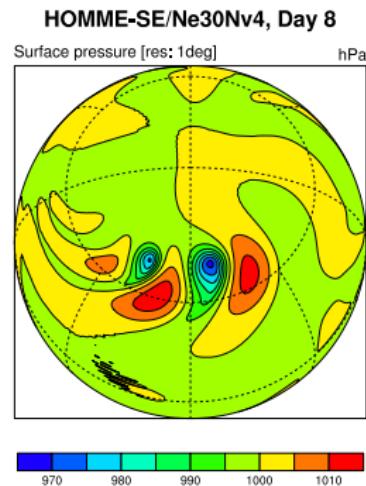
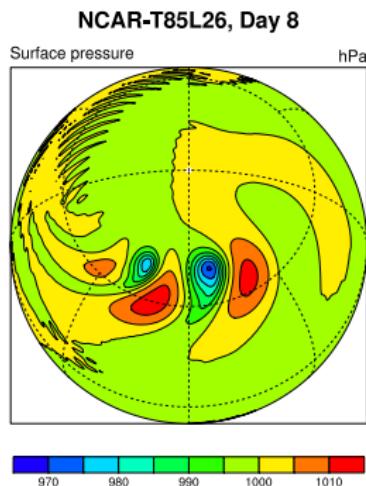
DGM Convergence: Gaussian Advection and Spectral Convergence

- Advection of a Gaussian Profile (*Levy et al. 2007; Hall & Nair, MWR, 2012*)



Beyond Advection: DG-3D Model Vs. CAM Spectral Models

- DG-3D Hydrostatic Dycore (*Nair et al. Comput. & Fluids, 2009*)
- JW-Baroclinic Instability Test, Day 8 Ps ($\approx 1^\circ$ resolution)
- The DG Solution is smooth and free from “spectral ringing”.



Part-II

- The quasi-Lagrangian coordinates for advection problems

Hydrostatic Equations in Flux Form: Curvilinear (x^1, x^2, η) coordinates

$$\frac{\partial u_1}{\partial t} + \nabla_c \cdot \mathbf{E}_1 + \dot{\eta} \frac{\partial u_1}{\partial \eta} = \sqrt{G} u^2 (f + \zeta) - R T \frac{\partial}{\partial x^1} (\ln p)$$

$$\frac{\partial u_2}{\partial t} + \nabla_c \cdot \mathbf{E}_2 + \dot{\eta} \frac{\partial u_2}{\partial \eta} = -\sqrt{G} u^1 (f + \zeta) - R T \frac{\partial}{\partial x^2} (\ln p)$$

$$\frac{\partial}{\partial t} (m) + \nabla_c \cdot (\mathbf{U}^i m) + \frac{\partial(m\dot{\eta})}{\partial \eta} = 0$$

$$\frac{\partial}{\partial t} (m\Theta) + \nabla_c \cdot (\mathbf{U}^i \Theta m) + \frac{\partial(m\dot{\eta}\Theta)}{\partial \eta} = 0$$

$$\frac{\partial}{\partial t} (mq) + \nabla_c \cdot (\mathbf{U}^i q m) + \frac{\partial(m\dot{\eta}q)}{\partial \eta} = 0$$

$$m \equiv \sqrt{G} \frac{\partial p}{\partial \eta}, \nabla_c \equiv \left(\frac{\partial}{\partial x^1}, \frac{\partial}{\partial x^2} \right), \eta = \eta(p, p_s), G = \det(G_{ij}), \frac{\partial \Phi}{\partial \eta} = -\frac{R T}{p} \frac{\partial p}{\partial \eta}$$

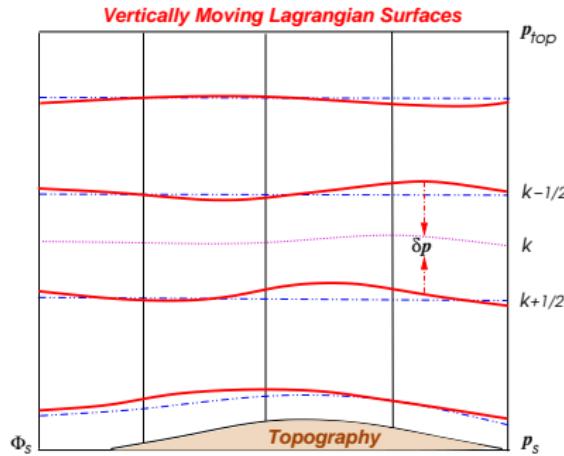
Where m is the mass function, Θ is the potential temperature and q is the moisture variable. $\mathbf{U}^i = (u^1, u^2)$, $\mathbf{E}_1 = (E, 0)$, $\mathbf{E}_2 = (0, E)$; $E = \Phi + \frac{1}{2} (u_1 u^1 + u_2 u^2)$ is the energy term. Φ is the geopotential, ζ is the relative vorticity, and f is the Coriolis term.

[Ref: HOMME/DG, Nair et al. Comput. & Fluids 2009]

Vertical (quasi) Lagrangian Coordinates (*Starr, J. Meterol. 1945*)

A “vanishing trick” for vertical advection terms!

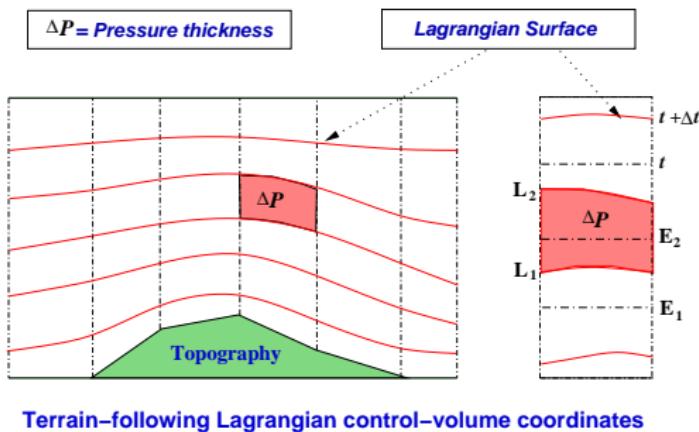
- Terrain-following Eulerian surfaces are treated as material surfaces ($\dot{\eta} = 0$).
- Simplified hydrostatic equations with no “vertical terms”
- The resulting **Lagrangian surfaces** are free to move up or down direction.



The Remapping of Lagrangian Variables

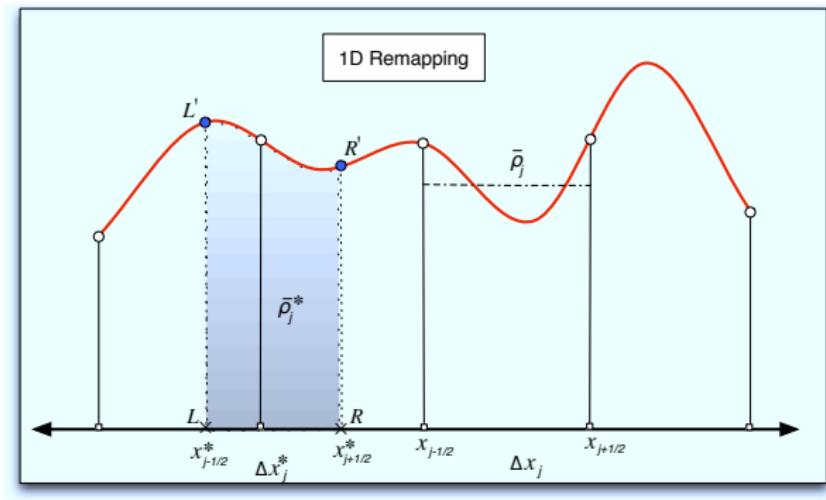
Vertically moving Lagrangian Surfaces

- Over time, Lagrangian surfaces deform and must be remapped.
- The velocity fields (u_1, u_2), and total energy (Γ_E) are **remapped** onto the reference coordinates.



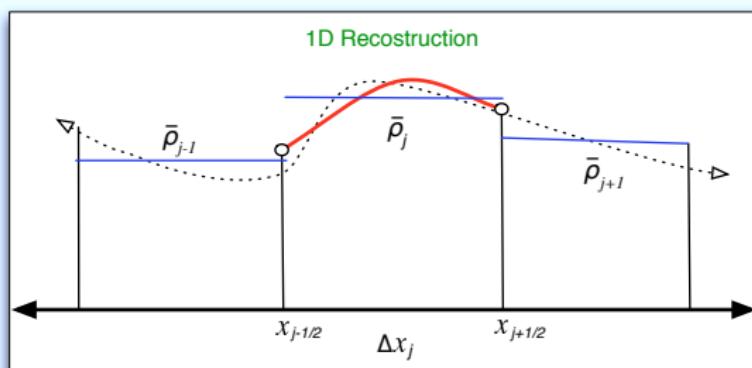
Remapping: Lauritzen & Nair, MWR, 2008; Norman & Nair, MWR, 2008)

Remapping (Rezoning or Re-gridding) on a 1D Grid



- Remapping: Interpolation from a source grid to target grid with constraints (conservation, monotonicity, positivity-preservation etc.).
- Application: Conservative semi-Lagrangian methods (e.g. CSLAM); Grid-to-grid data transfer for pre- or post-processing (GeCore).

Remapping (Rezoning or Re-gridding) on a 1D Grid

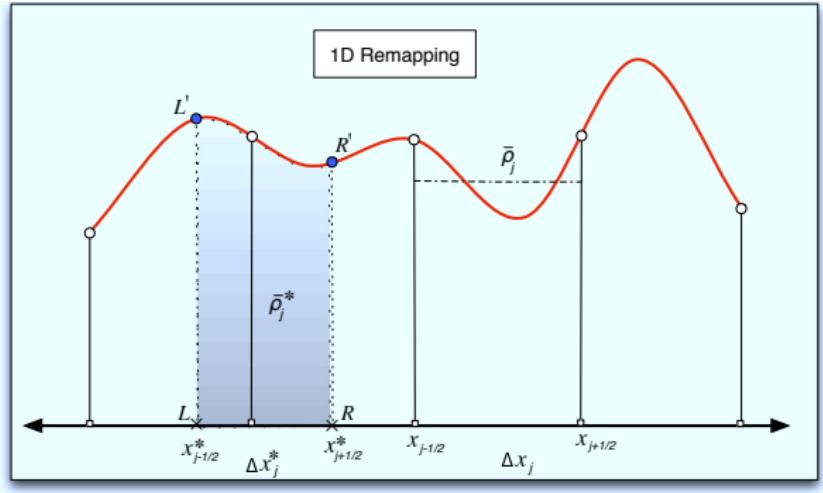


- **Reconstruction:** Fit a piecewise polynomial $\rho_j(x)$ for every cell $\Delta x_j = x_{j+1/2} - x_{j-1/2}$, using the known cell-average values $\bar{\rho}_j$ from the neighboring cells.
- The subgrid-scale distribution $\rho_j(x)$ must satisfy the conservation constraint:

$$\bar{\rho}_j = \frac{1}{\Delta x_j} \int_{x_{j-1/2}}^{x_{j+1/2}} \rho_j(x) dx, \quad \Rightarrow \text{Mass} = \bar{\rho}_j \Delta x_j$$

- $\rho_j(x)$ may be further modified to be monotonic (E.g: PLM, PPM, PCM, PHM)

Remapping 1D Grid



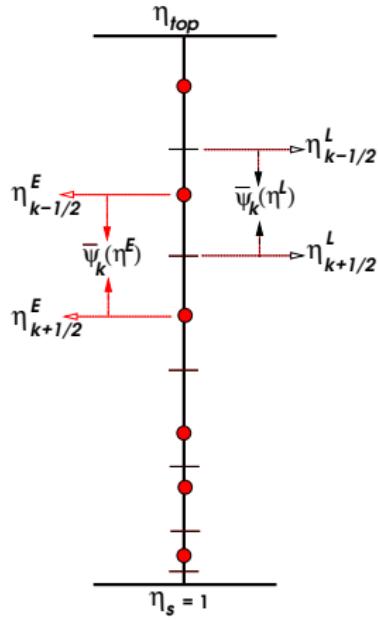
- Mass in the target cell ($\Delta x_j^* = x_{j+1/2}^* - x_{j-1/2}^*$) can be expressed as the difference of "Accumulated Mass" (\mathcal{A}_m):

$$\bar{\rho}_j^* \Delta x_j^* = \mathcal{A}_m(RR') - \mathcal{A}_m(LL') \implies \bar{\rho}_j^* = \frac{1}{\Delta x_j^*} [\mathcal{A}_m(RR') - \mathcal{A}_m(LL')]$$

where

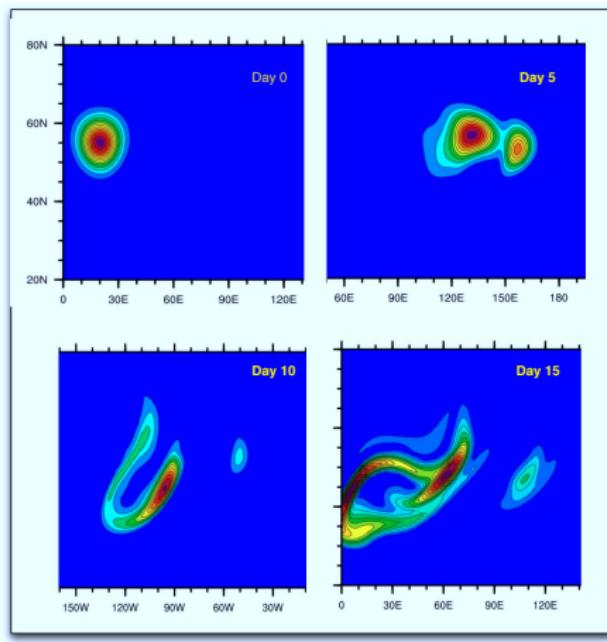
$$\mathcal{A}_m(RR') = \int_{x_{Ref}}^{x_{j+1/2}^*} \rho(x) dx = \sum_{k=1}^{j-1} \bar{\rho}_k \Delta x_k + \int_{x_{j-3/2}}^{x_{j+1/2}^*} \rho_{j-1}(x) dx$$

Vertical Advection with Lagrangian η -Coordinate



- Reference (initial) grid $\eta = \eta(P, P_s) \in [\eta_{top}, 1]$
- Source grid = Lagrangian η_k^L
- Target grid = Eulerian η_k^E ; $\sum \Delta \eta_k^E = \sum \Delta \eta_k^L$
- Lagrangian η_k^L can be computed from the predicted “pressure thickness” ΔP (CAM-FV)
- Remapping is performed at every advective Δt
- Every 1D vertical trajectory information can be “recycled” for all tracers

3D Transport (CAM-SE): SE horizontal + vertical remapping



- CAM-SE (1°): JW-Test divergent flow field. SE horizontal transport is quasi-monotonic
- $\Delta t_a = 4 \times 90$ s, vertical remapping by PCM (Zerroukat, 2005) for advection.

Figure courtesy: Christoph Erath

Summary & Conclusions :

- The DG method with moderate order (third or fourth) is an excellent choice for transport problems as applied in atmospheric sciences. DGM addresses:
 - 1 High-order accuracy
 - 2 Geometric flexibility
 - 3 Positivity-preserving advection
 - 4 High parallel efficiency
 - 5 Local and global conservation
- In comparison with finite-volume and finite-difference implementations of the Yin-Yang grid, the DG approach is considerably simpler as the overset interpolation is local, requiring information from the interior of a single element.
- In general, modified YY-P and YY meshes exhibited similar performance on most tests, while the YY-P mesh performed better on cases with strictly zonal flow.
- DG method is an ideal candidate for the new generation petascale-capable dynamical cores.
- The “moving” vertical Lagrangian (evolve and remap approach) method provides an efficient way for 3D conservative multi-tracer transport.

THANK YOU!