

Buoyancy Effects on Turbulence: 2

$$\frac{\partial \bar{q}^2}{\partial t} = \Phi + J_b - \epsilon$$

Simple shear flow, only $\partial U_1 / \partial x_3 \neq 0$

now it's stratified \rightarrow gradient of density, ρ
component eq'n's: $\bar{u}_1 \bar{u}_3 \partial U_1 / \partial x_3$

$$\frac{D \bar{u}_1^2}{Dt} = \Phi - \epsilon/3 + \overline{\rho} \frac{\partial \bar{u}_1}{\partial x_1}$$

$$\frac{D \bar{u}_2^2}{Dt} = -\epsilon/3 + \overline{\rho} \frac{\partial \bar{u}_2}{\partial x_2}$$

$$\frac{D \bar{u}_3^2}{Dt} = J_b - \epsilon/3 + \overline{\rho} \frac{\partial \bar{u}_3}{\partial x_3}$$

Φ - source ϵ, J_b - sink

pressure-work terms redistribute energy between $\bar{u}_1^2, \bar{u}_2^2, \bar{u}_3^2$

since J_b acts only on vertical term

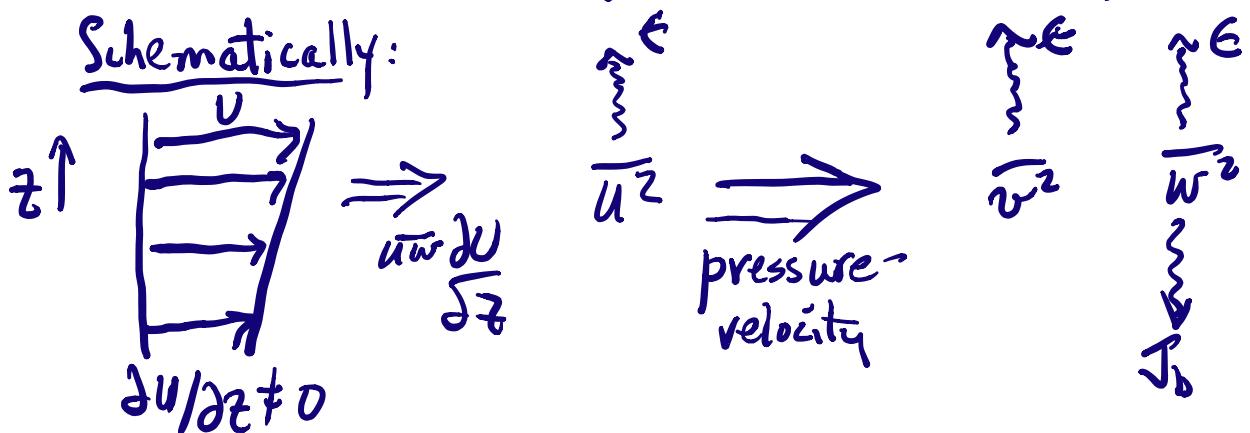
we might expect that \bar{u}_3^2 is preferentially suppressed at large scales - evidence from obs \Rightarrow revisit

when we discuss spectra.

Convection: heating from below, not sheared

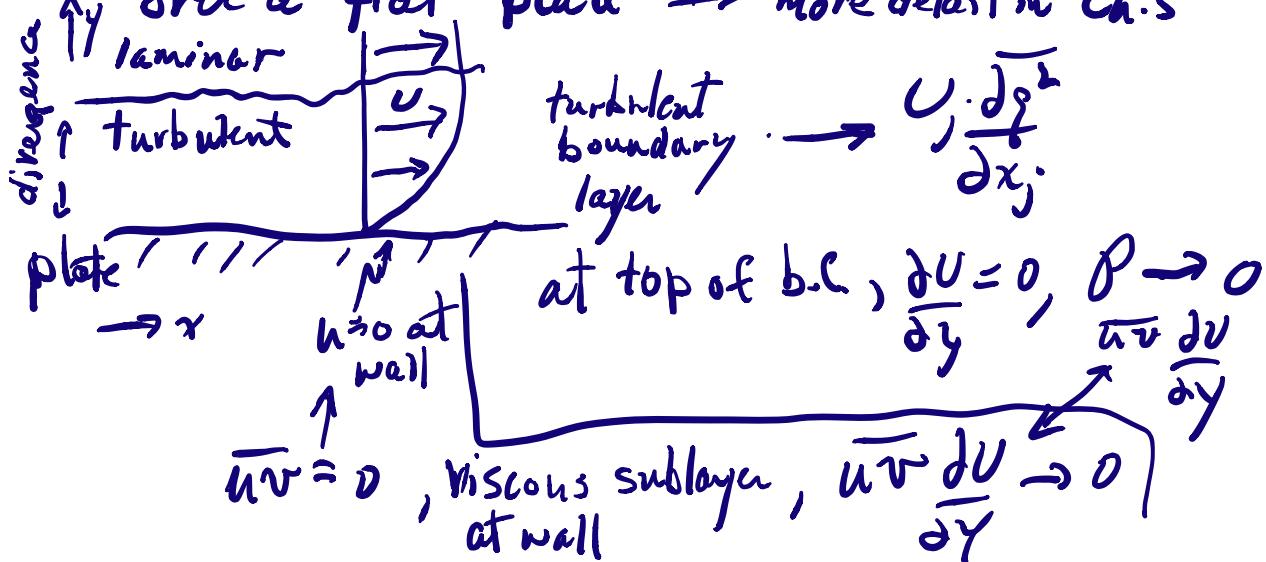
$$\frac{\partial U_i}{\partial x_j} = 0 \text{ all comp.}$$

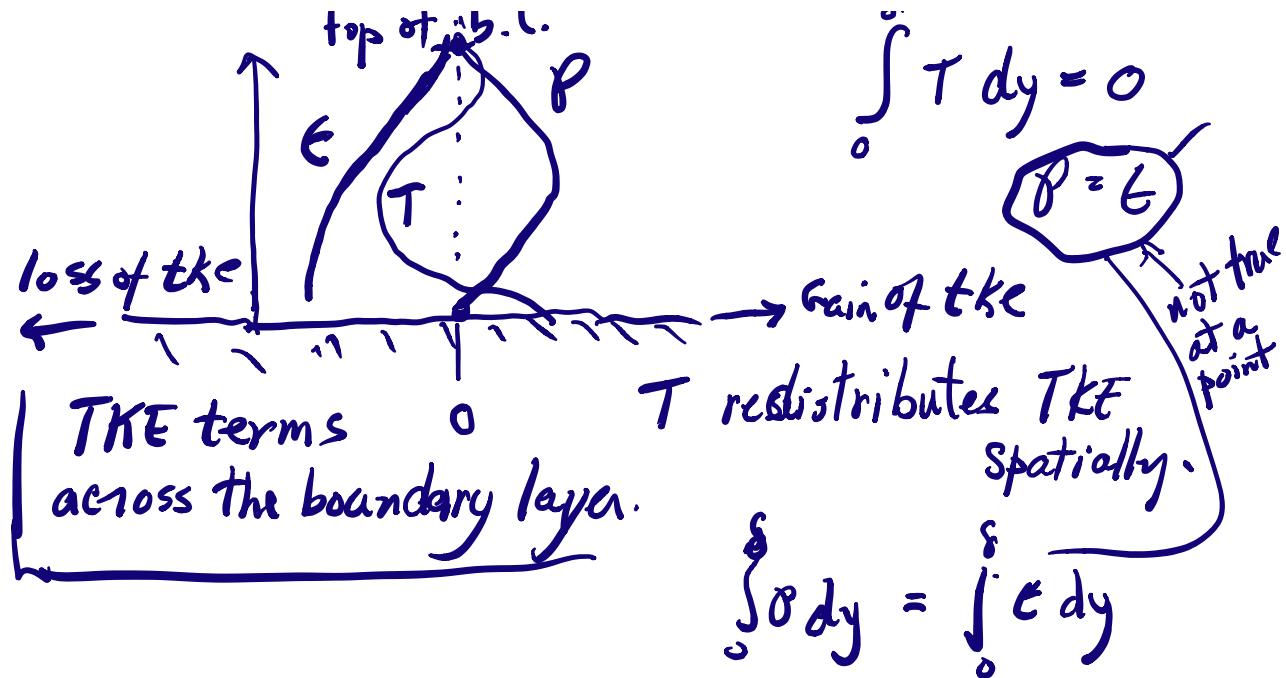
Then J_b is a source of tke. \rightarrow generates \bar{u}_3^2
and pressure-velocity redistribute to \bar{u}_1^2, \bar{u}_2^2



Consider transport(T), ρ , ϵ but not J_b yet

Consider example (schematically) of boundary layer
over a flat plate \rightarrow more detail in Ch.5





Buoyancy length scale l_b , L_o

L_o Ozmidov

consider a fluid parcel of dimension l , velocity u .
in a stratified fluid

inertial force $\rightarrow F_I = \hat{u} \cdot \nabla \hat{u} \sim u^2/l \leftarrow NL$
per unit mass if $u \sim (\epsilon l)^{1/3}$, $F_I = \frac{(\epsilon l)^{2/3}}{l} = \left(\frac{\epsilon^2}{l}\right)^{1/3}$

we integrate over the fluid parcel with volume $\sim l^3$

$$F_I^V = \rho_0 \int F_I dV \sim \rho_0 l^3 \left(\frac{\epsilon^2}{l}\right)^{1/3} = \rho_0 \epsilon^{2/3} l^{8/3}$$

buoyant force: $F_b = \delta\rho Vg$ $\delta\rho$ is the difference in ρ between parcel & surround.

define $N^2 = -g/\rho \frac{\partial \rho}{\partial z}$ \leftarrow a measure of stratification

units s^{-2} → squared frequency

$N \rightarrow$ frequency → natural frequency of oscillation of internal wave

$$N^2 \sim \frac{g}{\rho_0} \frac{\delta p}{l} \rightarrow \delta p = \rho_0 l N^2$$

$$\text{so } F_b \sim \rho_0 \frac{l N^2 \cdot l^3 \cdot g}{g} = \rho_0 N^2 l^4$$

This gives 2 scale-dependent estimates for F_b, F_I
in case buoyant forces = inertial forces

$$F_b = F_I$$
$$\rho_0 N^2 l^4 = \rho_0 \epsilon^{2/3} l^{8/3}$$
$$\boxed{l_b = (\epsilon/N^3)^{1/2}} \leftarrow \begin{matrix} \text{buoyancy} \\ \text{length} \\ \text{scale} \end{matrix}$$

at l_b in a stratified turbulence,
buoyant forces damp motion

scales $\gg l_b$, turbulence is damped by F_b

$\ll l_b$, should be free of buoyant damping
↑ locally isotropic turbulence

This scaling can be derived on dimensional grounds if ϵ, N are the only factors

that determine $u, l \leftarrow$ energy-containing scales.

$$\text{then } u_b = (\epsilon/N)^{1/2}, \quad k_b = \frac{2\pi}{l_b} = (N^3/\epsilon)^{1/2}$$

we will use l_b to help in gaining empirical-theoretical solutions to time-dependent tke equation \rightarrow

\rightarrow will arise later when we discuss spectra.