

GPHY 5513

3D Seismic Interpretation

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Coherence



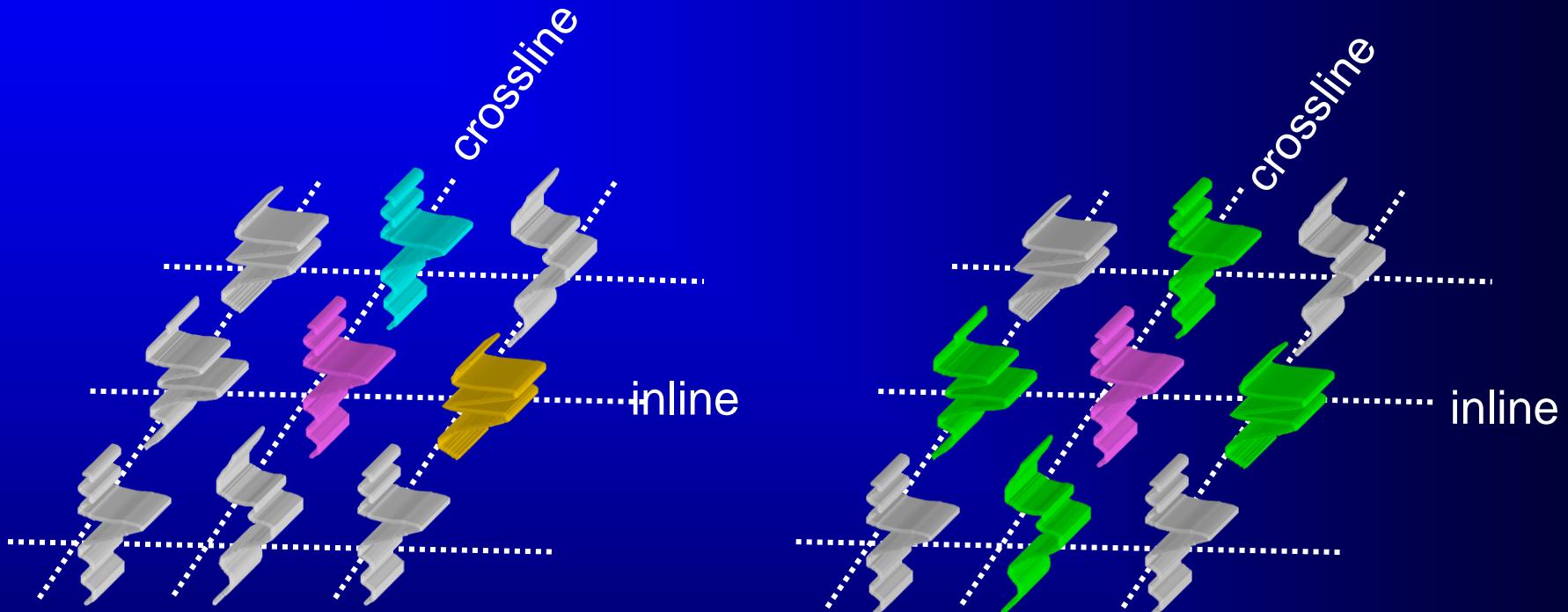
Coherence

After this section you will be able to:

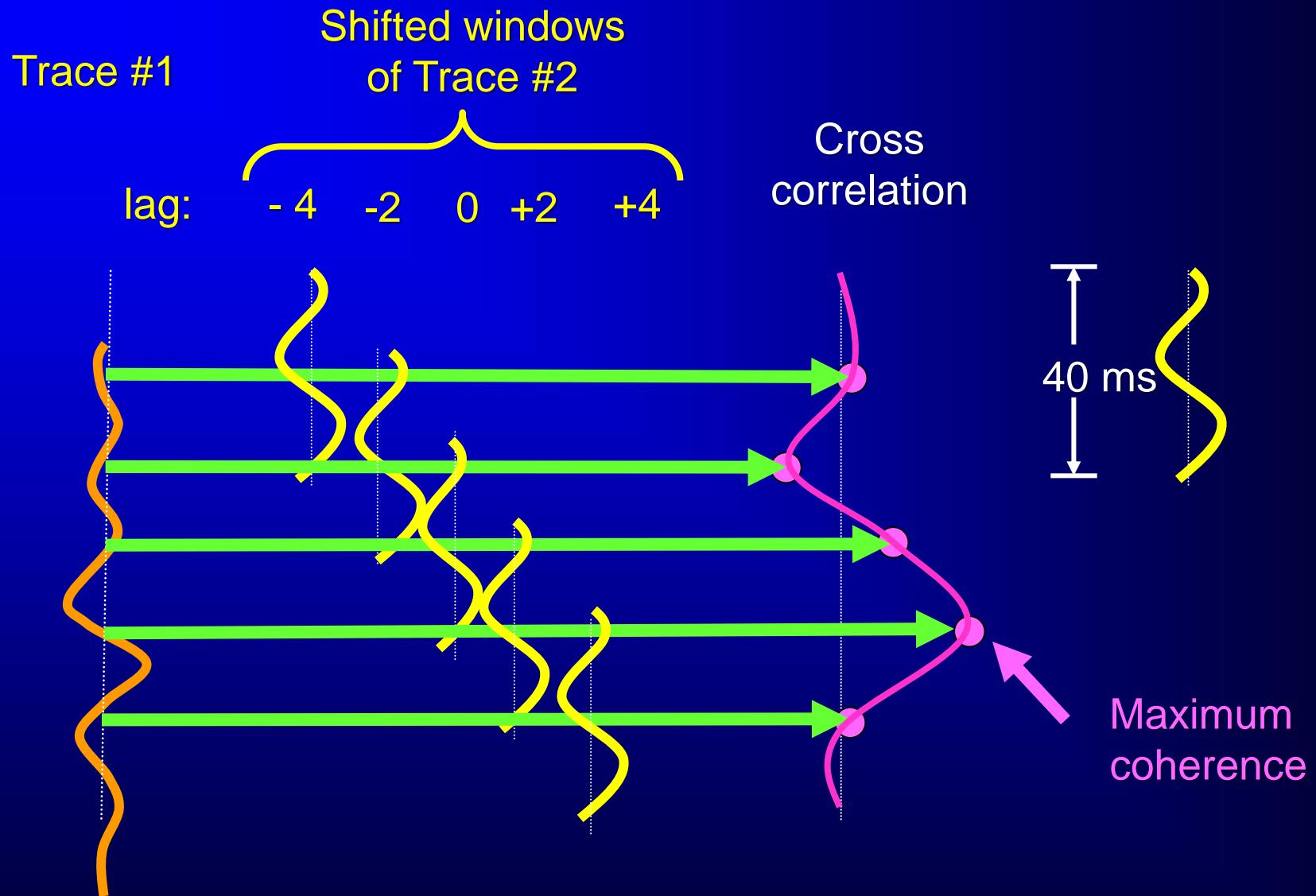
- Summarize the physical and mathematical basis of currently available seismic coherence algorithms,
- Evaluate the impact of spatial and temporal analysis window size on the resolution of geologic features,
- Recognize artifacts due to structural leakage and seismic zero crossings, and
- Apply best practices for structural and stratigraphic interpretation.



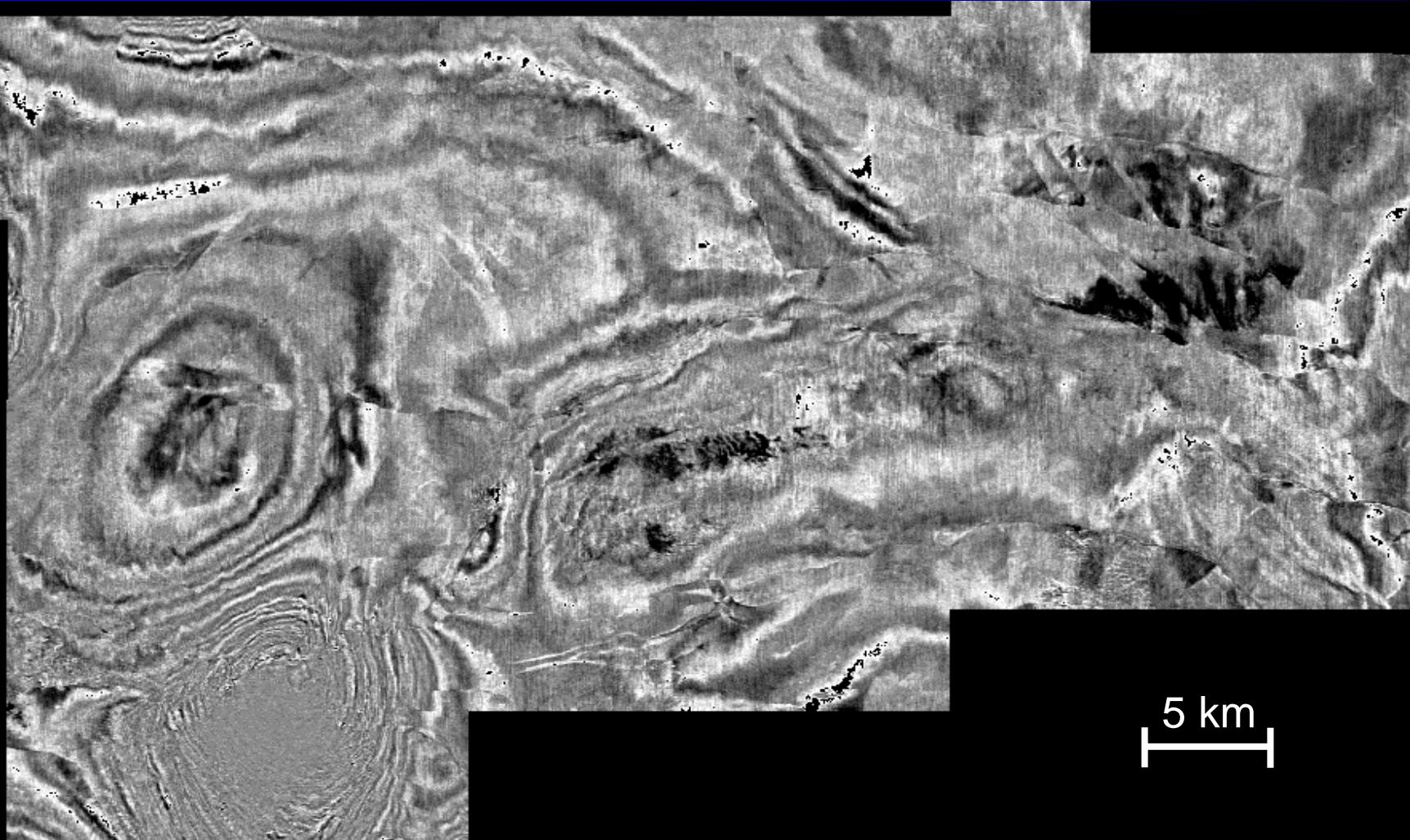
Coherence compares the waveforms of neighboring traces



Cross correlation of 2 traces



Seismic Time Slice

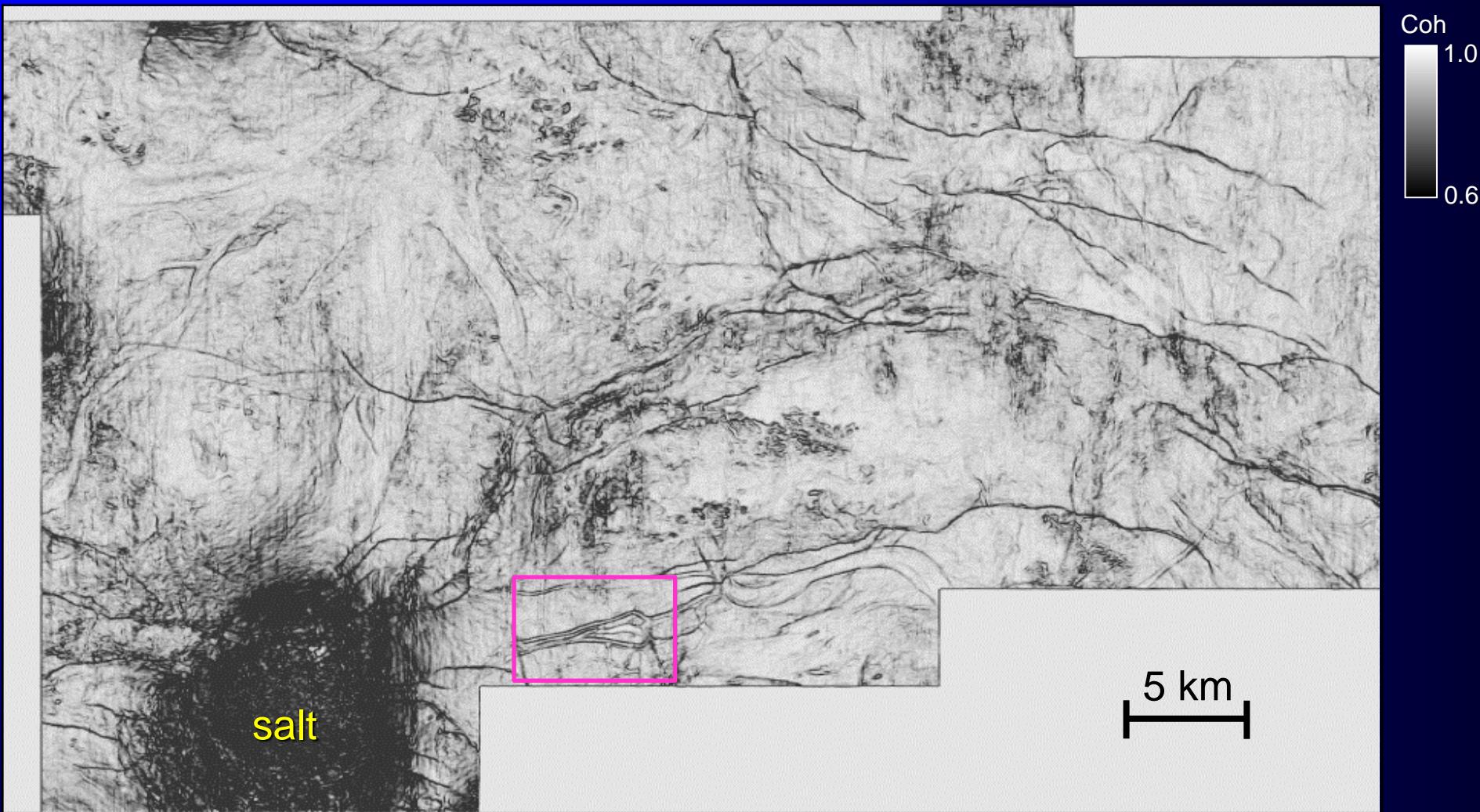


5 km



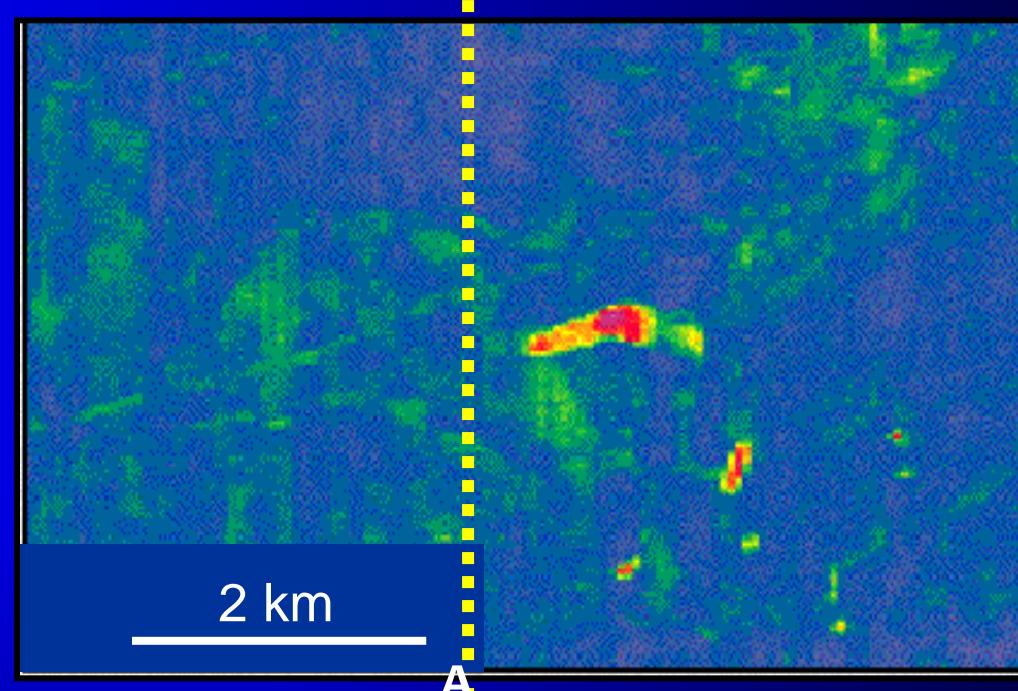
(Bahorich and Farmer, 1995)

Coherence Time Slice

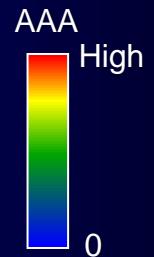
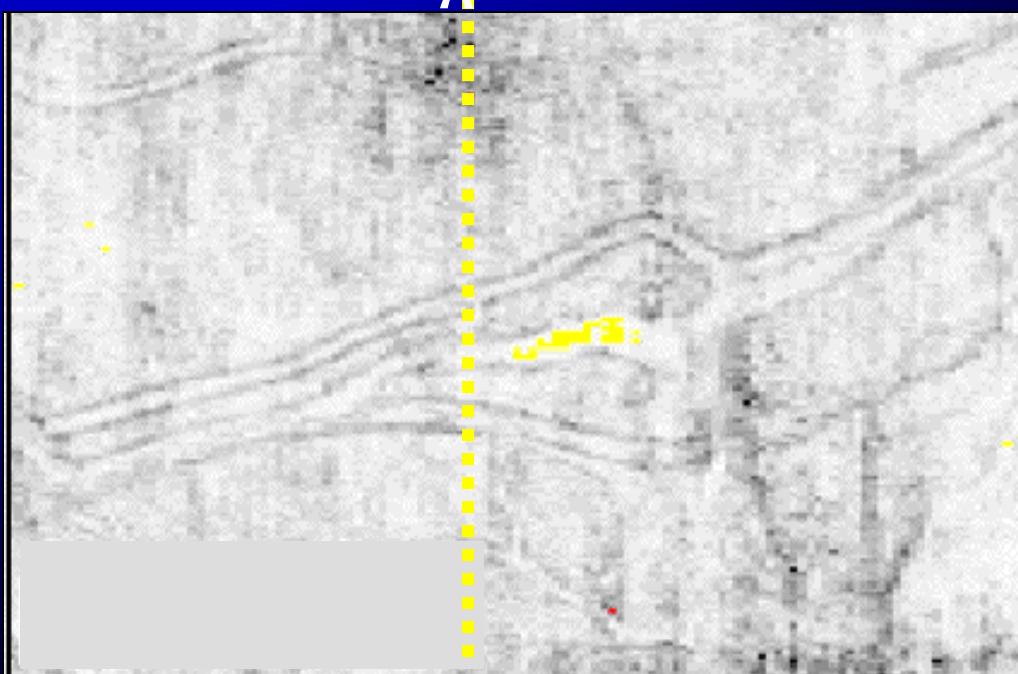


(Bahorich and Farmer, 1995)

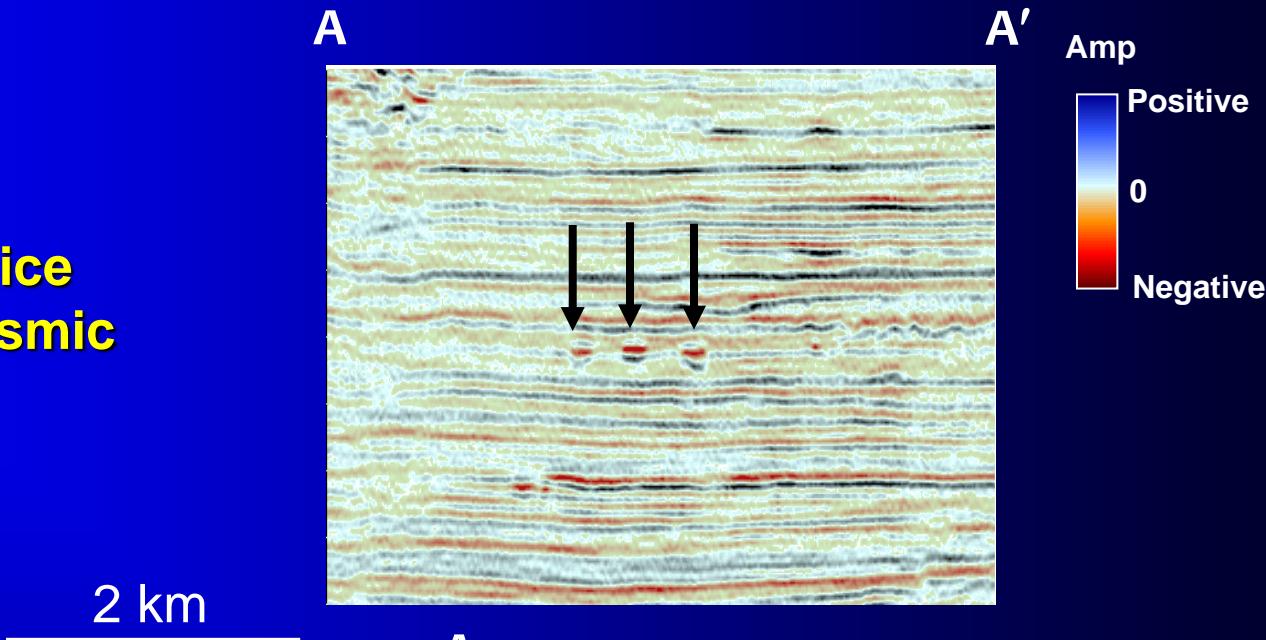
**Time slice through
average absolute
amplitude**



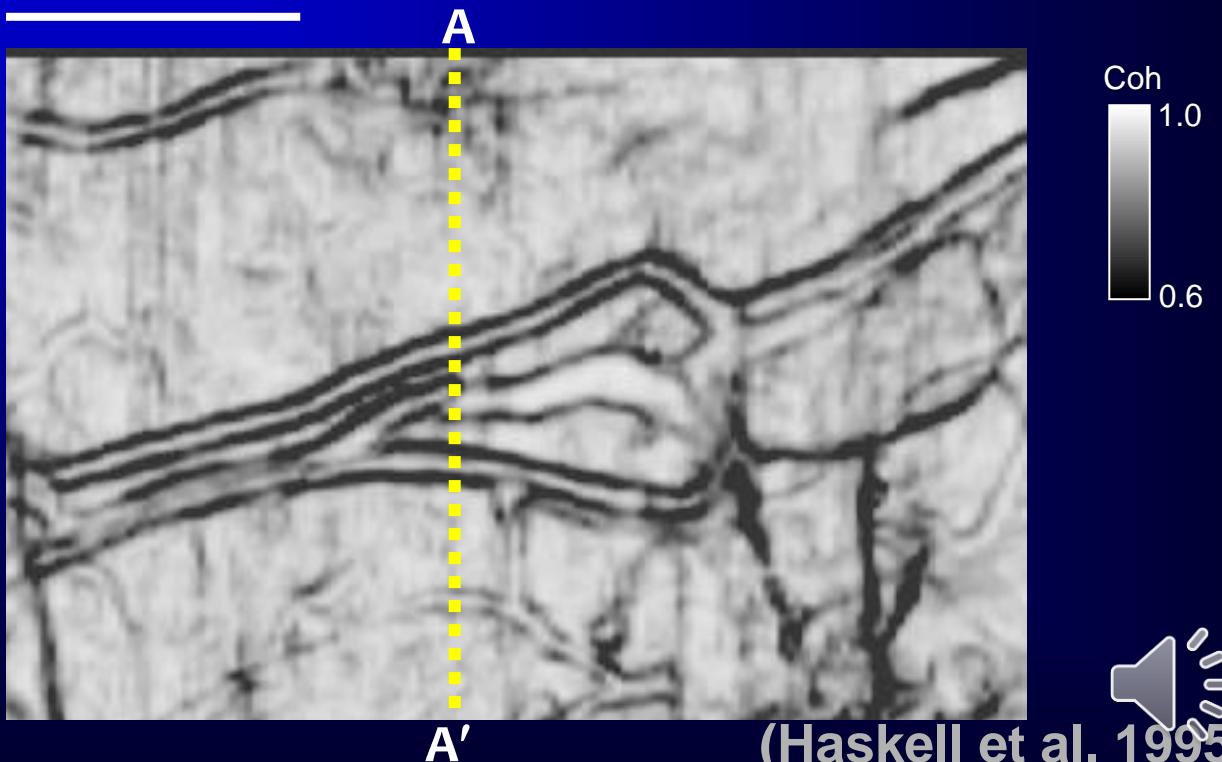
**Time slice through
coherence
(early algorithm)**



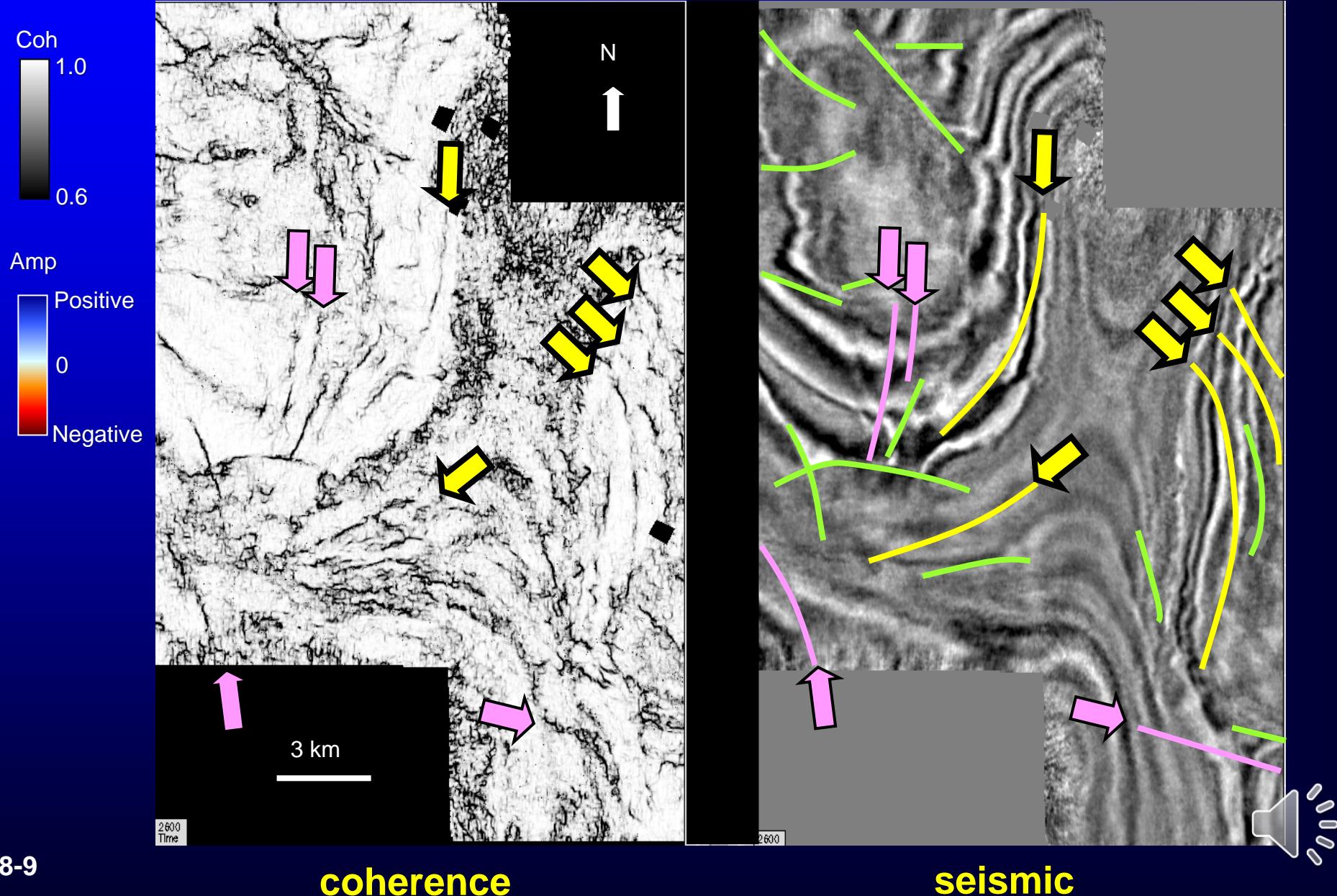
Vertical slice through seismic



Time slice through coherence (later algorithm)



Appearance faults perpendicular and parallel to strike



Alternative measures of waveform similarity

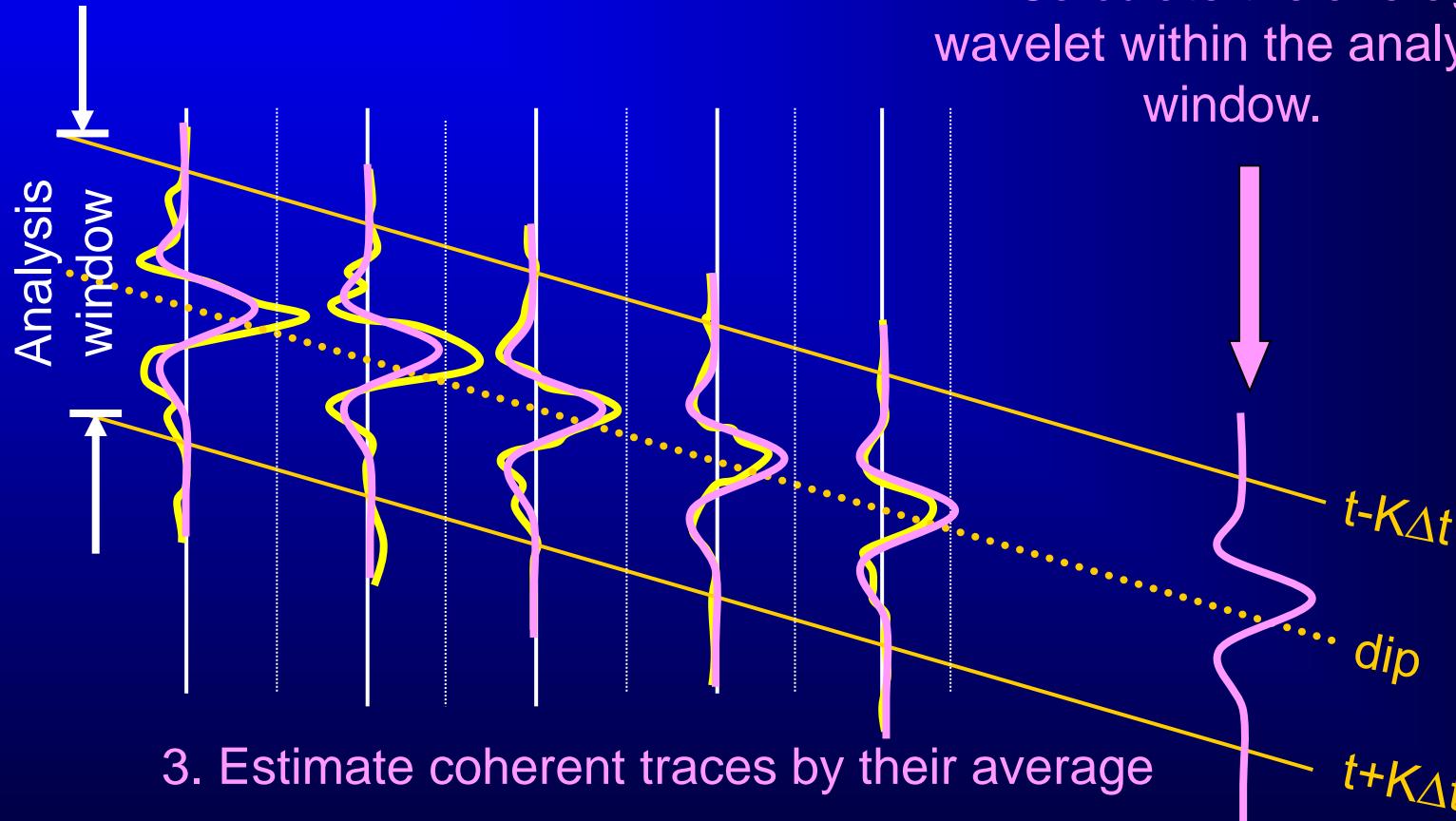
- cross correlation
- semblance, variance, and Manhattan distance
- eigenstructure
- Gradient Structural Tensors (GST)
- plane-wave destructors



Semblance estimate of coherence

$$5. \text{ coherence} = \frac{\text{energy of average traces}}{\text{Average energy of input traces}}$$

1. Calculate energy of input traces



Semblance estimate of coherence

$$c_s = \frac{\sum_{k=-K}^{+K} \left(\frac{1}{J} \sum_{j=1}^J [u(k\Delta t - px_j - qy_j)] \right)^2}{\sum_{k=-K}^{+K} \frac{1}{J} \left(\sum_{j=1}^J [u(k\Delta t - px_j - qy_j)]^2 \right)}$$

← Energy of the average trace

← Average of the energy of all the traces

Variance estimate of coherence

The diagram illustrates three different ways to derive the formula for the variance estimate of coherence:

- A blue arrow labeled "Classic definition of variance" points to the formula.
- A yellow arrow labeled "Statistical definition of variance" points to the formula.
- A yellow arrow labeled "'Fast' computation of variance" points to the formula.

$$\sigma^2 \equiv \frac{1}{J} \sum_{j=1}^J (u_j - m)^2 = \frac{1}{J} \left[\sum_{j=1}^J (u_j^2) - m^2 \right]$$

$$c_v = \frac{\sum_{k=-K}^{+K} \left(\frac{1}{J} \left\{ \sum_{j=1}^J [u(k\Delta t - px_j - qy_j)]^2 - \left[\frac{1}{J} \sum_{j=1}^J u(k\Delta t - px_j - qy_j) \right]^2 \right\} \right)}{\sum_{k=-K}^{+K} \frac{1}{J} \left(\sum_{j=1}^J [u(k\Delta t - px_j - qy_j)]^2 \right)} = 1 - c_s$$



The ‘Manhattan Distance’: $r=|x-x_0|+|y-y_0|$



The ‘as the crow flies’ (or Pythagorean) distance’
 $r=[(x-x_0)^2+(y-y_0)^2]^{1/2}$

New York City Archives



Manhattan distance estimate of coherence

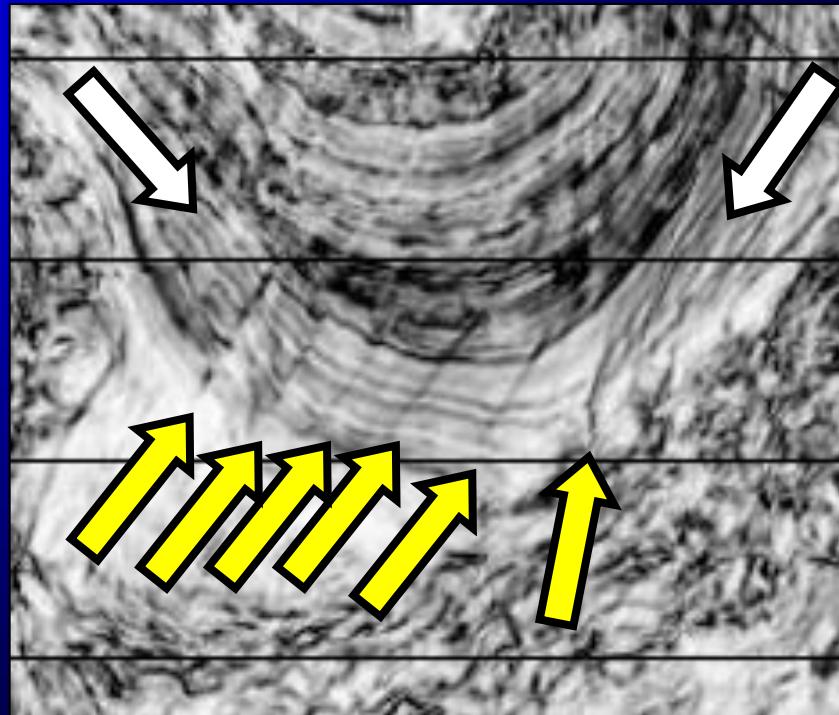
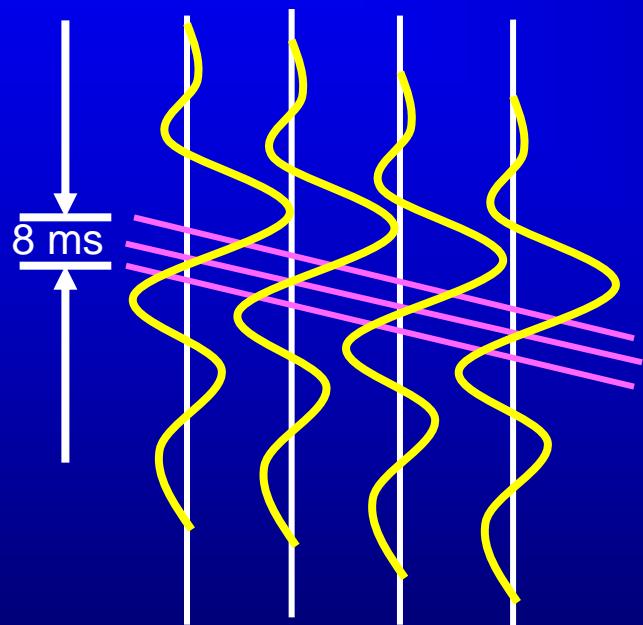
$$c_s = \frac{\sum_{k=-K}^{+K} \frac{1}{J} \left| \sum_{j=1}^J u(k\Delta t - px_j - qy_j) \right|}{\sum_{k=-K}^{+K} \frac{1}{J} \sum_{j=1}^J |u(k\Delta t - px_j - qy_j)|}$$

Absolute value of
the average trace

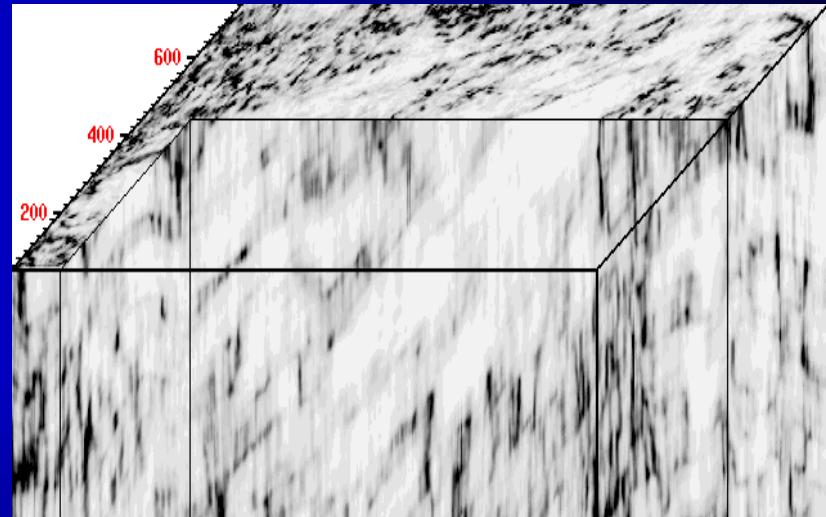
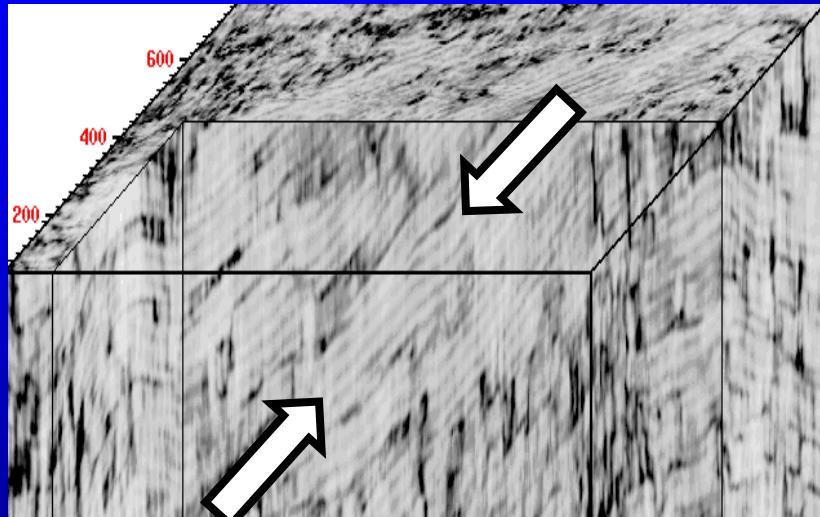
Average of the
absolute value of
all the traces



Pitfall: Banding artifacts near zero crossings



Solution: calculate coherence on the analytic trace



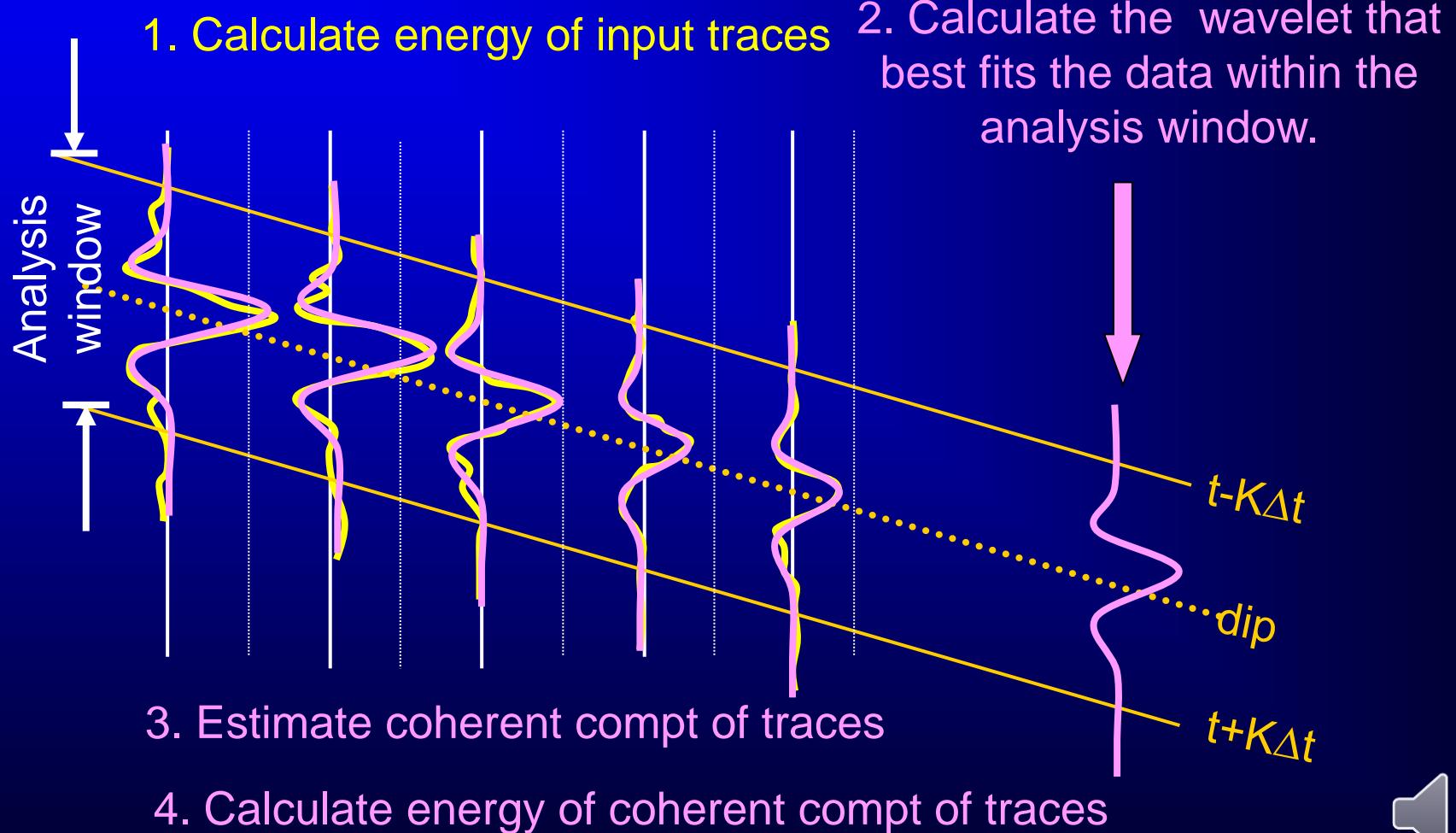
$$c_s = \frac{\sum_{k=-K}^{+K} \left(\frac{1}{J} \sum_{j=1}^J [u(k\Delta t - px_j - qy_j)] \right)^2 + \left(\frac{1}{J} \sum_{j=1}^J [u^H(k\Delta t - px_j - qy_j)] \right)^2}{\sum_{k=-K}^{+K} \frac{1}{J} \left(\sum_{j=1}^J [u(k\Delta t - px_j - qy_j)]^2 + \sum_{j=1}^J [u^H(k\Delta t - px_j - qy_j)]^2 \right)}$$

Coherence of real trace

Coherence of analytic trace

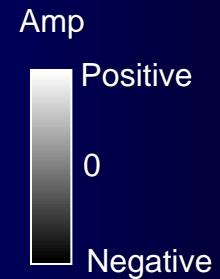
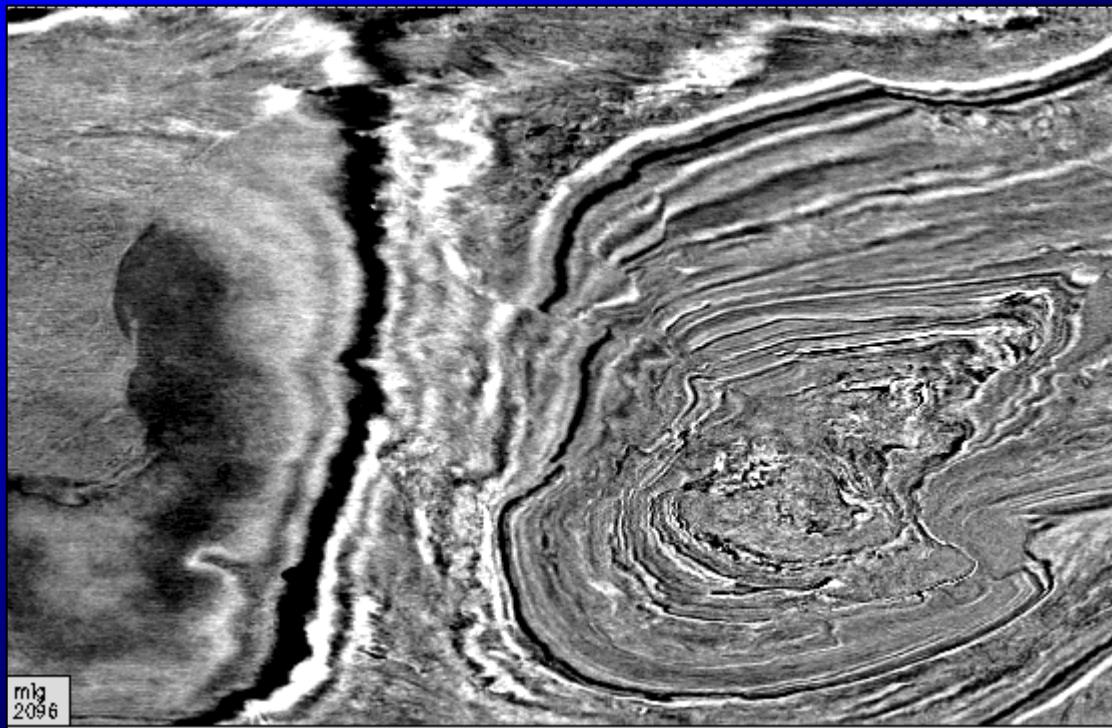
Eigenstructure estimate of coherence

$$5. \text{ coherence} = \frac{\text{energy of coherent comp}}{\text{energy of input traces}}$$



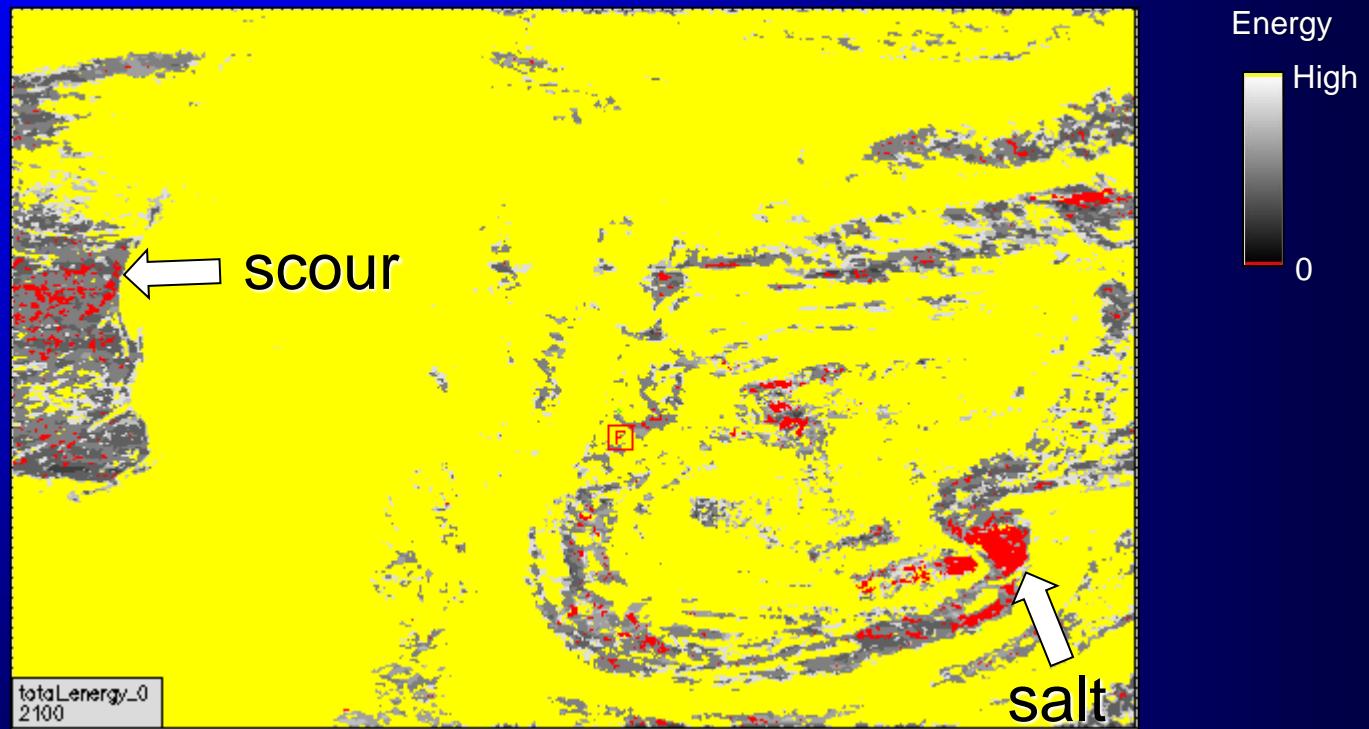
Eigenstructure coherence:

Time slice through seismic



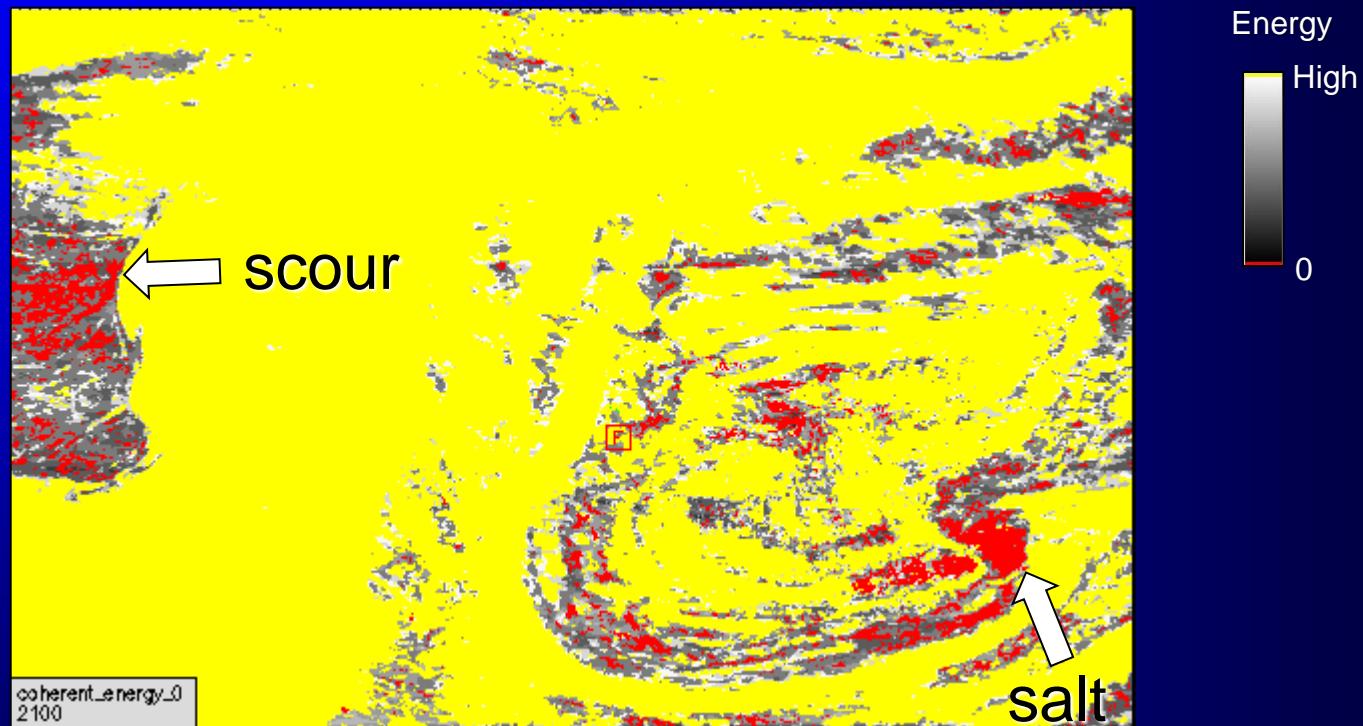
Eigenstructure coherence:

Time slice through total energy in 9 trace, 40 ms window



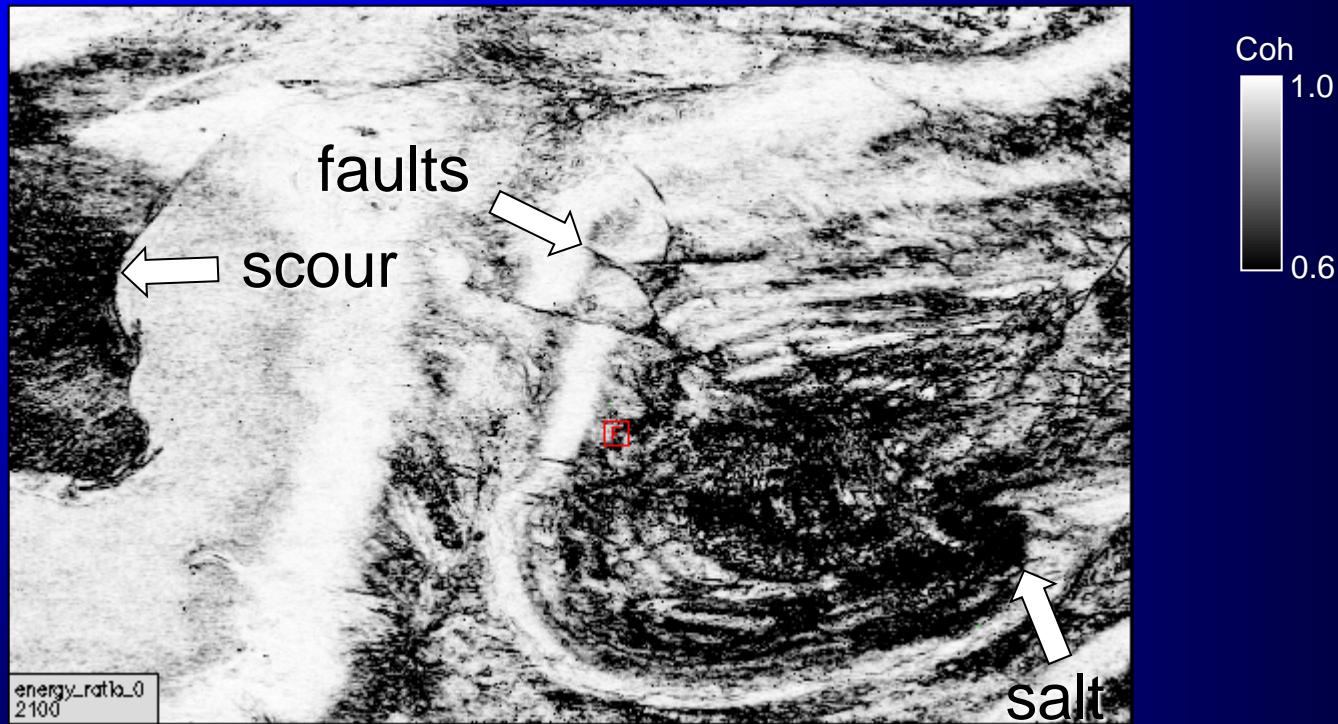
Eigenstructure coherence:

Time slice through coherent energy in 9 trace, 40 ms window

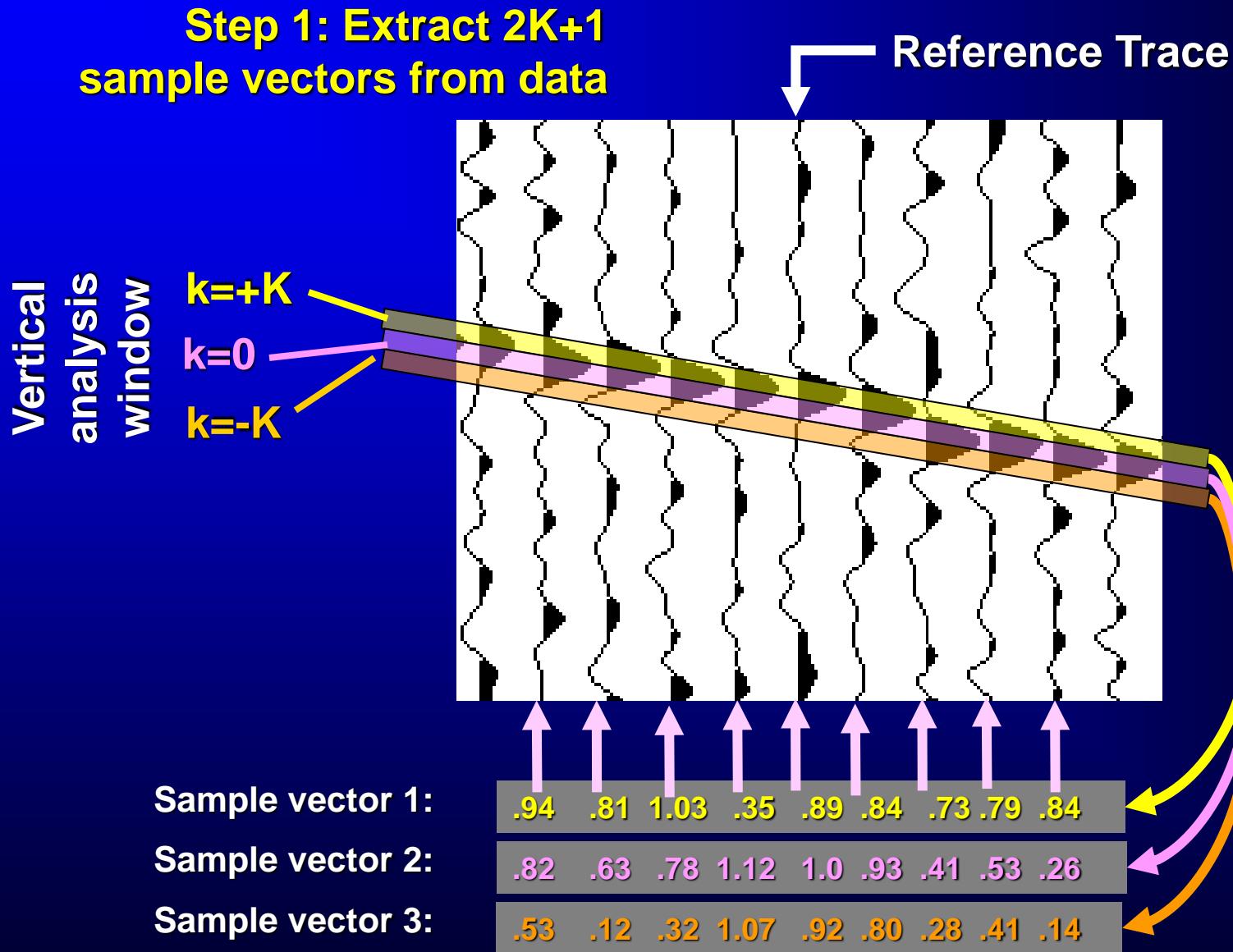


Eigenstructure coherence:

Time slice through ratio of coherent to total energy



Forming a covariance matrix



Forming a covariance matrix

Sample vector 1:

.94	.81	1.03	.35	.89	.84	.73	.79	.84
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Sample vector 2:

.82	.63	.78	1.12	1.0	.93	.41	.53	.26
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Sample vector 3:

.53	.12	.32	1.07	.92	.80	.28	.41	.14
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Step 2: Cross
Correlate each
column of the data
matrix with itself and
all other columns

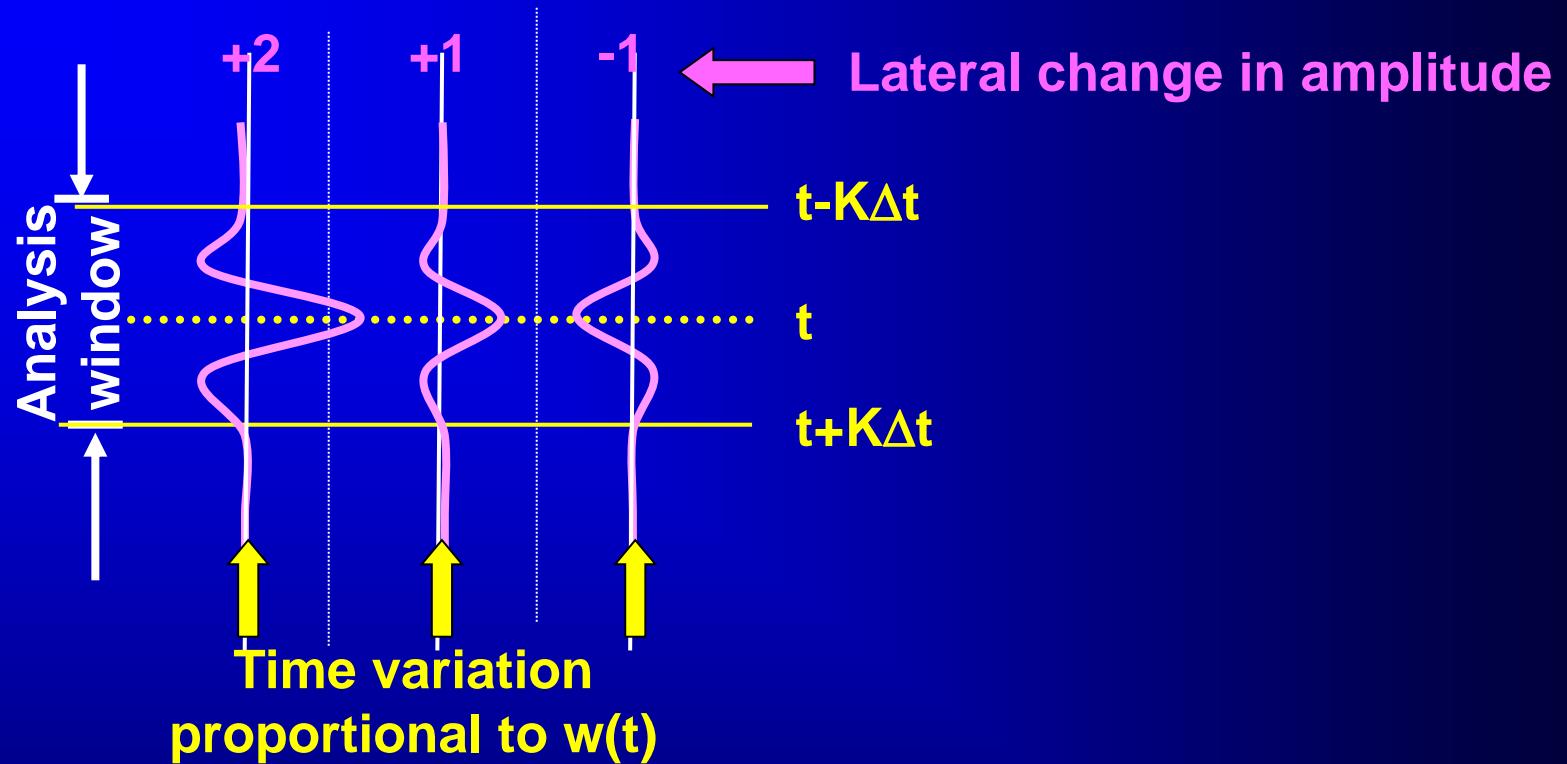


Step 3: Copy result
into corresponding
entry of the data
covariance matrix

$$\mathbf{C} = \begin{pmatrix} C_{11} & C_{12} & C_{13} & C_{14} & C_{15} & C_{16} & C_{17} & C_{18} & C_{19} \\ C_{21} & C_{22} & C_{23} & C_{24} & C_{25} & C_{26} & C_{27} & C_{28} & C_{29} \\ C_{31} & C_{32} & C_{33} & C_{34} & C_{35} & C_{36} & C_{37} & C_{38} & C_{39} \\ C_{41} & C_{42} & C_{43} & C_{44} & C_{45} & C_{46} & C_{47} & C_{48} & C_{49} \\ C_{51} & C_{52} & C_{53} & C_{54} & C_{55} & C_{56} & C_{57} & C_{58} & C_{59} \\ C_{61} & C_{62} & C_{63} & C_{64} & C_{65} & C_{66} & C_{67} & C_{68} & C_{69} \\ C_{71} & C_{72} & C_{73} & C_{74} & C_{75} & C_{76} & C_{77} & C_{78} & C_{79} \\ C_{81} & C_{82} & C_{83} & C_{84} & C_{85} & C_{86} & C_{87} & C_{88} & C_{89} \\ C_{91} & C_{92} & C_{93} & C_{94} & C_{95} & C_{96} & C_{97} & C_{98} & C_{99} \end{pmatrix}$$



Example of semblance coherence



$$c_s = \frac{\sum_{k=-K}^{+K} \left(\frac{1}{3} [2w(k\Delta t) + w(k\Delta t) - w(k\Delta t)] \right)^2}{\sum_{k=-K}^{+K} \frac{1}{3} ([+2w(k\Delta t)]^2 + [+w(k\Delta t)]^2 + [-w(k\Delta t)]^2)} = \frac{\left(\frac{4}{9} \right)}{\left(\frac{6}{3} \right)} = 0.33$$



Example of eigenstructure coherence

1. Form the 3x3 covariance matrix by cross-correlating each trace with itself and all other traces

$$\mathbf{C} = \sum_{k=-K}^{+K} \begin{pmatrix} (+2)w(k\Delta t)(+2)w(k\Delta t) & (+2)w(k\Delta t)(+1)w(k\Delta t) & (+2)w(k\Delta t)(-1)w(k\Delta t) \\ (+1)w(k\Delta t)(+2)w(k\Delta t) & (+1)w(k\Delta t)(+1)w(k\Delta t) & (+1)w(k\Delta t)(-1)w(k\Delta t) \\ (-1)w(k\Delta t)(+2)w(k\Delta t) & (-1)w(k\Delta t)(+1)w(k\Delta t) & (-1)w(k\Delta t)(-1)w(k\Delta t) \end{pmatrix}$$

Simplify to obtain

$$\mathbf{C} = \sum_{k=-K}^{+K} w^2(k\Delta t) \begin{pmatrix} +4 & +2 & -2 \\ +2 & +1 & -1 \\ -2 & -1 & +1 \end{pmatrix} \equiv E \begin{pmatrix} +4 & +2 & -2 \\ +2 & +1 & -1 \\ -2 & -1 & +1 \end{pmatrix}$$

where:

$$E \equiv \sum_{k=-K}^{+K} w^2(k\Delta t)$$



2. Guess at the first eigenvector, $\mathbf{v}^{(1)}$, that solves the equation:

$$\mathbf{C}\mathbf{v}^{(1)} = \lambda_1 \mathbf{v}^{(1)}$$

I claim $\mathbf{v}^{(1)}$ is proportional to the amplitude of the coherent part of the trace:

$$\mathbf{v}^{(1)} = \begin{pmatrix} +2 \\ +1 \\ -1 \end{pmatrix}$$

Let's test this claim:

$$E \begin{pmatrix} +4 & +2 & -2 \\ +2 & +1 & -1 \\ -2 & -1 & +1 \end{pmatrix} \begin{pmatrix} +2 \\ +1 \\ -1 \end{pmatrix} = E \begin{pmatrix} +12 \\ +6 \\ -6 \end{pmatrix} = 6E \begin{pmatrix} +2 \\ +1 \\ -1 \end{pmatrix} = \lambda_1 \begin{pmatrix} +2 \\ +1 \\ -1 \end{pmatrix}, \text{ which indicates: } \lambda_1 = 6E$$

To calculate coherence, we need the sum of the diagonal of the covariance matrix, \mathbf{C} :

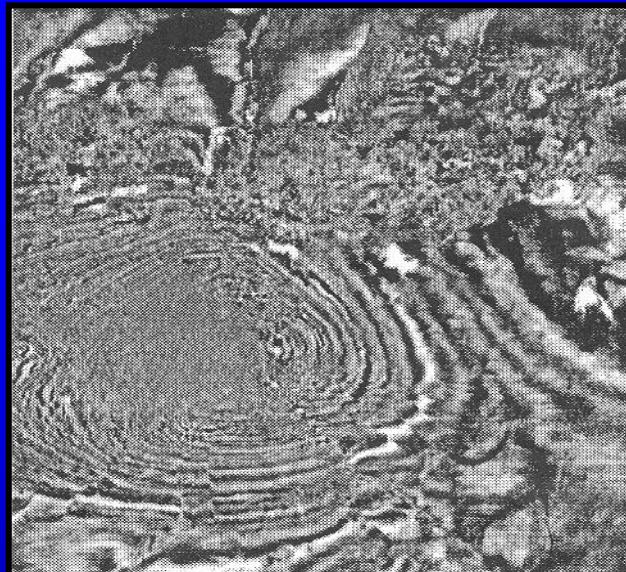
$$\sum_{j=1}^3 C_{jj} = C_{11} + C_{22} + C_{33} = E(4+1+1) = 6E$$

We can now form the eigenstructure estimate of coherence, c_e :

$$c_e \equiv \frac{\lambda_1}{\sum_{j=1}^3 C_{jj}} = \frac{6E}{6E} = 1$$



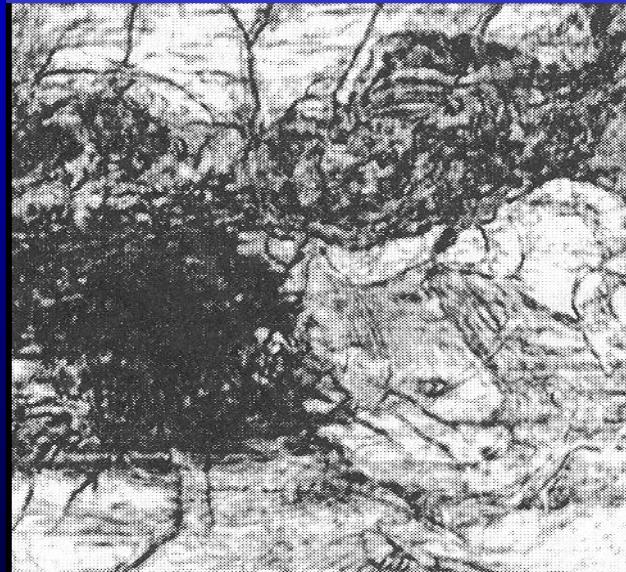
Coherence algorithm evolution



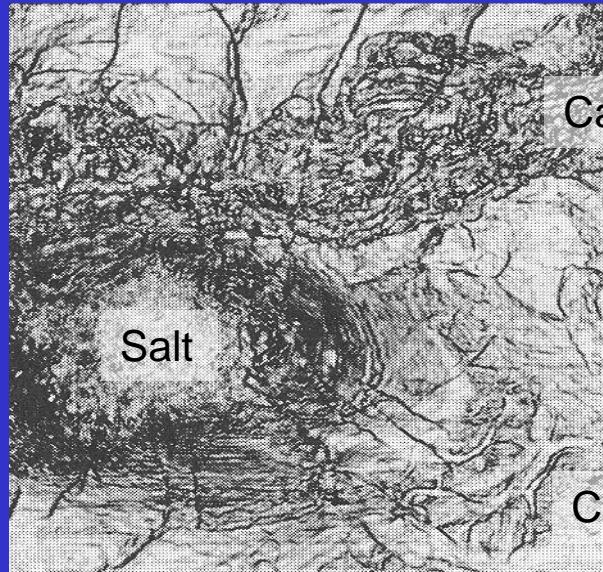
Seismic



Crosscorrelation



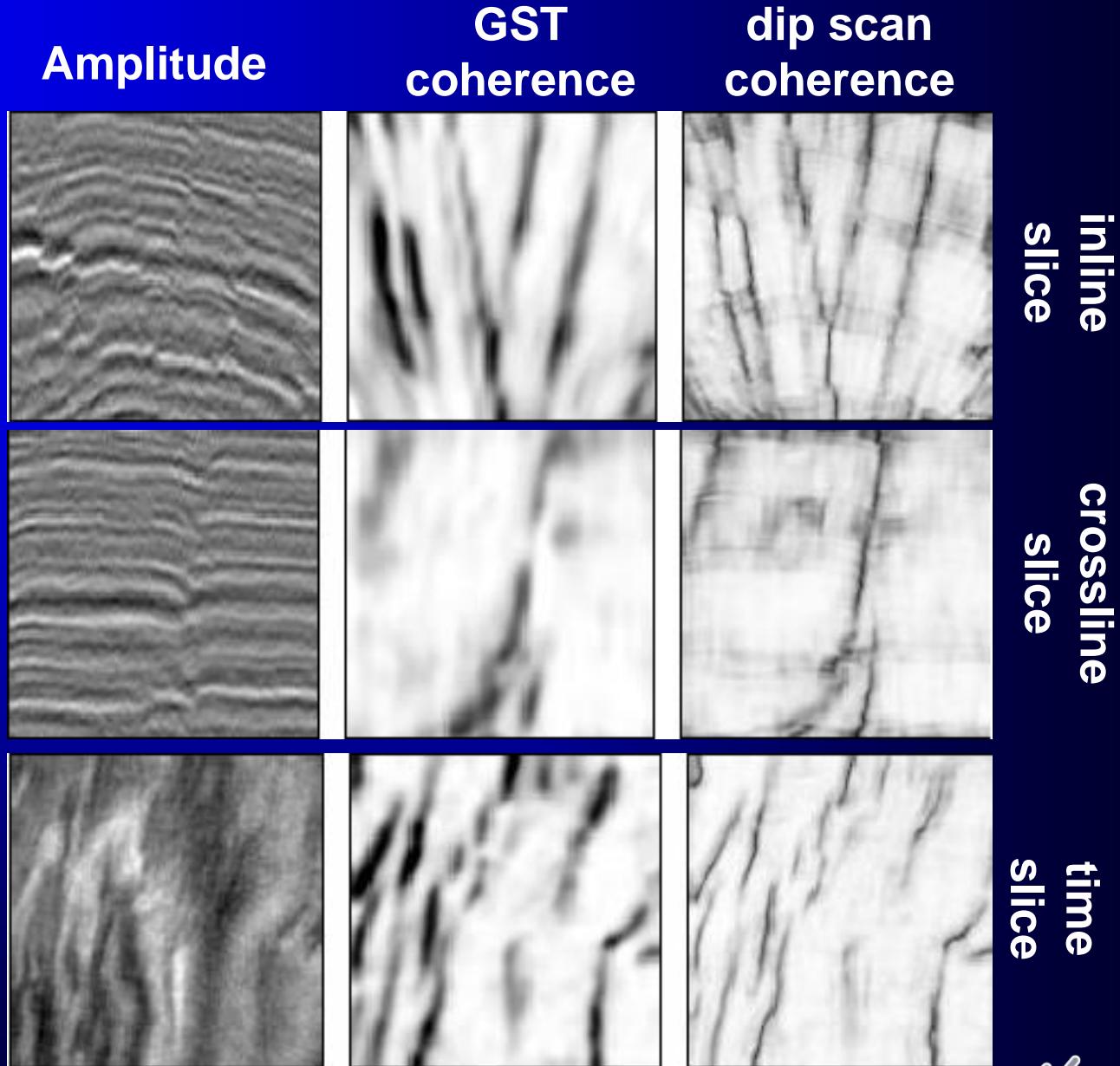
Semblance



Eigenstructure

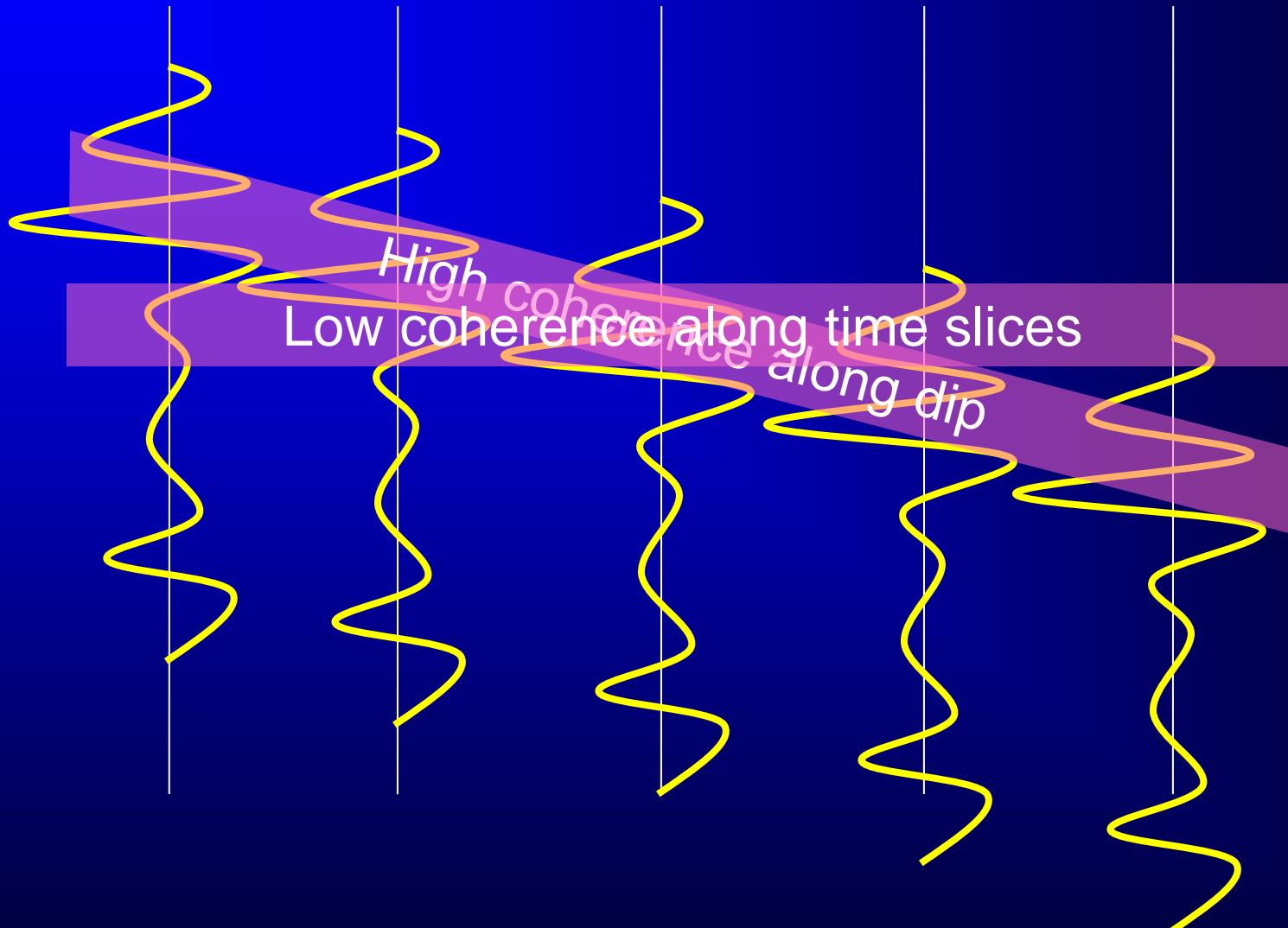


Comparison of Gradient Structure Tensor and dip scan eigenstructure coherence

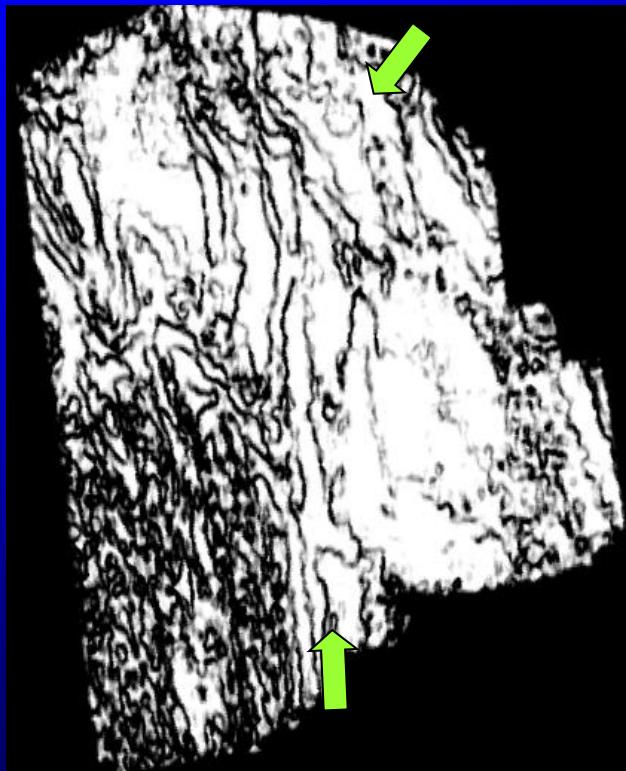


(Bakker, 2003)

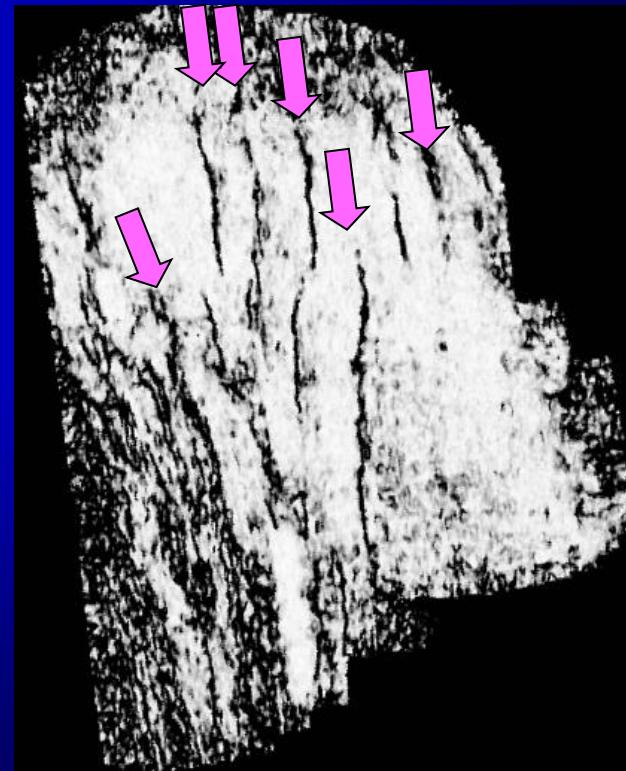
Importance of computing coherence along structural dip



Coherence artifacts due to an ‘efficient’ calculation without search for structure



Coherence computed
along a time slice



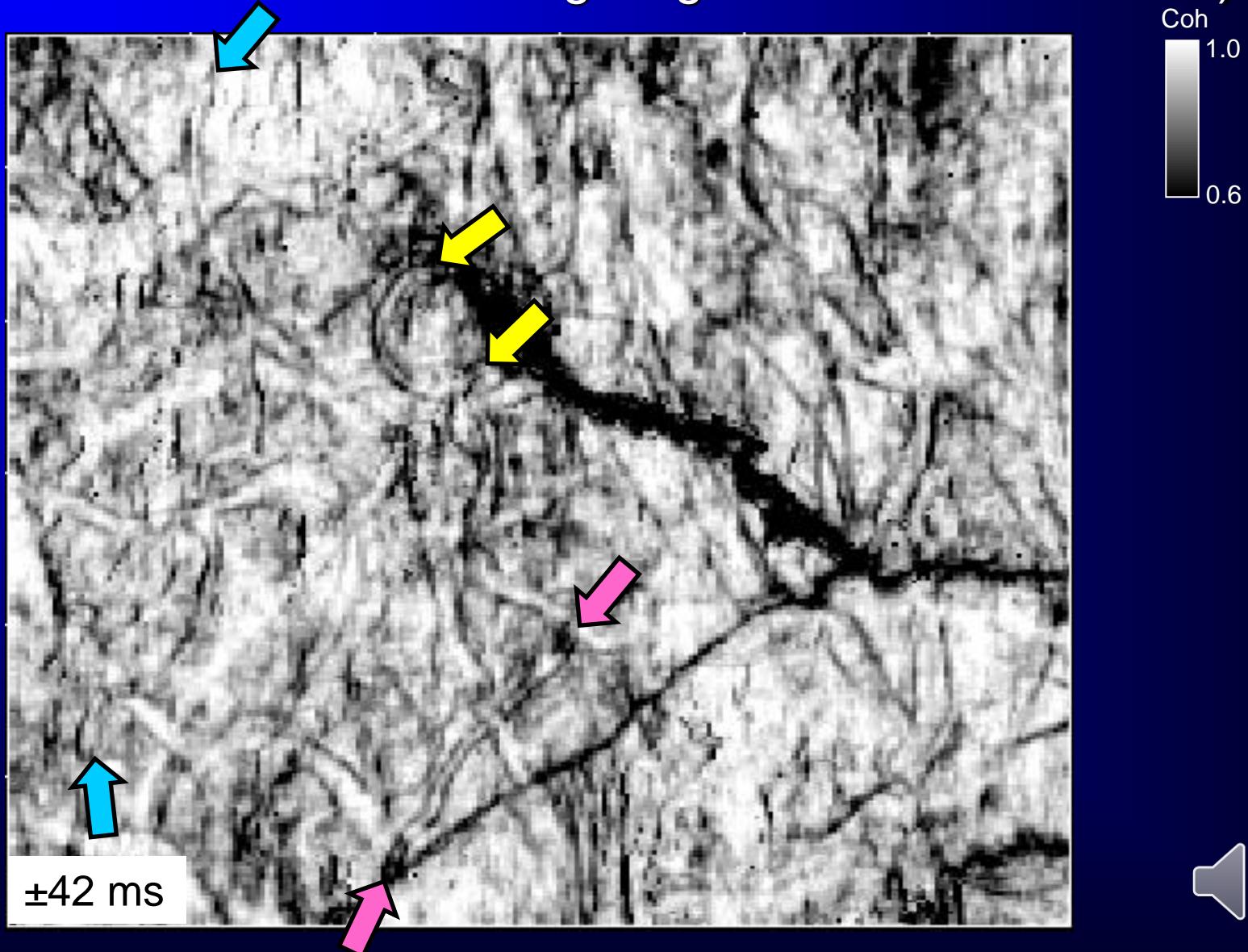
Coherence computed
along structure



(Chopra and Marfurt, 2008)

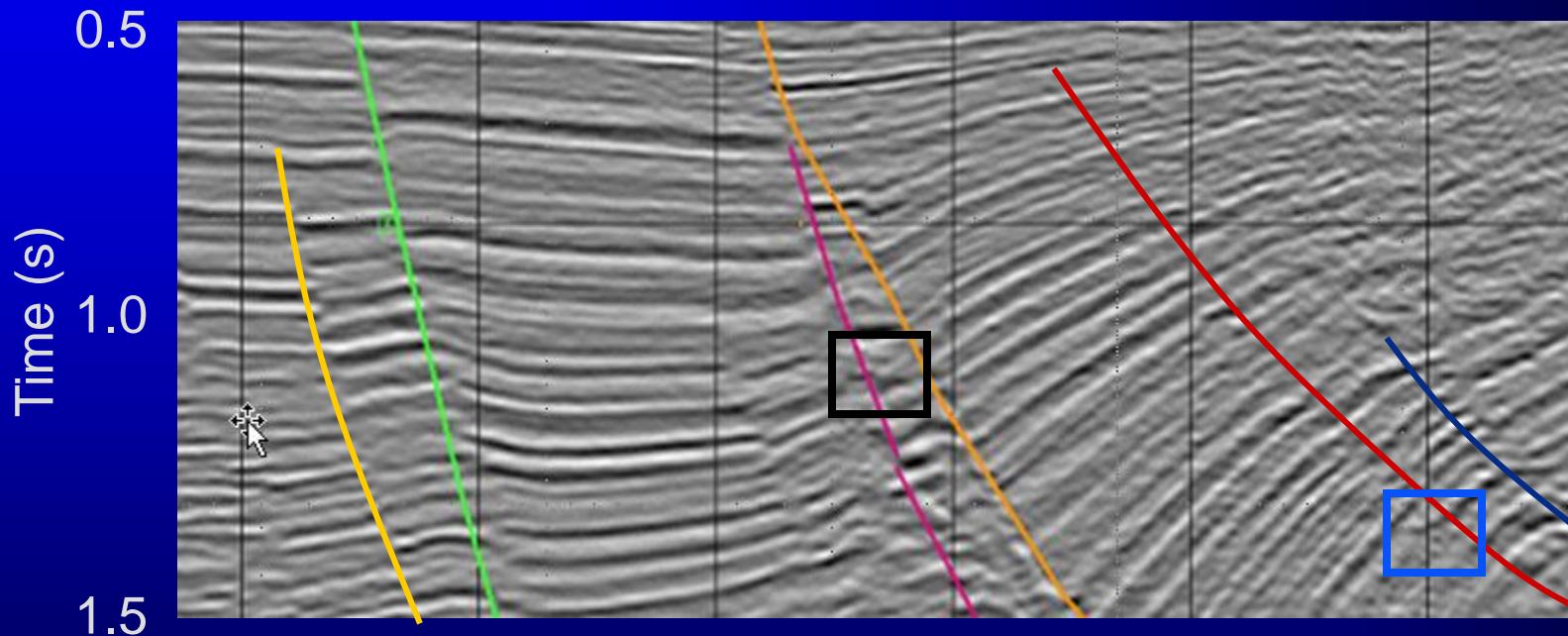
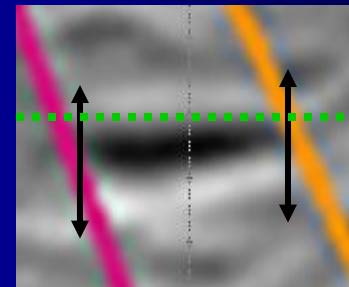
Impact of vertical analysis window

(phantom horizon slice through eigenstructure coherence)

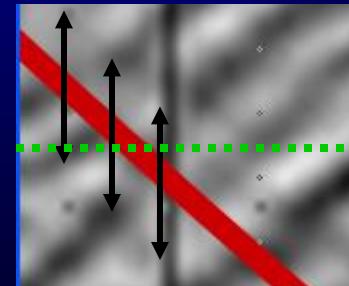


Impact of vertical analysis window

Fault on coherence green time slice is shifted by a stronger, deeper event



Steeply dipping faults will not only be smeared by long coherence windows, but may appear more than once!



Coherence

In summary, coherence:

- Is an excellent tool for delineating geological boundaries (faults, lateral stratigraphic contacts, etc.),
- Allows accelerated evaluation of large data sets,
- Provides quantitative estimate of fault/fracture presence,
- Often enhances stratigraphic information that is otherwise difficult to extract,
- Should always be calculated along dip – either through algorithm design or by first flattening the seismic volume to be analyzed, and
- Algorithms are local - Faults that have drag, are poorly migrated, or separate two similar reflectors, or otherwise do not appear locally to be discontinuous, will not show up on coherence volumes.

In general:

- Stratigraphic features are best analyzed on horizon slices,
- Structural features are best analyzed on time slices, and
- Large vertical analysis windows can improve the resolution of vertical faults, but smears dipping faults and mixes stratigraphic features.

