

# Elucidate long-term changes of ozone in Shanghai based on an integrated machine learning method

Jin Xue<sup>1,2</sup>, Fangting Wang<sup>1,2</sup>, Kun Zhang<sup>1,2</sup>, Hehe Zhai<sup>1,2</sup>, Dan Jin<sup>3</sup>, Yusen Duan (✉)<sup>3</sup>, Elly Yaluk<sup>1,2</sup>, Yangjun Wang<sup>1,2</sup>, Ling Huang<sup>1,2</sup>, Yuewu Li<sup>3</sup>, Thomas Lei<sup>4</sup>, Qingyan Fu<sup>3</sup>, Joshua S. Fu<sup>5</sup>, Li Li (✉)<sup>1,2</sup>

1 School of Environmental and Chemical Engineering, Shanghai University, Shanghai 200444, China

2 Key Laboratory of Organic Compound Pollution Control Engineering (MOE), Shanghai University, Shanghai 200444, China

3 Shanghai Environmental Monitoring Center, Shanghai 200235, China

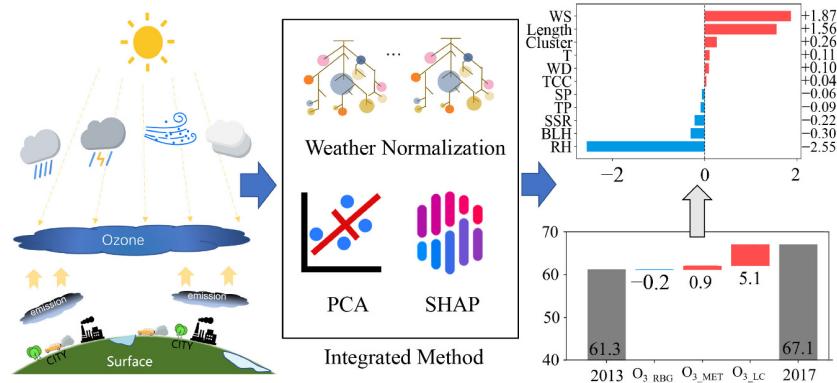
4 Institute of Science and Environment, University of Saint Joseph, Macao 999078, China

5 Department of Civil and Environmental Engineering, University of Tennessee, Knoxville, TN 37996, USA

## HIGHLIGHTS

- A novel integrated machine learning method to analyze O<sub>3</sub> changes is proposed.
- Various factors affecting long-term changes of O<sub>3</sub> in Shanghai are quantified.
- Meteorological, photochemical, and regional background O<sub>3</sub> are well separated.

## GRAPHIC ABSTRACT



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## ABSTRACT

Surface ozone (O<sub>3</sub>) is influenced by regional background and local photochemical formation under favorable meteorological conditions. Understanding the contribution of these factors to changes in O<sub>3</sub> is crucial to address the issue of O<sub>3</sub> pollution. In this study, we propose a novel integrated method that combines random forest, principal component analysis, and Shapley additive explanations to distinguish observed O<sub>3</sub> into meteorologically affected ozone (O<sub>3</sub>\_MET), chemically formed from local emissions (O<sub>3</sub>\_LC), and regional background ozone (O<sub>3</sub>\_RBG). Applied to three typical stations in Shanghai during the warm season from 2013 to 2021, the results indicate that O<sub>3</sub>\_RBG in Shanghai was 48.8 ± 0.3 ppb, accounting for 79.6%–89.4% at different sites, with an overall declining trend of 0.018 ppb/yr. O<sub>3</sub>\_LC at urban and regional sites ranged from 5.9–9.0 ppb and 8.9–14.6 ppb, respectively, which were significantly higher than the contributions of 2.5–7.4 ppb at an upwind background site. O<sub>3</sub>\_MET can be categorized into those affecting O<sub>3</sub> photochemical generation and those changing O<sub>3</sub> dispersion conditions, with absolute contributions to O<sub>3</sub> ranging from 13.4–19.0 ppb and 13.1–13.7 ppb, respectively. We found that the O<sub>3</sub> rebound in 2017, compared to 2013, was primarily influenced by unfavorable O<sub>3</sub> dispersion conditions and unbalanced emission reductions; while the O<sub>3</sub> decline in 2021, compared to 2017, was primarily influenced by overall favorable meteorological conditions and further emissions reduction. These findings highlight the challenge of understanding O<sub>3</sub> change due to meteorology and regional background, emphasizing the need for systematic interpretation of the different components of O<sub>3</sub>.

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## 1 Introduction

✉ Corresponding authors

duanys@sheemc.cn (Y. Duan); lily@shu.edu.cn (L. Li)

Surface ozone (O<sub>3</sub>) exerts significant impacts on air quality, climate, human health and ecosystems, displaying

deleterious effects (Lelieveld et al., 2015). As a secondary air pollutant, O<sub>3</sub> is formed via photochemical reactions among precursors under the presence of sunlight. Hence, meteorological variations play a vital role in the short- and long-term changes in O<sub>3</sub> (Wang et al., 2017). A quantitative analysis of the effects of meteorology and emissions is necessary to comprehend the fundamental relationship between O<sub>3</sub> and meteorology. Notably, when the effects of meteorological fluctuations are removed, variations in O<sub>3</sub> concentrations are governed by local photochemical reactions and background O<sub>3</sub>. The former from local emissions, whereas the latter is primarily formed from out-of-region emissions. Therefore, determining the influence of meteorology, emissions, and their specific factors on O<sub>3</sub> is fundamental to comprehend O<sub>3</sub> trends and evaluate the impact of emission control strategies.

Various methods are employed to estimate the contributions of meteorology and emissions to O<sub>3</sub>, including chemical transport models (CTMs), statistical methods, and machine learning (ML). CTMs involve manipulating emission inventories and meteorological fields to characterize the impact of emissions and meteorology on air quality, as demonstrated by previous studies (Li et al., 2019; Xu et al., 2019; Liu and Wang, 2020a, 2020b; Dang et al., 2021). However, CTMs methods have two main drawbacks. First, the computational resource requirements of numerical models, as well as the difficulty of ensuring the accuracy of emission inventories and meteorological fields, can pose challenges in terms of implementation and result in significant uncertainties. Second, determining the contribution of emissions and meteorology to O<sub>3</sub> requires extensive simulation scenarios that model the differences in O<sub>3</sub> concentrations resulting from interannual changes in emission inventories and meteorological fields across different years. This process can be arduous and time-consuming. Furthermore, this approach only allows for the overall contribution of meteorological conditions to O<sub>3</sub>, hindering our understanding of the impact of individual meteorological factors on O<sub>3</sub>. In contrast, statistical methods establish relationships between meteorological or emission data and air quality observations using multiple linear regression, empirical orthogonal functions, or Kolmogorov-Zurbenko (KZ) filtering (Camalier et al., 2007; Henneman et al., 2015; Yang et al., 2019; Li et al., 2020; Gao et al., 2021; Hu et al., 2021). These methods usually aim to fit or strip single influencing factors using mathematical equations. However, due to the nonlinearity between O<sub>3</sub> and its precursors, accurately capturing these nonlinear relationships can be difficult using statistics. Recently, the rapid rise of big data technology has led to the widespread application of ML methods for separating the effects of meteorology and emissions on air quality (Grange et al., 2018; Grange and Carslaw, 2019; Vu et al., 2019; Qu et al., 2020; Wang et al., 2020; Zhang

et al., 2020; Dai et al., 2021; Lovrić et al., 2021; Shi et al., 2021; González-Pardo et al., 2022; Hou et al., 2022; Tang et al., 2022; Wu et al., 2022; Zhou et al., 2022; Ding et al., 2023). For example, Grange et al. (2018) employed a random forest model to examine the trend of emissions on PM<sub>10</sub> in Switzerland from 1997 to 2016. Vu et al. (2019) evaluated the impact of a 5-year clean air action plan using the meteorological standardization technique. Qu et al. (2020) assessed the effectiveness of clean air policies on PM<sub>2.5</sub> reduction in China's "2+26" region using reinforcement tree models to remove meteorological effects. These ML methods have been demonstrated to effectively separate the effects of weather and emissions on air quality.

However, most of the existing research focuses on the overall impact of emissions or meteorology on O<sub>3</sub>, with less emphasis on systematically quantifying the impact of specific factors of emissions, meteorology, and background on O<sub>3</sub>. Recently, Lundberg et al. (2020) proposed a SHapley Additive exPlanation (SHAP) approach based on game-theoretic approach to explain ML models, which overcomes the black-box nature of ML methods and makes it easier to understand the attribution process of ML. The SHAP approach assigns predictions to each feature value, providing a clearer understanding of the degree of influence on the research target and quantifying the influence of each meteorological factor on O<sub>3</sub>. In addition to meteorological factors, it is also important to differentiate O<sub>3</sub> to local formation and regional background considering local and regional scale pollution control. Based on the characteristics of regional background O<sub>3</sub>, the dimensionality reduction property of principal component analysis (PCA) method can effectively isolate the regional background O<sub>3</sub> from the observed data (Suciu et al., 2017).

Here, we proposed a novel integrated method (IM) to quantify meteorological, emissions, and regional background contributions to changes of O<sub>3</sub> by integrating random forest (RF), PCA and SHAP. We applied the method to investigate changes of O<sub>3</sub> in Shanghai during the warm season (April–September) from year 2013 to 2021. During this period, we effectively stripped off the meteorological and chemical contribution to O<sub>3</sub>, assigned the meteorological contribution to each meteorological factor, and decomposed the chemically generated O<sub>3</sub> into local emission and regional background. Eventually, we systematically quantified the influence of each specific factor on O<sub>3</sub>. Our results demonstrate the reliability and provide strong scientific support for O<sub>3</sub> control.

## 2 Methodology

### 2.1 Data

For this study, three air quality monitoring stations located in Shanghai, China were selected. These stations

have continuous hourly observational data but diverse features. Pudong Huinan (PDHN, 31°3'N, 121°47'E), is an up-wind background site located in the eastern area of Shanghai and adjacent to the ocean, mainly influenced by regional transport. Putuo (PT, 31°14'N, 121°24'E) is an urban-center site strongly influenced by urban anthropogenic emissions such as vehicle exhaust, industries, and residential emissions, etc. Dianshan Lake (DSL, 31°5'N, 120°58'E) is a regional site located at the joint region of Jiangsu, Zhejiang provinces and Shanghai where O<sub>3</sub> accumulation occurs frequently. The reasons for selecting these three sites from multiple observational sites are detailed in Text S1, Figs. S1 and S2. Hourly observational data of air pollutants (NO<sub>x</sub> and O<sub>3</sub>) from 2013 to 2021 were collected from these three sites. O<sub>3</sub> concentration was monitored by an O<sub>3</sub> analyzer (Model 49i, Thermo Scientific, USA), and NO<sub>x</sub> concentration was monitored by an NO<sub>x</sub> analyzer (Model 42i, Thermo Scientific, USA). Outlier removal and data interpolation padding were performed on these data without changing the original data sample to meet the requirements of machine learning model training (Text S2).

Hourly reanalysis meteorological data for the years 2013–2021 were downloaded from the European Centre for Medium-Range Weather Forecasts (ECMWF). These data include 10 m u-component of wind (U10), 10 m v-component of wind (V10), 2 m dewpoint temperature (DT), 2 m air temperature (T), surface pressure (SP), surface net solar radiation (SSR), and total precipitation (TP) from ERA5-land with a spatial resolution of 0.1° × 0.1°. In addition, boundary layer height (BLH) and total cloud cover (TCC) from ERA5 were downloaded with a spatial resolution of 0.25° × 0.25°. To match the weather around the three sites above, all these meteorological variables were extracted from the grids covering these three stations in the gridded ERA5/ERA5-land data set. A comprehensive comparison between the ERA5/ERA5-land reanalysis data with surface observations (using temperature, dew point temperature, wind speed and, wind direction, at Pudong Airport and Hongqiao Airport) to verify the data reliability is presented and documented in detail in Text S3, Figs. S3 and S4.

The 72 h backward air mass trajectories arriving at an altitude of 100 m above ground level (a.g.l.) at PT, DSL and PDHN were calculated using the Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT 5.0.0) model (Stein et al., 2015). The HYSPLIT was driven by reanalysis meteorological data from the National Centers for Environmental Prediction (NCEP) and the National Center for Atmospheric Research (NCAR). We conducted trajectory clustering on the results generated by HYSPLIT, as depicted in Fig S5.

## 2.2 Trend analysis

The Mann-Kendall (MK) trend analysis is a useful tool

for detecting monotonic trends in time series data, which is widely employed in atmospheric sciences research (Hirsch et al., 1982; Pathakoti et al., 2021; Zhang et al., 2022). Compared to ordinary least squares (OLS), it is particularly appropriate when the data distribution is ambiguous or when the sample size is small. This is because the MK trend analysis does not make the assumption that error terms in the regression model are identically and independently distributed, which can lead to biased results. As a result, the MK trend analysis provides a more robust and reliable method for trend analysis in these scenarios. In this study, we employed a seasonal MK test method to determine the trend of O<sub>3</sub> while mitigating the influence of seasonal changes. A detailed explanation of the method can be found in the study of Hirsch et al. (1982). When the trend is statistically significant ( $p < 0.05$ ), a positive S-value indicates that the ozone is increasing (decreasing), and the magnitude of |S| is used to indicate the rate of increase (decrease) in the trend.

Furthermore, we also utilized the Theil-Sen estimator to examine the rate of O<sub>3</sub> increase/decrease. This method was proposed by Sen as a robust nonparametric statistical method for trend calculation (Sen, 1968). The method is impervious to measurement errors and outlier data, and it is suitable for long-term trend analysis. A detailed description of this method can be found in Zhang et al. (2022).

## 2.3 Random Forest (RF)-Based meteorological normalization

The RF model is a widely used ML model known for its high accuracy and stability. In this study, we employed RF to predict O<sub>3</sub> using time and meteorological variables. Specifically, the time variables, including Unix time, day of year, day of week, and hour, were used to represent the long-term, annual, weekly, and daily patterns of pollutant emissions, respectively. The meteorological variables used in the RF model included wind speed (WS, calculated from U10 and V10), wind direction (WD, calculated from U10 and V10), T, relative humidity (RH, calculated from T and DT, Alduchov and Eskridge, 1996), SSR, SP, TP, BLH, TCC, length of each air mass trajectory (Length) and cluster of each air mass trajectory (Cluster), to comprehensively consider the impact of meteorology on O<sub>3</sub>. The RF model was implemented in a Python environment, and underwent hyper-parameter selection to obtain the optimal model performance. The details of specific steps are described in Text S4, Figs. S6, and S7. The resulting RF model achieved an R<sup>2</sup> value ranging between 0.83 and 0.85, which is acceptable compared to previous studies (Vu et al., 2019; Huang et al., 2022; Li et al., 2022; Lu et al., 2022), and the other model performance characteristics are presented in Table S1.

The meteorological normalization method used in this study is similar to those utilized in previous studies (Grange et al., 2018; Wang et al., 2020; Zhang et al., 2020; Shi et al., 2021). This method predicts O<sub>3</sub> trends under historical average meteorological conditions. Specifically, we predict ozone by randomly sampling meteorological parameters from the historical meteorological data set without changing Unix time (the long-term pattern of pollutant emissions) to create a new data set for O<sub>3</sub> prediction. By making 1000 predictions in this manner, the predicted 1000 O<sub>3</sub> time series are averaged, resulting in a weather-normalized O<sub>3</sub> variation driven only by emissions, i.e., the changes in O<sub>3</sub> concentrations independent of meteorological effects.

#### 2.4 Principal component analysis (PCA)-based regional background O<sub>3</sub> estimation

PCA is a widely utilized method for dimensionality reduction in atmospheric environmental studies (Wren et al., 2020; Yang et al., 2020; Austin et al., 2021). Its principle lies in reducing high-dimensional data to a lower-dimensional space while retaining maximum information. This procedure facilitates the extraction of the main feature components of data, and it enables obtaining principal components that have the most significant impact on the data, thus enabling further data analysis and other subsequent operations (Jolliffe and Morgan, 1992; Zhu et al., 2021).

In this study, PCA was run and two components with eigenvalues greater than 1 were retained. Based on the relationship between component loadings and variables, we identified them as regional background ozone and others. PCA transforms data linearly, therefore, it is possible to calculate O<sub>3</sub>\_RBG from the linear relationship in PCA. We referred to previous studies' calculation of background concentration (Langford et al., 2009; Berlin et al., 2013; Suciu et al., 2017; Wang et al., 2022a) and input daytime (10:00–17:00) data, including O<sub>3</sub>, NO<sub>x</sub>, WS, WD and T, into the PCA model, to determine the corresponding components of O<sub>3</sub>\_RBG according to the scoring coefficient matrix. The detailed calculation steps are presented in Text S5. Using the results of PCA, we calculated the background concentration of O<sub>3</sub> by utilizing the average and standard deviation of O<sub>3</sub> observations, the score, and the variance contributions of corresponding components (see Tables S2 and S3).

#### 2.5 SHAP approach-based causal attribution

An explanatory model based on cooperative game theory, namely the SHAP approach (Lundberg and Lee, 2017; Song et al., 2022), is employed to attribute contributions from specific meteorological parameters to O<sub>3</sub>. This approach performs an algorithm on the ML model and the input features, resulting SHAP values that quantify the

impacts of each feature on the individual prediction results. The prediction results of the ML algorithm are the sum of the attributed values of each input feature (SHAP value). Therefore, using a ML model for predicting ozone (e.g., the RF model) enables calculation of the contribution of the input data (e.g., RH at 10:00 on July 1, 2013) to the predicted O<sub>3</sub>. The core algorithm is as follows (Eqs. (1) and (2)):

$$f(x) = \phi_0(f) + \sum_{i=1}^M \phi_i(f, x), \quad (1)$$

where  $f$  is the original model,  $\phi_0(f)$  is the baseline value of the model, which is the average of the model prediction results, and  $M$  is the number of model features.  $\phi_i(f, x)$  is the SHAP value of feature  $i$ , and it is the contribution of feature  $i$  to the prediction result  $f(x)$ . When  $\phi_i(f, x) > 0$  ( $\phi_i(f, x) < 0$ ), it indicates that feature  $i$  will increase (decrease)  $f(x)$ , and the absolute value of  $\phi_i(f, x)$  represents the importance of the feature  $i$ .

$$\phi_i(f, x) = \sum_{R \in \mathcal{R}} \frac{1}{M} [f_x(P_i^R \cup i) - f_x(P_i^R)], \quad (2)$$

where  $\mathcal{R}$  is the set of all feature orderings, and  $P_i^R$  is the set of all features that come before feature  $i$  in ordering  $\mathcal{R}$ .

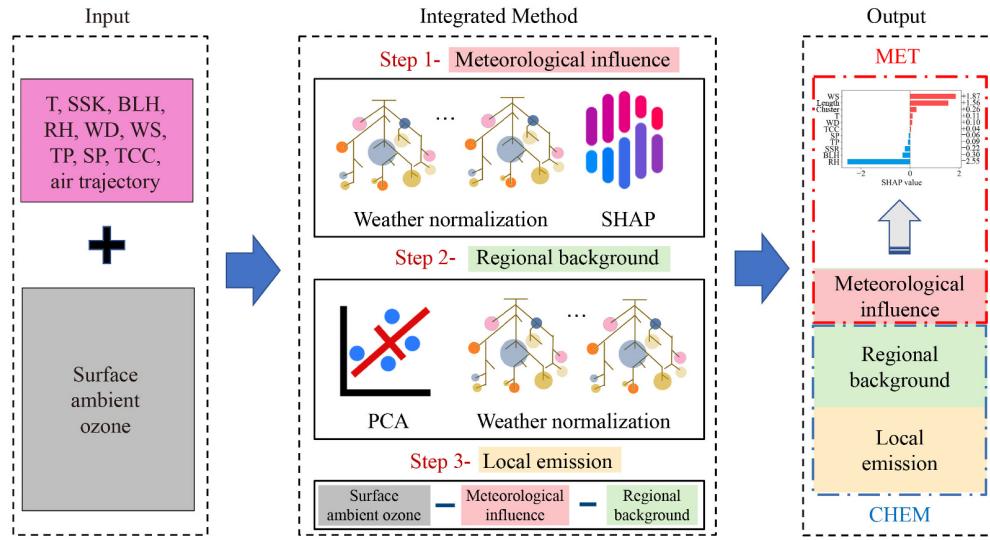
The SHAP value of each meteorological factor, including T, RH, SSR, WS, WD, BLH, TCC, SP, TP, Cluster, and Length, was determined based on the trained RF model and the SHAP method. The contribution of each meteorological component to O<sub>3</sub> was further analyzed.

#### 2.6 Integrated method

We propose an integrated method for identifying the contributions of meteorological, chemical / emissions and regional background factors to changes in O<sub>3</sub>, which includes three main steps. The overall framework of this IM is presented in Fig. 1.

**Step 1 (meteorology affected O<sub>3</sub>).** In this step, the meteorological normalization model is utilized to generate the time series of O<sub>3</sub> after meteorological normalization (denoted as "O<sub>3</sub>\_CHEM"). Then, the meteorologically normalized O<sub>3</sub> is subtracted from the observed O<sub>3</sub> to obtain the meteorological contribution to O<sub>3</sub> (denoted as "O<sub>3</sub>\_MET"). Next, the meteorological factors and the previously obtained meteorological contribution are used as the independent and dependent variables, respectively, in an RF model for training. Finally, utilizing the SHAP method, the SHAP values of each meteorological factor are calculated based on the trained RF model and the meteorological factors as input data. This step yields both the total meteorological contribution and the contribution of specific meteorological parameters.

**Step 2 (regional background O<sub>3</sub>).** The typical PCA



**Fig. 1** Framework of the Integrated Method.

approach for determining  $O_3$  is influenced by meteorology. To eliminate the meteorological influence and obtain the final meteorologically unaffected regional background  $O_3$  (denoted as “ $O_{3\_RBG}$ ”), we first compute the regional background  $O_3$  using the PCA model and then remove the meteorological impact using the meteorological normalization model from Step 1.

**Step 3 (local formation  $O_3$ ).** Surface  $O_3$  consists of the meteorological contribution, regional background  $O_3$ , and  $O_3$  generated from local emissions via photochemical reactions. In this step, we subtract the meteorological contribution (“ $O_{3\_MET}$ ”) and regional background ozone (“ $O_{3\_RBG}$ ”) from the observed ozone to obtain the ozone generated from local emissions (denoted as “ $O_{3\_LC}$ ”).

It should be noted that the relationship between  $O_3$  and its precursors is nonlinear and separating local, meteorological, and regional background factors from the perspective of precursor emissions would be complicated. However, our approach bypasses the complexity of the nonlinear relationship between  $O_3$  and its precursors by separating directly from the observed  $O_3$ , making the problem easier to resolve. Ultimately, this integrated method provides a systematic and quantitative separation of the observed  $O_3$  and clarifies the specific effects of meteorological factors, regional background, and local emissions on  $O_3$ .

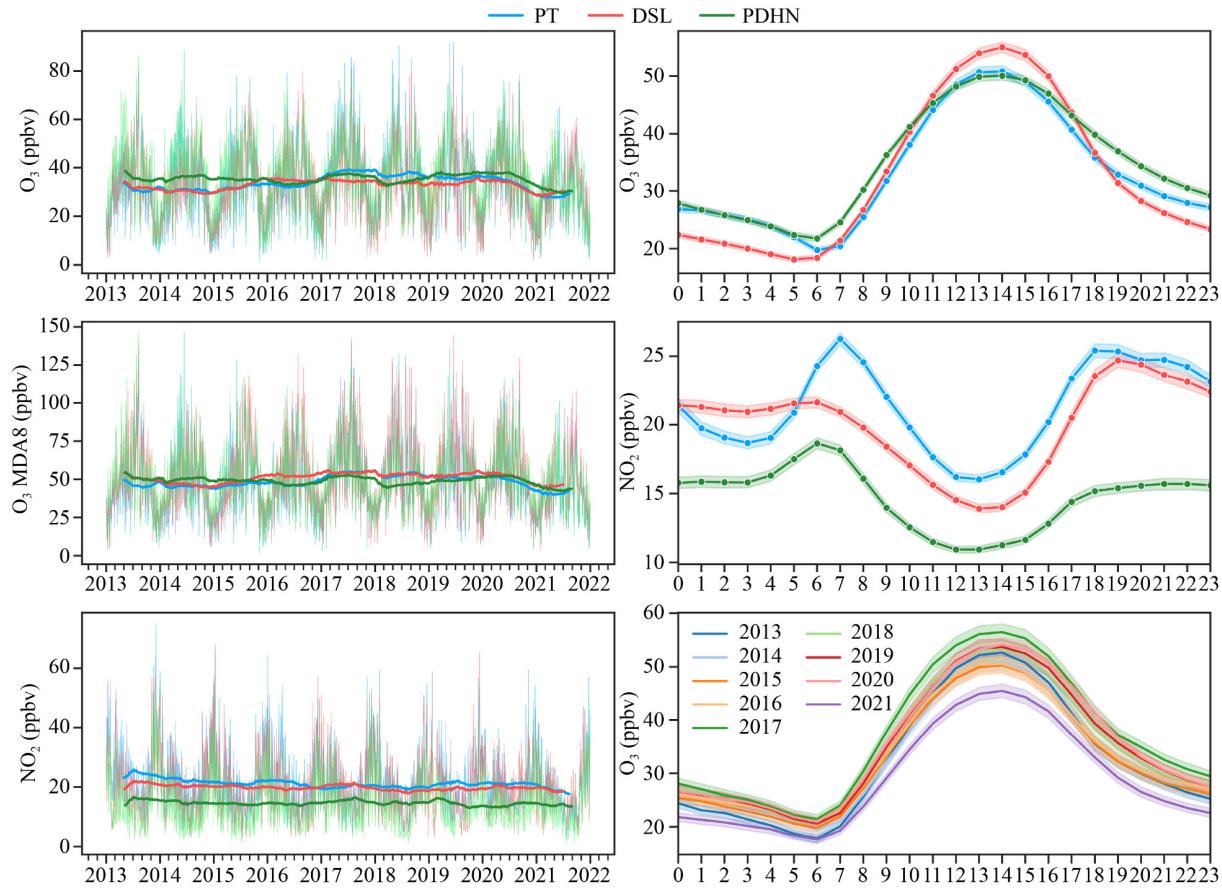
### 3 Results and discussion

#### 3.1 Trend of $O_3$ pollution

The long-term trends of  $O_3$  and nitrogen dioxide ( $NO_2$ ) at each site in Shanghai are presented in Fig. 2. The concentration of  $NO_2$  displayed a distinct pattern, with concentrations in descending order at PT, DSL, and

PDHN. From trend analysis, both PT and DSL showed a decreasing trend in  $NO_2$ , while PDHN did not show a clear trend. The decrease in  $NO_2$  at PT and DSL can be attributed to the implementation of  $NO_x$  emission control strategies in Shanghai and the Yangtze River Delta (YRD) region, respectively; while the low  $NO_2$  concentration at PDHN, which is located near the ocean and has a sparse population, may be due to the background. Although significant differences in  $NO_2$  concentrations were observed among the three stations, it was difficult to observe a clear trend in  $O_3$  concentrations over the long-term. Analyzing the diurnal variations of  $O_3$  at each site can provide a more reflective assessment of the  $O_3$  formation regime at the three stations. PT exhibited low daytime  $O_3$  concentration, significant diurnal variation, and sustained high  $NO_2$  concentration, indicating strong local generation of  $O_3$  and influence by  $NO$  titration. DSL had the highest  $O_3$  concentration during the day among the three sites, but  $NO_2$  was not prominent, and the diurnal variation of both  $NO_2$  and ozone was not significant, indicating that local photochemistry had a small contribution to ozone formation, and  $NO_x$  is relatively less limiting to  $O_3$ . The daytime  $NO_2$  concentration at PDHN was much lower than the other two sites, the  $O_3$  concentration was comparable to that at PT, indicating that the regional transport had a more significant impact on  $O_3$  around PDHN.

$O_3$  is primarily dominated by meteorology and photochemistry, therefore, meteorological conditions are crucial for  $O_3$  formation. We selected meteorological conditions for the warm season from April to September at 10–17 for analysis, as shown in Figs. 3 and 4. From the wind rose diagram, although the dominant wind during the warm season of PT, DSL and PDHN is south-east (SE), the wind aggregation and wind speed gradually



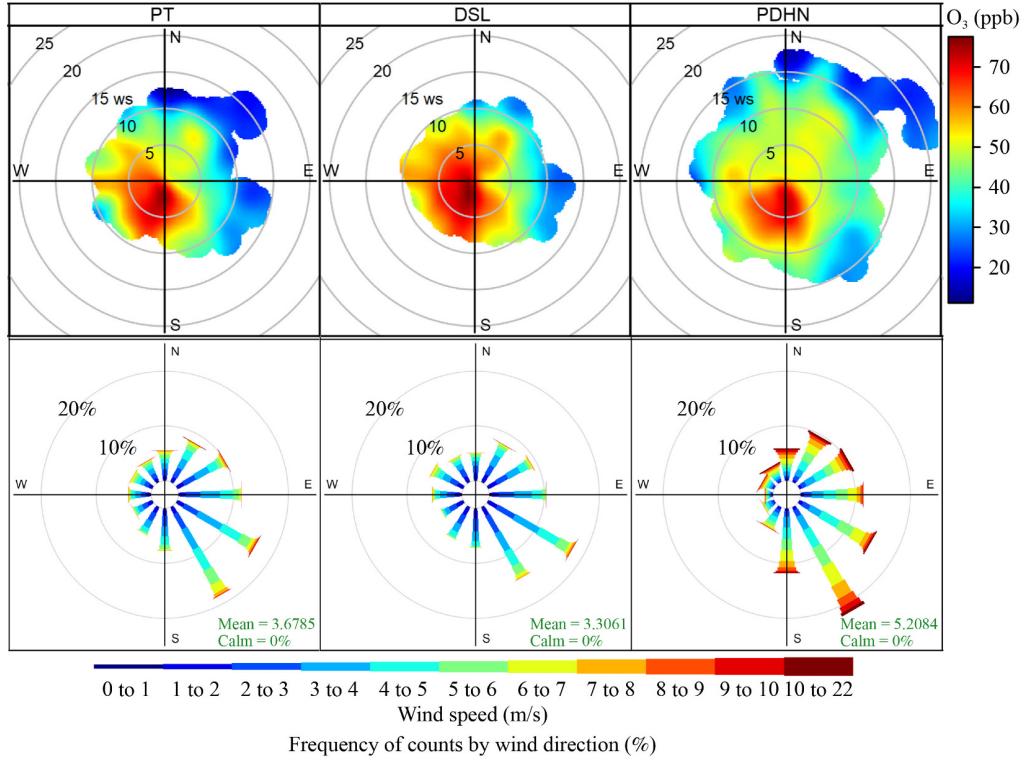
**Fig. 2** Time series of O<sub>3</sub>, maximum daily 8 hour average ozone (MDA8 O<sub>3</sub>) and NO<sub>2</sub> at PT (blue), DSL (red), PDHN (green). In the left figure, the thick line is the 360-day sliding average; the thin line is the daily average. In the right figure, the shaded area is the 95% confidence interval of corresponding data.

decrease from ocean to inland (Fig. 3). Therefore, PDHN is more susceptible to one-way wind, while the wind direction at DSL is more scattered. RH, T and SSR are considered the most important meteorological conditions that determine the O<sub>3</sub> photochemical processes. The analysis shows that RH at PDHN is significantly higher than the other two sites; T is significantly lower than the other two sites; and SSR is between PT and DSL, which may be related to the proximity of the PDHN site to the ocean. For the two inland sites, PT and DSL, the differences in RH and T are small, but DSL's SSR is significantly higher than PT. Regarding other parameters, PDHN's TP and TCC are significantly smaller than PT and DSL, and the difference in SP among the three sites is small. The BLH of DSL, PT, and PDHN decreases in turn. Overall, the dominant wind direction of PT is southeast, solar radiation is relatively small, and other meteorological conditions are mainly between the two stations. The wind direction of the DSL station is relatively scattered, with low RH, high T and high SSR, indicating that it is more favorable for ozone generation. The meteorological parameters at PDHN are significantly influenced by the ocean, leading to distinct variations in

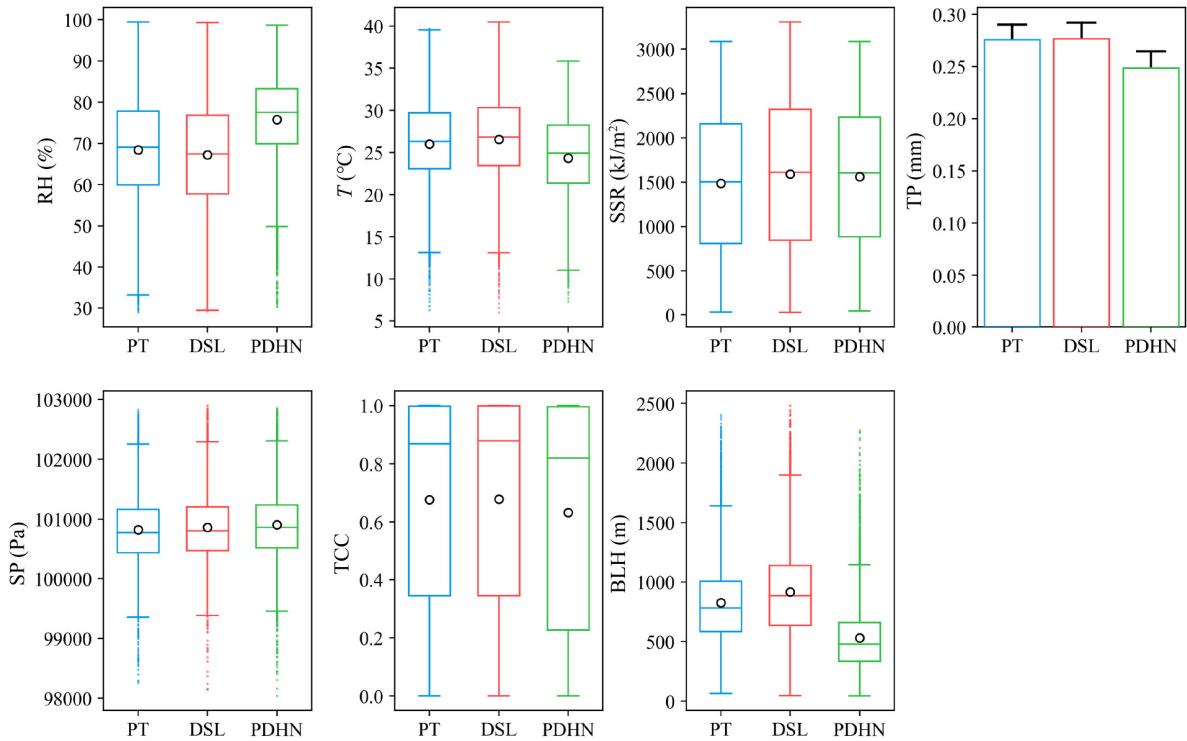
comparison to the other two stations. This influence is evident in the concentrated wind direction, high RH, and low temperature at PDHN. These factors contribute to the station's unique atmospheric conditions, setting it apart from the others.

### 3.2 Meteorological impact

The meteorological contributions to O<sub>3</sub> at various sites between 2013 and 2021 have been determined through the integrated method, as presented in Fig. 5, Text S6 and Figs. S8–S15. Meteorological conditions were found to have a positive impact on O<sub>3</sub> formation in 2013 (0.78–3.55 ppb) and 2017 (3.95–9.13 ppb), but unfavorable for O<sub>3</sub> formation in 2015 (−2.13 to −0.85 ppb), 2016 (−3.81 to −0.15 ppb) and 2021 (−8.86 to −4.41 ppb) at PT, DSL and PDHN stations, as depicted in Fig. 5. The remaining years displayed distinct features at each site. Although O<sub>3</sub><sub>MET</sub> displayed similar trends at the three locations in certain years, it varied significantly across most years due to their geographical locations, as illustrated in Fig. 5(a). Generally, the weather system affects O<sub>3</sub> in two ways: by altering T, RH, SSR, TCC and



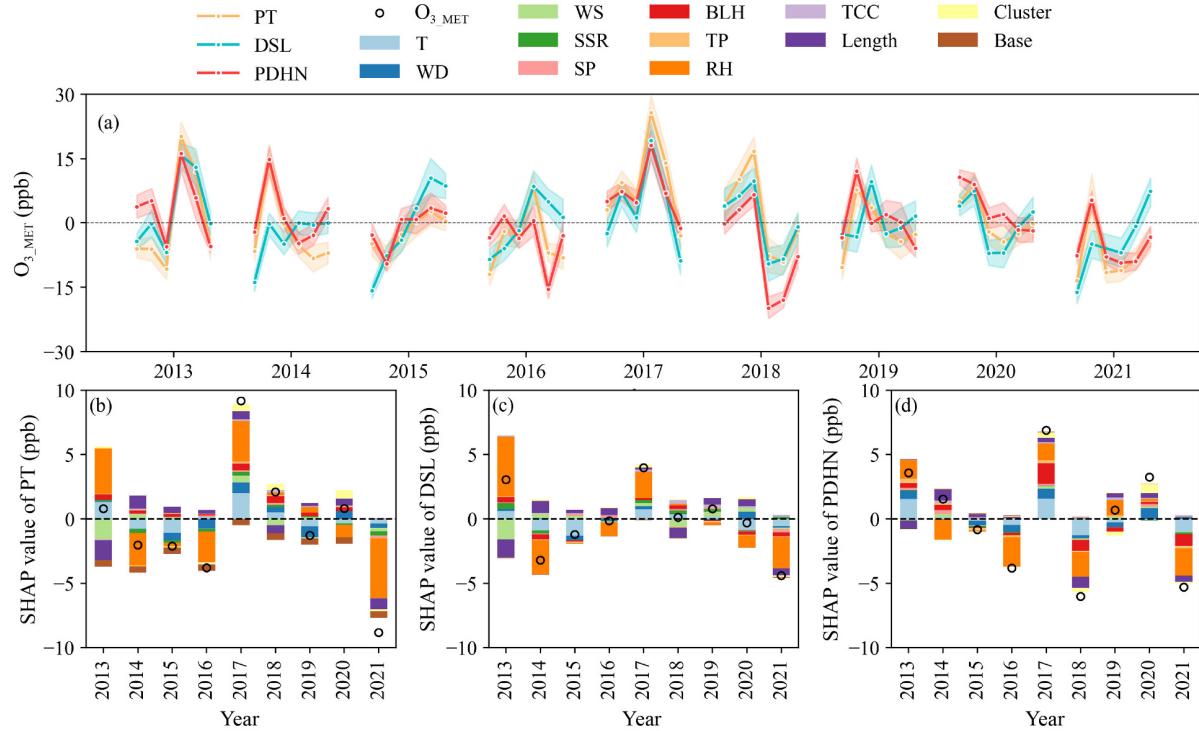
**Fig. 3** Bivariate polar plot and wind rose at PT, DSL and PDHN.



**Fig. 4** Comparisons of the major meteorological factors among the three observational sites.

TP to influence  $O_3$  photochemical generation (OPG), and by changing air mass trajectories, wind fields, BLH and SP to affect  $O_3$  dispersion conditions (ODC). The

cumulative absolute contributions to ozone ranged from 13.44–19.03 ppb and 13.18–13.66 ppb, respectively, as detailed in Table S4.



**Fig. 5** Meteorological contributions to O<sub>3</sub> at PT, DSL and PDHN. The bands beside the lines are the 95% confidence intervals in panel (a). Base is the base value of the SHAP approach, i.e. the average of the SHAP value in panels (b–d).

The overall trend of meteorological impact on O<sub>3</sub> is primarily through influencing OPG. For example, all PT, DSL and PDHN sites indicated favorable meteorological influences on O<sub>3</sub> photochemical formation in 2013 and 2017, as opposed to the unfavorable impact in 2021. Both RH and T contributed significantly to meteorological ozone, as illustrated in Figs. 5(b)–(d). In comparison, RH and T had a greater impact on O<sub>3</sub> levels at PT and DSL stations compared to PDHN stations. This discrepancy is likely due to the location of the PDHN station near the ocean, where RH and T are relatively stable, and anthropogenic activity levels are lower. In contrast, the higher levels of anthropogenic activities around PT and DSL stations, combined with their inland location, made them more susceptible to changes in RH and T that could affect the rate of ozone generation from NO<sub>x</sub> and volatile organic compounds (VOCs) emissions. Overall, our findings suggest that the impact of RH and T on O<sub>3</sub> levels is more significant at stations with high levels of anthropogenic activities, particularly in inland areas. Furthermore, this method effectively reflects the O<sub>3</sub> formation characteristics at different stations.

The impact of ODC at PT, DSL and PDHN varied significantly due to the different geographical locations of the stations. At the PT station, the cumulative absolute contribution of WD, Length, WS, and Cluster could reach 11.60 ppb (Table S4), signifying a clear influence of horizontal transport on O<sub>3</sub>. The concentration of O<sub>3</sub> at PT was affected by the arrival of inland air mass, clean air

mass from the ocean, and changes in wind direction (Fig. S13). At the DSL site, WS had an average absolute contribution of 4.85 ppb, which was significantly higher than the PT and PDHN stations (3.10 ppb and 0.69 ppb, respectively). The reason for this was that the DSL site was located at the conjunction area of the YRD region, where O<sub>3</sub> and its precursors from the three regions tended to converge around the site. As a result, the dilution effect and the diffusion effect of wind speed caused significant changes in O<sub>3</sub> (Fig. S13). The cumulative average absolute contribution of all meteorological parameters characterizing O<sub>3</sub> transport at the PDHN site was 12.51 ppb, with BLH accounting for 43.0%. This indicated that the effects of transport, especially vertical transport, had a significant impact on the local variability of O<sub>3</sub> at PDHN. As an up-wind background station, it has little anthropogenic emissions, resulting less *in situ* chemical reactions producing O<sub>3</sub>. However, the low concentration of NO<sub>x</sub> made it difficult to titrate off the O<sub>3</sub> transported from outside the station. Figure S14 showed a clear difference between the dependence plot of BLH at PDHN and other stations. The SHAP value of BLH at PDHN station decreased with the increase of T. This was probably because the increase of T strengthened the convective intensity within the troposphere, causing the BLH to increase and further strengthening the O<sub>3</sub> diffusion ability at PDHN station. This was consistent with the study of Tang et al. (2021) and a more detailed explanation could be found in Text S6.

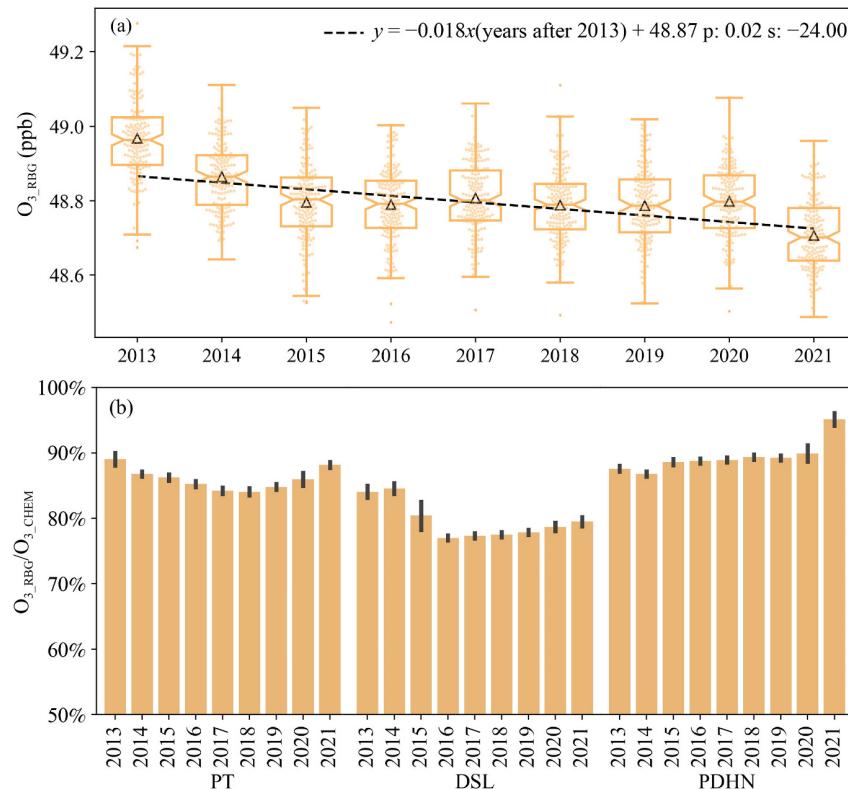
The results indicated that factors such as RH, T, and SSR, which affect OPG, determined the trend of the  $O_3$  MET, whereas the influence of ODC on  $O_3$  fully reflected the respective characteristics of PT, DSL, and PDHN sites. The meteorological impacts analyzed in this study were comparable to those reported by Lin et al. (2021). Meanwhile, Liu and Wang (2020a) revealed that weather conditions reduced ozone generation during 2014–2016, which was consistent with the meteorological findings at PT and DSL sites. Furthermore, Li et al. (2021b) investigated the daytime  $O_3$  pollution process in Shanghai and found that the contribution of downward  $O_3$  transport in the lower free troposphere at offshore sites could reach 49.1%. This indicated that BLH had a significant influence on ozone in PDHN, which was consistent with the results obtained in this study.

### 3.3 Regional background contribution

The concentration of  $O_3$  RBG was estimated to be  $48.8 \pm 0.3$  ppb, with a declining trend of  $-0.018$  ppb/yr ( $p = 0.02$ ), as presented in Fig. 6. This trend could be attributed to the stringent pollution control measures implemented in China between 2013 and 2020, which

primarily targeted the reduction of  $NO_x$  and VOCs, the precursors of  $O_3$ . The reduction rate of  $NO_x$  was higher than that of VOCs, which could be disadvantageous for  $O_3$  mitigation in urban areas but beneficial in rural and regional areas (Li et al., 2019). Thus, the minor decline in  $O_3$  RBG could be explained by the decrease in  $O_3$  in rural regions and the complexity of the regional background  $O_3$  components. The proportion of  $O_3$  RBG in  $O_3$  varied among the PT, DSL, and PDHN sites, with values of 86.4% (84.0% to 90.5%), 79.6% (76.7% to 85.4%), and 89.4% (86.6% to 96.0%), respectively. The contribution of  $O_3$  RBG at PT station increased and then decreased, indicating a change in the local  $O_3$  concentration in the opposite direction. The  $O_3$  RBG at DSL was lower than that at PT station, which could be attributed to favorable photochemical conditions at the DSL station. The contribution of  $O_3$  RBG at PDHN station had the highest percentage among the three stations, which was reasonable considering its regional location.

Our findings are overall consistent with previous studies. For example, Sahu et al. (2021) used the CMAQv5.2 model and found that the policy relevant background concentration of  $O_3$  in the Eastern China region accounted for 79% of the total  $O_3$ , primarily

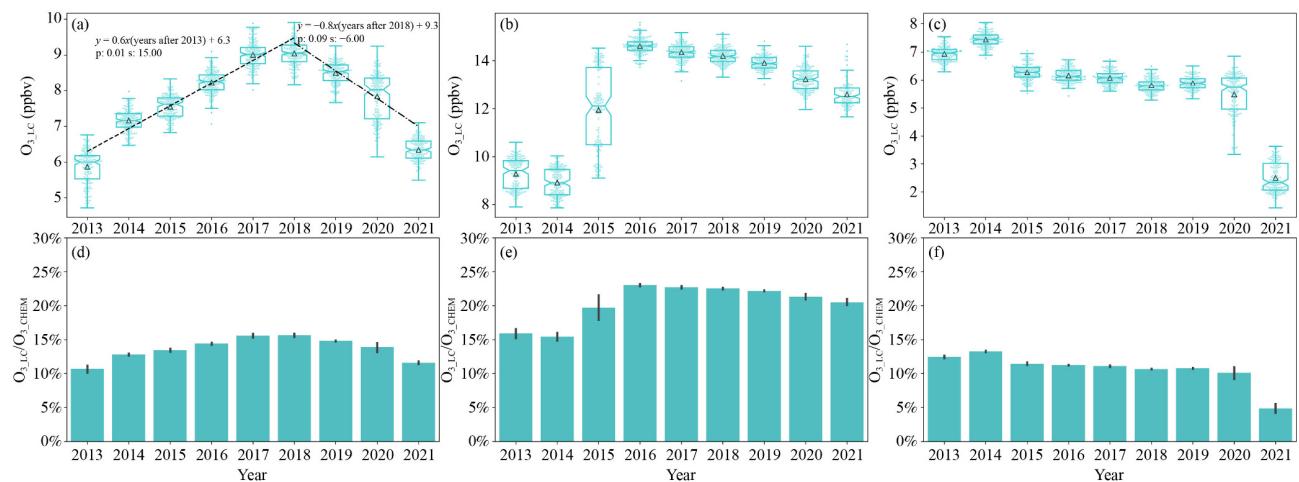


**Fig. 6** Contributions of  $O_3$  RBG at PT, DSL and PDHN. In panel (a), the upper and lower boundaries of the box represent the 25th percentile and 75th percentile, and the notches represent the confidence interval around the median. The line and hollow triangles inside boxes represent the average and median values of corresponding data, respectively. The error lines above and below the box represent the 25th percentile value minus  $1.5 * IQR$  (interquartile range) and the 75th percentile value plus  $1.5 * IQR$  (interquartile range), respectively. In panel (b), the error line indicates the 95% confidence interval of the corresponding data.

originating from South-east Asia and regions outside of Asia. Lu et al. (2019) conducted a study using the GEOS-Chem model and found that the average background concentration of  $O_3$  for March to October in the YRD region was about 41 ppb, with biogenic volatile organic compounds (BVOCs) being the primary contributors from natural sources. Ni et al. (2018) used the GEOS-Chem model to investigate the contribution of domestic and foreign  $O_3$  to China during the spring season and found that the background concentration of  $O_3$  in the Shanghai region was approximately 40–45 ppb. However, it is worth noting that the regional background  $O_3$  obtained from the aforementioned model studies was derived by simulating the shutdown of anthropogenic emissions solely within China. This approach failed to account for the contribution of  $O_3$  pollution transport from other regions to Eastern China and the YRD region. As a result, our findings indicate a higher regional background concentration than those reported in these studies. In addition to modeling approaches, some statistical methods have also been used to calculate the regional background concentration of  $O_3$ . Chen et al. (2022) established a relationship between surface  $O_3$  and temperature and obtained a regional background concentration of  $O_3$  in the YRD from 2013 to 2019 of  $53 \pm 13$  ppb. Suciu et al. (2017) used the same PCA method as our study to analyze the Houston-Galveston-Brazoria (TX) region for regional background  $O_3$ , resulting in a value of  $46.72 \pm 2.08$  ppb. Similarly, Wang et al. (2022b) used a similar PCA method to calculate regional background  $O_3$  concentrations of 51.0–52.1 ppb in Shandong, China. Overall, our results regarding the regional background  $O_3$  in Shanghai are reasonable.

### 3.4 Local emission contribution

**Figure 7** illustrates the variation of  $O_{3\_LC}$  at PT, DSL, and PDHN.  $O_{3\_LC}$  shows significant differences among these stations. Specifically, DSL has the highest local chemical contribution, ranging from 8.9–14.6 ppb (15.4%–23.0%), followed by PT and PDHN stations with concentrations of 5.9–9.0 ppb (10.7%–15.6%) and 2.5–7.4 ppb (4.6%–13.2%), respectively. Interestingly,  $O_{3\_LC}$  began to increase at the urban PT site at a rate of 0.6 ppb/yr since 2013, reaching 9.2 ppb in 2018. In contrast,  $NO_2$  decreased overall at a rate of 0.91 ppb/yr at PT during 2013–2018 (see Fig. S16).  $O_{3\_LC}$  then decreased at a rate of 0.8 ppb/yr to 6.5 ppb during 2018–2021, while  $NO_2$  continued to decrease at a rate of about 0.47 ppb/yr. This pattern of an increase followed by a decline in  $O_{3\_LC}$ , while  $NO_2$  levels are generally decreasing, may reveal the changes of photochemical regime related to  $O_3$  sensitivity. It is worth mentioning that China implemented the Air Pollution Prevention and Control Action Plan (APPCAP) during 2013–2017, which mainly targeted primary pollutants including  $NO_x$  and  $SO_2$ . Subsequently during 2018–2021, China implemented the Blue Sky Protection Campaign (BSPC), which continued the previous measures and strengthened VOCs control.  $O_3$  around the urban PT area continued to increase between 2013 and 2018, possibly related to Shanghai being under the VOC-limited regime. However, the decrease in  $O_3$  after 2018 may be due to the continuous reduction of  $NO_x$  making Shanghai shift from VOC-limited regime to a transition regime or  $NO_x$ -limited regime. Similar findings were also reported by previous studies (Xu et al., 2019; Mousavinezhad et al., 2021; Wang et al., 2022b). Notably, reducing VOCs mitigates



**Fig. 7** Contributions of  $O_{3\_LC}$  at (a) PT, (b) DSL and (c) PDHN. In panels (a–c), the upper and lower boundaries of the box represent the 25th percentile and 75th percentile, and the notches represent the confidence interval around the median. The line and hollow triangles inside boxes represent the average and median values of corresponding data, respectively. The error lines above and below the box represent the 25th percentile value minus  $1.5 * IQR$  (interquartile range) and the 75th percentile value plus  $1.5 * IQR$  (interquartile range), respectively. In panels (d–f), the error line indicates the 95% confidence interval of the corresponding data.

$O_3$ , irrespective of the  $O_3$  formation regime. Regarding DSL,  $O_{3\_LC}$  similarly rises and subsequently falls, but the rising trend is not monotonic compared to that at PT. During this period,  $NO_2$  decreased with a fluctuating speed of approximately 0.16 ppb/yr at DSL, which was much lower than the decrease rate observed at the urban station PT. This difference may be due to DSL being located at the border between Shanghai and other cities, which is influenced by the pollution reduction measures of both Shanghai and the YRD region. Hence, it is inconsistent with urban Shanghai. As for PDHN,  $O_{3\_LC}$  exhibits a fluctuating trend ranging between 5.5–6.9 ppb, dropping to only 2.5 ppb in 2021 due to a large decline in  $NO_x$  (Fig. S16). Looking at the period from 2013 to 2021,  $NO_2$  decreased overall at a rate of 0.38 ppb/yr at PDHN, which is also lower than the decline rate observed at the urban site PT. This phenomenon may be attributed to PDHN's location in an area of Shanghai with relatively low human activity, as illustrated by the significantly lower local  $NO_2$  concentrations compared to the other two stations (Fig. S16), and this could be the reason for PDHN's continuing stability of  $O_3$ .

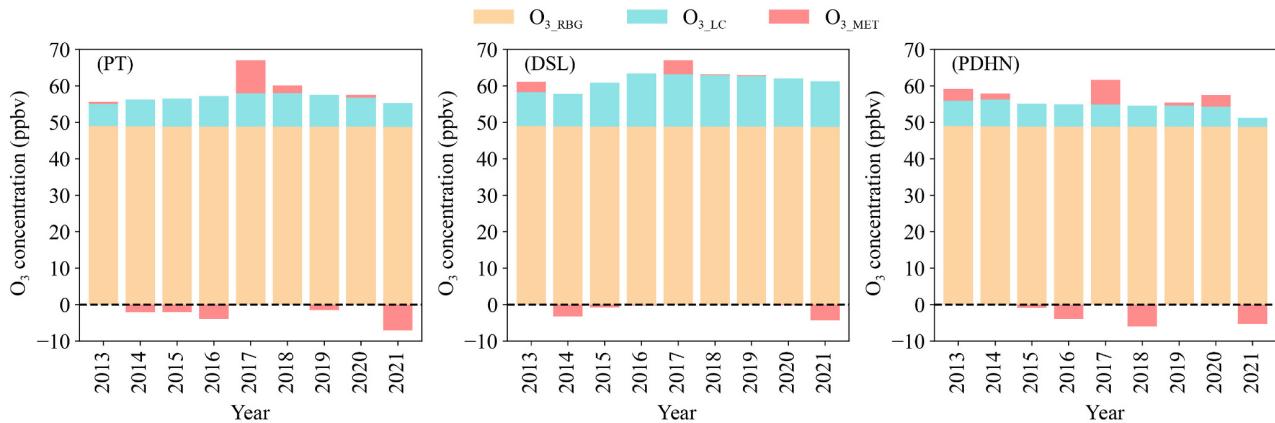
### 3.5 Investigation into $O_3$ changes in typical years

We quantified the numerous contributing elements of the observed  $O_3$  after unifying the data based on the integrated method, shown in Fig. 8. Based on the findings of the integrated method, we selected the years 2013, 2017, and 2021 as representative years to investigate changes in ozone-influencing variables at three representative locations in Shanghai.

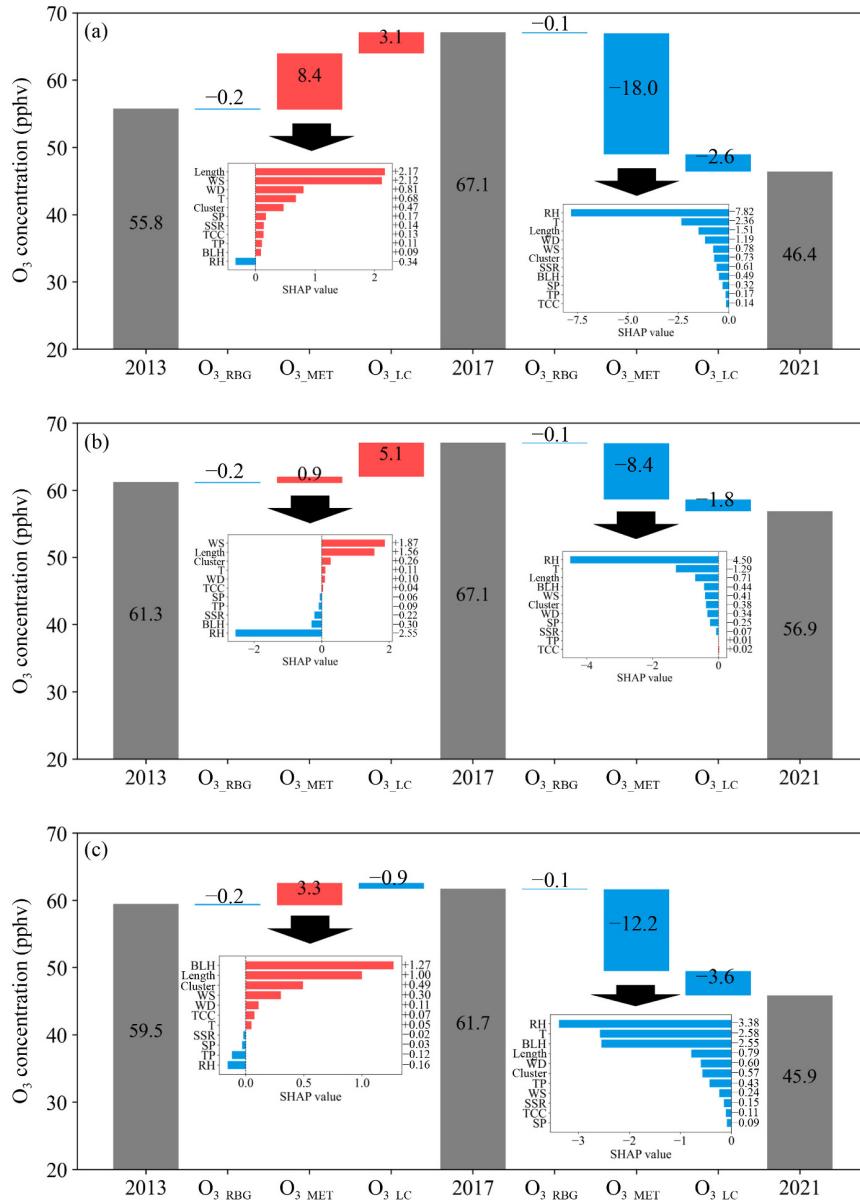
During the implementation of the APPCAP policy between 2013 and 2017, the concentrations of daytime  $O_3$  at the PT, DSL, and PDHN stations exhibited a rebound of 11.3 ppb (20.3%), 5.8 ppb (9.5%), and 2.2 ppb (3.6%), respectively, in 2017 compared to 2013. This trend suggests that the  $O_3$  rebound was more significant in urban areas than in regional or upwind areas. The changes

in  $O_{3\_LC}$  and  $O_{3\_MET}$  were the primary causes of the observed  $O_3$  rebound, as shown in Fig. 9. Apart from a slight decrease in  $O_{3\_LC}$  at PDHN, there was a substantial rebound in  $O_{3\_LC}$  of approximately 3.1 ppb and 5.1 ppb at PT and DSL, respectively, accounting for 27.4% and 87.9% of the total  $O_3$  rebound. This finding clearly indicates that the air pollution control measures implemented under the APPCAP policy were unfavorable for  $O_3$  mitigation in urban areas. Actually, since ozone formation around PT and DSL usually falls into the VOC-limited regime (Li et al., 2021a), the emission reduction measures caused a quick drop in ambient  $NO_x$  (see Fig. S16), leading to a weakened titration effect and hence the rebound of local ozone. Moreover, under high VOCs conditions, carbonyl compounds produced by VOCs oxidation, especially formaldehyde, can accelerate  $O_3$  production and VOCs oxidation (Li et al., 2021a). In addition, the decrease in  $PM_{2.5}$  slows down the uptake of aerosol particles on hydroperoxy radicals, thereby accelerating  $O_3$  generation (Li et al., 2019). These factors may have contributed to the increase in  $O_3$  levels. However, at the PDHN station, which is an upwind background station limited by  $NO_x$ , the reduction in  $NO_x$  was beneficial for  $O_3$  mitigation, explaining the drop in  $O_{3\_LC}$ .

For  $O_{3\_MET}$ , all the three stations exhibited rebound during the APPCAP period, although the magnitude of the rebound differed among the three sites due to the differences in location. Notably, the increase in meteorological  $O_3$  at all three sites was attributed to meteorological factors characterizing ODC, such as WS, Length, WD, Cluster, BLH and SP. We compare the variations in these factors between 2013 and 2017 (Tables S5–S7, Figs. S17 and S18) and found that the air mass trajectories from the ocean or primarily passing through the ocean (e.g., C7 and C11 at PT, C8 and C11 at DSL, C3 and C11 at PDHN) decreased significantly in 2017 compared to 2013, whereas air masses from inland, particularly south and west of Shanghai (e.g., C5 at PT,



**Fig. 8** Quantification of influencing factors on  $O_3$  changes during the warm season from 2013 to 2021 in Shanghai.



**Fig. 9** Variation of influencing factors on O<sub>3</sub> for typical years at PT (a), DSL (b) and PDHN (c).

C4 at DSL, C8 and C10 at PDHN), increased. Meanwhile, the length of most trajectories decreased significantly in 2017 (Tables S5–S7). The south-west direction of Shanghai is the hinterland of the Yangtze River Delta with a high level of anthropogenic activities. The increase in the number of trajectories from the south-west indicates a possible increase in local O<sub>3</sub> transport to Shanghai, while the decrease in the number of air masses passing oceans indicates a decrease in the mitigation effect of clean air masses on O<sub>3</sub> pollution. In addition, the length of air mass trajectories reflects the transport rate of air masses (Hou et al., 2022), and the decrease in the length of air mass trajectories indicates a slowdown in the transport rate of air masses. This may lead to more O<sub>3</sub> being transported to Shanghai from surrounding areas.

When the wind rose diagram and wind speed were analyzed (Figs. S17 and S18), the wind direction and wind speed in 2017 exhibited significant differences compared to those in 2013. From the bivariate polar plot (Fig. 3), it can be seen that O<sub>3</sub> pollution is heavier when south-west winds dominate and the lowest O<sub>3</sub> concentration occurs when north-east winds occur. According to the comparison of the wind rose diagram, the number of north-easterly winds was significantly reduced in 2017 compared to 2013, and the wind speed decreased at all three stations. The comparison results indicate that in 2017, compared to 2013, the frequency of north-east winds that can alleviate O<sub>3</sub> pollution decreased, and wind speed significantly decreased, leading to overall favorable conditions for ozone accumulation in the 2017

wind field. These results indicate that the unfavorable ODC was the primary cause of the unfavorable  $O_3$  MET in 2017 compared to 2013. In fact, according to Liu and Wang (2020a), the increase in  $O_3$  in 2017 was caused by a change in wind field, which is consistent with our findings.

After the implementation of the BSPC during 2018–2020, there was a significant reduction in daytime  $O_3$  at PT, DSL, and PDHN by 20.7 ppb (30.8%), 10.2 ppb (15.2%), and 15.8 ppb (25.6%), respectively, compared to 2017 (Fig. 9). Although the effects of this policy on  $O_3$  RBG were moderate, there was a decrease in  $O_3$  LC at PT (2.6 ppb), DSL (1.8 ppb), and PDHN (3.6 ppb), accounting for 12.6%, 17.6%, and 22.8% of the overall decline, respectively. This suggests a shift in the  $O_3$  formation regime (Xu et al., 2019). Furthermore,  $O_3$  MET decreased by 18 ppb, 8.4 ppb, and 12.2 ppb at PT, DSL, and PDHN, respectively, accounting for 87%, 82.4%, and 77.2% of the overall decline (Fig. 9). Therefore, it can be concluded that the significant reduction in  $O_3$  in 2021 was mainly driven by favorable weather conditions for  $O_3$  mitigation, but emission reduction also played an important role in controlling the  $O_3$  situation, particularly at the regional level. Further investigation into the contribution of meteorological factors revealed that all types of meteorological factors favored  $O_3$  mitigation, including decreasing SSR, decreasing T, increasing TP, increasing RH, and WS (Fig. S18). Based on the comparison of meteorological parameters, the overall decrease in  $O_3$  MET may have originated from a decrease in solar radiation or an increase in precipitation.

The implementation of the APPCAP and the BSPC has resulted in a significant reduction of  $O_3$ , particularly those that are photochemically formed. Furthermore, it may have led to a shift in the  $O_3$  formation regime in Shanghai. Therefore, if these precursor emission reduction policies are continued, they could eventually alleviate the  $O_3$  pollution problem. Weather also plays a significant role in  $O_3$  pollution during these periods. For example, unfavorable ODC in 2017 led to an increased  $O_3$  pollution condition, while a decrease in solar radiation or an increase in precipitation in 2021 amplified the effect of pollution control efforts. During both periods of emission reduction policies, the reduction of  $O_3$  RBG is sustained because rural and suburban areas are generally in the  $NO_x$ -limited regime, and  $NO_x$  reduction can lead to a reduction in suburban  $O_3$ , resulting in regional  $O_3$  mitigation (Li et al., 2021a). The persistent decrease in  $O_3$  RBG indicates that regional  $O_3$  precursor emission controls have some effect on  $O_3$  reduction. However, based on the changes in  $O_3$  LC, it is clear that stringent emission reduction of regional pollutants combined with precise characterization of  $O_3$  formation regime is more effective for  $O_3$  reduction.

Our results were compared with previous studies. For example, Wang et al. (2019) found that anthropogenic

emissions contributed positively to  $O_3$  during the summers of 2014 and 2015 in Shanghai, compared to 2013, while meteorology and natural sources had negative contributions. Similarly, Yin et al. (2021) used eXtreme Gradient Boosting (XGBoost) model and GEOS-Chem and found that, compared to the summer of 2019, emissions caused a decrease in  $O_3$  while meteorology caused an increase in  $O_3$  in Shanghai during the summer of 2020, which is similar to the results of our study. Li et al. (2020) found that meteorological factors caused an increase in  $O_3$  at a rate of 0.2 ppb/yr and emissions caused an increase in  $O_3$  at a rate of 1.5 ppb/yr for the YRD region between 2013 and 2017 using multilinear regression (MLR) model method and GEOS-Chem, which is similar to the results of the regional station DSL. Chen et al. (2020) used MLR model and found that, in the YRD region, emissions caused  $O_3$  to rise by 11.2  $\mu\text{g}/\text{m}^3$  in 2018 compared to 2014, which is comparable to the DSL station. Meanwhile, meteorology caused  $O_3$  to rise by 1.3  $\mu\text{g}/\text{m}^3$ , which is lower than the results from the DSL station. In addition, the meteorological impact, regional background  $O_3$ , and local formed  $O_3$  were also compared with previous research in previous sections. These comparative results can well support the reliability of our research results.

## 4 Conclusions

The formation of  $O_3$  is influenced by a combination of factors including meteorology, background, and local emissions, while its relationship with precursors is nonlinear. Therefore, isolating the effects of these influential factors is crucial for assessing the  $O_3$  pollution issue. In this study, we proposed an integrated method that combines RF-based meteorological normalization technology, PCA-based regional background  $O_3$  estimation, and SHAP-based causal attribution to separate observed  $O_3$  into meteorological, local formed, and regional background contributions. Our investigation of daytime  $O_3$  during the warm season between 2013 and 2021 in Shanghai revealed that meteorological factors that directly affects OPG include temperature, humidity, and solar radiation, etc., which play a significant role in  $O_3$  variations, with a cumulative absolute contribution of 13.44–19.03 ppb, and their effects on  $O_3$  have generally comparable yearly patterns. The overall effect of the ODC on  $O_3$  is 13.18–13.66 ppb, with individual characteristics among different sites. The regional background  $O_3$  is  $48.8 \pm 0.3$  ppb, accounting for 79.6%–89.4% of the observed  $O_3$  at different sites and shows an overall declining trend of 0.018 ppb/yr. Local formed  $O_3$  in urban and regional sites increased and then decreased with a concentration range between 5.9–9.0 ppb and 8.9–14.6 ppb, respectively, while local formed  $O_3$  in the upwind background site showed fluctuating variations with

concentrations ranging between 2.5–7.4 ppb. The O<sub>3</sub> rebound in 2017 compared to 2013 was primarily influenced by the unfavorable ODC and unbalanced emission reductions. In contrast, the O<sub>3</sub> decline in 2021 compared to 2017 was primarily influenced by overall unfavorable meteorological conditions for O<sub>3</sub> formation and further emissions reduction.

In summary, our method effectively captures the characteristics of O<sub>3</sub> variability and quantifies the effects of various factors on O<sub>3</sub> concentration. It should be noted that the intrinsic links between meteorological and chemical processes have been shown to accelerate chemical reaction rates and influence the abundance of VOCs in the troposphere, ultimately impacting surface O<sub>3</sub> concentrations. However, our research employs an integrated machine learning method to analyze the relationship between observed ozone and variables such as meteorology and emissions, prior to assigning the observed ozone to each influencing factor. This approach avoids potential impacts that may arise from the ozone formation process and enables the systematic quantification the impact of each factor, yielding relatively accurate results. These findings provide a more comprehensive and systematic understanding of the contribution of each factor to O<sub>3</sub> formation and identify which factors have a greater impact on O<sub>3</sub> concentrations. This information can aid in the development of more targeted and scientifically sound prevention and control strategies for local O<sub>3</sub> formation mechanisms, as well as proposing emergency plan based on weather forecasts. Although our method was applied to only three sites in Shanghai, these sites exhibited different geographically induced pollution characteristics, and the long-term time series provided good coverage of all types of O<sub>3</sub> pollution patterns. Therefore, the methods outlined in this work can be extended to other regions.

**Authorship Contribution** JX performed the data analysis and prepared the manuscript with contributions from all co-authors. FT Wang conducted the PCA analysis. KZ conducted the MK and ML analysis. DJ, YSD, YWL and QYF conducted the field observations. HHZ, EY, YJW, LH and TL helped to interpret the data analysis results. JSF reviewed the manuscript. LL formulated the research goals, led data analysis and discussions, edited and reviewed the manuscript. All authors contributed to data interpretations and discussions.

**Competing Interests** The authors declare no conflict of interest.

**Data and Code Availability** Data are available upon request to the corresponding authors. The IM model is available through a GitHub repository: <https://github.com/Xuejin66/An-integrated-machine-learning-approach-that-can-elucidate-factors-influencing-long-term-ozone-chang>.

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