EC1340 Topic #1

Introduction
and
Measuring climate change, CO_2 , and the link between them

Matthew A. Turner
Brown University
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Outline

Stabilizing atmospheric CO_2 I

- World emissions of C in 2013(most recent IPCC report) were about 13 Gt. Stabilizing atmospheric concentrations (not temp) requires cutting this about in half to 6.5Gt.
- There are about 7.4 bn people in the world as of 2015. Stabilization requires reducing emissions to 6.5Gt/7.5bn $<$ 1t C emissions per person.
- Current per capita emissions/incomes are about: US/CA/AU, 4.6t/55000$; China is 1.8/7600$; India is 0.5/1500$.
- The US needs between a 50% and 80% reduction to hope to reach this target.

This appears difficult to accomplish without reducing the number of people, their consumption, or being very clever. Being clever looks attractive here.

Some questions

- The Green New Deal proposes meeting 100% of US power demand with renewables. Is this a good idea?
- Given that the RGGI (Regional Greenhouse Gas Initiative) is in place, is the state carbon tax Aaron Regunberg proposed a good idea? RGGI website is here, `https://www.rggi.org/`. As of Q2 2019 the allowance auction cleared at \$5.62 per allowance. An allowance allows emission of 1 short ton of CO_2 by New England power plant $>$ 25MW.(short ton $=$ 2000lb $<$2200lb $=$1000kg$=$ 1 metric ton).
- How scary is this animation?

  `https://www.brown.edu/Departments/Economics/Faculty/Matthew_Turner/ec1340/lectures/pnas.1512482112.sm01.avi`

  (personal correspondence, Anders Leverrmann, Sept. 2018).

Content of the course.

We would like to think carefully about the questions that climate change raises. For example,

- How fast should we approach CO_2 stabilization?
- What are the trade-offs between economic welfare and climate?
- What policies should we use to achieve CO_2 reductions?

To think about these questions, it would be helpful to have a model in which the tradeoffs between consumption, emissions and climate at one time and another can be explicitly calculated and examined.

The model developed in ‘A Question of Balance’ does exactly this, and one of the main objectives of the course is to allow you to read this book and to understand what it does.

As a preview, here is some of the Nordhaus, DICE (Dynamic Integrated Climate Economy) model:

$$W = \sum_{t=0}^{24} L(t) \frac{c(t)^{1-\alpha}}{1-\alpha} \left(\frac{1}{1+\rho} \right)^{t} \tag{1}$$

$$Q(t) = \Omega(t) \left[1 - \Lambda(t) \right] A(t) K(t)^{\gamma} L(t)^{1-\gamma} \tag{2}$$

$$Q(t) = C(t) + I(t) \tag{3}$$

$$K(t) = I(t) + (1 - \delta_K) K(t-1) \tag{4}$$

$$E(t) = \sigma(t) \left[1 - \mu(t) \right] A(t) K(t)^{\gamma} L(t)^{1-\gamma} \tag{5}$$

$$\Lambda(t) = \pi(t) \theta_1(t) \mu(t)^{\theta_2} \tag{6}$$

plus a description of the way climate, carbon, population and technology evolve.

The DICE model has more stuff in it than we need to start thinking about the problem (and it's a bit hard).

What is the minimum amount of hardware that we need to discuss this problem? We need to describe (at least),

- how CO_2 affects climate (a climate model).
- the carbon cycle.
- how CO_2 comes from consumption.
- how climate affects consumption (and/or utility).
- how we can use resources to reduce CO_2 , i.e., a mitigation equation.
- how we are willing to make tradeoffs across time.

Notice several of these items describe physical science surrounding climate change. These are covered in the other main reading for the course, 'Storms of my grandchildren'.

Here is the math that goes with the list we just generated. It looks a lot like the 'consumer problem' you know, but with more complicated budget constraints. This will help us to organize ideas and keep track of our progress. Once we have worked our way through this problem, we'll be ready to tackle Nordhaus.

To start, we'll need some notation,

- $c_1, c_2 =$ per capita consumption now and in 100 years
- $W =$ per capita wealth/income today
- $I =$ investment today
- $M =$ expenditure on mitigation today
- $E_1 = (1 - \rho_4 \frac{M}{W})(\rho_5(c_1 + I)) =$ Emission of CO_2 today increases in s, c_1 and decreases in M
- $P_1, P_2 =$ Atmospheric concentration of CO_2 now and in 100 years
- $T_1, T_2 =$ climate now and in 100 years

Using this notation, we can state the 'baby DICE' model (BDICE) as

$$\max_{I,M} u(c_1, c_2) \qquad \text{utility} \qquad (7)$$

$$\text{s.t. } W = c_1 + I + M \qquad \text{budget} \qquad (8)$$

$$c_2 = (1 + r)I - \gamma(T_2 - T_1)I \qquad \text{production} \qquad (9)$$

$$E = (1 - \rho_4 \frac{M}{W})(\rho_5(c_1 + I)) \qquad \text{emissions} \qquad (10)$$

$$P_2 = \rho_0 E + P_1 \qquad \text{carbon cycle} \qquad (11)$$

$$T_2 = \rho_1(P_2 - P_1) + T_1 \qquad \text{climate model} \qquad (12)$$

That is, choose savings and mitigation to maximizes welfare u.

Note that physical quantities like climate or the relationship between CO_2 concentration and climate enter this problem like prices and endowments.

We're going to work towards an understanding this problem, one parameter and equation at a time. To do this, we'll need to study the following topics

- endowment of climate and carbon, P_1, T_1 (history of CO_2 and climate).
- the link between atmospheric CO_2 and climate, ρ_1 (climate model)
- the link between emissions and atmospheric CO_2 , ρ_0 (Carbon cycle)
- cost of climate change, γ, emissions per consumption, ρ_5
- cost of mitigation (reduction of emissions), $\frac{\partial E}{\partial M}$
- what should u look like (discounting and uncertainty)

Once we understand this, we'll be able to think about solving the global warming problem, and we'll be ready to tackle the Nordhaus model.

Other things we'll want to think about that aren't in the basic global warming problem (but are in the Nordhaus model)

- Population growth
- Economic growth
- Dynamics – this is a dynamic problem, so c, T are consumption paths and climate paths. A wise regulatory program will reflect the fact that investments in climate and economic growth have different returns at different times. This will turn out to suggest a 'ramping up' of mitigation expenditures.
- There is LOTS of uncertainty. This makes everything more difficult.

These are just generalizations of the basic model.
With the model in hand, we'll be able to think about policies to manage CO_2 . This will be the last third of the course.

Endowment of climate - outline

We want to try to understand our endowment of climate.

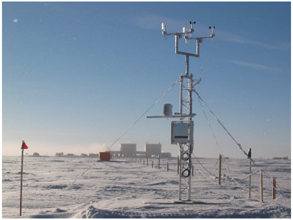
1. Measured temperature and the problem of measuring temperature.
2. Other measured aspects of climate.
3. Icecores and oxygen isotope ratios, other climate proxies.

Measured Temperature

This is what modern weather stations look like:

(JFK airport in NY)

http://www4.ncdc.noaa.gov/cgi-win/wwcgi.dll?wwDI~StnPhoto~20019418~a~000



(Antarctic weather station)

http://www.gdargaud.net/Antarctica/DC2005/20050406_10_WeatherStation.jpg

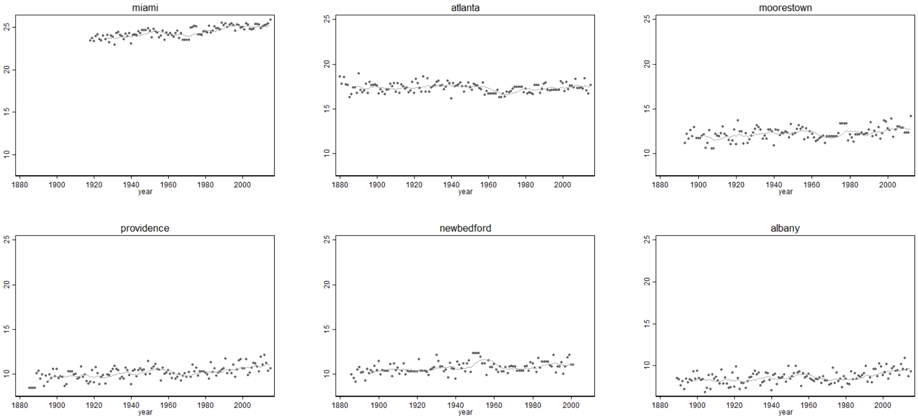
- We have measured temperature data going back about 100 years.
- There are also satellite, radiosonde and various ocean instrumental measurements, but for the most part they go back only 30-40 years.

Mean annual temperature, Durban and Shanghai

Black line is 10 year trailing average,
$T_{10,t} = 0.1 \times (T_{t-9} + T_{t-8} + ... + T_{t-0})$.

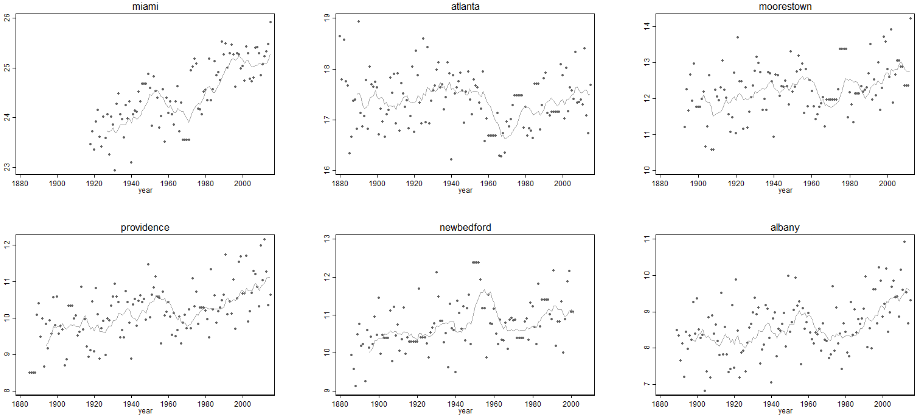
Source:
`http://data.giss.nasa.gov/gistemp/station_data/`

Mean annual temperature, US Atlantic coast I



Note: (1) Latitude affects temperature more than time. (2) More warming in the north.

Mean annual temperature, US Atlantic coast II



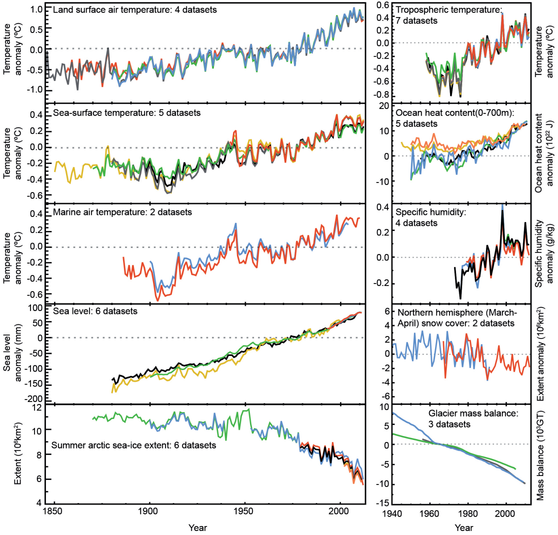
Same graphs with different axes. Change is about 1C for 100years.

Issues with measured temperature

- Measurement error within instrument.
- Different types of instruments.
- Sample of stations/instruments varies over time.
- What does 'world average temperature' mean? Average over what?(see next slide)
- Very little coverage of the ocean.
- Urban heat Islands (see IPCC 2013 TS 2.2 this effect is probably small).

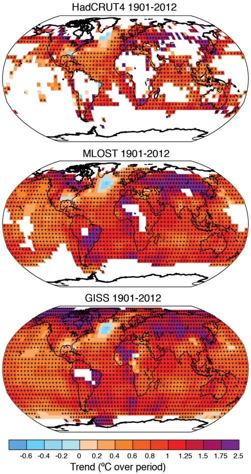
Many of these issues can be resolved with statistics. Some cannot.

IPCC summary of lots of instrumental measurements



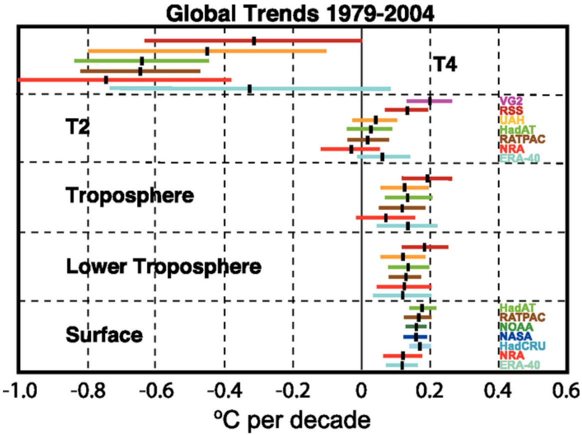
Note: warming hiatus in the last decade. IPCC 2013 figure TS.01

Distribution of surface warming



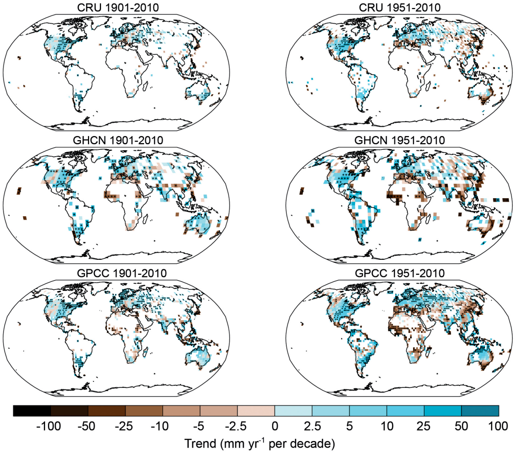
Distribution of change in surface temperature 1901-2012 as determined by linear trends for three data sets. White indicates missing data. IPCC 2013 Physical Science Basis figure TS.02

Stratospheric cooling



Stratosphere begins about 3 miles up. It has been cooling. (1) The world is complicated, some places on the surface cool too. (2) There is very little energy stored in the Stratosphere. IPCC 2007 figure 3-18

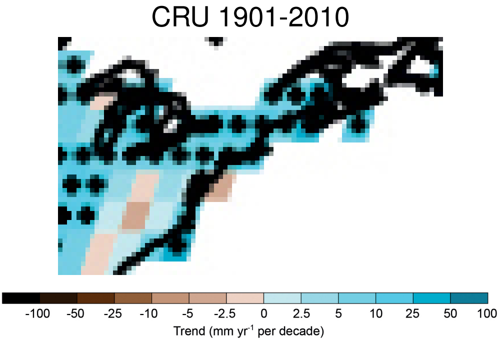
Changes in rainfall I.



Trends in rainfall from 3 data sets. "+" 's indicate changes statistically different from zero. Mean annual rainfall in Providence is about 47 inches $\sim$ 1175mm. In Salt Lake, annual average precipitation is 455. Are these changes big? IPCC Physical Science Basis 2013 figure TS.TFE.1.2

Changes in rainfall II.

Zoom in on Providence:



Providence had 2.5-5mm increase in annual precipitation each decade, for 11 decades $\sim$ 27.5-55mm increase on a base of about 1175mm. This is a 2-5% increase.

Sea level rise ≈20cm since 1860

From 1961 on, measured sea level rise is just less than 2mm/year. This occurs because water expands as it gets warmer and ice melts. The IPCC attributes about 60% of sea level rise to warming.

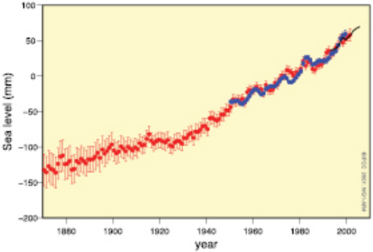


Figure TS.18. Annual averages of the global mean sea level based on reconstructed sea level fields since 1870 (red), tide gauge measurements since 1950 (blue) and satellite altimetry since 1992 (black). Units are in mm relative to the average for 1961 to 1990. Error bars are 90% confidence intervals. IPCC 2007, fig. TS.18

Ocean acidity is also increasing.

Long run climate

Climate is something that happens on geological time scales. Some of the recent rise in global measured temperature may be part of normal climate fluctuations. To assess this, we need longer time series. Since no longer series of measured temperatures exist, we use proxies. There are many sorts of proxy data that can give us information about historical temperature.

- Boreholes.
- Tree rings.
- Glacial extent.
- Species range.

Generally, all point toward warming. Among these ice cores are of particular interest.

Icecores I

(Extracting an ice core)

http://kaira.sgo.fi/2011/05/ice-cores-part-1.html

(Extracting an ice core)

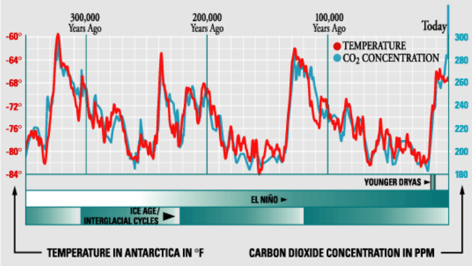
http://kaira.sgo.fi/2011/05/ice-cores-part-1.html

(sections of an icecore)

http://www.gdargaud.net/Antarctica/DC2005/20050406_10_

Icecores are records of the water deposited on the site over many years. Water consists of several isotopes. In particular, some water molecules contain oxygen atoms with 8 neutrons, while other contain 10. The ratio of these isotopes in polar snow varies systematically with temperature. Thus, icecores allow us to track temperature back into the very distant past, about 800,000 years in some cases.

Isotope ratio from Vostok Antarctica ice core



Red is temperature imputed from isotope ratio. This core goes back about 400,000 years, others go back 800,000 and show the same thing. It is warm now by the standards of the last several hundred thousand years.(NB: blue is CO_2 from the same core).

http://www.koshland-science-museum.org/exhibitgcc/historical02.jsp

The current state of climate

We've measured T with instrumental measurements and proxies

- Measured temperature data, though subject to some problems, suggests dramatic warming.
- Other short run proxies, glacial extent and species range, confirm measured temperature data.
- Ice core data shows that the last 200 years are warmer than the preceding 650,000+. Other long run proxies (tree rings, boreholes) are broadly consistent, though they also have problems.
- Weak evidence of increased rainfall in North America.

Endowment of Carbon - outline

1. Basic atmospheric chemistry.
2. Measured CO_2 .
3. Icecore CO_2 .

Basic atmospheric chemistry

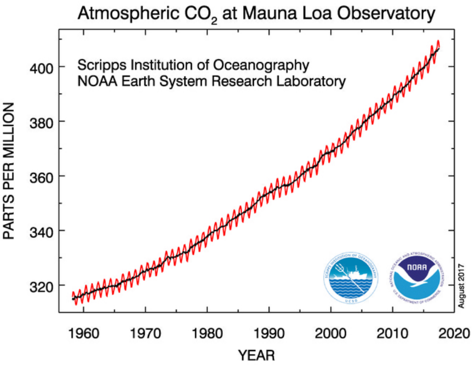
- Nitrogen 78%, 780,000 ppm
- Oxygen 21%, 210,000 ppm
- Argon 0.93% 930 ppm
- CO_2 0.0365% , 365 ppm
- Methane (CH_4) 1.7 ppm

and lots of other trace gases. From: Introduction to Atmospheric Chemistry, D. J. Jacob, Princeton University press, 1999.

CO_2 concentration $=$ 409ppm in July 2018, $=$ 412ppm in July 2019.

Atmospheric Carbon Measurements

Since 1959, the Mauna Loa observatory in Hawaii has measured atmospheric concentration of CO_2 daily. CO_2 disperses rapidly through the atmosphere, so a single observatory gives a good description of the whole world.

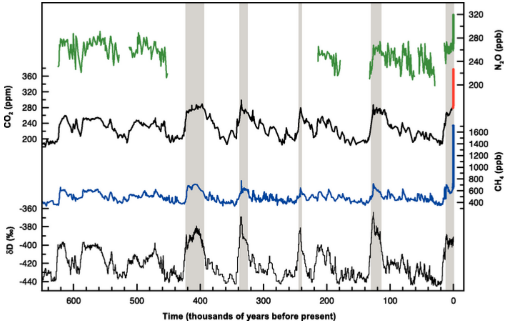


http://www.esrl.noaa.gov/gmd/webdata/ccgg/trends/co2_data_mlo.pdf

Icecore CO_2 Measurements

Small air bubbles are trapped in icecores. By checking the CO_2 concentration in this air, it is possible to create much longer series of atmospheric CO_2 , though with somewhat lower frequency (we can't reconstruct daily CO_2 concentrations from icecores). Similarly for other trace gases.

Trends in CO_2 and other trace gases



Antarctic ice and recent atmospheric measurements of N_2O (top), CO_2 (second), CH_4 (third). Data cover 650,000 years and the shaded bands indicate current and previous interglacial warm periods. Bottom graph is isotope ratio

CO_2 is very high now.

IPCC 2007 Physical Science Basis, figure TS 1

The current state of atmospheric CO_2

We've measured CO_2 directly with instruments and with icecore bubbles,

- Mauna Loa measured CO_2 shows steady increase since 1959.
- Icecore record of CO_2 and other trace gases show that the levels of last 200 years are without precedent in the preceding 650,000+ years.

Next step is to look at how atmospheric CO_2 affects temperature.

Summary of progress

Recall BDICE :

$$\max_{I,M} u(c_1, c_2) \tag{13}$$

$$\text{s.t. } W = c_1 + I + M \tag{14}$$

$$c_2 = (1 + r)I - \gamma(T_2 - T_1)I \tag{15}$$

$$E = (1 - \rho_4 \frac{M}{W})(\rho_5(c_1 + I)) \tag{16}$$

$$P_2 = \rho_0 E + P_1 \tag{17}$$

$$T_2 = \rho_1(P_2 - P_1) + T_1 \tag{18}$$

We have just described,

- measured and proxy climate data $\sim T_1$
- historical and modern CO_2 concentration data $\sim P_1$

CO_2 and climate – Outline

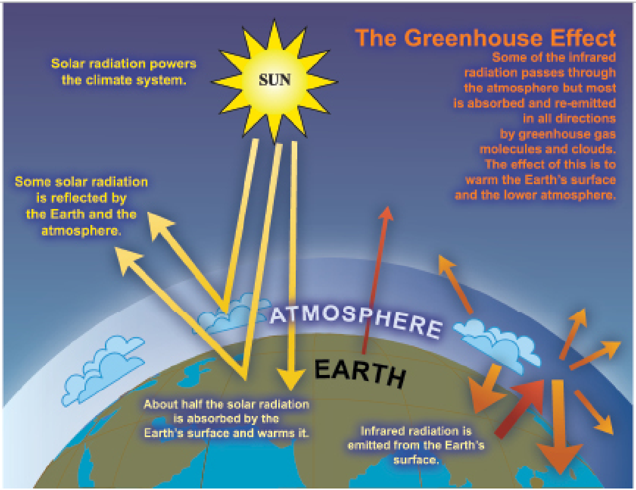
We now begin to investigate the relationship between atmospheric CO_2 and future climate (In BDICE , this relationship is determined by the parameter ρ_1). More specifically, we'll look at

1. Basic physics of the greenhouse effect
2. Correlation of carbon levels and temperature
3. Complications: Astronomical cycles, aerosols, heat transmission
4. Climate models
5. Climate projections
6. Sea level rise

Physics of the greenhouse effect

1. Solar radiation arrives at Earth at a known and nearly constant rate.
2. About 59% of this radiation is absorbed by Earth (or its atmosphere), the rest is reflected back to space.
3. All radiation of heat from the Earth occurs at the upper edge of the atmosphere.
4. All radiation from the surface occurs at the surface.
5. GHG molecules reflect heat radiating from the surface back to the surface and prevent it from radiating to space.

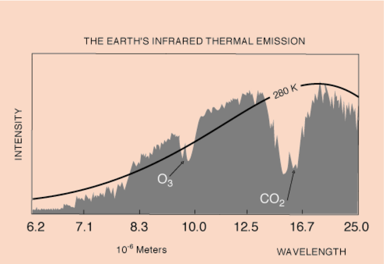
The physics of this process are elementary, are well understood and uncontroversial. See ‘Principles of Atmospheric Science’, J. E. Frederick, Jones and Bartlett, Sudbury MA, 2008.



http://www.ipcc.ch/publications_and_data/ar4/wg1/en/faq-1-3.html

Spectral evidence for CO_2 warming effect

CO_2 only stops the radiation of certain wavelengths of light from making out to space. These wavelengths can be measured in the laboratory. When satellites look back at Earth and measure the intensity of reflected spectra, there is a gap where CO_2 absorbs reflected light in the relevant wavelength. That is, the physicists are right. CO_2 traps heat.



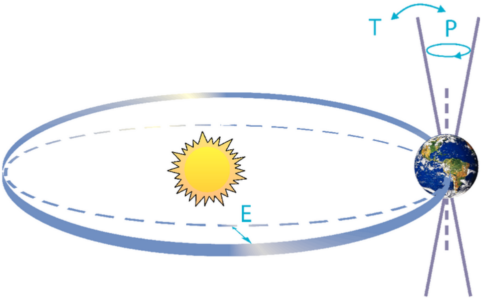
Graph taken from online physics forum, but it looks just like figure 2.11 in ‘Principles of Atmospheric Science’ J. Frederick, 2008

Complications

That more atmospheric CO_2 causes warming seems clear, other issues are less clear. Here are some of the problems:

- non-constant solar radiation
- Aerosols
- transmission of heat in the atmosphere
- transmission of heat to and from the Ocean

Cycles in solar radiation



Schematic of the Earth's orbital changes (Milankovitch cycles) that drive the ice age cycles. 'T' denotes changes in the tilt (or obliquity) of the Earth's axis, 'E' denotes changes in the eccentricity of the orbit (due to variations in the minor axis of the ellipse), and 'P' denotes precession, that is, changes in the direction of the axis tilt at a given point of the orbit.
http://www.ipcc.ch/publications_and_data/ar4/wg1/en/faq-6-1-figure-1.html

Aerosols

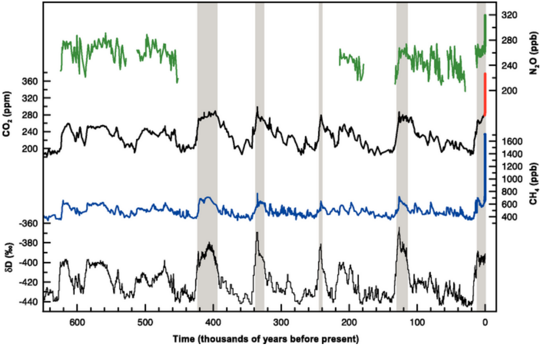
- Aerosols are small particles that float in the atmosphere, e.g., dust, diesel exhaust. They reflect light back out into space and thus promote global COOLING.
- Evidence in support of this theory is found in the cooler years that follow major volcanic eruptions.
- Aerosols are also increasing.

Heat transmission

- Radiative heat transmission appears to be a fairly straightforward problem for physicists. Heat transmission by convection is very hard.
- Heat transmission in the atmosphere COULD be quite different from what models suggest. Stay tuned for more on problems with climate models.
- Heat transmission to the ocean and within the ocean is not very well understood and is important. Water holds a lot more energy than air.

CO_2 and climate, empirical relationship

One way to guess at the total effect of CO_2 and complicating factors is to look at the historical record:



This shows a clear relationship between temperature proxies and CO_2 . BUT, the current warming cycle is clearly not like the ones before it ...

CO_2 and climate forcing I

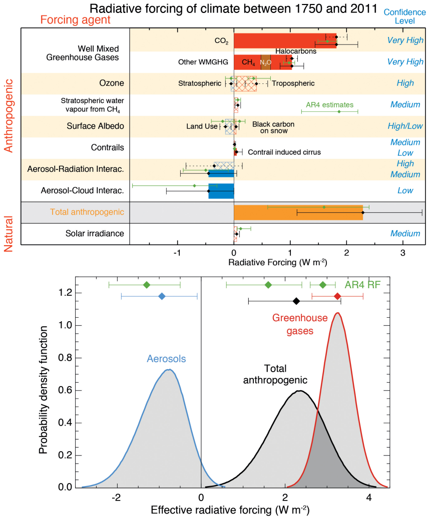
The other way to guess at the total effect of CO_2 and complicating factors is to try to add them all up. The effect of different forces on climate is measured in Watts/m^2. (To compare different heat sources, climate scientists measure 'climate forcing' in Watts/m^2.)

- baseline solar radiation ≈ 240 Watts/m^2.
- $CO_2 \approx +1.82$ Watts/m^2, $CH_4 \approx +0.48$ Watts/m^2, $N_2O \approx +0.17$ Watts/m^2
- changes in solar radiation $\approx +0.12$ Watts/m^2
- aerosols ≈ -1.0 Watts/m^2
- reflectivity from buildings ≈ -0.2 Watts/m^2

Others: airplane contrails, Montreal Protocol gases, volcanic eruptions, changes to water vapor content of air, black carbon.

IPCC Physical Science basis 2007, p23-31 and IPCC Physical Science basis 2013, TS 3.2

CO_2 and climate forcing II



IPCC Physical Science basis 2013, Fig TS.06

Models of CO_2 and climate I

If we want to think about what will happen to the climate, rather than just account for change in heat, then we need to worry about feedbacks. For example:

- As polar ice melts, reflectivity decreases.
- As climate warms, atmospheric water vapor increases (water vapor is a GHG), tundra melts and gives off CH_4 , and methane hydrates give off CH_4 .
- As temperature changes weather changes, and with it cloud cover. Clouds reflect sunlight and trap heat.
- How does the ocean absorb heat? Will this change as the climate warms? (This seems to be a very hard problem).

There are LOTS of others (I think).

Models of CO_2 and climate II

To learn the effect of forcings and feedback on climate we need to do something similar, that is, make a model of the climate in which all of the different forces can act and see what happens.

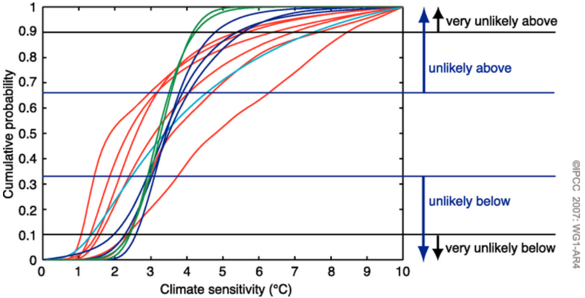
- This is going to be much harder than just adding up all of the forcings.
- There is going to be LOTS of uncertainty around model predictions.

Five minutes on climate models I

Suppose atmospheric CO_2 stabilizes at 550ppm. (Recall: Current level is 407, growth rate about 2ppm/year.) What is our best guess at ultimate equilibrium temperature based on:

- Extrapolation from 20th century measured temperature and CO_2 (red)
- AOGCM's Atmosphere Ocean General Circulation Models
- Extrapolation from long run proxy data (light blue).
- Climate models(green).

Five minutes on climate models II



Baseline is pre-industrial climate, about 0.7 degrees Celsius cooler than current. Light blue line (historical proxy data), if CO_2 stabilizes at 550 ppm the probability of an ultimate equilibrium climate 7 degrees Celsius warmer than pre-industrial norm is about 0.1.

IPCC Physical Science basis 2007 Figure TS.25

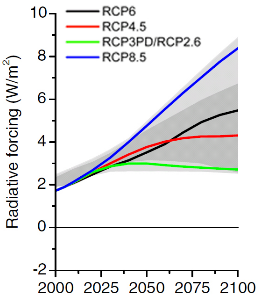
Five minutes on climate models III

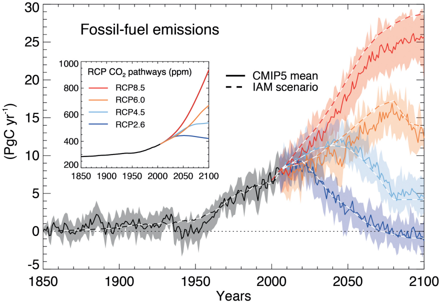
Issues:

- How to choose between models? If each is equally likely then resulting uncertainty is much larger
- These types of models, by construction, understate uncertainty.
- Remember, these models are trying to do something really hard, they're probably not very accurate. It looks like they may predict too much warming, they failed to predict the warming hiatus.

Representative Concentration Pathways (RCPs)

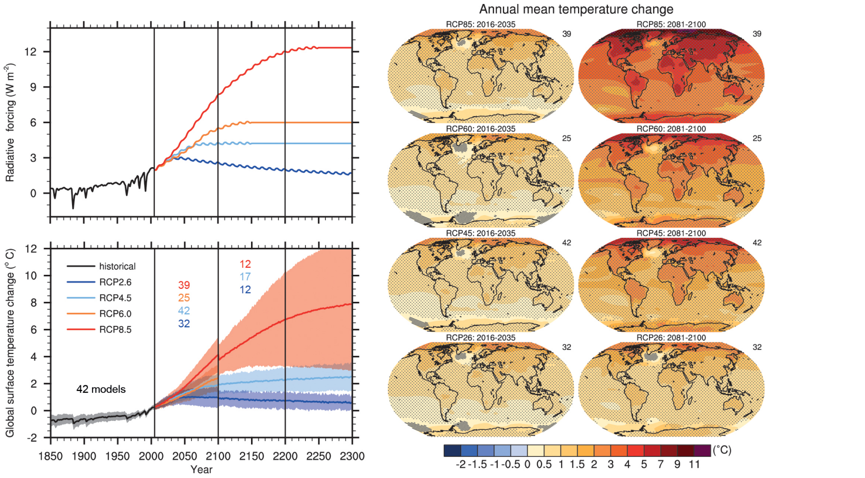
The IPCC fifth assessment report makes some choice of their favorite climate model and then makes climate projections on the basis of RCPs. These describe carbon paths that lead to different levels of climate forcing.



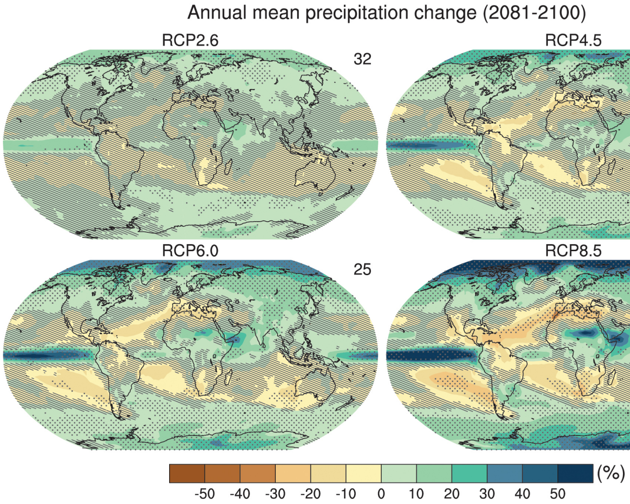


RCP 8.5 is 'business as usual', while the others involve varying degrees of mitigation.

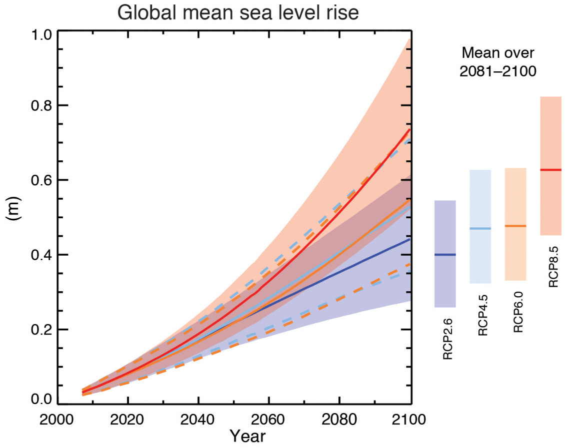
Projected temperature from Fifth Assessment Report



Projected rainfall from Fifth Assessment Report



Projected sea level rise from Fifth Assessment Report



NB: (1) Really scary sea level rise is probably not in the cards over this time frame. (2) Figure reports ‘credibility bounds’, not ‘confidence intervals’.

Stern Report steady state climate projections I

The Stern report also summarizes the output of a bunch of climate models, in what is probably a more useful way (for our purposes). The following table relates atmospheric concentrations of CO_2 to the PROBABILITY that climate stabilizes at a particular level. NB: We're ignoring the 2-300 year adjustment period, which is when most of us will live.

Stern Report steady state climate projections II

This table provides an indicative range of likelihoods of exceeding a certain temperature change (at equilibrium) for a given stabilisation level (measured in CO_2 equivalent). For example, for a stock of greenhouse gases stabilised at 550 ppm CO_2e, recent studies suggest a 63 - 99 % chance of exceeding a warming of 2°C relative to the pre-industrial.

The data shown is based on the analyses presented in Meinshausen (2006), which brings together climate sensitivity distributions from eleven recent studies (chapter 1). Here, the 'maximum' and 'minimum' columns give the maximum and minimum chance of exceeding a level of temperature increase across all eleven recent studies. The 'Hadley Centre' and 'IPCC TAR 2001' columns are based on Murphy *et al.* (2004) and Wigley and Raper (2001), respectively. These results lie close to the centre of the range of studies (Box 1.2). The 'IPCC TAR 2001' results reflect climate sensitivities of the seven coupled ocean-atmosphere climate models used in the IPCC TAR. The individual values should be treated as approximate.

The red shading indicates a 60 per cent chance of exceeding the temperature level; the amber shading a 40 per cent chance; yellow shading a 10 per cent chance; and the green shading a less than a 10 per cent chance.

Stabilisation Level (CO_2e)	Maximum	Hadley Centre Ensemble	IPCC TAR 2001 Ensemble	Minimum
Probability of exceeding 2°C (relative to pre-industrial levels)				
400	57%	33%	13%	8%
450	78%	78%	38%	26%
500	96%	96%	61%	48%
550	99%	99%	77%	63%
650	100%	100%	92%	82%
750	100%	100%	97%	90%
Probability of exceeding 3°C (relative to pre-industrial levels)				
400	34%	3%	1%	1%
450	50%	18%	6%	4%
500	61%	44%	18%	11%
550	69%	69%	32%	21%
650	94%	94%	57%	44%
750	99%	99%	74%	60%
Probability of exceeding 4°C (relative to pre-industrial levels)				
400	17%	1%	0%	0%
450	34%	3%	1%	0%
500	45%	11%	4%	2%
550	53%	24%	9%	6%
650	66%	58%	25%	16%
750	82%	82%	41%	29%
Probability of exceeding 5°C (relative to pre-industrial levels)				
400	3%	0%	0%	0%
450	21%	1%	0%	0%
500	32%	3%	1%	0%
550	41%	7%	2%	1%
650	53%	24%	9%	5%
750	62%	47%	19%	11%

Nordhaus' rule of thumb

Doubling CO_2 from preindustrial level (280ppm) to 560ppm give a temperature increase of about 3 degrees Celsius above pre-industrial level.e.g. Nordhaus p54.

From our model of the climate change problem,

$$T_2 = \rho_1 (P_2 - P_1) + T_1 \tag{19}$$

or

$$\rho_1 = \frac{T_2 - T_1}{P_2 - P_1} \tag{20}$$

$$= \frac{3}{280} \frac{\text{Degrees}}{\text{ppm } CO_2} \tag{21}$$

Summary: CO_2 and climate

Nordhaus, Stern report and IPCC climate projections give us the relationship between atmospheric CO_2 and climate.

- Nordhaus' rule of thumb fits our framework neatly.
- Stern and IPCC are more complicated because time is not just two periods, climate is not just temperature, and climate will not change the same way in all places.
- Other aspects of our environment will change as a consequence of higher CO_2 levels and warming:
  - Sea level will rise
  - Ocean acidification
  - Ecological changes/ species extinctions
  - Changes to regional weather, e.g., monsoon patterns, heat spells, droughts, extreme weather.

Conclusion I

- We posed the problem of climate change as a simple optimization problem, ‘the Baby Dynamic Integrated Climate Economy model’. To solve this problem we need to understand our endowment of climate and carbon, and how they are related. We’ve started on this today.
- Climate:
  - Measured and Icecore data show increases in temperature over past 650,000+years.
  - Climate is complicated. Measured and proxy data show other changes in climate; rainfall, sea level, ocean acidity, ice extent, etc.
- CO_2 :
  - Measured data show increases in CO_2 .
  - Icecore data show increases in CO_2 .

Conclusion II

- Climate and CO_2
  - Visible in icecore data (homework)
  - The physics relating atmospheric CO_2 to warming is elementary and uncontroversial. But there are many complications; aerosols, heat transmission, feedbacks.
  - Stern, IPCC and Nordhaus all make climate projections, we'll mostly use Nordhaus just because it is simple.

Our next projects are to understand the path of CO_2 emissions, the relationship between CO_2 emissions and atmospheric concentration, and the relationship between consumption and emissions.