

GEOTHERMAL GRADIENT AND ITS ROLE ON INDUCED SEISMICITY IN RATON
BASIN, COLORADO AND NEW MEXICO

By

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A thesis submitted to the
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Master of Science
Department of Geology
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Geothermal Gradient and its role on induced seismicity in Raton Basin, Colorado and New Mexico

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ABSTRACT

Pfeiffer, Katherine (M.S., Geological Sciences)

Geothermal Gradient and its role on Induced Seismicity in Raton Basin, Colorado and New Mexico

Thesis directed by Shemin Ge

Raton Basin, in southeastern Colorado and northern New Mexico, is known for coal bed methane (CBM) production. Raton Basin also has been a site of wastewater injection, a byproduct of the CMB production since the 1990s. Raton Basin has experienced an increase in earthquakes since wastewater injection began with over $130 M \geq 3$ events between 1990 and 2016. Seismicity has been attributed to anthropogenic causes because of a significant increase since 2001 following commencement of wastewater disposal. The largest earthquake is a moment magnitude (M_w) 5.3 in 2011. Peer-reviewed studies have attributed the increase in earthquakes to pore pressure increase due to injection. However, other studies have presented the possibility of a naturally high geothermal gradient to the increase in earthquake activity. The average geothermal gradient is $49 ^\circ\text{C}/\text{km}$ and may also be the cause of seismicity in the Basin. The goal of this study was to assess if there is a spatial relationship between wastewater injection wells, seismicity, and the geothermal gradient. Cool wastewater has the ability to change the stress of the warmer rocks in the subsurface, thus leading to failure.

I hypothesized that injection wells in areas with a higher geothermal gradient would have a stronger correlation with earthquakes. Results of coupled fluid and heat transport modeling do not show a clear correlation between the role of the geothermal gradient in induced seismicity from wastewater injection. Modeling results show a temperature change of a few degrees $^\circ\text{C}$ at an observation point 100 m away from the disposal wells. This temperature change is at the same

depth of disposal. The modeling results indicate most fluid moves laterally and thus the temperature decrease does not reach the depths at which earthquakes are occurring.

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1. INTRODUCTION

Induced seismicity has been gaining notoriety throughout the United States during the past few decades due to the sharp increase in the number of midcontinent earthquakes. In fact, during the past decade there has been a six-fold increase in earthquakes in this region [Zhang *et al.*, 2013]. Since the 1920s, pumping fluids in and out of the subsurface has been recognized to potentially cause earthquakes [Hubbert and Rubey, 1959, Folger and Tiemann, 2015, Foulger *et al.*, 2017]. Colorado, as well as Texas and Oklahoma, are at the forefront of this issue as they are actively producing a large amount of oil and gas, which requires disposal of large amounts of wastewater from hydraulic fracturing. Disposal has statistically correlated to increases in seismicity [Folger and Tiemann, 2015, Foulger *et al.*, 2017]. The prominent increase in the number of earthquakes in these regions has been attributed to the disposal of wastewater. Wastewater - water that is unsuitable for agricultural or drinking water use - is a byproduct of natural gas and oil extraction [Rubinstein and Mahani, 2015]. One way to dispose of this produced wastewater is to inject it back into deep formations in the subsurface.

Studies have shown that pore pressure may increase with the injection of these fluids [Rubinstein *et al.*, 2014, Weingarten *et al.*, 2015, Nakai *et al.*, 2017]. The migration of pore pressure increase to faults has allowed for seismicity. In the fault zone, the effective stress and the frictional resistance may be lowered, allowing faults to slip [Hubbert and Rubey, 1959; Rubinstein *et al.*, 2014]. The shock wave produced from fault slippage has the potential to reach the surface and cause damage to infrastructure if sufficiently large [Folger and Tiemann, 2015]. However, there is an ongoing debate in some regions as to the cause of the seismicity, with experts pointing to other potential causes. Raton Basin, in southern Colorado, is an example of

such an area (Figure 1). The most widely expressed alternative to injection-induced seismicity is the depth at the occurrence of the earthquakes relative to the injection depth. Specifically, a majority of the earthquake locations in the Raton Basin are traced to the crystalline basement rock, with locations ranging from 0.5 to 7 km below the injection well depths (Figure 2) [Weingarten *et al.*, 2015]. In addition, Raton Basin is being explored as a potential geothermal resource because of its high temperature gradient [Macartney and Morgan, 2011]. Elevated temperatures in the subsurface can change fluid properties such as viscosity and hydraulic conductivity and allow them to migrate more easily [De Simone *et al.*, 2013]. We explore whether the link to seismicity could be due to the ease at which the pore pressure is able to migrate to faults and result in fault slippage as well as thermal contrast that results in thermo-elastic effects.

For this study, temperature and fluid flow were modeled in Raton Basin to see if there is a spatial relationship between injection wells, seismicity, and geothermal activity as well as to further understand the physics of induced seismicity.

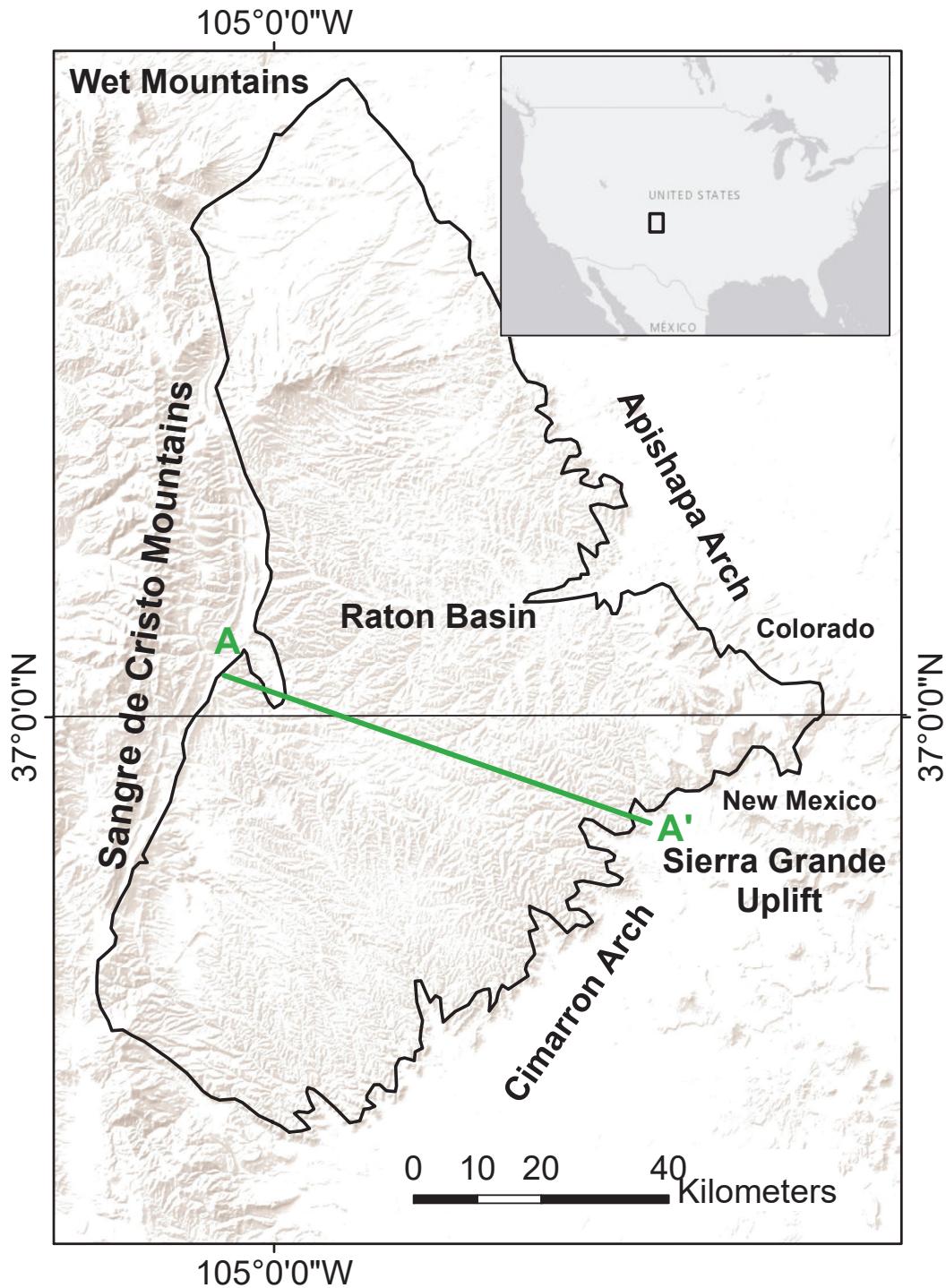


Figure 1. Outline of Raton Basin with structural features. A to A' shows the cross section line that is represented in Figure 6.

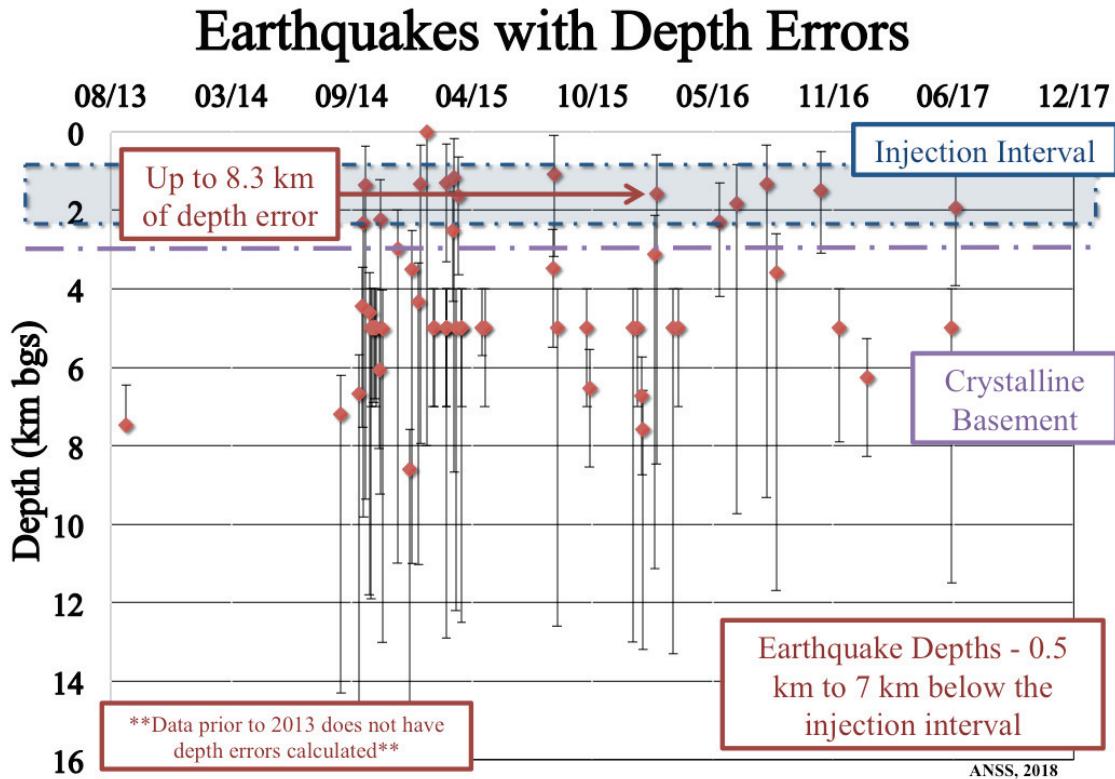


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2. BACKGROUND

2.1 Study Site

Raton Basin is an asymmetrical synclinal basin (Figure 1 and 5). The Basin is approximately delineated by the outcrop of the Cretaceous Trinidad Formation and is estimated to be approximately 2,200 mi² (5,698 km²) (Table 1). Originally explored by Stephen Long in his Great Plains Expedition, the Basin was later commercially mined for coal from 1870 to 2002 [Osterhout *et al.*, 2015]. As a result of a change in mining practices in the 1970's, implemented by the U.S. Bureau of Mines, coal bed methane (CBM) became more popular. This change involved extracting the methane before mining coal to improve mine safety. Since the early 1990s, companies produce CBM from the Upper Cretaceous aged formations (Table 1). A

byproduct of this CBM production is wastewater due to the dewatering procedure. Methane is adsorbed into the coal macerals so the coal seam has to be depressurized to be produced. Macerals are organic components of coal. As you pump water from the well you decrease the pressure which allows methane to desorb from the coal and flow as gas state to the surface [Weingarten, 2015]. Wastewater injection began in 1994 on the Colorado side of Raton Basin and 1999 on the New Mexico side, while potential injection driven seismicity began in 2001 [Weingarten, 2015]. Few seismic events were present in the Basin prior to 2001 [ANSS, 2016]. Wastewater in Raton is injected primarily into the Dakota Formation, a 68-m thick sandstone that lies approximately 1,500 m below the surface (Table 1) [Geldon, 1989, COGCC, 2016, Weingarten, 2015].

2.2 Geology

For this study, the boundary of Raton Basin, as defined by the outcrop of the Trinidad Formation, is approximately 130 km long north - south and 60 km wide east - west. The Basin is an asymmetrical syncline with steeply dipping beds along its western limb and much shallower dips along the eastern limb adjacent to the Great Plains region [Baltz, 1965; Topper, 2011] (Figure 5). Elevations range from 4,000 ft (1,219 m) in the east to 11,000 ft (3,352 m) in the west. Laramide faulting and thrusting during the Paleocene, approximately 70 - 55 my ago, resulted in uplift of the western beds to their current position. The Basin is bounded to the west by the Sangre de Cristo Mountains (Figure 1) and the western edge is heavily deformed by thrust faults and several major folds (Figure 3) [Baltz, 1965]. As a result, Precambrian crystalline rocks are juxtaposed against Lower Permian sedimentary rocks [Nelson et al., 2013; Baltz, 1965]. The north plunging axis, the La Veta syncline, runs parallel to the Sangre de Cristo (Figure 3) [Baltz, 1965]. The southern extension of the Wet Mountains is the northern boundary of the Basin

(Figure 1). To the north, the north-south trending Wet Mountains split into two different branches, the Apishapa and Greenhorn (Figure 3). The northern extension plunges east and southeast, known as the Apishapa Uplift and the Apishapa Arch forms the northeast boundary of Raton Basin. The southern spur, known as the Greenhorn Nose, plunges south into the Basin (Figure 3). Raton Basin is separated to the south from the Las Vegas sub-basin by an interbasinal arch, known as the Cimarron Arch. The Las Animas Uplift forms the eastern extent of Raton Basin. The Sierra Grande Uplift in northeastern New Mexico also bounds Raton Basin to the east, and is known as the Sierra Grande Arch [Baltz, 1965].

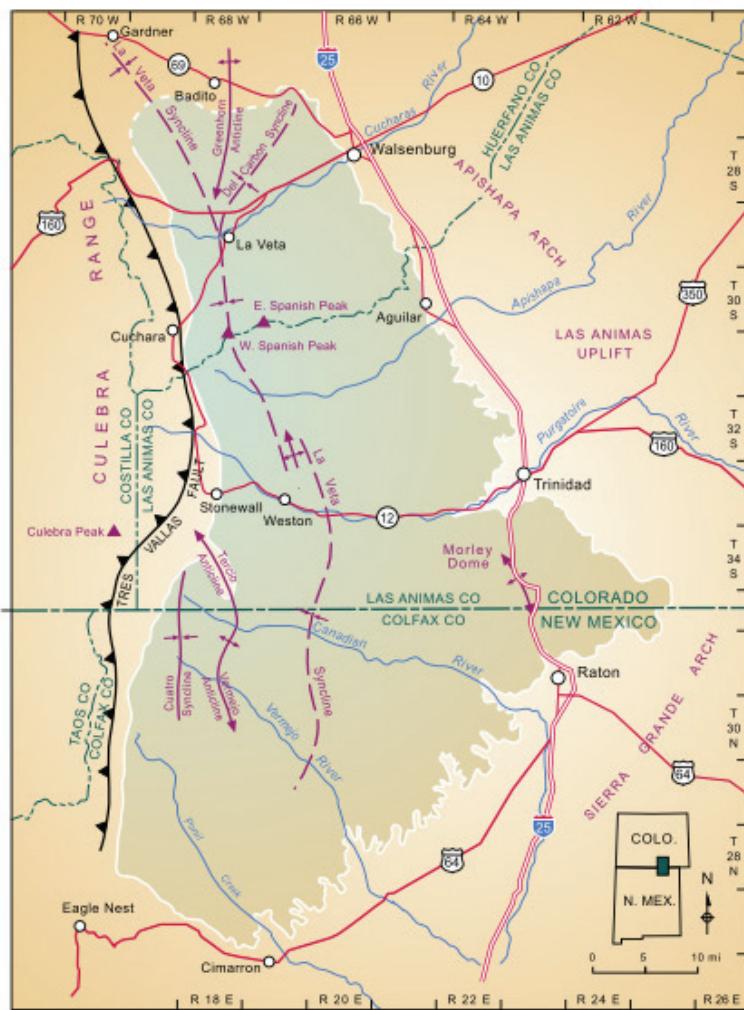


Figure 3. Raton Basin with structural components, rivers, and county lines [Topper, 2011].

The tectonic history of Raton Basin can be divided into four different events [Bohlen, 2012]. Uplift of Raton's crystalline basement rocks occurred in the PreCambrian Era. The formation of the Ancestral Rocky Mountains during the Pennsylvanian and Permian resulted in widespread uplift and tilted block faulting in Colorado. The formation of the Ancestral Rocky Mountains in the early Pennsylvanian originally formed Raton Basin [Bohlen, 2012]. The present - day configuration of Raton Basin was formed during the third phase of tectonic activity, the Laramide Orogeny, during the late Cretaceous through the Early Eocene.

The San Juan Mountains, located to the 230 miles west of Raton Basin, formed during the volcanic period from the Oligocene to Miocene [Bohlen, 2012]. This volcanic activity shaped the present day landscape. There were up to 18 late Tertiary magmatic events. Later east – west extensional tectonic activity formed the Rio Grande Rift [Carlton, 2006]. Post Laramide uplift and dike eruptions led to the creation of the nearby Spanish Peaks (Figure 5) [Bohlen, 2012]. The Spanish Peaks, located in Huerfano County, Colorado, reach heights up to 13,000 ft (3,962 m) and were intruded 22 Ma (Figure 3) [Abbott *et al.*, 1983]. Radial dikes extend from the Spanish Peaks throughout northern Raton and form wall – like features due to differential erosion [Geldon, 1989]. In addition, other Tertiary – aged intrusions have formed dikes in the northern portion of the Basin while sills are prominent in the southern and central regions. The dates for these intrusions range from approximately 32 to 25 Ma (Oligocene to Miocene). During the Miocene to Pleistocene, basin and range block faulting depressed the San Luis Mountains and Rio Grande Valleys [Gabelman, 2005]. The boundaries of Raton Basin, specifically, the Wet Mountains, Apishapa Arch, and Las Animas and Sierra Grande Uplifts were formed during the Tertiary [Baltz, 1965]. Uplift and erosion of the Sangre de Cristo Mountains continued during the Pliocene, which led to a number of formations being deposited in the Basin.

The heat flow anomaly in Colorado is the result of intrusive igneous activity and is the second most areally extensive heat flow anomaly in the U.S. (Figure 4) [Sares *et al.*, 2009]. Heat flow at the edge of the Basin is approximately 100 mW/m^2 while at the center of the Basin it is 200 mW/m^2 [Viegas *et al.*, 2012]. The mean heat flow is estimated to be between 115 and 165 mW/m^2 , which is significantly higher than the High Plains average of $60 - 80 \text{ mW/m}^2$ [Morgan, 2009].

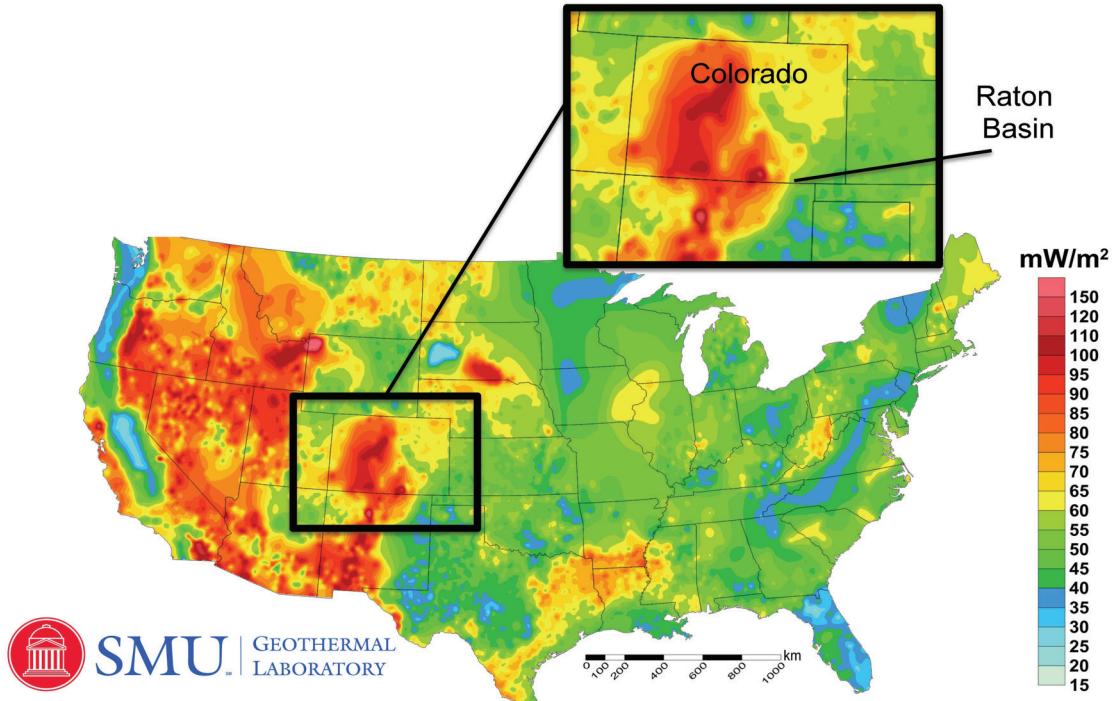
The oldest sedimentary rocks present in the area range in age from the Mississippian to Permian [Osterhout *et al.*, 2015]. The Pennsylvanian to Early Permian was marked by sediment fill from uplift of the San Luis Valley, which lies to the West of Raton Basin. These sediments are known as the Sangre de Cristo Formation. Table 1 shows a stratigraphic column of Raton Basin. The Sangre de Cristo Formation (PRsc) sits above the crystalline basement and is a 1,000 m thick sandstone unit (Figure 5, 6). The Triassic – Permian unit, known as the Dockum Group, is 125 m thick and consists of sandstones and shales. The Triassic Period was tectonically quiet in the region, resulting in the deposition of shallow marine and terrestrial sediment [Osterhout *et al.*, 2014]. The Jurassic aged Morrison (Jmra) and Entrada (Je) Formations are shale, limestone, and sandstone units that are approximately 150 m thick in total. During the Late Cretaceous to Paleocene, the Sevier and Laramide Orogenies contributed to the accumulation of sediment in the Basin, known as the Dakota and Purgatoire Formations [Osterhout *et al.*, 2014]. The Dakota (Kd) and Purgatoire (Kp) are the main targets for injection due to their permeable sandstone lithology and greater depth. Most injection wells target the Dakota Formation to dispose wastewater, but the Purgatoire, Morrison, Entrada and Dockum formations are also included in the screened injection well interval.

| Era | Period | Formation (Members) | Lithology | Average Thickness (m) | Hydrologic Unit (modeled thickness - m) |
|------------|-----------------------|---|-----------------------------------|-----------------------|---|
| Cenozoic | Tertiary | Poison Canyon | Sandstone, Conglomerate, Shale | 381 | Not Modeled |
| | Upper Cretaceous | Raton | Shale, Siltstone, Sandstone, Coal | 304 | Not Modeled |
| | | Vermejo | Shale, Siltstone, Sandstone, Coal | 49 | Not Modeled |
| | | Trinidad | Sandstone, Shale | 40 | Not Modeled |
| | Mid Cretaceous | Pierre Shale | Shale | 600 | 1 (600 m) |
| | | Niobrara (Smoky Hills and Fort Hays) | Limestone, Chalk, Sandy Shale | 300 | 2 (300 m) |
| Mesozoic | Early Cretaceous | Benton (Carlile Shale, Codell, Greenhorn, and Graneros) | Sandstone, Limestone, Shale | 100 | 3 (100 m) |
| | | Dakota | Sandstone | 70 | 4 (1,300 m) |
| | | Purgatoire | Sandstone, Shale | 30 | 4 |
| | Jurassic | Morrison | Shale, Limestone, Sandstone | 109 | 4 |
| | | Entrada | Sandstone | 36 | 4 |
| | Triassic-Permian | Dockum | Sandstone, Shale | 125 | 4 |
| Palaeozoic | Permian-Pennsylvanian | Sangre de Cristo | Sandstone | 1000 | 5 (1,000 m) |
| | Precambrian | Crystalline Basement | Gneiss, Schist, Granite | | 6 (1,700 m) |

Table 1. Stratigraphic column in Raton Basin. Coal bed methane production occurs in the Trinidad, Raton, and Vermejo Formations (shown by red box). Disposal of wastewater occurs in the Dakota, Purgatoire, Morrison, Entrada, and Dockum Formations (shown by orange box). [Geldon, 1989; Bohlen, 2012; Osterhout *et al.*, 2015; Weingarten, 2015]

The retreating Western Interior Seaway during the Cretaceous also created the environments in which the Benton through Raton Formations were deposited [Baltz, 1965]. The Benton (Kb) is Early Cretaceous in age and sits above the Dakota Sandstone (Kd). It is composed of relatively thin sandstone, limestone, and shale [Geldon, 1989]. Average thickness of Benton Formation in Raton Basin is 100 m [Weingarten, 2015]. The Niobrara (Kn) is a 300 m thick limestone, chalk, and sandy shale unit [Geldon, 1989; Weingarten, 2015]. The Mid – Cretaceous aged Pierre Shale and Niobrara Formations are found extensively throughout Colorado. The Pierre Shale (Kpn) is a shale unit that is 600 m thick in Raton Basin [Weingarten, 2015].

The 40 m thick Trinidad Formation (Kt) is composed of sandstone deposits from regressive beach and shelf deposits [Billingsley, 1997]. The Vermejo (Kv) and Raton (TKr) formations produce 98% of the CBM for Raton Basin [Osterhout *et al.*, 2014]. These Upper Cretaceous aged units are composed of shale, siltstone, sandstone, and coal. The overlying Raton Formation is thicker with an average thickness of 300 m while the underlying Vermejo Formation is approximately 50 m thick [Geldon, 1989].



SMU | GEOTHERMAL
LABORATORY

Reference: Blackwell, D.D., Richards, M.C., Frone, Z.S., Batir, J.F., Williams, M.A., Ruzo, A.A., and Dingwall, R.K., 2011, "SMU Geothermal Laboratory Heat Flow Map of the Conterminous United States, 2011". Supported by Google.org. Available at <http://www.smu.edu/geothermal>.

Figure 4. Heat flow map of the United States. Colorado is shown with a zoomed in view with the black box [Blackwell, 2011]

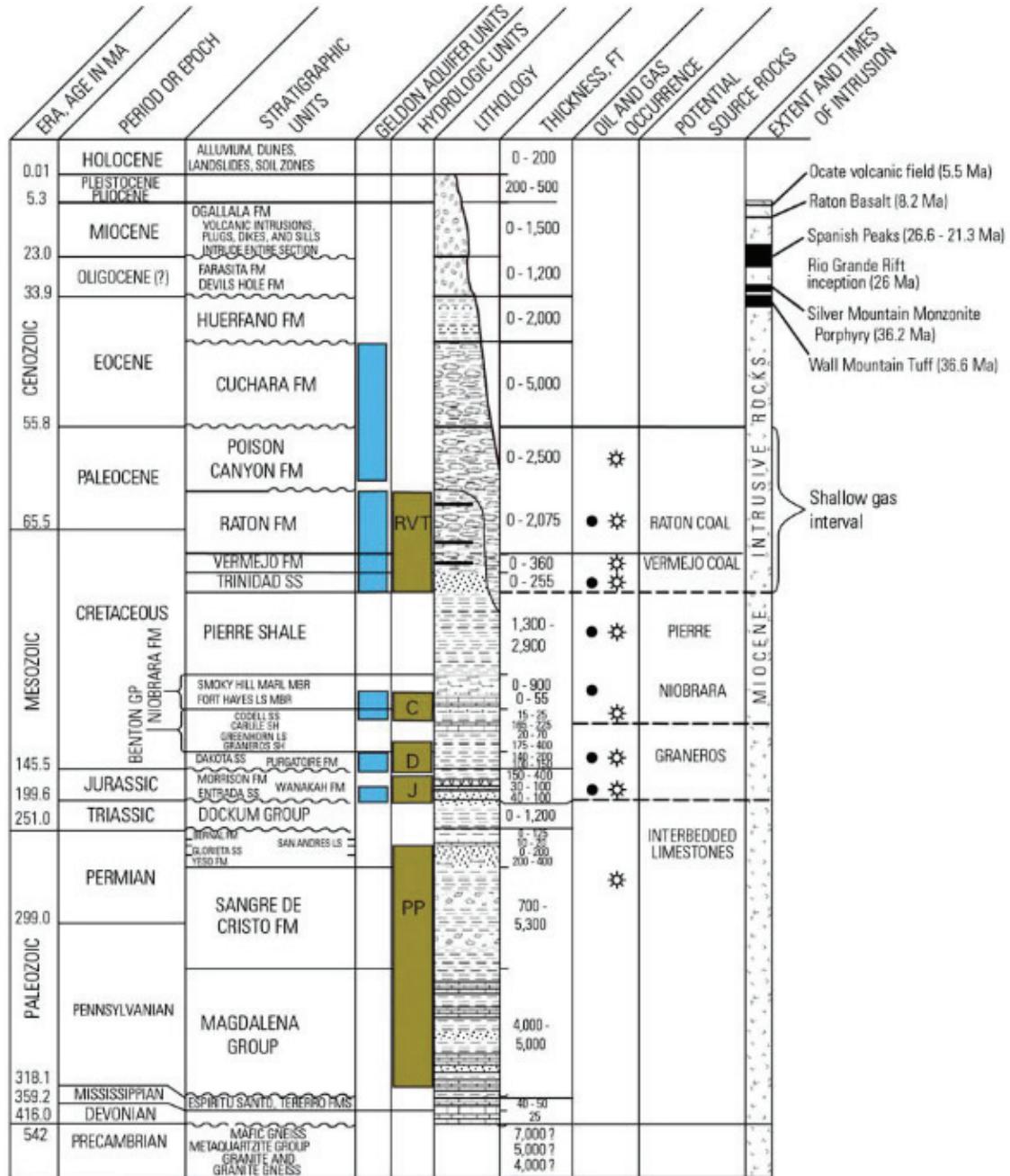


Figure 5. Stratigraphic Column, Petroleum Systems, and Intrusions [Higley, 2005].

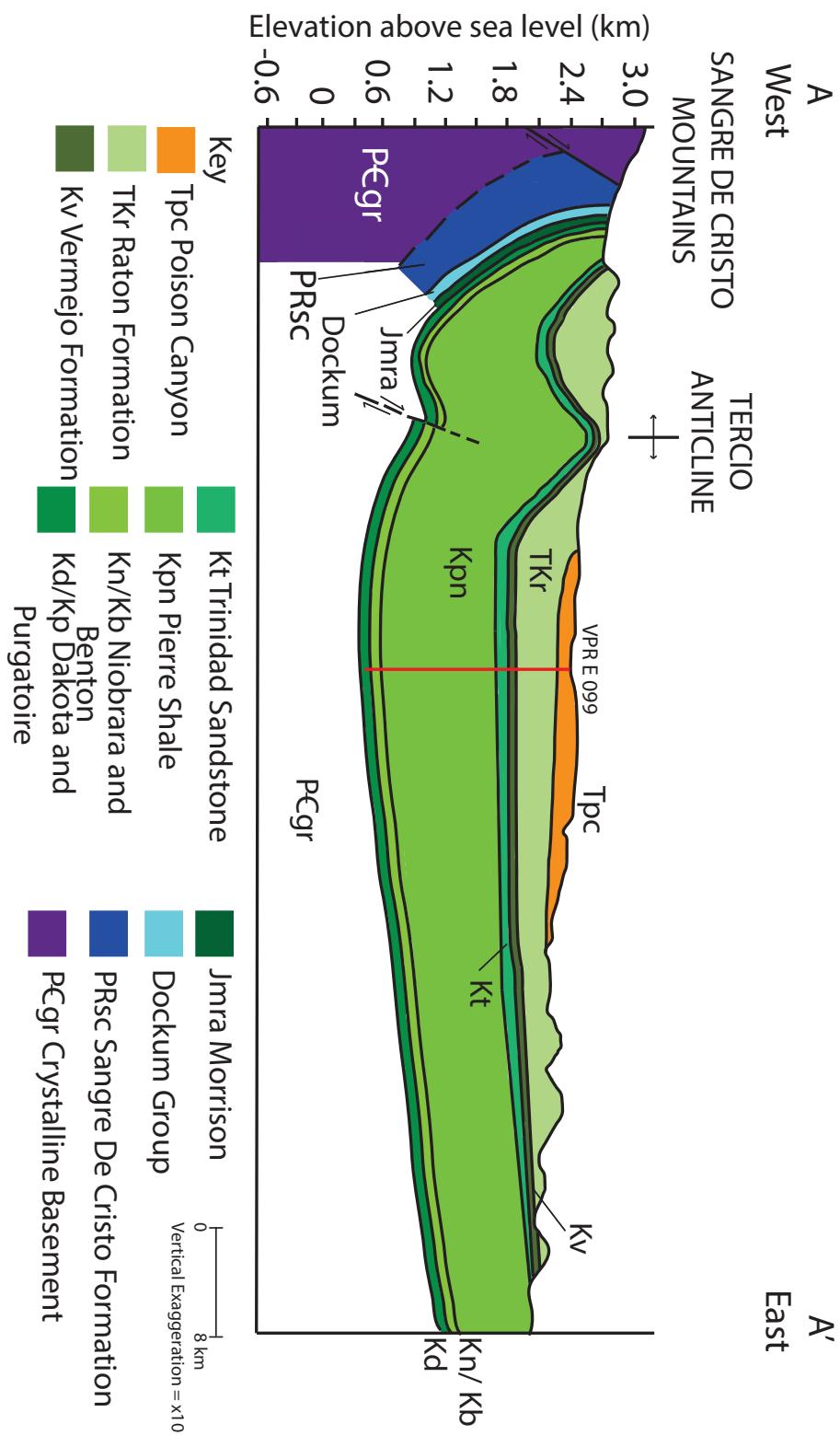


Figure 6. Cross section of Raton Basin. A majority of injection occurs in the Kd (Dakota Sandstone) with seismicity occurring at depths from 3 – 5 km below ground surface in the crystalline basement. VPR E 099 is an injection well in the basin and is shown injecting into the Dakota. Modified from Stevens et al., 1992

2.3 Hydrogeology

Groundwater recharge to bedrock aquifers in the Raton Basin occurs due to precipitation in the upland areas to the west and seepage from alluvium in stream valleys [Geldon, 1989]. The Purgatoire and Apishapa Rivers are the primary rivers and both are tributaries of the Arkansas River (Figure 3) [Geldon, 1989]. Approximately 90% of the CBM production in Raton Basin is within the Purgatoire River watershed located in Las Animas County, Colorado (Figure 3) [Topper, 2011].

The groundwater of the central Raton Basin moves regionally from west to east with local flow being intercepted by stream divides to local valleys (Figure 7) [Geldon, 1989]. Therefore, subsurface flow is north and south inward towards the Purgatoire River drainage. Groundwater flows down dip in the western side of the basin, and up dip in the east, following the synclinal structure of the basin. The movement of groundwater plays a role in the distribution of Raton's geothermal gradient. The average geothermal gradient in Raton Basin is $49 \pm 12 \text{ }^{\circ}\text{C}/\text{km}$ [Morgan and Witcher, 2011]. The highest geothermal gradients are located on the east - central side of the Basin beneath the Purgatoire River drainage (Figure 8). There is a negative correlation between the geothermal gradient and elevation (depth below ground surface (bgs)), which is interpreted as groundwater flow down dip in the west and up dip in the east (Figure 9). This groundwater movement results in thermal convection, which explains the lateral distribution of the geothermal gradient in Raton Basin. This pattern is seen in wells deeper than 1,150 m as well as shallower. Heat transport by groundwater convection extends to depths of at least 2 km in Raton Basin [Morgan, 2009].

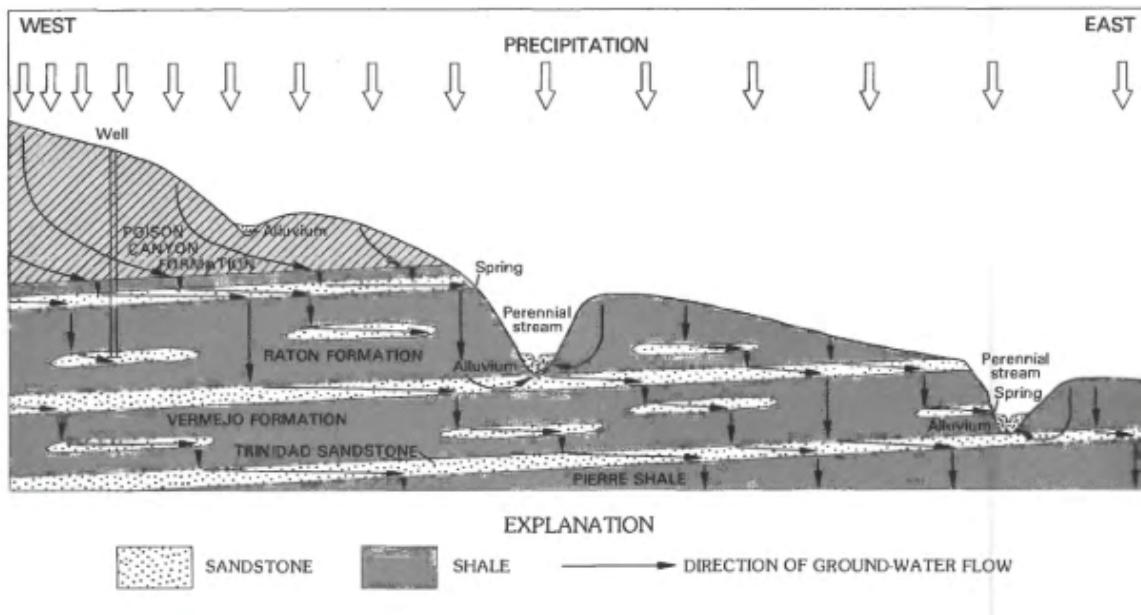


Figure 7. Groundwater flow from west to east in Raton Basin [Geldon, 1989].

There is a projected temperature of 149 °C at 2,438.4 m (300°F at a depth of 8,000 ft) over 960 km² (350 mi²) in Raton Basin. It is worth noting that there is a negative correlation between geothermal gradient and elevation in Raton Basin, which means the deepest part of the Basin, to the West, has a lower geothermal gradient than the eastern part of Raton Basin [Morgan, 2009] (Figure 6). This has been attributed to thermal convection from groundwater. Heat transport by groundwater convection extends to depths of at least 609 ft (2,000 m) [Morgan, 2009].

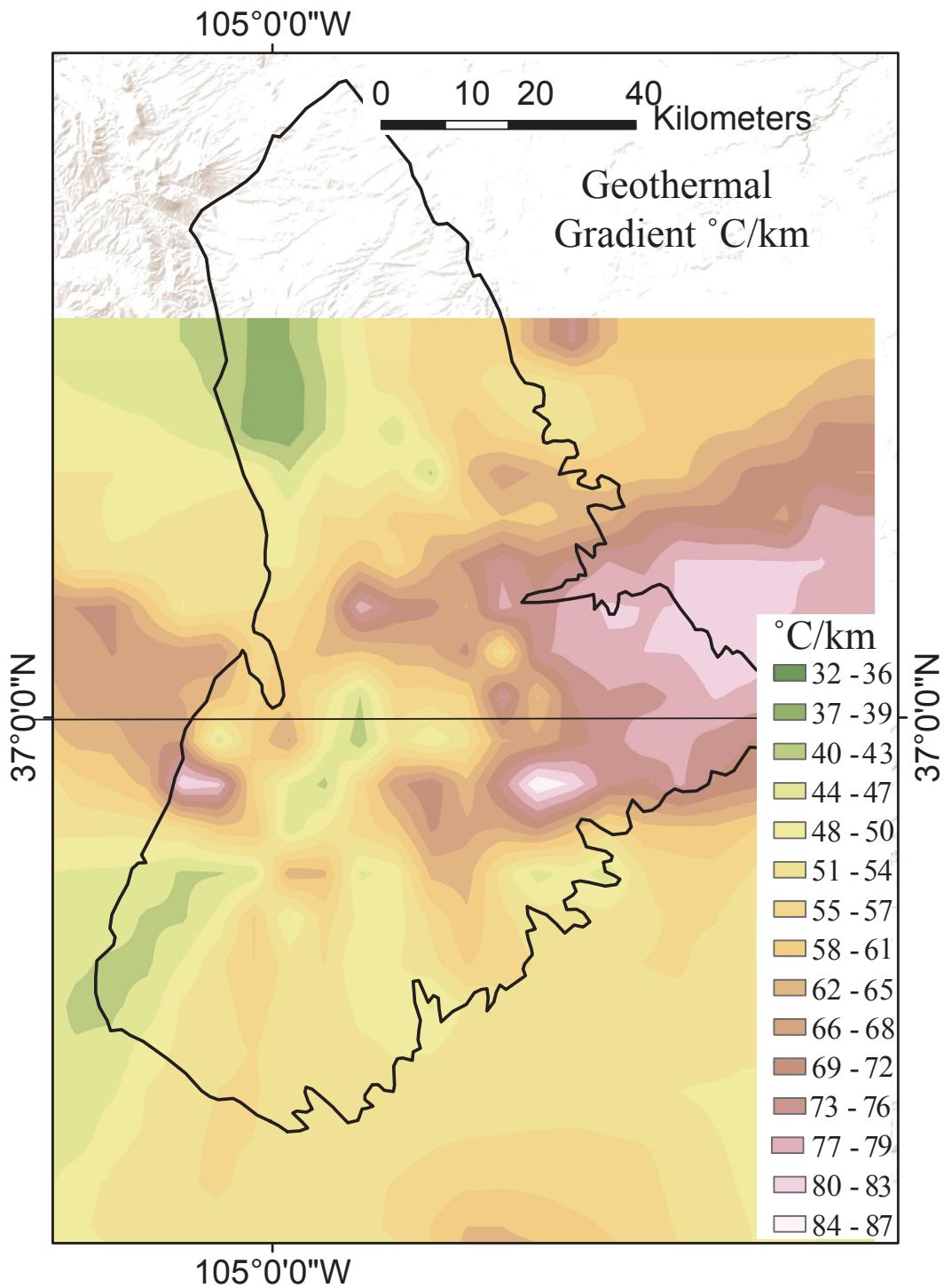


Figure 8. Geothermal Gradient of Raton Basin from BHT measurements [COGCC, 2016, Kelley, 2015 and Kelley, 2011, National Geothermal Data System, 2017, and Morgan, personal comm.].

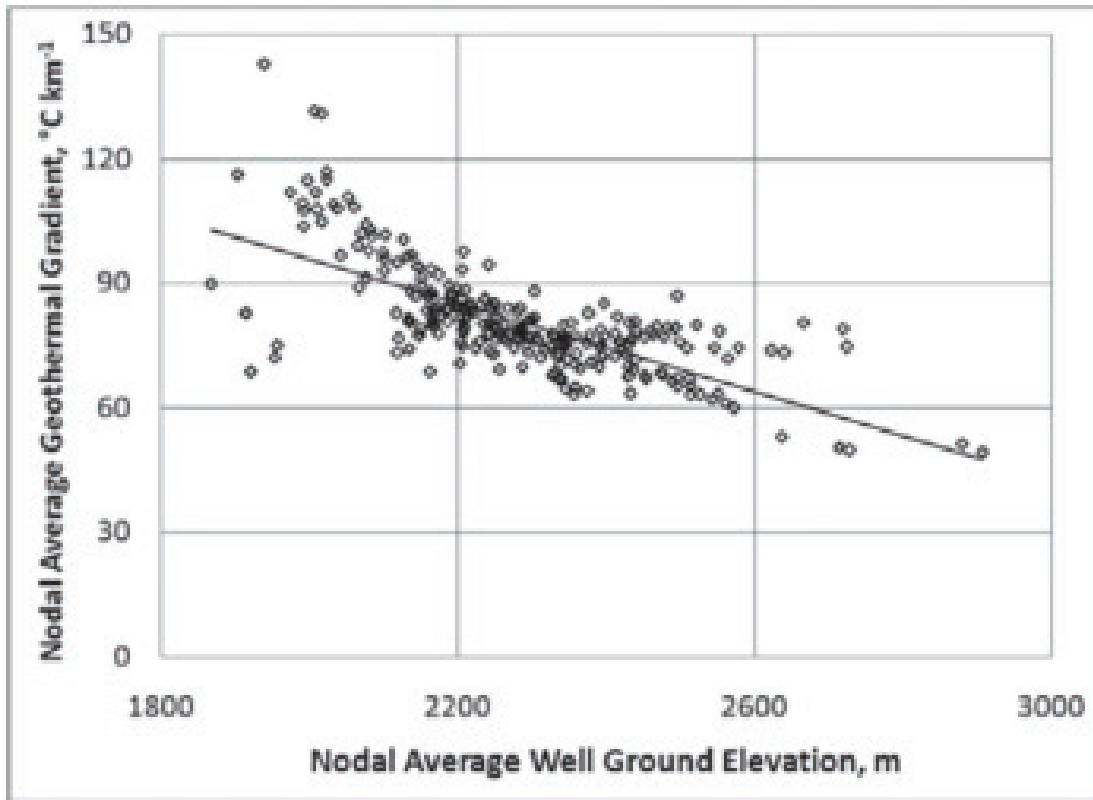


Figure 9. Geothermal Gradient vs. Elevation in Raton Basin. There is a negative correlation between elevation and geothermal gradient, meaning higher elevations have lower geothermal gradients [Morgan, 2011].

The Raton Basin, like many other basins in the area such as the San Luis and Denver Julesburg Basin, is hydrostatically underpressured as discovered by drill- stem tests conducted from 1956 – 1982 [Belitz and Bredehoeft, 1988; Nelson *et al.*, 2013]. Underpressured reservoirs have pressures well below normal values of hydrostatic and lithostatic pressures. The underpressure in Raton Basin is thought to be caused by uplift, erosion, and equilibration of fluid pressure to atmospheric pressure where the formations crop out [Nelson *et al.*, 2013]. Underpressure developed as strata were exposed to Earth's surface by uplift and erosion. Strata allowed for equilibrium of pore pressure in the formation with atmospheric pressure. Low permeability Pierre Shale, Niobrara, and Benton formations and limited up-gradient outcrops and faults in the subsurface restrict recharge. Most induced seismicity across the country has been

attributed to change in pore pressures due to disposal [Rubinstein *et al.*, 2015]. However, because Raton Basin is hydrostatically underpressured, wastewater is gravity fed into the disposal wells, meaning that fluid is drained down the well without any added pressure [Barnhart *et al.*, 2014]. Of the 21 wells on the Colorado side, 5 have been injected under pressure at one point in time (Table 2) [Rubinstein *et al.*, 2014].

2.4 Coal-bed methane

The coal industry was prominent in the region until 2002 when it was no longer economically viable to mine [Hoffman and Brister, 2003]. Raton Basin was divided into two major coalfields, the Trinidad Coal Field and Walsenburg Coal Field named after their respective towns [Tremain, 1980]. The local Spanish Peaks igneous complex and associated intrusions contributed to the elevated coal rank in the area as high heat flow elevated the geothermal gradient (49 ± 12 °C/km) [Carlton, 2006; Morgan and Witcher, 2011]. The natural gas from the coal beds is thermogenic in origin, meaning it contains a higher percentage of non-methane hydrocarbon at higher pressure, temperature, and greater depth [Batz, 1965]. Methane gas is generated from thermal alteration or maturation of the interbedded coals and carbonaceous shale in the area [Dolly and Meissner, 1977].

The extraction of methane occurs after water is pumped out of coal seams to lower the water pressure, which in turn allows the methane to desorb, i.e. be released from the coal seam. The methane then coalesces and bubbles into the pumped water [Weingarten, 2015]. The Raton and Vermejo formations produce 98% of the CBM in Raton Basin [Osterhout *et al.*, 2015] (Table 1). Most the production takes place in Las Animas County, CO with a total number of 2,771 producing wells (Figure 3). The total number of producing CBM wells in Raton Basin is approximately 3,500 [COGCC, 2016].

2.5 Wastewater Disposal

Wastewater is a byproduct of natural gas, oil, and CBM production. The formation brine, also known as produced water or wastewater, is ancient groundwater that has been trapped in the same pore spaces as hydrocarbons [Rubinstein and Mahani, 2015]. This salty water is produced during oil and gas extraction. Formation water typically contains dissolved minerals or other materials that make it unsuited for agricultural or human uses [Rubinstein and Mahani, 2015].

Wastewater injection typically takes place in high-permeability formations that are deeper than the producing formations, which is the case in the Raton Basin as the Dakota Formation is located deeper than the CBM producing formations (Table 1). The wastewater injection is typically isolated from the fluid from both drinking-water aquifers and oil and gas producing reservoirs [Rubinstein and Mahani, 2015].

Wastewater injection wells are known as Class II wells, as defined in the Underground Injection Control Program (UIC) in association with the Safe Drinking Water Act [EPA, 2018]. Class II injection well depths in Raton Basin range from 1,250 m to 2,100 m [COGCC, 2016 - Appendix]. Most wells target the Dakota Formation, but the underlying Purgatoire, Morrison, Entrada and Dockum Formations are also included in the injection interval [COGCC, 2016 - Table 1 and Figure 6]. Wastewater disposal began in 1994 in the Colorado portion of the Basin and expanded into New Mexico in 1999 (Table 2). Disposal volume began to increase significantly in 2001. There were ten disposal wells in 2001, which increased to 21 in 2002 [Viegas *et al.*, 2012]. As of December 2016, there are 28 active disposal wells and more than 3,000 CBM producing wells [COGCC, 2016; NMOCD, 2016]. Of the 28 disposal wells, 22 are located within the Purgatoire River field in Las Animas County, CO and the other six are in the Raton Coal Field in Colfax County, NM (Figure 3). It is worth noting that the disposal well

information for the New Mexico portion of the Basin is unavailable prior to June 2006. Six wells have injected waster as opposed to gravity disposal from 1999 – 2016 (Table 2).

Table 2. Raton Basin Salt Water Disposal Well Information, New Mexico and Colorado. Wells that have injected under pressure at one time are shown by *. Wells that outside of the basin outline are marked with ~ .

| Well Name | Disposal Start Date | Years Active | Average Disposal Rate (bbl/month) | Maximum Disposal Rate (bbl/month) | Cumulative Disposal Volume (barrels) |
|--------------------------------|---------------------|--------------|-----------------------------------|-----------------------------------|--------------------------------------|
| Jarosa 32-33 WD | May-07 | 2007-2016 | 149,187 | 334,640 | 17,305,692 |
| SouthPaw | Apr-09 | 2009-2016 | 124,516 | 282,880 | 12,451,600 |
| Cottontail Pass Disposal 32-33 | Nov-94 | 1999-2016 | 76,945 | 448,374 | 16,543,178 |
| Cimarron 32-18 WD | Jan-01 | 2005-2016 | 29,910 | 157,207 | 4,247,178 |
| Ferminia | Sep-07 | 2007-2016 | 65,894 | 146,789 | 7,380,101 |
| Sawtooth | 4/1/00 | 2000-2016 | 45,942 | 131,967 | 9,234,370 |
| La Garita | Aug-01 | 2001-2016 | 29,648 | 203,122 | 548,914 |
| San Pablo 11-4 WD | Dec-14 | 2014-2016 | 138,232 | 249,592 | 4,976,383 |
| Weston 24-23 A WD | Jan-04 | 2004-2016 | 93,723 | 211,903 | 14,620,852 |
| Long Canyon 43-12 WD | Apr-01 | 2004-2016 | 81,908 | 166,039 | 15,480,652 |
| Apache Canyon 19-10* | Jan-95 | 1999-2016 | 46,012 | 265,405 | 9,892,512 |
| Apache Canyon 10-3* | Jan-95 | 1999-2016 | 14,960 | 69,382 | 3,231,394 |
| VPR C 204 WDW | Mar-12 | 2012-2016 | 145,839 | 230,500 | 6,854,448 |
| Garcia (EPA) 1-WD ~ | Jan-99 | 1999-2016 | 1,240 | 4,800 | 262,948 |
| PCW | Jul-97 | 1999-2016 | 115,823 | 327,572 | 25,017,857 |
| Wild Boar | Aug-00 | 1999-2016 | 118,928 | 491,058 | 23,428,799 |
| Hill Ranch Deep | Jul-05 | 2005-2016 | 17,090 | 67,081 | 2,358,403 |
| Beardon | Jan-01 | 2001-2016 | 67,316 | 593,835 | 12,386,139 |
| Polly | Jul-09 | 2009-2015 | 4,014 | 74,285 | 389,438 |
| Del Agua | Jul-05 | 2005-2012 | 107,767 | 196,239 | 9,160,226 |
| Lopez Canyon SWD* | Sep-10 | 2010-2016 | 6,236 | 33,871 | 461,434 |
| VPR A 007 | Jun-06 | 2006-2016 | 82,195 | 237,672 | 10,438,782 |
| VPR C 39 | May-00 | 2000-2016 | 115,757 | 356,826 | 21,993,787 |
| VPR C 14* | Sep-99 | 1999-2002 | 84,416 | 378,554 | 17,474,060 |
| VPR A 182* | Jun-06 | 2006-2016 | 324,903 | 485,190 | 41,262,655 |
| VPR A 500 | Jun-08 | 2006-2016 | 125,620 | 228,404 | 12,938,869 |
| VPR B 027 | Jun-06 | 2006-2016 | 63,260 | 115,768 | 8,033,957 |
| VPR D 025 | Jun-06 | 2006-2016 | 181,534 | 714,590 | 23,054,880 |
| VPR A 042 | Jun-06 | 2006-2016 | 126,944 | 387,704 | 16,121,901 |
| VPR E 099* | Jun-06 | 2006-2016 | 258,602 | 367,588 | 32,842,506 |

2.6 Seismicity

Induced seismicity has been a focus of scientific research for the past few decades [Ellsworth, 2013; Rubinstein and Mahani, 2015]. At least 50% of the 4.5 magnitude (M) or larger earthquakes in the interior of the U.S. during the past decade are in areas with injection induced seismicity [van der Elst et al., 2013]. There are a number of ways in which fluid injection can cause earthquakes. The most widely researched includes the increase of pore-fluid pressure along a fault, which lowers the effective stress resulting in fault failure. The Mohr – circle can be used to explain the reduced angle of internal friction [Figure 10 – Brown, 2018, Rubinstein and Mahani, 2015]. Deformation may occur when fluids are compressed within pore spaces resulting in poro-elastic effects [Rubinstein and Mahani, 2015; Foulger et al., 2017]. Poroelasticity deals with the interaction of fluid flow and solids deformation within a porous medium. An external load affects the volume fraction of pores. Fluid-filled pores experience an increase in pressure due to this mechanical stress of pumping in wastewater. This leads to fluid migration. The pore volume is changed which shifts the solid material and causes deformation [Comsol, 2017]. Thermoelastic deformation, specifically compressions, may result when cool fluid is injected into a much warmer rock [De Simone et al., 2013; Gaucher et al., 2015].

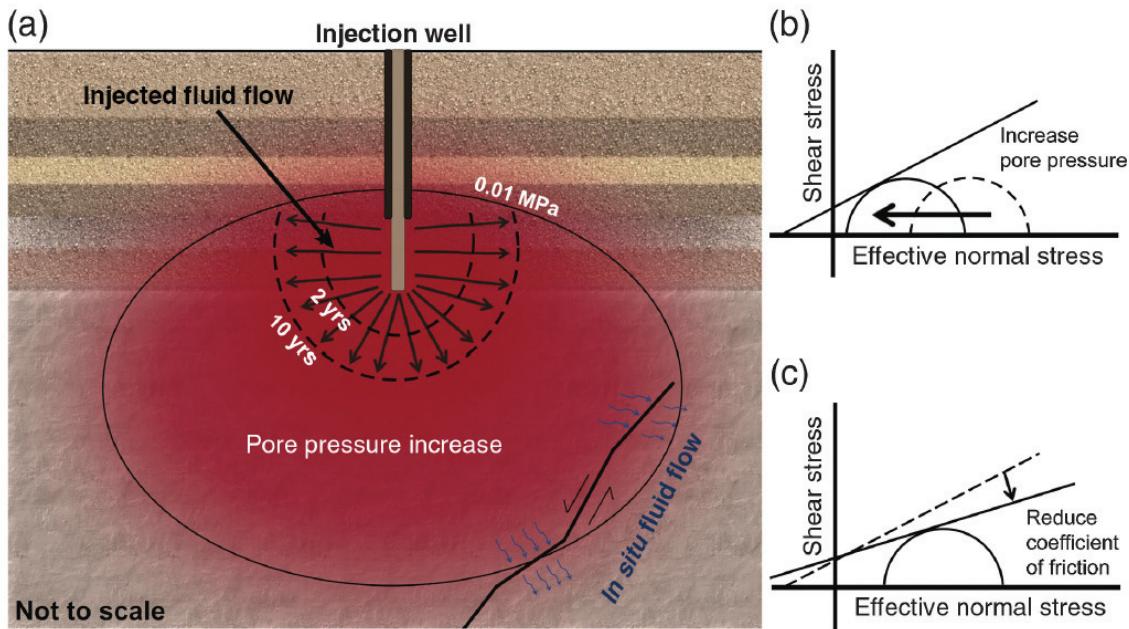


Figure 10. Conceptual figure of the mechanisms of fluid induced seismicity. a) Red area is pore pressure increase from disposal with a contour of 0.01MPa after 10 years. The dashed lines are the area that injected fluid flow reaches after 2 years of injection and 10 years. A fault is shown in the solid black line with blue arrows representing in situ fluid flow near the fault. b) Mohr – Coulomb diagram triggering via pore pressure increase along a pre – existing fault that is optimally oriented. c) Mohr – Coulomb showing induced seismicity from reduction of coefficient of friction from in situ fluid flow caused by injection induced seismicity pore pressure gradient [Brown, 2018].

In the United States, there are approximately 35,000 active wastewater disposal wells. However, only a few dozen wells have been proven to induce earthquakes [Rubinstein and Mahani, 2015]. The issue of fluid injection induced earthquakes was first documented in Colorado during the 1960s [Evans, 1966]. Rocky Mountain Arsenal (RMA), located in northeastern Denver, was the site of chemical waste disposal into the subsurface with a single injection well. Fluids were injected into the heavily fractured Precambrian metamorphic rocks. As a result of injecting a large amount of liquid waste (approximately 4.7 Mbbl), RMA experienced earthquakes that were unknown to the area prior to injection. In fact, prior to 1962 the Denver area had not experienced a noticeable earthquake since November 7, 1882 [Evans,

1966]. A total of 710 earthquakes were recorded in the Denver area from March 1962 to October 1965, with a majority of epicenters within five miles of the RMA injection well [Evans, 1966]. Other seismic events were located at distances greater than 10 km and at depths greater than four km from the RMA's well [Rubinstein and Mahani, 2015].

In the case of Raton Basin, the injection interval includes the Dakota, Purgatoire, Dockum, and Entrada formations, and the majority of wastewater is injected into the Dakota. In contrast to the RMS, fluids are not dumped directly into Precambrian basement rock. An increase in the injection of produced wastewater began in 1999 in Raton Basin with seismicity beginning in 2001 (Figure 11). In 1999 there were 22.9 million barrels ($3,640,809 \text{ m}^3$) of wastewater produced, increasing to 119 million barrels ($1.89 \times 10^7 \text{ m}^3$) in 2007 [Topper, 2011]. The closest known active faults are within the Quaternary sediments present in the western part of the Basin [Rubinstein et al., 2014]. There are a number of faults in the crystalline basement, with some cutting into the permeable overlying sedimentary layers (Figure 6).

Historical and instrumental data show that the increase in seismicity in Raton Basin is temporally correlated to wastewater injection [Rubinstein et al., 2014; Barnhart et al., 2014; Meremonte et al., 2002]. Between 1972 and 2001 only one $M \geq 4$ earthquake occurred in the area. However, between August 2001 and 2013 there were 12 $M \geq 4$ recorded earthquakes in the Raton Basin [Rubinstein et al., 2014]. Moment magnitude (M) is a measurement of the earthquake size or energy. It is similar to the Richter magnitude but is more accurate in describing the size of earthquakes [Folger and Tiemann, 2015].

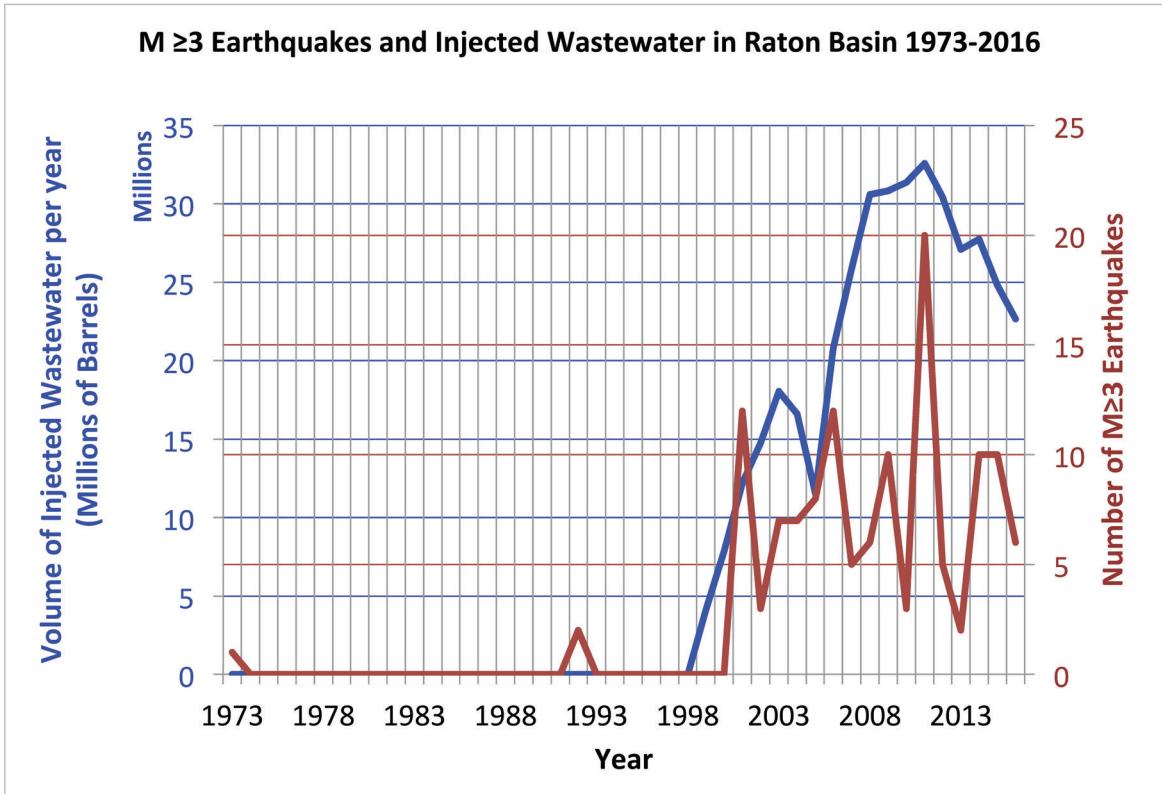


Figure 11. Earthquakes $M \geq 3$ and Volume of Injected Wastewater in Raton Basin from 1973 and 2016 [COGCC, 2016; ANSS, 2016]

The largest earthquake occurred on August 23, 2011 when a M_w 5.3 earthquake damaged buildings in the town of Trinidad, CO [Viegas *et al.*, 2012; Barnhart *et al.*, 2014] (Figure 13). Approximately 41 million barrels (approximately 4.9 million m^3) of wastewater was injected from two disposal wells (VPR C 39 and VPR C 14 – Figure 24) before the earthquake, which is enough water to fill 1,960 Olympic sized swimming pools and seven times the total amount disposed of at RMA in 4 years [Rubinstein *et al.*, 2014]. The United States Geological Survey (USGS) located 584 aftershocks between the August 23rd event and December 15th, 2011 (Figure 14) [Barnhart *et al.*, 2014]. These series of earthquakes are referred to as an earthquake swarm. Raton has had two major earthquake swarms in the same area in 2001 and 2011. Meremonte et al. [2002] discovered a spatial correlation between the location of brine reinjection from CBM

and the earthquake swarm in August 2001. A majority of the events were recorded within five km of active disposal wells [Meremonte *et al.*, 2002]. Rubinstein *et al.* [2014] noted that seismicity tended to occur within two km of high volume and high injection rate wells. Weingarten *et al.* [2015] defined high rate injection wells as having injection rates greater than 300,000 barrels/ month, which is higher than the injection rates in Raton Basin. Viegas *et al.* [2012] showed that wells do not have to cross a fault to trigger an earthquake and may be up to six km away from the epicenter.

Most induced seismicity has been attributed to changes in pore pressure because of the injection pressure at which the wastewater is disposed. However, Raton Basin is hydrostatically underpressured so the wastewater is gravity fed in the disposal wells [Barnhart *et al.*, 2014]. Although, studies have shown that even though the Basin is underpressured, pore pressure may still increase due to wastewater disposal [Weingarten *et al.*, 2015]. However, Seismic events have occurred prior to the commencement of wastewater injection in the Basin suggesting natural causes, such as ambient tectonic activity. For example, earthquakes in 1966 and 1973 in Raton Basin are two examples attributed to tectonic activity [COGCC, 2016] (Figure 11).

There are differences between induced seismicity in geothermal reservoirs versus wastewater disposal. Geothermal reservoirs will stimulate reservoirs on purpose in order to create permeable rock that allows fluid to flow. Therefore, they are aiming to reduce rocks strength and therefore microseismicity. Wastewater disposal areas are not aiming to stimulate permeability, as they are already disposing in formations with high permeability and porosity. In addition, the location of geothermal reservoirs and disposal sites are different. Geothermal reservoirs are normally in tectonically active areas that are more prone to seismicity, unlike the center of continents where wastewater disposal occurs [Majer *et al.* 2007].

The mechanisms for induced seismicity in geothermal reservoirs and wastewater disposal are similar. Pore pressure increase is observed from both disposal wells and geothermal injection wells. Geothermal reservoirs add another component that wastewater disposal sites do not – production. Water is also withdrawn from the reservoir, which can create volume changes and affect the failure (Figure 12) [Majer *et al.* 2007].

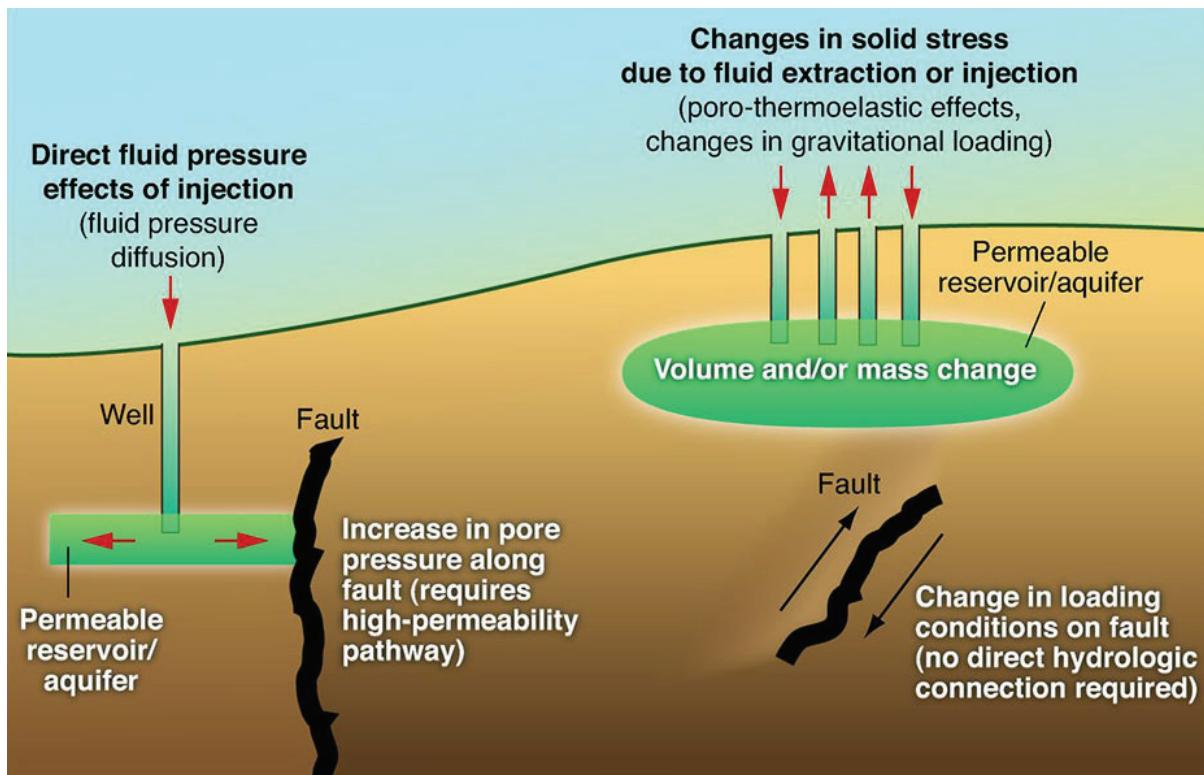


Figure 12. Mechanisms for induced seismicity. Wastewater disposal is illustrated on the left into a permeable reservoir while enhanced geothermal reservoirs are illustrated with the mass/volume change in a reservoir [Ellsworth *et al.*, 2003].

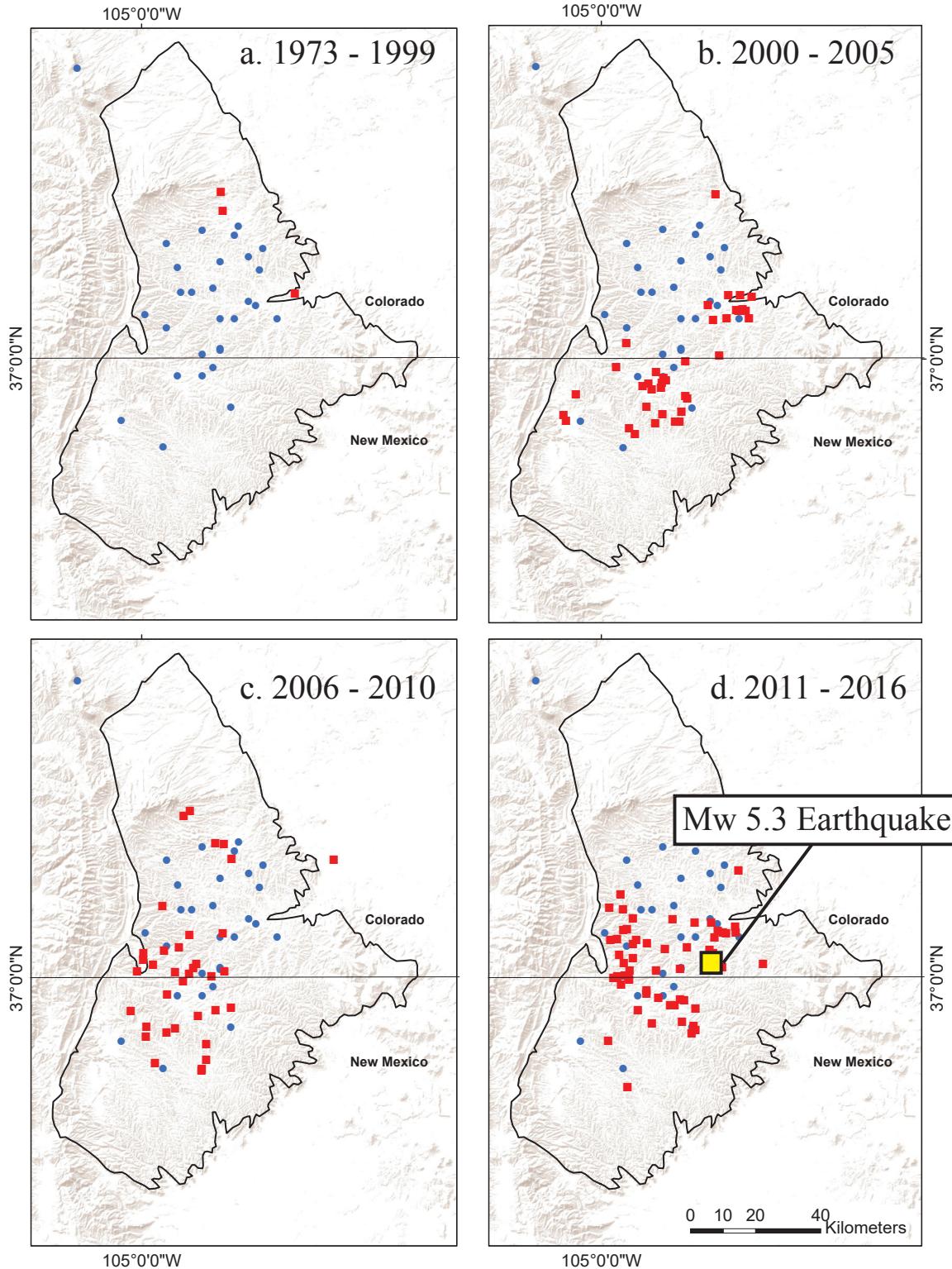


Figure 13. Map showing $M \geq 3$ Earthquakes in Raton Basin (red squares) though time. Injection wells are shown in blue. The Mw 5.3 earthquake is shown in yellow [COGCC, 2016; ANSS, 2016].

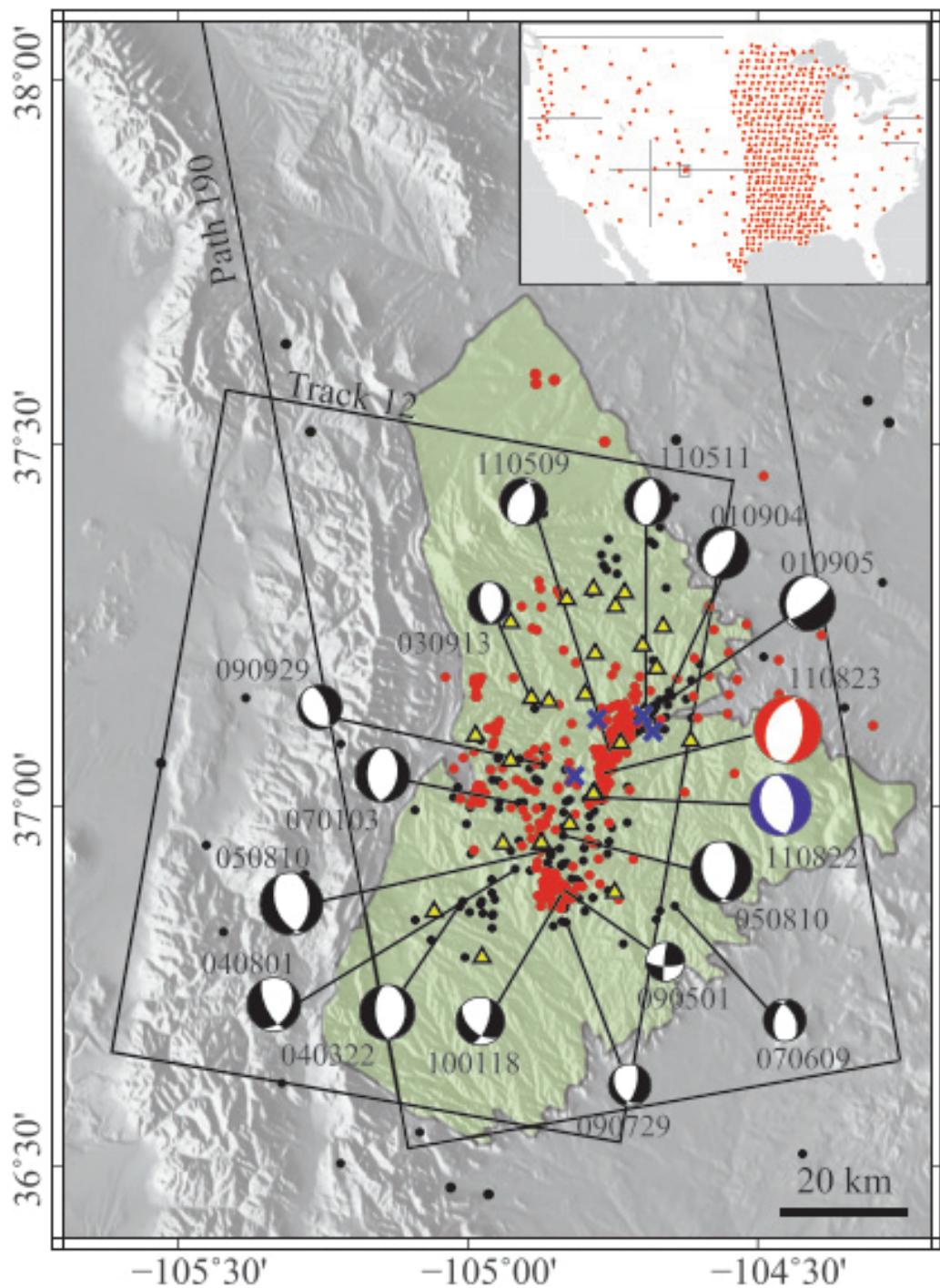


Figure 14. Regional moment tensors for seismicity in Raton Basin. Black dots are seismicity prior to the 2011 Trinidad earthquake (shown as the red moment tensor). Red dots are seismicity that Rubinstein et al. 2014 located. Blue crosses are temporary seismometers that the USGS deployed after the 2011 Mw 5.3 earthquake. Yellow triangles are the wastewater injection wells. Black outlined boxes are spatial extends for InSAR. [Barnhart et al., 2014.]

2.7 Geothermal Gradient and Enhanced Geothermal Systems

There is much debate about the causes of the recent seismicity in Raton Basin. The Raton Basin has an average geothermal gradient of $49 \pm 12^{\circ}\text{C/km}$, which is much higher than the average for the western United States (34°C/km) [Morgan and Witcher, 2011; Nathenson and Guffanti, 1988]. Although some studies have shown a spatial and temporal correlation between wastewater injection and seismicity, no unanimity as to the origin of the Basin's seismic activity has been reached [Rubinstein et al., 2014; Barnhart et al., 2014; Meremonte et al., 2002] (Figure 11). Most research into thermal effects on induced seismicity is focused on Enhanced Geothermal Systems (EGS) [Foulger et al., 2017, Gaucher et al., 2015, Brodsky et al., 2013]. EGS are reservoirs in areas of high geothermal gradients that are stimulated by either hydraulic fracturing, water injection, or chemical stimulation to extract heated water for energy production [Gaucher et al., 2015, Majer et al. 2007]. Water is injected into the fractured rocks in the subsurface, which is then heated by surrounding warm rock, and is then extracted in separate producing wells (Figure 15). The goal is to enhance and create permeable pathways for fluids to flow. This process creates microearthquakes ($M < 3$) in most cases, but some larger earthquakes have taken place, such as the $M 4.6$ that occurred in The Geysers field in northern California in the 1980s [Majer et al. 2007, U.S. DOE, 2006].

The injection of cool water into much warmer rock results in thermal contraction and thus rock failure [De Simone et al., 2013]. The temperature difference between the injected water, at atmospheric conditions at the surface, and hot geothermal reservoir causes thermal contraction of the fracture surfaces, which affects the in situ stress state [Majer et al., 2007] The higher temperature of Raton Basin, along with wastewater injection of fluids that are significantly colder than the surrounding rock, could lead to thermal contraction. Temperature variation

affects fluid viscosity and deforms the rock mass thermoelasticity [Holtzman *et al.*, 2018]. Hydrologic parameters, such as porosity and permeability, can be changed due to stress and strain due to increased water pressure. Increased fluid pressure extends farther into fractured zones than in surrounding rock due to larger contrasts in the hydraulic conductivity. Temperature decrease lowers the total stress in addition to the effect stress by the thermo-elastic effect. This can then lead to tensile or shear failure [Brodsky *et al.*, 2013; De Simone *et al.*, 2013; Gaucher *et al.*, 2015].

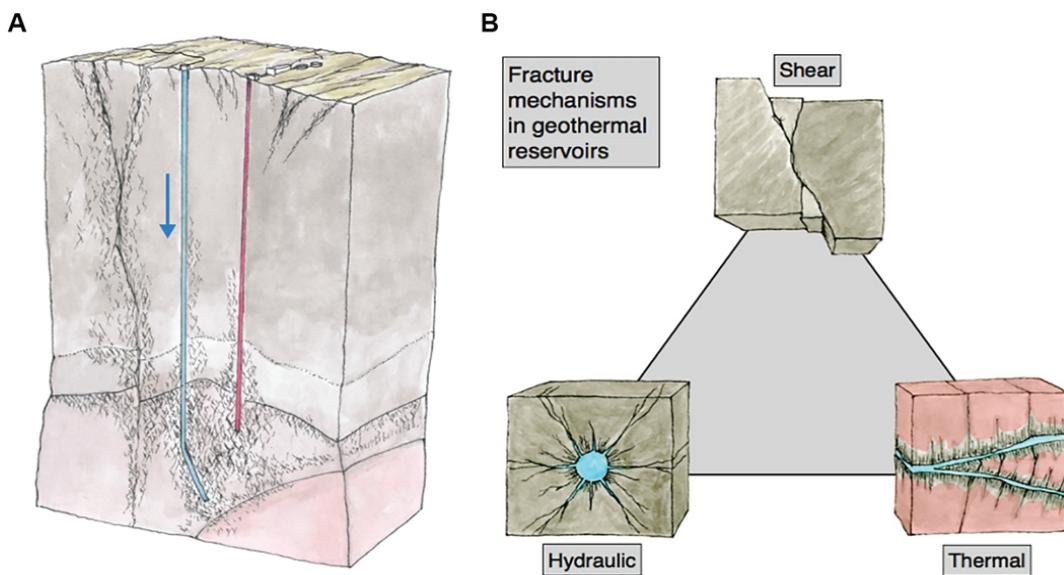


Figure 15. Mechanisms for Induced seismicity in EGS [Holtzman, 2018].

Elevated temperatures enhance the hydraulic conductivity and allow for fluids to migrate more easily [Weingarten, 2015]. Contraction and expansion of rock, resulting from temperature changes in both rock and fluids, create poro-thermoelasticity effects, driving changes in porosity, intrinsic permeability and hydraulic conductivity [De Simone *et al.*, 2013; Gaucher *et al.*, 2015]. Temperature decreases in rock due to cooler water injection, combined with an increase in pore pressure, can potentially cause rock failure due to the lowering of the effective normal stress, as can be shown by a Mohr circle [De Simone *et al.*, 2013; Gaucher *et al.*, 2015] (Figure 5,

Rubinstein et al., 2015). In addition to contraction, expansion of rock due to higher temperatures can cause pore pressure to increase due to confinement.

In order to address the relationship between the geothermal gradient and seismicity in Raton Basin, modeling was performed. Prior to modeling, multiple datasets were collected from a variety of sources. Disposal well information was collected from state websites for both New Mexico and Colorado. Secondly, geothermal data were collected from bottom hole temperatures. Pioneer Natural Resources provided wastewater disposal temperature while formation lithologies and parameters were collected from literature. Lastly, earthquake locations were collected from publically available earthquake catalogs.

3. DATA ACQUISITION and ANALYSIS

3.1 Injection Wells

Wastewater injection well information was for Raton Basin. The Colorado Oil and Gas Conservation Commission (COGCC) database provided injection rates, locations, current status, and dimensions for the injection wells on the Colorado side of the Basin [COGCC, 2016 - Appendix]. The New Mexico injection well information was collected from the New Mexico – Oil Conservation Division (NMOCD) [NMOCD, 2016]. Well logs from disposal wells and producing wells also provided top and bottom depths of formations and porosity in the Basin. In 2001, there were 10 injection wells, which increased to 21 in 2002 [*Viegas et al.*, 2012]. As of December 2016, there are 28 active injection wells (Table 2) [COGCC, 2016; NMOCD, 2016]. Note that injection well data prior to 2006 are not available for New Mexico injection wells and thus New Mexico injection rates were not included from 1999 to 2006 in the model study described later in this thesis.

3.2 Geothermal data

Bottom Hole Temperatures (BHT) data from the bottom of the producing CBM and disposal wells in the area were collected [Kelley, 2015 and Kelley, 2011, National Geothermal Data System, 2017, and Morgan, personal comm.]. This allows calculation of the spatially varying geothermal gradient (Figure 7). BHT are typically corrected because the measurement is taken within a few hours of drilling, which results in inaccurate temperatures due to drilling fluid convecting heat into or out of the hole [Morgan, 2009]. BHT are corrected using a standard correction equation to eliminate discrepancies between the temperatures recorded during drilling and days after drilling was completed. We corrected the BHT data using the following equation [Kelley, 2015].

$$\text{BHT}_{\text{corrected}} \text{ } ^\circ\text{C} = \text{BHT}_{\text{uncorrected}} \text{ } ^\circ\text{C} + 0.0056125 \times \text{Depth (m)} + 3.369$$

The COGCC provided BHT for Colorado side, while New Mexico Bureau of Geology and Mineral Resources through The National Geothermal Data Systems (NGDS) provided BHT data for the New Mexico side [Kelley, 2011; Kelley, 2015; COGCC, 2016,]. In total, 1,627 BHT data measurements were used to interpolate the geothermal gradient spatially for the model in this study.

3.3 Injection Temperature

The surface temperature of disposed wastewater was provided by Pioneer Natural Resources, a main operator in Raton Basin. The average temperature is 20°C. The temperatures were sampled between 2000 and 2014 at surface disposal tanks at different times of the year (Figure 16). These tanks store the wastewater before it is transported for disposal in the subsurface [Macartney, personal comm.].

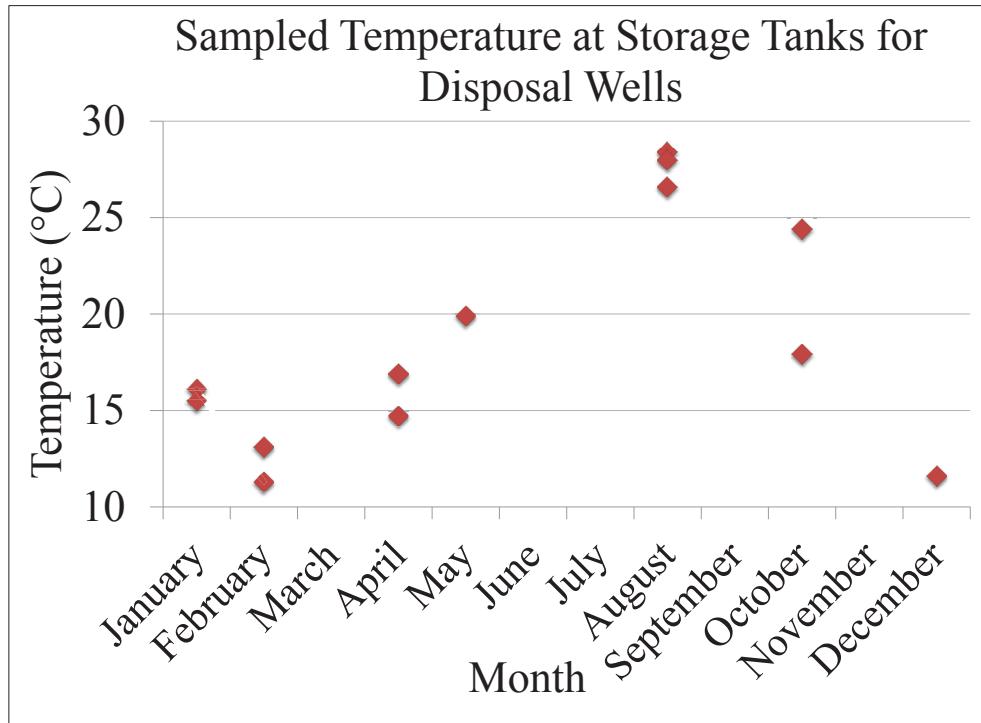


Figure 16. Wastewater temperature recorded at storage tanks on the surface prior to injection. There were collected from 2000 – 2014. Appendix has the years that each were sampled. [Pioneer Natural Resources; McCartney, personal communication]

3.4 Formation Geology

Formation lithologies were used to estimate porosity and permeability for the model.

Porosity values for the injection intervals were taken from various injection well logs from the COGCC [COGCC, 2016]. Permeability from core and outcrop provided permeability values for formations [Bohlen, 2012; COGCC, 2016]. Literature values were also used when describing the hydraulic parameters of the different model layers, which varied due to lithology differences. These include density (kg/m^3), wet heat conductivity (saturated heat conductivity) ($\text{W}/\text{m} \cdot ^\circ\text{C}$), specific heat of the rock ($\text{J}/\text{kg} \cdot ^\circ\text{C}$), pore compressibility ($1/\text{Pa}$), and pore expansivity ($1/\text{^\circ C}$) [Robertson, 1988 - Appendix].

The CBM producing formations - Raton, Vermejo, and Trinidad - have an average porosity of 14% and permeability ranging from <0.1 - 28.5 millidarcys (md) ($1 \times 10^{-16} \text{ m}^2$ - $2.81 \times 10^{-14} \text{ m}^2$) [Dolly and Meissner, 1977] (Table 1). These formations are not modeled in this study because they shallower then the depths that wastewater injection is occurring. The Pierre Shale, Niobrara and Benton Formations act as a confining layer to the underlying injection interval because of their low permeability. These formations' permeability ranges from 0.00475 md – 0.4 md ($4 \times 10^{-18} \text{ m}^2$ to $4 \times 10^{-16} \text{ m}^2$) and restrict upward migration of groundwater from the injection interval into overlying formations. Porosities in the injection interval range from 5 to 16% with permeability ranging from 2 to 20 md ($1.97 \times 10^{-15} \text{ m}^2$ – $1.97 \times 10^{-14} \text{ m}^2$ [COGCC, 2016]). Parameters used in the numerical modeling to represent these formations are discussed in Chapter 4.

3.5 Earthquakes

Earthquake locations and magnitudes in Raton Basin from 1973 to 2016 are shown in Figure 17. Earthquake data were collected from the Advanced National Seismic System (ANSS), which is operated by the United States Geological Survey (USGS) [ANSS, 2016]. Only earthquakes $M \geq 3$ were examined because that is the magnitude at which humans can feel the effects at the surface. In total, this catalog has recorded 142 $M \geq 3$ earthquakes in Raton Basin from January 1999 to April 2018, of which 130 of the 142 occurred from January 1999 to December 2016. Seismicity is on going in Raton Basin with the most recent earthquake (as of June 17, 2018) a M 3.5 occurred on February 25th 2018. Depths of seismicity epicenter from ground surface range from 3 to 6 km and thus primarily occur in the crystalline basement. Fault locations can be assumed from the locations of earthquakes as earthquakes highlight the location of preexisting faults.

Other earthquake logs were examined to compare depth uncertainty. These include Rubinstein et al. [2014], Barnhart et al., [2014], Meremonte et al., [2002], and Nakai et al., [2017]. The limitations and future work section (5.1) addresses the comparisons between these catalogs with the ANSS. Most earthquake epicenters are located within the crystalline basement in these catalogs with depths ranging from 4 – 6 km and uncertainties of up to 15 km [Nakai et al., 2017; Rubinstein et al., 2014]. These catalogs focus on certain earthquake sequences such as 2001 and 2011 and utilize temporary seismic arrays. Therefore, the relocations have less error in these events. In addition, a number of seismic models contribute to depth locations and associated errors [Rubinstein et al., 2014, Barnhart et al., 2014].

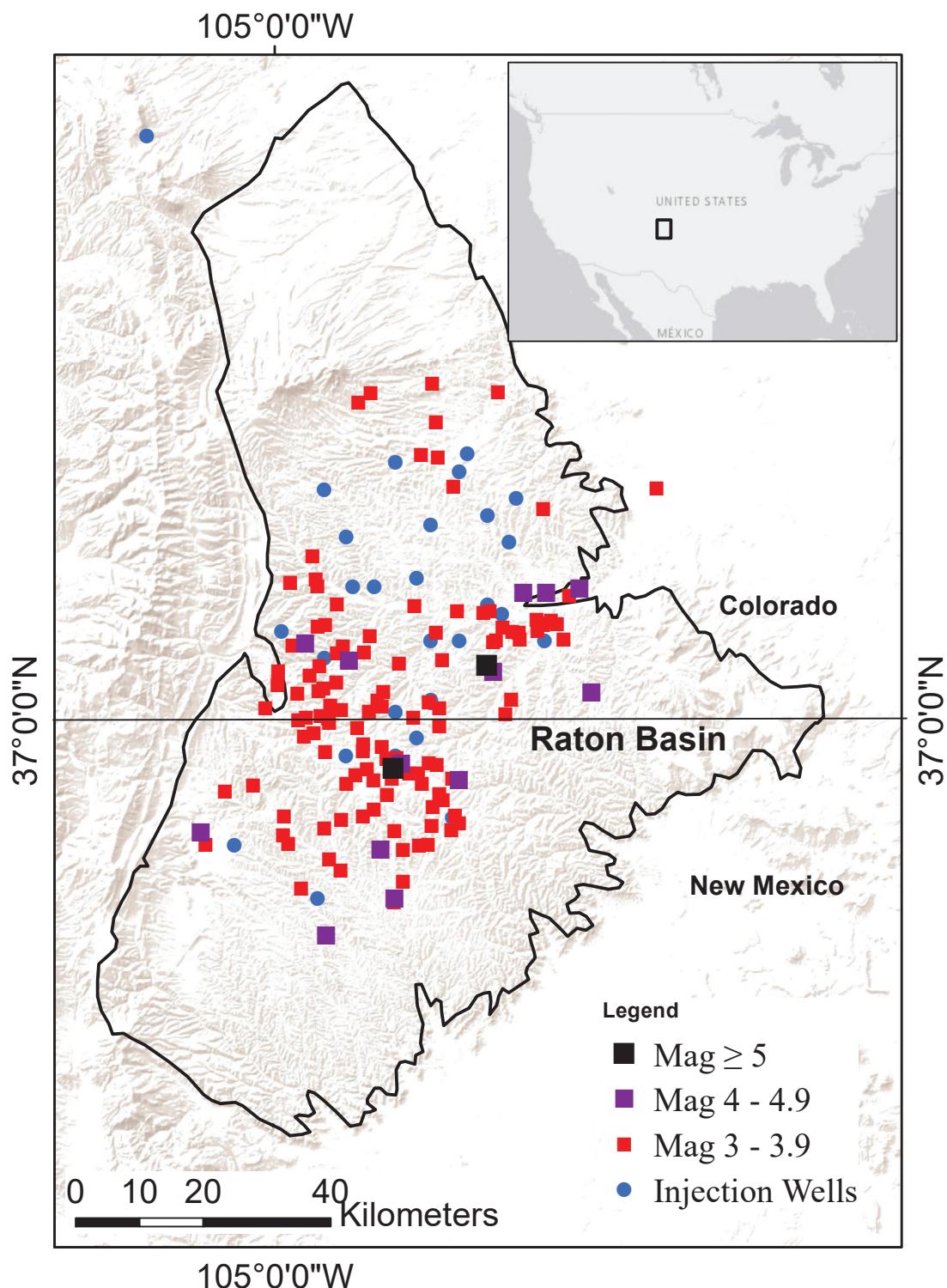


Figure 17. Wastewater disposal wells and earthquakes in Raton Basin. Earthquakes M 3 – 3.9 are shown by red squares. M 4 – 4.9 are purple squares and M ≥ 5 are black squares. Disposal wells are shown by blue squares.

4. GROUNDWATER FLOW and GEOTHERMAL MODELING

4.1 Objective

The scientific question addressed here is whether a spatial relationship exists between earthquakes, the injection wells, and the geothermal gradient. By using a coupled heat and fluid model, the objective is to determine if areas of high geothermal gradient, specifically the east central area, correspond to a change in earthquake activity (Figure 8). The hypothesis is that wells in high geothermal gradient areas have higher correlation with earthquake occurrence. Modeling was done to determine if any temperature contrasts due to cooler wastewater has the ability to reach the basement.

4.2 Numerical Model and Governing Equations

The Transport Of Unsaturated Groundwater and Heat (TOUGH2) code models pore pressure and temperature in three dimensions [Pruess *et al.*, 2012; Thunderhead Engineering, 2016]. TOUGH2 solves mass and energy balance equations while simulating non-isothermal multiphase flow in fractured and porous media. It was developed at the Lawrence Berkley National Laboratory and is written in FORTRAN77. TOUGH2 uses the integral finite difference method and simulates fluid and heat transfer in porous media. It uses a linear triangulation algorithm to interpolate between temperature data points and has been used in geothermal reservoir studies. PetraSim is the graphical interface that TOUGH2 is run in.

4.2.1 Mathematic Model

The mass and energy balance equation that is solved in TOUGH2 is

$$\frac{d}{dt} \int_{V_n} M^\kappa dV_n = \int_{\Gamma_n} F^\kappa \cdot n d\Gamma_n + \int_{V_n} q^\kappa dV_n \quad (1)$$

The integration is over an arbitrary subdomain V_n of the flow system. Γ_n is the boundary of the closed surface. M is the accumulation term and represents the mass or energy per volume. κ is

the heat or fluid component. F symbolizes the mass or heat flux. q are the sinks and sources. η is the normal vector on surface element $d\Gamma_n$ pointing inward into V_n [Pruess *et al.*, 2012].

The mass accumulation is expressed as

$$M^\kappa = \phi \sum_\beta S_\beta \rho_\beta X_\beta^\kappa \quad (2)$$

κ is the total mass component obtained by summing over the fluid phases (β = liquid, gas, and non – aqueous phase liquids). ϕ is porosity. S_β is the saturation of phase β or the fraction of pore volume occupied by phase β . ρ_β is the density of phase β . X_β^κ is the mass fraction of component κ present in phase β [Pruess *et al.*, 2012].

Heat accumulation term in a multiphase system is:

$$M^{N\kappa+1} = (1 - \phi)\rho_R C_R T + \phi \sum_\beta S_\beta \rho_\beta u_\beta \quad (3)$$

ρ_R grain density. N is the mass components (water, air, H₂, solutes). C_R is the specific heat of the rock. T is temperature. u_β is specific internal energy in phase β [Pruess *et al.*, 2012].

Advective mass flux is sum over phases. Individual phase fluxes are given by a multiphase version of Darcy's law:

$$F_\beta = \rho_\beta u_\beta = -k \frac{k_{r\beta} \rho_\beta}{\mu_\beta} (\nabla P_\beta - \rho_\beta g) \quad (4)$$

u_β is the Darcy velocity (volume flux) in the phase β . k is the absolute permeability. $k_{r\beta}$ is the relative permeability to phase β . μ_β is viscosity. P_β is the fluid pressure in phase β , which is the sum of pressure of a reference phase (usually taken to be the gas phase) and the capillary pressure. g is vector of gravitational acceleration.

Heat flux includes conductive and convective components

$$F^{N\kappa+1} = -\lambda \nabla T + \sum_\beta h_\beta F_\beta \quad (5)$$

λ is thermal conductivity. h_β is specific enthalpy in phase β .

4.3 Model Setup and Boundary Conditions

4.3.1 Domain and Discretization

I created a three dimensional coupled fluid and heat transport model to simulate pore pressure and temperature in Raton Basin from January 1999 to December 2016. The model domain extends 110 km north - south and 100 km east - west. Model depths range from 0.5 km (top of the confined aquifer) to 5 km (Figure 18). The Sangre de Cristo Mountains act as the western boundary; the Sierra Grande Uplift and Apishapa and Cimarron Arches mark the eastern, north, and south boundaries of Raton Basin (Figure 1). A spatially varying polygonal grid with refinement around injection wells in the horizontal direction allowed for finer grid dimensions in the injection interval in the vertical direction.

The model top is marked by the no flow boundary of the Pierre Shale, which has an average depth of 0.5 km bgs. The model extends into the crystalline basement to a total depth of 5 km. The average crystalline basement depth starts at 3.3 km bgs. The modeled Pierre Shale, Niobrara, and Benton Formations act as a confining layer to upward flow (Figures 6,18). Six hydrologic units were modeled with varying thicknesses to represent different lithologies in Raton Basin (Figure 18, 19).

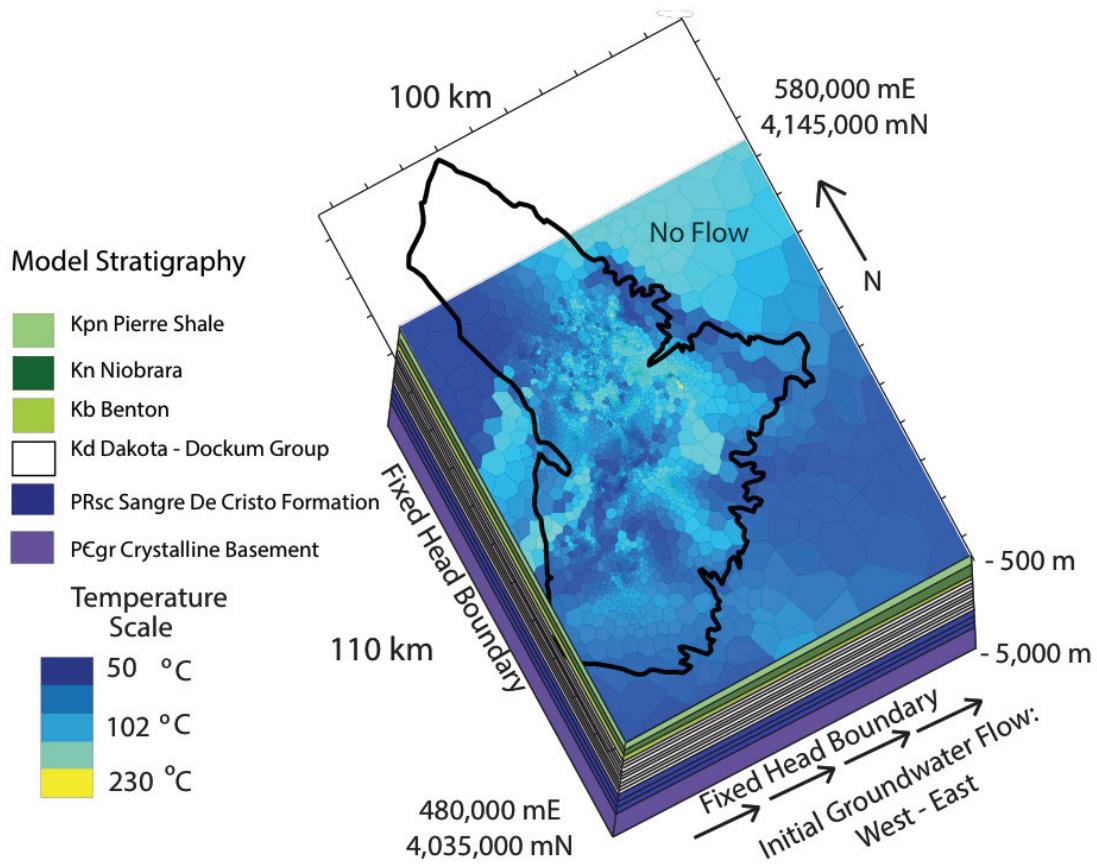


Figure 18. Domain and Boundary Conditions for the model. Plan view shows the model grid with Raton Basin's outline with the black line. The grid is refined around the disposal wells. The temperature for the top of the Dakota Formation is shown and interpolated from the Bottom Hole Temperature data points. Cross section view shows the layers of the model. UTM (mE and mN) was used as the coordinate system in the model.

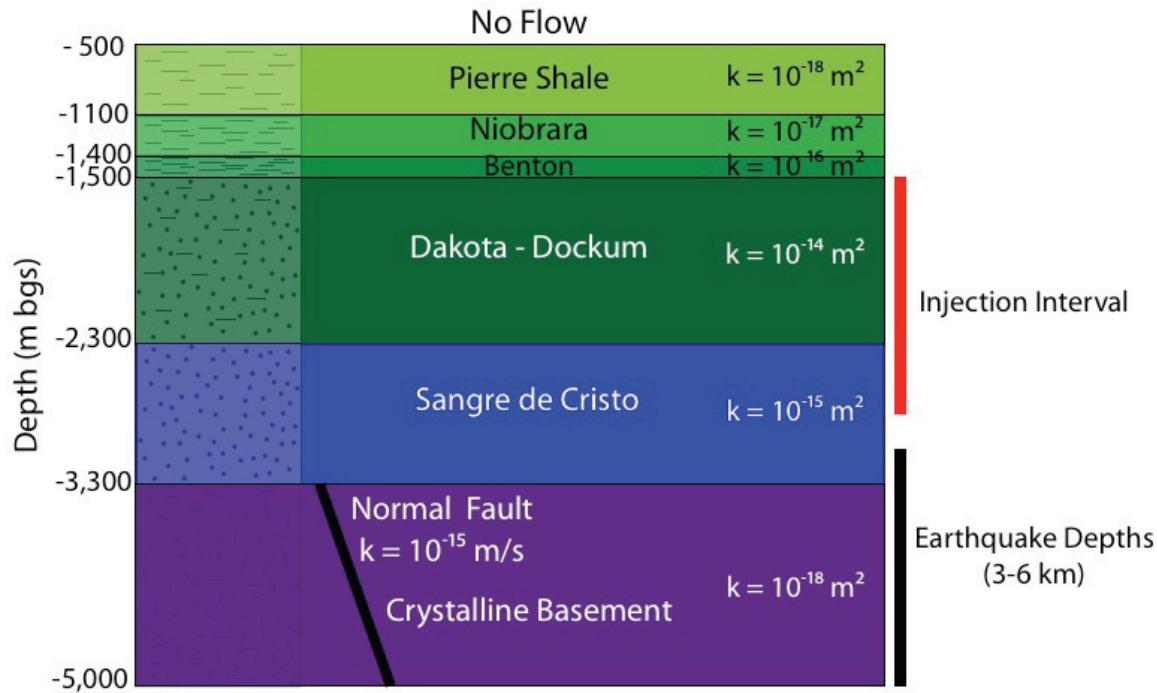


Figure 19. The six hydrogeologic units in the TOUGH2 model. Lithologies are shown and vary from shale, sandstone, and gneiss/granite in the crystalline basement. Permeability values are shown for each layer [Geldon, 1989; Bohlen, 2013, Weingarten, 2015]

4.3.2 Boundary and Initial Conditions

Fixed head boundaries are used on all boundaries of the domain. Drillstem test data from Nelson et al. [2013] provided the hydraulic head values in the Dakota Formation. These were used to determine the initial conditions for pressure in the Basin and it reflects overall groundwater flow from west to east (Figure 20 and 21). Pressure values equivalent to approximately 1,772 m of head (17.37 MPa) are shown on the western side of the Basin at a depth of 1,500 m while 1,547 m of head (15.16 MPa) is found on the eastern boundary of the Basin (Figure 20). The top is a zero flux boundary because the permeability values used in the top confining layers (Pierre Shale – Benton) are two to four orders of magnitude less than the injection interval (Figure 19). The boundaries of the model are far enough from injection wells that there will be little boundary effect on the model results.

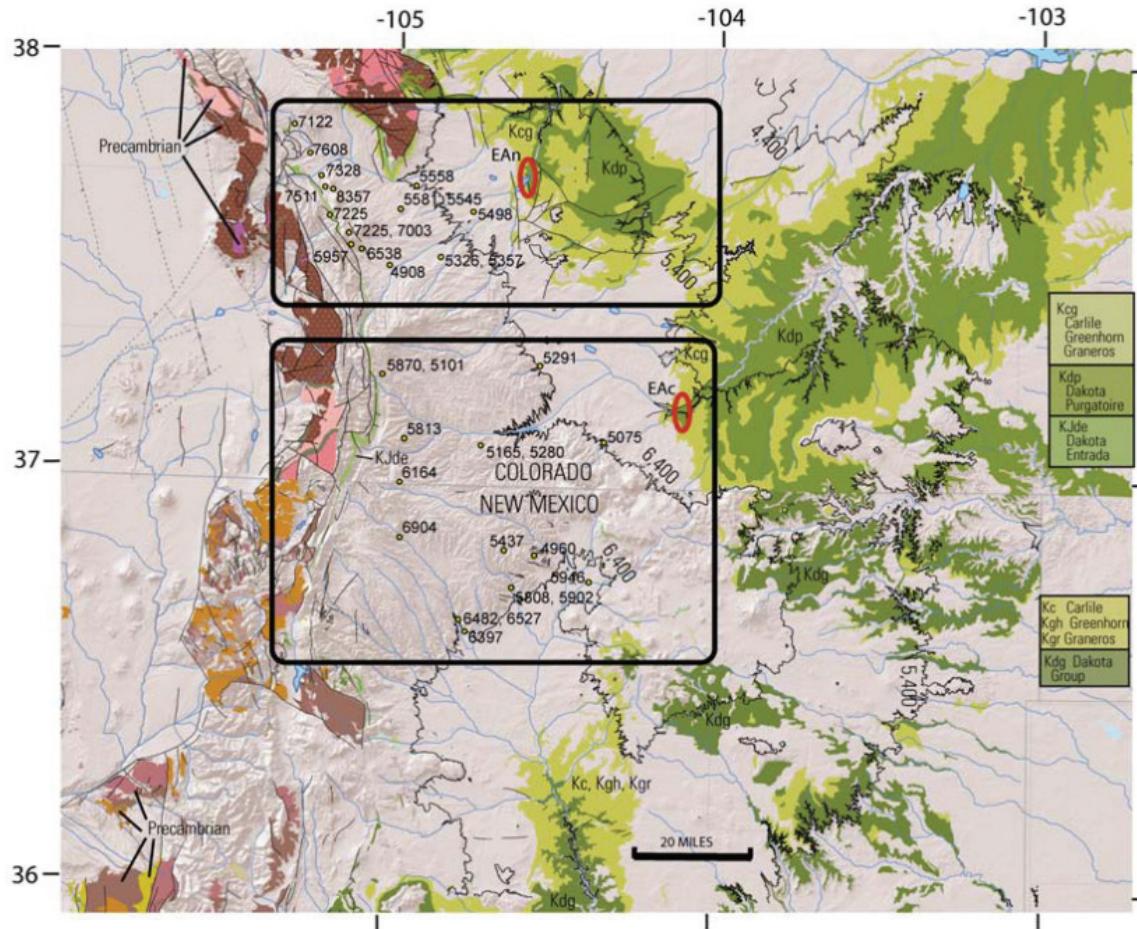


Figure 20. Hydraulic head in ft at the Dakota Formation [Nelson *et al.*, 2013]

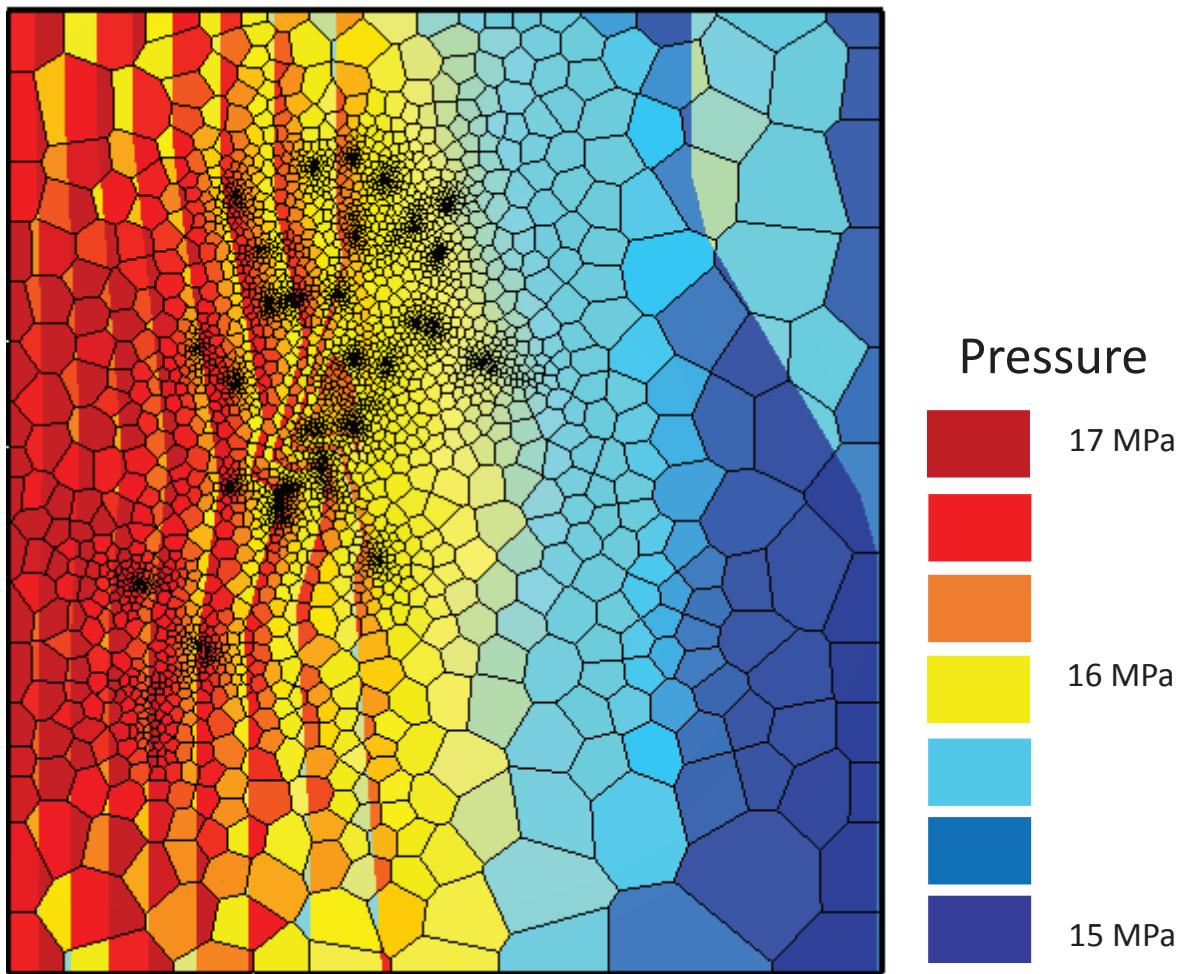


Figure 21. Initial pressure at the top of the Dakota Formation. Groundwater flows from high pressure on the west side to low on the east.

The initial temperature regime was calculated from the interpolation of BHT data (Figure 22). 1,627 points were input representing the temperature at the top of each hydrologic unit. The interpolation used in the model is a linear triangulation algorithm.

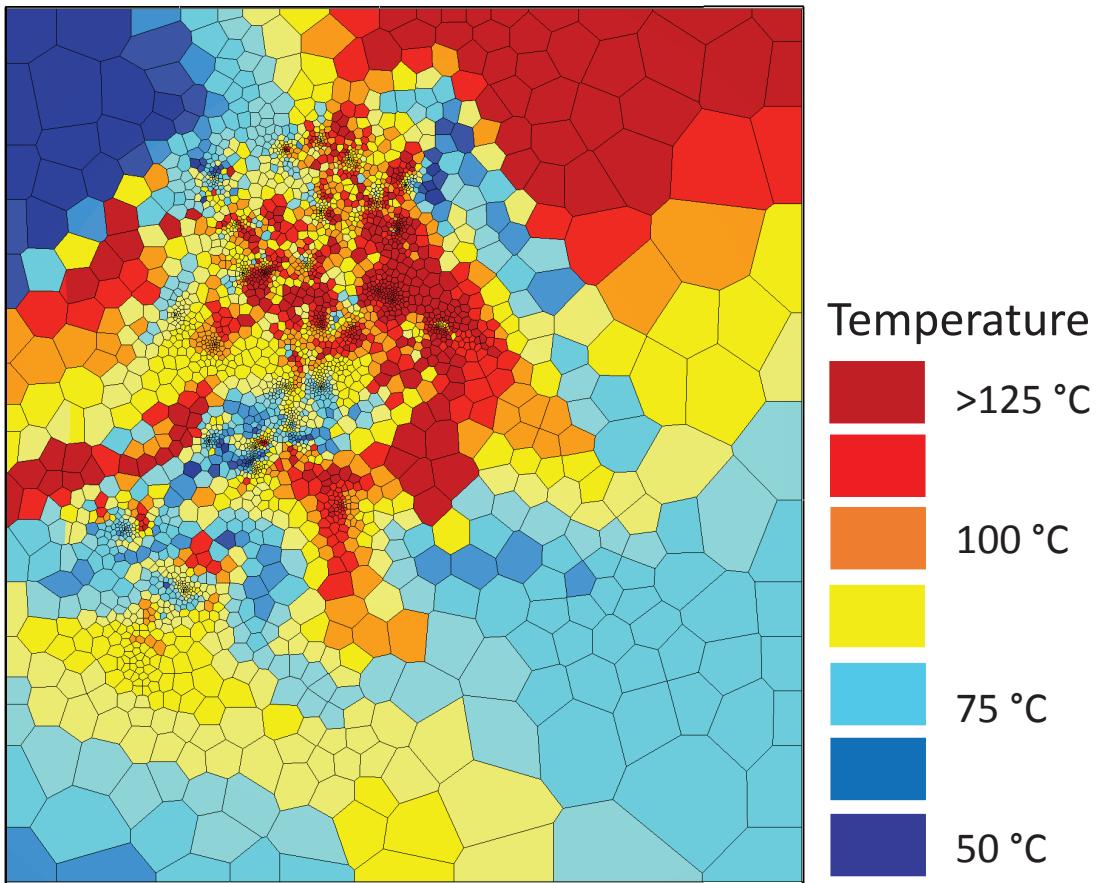


Figure 22. Initial Temperature at the top of the Dakota Formation.

4.3.3 Faults

There are a number of faults in the crystalline basement, with some cutting into the permeable sedimentary layers that lie above it [Meremonte *et al.*, 2002, Barnhart *et al.*, 2014, Nakai *et al.*, 2017] These crystalline basement faults are critical for earthquake occurrence because of thermal contraction due to temperature differences between the cooler wastewater and hot subsurface moving through the faults. This could allow a pathway for stress differences due to thermo-elasticity effects. Faults were added into the model to facilitate a conduit for temperature and pore pressure changes to reach the crystalline basement, which is the depth at which a majority of the earthquakes occur. Faults are important in facilitating seismicity. Fluid

pressure increase transmitted to a fault lowers the stress holding the fault closed (normal stress) and then results in a lower effective stress. Faults are more prone to failure or slip because the lower effective normal stress lowers the frictional resistance to slip [Rubinstein *et al.*, 2015]. Because PetraSim does not allow faults to be put into the model, two parallel vertical planes (i.e. layers) created an internal boundary that cut through the crystalline basement layer of the model (Appendix – Figure 42). Parameters for the rock between the two areas, mainly permeability, were changed to simulate a fault area. The permeability within the faults was set to 10^{-14} m^2 . Sensitivity analysis was also performed with multiple permeability values ranging from 10^{-16} m^2 to 10^{-14} m^2 . Previous studies have shown that this fault can provide a conduit to flow, which allows for pore pressure to increase within the range to induce seismicity [Barnhart *et al.*, 2014, Nakai *et al.*, 2017]. These modeled normal faults begin at 3 km bgs, or slightly into the Sangre de Cristo sandstone, and extends through the crystalline basement to the base of the model at 5 km.

An internal boundary was placed where known faults and earthquakes have occurred [Barnhart *et al.*, 2014; Nakai *et al.*, 2017]. This fault represents a normal fault starting at a depth of 3.3 km, or the top of the crystalline basement. The fault strikes 35 NE and dips 65 SE through the central part of Raton Basin (Figure 23, Barnhart *et al.*, 2014). The Wild Boar well, which is located 1 km away from the fault location is in a high geothermal gradient area (Figure 23).

A second fault zone, the Tercio, was modeled on the western edge of the Basin [Nakai *et al.*, 2017; Figure 23]. The Tercio fault has 129 earthquakes associated with it. Lineation of seismicity and mapped surface faults 10 – 15 km of Tercio have been used as evidence for the presence of this fault. The Mw 5.3 occurred 4 km east of the Tercio fault with a moment tensor indicating an east- west normal fault. In addition, a moment tensor for a Mw 3.53 earthquake that

occurred on September 29, 2009 at the northeast terminus of the Tercio fault indicates a northeast-southwest normal fault [Nakai et al., 2017].

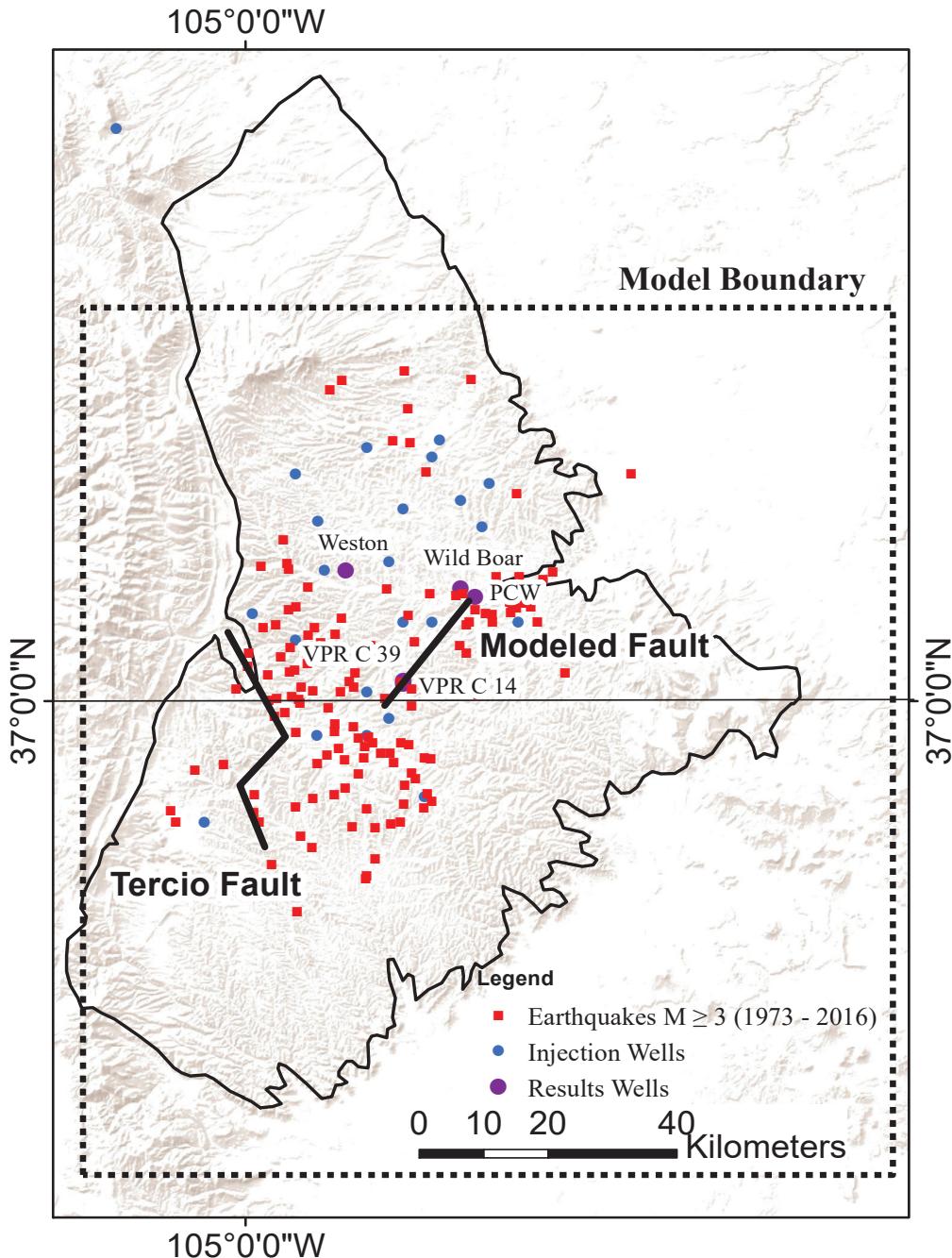


Figure 23. Fault locations that were added into the model. The Tercio fault has been named while the earthquake labeled as “modeled fault” does not have a specific name in literature. Red squares represent earthquakes $M \geq 3$ while injection wells are in blue. The five wells that will be analyzed in the results section are shown as purple dots and labeled with their names.

4.4 Model Sources and Input Parameters

Input data sources for the model include the 28 injection wells that were active from January 1999 to December 2016. Model inputs include well locations and monthly injection rates. Hydrologic and geologic parameters were collected from literature. The average temperature of wastewater injection at the time of injection is 20°C (Figure 16).

4.5 Model Results

Pore pressure and temperature change from injection were modeled using TOUGH2 [Pruess *et al.*, 2012] (Figures 24 – 35). Modeling shows that pore pressure increased up to 0.2 MPa, at a depth of 3.3 km bgs (Figure 35). These model results support previous studies that point to pore pressure increase as the main cause for induced seismicity. These studies have shown pore pressure increase between 0.5 and 1 MPa in the basement, where the earthquake nucleation has occurred [Weingarten, 2015; Nakai *et al.*, 2017; Rubinstein *et al.*, 2014]. In addition, pore pressure increases up to 0.8 MPa within 100 m of the injection well at the injection well depth (Figure 26). This is showing that pore pressure is highest at the injection depth and diffuses towards the basement. However, temperature does not seem to show any change at crystalline basement depths and therefore, there is no temperature contrast created from the wastewater injection and the rock in the subsurface at the depths at which earthquakes are occurring. Five of the twenty-eight injection wells modeled had a temperature decrease 100 m away in the injection interval depth. However, results show that the temperature remains unchanged directly below all injection wells at the top of the crystalline basement (3.3 km bgs).

Faults were included in the model as a conduit to see if the temperature difference between the wastewater injection and rock would be able to reach the basement depth (Figure 23). With the addition of fault – induced permeability, pore pressures increased approximately

0.1 MPa at the crystalline basement, but temperature remained constant in all scenarios (Figures 30, 32 - 35). Weston Injection well had a pore pressure increase of 0.075 MPa at the crystalline basement (3.3 km bgs) with a constant temperature of approximately 160°C (Figure 30).

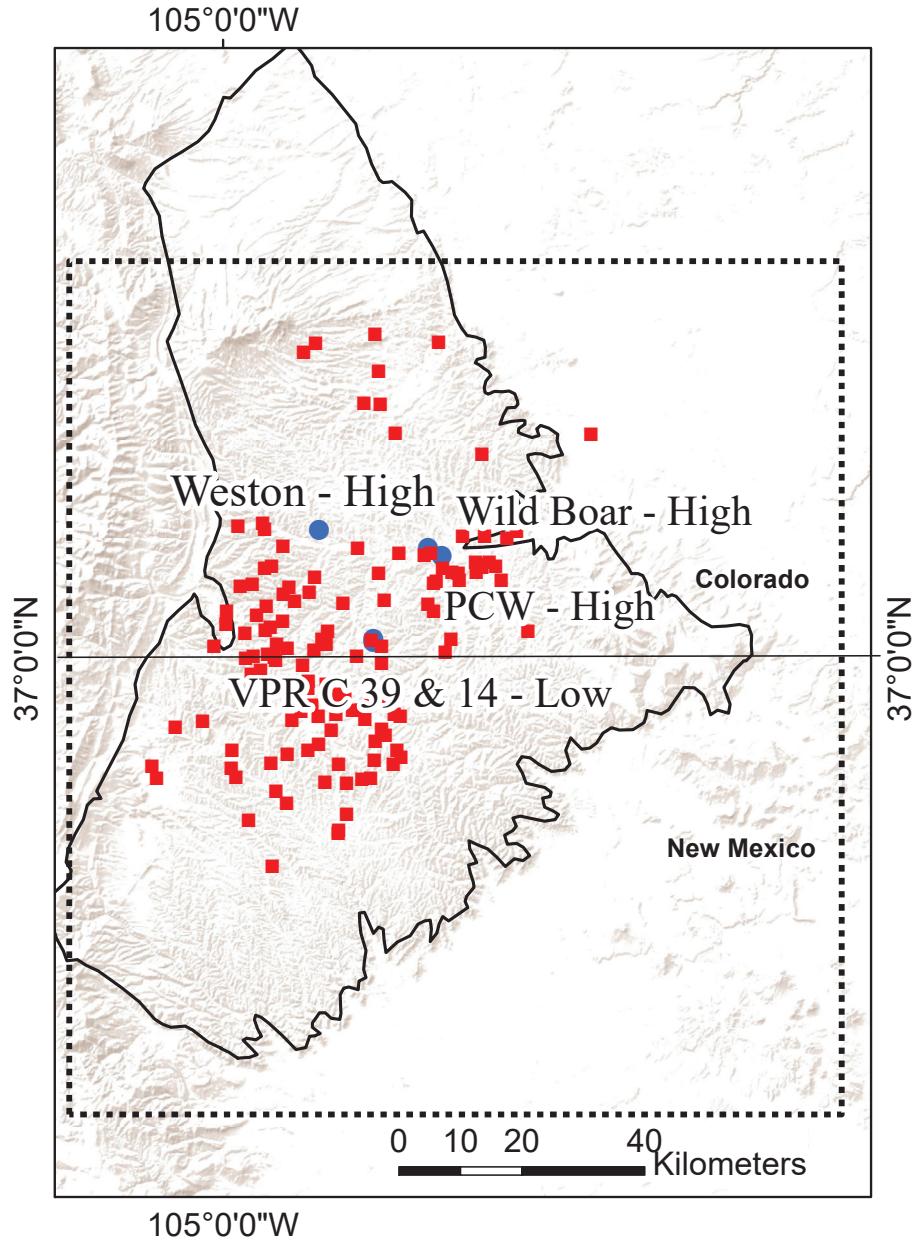


Figure 24. Location of wells analyzed for results. It is annotated if they are in low or high geothermal gradient areas.

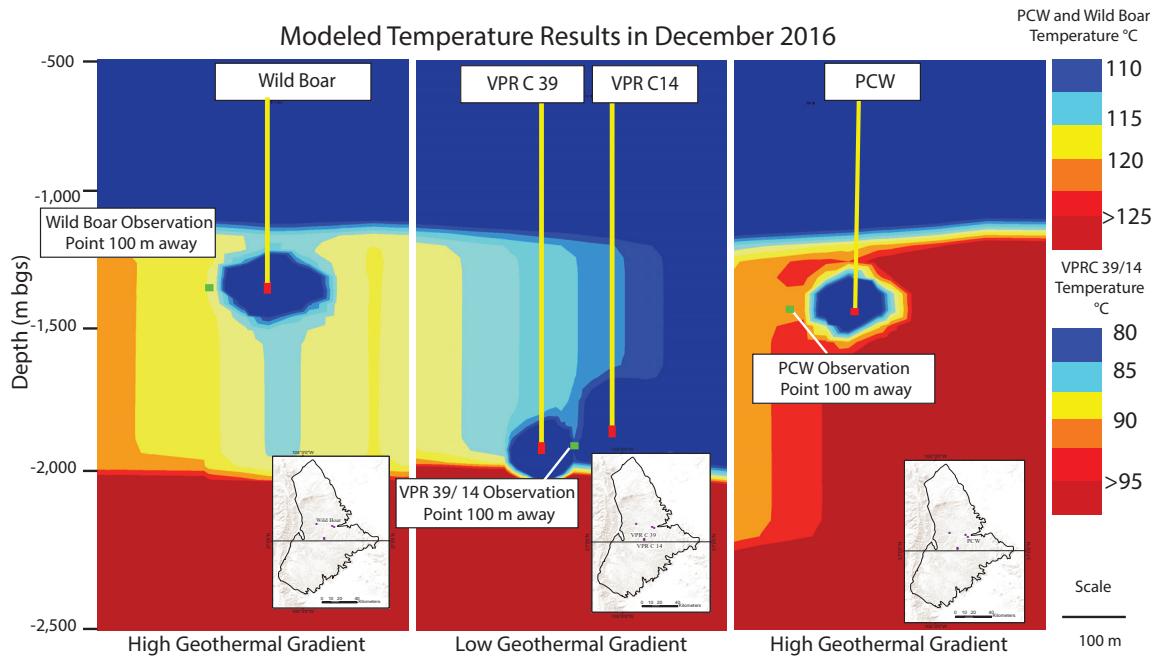


Figure 25. Modeled Temperature surrounding four injection wells (Wild Boar, VPR 39/14, and PCW).

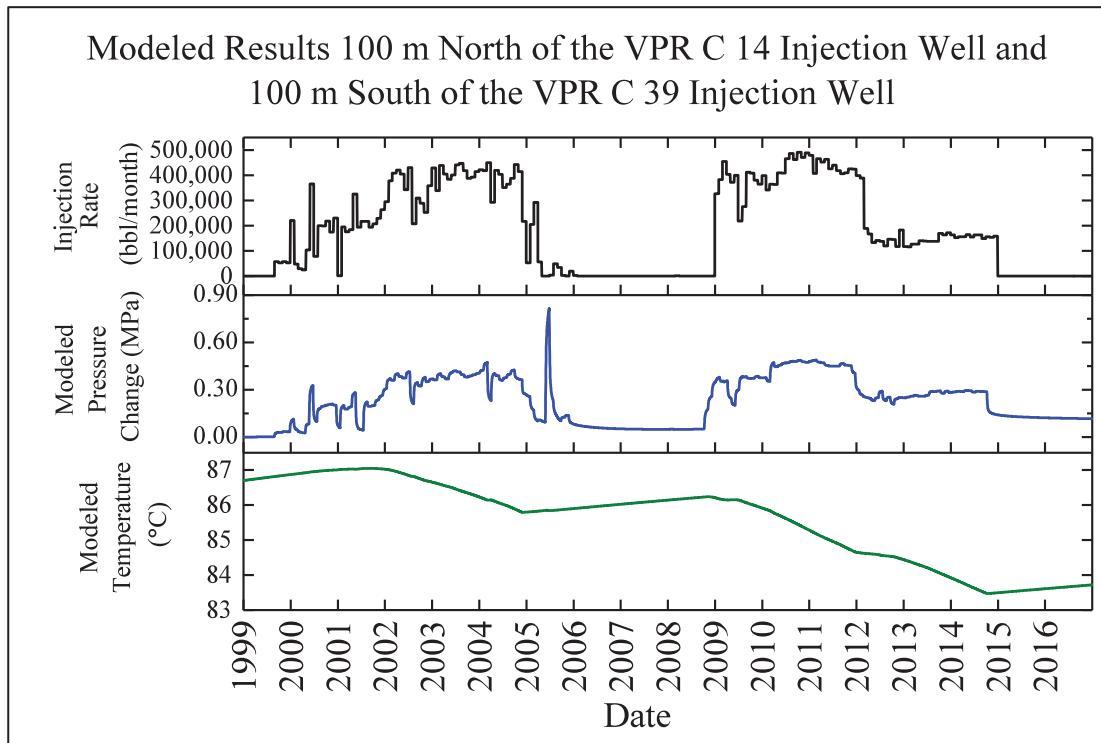


Figure 26. Modeled pressure and temperature for VPR C 14 and VPR C 39 injection wells at an observation point 100 m away from both wells in Raton Basin from 1999 – 2016. As shown in Figures 8 and 24, they are in a low geothermal gradient area. There is a break in injection due to the status of the wells being either closed or shut in for a few years.

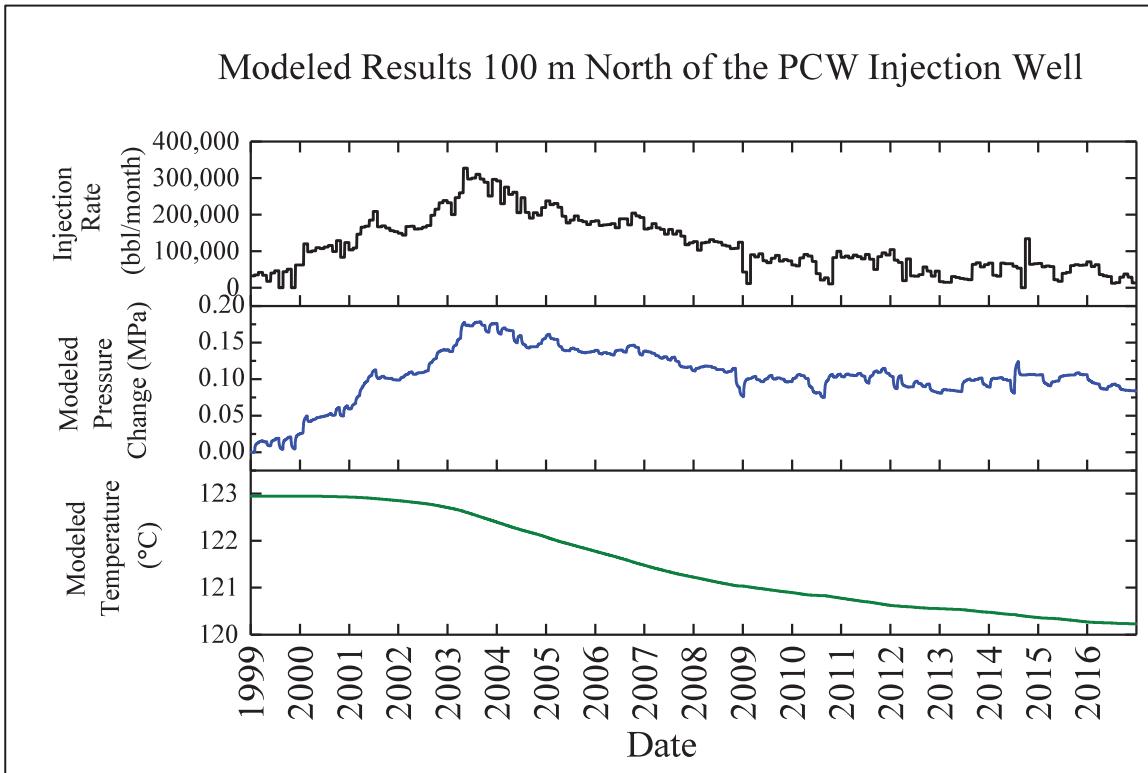


Figure 27. Modeled pressure and temperature at an observation point 100 m north of the PCW injection well in Raton Basin from 1999 – 2016. As shown in Figures 8 and 24, PCW is in a high geothermal gradient area.

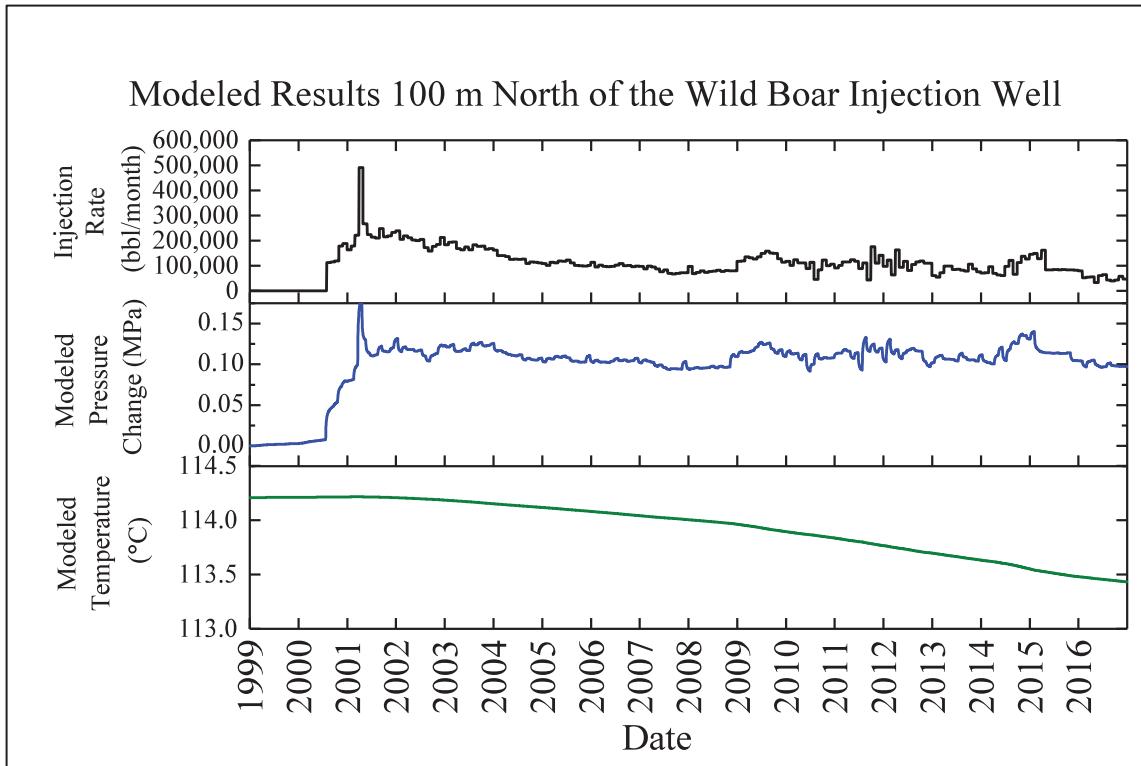


Figure 28. Modeled pressure and temperature 100 m north of the Wild Boar injection well in Raton Basin from 1999 – 2016. As shown in Figures 8 and 24, Wild Boar is in a high geothermal gradient area.

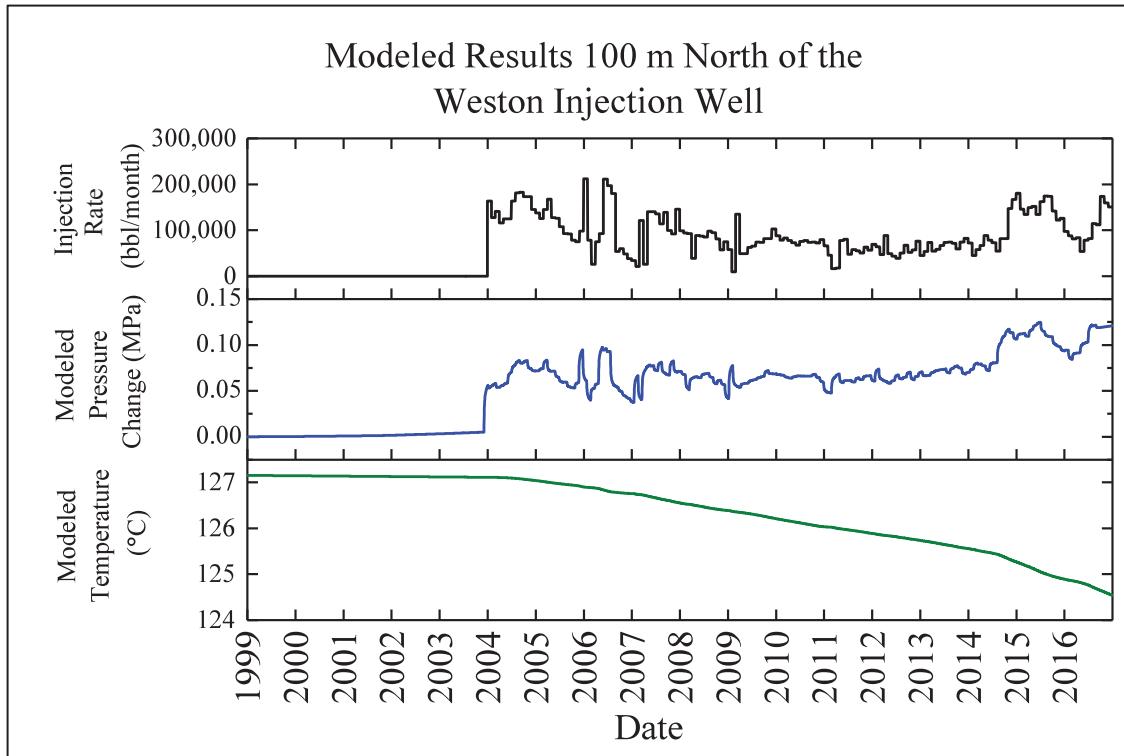


Figure 29. Modeled pressure and temperature 100 m north of the Weston injection well in Raton Basin from 1999 – 2016. As shown in Figure 8 and 24, Weston is in a high geothermal gradient area.

4.5.1 High Geothermal Gradient Wells

Disposal at the Weston well began in January 2004, and continues as of December 2016.

The average disposal rate is 93,723 bbls/month (~3,000 bbl/day) and a maximum injection rate of 203,122 bbl/month (~7,000 bbl/day). The depth of the disposal is from 1,879 and 2,077 m bgs, primarily into the Purgatoire Formation (Appendix). Modeling indicated pore pressure increased up to 0.1 MPa at 100 m from the injection interval depth and 0.075 MPa at a depth of 3.3 km bgs directly below the well at the top of crystalline basement (Figures 29 and 30). Temperature decreased by approximately 3°C in the injection interval, but did not change at the basement depth (Figure 30).

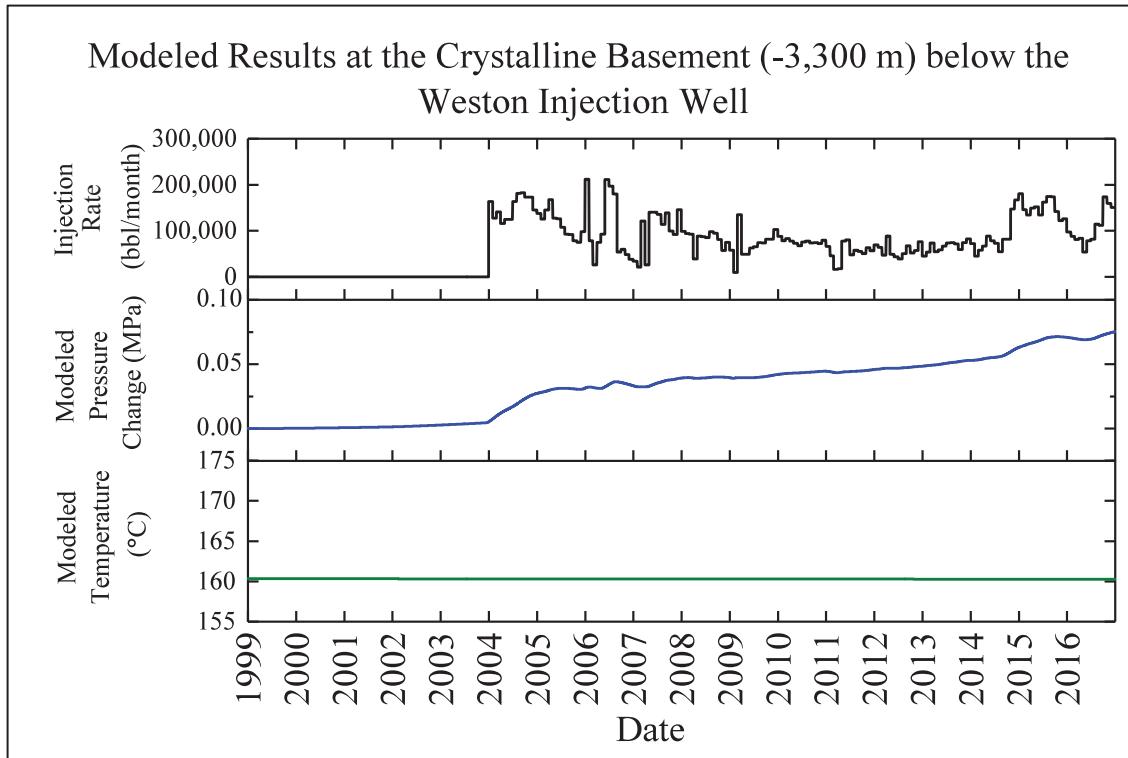


Figure 30. Modeled results at the crystalline basement below the Weston Injection well.

The PCW injection well is in the area of higher geothermal gradient within Raton Basin. PCW has been actively disposing since 1999. Its average injection rate is slightly higher than the Weston Well with a rate of 115,832 bbls/month (~3,900 bbl/day). Its injection interval is from a depth of 1,208 and 1,225 m bgs, into the Dakota Formation. Per the modeling results, pore pressure increased up to 0.15 MPa during its peak injection period in the injection interval (Figure 27). Temperature decreased up to 2°C at the injection interval depth and did not change at the basement depth.

A “worst case” scenario was run with this well injecting wastewater at 500,000 bbl/month at 11°C, instead of 20°C. The minimum temperature that wastewater is injected as recorded at disposal tanks is 11°C [Macartney, personal comm.]. This resulted in a modeled pressure change of 0.3 MPa and a temperature decrease of 20°C, 100 m from the injection well within the injection interval (Figure 31). However, the temperature remained constant below the

PCW injection well at the crystalline basement depth (Figure 32).

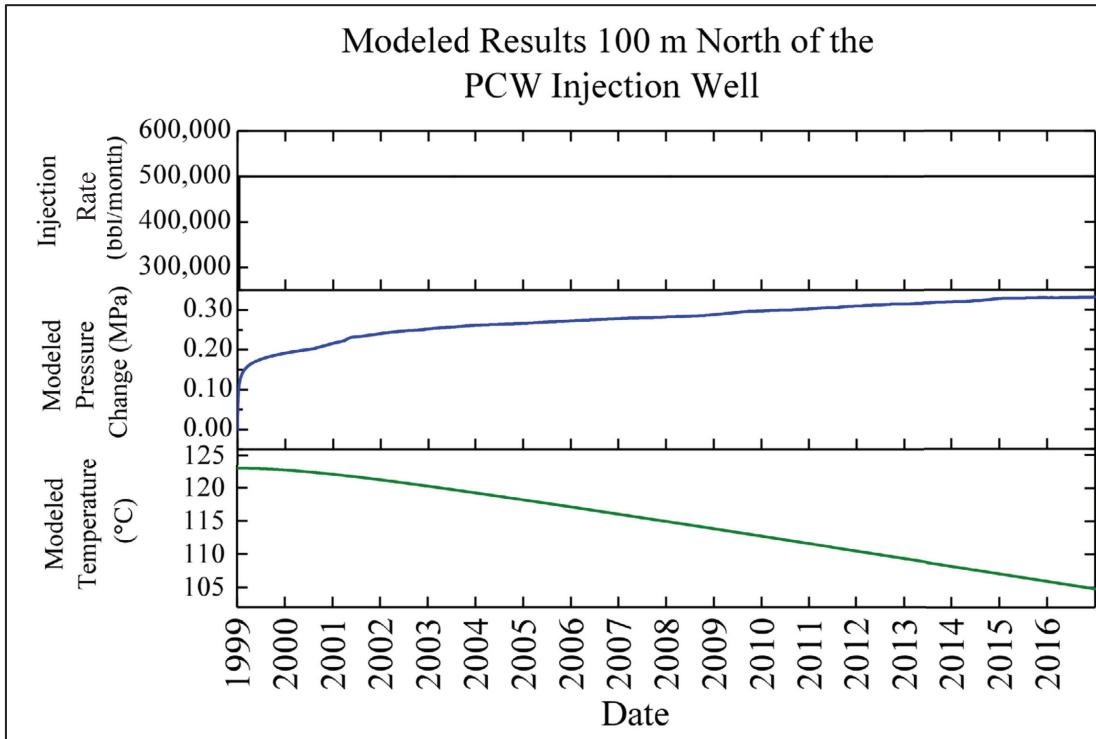


Figure 31. “Worst Case Scenario” for the PCW Injection Well. Injection was set to 500,000 bbl/month with an injection temperature of 11°C.

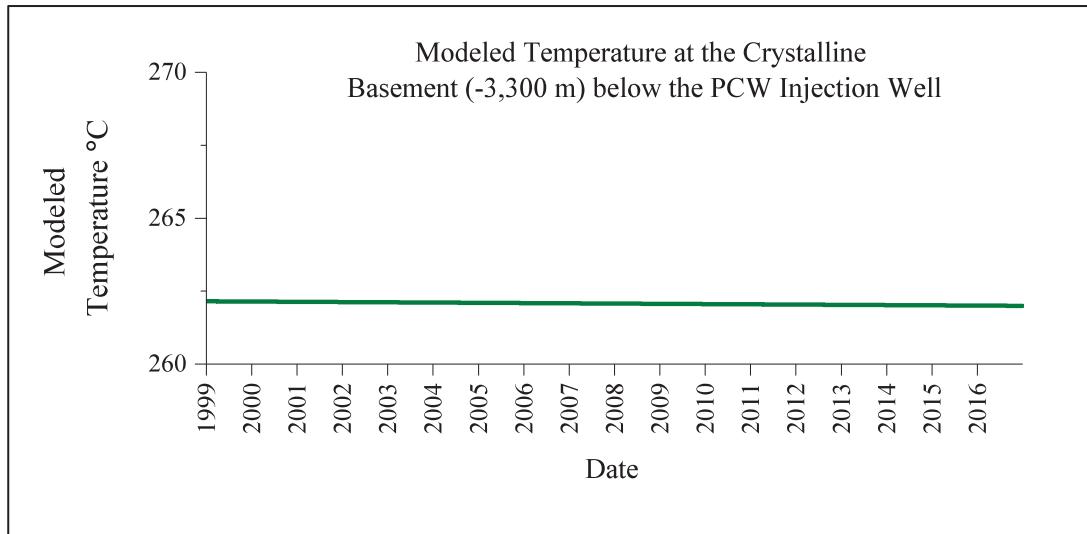


Figure 32. Modeled temperature at the basement for the PCW Injection Well.

The Wild Boar well, which began disposing in August 2000 has an average injection rate of 118,928 bbls/month (~3,900 bbl/day), with a maximum of 491,058 bbls/month (16,368

bbl/day). The cumulative volume is 23,428,799 barrels. Its injection interval ranges from 1,257 and 1,292 m bgs into the Dakota Formation. Modeling results indicate pore pressure increased by 0.15 MPa in the injection interval depth while temperature decreased by less than 1°C (Figure 16). Pore pressure increased approximately to 0.075 MPa in 17 years and temperature remained constant in the crystalline basement (Figure 33). The Wild Boar well is located 1 km from the normal fault that was added to the model (Figure 23). Modeling results showed similar ranges of pressure change below the injection well in the crystalline basement relative to the results from model simulations without the fault. Temperature did not change (Figure 34).

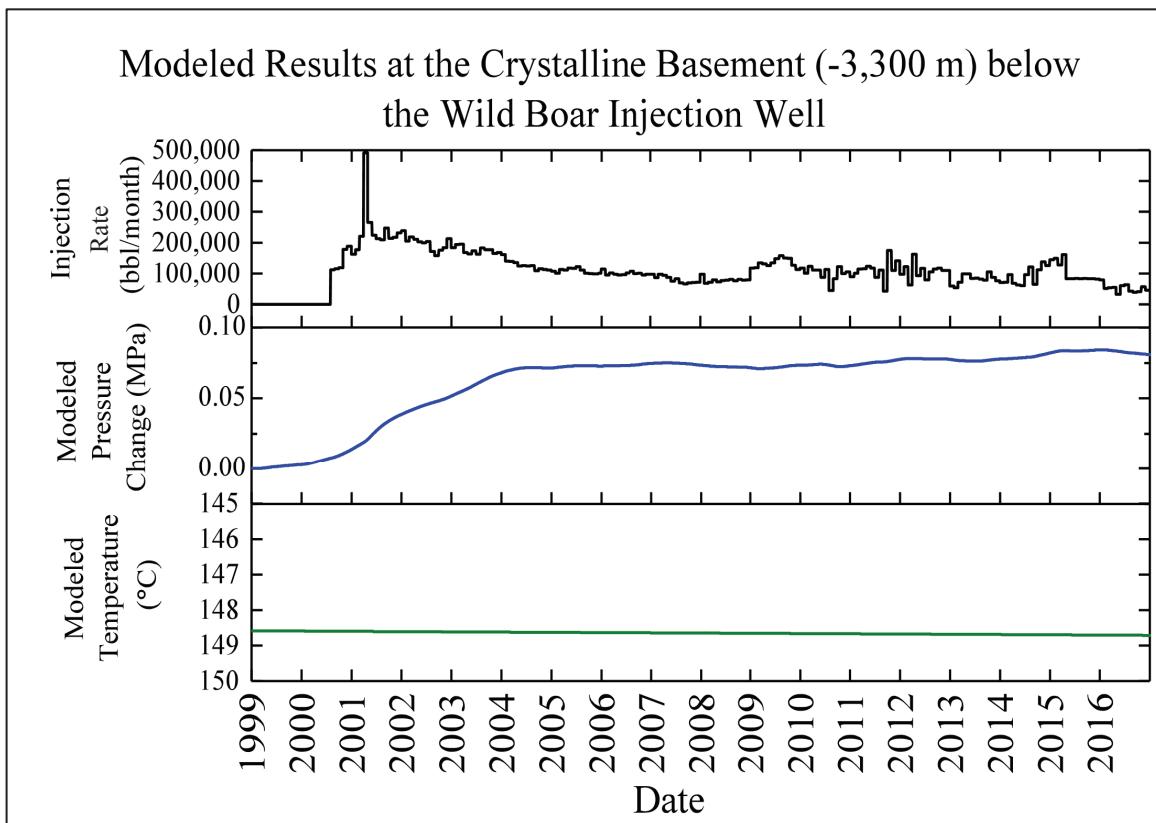


Figure 33. Modeled Results at the crystalline basement for the Wild Boar Injection Well.

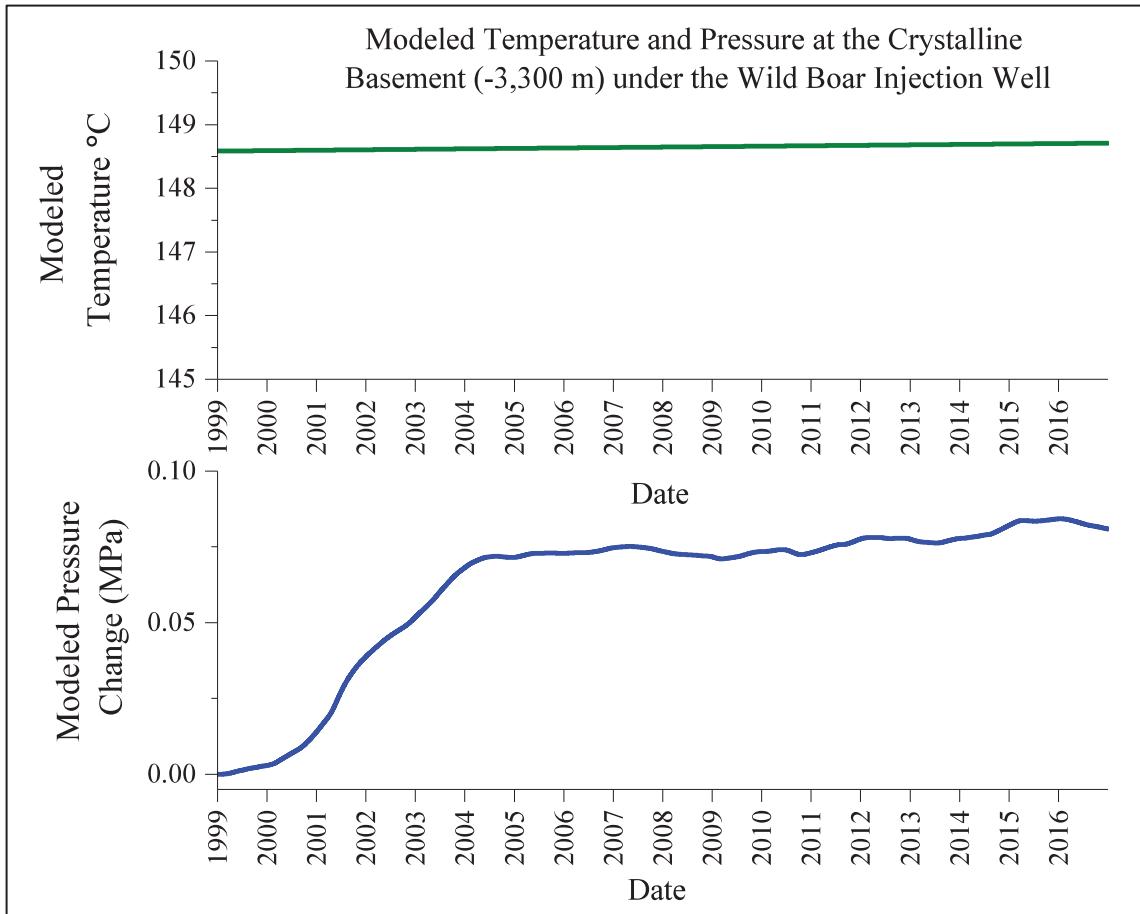


Figure 34. Modeled pressure and temperature at the crystalline basement under the Wild Boar Injection Well after the normal fault was added. See Figure 23 for the location of Wild Boar in relation to the fault. It is 1 km away from the fault.

4.5.2 Low Geothermal Gradient Wells

The VPR C 39 and VPR C 14 disposal wells are both located close to the Colorado-New Mexico border and are wells in a low geothermal gradient area (Figure 24). The injection intervals are at depths of 1,973 and 2,007 m (VPR C 39) and 1,682 and 2,020 m (VPR C 14) in the Entrada and Glorieta formations. VPR 39 started injecting in May 2000 and was shut-in in 2015. Its average injection rate is 115,757 bbls/month with a maximum rate of 356,826 bbls/month. VPR C 14 was an active injection well from September 1999 to 2002 and was closed in 2002. The average injection rate was 84,416 bbls/month and had a maximum injection

of 378,554 bbls/month. Cumulative injection for these wells was 32,432,569 barrels. Modeling results indicate that pore pressure increased up to 0.6 MPa with a temperature decrease of 3°C (Figure 14). Pore pressure at the top of the crystalline basement (-3,300 m bgs) increased approximately 0.2 MPa and did not show any temperature difference between steady state temperature when cooler wastewater was injected (Figure 35).

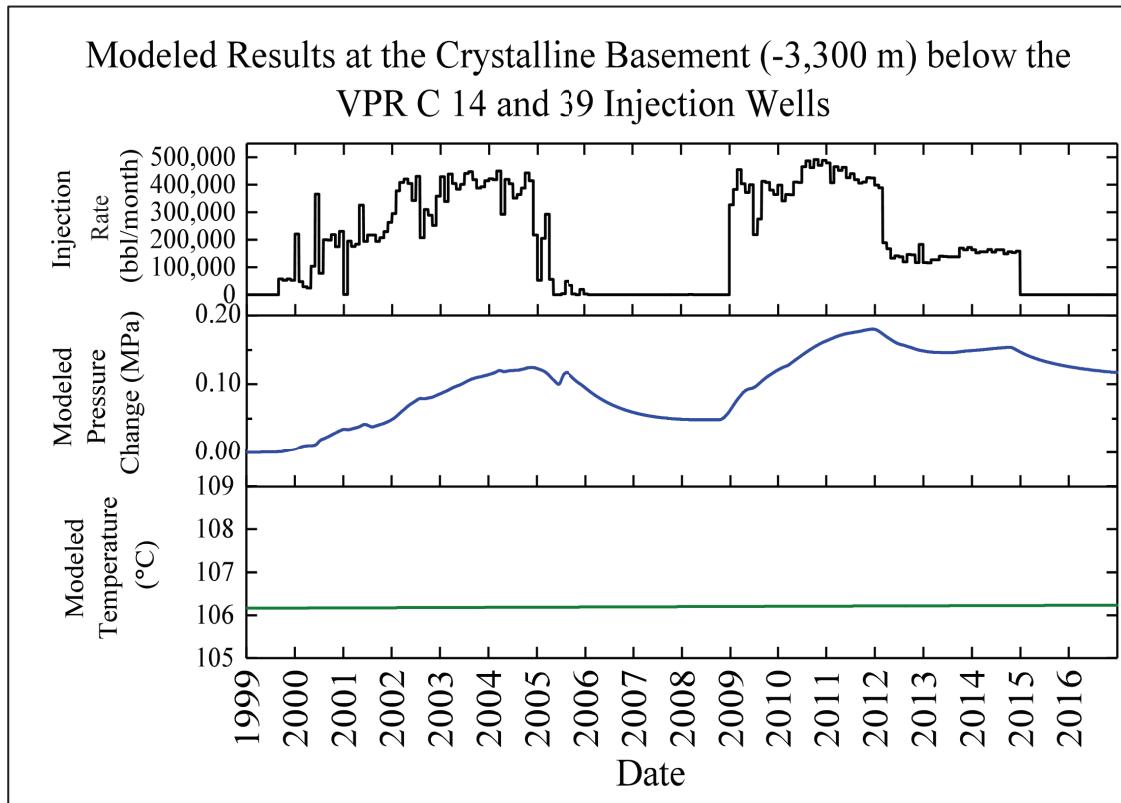


Figure 35. Modeled results at the basement below the VPR C 14 and 39 Injection Wells.

4.6 Sensitivity Analysis

Varying wastewater temperatures were used to determine if cooler wastewater would reach the basement. These temperature values were chosen from the varying temperatures that Pioneer Natural Resources had sampled over a 14 - year period. Since the wastewater is stored at above ground storage tanks, they are subject to the ambient air temperature. Therefore, the temperature of the water that is injected varies seasonally. Sensitivity tests were run using an

injection temperature of 11°C and 28 °C before converging on an overall average of 20 °C (Figure 36). The injection temperature did not show any significant change to the overall temperature distribution at the injection well or the surrounding area [Figure 42-44 – Appendix].

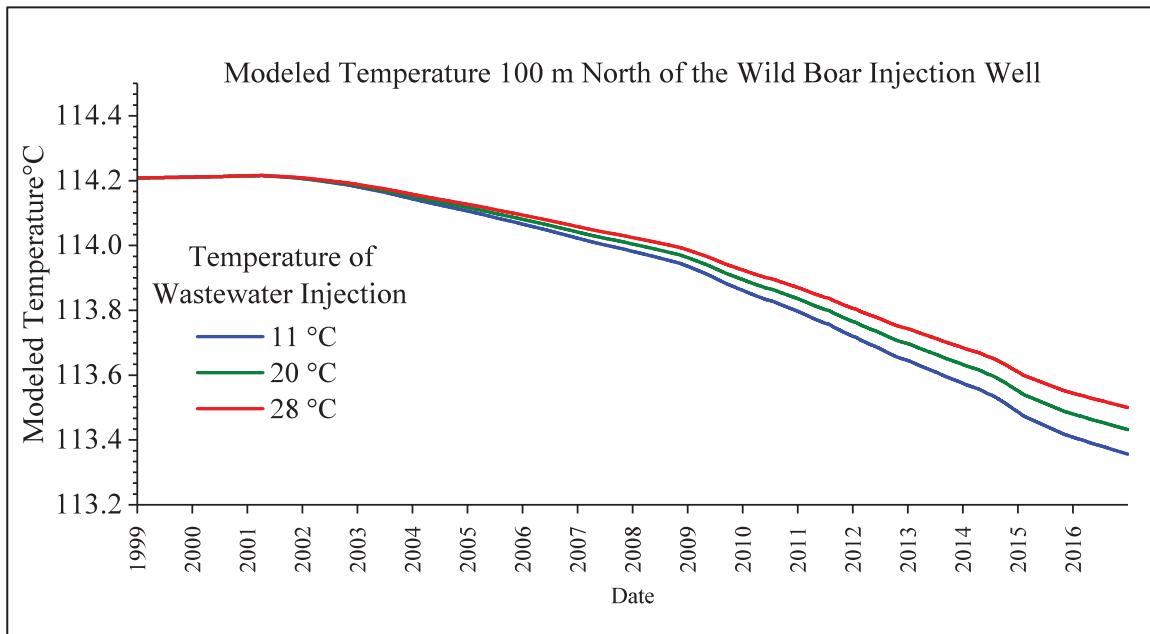


Figure 36. Sensitivity for Wastewater Injection Temperature on the Wild Boar Injection Well.

In addition to injection temperature, the permeability was varied for the Dakota Formation, as well as the Sangre de Cristo Formation. This was to see if water could migrate farther by increasing permeability in the formations. Similarly to the injection temperature, the change in permeability of the formations did not provide any significant difference for the temperature contrast and its inability to reach crystalline basement. The permeability of the fault was also changed to see if that would facilitate temperature changes or higher pore pressure diffusion, as shown by previous studies. However, modeling showed that there was not a notable difference (Figure 37).

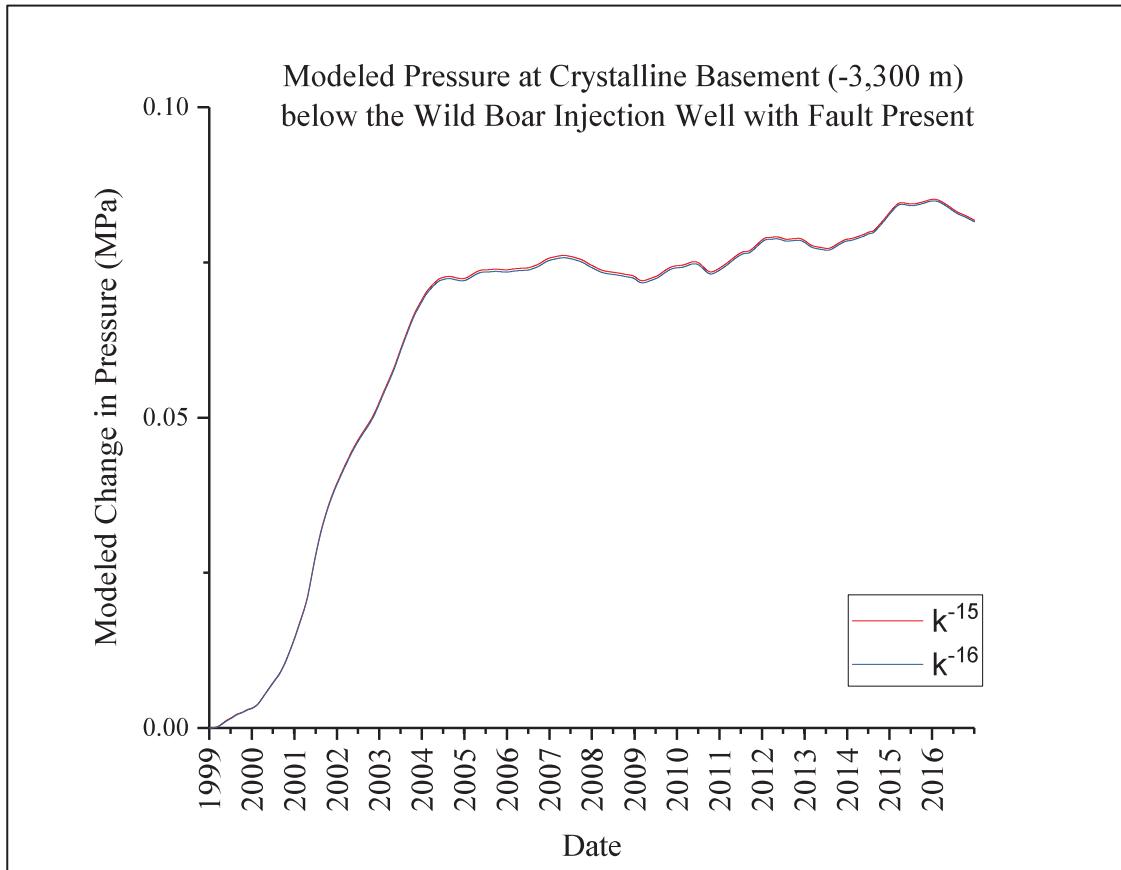


Figure 37. Modeled pressure at the crystalline basement below the Wild Boar Injection Well. Hydraulic conductivity by adding a fault was changed to see if it caused pore pressure change in the basement.

5. DISCUSSION

Diffusivity is the measurement of the area of influence (m^2) per unit of time (s). By comparing hydraulic diffusivity and thermal diffusivity in rock over the 17-year time span that the model was run, it becomes evident why the pore pressure was able to reach the depth of the crystalline basement, but any temperature change was not. Hydraulic Diffusivity (D_h) is calculated from hydraulic conductivity (K) over specific storage (S_s) ($D_h = K/S_s$). K is 10^{-7} m/s in the Dakota Sandstone, while specific storage is 10^{-6} m^{-1} , which gives a D of $10^{-1} (\text{m}^2/\text{s})$. Diffusivity (D) is $\sim L^2 / T$ (m) / T (s) where L^2 is the area of influence and T is the time that the model ran (10^8 s or 17 years). Therefore $L^2 = 10^7 \text{ m}^2$ which translates to an affected area for

hydraulic diffusivity on the scale of a few kilometers. Therefore, pore pressure has the ability to reach the crystalline basement, or depths that earthquakes are occurring.

Thermal diffusivity is calculated in a similar way. Thermal diffusivity of sandstones (D_T) is approximately $10^{-6} \text{ m}^2/\text{s}$ [Robertson, 1988]. $D_T \sim L^2 (\text{m}) / T (\text{s})$ where T is 10^8 s (17 years that the model ran). This results in an area of influence L^2 of only tens of meters. This accounts for the inability for any temperature difference to travel the $\sim 1.5 \text{ km}$ to the crystalline basement from where the cooler water is being injected.

Thermal stress is calculated as $\sigma = E \alpha (\Delta T)$ where σ is the thermal stress (Pa), E is Young's modulus (Pa), α is coefficient of thermal expansion ($1/\text{ }^\circ\text{C}$), and ΔT is the change in temperature in $^\circ\text{C}$. The total thermal stress 100 m away from the well is 0.09 MPa, which is enough to induce seismicity. α is $3 \times 10^{-5} 1/\text{ }^\circ\text{C}$ and E ranges from 1 – 20 GPa (1,000 – 20,000 MPa) for sandstones [Robertson, 1988, UT JSG, 2018]. ΔT is 3°C at its maximum at an observation point 100 m away, resulting in a stress change between 0.09 – 1.8 MPa 100 m away from the disposal well.

Advection of heat with groundwater flow is critical to get the temperature decrease away from the disposal well. Velocity of the groundwater $V = K (dh/dl) / n$. K is the Hydraulic Conductivity (m/s) and is 10^{-7} m/s in the Dakota Sandstone. The Hydraulic Gradient (dh/dl) is 2×10^{-3} , which is calculated from the pressure gradient as discovered from Nelson et al., [2013]. The western edge of the model is 17 MPa and the eastern edge is 15 MPa. Therefore, the gradient of $1733.96 - 1529.96 \text{ m} / 100,000 \text{ m}$ of the length of the basin is 2×10^{-3} . By dividing this by the effective porosity, of 0.14 for the Dakota Sandstone, the velocity of the groundwater in our model is calculated to be $1.43 \times 10^{-9} \text{ m/s}$. Therefore, the area of influence with heat that will move with groundwater over the 17 years that the model is ran (10^8 s) is approximately 0.14 m.

Our hypothesis that wells in high geothermal gradient areas have higher correlation with earthquake occurrence due to thermal contrast from relatively cold wastewater injection into warmer formation rock was not validated by the TOUGH2 modeling that was performed. Although three of the wells that show a temperature decrease were located in high geothermal gradient areas, the injected water did not result in a temperature decrease close to the depth where earthquakes are occurring.

5.1 Limitations and Future Work

A number of limitations were encountered during this study. Regarding the modeling, limitations include the number of cells that the graphical interface, PetraSim, allows. The number of cells corresponds to the model grid. A finer grid (smaller area per individual cell) allows for a more accurate depiction of changes in temperature and pressure in a given area. Therefore, ideally more refined grids in the area of the disposal well sites may show the most change. The maximum number of cells that can be used in PetraSim is 200,000. Therefore, grid refinement had to be manipulated to show finer resolution at well sites without losing accuracy or exceeding the 200,000 cell limit. There is an upgraded version of PetraSim that allows for millions of cells, but was not feasible for this study due to the cost of the upgrade.

In addition to model limitations, there are also data limitations. For example, disposal well information prior to 2006 was not available for the New Mexico side of Raton Basin [NMOCD, 2016]. In addition, not all wastewater produced from CBM production is disposed of in the sub-surface. Therefore, it is not an accurate assumption that all produced water in Raton Basin is being disposed of in these disposal wells. Along with injection well data limitations, fault locations are another major gap in data collection. Seismic data would have been very

helpful to accurately determine the total number and the location of faults in the subsurface.

There was not any publically available seismic data.

Another limitation is uncertainty in the earthquakes' horizontal and depth locations. The USGS ANSS catalog was used in this study because it included a wide variety of earthquakes $M \geq 3$ that occurred during the time frame investigated for this study (1999 – 2016). The USGS has depth error up to 8 km (Figure 2). Horizontal error is also a limitation and corresponds to as much as 5.7 km of potential error. [Table 6 - Appendix] There are other catalogs that have focused on specific sequences of earthquakes. These catalogs include *Meremonte et al.* [2002], *Rubinstein et al.*, [2014], *Barnhart et al.*, [2014], and *Nakai et al.*, [2017].

Meremonte et al. [2002] investigated aftershocks and earthquakes associated with the August 2001 swarms in Raton Basin. Meremonte deployed 12 seismographs on September 6th and recorded events from September 7th – October 15th in the Fall of 2001. A total of 39 earthquakes were located and Meremonte was able to identify a 6 km long north-east trending normal fault with a dip of 70 – 80 ° to the southeast [Table 10 – Appendix]. This fault extends from 3 – 6 km bgs. Most earthquakes in this 2001 catalog occur at these depths and their depth errors are +/- 1.5 km (Figure 38). This catalog supports the hypothesis that earthquakes are occurring along faults at crystalline basement depths.

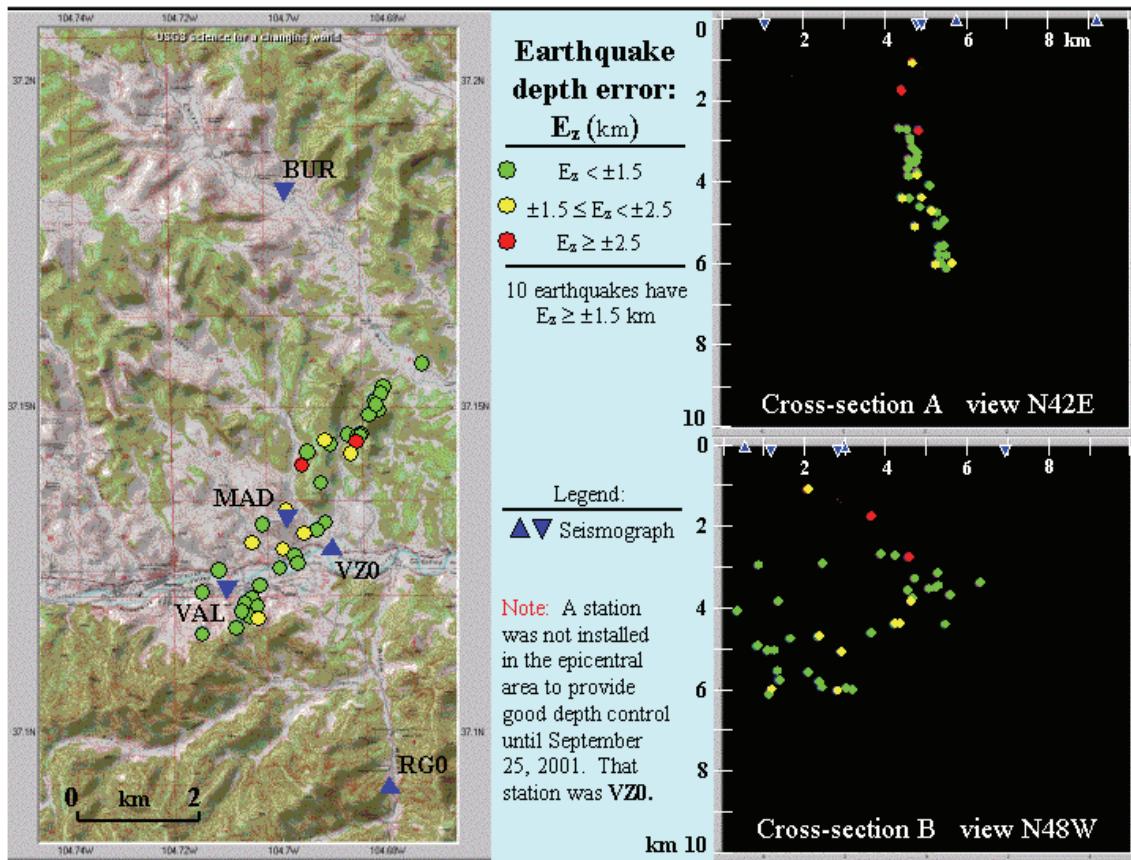


Figure 38. Depth errors in Meremonte catalog from aftershocks to the August 2001 earthquakes [Meremonte *et al.*, 2002].

Barnhart *et al.*, [2014] looked at earthquakes associated with the 2011 Mw 5.3 event.

Following this event, four more seismometers were deployed in the region to record aftershocks. The USGS National Earthquake Information Center (NEIC) located 584 aftershocks from August 23, 2011 to December 15, 2011. These were traced to the crystalline basement along an 8 – 10 km normal fault. Depths range from 1.5 to 6 km. Rubinstein *et al.*, [2014] analyzed $M \geq 3.8$ earthquakes from August 2001 – 2013 in Raton Basin. 16 events were observed and have epicenter uncertainties of +/- 15 km. Although the earthquakes were located without the use of high density array, the uncertainties for the 2001 and 2011 sequences have lower uncertainties of 2 km. Epicenters for these events are within 2 km of high, volume, high injection rate wells.

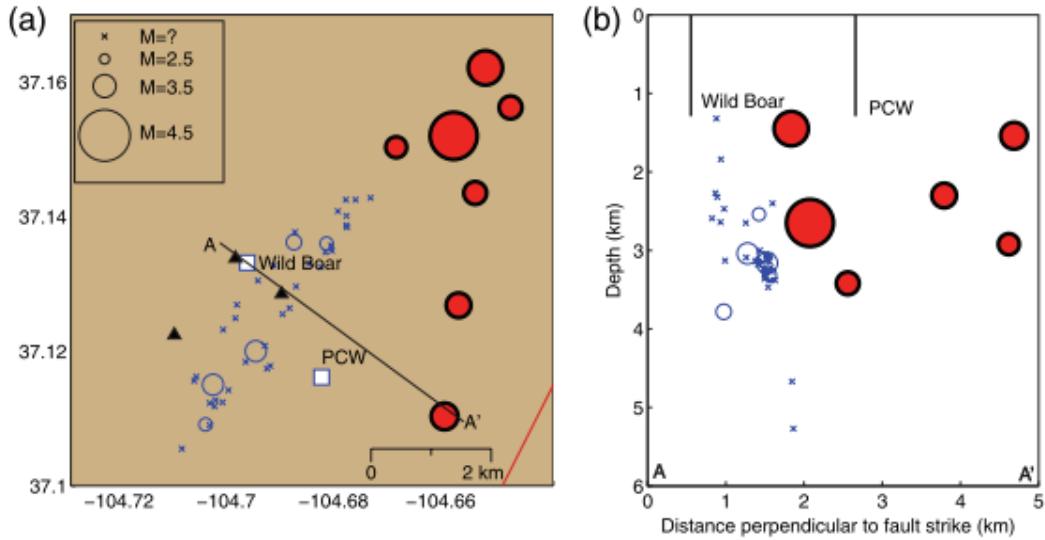


Figure 39. Earthquake epicenters near the Wild Boar and PCW disposal wells. A) Plan view of the earthquakes. B) Cross section view with the two disposal wells. The lineation of earthquakes corresponds to the modeled fault that was added into the model (Figure 23) [Rubinstein *et al.* 2014]

Nakai provided the last catalog that was used to determine epicenter depth. The EarthScope Transportable Array (TA) provided additional seismic coverage that allowed for improved earthquake epicenter locations. Nakai et al., [2017] manually located 1,881 earthquakes with P and S waves from 2008 - 2010. A majority of the earthquakes were located between 4 – 6 km, at basement depths, but had depths extending from 1 – 20 km depending on the velocity model that was used. The epicentral errors for Nakai et al., 2017 catalogs are +/- 2 km for the catalog. The depths Nakai determined are similar to Rubinstein *et al.*, [2014] and Barnhart *et al.* [2014].

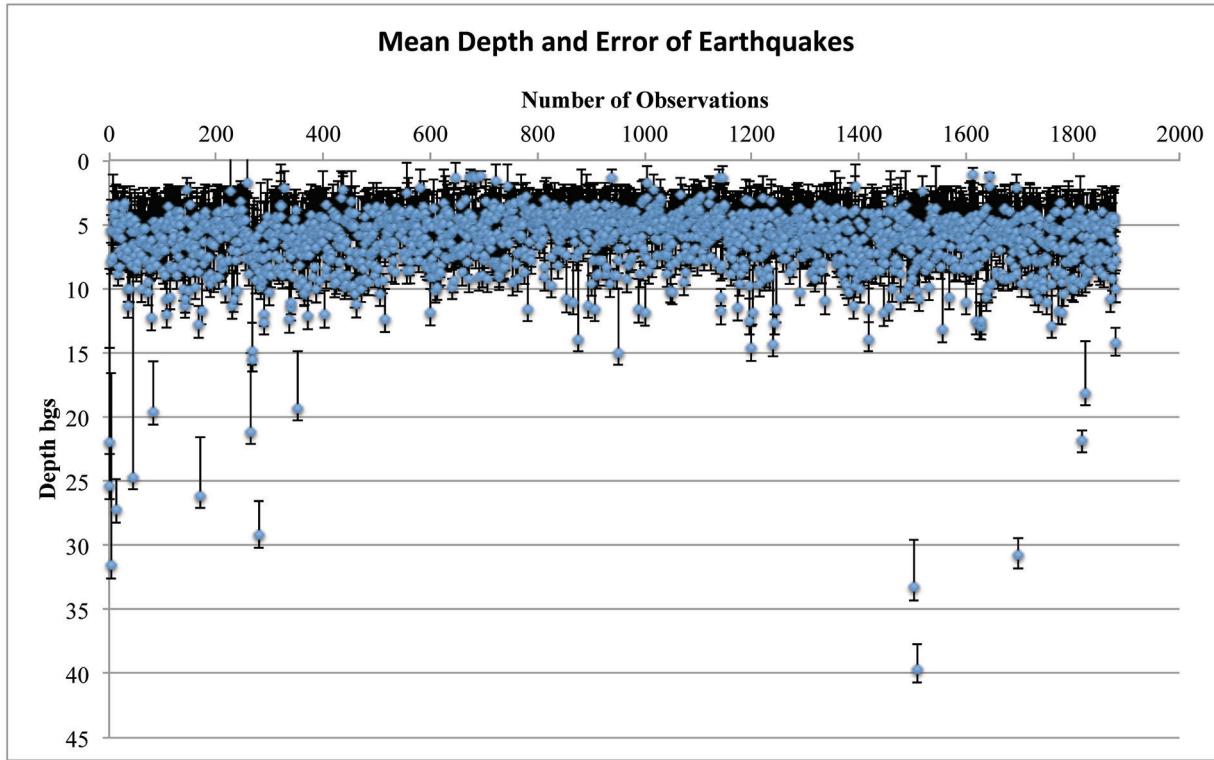


Figure 40. Mean depth with vertical depth error for earthquakes from Nakai's Catalog [Nakai et al., 2017]

6. CONCLUSION

Model results show that temperature can decrease up to 3°C at 100 m from injection wells over a period of 17 years in the injection interval of Raton's wastewater disposal wells. This cooling does not appear to cause significant thermal contraction and therefore rock failure because it is not reaching the depths at which earthquakes are occurring. Results from the crystalline basement directly below the injection wells show that temperature remains unchanged throughout the 17 years. (Figures 18 – 23). Thus the geothermal gradient does not seem to play any significant role with respect to induced seismicity as shown by hydrogeologic and thermal modeling. Modeling shows that there is only a limited spatial relationship between the geothermal gradient, injection wells, and earthquakes. Our modeling further solidified the theory that pore pressure increase in the basement, as a result of wastewater injection, is a major

mechanism for induced seismicity in Raton Basin as it is within the range to create earthquakes (>0.01 MPa) [*King et al.*, 1994].

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Appendix

Table 3. Disposal Well Locations

| Name | Latitude | Longitude | Top Injection Depth(m) | Bottom Injection depth (m) |
|--------------------------------|----------|-----------|------------------------------|----------------------------------|
| Jarosa 32-33 WD | 37.09 | -104.78 | 1667 | 1710 |
| SouthPaw | 37.3 | -104.729 | 1711 | 1710 |
| Cottontail Pass Disposal 32-33 | 37.22 | -104.78 | 1899 | 1986 |
| Cimarron 32-18 WD | 37.26 | -104.93 | 2116 | 2149 |
| Ferminia | 37.29 | -104.83 | 1834 | 1899 |
| Sawtooth | 37.2 | -104.67 | 1344 | 1383 |
| La Garita | 37.16 | -104.8 | 1624 | 1632 |
| San Pablo 11-4 WD | 37.206 | -104.899 | 1879 | 2077 |
| Weston 24-23 A WD | 37.15 | -104.86 | 1680 | 1710 |
| Long Canyon 43-12 WD | 37.09 | -104.62 | 1128 | 1172 |
| Apache Canyon 19-10 | 37.07 | -104.93 | 1879 | 1938 |
| Apache Canyon 10-3 | 37.1 | -104.99 | 2147 | 2179 |
| VPR C 204 WDW | 37.01 | -104.83 | 1648 | 1929 |
| PCW | 37.12 | -104.68 | 1208 | 1225 |
| Wild Boar | 37.13 | -104.7 | 1257 | 1292 |
| Hill Ranch Deep | 37.09 | -104.74 | 1454 | 1489 |
| Beardon | 37.25 | -104.66 | 1459 | 1489 |
| Polly | 37.23 | -104.7 | 1431 | 1612 |
| Del Agua | 37.28 | -104.74 | 1733 | 1737 |
| Lopez Canyon SWD | 37.15 | -104.89 | 1837 | 1985 |
| VPR A 007 | 36.955 | -105.4 | 1951 | 2001 |
| VPR C 39 | 37.023 | -104.78 | 1973 | 2007 |
| VPR C 14 | 37.02 | -104.78 | 1682 | 2020 |
| VPR A 182 | 36.98 | -104.8 | 1842 | 2148 |
| VPR A 500 | 36.89 | -104.75 | 2078 | 2171 |
| VPR B 027 | 36.8 | -104.94 | 1950 | 2240 |
| VPR D 025 | 36.86 | -105.057 | 1935 | 2240 |
| VPR A 042 | 36.96 | -104.83 | 2063 | 2273 |
| VPR E 099 | 36.96 | -104.9 | 2163 | 2301 |

Table 4. Disposal Operator and target formation for disposal [COGCC, 2016; NMOCD, 2016].

| Name | Operator | Years Active | Formation of Injection |
|--------------------------------|---------------------------|--------------|---------------------------------------|
| Jarosa 32-33 WD | Pioneer Natural Resources | 2007-2016 | Dakota, Purgatoire |
| SouthPaw | Pioneer Natural Resources | 2009-2016 | Dakota |
| Cottontail Pass Disposal 32-33 | Pioneer Natural Resources | 1999-2016 | Dakota, Purgatoire |
| Cimarron 32-18 WD | Pioneer Natural Resources | 2005-2016 | Dakota, Purgatoire |
| Ferminia | Pioneer Natural Resources | 2007-2016 | Dakota |
| Sawtooth | Pioneer Natural Resources | 2000-2016 | Dakota, Purgatoire |
| La Garita | Pioneer Natural Resources | 2001-2016 | Dakota |
| San Pablo 11-4 WD | Pioneer Natural Resources | 2014-2016 | Dakota, Entrada, Morrison, Purgatoire |
| Weston 24-23 A WD | Pioneer Natural Resources | 2004-2016 | Dakota (primary Purgatoire) |
| Long Canyon 43-12 WD | Pioneer Natural Resources | 2004-2016 | Dakota, Purgatoire |
| Apache Canyon 19-10 | XTO Energy | 1999-2016 | Dockum |
| Apache Canyon 10-3 | XTO Energy | 1999-2016 | Dakota, Entrada |
| VPR C 204 WDW | ARP Production Company | 2012-2016 | Dakota, Entrada, Glorieta |
| PCW | Pioneer Natural Resources | 1999-2017 | Dakota |
| Wild Boar | Pioneer Natural Resources | 1999-2018 | Dakota, Purgatoire |
| Hill Ranch Deep | Pioneer Natural Resources | 2005-2016 | Dakota |
| Beardon | Pioneer Natural Resources | 2001-2016 | Dakota, Purgatoire |
| Polly | Pioneer Natural Resources | 2009-2015 | Dakota, Entrada |
| Del Agua | Pioneer Natural | 2005-2013 | Dakota |

| | Resources | | |
|------------------|---------------------------|-----------|-----------------------------------|
| Lopez Canyon SWD | XTO Energy | 2010-2016 | Dakota, Entrada, Purgatoire |
| VPR A 007 | ARP Production Company | 2006-2016 | Dakota |
| VPR C 39 | ARP Production Company | 2000-2016 | Entrada, Glorieta |
| VPR C 14 | SONAT Raton LLC | 1999-2002 | |
| VPR A 182 | ARP Production Company | 2006-2016 | Entrada and Glorieta |
| VPR A 500 | ARP Production Company | 2006-2017 | Glorieta, Dakota |
| VPR B 027 | ARP Production Company | 2006-2018 | Entrada and Glorieta |
| VPR D 025 | ARP Production Company | 2006-2019 | Dakota, Entrada |
| VPR A 042 | ARP Production Company | 2006-2020 | Dakota, Entrada |
| VPR E 099 | ARP Production Company | 2006-2021 | Entrada and Glorieta |

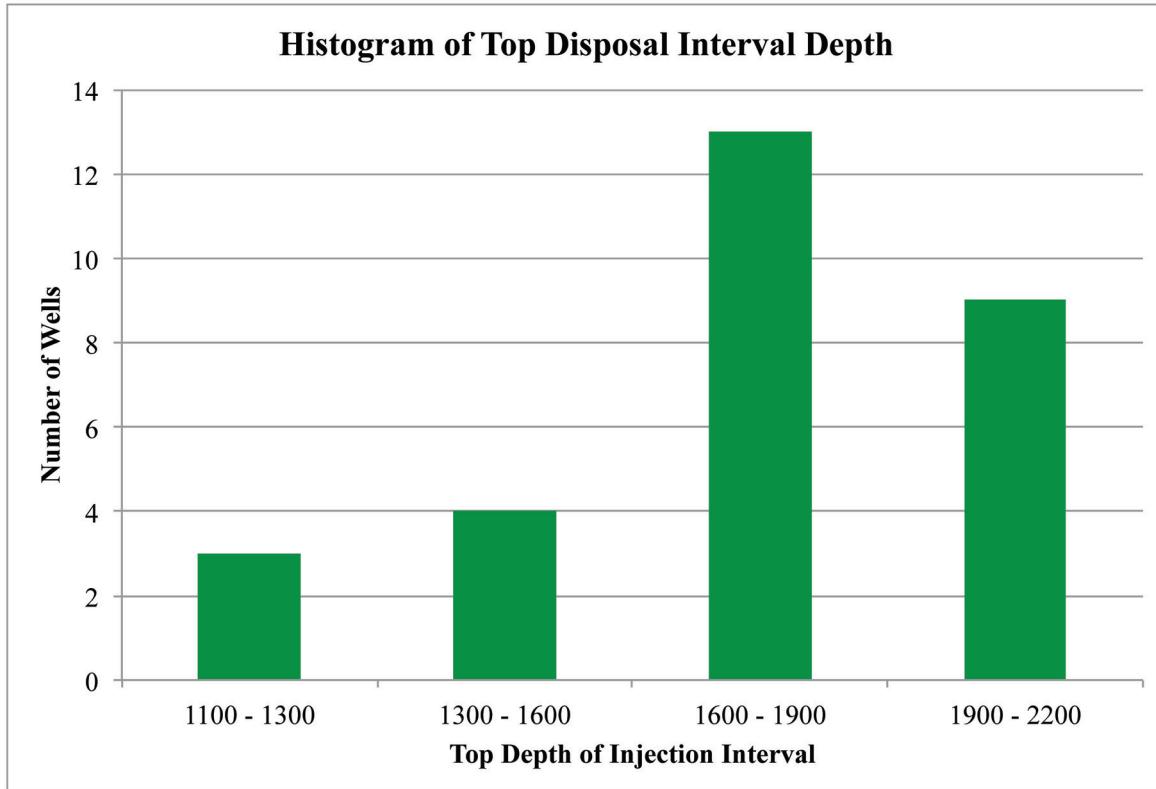


Figure 41. Histogram of injection well depths [COGCC, 2016; NMOCD, 2016].

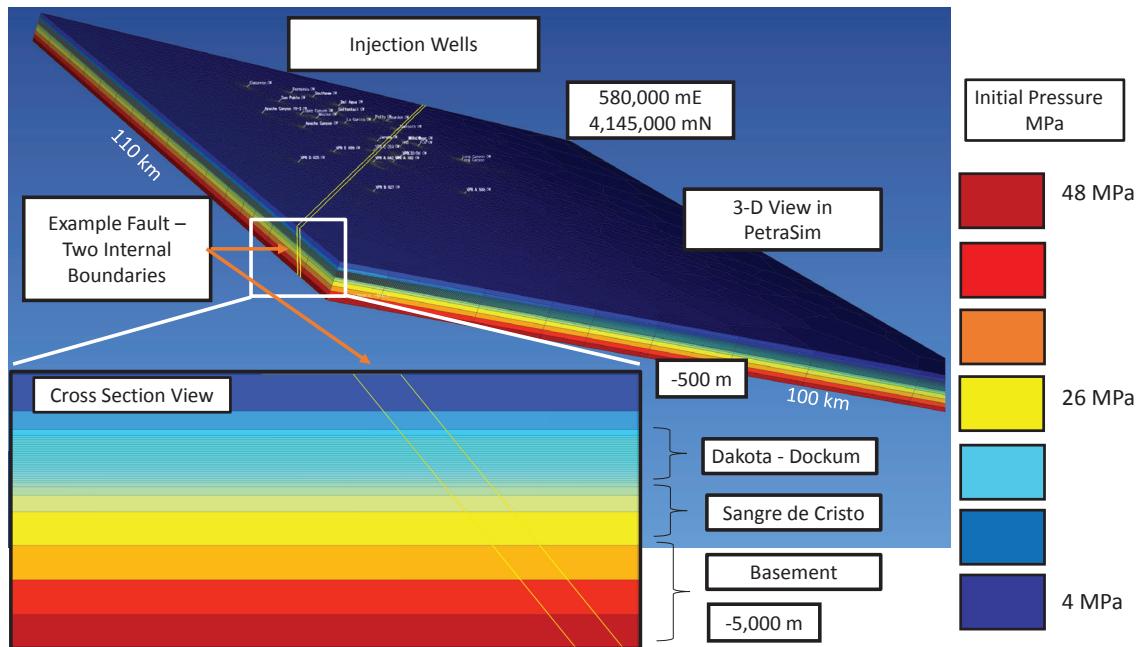


Figure 42. Model in 3-D with Initial Pressure shown in MPa. Cross Section view shows the grid refinement in the disposal interval.

Table 5. Earthquake Locations from USGS Catalog [ANSS, 2018]

| Date | Time | Lat | Long | Depth | Magnitude | Measurement Type |
|----------|---------|--------|----------|-------|-----------|------------------|
| 9/23/73 | 58:54.9 | 37.148 | -104.571 | 5 | 4.2 | Mb |
| 4/15/92 | 46:05.1 | 37.335 | -104.773 | 5 | 3.3 | Unknown |
| 5/2/92 | 19:29.8 | 37.378 | -104.778 | 5 | 3.1 | Unknown |
| 8/28/01 | 16:09.5 | 37.088 | -104.692 | 5 | 3.4 | ML |
| 8/28/01 | 22:00.3 | 37.091 | -104.655 | 5 | 3.5 | ML |
| 9/4/01 | 22:45.0 | 37.107 | -104.622 | 5 | 3.4 | ML |
| 9/4/01 | 45:53.2 | 37.143 | -104.65 | 5 | 4 | ML |
| 9/5/01 | 52:07.9 | 37.143 | -104.618 | 5 | 4.5 | ML |
| 9/5/01 | 48:58.3 | 37.112 | -104.611 | 5 | 3.7 | ML |
| 9/6/01 | 41:43.6 | 37.11 | -104.628 | 5 | 3.6 | ML |
| 9/6/01 | 28:26.5 | 37.14 | -104.585 | 5 | 3.5 | ML |
| 9/10/01 | 56:00.4 | 37.108 | -104.602 | 5 | 3.4 | ML |
| 9/13/01 | 39:05.4 | 37.091 | -104.593 | 5 | 3 | ML |
| 9/21/01 | 10:59.7 | 37.121 | -104.706 | 5.4 | 3.4 | ML |
| 12/15/01 | 58:31.4 | 36.859 | -104.797 | 5 | 3.3 | ML |
| 1/26/02 | 06:03.9 | 36.86 | -104.784 | 5 | 3.4 | ML |

| | | | | | | |
|----------|---------|--------|----------|---|-----|----|
| 6/18/02 | 12:36.7 | 36.881 | -104.779 | 5 | 3.5 | ML |
| 11/14/02 | 56:52.3 | 36.917 | -104.768 | 5 | 3.2 | ML |
| 4/28/03 | 32:26.0 | 36.844 | -104.923 | 5 | 3.6 | ML |
| 6/3/03 | 09:27.8 | 36.994 | -104.768 | 5 | 3.3 | ML |
| 6/15/03 | 22:18.0 | 36.91 | -104.763 | 5 | 3.6 | ML |
| 8/14/03 | 11:09.0 | 36.945 | -104.87 | 5 | 3.3 | ML |
| 9/8/03 | 02:49.3 | 37.369 | -104.685 | 5 | 3 | ML |
| 9/13/03 | 22:41.0 | 36.831 | -104.907 | 5 | 3.8 | ML |
| 11/24/03 | 05:57.7 | 36.958 | -104.828 | 5 | 3.1 | ML |
| 2/3/04 | 34:22.6 | 36.932 | -104.861 | 5 | 3.4 | ML |
| 3/22/04 | 09:56.5 | 36.855 | -104.851 | 5 | 4.4 | Mb |
| 3/30/04 | 02:55.4 | 36.892 | -104.876 | 5 | 3 | ML |
| 3/30/04 | 23:37.9 | 36.876 | -104.831 | 5 | 3.1 | ML |
| 3/30/04 | 41:04.1 | 37.036 | -104.931 | 5 | 3.5 | ML |
| 5/31/04 | 27:43.8 | 36.935 | -104.835 | 5 | 3.3 | ML |
| 8/1/04 | 50:47.6 | 36.874 | -105.104 | 5 | 4.3 | Mw |
| 1/10/05 | 14:59.1 | 37.007 | -104.675 | 5 | 3.4 | ML |
| 4/24/05 | 02:35.9 | 36.92 | -105.07 | 5 | 3.4 | ML |
| 7/4/05 | 45:24.5 | 36.86 | -105.097 | 5 | 3 | ML |
| 7/8/05 | 24:01.1 | 36.938 | -104.886 | 5 | 3 | ML |
| 8/10/05 | 08:17.0 | 36.952 | -104.822 | 5 | 4.1 | Mb |
| 8/10/05 | 08:22.6 | 36.947 | -104.833 | 5 | 5 | Mw |
| 8/10/05 | 24:33.9 | 36.982 | -104.959 | 5 | 3 | ML |
| 10/20/05 | 15:36.6 | 36.97 | -104.849 | 5 | 3 | ML |
| 1/27/06 | 48:49.2 | 37.03 | -104.968 | 5 | 3.3 | ML |
| 5/6/06 | 07:01.3 | 37.014 | -104.768 | 5 | 3.1 | ML |
| 5/26/06 | 14:25.1 | 36.795 | -104.832 | 5 | 3.1 | ML |
| 7/11/06 | 53:37.8 | 36.964 | -104.929 | 5 | 3.2 | ML |
| 8/24/06 | 04:25.9 | 37.014 | -105.013 | 5 | 3.1 | ML |
| 9/9/06 | 54:06.6 | 37.296 | -104.77 | 5 | 3.2 | ML |
| 9/9/06 | 53:14.2 | 37.368 | -104.865 | 5 | 3 | ML |
| 9/9/06 | 14:35.5 | 37.298 | -104.794 | 5 | 3.6 | ML |
| 9/14/06 | 03:24.3 | 37.01 | -104.867 | 5 | 3 | ML |
| 10/30/06 | 35:13.5 | 36.811 | -104.963 | 5 | 3.5 | ML |
| 11/24/06 | 22:24.1 | 37.04 | -104.996 | 5 | 3.1 | ML |
| 12/24/06 | 50:21.5 | 36.935 | -104.75 | 5 | 3.6 | ML |
| 1/3/07 | 34:38.5 | 37.067 | -104.895 | 5 | 4.4 | Mw |
| 1/14/07 | 17:36.7 | 36.878 | -104.93 | 5 | 3.2 | ML |
| 2/25/07 | 24:19.1 | 37.099 | -104.773 | 5 | 3.1 | ML |
| 3/12/07 | 32:14.6 | 37.061 | -104.937 | 5 | 3.4 | ML |

| | | | | | | |
|----------|---------|--------|----------|---|-----|----|
| 6/9/07 | 45:44.7 | 36.929 | -104.793 | 1 | 3.3 | Mw |
| 1/29/08 | 30:24.3 | 36.871 | -104.988 | 5 | 3.1 | ML |
| 4/21/08 | 36:29.9 | 37.158 | -104.942 | 5 | 3.2 | ML |
| 4/24/08 | 21:51.4 | 37.032 | -104.847 | 5 | 3.1 | ML |
| 8/24/08 | 48:31.5 | 37.095 | -104.866 | 5 | 3.4 | ML |
| 9/25/08 | 55:35.3 | 37.357 | -104.882 | 5 | 3.2 | ML |
| 10/4/08 | 41:20.9 | 37.263 | -104.748 | 5 | 3.4 | ML |
| 2/3/09 | 27:10.3 | 36.992 | -104.884 | 5 | 3 | ML |
| 3/22/09 | 14:40.1 | 37.261 | -104.462 | 5 | 3 | ML |
| 5/1/09 | 34:03.9 | 36.818 | -104.819 | 5 | 3.2 | Mw |
| 6/27/09 | 44:39.9 | 36.888 | -104.906 | 5 | 3 | ML |
| 7/29/09 | 00:36.7 | 36.799 | -104.831 | 5 | 4.1 | Mw |
| 9/29/09 | 20:27.8 | 37.055 | -104.995 | 5 | 3.1 | ML |
| 9/29/09 | 54:07.5 | 37.003 | -104.805 | 5 | 3.5 | Mw |
| 10/3/09 | 45:31.9 | 37.022 | -104.855 | 5 | 3.6 | ML |
| 11/20/09 | 54:30.1 | 36.892 | -104.987 | 5 | 3.7 | ML |
| 12/11/09 | 32:27.2 | 36.927 | -105.03 | 5 | 3.2 | ML |
| 1/18/10 | 41:07.4 | 36.854 | -104.819 | 5 | 3.8 | Mw |
| 4/8/10 | 36:57.4 | 36.916 | -104.842 | 5 | 3.5 | ML |
| 5/27/10 | 43:11.1 | 37.012 | -104.906 | 5 | 3.6 | ML |
| 2/13/11 | 44:52.6 | 37.005 | -104.935 | 5 | 3.6 | ML |
| 3/12/11 | 16:05.6 | 36.861 | -104.981 | 5 | 3.2 | ML |
| 5/9/11 | 28:52.8 | 37.021 | -104.783 | 5 | 3.7 | Mw |
| 5/11/11 | 06:15.3 | 37.1 | -104.665 | 5 | 3.8 | Mw |
| 8/22/11 | 30:19.9 | 37.032 | -104.554 | 5 | 4.7 | Mw |
| 8/23/11 | 48:51.0 | 37.023 | -104.667 | 5 | 3 | ML |
| 8/23/11 | 46:18.2 | 37.063 | -104.701 | 4 | 5.4 | Mw |
| 8/23/11 | 56:58.5 | 37.113 | -104.631 | 5 | 3.5 | ML |
| 8/23/11 | 01:34.1 | 37.105 | -104.679 | 5 | 3.2 | ML |
| 8/23/11 | 17:58.3 | 37.101 | -104.63 | 5 | 3.5 | Mw |
| 8/23/11 | 37:56.8 | 37.068 | -104.764 | 5 | 3.2 | ML |
| 8/23/11 | 11:12.8 | 37.055 | -104.692 | 5 | 4 | Mw |
| 8/24/11 | 15:57.3 | 37.129 | -104.803 | 5 | 3 | ML |
| 8/25/11 | 44:38.4 | 37.123 | -104.697 | 5 | 3.2 | ML |
| 9/13/11 | 37:18.5 | 36.94 | -104.798 | 5 | 3.5 | ML |
| 9/13/11 | 24:39.6 | 36.933 | -104.741 | 5 | 4 | ML |
| 9/16/11 | 51:51.3 | 36.884 | -104.74 | 5 | 3.9 | Mw |
| 10/10/11 | 26:29.2 | 37.09 | -104.688 | 5 | 3.2 | ML |
| 11/22/11 | 20:02.9 | 37.003 | -104.956 | 5 | 3.4 | ML |
| 12/9/11 | 54:15.8 | 37.099 | -104.656 | 5 | 3.4 | ML |

| | | | | | | |
|----------|---------|---------|-----------|------|-----|----|
| 12/17/11 | 16:49.3 | 37.043 | -104.913 | 5 | 3 | ML |
| 3/6/12 | 10:21.7 | 37.002 | -104.925 | 5 | 3.3 | ML |
| 5/24/12 | 44:37.6 | 37.123 | -104.743 | 5 | 3.2 | ML |
| 9/16/12 | 15:31.7 | 37.064 | -104.825 | 5 | 3 | ML |
| 12/4/12 | 15:26.8 | 37.033 | -104.938 | 5 | 3.5 | Mw |
| 12/13/12 | 05:52.2 | 36.952 | -104.783 | 8.3 | 3.3 | ML |
| 4/20/13 | 47:15.9 | 36.973 | -104.875 | 5 | 3 | ML |
| 9/8/13 | 15:31.1 | 36.9651 | -104.8757 | 5 | 3.6 | Mw |
| 1/22/14 | 45:03.5 | 37.2381 | -104.6215 | 1.26 | 3.6 | ML |
| 4/11/14 | 47:46.2 | 37.1848 | -104.9467 | 7.44 | 3.2 | ML |
| 7/10/14 | 22:47.6 | 37.0008 | -104.9669 | 3.91 | 3.1 | ML |
| 10/10/14 | 57:21.9 | 37.0836 | -104.9037 | 1.36 | 3.3 | Mb |
| 10/17/14 | 01:05.4 | 37.0163 | -104.9216 | 4.6 | 3.8 | ML |
| 10/24/14 | 50:47.8 | 36.9972 | -104.9233 | 5 | 3.3 | ML |
| 11/2/14 | 16:18.7 | 37.0753 | -104.9121 | 6.07 | 3 | ML |
| 12/2/14 | 57:20.3 | 37.1054 | -104.9393 | 2.99 | 3 | ML |
| 12/23/14 | 33:36.4 | 36.9501 | -104.7718 | 8.59 | 3.4 | ML |
| 12/27/14 | 23:24.8 | 37.0159 | -104.8487 | 3.51 | 3 | ML |
| 1/7/15 | 40:26.6 | 36.8765 | -104.751 | 4.33 | 3.7 | ML |
| 2/2/15 | 10:13.0 | 37.0845 | -104.975 | 5 | 3.1 | ML |
| 2/22/15 | 31:48.2 | 37.1545 | -104.978 | 5 | 3.4 | ML |
| 2/22/15 | 20:34.7 | 37.0503 | -104.9507 | 5 | 3.6 | ML |
| 3/5/15 | 49:01.0 | 36.9283 | -104.8993 | 1.81 | 3.4 | Mb |
| 3/7/15 | 59:35.5 | 37.0771 | -104.8742 | 5.74 | 3.5 | ML |
| 3/14/15 | 17:30.6 | 36.9028 | -104.7774 | 1.64 | 3.6 | ML |
| 3/20/15 | 19:18.9 | 36.8932 | -104.7459 | 5 | 3.8 | ML |
| 8/20/15 | 14:09.5 | 37.1078 | -104.9291 | 1.08 | 3.9 | Mw |
| 12/29/15 | 25:32.5 | 37.1308 | -104.9126 | 5 | 3.2 | ML |
| 2/3/16 | 59:19.0 | 37.1511 | -104.9396 | 1.42 | 3.5 | ML |
| 2/6/16 | 09:11.0 | 37.0868 | -104.957 | 1.57 | 4 | Mb |
| 5/21/16 | 51:16.8 | 36.9398 | -104.8106 | 2.3 | 3.2 | ML |
| 8/23/16 | 56:11.5 | 36.9863 | -104.9449 | 3.6 | 3.9 | Mw |
| 11/6/16 | 48:36.4 | 36.8997 | -104.8605 | 5 | 3.2 | ML |
| 12/23/16 | 31:13.0 | 36.7571 | -104.9277 | 3.55 | 4.2 | Mw |
| 1/21/17 | 57:54.3 | 36.9566 | -104.8425 | 6.27 | 3.1 | ML |

Table 6. Earthquake Depth and Horizontal Errors from USGS Catalog [ANSS, 2018].

| Date | Lat | Long | Depth | Magnitude | Horizontal Error (km) | Depth Error (km) |
|----------------|---------|-----------|-------|-----------|-----------------------|------------------|
| 9/8/13 9:04 | 36.9372 | -104.9105 | 7.46 | 2.6 | 4.7 | 7.1 |
| 8/31/14 10:20 | 37.0434 | -104.9181 | 7.2 | 2.7 | 4.3 | 8.1 |
| 9/30/14 9:55 | 37.0584 | -104.9273 | 6.67 | 2.9 | 3.7 | 3.1 |
| 10/6/14 8:10 | 37.0937 | -104.8877 | 4.44 | 2.9 | 4 | 7.5 |
| 10/8/14 0:16 | 37.1423 | -104.9102 | 2.32 | 2.6 | 2 | 8 |
| 10/10/14 22:57 | 37.0836 | -104.9037 | 1.36 | 3.3 | 5.7 | 7.2 |
| 10/17/14 17:01 | 37.0163 | -104.9216 | 4.6 | 3.8 | 4.3 | 6.9 |
| 10/20/14 4:01 | 37.172 | -104.9084 | 5 | 2.5 | 1.8 | 2 |
| 10/21/14 4:33 | 37.0675 | -104.9663 | 5 | 2.6 | 2.2 | 1.9 |
| 10/24/14 12:50 | 36.9972 | -104.9233 | 5 | 3.3 | 3.3 | 1.8 |
| 10/26/14 20:11 | 37.0106 | -104.9475 | 5 | 2.5 | 2.3 | 2 |
| 10/28/14 6:26 | 37.0768 | -104.9233 | 5 | 2.7 | 3.4 | 2 |
| 11/2/14 13:16 | 37.0753 | -104.9121 | 6.07 | 3 | 4.7 | 7 |
| 11/4/14 16:59 | 36.9459 | -104.8544 | 2.23 | 2.7 | 3.3 | 8 |
| 11/8/14 1:31 | 37.0645 | -104.9411 | 5.02 | 2.8 | 4.9 | 8 |
| 12/2/14 21:57 | 37.1054 | -104.9393 | 2.99 | 3 | 3.1 | 7 |
| 12/23/14 12:33 | 36.9501 | -104.7718 | 8.59 | 3.4 | 5.3 | 7.5 |
| 12/27/14 8:23 | 37.0159 | -104.8487 | 3.51 | 3 | 4.2 | 6.7 |
| 1/7/15 0:40 | 36.8765 | -104.751 | 4.33 | 3.7 | 5 | 6.6 |
| 1/10/15 19:04 | 37.1436 | -104.8777 | 1.34 | 2.5 | 1.9 | 8 |
| 1/20/15 13:53 | 36.8765 | -105.0072 | 0 | 3.3 | 5.5 | 2 |
| 1/31/15 10:48 | 36.9807 | -104.8584 | 5 | 2.6 | 4.6 | 2 |
| 2/2/15 14:10 | 37.0845 | -104.975 | 5 | 3.1 | 3 | 2 |
| 2/22/15 7:31 | 37.0797 | -105.0638 | 1.31 | 2.6 | 4.1 | 7.9 |
| 2/22/15 7:31 | 37.1555 | -104.9824 | 5 | 3.6 | 4.6 | 2 |
| 2/22/15 7:36 | 37.1008 | -105.0038 | 5 | 2.7 | 4.1 | 2 |
| 2/22/15 8:20 | 37.0503 | -104.9507 | 5 | 3.6 | 2.5 | 1.8 |
| 3/5/15 14:49 | 36.9461 | -104.8989 | 2.52 | 3.4 | 4.4 | 7.5 |
| 3/7/15 8:59 | 37.0719 | -104.9376 | 1.17 | 3.5 | 2.9 | 7.2 |
| 3/9/15 18:14 | 37.0553 | -104.9871 | 5 | 2.7 | 2.2 | 2 |
| 3/14/15 20:17 | 36.9028 | -104.7774 | 1.64 | 3.6 | 5.3 | 7.5 |
| 3/19/15 11:48 | 36.8128 | -104.6004 | 5 | 2.6 | 3.5 | 2 |
| 3/20/15 5:19 | 36.8932 | -104.7459 | 5 | 3.8 | 5 | 0.7 |
| 4/22/15 18:28 | 37.091 | -104.9699 | 5 | 2.5 | 3.5 | 2 |
| 4/27/15 4:49 | 37.0527 | -104.985 | 5 | 2.9 | 3.3 | 2 |

| | | | | | | |
|----------------|---------|-----------|------|-----|-----|------|
| 8/18/15 2:37 | 36.9044 | -104.8848 | 3.49 | 2.7 | 3.1 | 2.1 |
| 8/20/15 5:14 | 37.1001 | -104.921 | 1.08 | 3.9 | 4.9 | 7.6 |
| 8/25/15 23:08 | 37.162 | -104.9149 | 5 | 2.8 | 3 | 2 |
| 10/14/15 8:23 | 37.0651 | -104.8339 | 5 | 2.8 | 3.1 | 2 |
| 10/18/15 1:47 | 37.0899 | -104.8818 | 6.54 | 2.9 | 1.9 | 8 |
| 12/29/15 12:25 | 37.1308 | -104.9126 | 5 | 3 | 4.1 | 2 |
| 1/6/16 12:45 | 36.9218 | -104.8414 | 5 | 2.6 | 2.7 | 2 |
| 1/13/16 18:58 | 36.9677 | -104.8247 | 6.74 | 2.6 | 4.3 | 5.6 |
| 1/15/16 1:54 | 36.928 | -104.7711 | 7.59 | 3 | 5.3 | 8 |
| 2/3/16 14:59 | 37.1533 | -104.9544 | 3.13 | 3.3 | 3.2 | 6.9 |
| 2/6/16 23:09 | 37.0868 | -104.957 | 1.57 | 4 | 2.4 | 8.3 |
| 3/5/16 3:13 | 37.2732 | -104.9779 | 5 | 2.9 | 1.9 | 2 |
| 3/13/16 22:17 | 37.054 | -104.9731 | 5 | 2.8 | 2 | 1.9 |
| 5/21/16 20:51 | 36.9398 | -104.8106 | 2.3 | 3.2 | 4.4 | 7.9 |
| 6/19/16 5:46 | 36.9366 | -104.8569 | 1.83 | 2.9 | 3.4 | 8 |
| 8/7/16 9:28 | 36.8459 | -105.1098 | 1.33 | 2.6 | 3.6 | 8.1 |
| 8/23/16 16:56 | 36.9863 | -104.9449 | 3.6 | 3.9 | 3.2 | 1.6 |
| 11/6/16 4:48 | 36.9233 | -104.889 | 1.49 | 3.2 | 5.4 | 2.9 |
| 12/6/16 16:41 | 36.9968 | -105.0412 | 5 | 2.7 | 1.2 | 2 |
| 1/21/17 3:57 | 36.9566 | -104.8425 | 6.27 | 3.1 | 3.3 | 6.5 |
| 6/11/17 8:58 | 37.0704 | -104.9702 | 5 | 2.8 | 2.8 | 2 |
| 6/17/17 6:42 | 37.0352 | -104.8986 | 1.93 | 3.5 | 4.9 | 11.1 |

Table 7. Nakai Catalog with depth errors. Depth_sd represents km vertical error. ml is magnitude (-999 is a magnitude that was not determined) [Nakai, personal comm.]

| year | mo | day | lat_mean | lon_mean | depth_mean | depth_sd | ml |
|------|----|-----|----------|-----------|------------|----------|------|
| 2009 | 6 | 5 | 36.8283 | -104.1325 | 25.4103 | 10.8221 | 1.07 |
| 2009 | 7 | 29 | 36.8537 | -104.1017 | 21.9162 | 13.5052 | 1.01 |
| 2009 | 11 | 2 | 36.9621 | -104.5801 | 31.5914 | 15.0534 | -999 |
| 2009 | 12 | 13 | 37.3051 | -104.9362 | 5.4287 | 2.3968 | -999 |
| 2009 | 12 | 15 | 37.3051 | -104.9335 | 7.8719 | 3.6 | 0.76 |
| 2009 | 12 | 16 | 36.8482 | -104.861 | 5.602 | 2.5561 | 1.48 |
| 2009 | 12 | 17 | 36.8726 | -104.8715 | 7.5483 | 2.8499 | 1.52 |
| 2009 | 12 | 17 | 36.8221 | -104.8482 | 3.6076 | 2.4683 | 1.51 |
| 2009 | 12 | 17 | 36.8681 | -104.8711 | 4.6788 | 2.2779 | 1.39 |
| 2009 | 12 | 17 | 36.8691 | -104.871 | 4.4205 | 2.0136 | 1.36 |
| 2009 | 12 | 19 | 37.0487 | -104.9785 | 4.1866 | 2.1012 | 1.41 |

| | | | | | | | |
|------|----|----|---------|-----------|---------|---------|------|
| 2009 | 12 | 19 | 36.9342 | -104.877 | 27.2308 | 2.3455 | 1.16 |
| 2009 | 12 | 20 | 37.1109 | -104.717 | 3.4321 | 1.4809 | 1.33 |
| 2009 | 12 | 22 | 37.0435 | -104.9665 | 5.466 | 2.9329 | 1.61 |
| 2009 | 12 | 23 | 37.3078 | -104.9333 | 4.9545 | 2.3525 | 1.84 |
| 2009 | 12 | 24 | 36.8553 | -104.8695 | 5.84 | 2.6247 | 1.37 |
| 2009 | 12 | 24 | 36.885 | -104.868 | 5.9135 | 3.3133 | 1.52 |
| 2009 | 12 | 24 | 36.8568 | -104.8723 | 8.8322 | 1.8332 | 1.58 |
| 2009 | 12 | 25 | 36.8548 | -104.87 | 8.2911 | 1.0354 | 1.85 |
| 2009 | 12 | 25 | 36.8961 | -104.8756 | 5.3818 | 2.3879 | 1.35 |
| 2009 | 12 | 25 | 36.8958 | -104.8716 | 7.0438 | 3.3564 | -999 |
| 2009 | 12 | 26 | 37.3066 | -104.9353 | 4.891 | 2.332 | 1.64 |
| 2009 | 12 | 27 | 36.8541 | -104.8721 | 6.0371 | 2.9781 | 0.99 |
| 2009 | 12 | 28 | 37.1942 | -104.7527 | 7.3589 | 4.091 | -999 |
| 2009 | 12 | 28 | 37.2317 | -104.6277 | 5.0993 | 1.9785 | 2.54 |
| 2009 | 12 | 28 | 37.2322 | -104.6265 | 4.1972 | 1.7365 | 2.08 |
| 2009 | 12 | 28 | 37.2315 | -104.6248 | 3.7986 | 1.69 | 1.76 |
| 2009 | 12 | 28 | 37.2337 | -104.6286 | 3.3575 | 1.3736 | 1.66 |
| 2009 | 12 | 28 | 37.2311 | -104.6283 | 4.6358 | 1.8596 | 2.3 |
| 2009 | 12 | 29 | 37.3401 | -104.2853 | 6.7859 | 3.0473 | 0.95 |
| 2009 | 12 | 30 | 37.3079 | -104.9362 | 7.5057 | 3.7569 | 1.26 |
| 2009 | 12 | 30 | 37.3047 | -104.9352 | 7.3993 | 3.8666 | -999 |
| 2009 | 12 | 31 | 37.3069 | -104.937 | 7.6131 | 1.1569 | 1.92 |
| 2010 | 1 | 1 | 37.3079 | -104.9294 | 6.9553 | 2.3081 | 1.69 |
| 2010 | 1 | 1 | 36.8083 | -104.8635 | 10.0324 | 4.2971 | -999 |
| 2010 | 1 | 1 | 37.3276 | -104.9158 | 11.3146 | 5.4677 | 1.32 |
| 2010 | 1 | 1 | 36.8789 | -104.8432 | 4.9214 | 2.3034 | -999 |
| 2010 | 1 | 2 | 37.3355 | -104.9035 | 10.0614 | 4.8016 | 1.14 |
| 2010 | 1 | 2 | 36.8469 | -104.8552 | 4.437 | 2.3736 | 1.34 |
| 2010 | 1 | 2 | 37.31 | -104.9356 | 7.7846 | 2.9593 | 1.4 |
| 2010 | 1 | 3 | 36.8787 | -104.8402 | 4.9789 | 2.4557 | 1.72 |
| 2010 | 1 | 3 | 36.8718 | -104.8313 | 8.1538 | 3.9888 | -999 |
| 2010 | 1 | 3 | 36.8983 | -104.8723 | 5.598 | 2.6631 | 1.37 |
| 2010 | 1 | 6 | 37.0908 | -105.0036 | 24.6865 | 17.6643 | 0.99 |
| 2010 | 1 | 6 | 36.8483 | -104.8671 | 6.5554 | 2.6802 | 1.77 |
| 2010 | 1 | 6 | 37.3117 | -104.9328 | 6.6124 | 2.5758 | -999 |
| 2010 | 1 | 7 | 36.8707 | -104.8494 | 5.5895 | 2.3835 | 1.37 |
| 2010 | 1 | 7 | 36.8697 | -104.8676 | 7.4283 | 3.9903 | -999 |
| 2010 | 1 | 7 | 36.8727 | -104.8519 | 4.6987 | 2.4014 | 1.29 |
| 2010 | 1 | 8 | 37.3123 | -104.936 | 8.1758 | 1.1975 | 2.46 |
| 2010 | 1 | 9 | 36.8843 | -104.8542 | 6.7835 | 2.8071 | 1.62 |

| | | | | | | | |
|------|---|----|---------|-----------|---------|--------|------|
| 2010 | 1 | 9 | 36.9029 | -104.8613 | 8.4442 | 4.4905 | -999 |
| 2010 | 1 | 10 | 37.0246 | -104.9735 | 5.2045 | 2.6132 | 1.63 |
| 2010 | 1 | 10 | 37.3143 | -104.9362 | 6.9901 | 1.4545 | 2.14 |
| 2010 | 1 | 11 | 36.9143 | -104.8482 | 7.9475 | 3.5323 | 1.64 |
| 2010 | 1 | 12 | 36.8885 | -104.8732 | 5.549 | 2.5596 | 2.46 |
| 2010 | 1 | 12 | 36.8912 | -104.8728 | 6.3723 | 3.1633 | 1.33 |
| 2010 | 1 | 12 | 36.8907 | -104.8725 | 6.3065 | 3.5374 | 1.08 |
| 2010 | 1 | 12 | 36.8798 | -104.8456 | 6.0214 | 2.6267 | 1.85 |
| 2010 | 1 | 12 | 36.8708 | -104.8299 | 10.0587 | 4.4109 | -999 |
| 2010 | 1 | 12 | 36.893 | -104.8735 | 7.405 | 2.5762 | 1.61 |
| 2010 | 1 | 12 | 36.8915 | -104.8734 | 6.3142 | 2.6534 | 2.17 |
| 2010 | 1 | 12 | 37.3044 | -104.9416 | 6.0944 | 2.797 | 1.23 |
| 2010 | 1 | 13 | 36.8988 | -104.8782 | 5.3853 | 2.2395 | 1.51 |
| 2010 | 1 | 13 | 36.8967 | -104.9046 | 5.4969 | 3.2713 | -999 |
| 2010 | 1 | 13 | 36.8969 | -104.8747 | 5.3224 | 2.7474 | 1.17 |
| 2010 | 1 | 14 | 36.8946 | -104.875 | 5.338 | 2.2408 | 1.49 |
| 2010 | 1 | 14 | 36.8075 | -104.8601 | 5.3312 | 2.7929 | 1.5 |
| 2010 | 1 | 14 | 36.8946 | -104.874 | 6.8103 | 2.8101 | -999 |
| 2010 | 1 | 15 | 36.8826 | -104.8377 | 7.8304 | 1.4056 | 2.48 |
| 2010 | 1 | 15 | 36.8811 | -104.8367 | 6.6137 | 2.0286 | 1.92 |
| 2010 | 1 | 15 | 36.884 | -104.8393 | 6.8003 | 3.3256 | -999 |
| 2010 | 1 | 15 | 36.8888 | -104.8376 | 9.4541 | 2.6024 | 1.43 |
| 2010 | 1 | 16 | 37.3388 | -104.9037 | 10.2949 | 4.5435 | -999 |
| 2010 | 1 | 16 | 37.3137 | -104.9368 | 8.0004 | 1.2466 | 2.19 |
| 2010 | 1 | 16 | 36.8093 | -104.8657 | 8.4417 | 4.1575 | -999 |
| 2010 | 1 | 17 | 36.8052 | -104.861 | 6.7424 | 3.809 | -999 |
| 2010 | 1 | 17 | 36.8095 | -104.8664 | 12.2534 | 2.854 | 0.84 |
| 2010 | 1 | 17 | 37.3081 | -104.928 | 6.2933 | 3.8422 | -999 |
| 2010 | 1 | 17 | 36.8073 | -104.8618 | 8.2918 | 3.7206 | 1.39 |
| 2010 | 1 | 17 | 36.7955 | -104.8538 | 7.0348 | 4.6932 | 1.22 |
| 2010 | 1 | 17 | 36.8081 | -104.865 | 7.826 | 2.1071 | 1.92 |
| 2010 | 1 | 17 | 36.893 | -104.8713 | 4.8197 | 2.174 | -999 |
| 2010 | 1 | 17 | 36.9146 | -104.8007 | 19.5946 | 3.9812 | 1.32 |
| 2010 | 1 | 18 | 36.8511 | -104.8681 | 5.3719 | 1.4839 | 4.27 |
| 2010 | 1 | 18 | 36.8497 | -104.8706 | 4.7713 | 2.574 | -999 |
| 2010 | 1 | 18 | 36.8571 | -104.873 | 6.9852 | 3.733 | 1.17 |
| 2010 | 1 | 18 | 36.8526 | -104.8723 | 8.2024 | 4.3162 | 1.31 |
| 2010 | 1 | 18 | 36.9134 | -104.8473 | 6.8351 | 3.6485 | 1.11 |
| 2010 | 1 | 18 | 36.8514 | -104.8684 | 5.1612 | 2.3621 | 1.14 |
| 2010 | 1 | 18 | 36.8513 | -104.8697 | 5.5197 | 2.358 | 1.51 |

| | | | | | | | |
|------|---|----|---------|-----------|---------|--------|------|
| 2010 | 1 | 18 | 36.8511 | -104.8699 | 7.4214 | 2.8955 | 1.42 |
| 2010 | 1 | 18 | 36.8555 | -104.8683 | 7.1511 | 3.9811 | 1.12 |
| 2010 | 1 | 18 | 36.867 | -104.8707 | 5.8949 | 2.2351 | 1.57 |
| 2010 | 1 | 18 | 36.8492 | -104.8703 | 4.9923 | 2.4403 | -999 |
| 2010 | 1 | 18 | 36.8555 | -104.8711 | 5.9072 | 2.5317 | -999 |
| 2010 | 1 | 18 | 36.8501 | -104.8709 | 4.8017 | 1.7594 | 1.8 |
| 2010 | 1 | 18 | 36.8508 | -104.8669 | 6.5224 | 3.6002 | -999 |
| 2010 | 1 | 18 | 36.8632 | -104.8609 | 6.9504 | 3.7853 | -999 |
| 2010 | 1 | 18 | 36.8702 | -104.871 | 6.5707 | 3.1035 | 1.19 |
| 2010 | 1 | 18 | 36.8546 | -104.8699 | 8.6215 | 1.9123 | 1.64 |
| 2010 | 1 | 18 | 36.8514 | -104.8727 | 6.1159 | 2.5827 | 1.74 |
| 2010 | 1 | 18 | 36.8669 | -104.87 | 7.0782 | 2.9531 | -999 |
| 2010 | 1 | 18 | 36.8528 | -104.8704 | 5.2295 | 2.8275 | -999 |
| 2010 | 1 | 18 | 36.8545 | -104.8699 | 9.0337 | 0.8144 | 2.26 |
| 2010 | 1 | 18 | 36.8535 | -104.8575 | 10.7543 | 6.0061 | -999 |
| 2010 | 1 | 18 | 36.7943 | -104.857 | 12.0322 | 3.9624 | 1.56 |
| 2010 | 1 | 18 | 36.8843 | -104.8758 | 5.2932 | 2.0716 | 1.81 |
| 2010 | 1 | 18 | 36.8858 | -104.8746 | 5.1493 | 2.6218 | 1.5 |
| 2010 | 1 | 18 | 36.8518 | -104.8682 | 3.6689 | 1.709 | 2.1 |
| 2010 | 1 | 19 | 36.8533 | -104.8692 | 4.503 | 2.3192 | 1.02 |
| 2010 | 1 | 19 | 36.854 | -104.8702 | 4.8103 | 2.4961 | 1.32 |
| 2010 | 1 | 19 | 36.8083 | -104.8646 | 4.3076 | 2.2227 | 1.98 |
| 2010 | 1 | 19 | 36.865 | -104.8653 | 10.6591 | 4.6497 | -999 |
| 2010 | 1 | 19 | 36.8795 | -104.8445 | 8.2085 | 2.7247 | 1.69 |
| 2010 | 1 | 20 | 36.8104 | -104.8627 | 8.4387 | 3.2865 | 1.09 |
| 2010 | 1 | 20 | 36.8978 | -104.8771 | 6.3861 | 3.197 | 1.28 |
| 2010 | 1 | 20 | 36.8935 | -104.8752 | 5.2868 | 2.8765 | 1.59 |
| 2010 | 1 | 20 | 36.8925 | -104.8745 | 4.6903 | 2.3267 | 1.5 |
| 2010 | 1 | 20 | 36.894 | -104.8769 | 5.5795 | 2.5045 | 1.69 |
| 2010 | 1 | 20 | 36.8925 | -104.8754 | 5.9141 | 2.5671 | 1.77 |
| 2010 | 1 | 20 | 36.8933 | -104.8803 | 4.913 | 2.7262 | 0.89 |
| 2010 | 1 | 20 | 36.9973 | -104.9753 | 8.3889 | 4.1238 | 1.27 |
| 2010 | 1 | 20 | 36.8922 | -104.8764 | 6.9728 | 3.5552 | 1.52 |
| 2010 | 1 | 21 | 36.8931 | -104.8765 | 4.047 | 1.8212 | 1.8 |
| 2010 | 1 | 21 | 36.8786 | -104.8717 | 9.034 | 1.2981 | 3.28 |
| 2010 | 1 | 21 | 36.8781 | -104.8716 | 4.0175 | 1.9835 | 2.55 |
| 2010 | 1 | 21 | 36.8749 | -104.8677 | 4.1366 | 2.0609 | 1.19 |
| 2010 | 1 | 21 | 36.8774 | -104.8709 | 4.5776 | 2.2695 | 1.49 |
| 2010 | 1 | 21 | 36.8746 | -104.8738 | 7.5894 | 2.4187 | 2.38 |
| 2010 | 1 | 21 | 36.8759 | -104.8712 | 6.7006 | 2.1485 | 2.02 |

| | | | | | | | |
|------|---|----|---------|-----------|---------|--------|------|
| 2010 | 1 | 21 | 36.8796 | -104.8699 | 4.0396 | 1.9849 | 1.42 |
| 2010 | 1 | 21 | 36.8812 | -104.8714 | 4.5845 | 2.4526 | 1.72 |
| 2010 | 1 | 21 | 36.8795 | -104.8741 | 5.1002 | 2.3481 | 1.83 |
| 2010 | 1 | 21 | 36.8496 | -104.8694 | 5.8418 | 2.823 | 1.21 |
| 2010 | 1 | 22 | 36.877 | -104.8702 | 6.2714 | 2.7829 | 1.73 |
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| 2010 | 2 | 27 | 37.0025 | -104.9721 | 8.1147 | 1.9938 | 2.11 |
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| 2008 | 5 | 31 | 36.9138 | -104.871 | 5.8875 | 3.3424 | 1.54 |
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| 2008 | 6 | 5 | 36.8745 | -104.8709 | 8.481 | 2.1112 | 2.51 |
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| 2008 | 6 | 19 | 36.8796 | -104.8714 | 5.0734 | 2.7914 | 1.09 |
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| 2008 | 7 | 6 | 36.9486 | -104.8667 | 9.0954 | 4.4639 | -999 |
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| 2008 | 7 | 17 | 36.8915 | -104.8691 | 6.9388 | 3.6087 | 1.51 |
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| 2008 | 9 | 6 | 36.895 | -104.8744 | 5.1167 | 2.0015 | 1.04 |
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| 2008 | 9 | 9 | 36.8746 | -104.8712 | 3.9832 | 2.0544 | 0.94 |
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| 2008 | 9 | 16 | 37.0258 | -104.9721 | 2.244 | 1.2201 | 1.68 |
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| 2008 | 9 | 18 | 37.3814 | -104.6115 | 4.669 | 2.239 | 1.92 |
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| 2008 | 9 | 21 | 36.9162 | -104.8801 | 5.2138 | 3.1383 | 1.14 |
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| 2008 | 9 | 21 | 37.3044 | -104.7791 | 9.8299 | 0.7629 | 1.3 |
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| 2008 | 10 | 1 | 36.8909 | -104.8766 | 5.2405 | 2.2764 | 0.8 |
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| 2008 | 10 | 1 | 37.2959 | -104.7687 | 9.5487 | 1.6652 | 0.97 |
| 2008 | 10 | 2 | 37.2959 | -104.7708 | 8.516 | 1.1415 | 0.74 |
| 2008 | 10 | 3 | 36.892 | -104.8771 | 4.9104 | 2.1108 | 1.07 |
| 2008 | 10 | 3 | 36.8929 | -104.8778 | 4.6579 | 2.1823 | 1.06 |
| 2008 | 10 | 3 | 36.891 | -104.8767 | 10.3977 | 4.641 | -999 |
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| 2008 | 10 | 3 | 36.8946 | -104.8792 | 4.6901 | 2.0757 | 1.64 |
| 2008 | 10 | 4 | 37.3831 | -104.6079 | 4.6267 | 2.1217 | 0.79 |
| 2008 | 10 | 4 | 36.8618 | -104.8552 | 6.5059 | 2.8563 | 1.26 |
| 2008 | 10 | 4 | 37.2971 | -104.7715 | 9.3793 | 0.5314 | 3.57 |
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| 2008 | 8 | 27 | 36.8503 | -104.8704 | 7.2674 | 2.7267 | 0.47 |
| 2008 | 10 | 5 | 36.8764 | -104.8445 | 5.9186 | 2.425 | 1.39 |
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| 2008 | 8 | 27 | 36.8838 | -104.8695 | 4.2253 | 2.1112 | 0.71 |
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| 2008 | 8 | 27 | 36.8921 | -104.8742 | 6.5014 | 3.2811 | 1.05 |
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| 2008 | 10 | 20 | 37.0204 | -104.9776 | 4.3524 | 0.912 | 1.46 |
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| 2008 | 10 | 20 | 37.021 | -104.9782 | 4.2007 | 0.857 | 1.78 |
| 2008 | 10 | 20 | 37.0234 | -104.9794 | 4.1285 | 0.8414 | 1.05 |
| 2008 | 10 | 20 | 36.8906 | -104.8774 | 4.8255 | 2.3356 | 1.14 |
| 2008 | 10 | 20 | 37.2981 | -104.7744 | 5.8494 | 2.3841 | 1.44 |
| 2008 | 10 | 20 | 37.0179 | -104.9762 | 4.4849 | 1.4581 | 1.48 |
| 2008 | 10 | 20 | 36.8856 | -104.875 | 5.0424 | 2.7284 | 0.7 |
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| 2008 | 10 | 21 | 36.8579 | -104.8471 | 7.2203 | 3.0364 | 0.6 |
| 2008 | 10 | 21 | 36.8582 | -104.8434 | 5.4145 | 3.0202 | 0.91 |
| 2008 | 10 | 21 | 36.8553 | -104.8472 | 7.178 | 1.2687 | 2.3 |
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| 2008 | 10 | 21 | 37.2953 | -104.7737 | 7.6089 | 0.7453 | 1.53 |
| 2008 | 10 | 22 | 36.9253 | -104.802 | 2.5308 | 2.3933 | 1.26 |
| 2008 | 10 | 22 | 36.8569 | -104.8469 | 8.6234 | 0.9504 | 2.53 |
| 2008 | 10 | 22 | 36.9254 | -104.8014 | 4.9972 | 3.2781 | 1.58 |
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| 2008 | 10 | 27 | 36.8916 | -104.8754 | 4.7093 | 1.9241 | 1.06 |
| 2008 | 10 | 27 | 36.8902 | -104.873 | 4.3493 | 2.1243 | 0.3 |
| 2008 | 10 | 28 | 36.8643 | -104.8667 | 6.0997 | 3.2177 | -999 |
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| 2008 | 10 | 29 | 36.8786 | -104.8407 | 7.2377 | 3.0874 | -999 |
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| 2008 | 10 | 29 | 36.8783 | -104.8415 | 6.4977 | 2.4534 | 1.39 |
| 2008 | 8 | 30 | 37.0261 | -104.9737 | 2.2119 | 1.4535 | 0.62 |
| 2008 | 10 | 29 | 36.896 | -104.8758 | 6.5195 | 2.1433 | 1.89 |
| 2008 | 10 | 30 | 36.894 | -104.8727 | 6.1815 | 1.9035 | 1.14 |
| 2008 | 10 | 31 | 36.8842 | -104.8708 | 6.4721 | 3.3781 | 1.04 |
| 2008 | 10 | 31 | 36.8837 | -104.8722 | 4.2978 | 1.8517 | 1.37 |
| 2008 | 10 | 31 | 36.882 | -104.8722 | 3.5806 | 1.8859 | 1.51 |
| 2008 | 10 | 31 | 37.1574 | -104.9006 | 6.1188 | 2.9072 | 0.51 |
| 2008 | 10 | 31 | 36.8919 | -104.8725 | 7.0135 | 3.4866 | -999 |
| 2008 | 10 | 31 | 36.8944 | -104.8704 | 6.1442 | 3.7678 | 1.07 |
| 2008 | 10 | 31 | 36.8971 | -104.8773 | 5.2607 | 2.6187 | 0.81 |
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| 2008 | 8 | 30 | 36.8802 | -104.8461 | 8.0761 | 3.5351 | 0.69 |
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| 2008 | 11 | 1 | 37.6722 | -104.7136 | 8.3123 | 2.9612 | 0.57 |
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| 2008 | 11 | 5 | 37.153 | -104.9017 | 9.0757 | 0.5693 | 1.11 |
| 2008 | 11 | 5 | 37.2958 | -104.7738 | 8.0588 | 0.7975 | 2.28 |
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| 2008 | 11 | 7 | 37.1561 | -104.9016 | 5.1548 | 2.4686 | 0.56 |
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| 2009 | 3 | 4 | 36.8752 | -104.8717 | 4.2309 | 2.113 | 1.54 |
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| 2009 | 5 | 19 | 36.8785 | -104.8702 | 5.0903 | 3.0568 | 0.82 |
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| 2009 | 5 | 26 | 37.6891 | -104.6493 | 2.9387 | 1.6691 | 0.96 |
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| 2009 | 7 | 10 | 36.8231 | -104.8514 | 7.279 | 3.0317 | 1.53 |
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| 2009 | 11 | 19 | 36.8876 | -104.8786 | 4.0246 | 2.3958 | 1.43 |
| 2009 | 11 | 19 | 36.8881 | -104.8764 | 4.8981 | 2.419 | 1.28 |
| 2009 | 11 | 19 | 36.8943 | -104.8814 | 4.9464 | 2.8615 | 0.87 |
| 2009 | 11 | 19 | 36.8031 | -105.0209 | 9.44 | 1.1271 | 3.22 |
| 2009 | 11 | 19 | 36.8873 | -104.88 | 5.7376 | 2.8478 | 1.36 |
| 2009 | 11 | 19 | 36.8003 | -105.0244 | 7.9585 | 2.9252 | 1.91 |
| 2009 | 11 | 20 | 36.8559 | -104.8685 | 7.5463 | 3.1223 | -999 |
| 2009 | 11 | 20 | 36.9607 | -104.8669 | 9.9002 | 3.3884 | 1.2 |
| 2009 | 11 | 20 | 36.979 | -104.9767 | 4.8594 | 2.6288 | 1.11 |
| 2009 | 11 | 20 | 36.8081 | -105.0225 | 6.8024 | 3.014 | 2.33 |
| 2009 | 11 | 20 | 36.8073 | -105.0227 | 10.0466 | 0.9485 | 3.85 |
| 2009 | 11 | 21 | 36.8444 | -104.8633 | 9.216 | 3.0886 | 1.39 |
| 2009 | 11 | 21 | 36.8883 | -104.8737 | 5.4184 | 2.5733 | 1.18 |
| 2009 | 11 | 21 | 37.2247 | -104.6637 | 4.7635 | 1.591 | 1.67 |
| 2009 | 11 | 21 | 37.2272 | -104.6421 | 7.7416 | 4.0227 | 1.22 |
| 2009 | 11 | 21 | 37.2258 | -104.6614 | 4.4497 | 1.7532 | 1.65 |
| 2009 | 11 | 21 | 37.2221 | -104.6659 | 4.6666 | 2.0201 | 1.47 |
| 2009 | 11 | 21 | 37.2316 | -104.6432 | 7.3373 | 3.7059 | 1.13 |
| 2009 | 11 | 22 | 37.224 | -104.6593 | 4.9489 | 1.8896 | 1.54 |
| 2009 | 11 | 22 | 37.2229 | -104.6647 | 9.5077 | 3.8195 | 2.22 |
| 2009 | 11 | 22 | 37.2254 | -104.663 | 4.4701 | 1.7479 | 2 |

| | | | | | | | |
|------|----|----|---------|-----------|---------|--------|------|
| 2009 | 11 | 22 | 37.2254 | -104.6611 | 4.396 | 1.922 | 1.6 |
| 2009 | 11 | 22 | 36.9774 | -104.9882 | 6.8964 | 5.695 | 1.65 |
| 2009 | 11 | 22 | 37.2238 | -104.6668 | 4.3428 | 2.0839 | 0.81 |
| 2009 | 11 | 22 | 37.2244 | -104.6632 | 3.8546 | 1.5625 | 1.72 |
| 2009 | 11 | 23 | 37.2256 | -104.6552 | 5.9047 | 3.1827 | 0.9 |
| 2009 | 11 | 23 | 37.2338 | -104.3723 | 21.7956 | 0.7343 | 0.29 |
| 2009 | 11 | 24 | 36.7982 | -105.0309 | 9.6096 | 3.6016 | 1.41 |
| 2009 | 11 | 24 | 36.8832 | -104.837 | 8.486 | 3.7705 | 0.82 |
| 2009 | 11 | 25 | 36.8888 | -104.8441 | 7.028 | 3.0108 | 1.23 |
| 2009 | 11 | 25 | 37.0405 | -104.9708 | 5.1338 | 2.42 | 1.57 |
| 2009 | 11 | 25 | 37.211 | -104.4552 | 4.7063 | 1.0525 | 1.01 |
| 2009 | 11 | 25 | 36.8056 | -105.0328 | 9.3038 | 5.1021 | 1.34 |
| 2009 | 11 | 25 | 36.8022 | -105.0322 | 18.1148 | 4.0544 | 1.15 |
| 2009 | 11 | 26 | 37.2107 | -104.4568 | 4.9324 | 0.5437 | 1.64 |
| 2009 | 11 | 27 | 36.8823 | -104.8705 | 8.794 | 3.2356 | -999 |
| 2009 | 11 | 28 | 37.3117 | -104.9291 | 5.952 | 3.3905 | 0.67 |
| 2009 | 11 | 28 | 36.8067 | -105.016 | 8.7106 | 2.6005 | 1.69 |
| 2009 | 11 | 28 | 36.8825 | -104.8453 | 7.7275 | 3.381 | 1.3 |
| 2009 | 11 | 28 | 36.8775 | -104.8436 | 6.5125 | 1.9998 | 2.39 |
| 2009 | 11 | 29 | 36.8786 | -104.8428 | 5.2529 | 2.637 | -999 |
| 2009 | 11 | 29 | 36.877 | -104.8392 | 7.8791 | 3.7277 | -999 |
| 2009 | 11 | 29 | 36.8778 | -104.8433 | 6.5257 | 1.9601 | 2.19 |
| 2009 | 11 | 29 | 36.8808 | -104.8441 | 6.8033 | 3.3251 | 1.29 |
| 2009 | 11 | 29 | 36.8763 | -104.8376 | 7.5261 | 3.8149 | -999 |
| 2009 | 11 | 29 | 36.8795 | -104.8443 | 7.0976 | 2.653 | 1.56 |
| 2009 | 11 | 29 | 36.8764 | -104.8678 | 4.2434 | 2.0157 | 1.33 |
| 2009 | 11 | 29 | 36.8762 | -104.8759 | 5.9771 | 3.4802 | 1.45 |
| 2009 | 11 | 29 | 36.8832 | -104.8713 | 4.6303 | 2.3469 | 1.61 |
| 2009 | 11 | 29 | 36.8795 | -104.8651 | 5.0967 | 2.7738 | -999 |
| 2009 | 11 | 29 | 36.878 | -104.8663 | 4.5557 | 2.3003 | 1.8 |
| 2009 | 11 | 29 | 36.8813 | -104.8663 | 5.3405 | 3.0422 | 1.28 |
| 2009 | 11 | 29 | 36.8783 | -104.8684 | 5.2034 | 2.3639 | 1.78 |
| 2009 | 11 | 29 | 37.0397 | -105.0143 | 7.0278 | 3.2047 | 1.3 |
| 2009 | 11 | 30 | 37.2977 | -104.6351 | 5.3293 | 1.8713 | 2.07 |
| 2009 | 11 | 30 | 37.0109 | -104.9722 | 7.8527 | 3.4711 | 1.54 |
| 2009 | 11 | 30 | 36.878 | -104.8448 | 7.2816 | 2.3635 | 1.2 |
| 2009 | 11 | 30 | 36.8776 | -104.8415 | 8.45 | 2.2102 | 1.49 |
| 2009 | 11 | 17 | 36.8813 | -104.8404 | 7.778 | 2.9112 | -999 |
| 2009 | 12 | 1 | 36.8169 | -104.858 | 8.712 | 3.6387 | 1.25 |
| 2009 | 12 | 1 | 36.8851 | -104.8585 | 6.7941 | 3.5131 | 1.91 |

| | | | | | | | |
|------|----|----|---------|-----------|---------|--------|------|
| 2009 | 12 | 1 | 36.9499 | -104.8587 | 5.6831 | 2.6441 | 1.49 |
| 2009 | 12 | 2 | 37.3066 | -104.9345 | 5.6987 | 2.976 | -999 |
| 2009 | 12 | 3 | 36.8658 | -104.8704 | 4.9732 | 2.2855 | 1.7 |
| 2009 | 12 | 3 | 36.8773 | -104.8684 | 4.0458 | 2.0112 | 1.93 |
| 2009 | 12 | 3 | 36.8783 | -104.8445 | 7.4721 | 2.9291 | 1.27 |
| 2009 | 12 | 4 | 36.8783 | -104.842 | 7.3568 | 2.873 | 1.58 |
| 2009 | 12 | 4 | 36.9328 | -104.8698 | 8.9336 | 2.7316 | 1.99 |
| 2009 | 12 | 5 | 36.8791 | -104.8447 | 5.6499 | 2.5429 | 1.39 |
| 2009 | 12 | 5 | 36.9989 | -104.9714 | 4.883 | 2.4115 | 1.04 |
| 2009 | 12 | 6 | 36.8758 | -104.8386 | 7.6593 | 3.1487 | -999 |
| 2009 | 12 | 6 | 36.954 | -104.8666 | 5.2658 | 2.7281 | -999 |
| 2009 | 12 | 8 | 36.8826 | -104.8396 | 7.2432 | 2.8674 | 1.4 |
| 2009 | 12 | 9 | 36.8919 | -104.8726 | 4.7897 | 2.6685 | 1.3 |
| 2009 | 12 | 10 | 37.3036 | -104.9363 | 7.0233 | 3.6565 | -999 |
| 2009 | 12 | 10 | 36.8848 | -104.8709 | 5.8245 | 2.9009 | -999 |
| 2009 | 12 | 10 | 36.8717 | -104.8708 | 4.7806 | 1.905 | 2.08 |
| 2009 | 12 | 10 | 36.8698 | -104.8684 | 5.4287 | 2.6428 | -999 |
| 2009 | 12 | 11 | 36.9548 | -104.8513 | 10.8217 | 5.1106 | -999 |
| 2009 | 12 | 11 | 36.8704 | -104.8732 | 7.8726 | 1.7854 | 1.5 |
| 2009 | 12 | 11 | 37.3745 | -104.6079 | 4.4487 | 2.244 | 0.97 |
| 2009 | 12 | 11 | 36.8701 | -104.8708 | 9.2064 | 0.739 | 3.56 |
| 2009 | 12 | 11 | 36.8741 | -104.8697 | 8.3181 | 3.6535 | 2.14 |
| 2009 | 12 | 11 | 36.8696 | -104.8666 | 6.2303 | 2.8792 | 1.46 |
| 2009 | 12 | 11 | 36.8683 | -104.8703 | 4.4197 | 2.0468 | 1.3 |
| 2009 | 12 | 12 | 36.8671 | -104.8676 | 4.9289 | 2.3642 | -999 |
| 2009 | 12 | 13 | 36.8911 | -104.8759 | 4.4733 | 2.2264 | 1.51 |
| 2009 | 12 | 23 | 37.0252 | -105.0836 | 7.8186 | 5.8515 | -999 |
| 2010 | 1 | 5 | 37.7946 | -105.4893 | 14.2281 | 1.2318 | 1.57 |
| 2009 | 5 | 19 | 37.7307 | -105.3773 | 6.9134 | 1.3928 | 1.31 |
| 2009 | 9 | 1 | 37.5658 | -105.2973 | 10.0737 | 1.4395 | 1.2 |

Table 8. Barnhart's relocated earthquakes for the 2011 earthquake swarm [Barnhart *et al.*, 2014]

| Date ^b | OT1 | Lon1 | Lat1 | OT2 | Lon2 | Lat2 | Z (km) |
|------------------------|---------|-----------|---------|---------|------------|----------|----------------|
| 2011/5/9 | 28:52.8 | -104.783 | 37.021 | 28:51.2 | -104.76534 | 37.06877 | 3 ^d |
| 2011/5/11 | 06:15.3 | -104.665 | 37.1 | 06:13.2 | -104.76775 | 37.05136 | 3 ^d |
| 2011/8/22 | 30:18.1 | -104.8114 | 36.9751 | 30:17.1 | -104.77892 | 37.05834 | 3 ^d |
| 2011/8/23 | 48:50.3 | -104.7238 | 36.9226 | 48:48.7 | -104.76071 | 37.01174 | 3 ^d |
| 2011/8/23 ^c | 46:17.8 | -104.8722 | 37.0394 | 46:15.4 | -104.75298 | 37.03811 | 3 ^d |
| 2011/8/23 | 12:18.3 | -104.715 | 37.1222 | 12:58.6 | -104.71761 | 37.12538 | 3 ^d |
| 2011/8/23 | 56:57.6 | -104.6608 | 37.0849 | 56:55.9 | -104.73306 | 37.09641 | 3 ^d |
| 2011/8/23 | 01:33.5 | -104.7282 | 37.1493 | 01:31.9 | -104.73334 | 37.11427 | 3 ^d |
| 2011/8/23 | 17:59.1 | -104.7098 | 37.1266 | 17:56.2 | -104.73158 | 37.10547 | 3 ^d |
| 2011/8/23 | 34:39.7 | -104.7693 | 37.1155 | 34:37.3 | -104.73845 | 37.09292 | 3 ^d |
| 2011/8/23 | 32:17.1 | -104.6852 | 37.1633 | 32:15.7 | -104.72594 | 37.14349 | 3 ^d |
| 2011/8/23 | 37:57.1 | -104.7984 | 37.0082 | 37:54.4 | -104.77573 | 37.00424 | 3 ^d |
| 2011/8/23 | 29:27.9 | -104.8184 | 36.9205 | 29:29.6 | -104.74202 | 37.08241 | 3 ^d |
| 2011/8/23 | 03:50.4 | -104.7569 | 37.1376 | 03:17.9 | -104.76328 | 37.03284 | 3 ^d |
| 2011/8/23 | 11:12.9 | -104.8015 | 36.9967 | 11:10.6 | -104.76907 | 37.03746 | 3 ^d |
| 2011/8/24 | 40:43.2 | -104.7672 | 37.0886 | 40:42.5 | -104.76219 | 37.05419 | 5.95 |
| 2011/8/24 | 07:49.4 | -104.7354 | 37.1224 | 07:48.8 | -104.7338 | 37.08096 | 5.31 |
| 2011/8/24 | 15:52.6 | -104.7192 | 37.1589 | 15:51.7 | -104.73122 | 37.0794 | 4.42 |
| 2011/8/24 | 39:48.4 | -104.7278 | 37.122 | 39:47.1 | -104.72931 | 37.09176 | 5.72 |
| 2011/8/25 | 44:35.8 | -104.7601 | 37.0966 | 44:35.6 | -104.74323 | 37.08858 | 4.94 |
| 2011/8/25 | 34:33.8 | -104.7764 | 37.0235 | 34:34.1 | -104.74538 | 37.05739 | 5.65 |

^aOT1, Lon1, and Lat1 indicate origin time, longitude, and latitude from the NEIC/Hydra location. OT2, Lon2, and Lat2 indicate origin time, longitude, and latitude of the relocated event.

^bDates are formatted as yy/mm/dd.

^cThe 2011 Trinidad earthquake main shock.

^dFixed values of depth. Relative position between locations for each event are shown in Figure 3b. All depths are relative to mean sea level. The surface elevation in the Raton Basin near the 2011 Trinidad epicenter is ~2.1 km above sea level.

Table 9. Magnitude 3.8 > Earthquakes in Raton Basin from 1973 – 2011 [Rubinstein *et al.*, 2014]

| Date (yyyy/mm/dd) | Time (hh:mm:ss.ss) | Latitude ($^{\circ}$)* | Longitude ($^{\circ}$) | Depth (km) | Magnitude [†] |
|--------------------------|--------------------|--------------------------|--------------------------|------------|------------------------|
| 1973/09/23 | 03:58:54.5 | 37.178 | -104.61 | 1.05 | 4.2 |
| 2001/09/04 [‡] | 12:45:52.1 | 37.162 | -104.65 | 1.45 | 4.0 |
| 2001/09/05 [‡] | 10:52:07.5 | 37.152 | -104.66 | 2.65 | 4.5 |
| 2003/09/13 | 15:22:47.9 | 37.139 | -104.88 | 2.99 | 3.8 |
| 2004/03/22 [§] | 12:09:54.9 | 36.852 | -104.96 | 5.43 | 4.4 |
| 2004/08/01 [§] | 06:50:44.8 | 36.865 | -105.01 | 2.41 | 4.3 |
| 2005/08/10 | 22:08:14.6 | 36.963 | -104.83 | 2.77 | 4.1 |
| 2005/08/10 | 22:08:20.9 | 36.946 | -104.84 | 5.95 | 5.0 |
| 2007/01/03 | 14:34:37.17 | 37.009 | -104.911 | 5.17 | 4.4 |
| 2009/07/29 | 10:00:35.7 | 36.843 | -104.83 | 13.33 | 4.1 |
| 2010/01/18 | 08:41:06.6 | 36.886 | -104.83 | 10.04 | 3.8 |
| 2011/05/11 | 19:06:12.6 | 37.123 | -104.69 | 1.1 | 3.8 |
| 2011/08/22 [#] | 23:30:18.1 | 37.019 | -104.77 | 3.5 | 4.7 |
| 2011/08/23 [#] | 05:46:17.8 | 37.054 | -104.76 | 3.5 | 5.3 |
| 2011/08/23 [#] | 14:11:13.0 | 37.048 | -104.76 | 3.5 | 4.0 |
| 2011/09/13 [#] | 05:24:37.26 | 36.880 | -104.869 | 3.58 | 4.0 |
| 2011/09/16 | 14:51:50.0 | 36.880 | -104.88 | 2.25 | 3.9 |

* 2σ location uncertainties are ± 15 km for all events except the earthquakes on 22 and 23 August 2011.

[†]Magnitude is the authoritative magnitude in ComCat. Moment magnitudes (**M**) are indicated with the magnitude printed in boldface type. The other magnitudes are a combination of amplitude based measures: M_b , M_{blg} , and M_L .

[‡]Cluster of $M \geq 4$ earthquakes in 2001. These earthquakes are located close in space and time, such that they cannot be considered to be independent in space and time. As such, they are counted as an individual earthquake sequence in the statistical calculations.

[§]Cluster of $M \geq 4$ earthquakes in 2004.

^{||}Cluster of $M \geq 4$ earthquakes in 2005.

[#]Cluster of $M \geq 4$ earthquakes in 2011.

Table 10. Earthquakes located with Meremonte's array from September 7 – October 15th [Meremonte et al., 2002].

| Earthquakes divided into 5 time periods: | | DATE YR MO DY | TIME HR MN SEC | LATITUDE DEG MIN | LONGITUDE DEG MIN | DEPTH KM | PREF. HR.G | RMS SEC. | GAP DEG | PHASES P S | MIN Q |
|--|--|------------------|-------------------|---------------------|----------------------|-------------|---------------|-------------|------------|---------------|-------|
| Stations PT0, PD0, TK0, BHO, DK0, CKD, TKD installed | | 1 9 10 | 18 56 1.12 | 37 7.70N | 104 42.01W | 4.7 | 3.4mb | 0.05 | 117 | 7 7 | 7 B |
| 1 9 12 3 52 0.63 | | 37 8.75N | 104 41.15W | 3.8 | 0.0 | 0.03 | 135 | 5 5 | 9 D | | |
| 1 9 13 6 46 12.98 | | 37 8.75N | 104 41.14W | 3.7 | 0.0 | 0.04 | 136 | 7 7 | 6 C | | |
| 1 9 13 11 22 14.97 | | 37 8.70N | 104 41.16W | 2.7 | 2.8mb | 0.03 | 135 | 7 7 | 6 A | | |
| 1 9 13 11 22 29.48 | | 37 8.77N | 104 41.12W | 3.3 | 0.0 | 0.03 | 136 | 7 7 | 6 D | | |
| 1 9 13 11 38 5.63 | | 37 8.48N | 104 41.79W | 1.7 | 0.0 | 0.03 | 131 | 5 5 | 9 A | | |
| 1 9 13 16 39 5.48 | | 37 8.71N | 104 41.53W | 4.4 | 3.0mb | 0.05 | 135 | 7 7 | 6 B | | |
| 1 9 14 19 19 2.50 | | 37 7.76N | 104 42.36W | 1.1 | 0.0 | 0.06 | 118 | 5 5 | 9 A | | |
| 1 9 14 22 22 54.54 | | 37 7.85N | 104 41.76W | 6.0 | 0.0 | 0.03 | 91 | 6 6 | 6 B | | |
| 1 9 17 12 33 47.79 | | 37 9.14N | 104 40.87W | 4.4 | 0.0 | 0.03 | 93 | 6 6 | 4 B | | |
| 1 9 19 6 58 52.87 | | 37 8.59N | 104 41.23W | 4.4 | 0.0 | 0.02 | 133 | 5 5 | 6 C | | |
| 1 9 19 8 16 52.83 | | 37 8.92N | 104 42.94W | 4.1 | 0.0 | 0.02 | 154 | 5 5 | 4 B | | |
| 1 9 19 20 19 9.41 | | 37 9.04N | 104 40.93W | 3.1 | 0.0 | 0.03 | 141 | 6 6 | 7 D | | |
| 1 9 19 20 21 52.53 | | 37 7.95N | 104 41.52W | 6.0 | 0.0 | 0.04 | 122 | 6 6 | 5 A | | |
| 1 9 20 2 47 6.96 | | 37 7.89N | 104 41.60W | 5.9 | 0.0 | 0.04 | 121 | 6 6 | 5 A | | |
| 1 9 21 3 17 5.12 | | 37 7.53N | 104 42.04W | 5.6 | 0.0 | 0.04 | 115 | 6 6 | 4 A | | |
| 1 9 21 19 10 59.76 | | 37 7.25N | 104 42.36W | 5.8 | 3.4mb | 0.04 | 110 | 8 8 | 4 A | | |
| 1 9 21 19 14 46.87 | | 37 7.13N | 104 42.49W | 5.0 | 0.0 | 0.03 | 161 | 5 5 | 4 A | | |
| 1 9 21 19 39 13.14 | | 37 6.98N | 104 42.54W | 4.9 | 2.8mb | 0.05 | 106 | 7 7 | 4 A | | |
| 1 9 21 19 49 23.67 | | 37 8.31N | 104 41.57W | 4.6 | 0.0 | 0.02 | 128 | 5 5 | 5 B | | |
| 1 9 21 20 9 23.40 | | 37 7.08N | 104 42.38W | 6.1 | 0.0 | 0.06 | 108 | 6 6 | 4 A | | |
| 1 9 21 20 36 1.86 | | 37 7.21N | 104 42.44W | 5.0 | 0.0 | 0.04 | 109 | 6 6 | 4 A | | |
| 1 9 22 2 11 12.53 | | 37 7.18N | 104 42.31W | 5.5 | 0.0 | 0.05 | 109 | 6 6 | 4 A | | |
| 1 9 22 7 5 43.85 | | 37 7.06N | 104 42.28W | 6.0 | 0.0 | 0.04 | 108 | 6 6 | 4 B | | |
| 1 9 23 5 2 12.75 | | 37 7.51N | 104 42.74W | 3.8 | 0.0 | 0.04 | 113 | 6 6 | 5 B | | |
| 1 9 23 7 38 49.92 | | 37 8.07N | 104 41.97W | 5.1 | 0.0 | 0.05 | 124 | 5 5 | 5 B | | |
| 1 9 25 16 22 35.60 | | 37 7.37N | 104 42.26W | 4.7 | 0.0 | 0.04 | 112 | 7 7 | 1 A | | |
| 1 9 27 2 56 48.24 | | 37 7.93N | 104 42.23W | 2.9 | 0.0 | 0.05 | 75 | 10 10 | 1 A | | |
| 1 9 27 21 2 18.59 | | 37 9.41N | 104 40.39W | 3.4 | 0.0 | 0.01 | 264 | 5 5 | 3 B | | |
| 1 10 5 0 51 52.32 | | 37 8.67N | 104 41.46W | 2.7 | 0.0 | 0.06 | 75 | 8 8 | 1 A | | |
| 1 10 6 18 56 30.73 | | 37 9.20N | 104 40.85W | 3.7 | 0.0 | 0.03 | 144 | 7 7 | 3 A | | |
| 1 10 6 23 8 55.35 | | 37 8.95N | 104 41.01W | 3.5 | 0.0 | 0.03 | 86 | 11 11 | 2 A | | |
| 1 10 7 0 22 0.53 | | 37 8.99N | 104 40.93W | 3.5 | 0.0 | 0.02 | 140 | 7 7 | 3 A | | |
| 1 10 7 2 42 22.15 | | 37 9.08N | 104 40.96W | 3.4 | 0.0 | 0.03 | 142 | 7 7 | 2 A | | |
| 1 10 7 2 53 21.05 | | 37 8.76N | 104 41.27W | 3.5 | 0.0 | 0.02 | 205 | 6 6 | 2 A | | |
| 1 10 9 1 27 59.02 | | 37 7.50N | 104 42.94W | 2.9 | 0.0 | 0.05 | 69 | 7 7 | 2 A | | |
| 1 10 11 7 41 46.01 | | 37 8.60N | 104 41.73W | 2.7 | 0.0 | 0.04 | 197 | 5 5 | 2 A | | |
| 1 10 12 13 47 8.97 | | 37 7.65N | 104 41.87W | 5.9 | 0.0 | 0.03 | 117 | 8 8 | 1 A | | |
| 1 10 15 10 24 16.59 | | 37 7.57N | 104 41.84W | 5.6 | 0.0 | 0.04 | 116 | 7 7 | 1 A | | |

Table 11. Modeled Parameters for different formations [Bohlen, 2013, COGCC, 2016, Robertson, 1988]

| Modeled Formation(s) | Lithology | POROSITY | XYZ PERM | Wet Heat Conductivity (W/m·°C) | Specific Heat (J/kg·°C) | Pore Compressibility (1/Pa) | Pore Expansivity (1/°C) |
|----------------------|-------------------------|----------|----------|--------------------------------|-------------------------|-----------------------------|-------------------------|
| Pierre Shale | Shale | 0.095 | 4.69E-18 | 1.2-3 | 777 | 1.00E-10 | 2.70E-05 |
| Niobrara | Limestone | 0.09 | 9.87E-17 | 2-3.4 | 1000 | 1.11E-10 | 2.40E-05 |
| Benton | Shale | 0.11 | 4.93E-16 | 1.2-3 | 777 | 1.00E-10 | 2.70E-05 |
| Dakota-Dockum | Sandstone | 0.1 | 1.25E-14 | 1.5-4.2 | 930 | 1.43E-09 | 3.00E-05 |
| Sangre De Cristo | Sandstone | 0.06 | 1.06E-15 | 1.5-4.2 | 930 | 1.43E-09 | 3.00E-05 |
| Basement | Granite, Gneiss, Schist | 0.016 | 1.58E-18 | 2.4-3.8 | 1000 | 2.00E-11 | 2.40E-05 |

Table 12. Sampled temperatures at the storage tanks on the disposal well sites (Hal Mccartney, personal comm.).

| Well | Test Date | Temperature °C |
|------|-----------|----------------|
| 1 | 12/14/00 | 11.6 |
| 1 | 1/31/04 | 15.5 |
| 1 | 5/18/05 | 19.9 |
| 2 | 8/14/01 | 28 |
| 2 | 8/23/01 | 28.4 |
| 3 | 10/29/01 | 24.4 |
| 4 | 10/1/02 | 17.9 |
| 4 | 2/28/06 | 11.3 |
| 5 | 1/31/04 | 16.1 |
| 6 | 8/11/10 | 26.6 |
| 7 | 2/15/11 | 13.1 |
| 8 | 4/2/14 | 16.9 |
| 8 | 4/2/14 | 14.7 |

Results

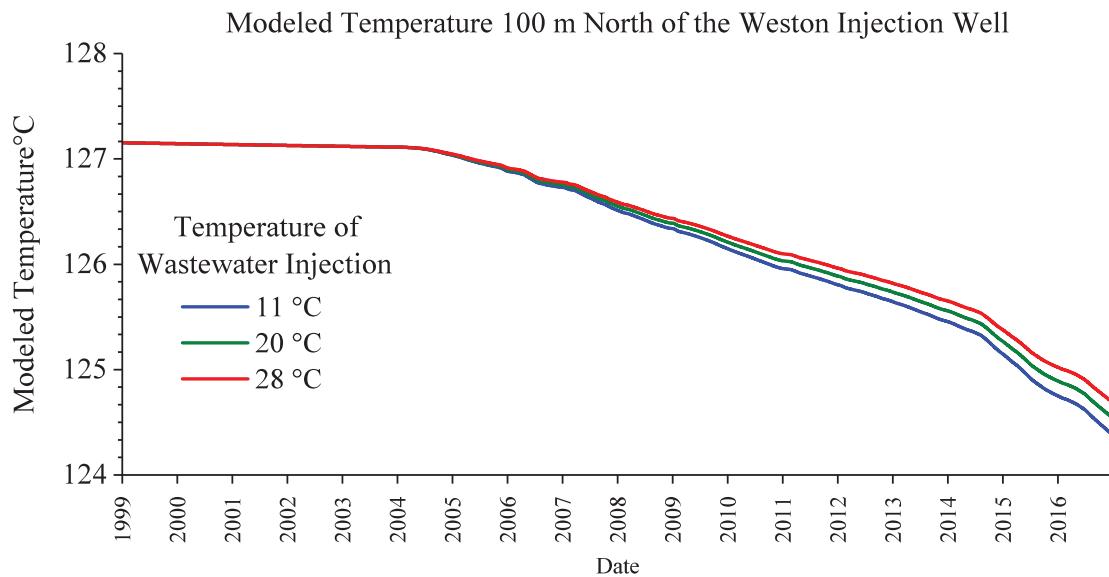


Figure 43. Sensitivity analysis for wastewater injection temperature for the Weston injection well

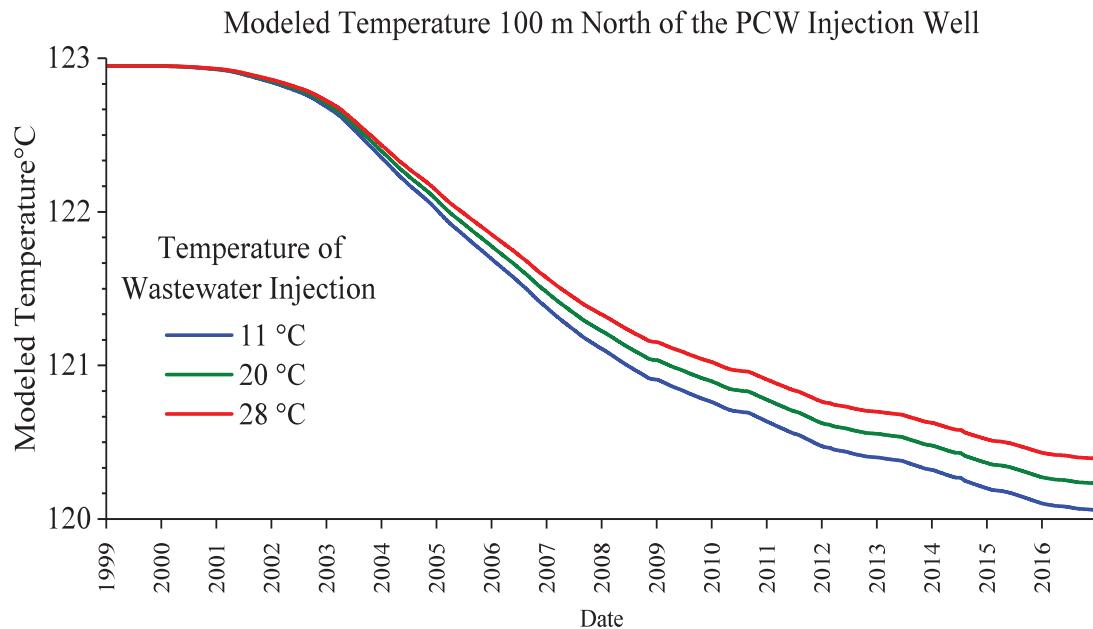


Figure 44. Sensitivity analysis for wastewater injection temperature for the PCW injection well

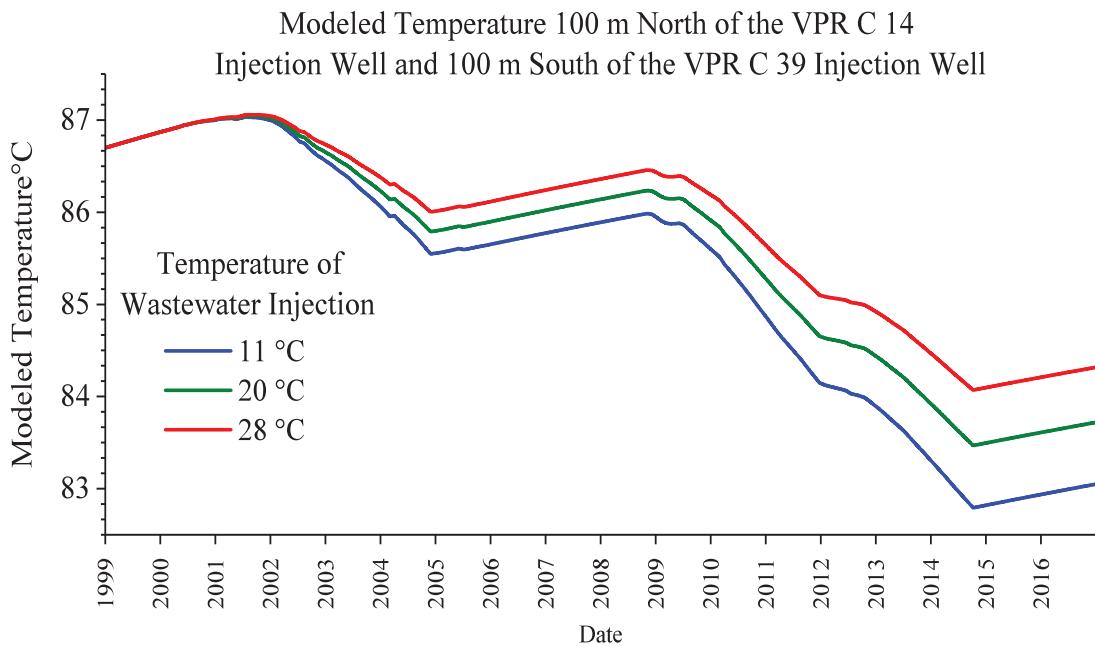


Figure 45. Sensitivity analysis for wastewater injection temperature for the VPR C 14 and 39 injection wells.