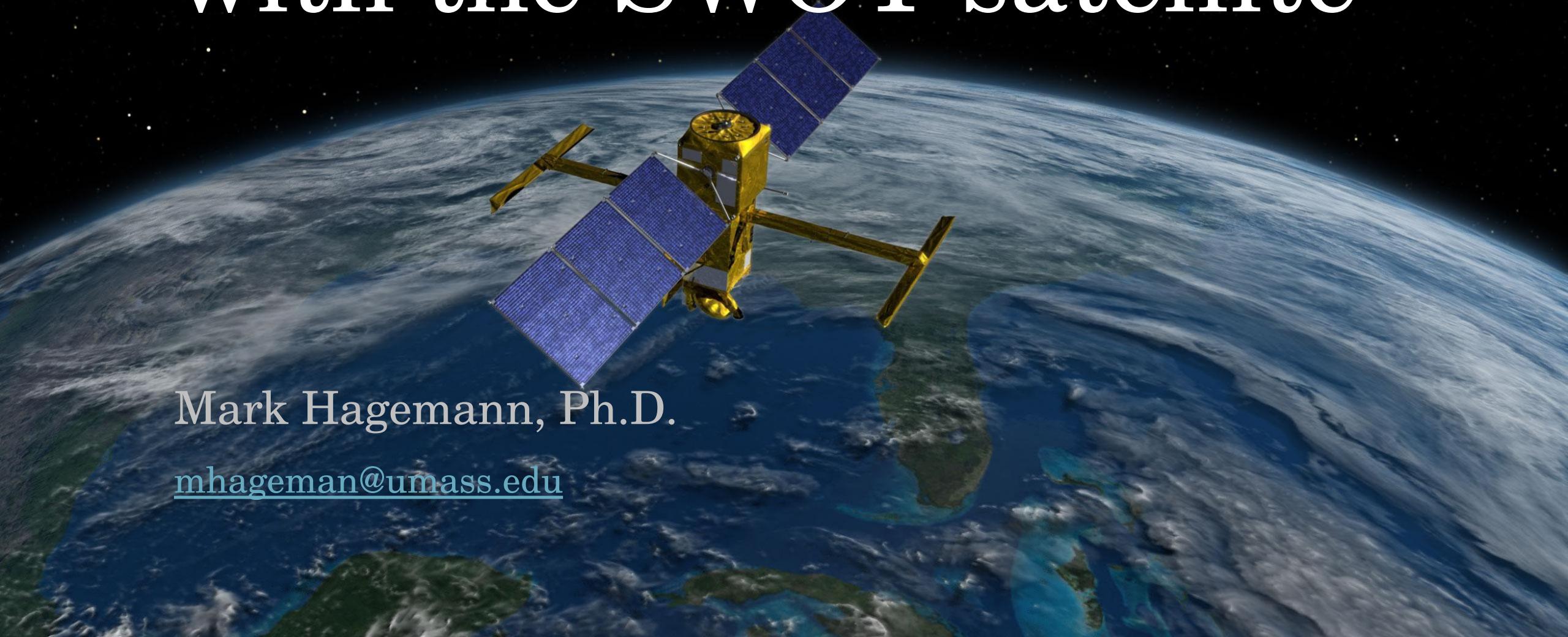


Estimating streamflow with the SWOT satellite

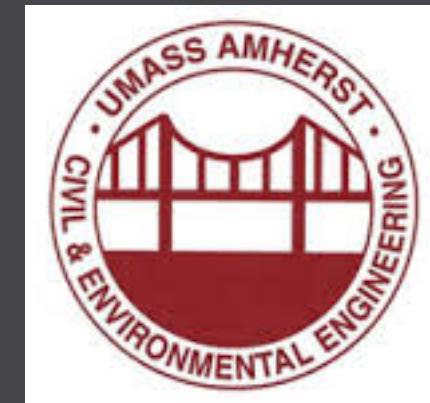
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- NASA Jet Propulsion Lab
- UMass Fluvial hydrology group, Dr. Colin Gleason



Overview

- SWOT satellite mission
- Bayesian model for river discharge using SWOT observations
- HydroSWOT proxy dataset
- Results

Bayes and data assimilation

-

SWOT satellite mission

- Surface Water & Ocean Topography
- NASA, CNES
- 2021 Launch

Measurements

- Height (H) via RADAR interferometry
- Width (W) via #####
- Slope via height / distance
- dA via integration of $W \, dH$

The model (1)

$$Q = \frac{1}{n} A^{5/3} W^{-2/3} S^{1/2}$$

- Manning's equation

The model (2)

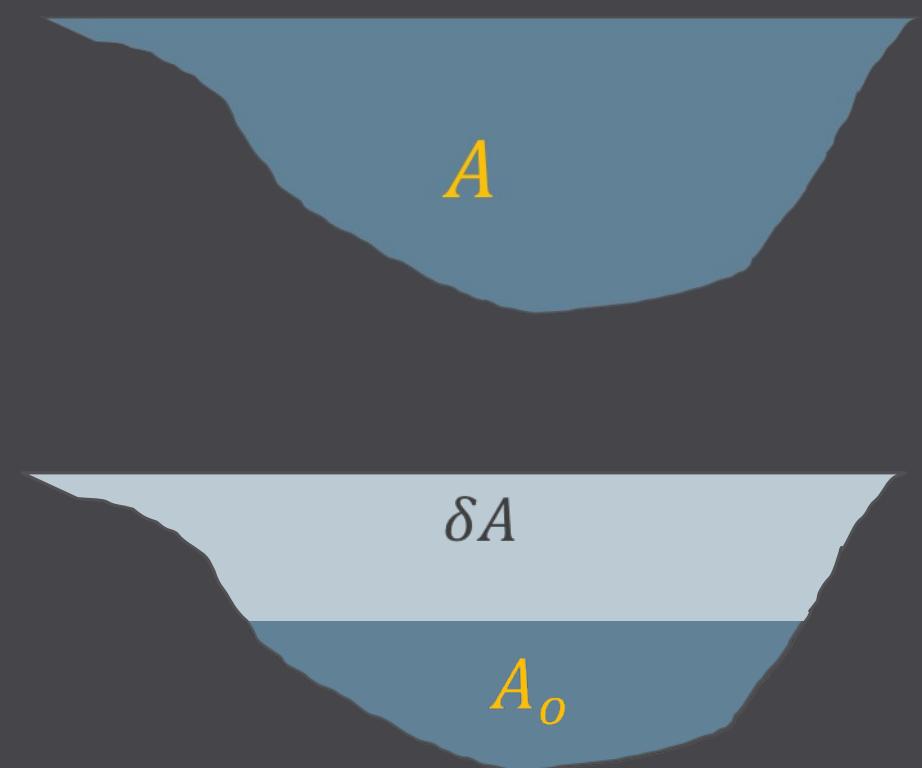
$$\log Q = -\log n + \frac{5}{3} \log A - \frac{2}{3} \log W + \frac{1}{2} \log S$$

- Manning's equation
- Log-transform

The model (3)

$$\log Q = -\log n + \frac{5}{3} \log(A_0 + \delta A) - \frac{2}{3} \log W + \frac{1}{2} \log S$$

- Manning's equation
- Log-transform
- Adjust for SWOT observables



The model (4)

$$\log Q_{it} = -\log n + \frac{5}{3} \log(A_{0,i} + \delta A_{it}) - \frac{2}{3} \log W_{it}$$

- Manning's equation
- Log-transform
- Adjust for SWOT observables
- Index in time and space



The model (5)

$$\log Q_t = -\log n + \frac{5}{3} \log(A_{0,i} + \delta A_{it}) - \frac{2}{3} \log W_{it}$$

- Manning's equation
- Log-transform
- Adjust for SWOT observables
- Index in time and space
- Mass conservation

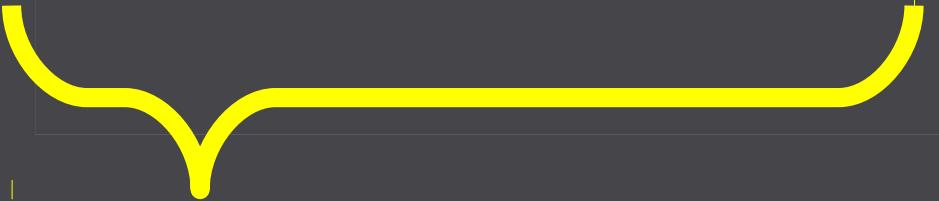


The model (6)

$$\frac{2}{3} \log W_{it} - \frac{1}{2} \log S_{it} = -\log n + \frac{5}{3} \log(A_{0,i} + \delta A_{it})$$
$$- \log Q_t + \epsilon_{it}$$

- Manning's equation
- Log-transform
- Adjust for SWOT observables
- Index in time and space
- Mass conservation
- Rearrange, acknowledge error

The model (7)

$$\frac{2}{3} \log W_{it} - \frac{1}{2} \log S_{it} = -\log n + \frac{5}{3} \log(A_{0,i} + \delta A_{it})$$

$$-\log Q_t + \epsilon_{it}$$

$$f(y_{it}|n, Q_t, A_{0,i}, \delta A_{it}) = N(\mu = \text{RHS}, \sigma)$$

- Observables vary in time and space
- Quantities of interest vary in time or space (not both)
- Random error

The model (8)

$$f(y_{it}, \delta A_{it} | n, Q_t, A_{0,i})$$

$$f(y_{it} | n, Q_t, A_{0,i}, \delta A_{it}) = N(\mu = \text{RHS}, \sigma)$$

- Observables vary in time and space
- Quantities of interest vary in time or space (not both)
- Random error

The model (9)

$$f(y_{it}, \delta A_{it} | n, Q_t, A_{0,i}) = f(y_{it} | n, Q_t, A_{0,i}, \delta A_{it}) \\ \times f(\delta A_{it} | n, Q_t, A_{0,i})$$

$$f(y_{it} | n, Q_t, A_{0,i}, \delta A_{it}) = N(\mu = \text{RHS}, \sigma)$$

- Observables vary in time and space
- Quantities of interest vary in time or space (not both)
- Random error

$$f(\delta A_{it} | \textcolor{blue}{n}, Q_t, A_{0,i})$$

$$f(\delta A_{it} | \textcolor{blue}{n}, \textcolor{red}{Q_t}, A_{0,i}) = f(\delta A_{it} | A_{0,i}) = f(\delta A_{it} | A_{0,i}, \sigma_A)$$

$$\begin{aligned} &= \frac{1}{(\delta A + A_0) \sqrt{2\pi} \sigma_A} \\ &\times \exp \Big(- \frac{1}{2\sigma_A^2} [\log(\delta A + A_0) - \log A_0]^2 \Big), \\ & (\delta A > -A_0) \end{aligned}$$

HydroSWOT database

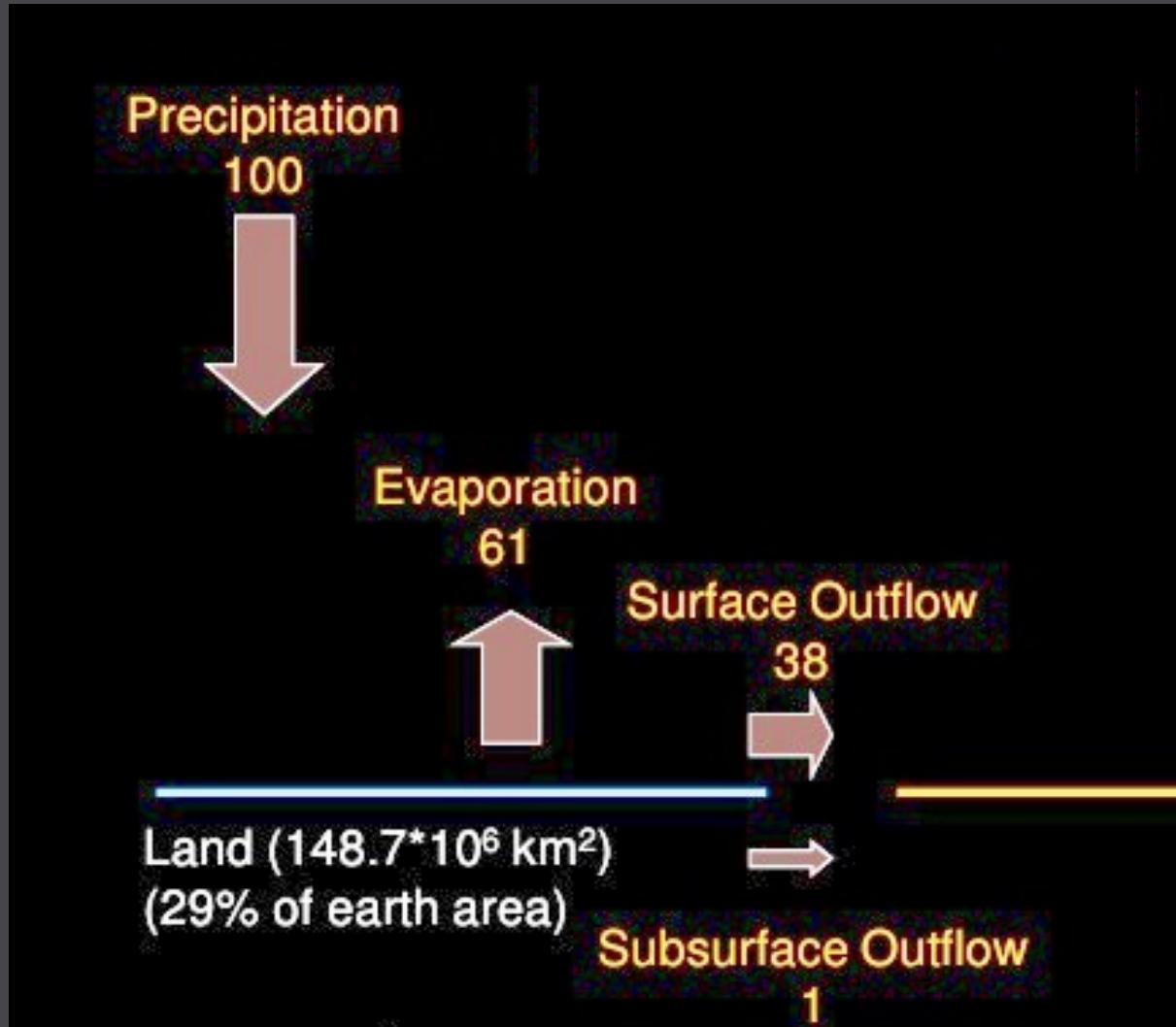
- Direct measurement
- Stream gages

Results

How is Q measured?

- Direct measurement
- Stream gages
 - Great for USA
 - Less great for Africa

$$Q = P - ET$$



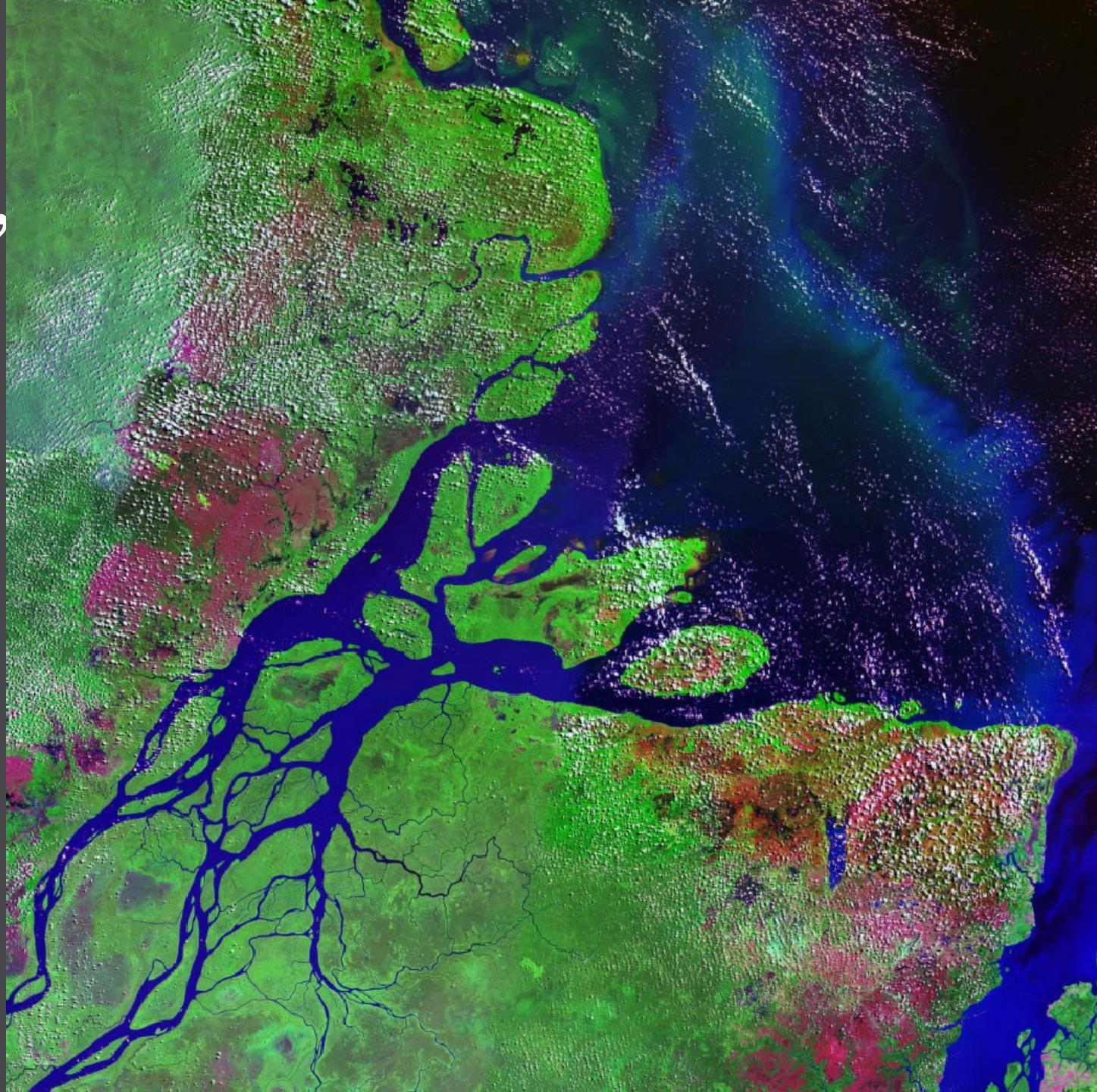
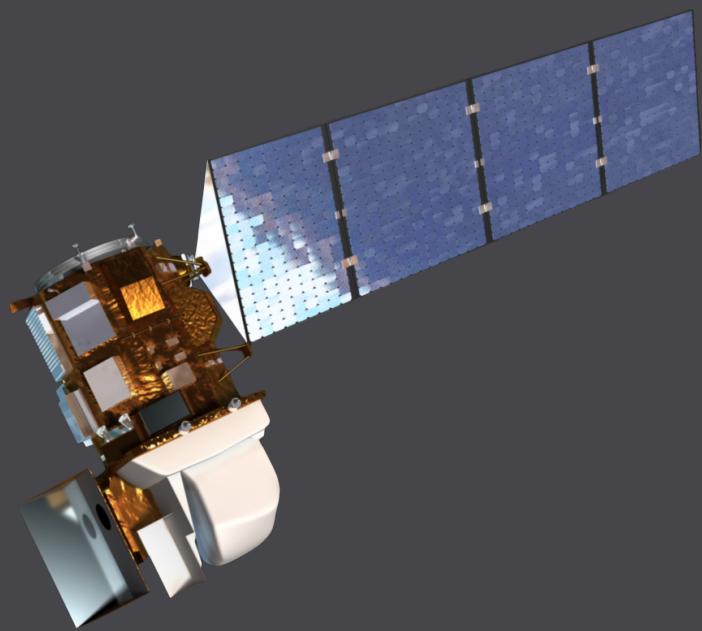
How is Q measured?

- Direct measurement ← Labor-intensive
 - Stream gages ← Labor-, \$-intensive
 - Water-balance models ← Imprecise
-
- ← Insufficient data

How is Q measured?

- Direct measurement
- Manning's equation: $Q = \frac{1}{n} \frac{A^{5/3}}{W^{2/3}} S^{1/2}$
- Streamgages: sectional area
- Water-balance models
 - W: stream width
- "Flow laws"
 - Relate Q to slope, bed material, channel geometry
- At-a-station hydraulic geometry: $Q = aW^b$
 - W: stream width
 - a, b: parameters relating to channel geometry

More on "flow laws"



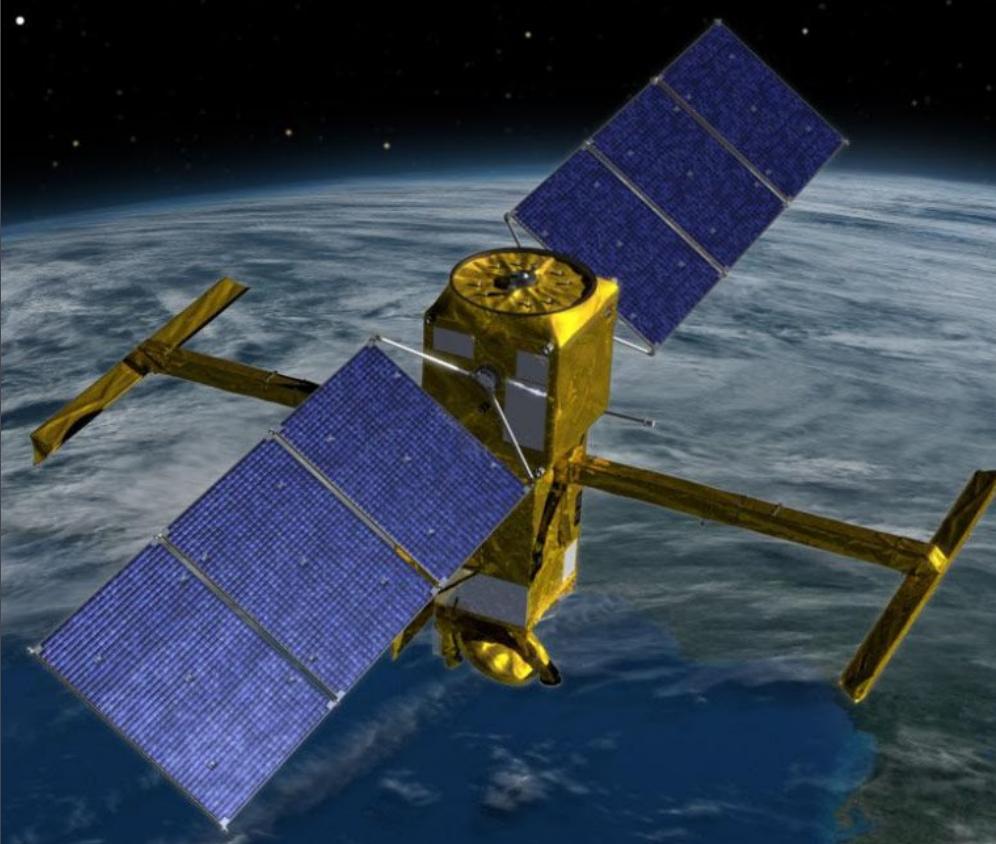
Satellite Remote Sensing

- Pictures of Earth from space
- Upcoming NASA
- E.G. LANDSAT (1972 present)
satellite mission
(2020 launch)
- Stream width $\pm 15m$
- Stream height
 $\pm 10 cm$
- Stream slope
 $\pm 2 cm/km$



For any stream reach worldwide!

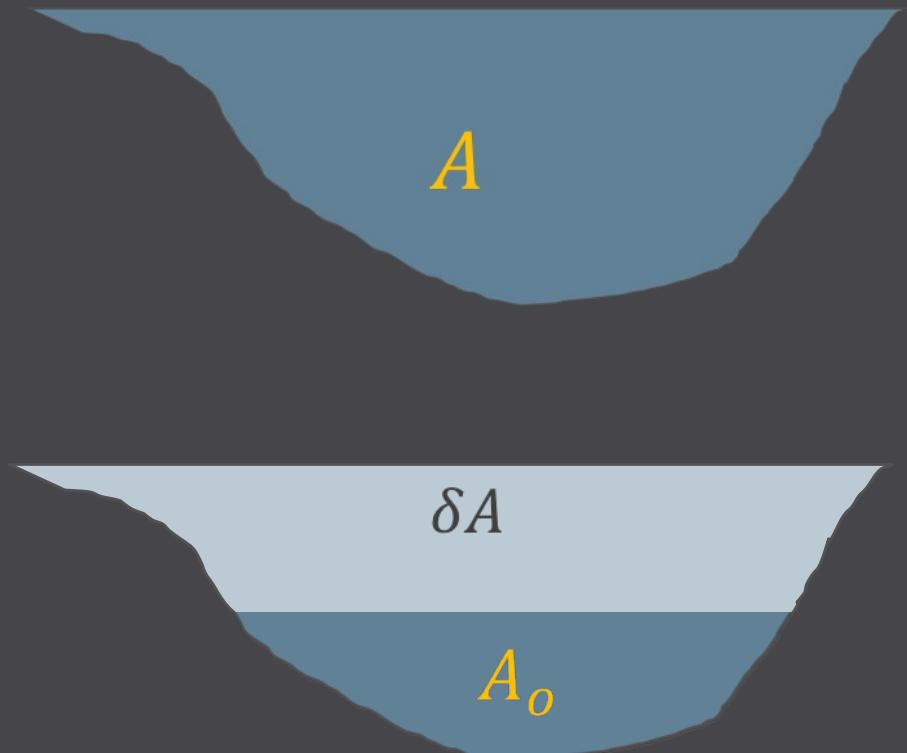
Surface Water and Ocean Topography (SWOT)



Can we get Q from SWOT measurements?

• Using flow-law equations
Manning's equation: $Q = \frac{1}{n} \frac{(A_o + \delta A)^{5/3}}{W^{2/3}} S^{1/2}$

- A_o : cross-sectional area at smallest measured width
- δA : change in area from A_o
- n : “manning’s n” (relates to bed roughness)
- W : stream width
- S : stream slope



Can we solve the flow-law equations?

- Manning's equation: $Q = \frac{1}{n} \frac{(A_o + \delta A)^{5/3}}{W^{2/3}} S^{1/2}$
 - Multiple cross-sections
- Conserve flow in reach: $Q_1 = Q_2$
- More unknowns than equations!



Can we solve the flow-law
Bayes briefly $p(A|B) \propto p(B|A)p(A)$
equations?





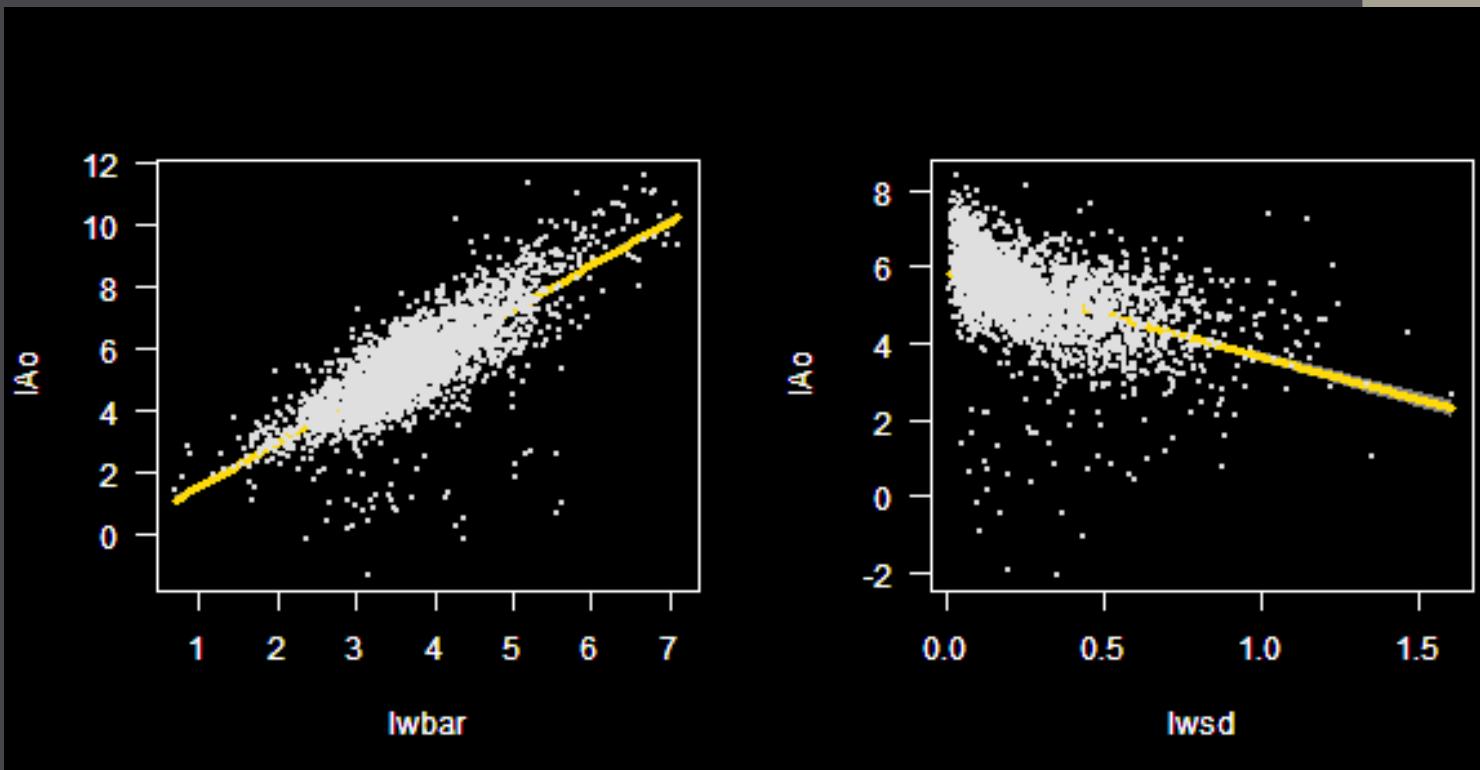
- Describe and constrain uncertain quantities
- Get a *distribution* for unknowns
- Computationally expensive

```
parameters {
    vector[nt] logQ;
    real logn;
    real<lower=30> Ao[nx];
}

model {
    logn ~ normal(logn_hat, logn_sd);
    logQ ~ normal(logQ_WBM,
cv2sigma(cvQ));
    for (i in 1:nx) {
        man_lhs[i] ~ normal(man_rhs[i],
sigma_man);
    }
    Ao ~ lognormal(logAo_hat, logAo_sd);
}
```

Bayes with Stan language

- Code looks like math!
- Bayes-specific, Bayes-optimized
- Open-source
- Plays nice with scientific computing environments



Constraining “prior distributions”

- Ao relates to width statistics
- Wider rivers are deeper on average
- More width-variable rivers are shallower on average
- n: from literature

 **AGU PUBLICATIONS**

Water Resources Research

RESEARCH ARTICLE
10.1002/2015WR018434

Key Points:

- SWOT discharge algorithms were tested on synthetic observations for 19 rivers
- Algorithms accurately characterized temporal dynamics of river discharge
- At least one algorithm estimated discharge to <35% relative RMSE on 14/16 of nonbraided rivers

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Citation:
Durand, M., et al. (2016), An intercomparison of remote sensing river discharge estimation algorithms from measurements of river height, width, and slope, *Water Resour. Res.*, 52, 4527–4549, doi:10.1002/2015WR018434.

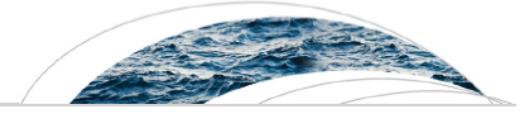
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Published online 15 JUN 2016

An intercomparison of remote sensing river discharge estimation algorithms from measurements of river height, width, and slope

M. Durand¹, C. J. Gleason², P. A. Garambois³, D. Bjerklie⁴, L. C. Smith⁵, H. Roux^{6,7}, E. Rodriguez⁸, P. D. Bates⁹, T. M. Pavelsky¹⁰, J. Monnier¹¹, X. Chen¹², G. Di Baldassarre¹³, J.-M. Fiset¹⁴, N. Flipo¹⁵, R. P. d. M. Frasson¹, J. Fulton¹⁶, N. Goutal¹⁷, F. Hossain¹⁸, E. Humphries¹⁰, J. T. Minear¹⁹, M. M. Mukolwe²⁰, J. C. Neal⁹, S. Ricci²¹, B. F. Sanders²², G. Schumann^{9,23}, J. E. Schubert²², and L. Vilmin¹⁵

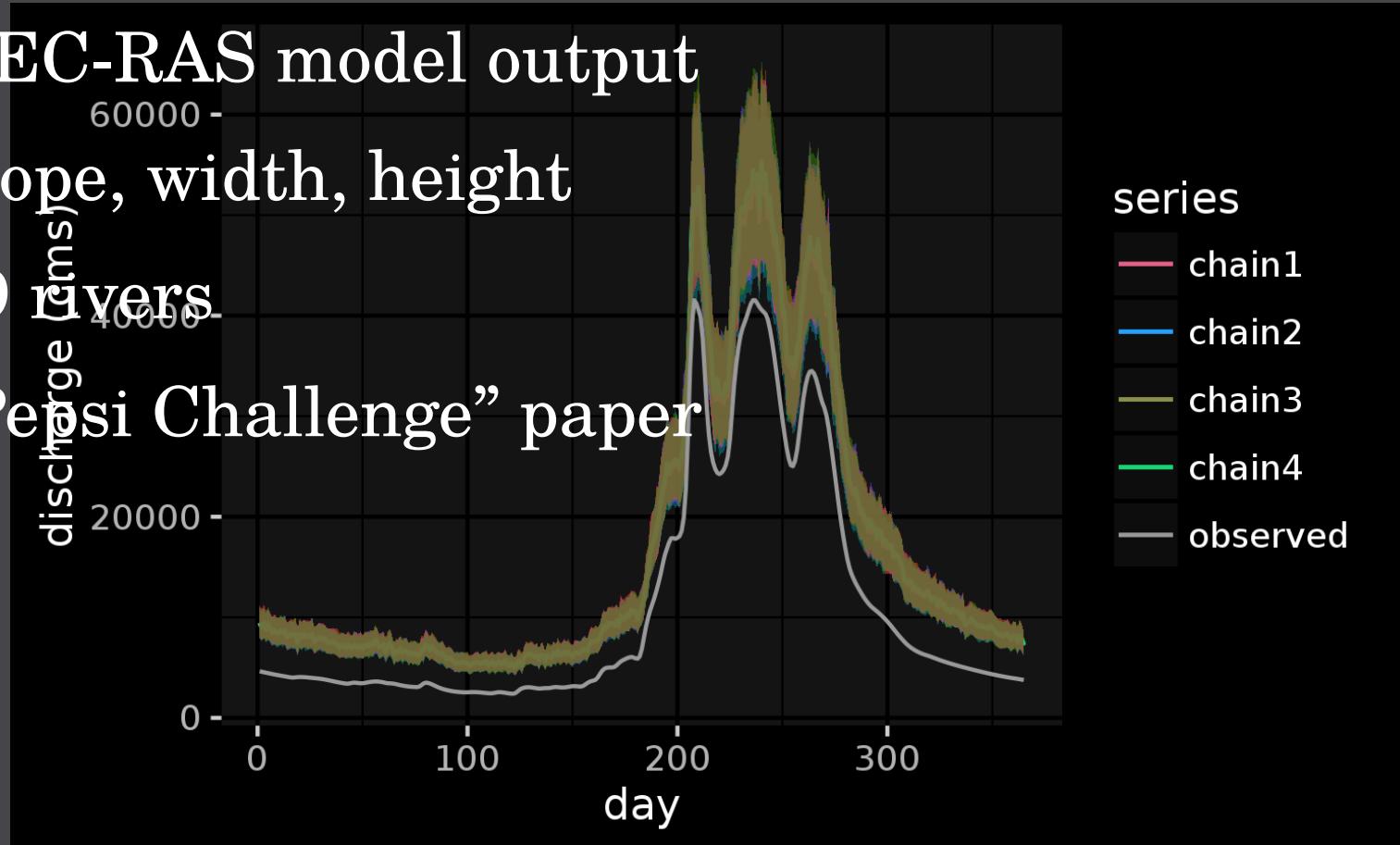
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Abstract The Surface Water and Ocean Topography (SWOT) satellite mission planned for launch in 2020 will map river elevations and inundated area globally for rivers >100 m wide. In advance of this launch, we

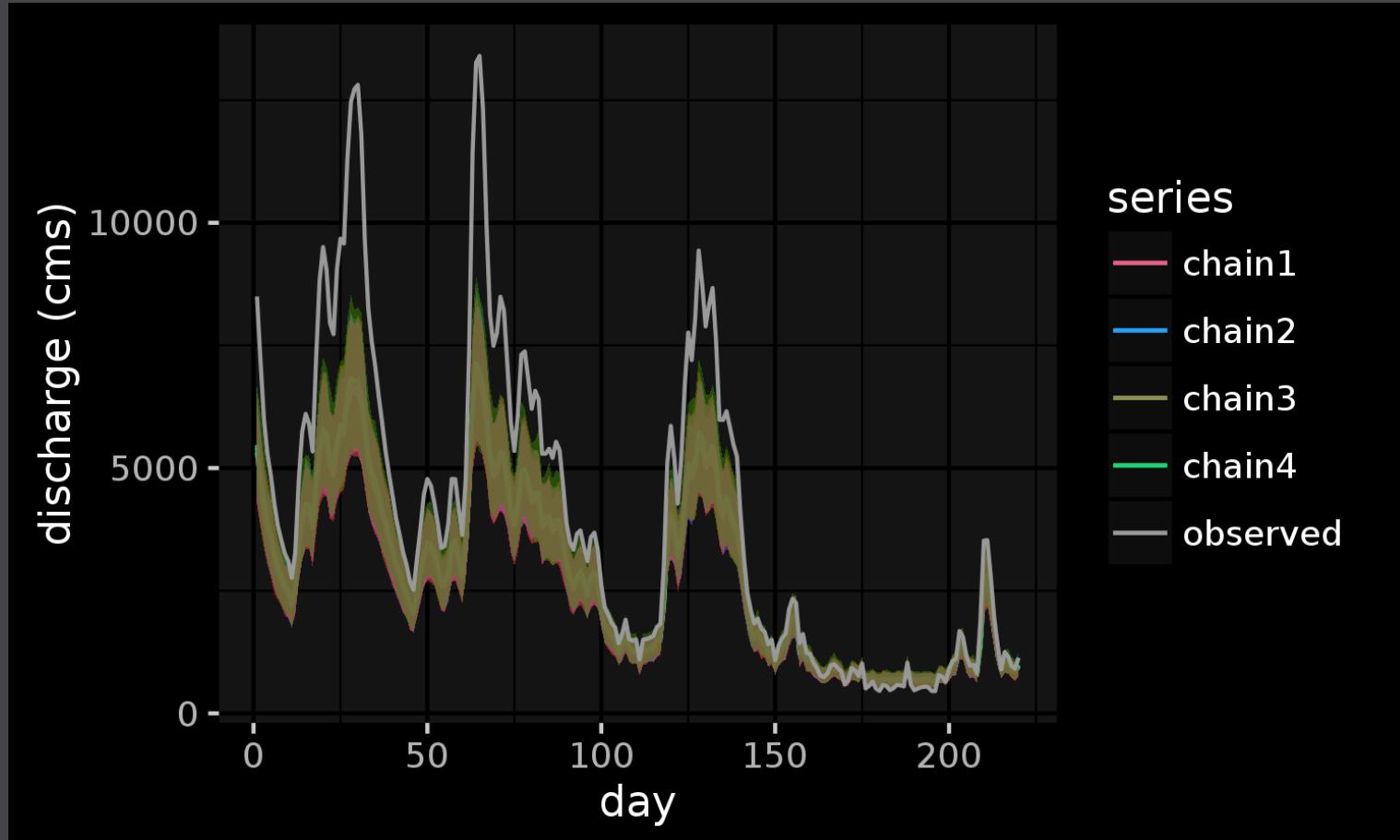


“Data”

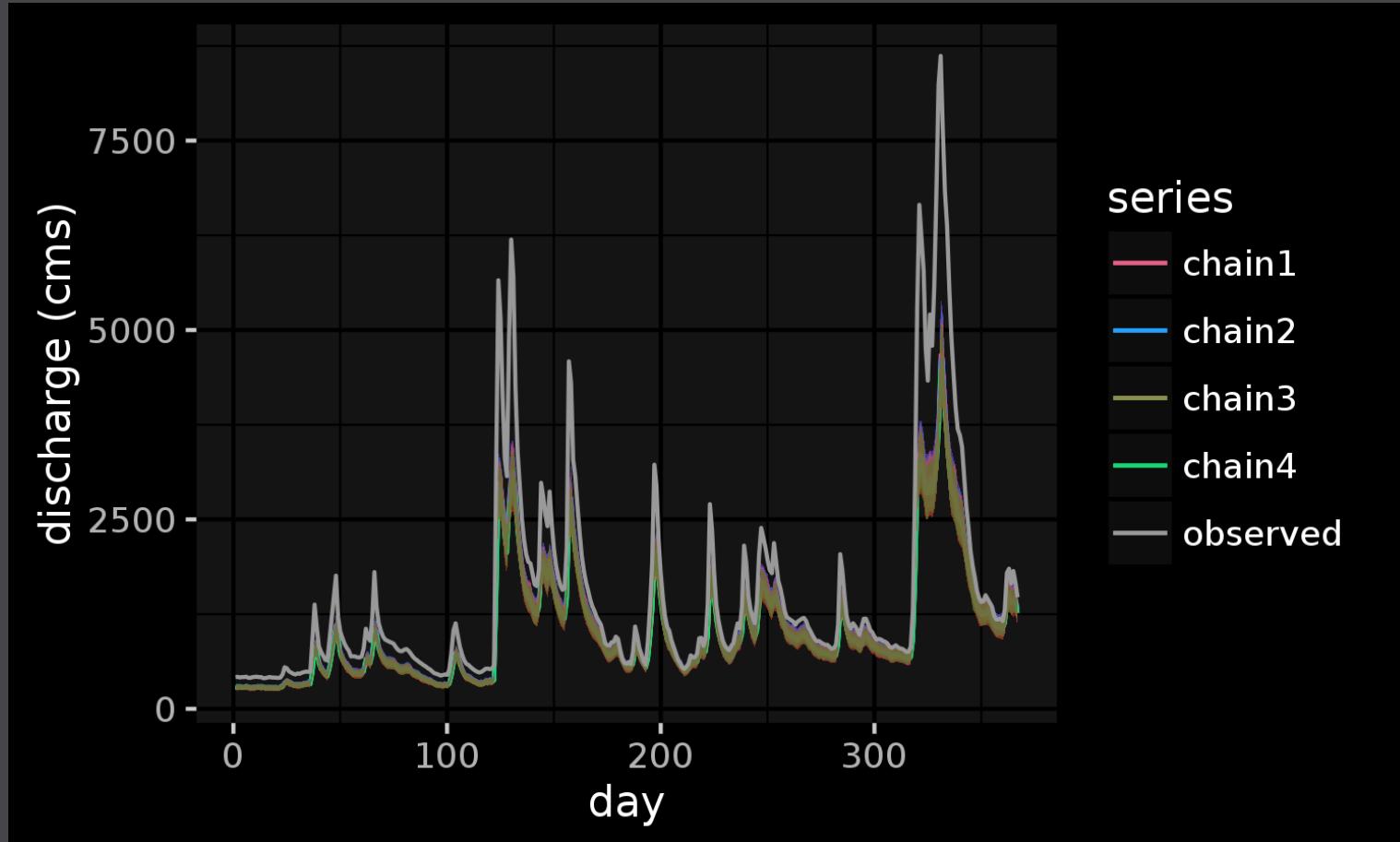
- HEC-RAS model output
- Slope, width, height
- 19 rivers
- “Pepsi Challenge” paper



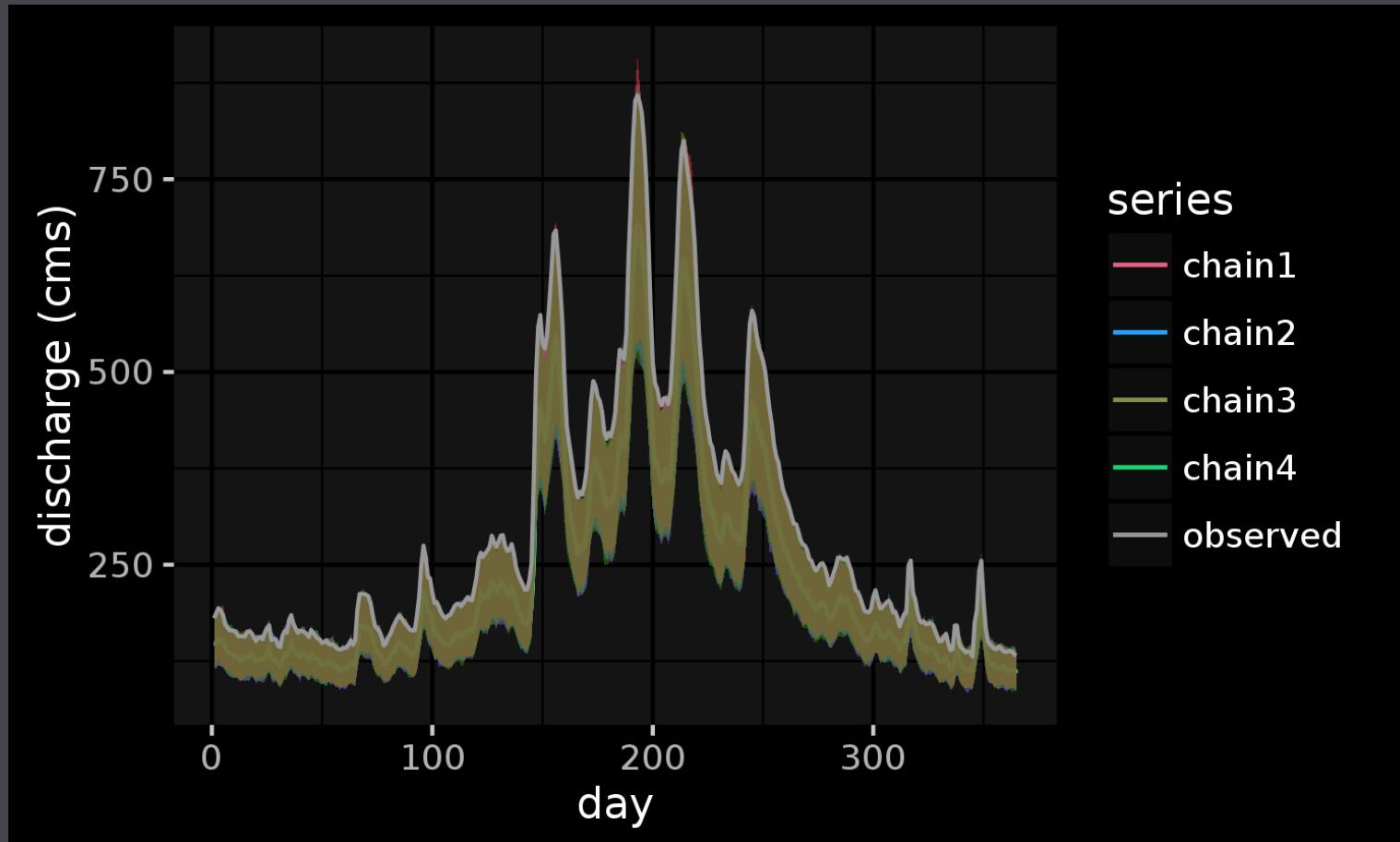
Results - Ganges



Results - Ohio



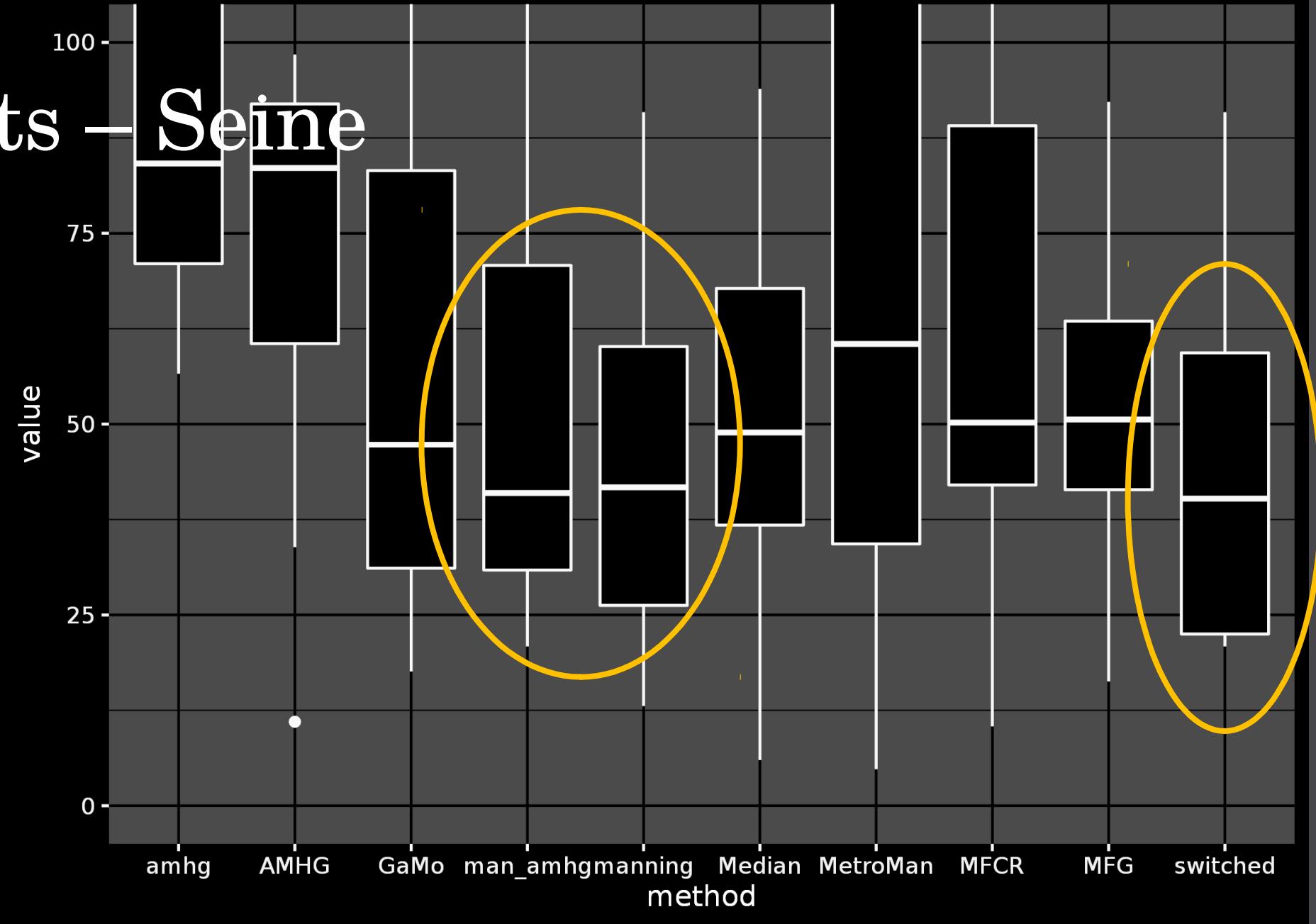
Results - Po



Results - Seine



Better



Comparing different algorithms (Pepsi Challenge)

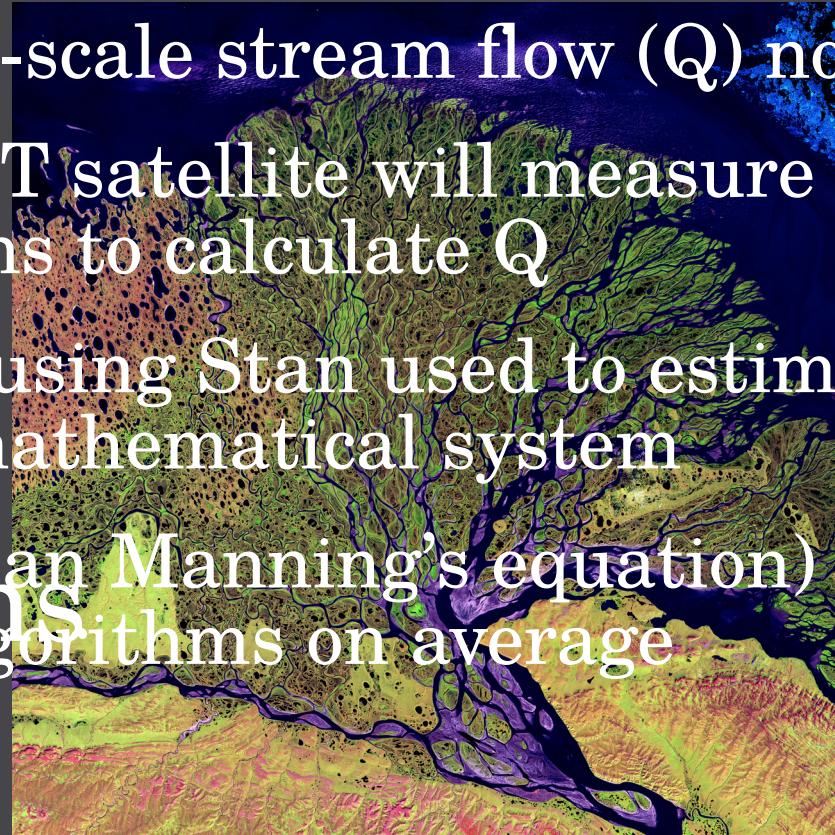
AirSWOT



Next steps

- Validate using *actual* data
 - Field datasets
 - airSWOT
- Better constrain Manning's n
 - Not actually constant
- Computer vision / deep learning for Q

- Global-, continental-scale stream flow (Q) not well constrained
- The upcoming SWOT satellite will measure many quantities used by flow-law equations to calculate Q
- Bayesian inference using Stan used to estimates Q in an underconstrained mathematical system
- Our method (Bayesian Manning's equation) outperforms other SWOT discharge algorithms on average



Conclusions

Thanks!

Thanks!