

The Effect of Bathymetry Changes on Meridional Overturning Currents

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Abstract

1 Introduction

The geometry and resulting bathymetry of our planet is an ever changing phenomenon[1]. In the last 120 Ma the earth moved from having one major oceanic system in the Pacific with a single large continent, to the current 3 ocean system. The Bathymetry changes that occurred in this time period are characterized by the opening and closing of certain passages through which exchange of water between the oceanic basins is characterized. The exact timing of passage openings is a topic of rigorous debate in literature[5][6].

One of the changes on which there is general consensus is the inception and expansion of the Atlantic ocean and the resulting decrease in size of the Pacific basin. The creation of the Atlantic basin has had major effects on the earth's climate especially resulting in massive localized changes such as the temperate European climate due to the north Atlantic meridional overturning current (AMOC). This current creates the current northern sinking situation in the Atlantic basin. These overturning currents are a large driver for the climates in the coastal regions of the earth moving a lot of thermal energy to higher latitudes. However it is unknown when exactly this northern sinking started. With the past non-existence of the Atlantic it must have started some time in the last 40Ma with the advent of a larger Atlantic.

The result of these bathymetry changes on the oceanic stream function and the resulting overturning currents is something that has been previously

studied by Mulder et al.[3] on simplified models. Here for each time step of 5Ma the previous model's outcome were used as initial conditions for the next bathymetry. This results first of all in having to interpolate the forcings at locations where there was previously no ocean and it may also result in finding different equilibrium than those that may be found by studying the changes with exactly the same forcings for every model.

Ocean modeling has been an area of continued progress. The resolutions of the models have been steadily increasing since the inception of the first computerized ocean models. However, due to the age of some of these models and the continued adaptation of often old legacy Fortran code, many models have become enormous hurdles to get started with often resulting in frustration. The Veros[2] ocean model project is trying to tackle this problem with a totally new code base written entirely in Python. Veros allows easy editing of forcing and geometry input and infinite flexibility in the model's setup without the hassle of learning Fortran.

Veros also allows a lot of flexibility in the resolution of ocean models. It allows easy scaling which results in being able to first test some changes on a low resolution model and then slowly scaling up to higher resolutions.

This paper will focus solely on changes in bathymetry using very simplified zonally averaged global forcings. The results of the model will be used to estimate global changes in oceanic through

flow at the critical passages. Furthermore the strength of the meridional overturning currents (MOC) will be studied.

Horizontal wind driven circulation

The gateways in the oceans are the connections between the different oceans.

2 Methods

2.1 On Veros

2.2 Simplified Forcings

The simplified forcings used in this paper will mostly consist of zonally averaged forcings taken from existing forcings related to the current ocean. The biggest and most obvious drawback of this approach is ignoring the massive changes in the climate in the time period on which this paper is focused. Another is the fact that zonally averaged forcings are often harder to stabilize for finer grids.

2.3 Scaling oceanic basins

2.4 Measuring overturning currents

About the overturning currents

converting the data from [4] to veros

ignoring many things idealized global temperature profile idealized global salinity idealized global wind stress

3 Results

Drake passage Titis Panama

Specifically look at changes over time.

4 Summary

References

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