

Exercise 3, TFY4235 Computational physics

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Introduction

This paper documents the implementation and results of the simulation as given in [1]. By using a finite difference scheme, the process of absorption of CO₂ by the global ocean in the form of dissolved inorganic carbon (DIC). A simplified model of the physics is used, where both the atmospheric and oceanic concentration of carbon is constant at a given height, meaning there is in effect only one spatial dimension. The method is tested against known solutions, and the convergence is monitored. It is then used to extract results of how the ocean absorbs carbon, and at what rate.

Theory and implementation

The implementation of the simulation largely follows [1], as laid out in the appendix. This section describes the particular choices made, and in a notation consistent with the code. The diffusion equation can be written as

$$\Delta t \frac{\partial}{\partial t} C(z, t) = \Delta t \left(K(z) \frac{\partial^2}{\partial z^2} + \frac{dK(z)}{dz} \frac{\partial}{\partial z} \right) C(z, t) = \mathcal{D}C(z, t).$$

\mathcal{D} is thus a linear operator which can be approximated by a discretization. With the scheme as described in [1], and including the boundary condition this gives

$$\Delta t \frac{\partial}{\partial t} C_n(t) = \mathcal{D}_{nm} C_n(t) + S_n(t),$$

where summation over repeated indices are implied, and

$$\mathcal{D} = \begin{pmatrix} -4\alpha K_0 - 2\Gamma & 4\alpha K_0 & \dots & 0 & 0 \\ \vdots & \ddots & \ddots & 0 & 0 \\ 0 & -\frac{\alpha}{2} K'_i + 2\alpha K_i & -4\alpha K_i & \frac{\alpha}{2} K'_i + 2\alpha K_i & 0 \\ 0 & 0 & \ddots & \ddots & \vdots \\ 0 & 0 & \dots & 4\alpha K_N & -4\alpha K_N \end{pmatrix}, \quad \alpha = \frac{\Delta t}{2\Delta z^2},$$
$$S(t) = (2\Gamma C_{\text{eq}}(t) \quad 0 \quad \dots \quad 0)^T \quad \Gamma = 2 \frac{\alpha k_w \Delta z}{K_0} \left(K_0 - \frac{1}{2} \left(-\frac{3}{2} K_0 + 2K_1 - \frac{1}{2} K_2 \right) \right), \quad K'_n = K_{n+1} - K_{n-1}.$$

Using the Cranck-Nichelson scheme to discretize the time part yields

$$C_n^{i+1} = C_n^i + \frac{1}{2} (\mathcal{D}_{nm} C_m^i + S_n^i) + \frac{1}{2} (\mathcal{D}_{nm} C_m^{i+1} + S_n^{i+1}),$$

so the equation to be solved to get the next timestep is

$$A_{nm} C_m^{i+1} = V_n^i, \quad V_n^i = \left(\delta_{nm} + \frac{1}{2} \mathcal{D}_{nm} \right) C_m^i + \frac{1}{2} (S_n^i + S_n^{i+1}), \quad A_{mn} = \left(\delta_{nm} - \frac{1}{2} \mathcal{D}_{nm} \right).$$

The implementation of this system of equation uses SciPy’s sparse matrix library. After creating sparse realizations of A , SciPy’s `splu` is used to generate the LU decomposition LU of A . This is an object with methods such as `.solve()`, which utilizes the LU decomposition. The wrapper `simulate` then loops over $N_t - 1$ steps, using `solve(V) = lambda V: LU.solve(V)`, where V is as given above. The results of the simulation of C is kept in a NumPy array. However, it is not necessary to keep all steps made in time. `simulate` can be passed a integer argument `save`, the number of time steps to be kept in the array. This entails that $N_t - 1$ is divisible by `save - 1`, which is checked by an `assert` statement.

To check convergence of the method, both in time and space, a way of measuring error is needed. In this project, relative root mean square (RMS) error is used. Given a concentration $C(z)$ and a reference $C_0(z)$ at a given time, sampled at a set of points $\{z_i\}_{i=1}^N \in [0, L]$, relative RMS error is given by

$$\Delta_{\text{rms}} = \sqrt{\frac{1}{N} \sum_i \left(\frac{C(z_i) - C_0(z_i)}{C_0(z_i)} \right)^2}.$$

When comparing concentrations with different values for N_z , as is necessary to find the convergence of the method in space, one must make sure that the points of comparison are the same for both C and C_0 . This is done by setting $N_z = 5 \cdot 2^N a + 1$ for C_0 , and $N_z = 2^n a + 1$. Here, a and N are integers, and n is some integer less or equal to N . This ensures that the set of points $\{z_i\}_{i=1}^{2^n a + 1}$ that $C(z)$ is simulated on, $C_0(z)$ is as well, and they can be compared. This is implemented by `get_rms` in `utilities.py`, which takes a list of arrays and compare each with the last element. The task also requires the spatial integral of the concentration and related quantities. This is done with the SciPy implementation of Simpsons method, `simpson`.

Tests

Make sure the implementation gives good answers, a series of tests are run, where the answer is known or there is a method to check error. The method used in this implementation has quadratic convergence, both in time and space. A convergence test was implemented for a simple test case with constant diffusivity K and a Gaussian initial concentration. The result is shown in Figure 1. A constant concentration of CO_2 should remain constant, regardless of $K(z)$. This test is shown in Figure 2, with a both a constant $K(z)$, and a smooth step function between two values of K as described in [1], problem 2. For both, the constant initial concentration is close to unchanged, save for numerical errors. The systems should also, given $k_w = 0$, conserve mass. To test this, a initial distribution of two Gaussian functions were evolved in time, both with constant and non-constant diffusivity. The result is shown in Figure 3, where the mass is conserved to a good approximation.

A sharply peaked Gaussian package should have a variance that increases linearly with time, then approach a steady state. This is shown in Figure 4. For a small Biot nuber and given a zero partial pressure in the atmosphere, the depletion of CO_2 from the ocean should follow a exponential decay. Furthermore, the concentration should remain constant throughtout the depht. This is the case, as Figure 5 showscases. Lastly, Figure 6 shows how the ocean reaches a equilibrium with the atmosphere, given a non-zero, positive mass-transfer coefficient k_w and either constant or non-constant diffusivity.

Results

Shallow waters

To simulate the uptake of CO_2 in shallow areas, $L = 100\text{m}$, diffusivity $K(z)$ is used as given in equation (11) in [1]. Figure 7 illustrates the convergence of the simulation, which shows that the simulation still has quadratic convergence both in time and space. The convergence test are run for a time of 10 days.

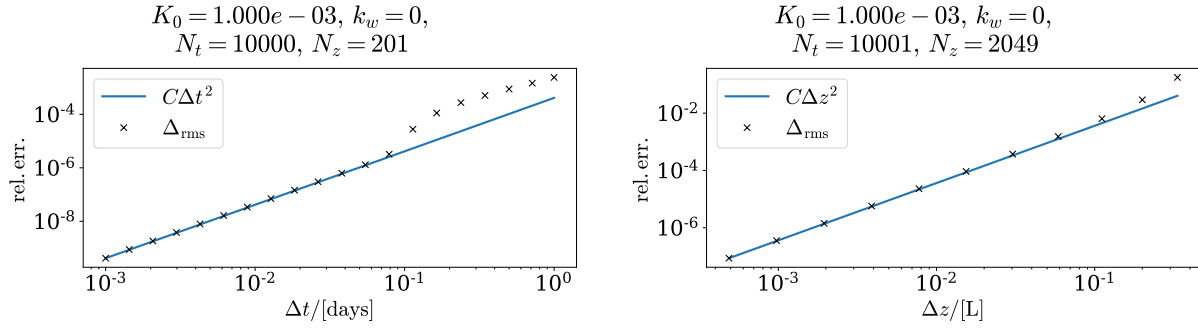


Figure 1: Error, measured as the relative RMS deviation from a reference value, after 1 day simulation of an Gaussian initial concentration.

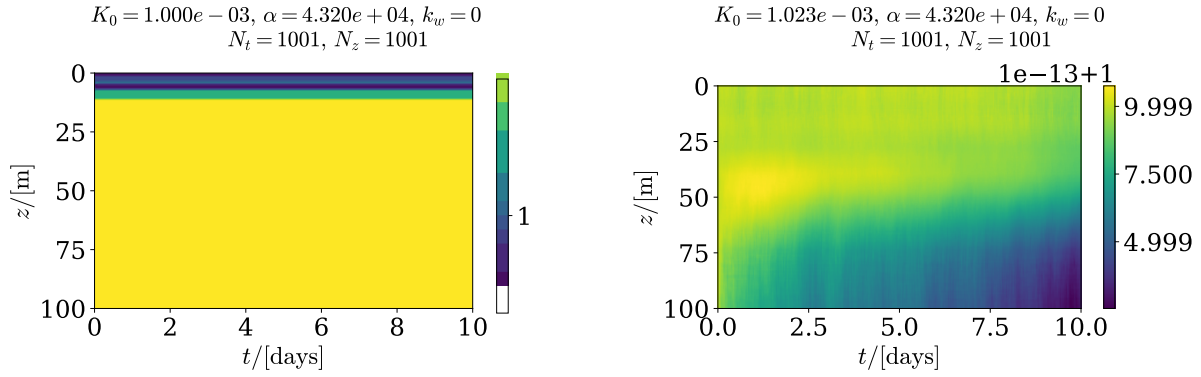


Figure 2: The time evolution of an constant concentration. $K(z)$ is constant in the system on the left, and varies in the system on the right. The largest deviation from the initial values of the systems are of order 10^{-15} and 10^{-13} , respectively.

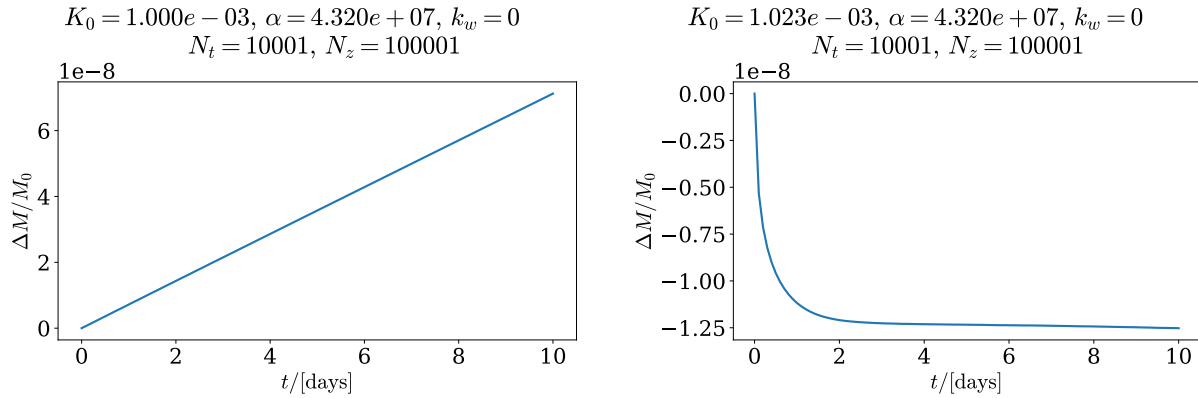


Figure 3: Relative change in mass of a system with constant (left) and varying (right) diffusivity.

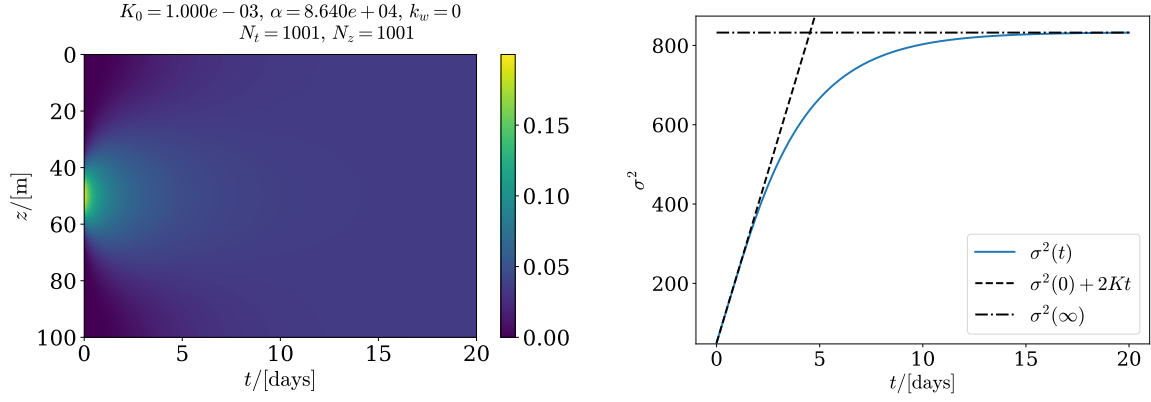


Figure 4: A system with an initial Gaussian distribution. On the right, the variance as a function of time is shown. The increase is initially constant, but then approaches a steady state.

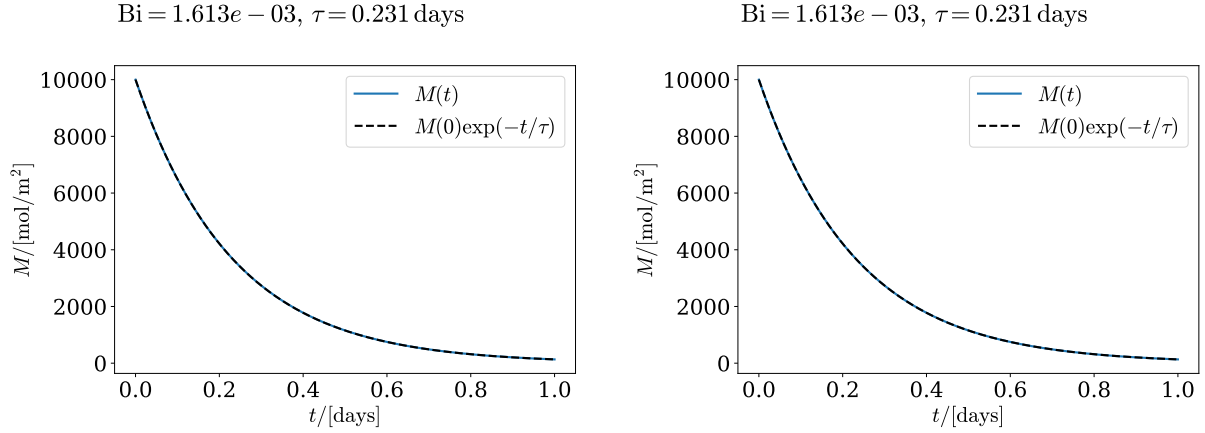


Figure 5: The slow removal of CO₂ from the ocean, when the atmosphere contains a partial pressure of 0. Constant (left) and varying (right) diffusivity.

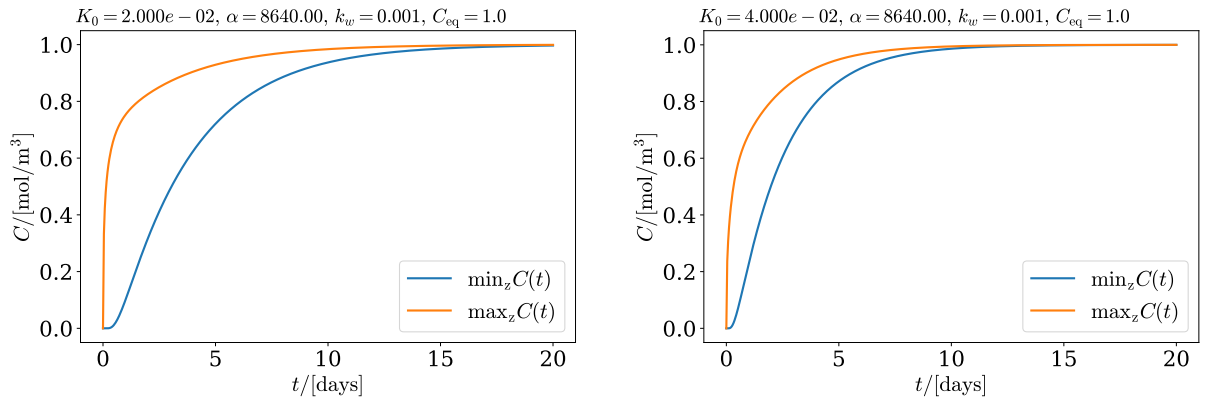


Figure 6: Equilibration of the ocean with an atmosphere with $C_{\text{eq}} = 0.5$. The figure on the left is a system with constant diffusivity, the left side has an oscillating but always positive diffusivity.

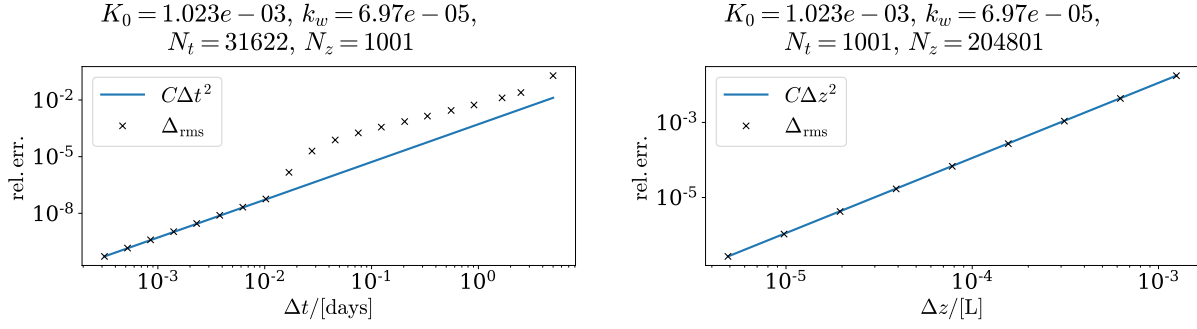


Figure 7: Error, as a function of step length in both time and space.

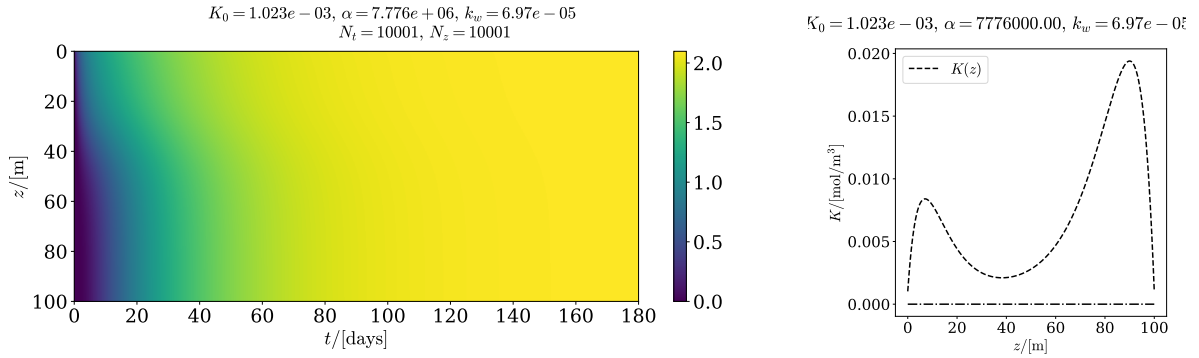


Figure 8: cap

The simulation is then run for 180 days, with $N_t = N_z = 10000$. This corresponds to $\Delta t = 180/10000 = 1.8 \cdot 10^{-2}$ days, and $\Delta z = L/10000 = 10^{-4}L$. The results are shown in Figure 8, together with the form of the diffusivity, $K(z)$. Figure 9 shows how the minimum and maximum concentration evolves as a function of time, together with some snap shots of the concentration as it evolves. The surface of the ocean, where wind dominates the diffusion, first absorbs the CO_2 . The deeper parts of the ocean lacks behind some, as there low diffusivity around 40 meters are a bottle neck, but it does not take long before it penetrates down to the ocean floor.

At 90 days, the ocean is nearly in equilibrium with the atmosphere, and at 180 it is so close as does not make a difference. This gives confidence to an assumption that the carbon in the ocean is in sync with the atmosphere.

Deep waters

In this section, the ocean is modeled as 4000 m deep, and with a 2-layered diffusivity. This model is run for a longer time of several years, and therefore includes the increase in CO_2 in the atmosphere. Figure 10 shows the convergence of the simulation, confirming the simulation still converges quadratically.

Figure 11 shows the concentration of CO_2 as a function of time, as given by the simulation. The units of C in this exercise is mol/m^{-3} , so the "mass" given by the formula $M = \int dz C(z)$ is mol per square meter of ocean. Using $0.36 \cdot 10^{15} \text{m}^2$ as the value of the surface area of the global oceans, and 24 g/mol as the molar weight of carbon in CO_2 , the total weight of the carbon stored in the ocean is plotted in Figure 12. This simulation results in a uptake of $\Delta M = 504.8 \cdot 10^{15} \text{g}$ carbon over 10 years, i.e. an average absorption rate of

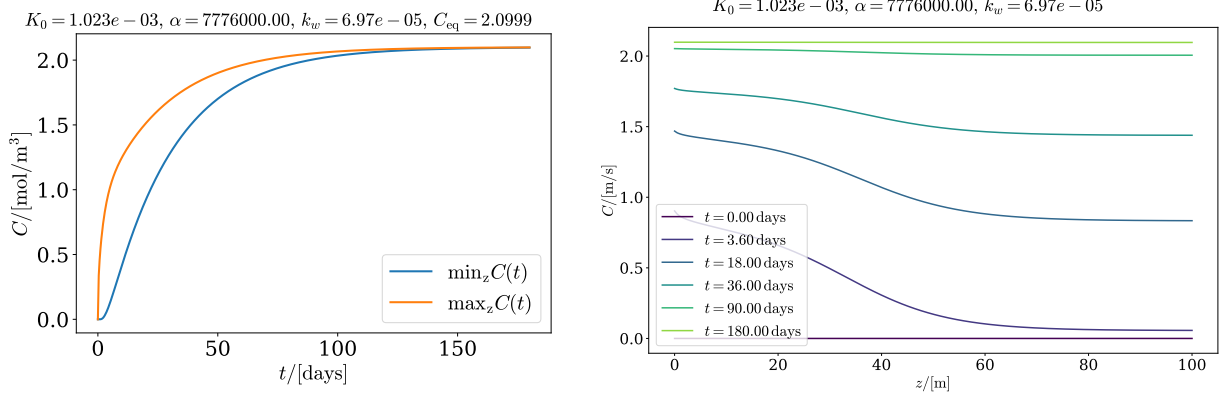


Figure 9: cap

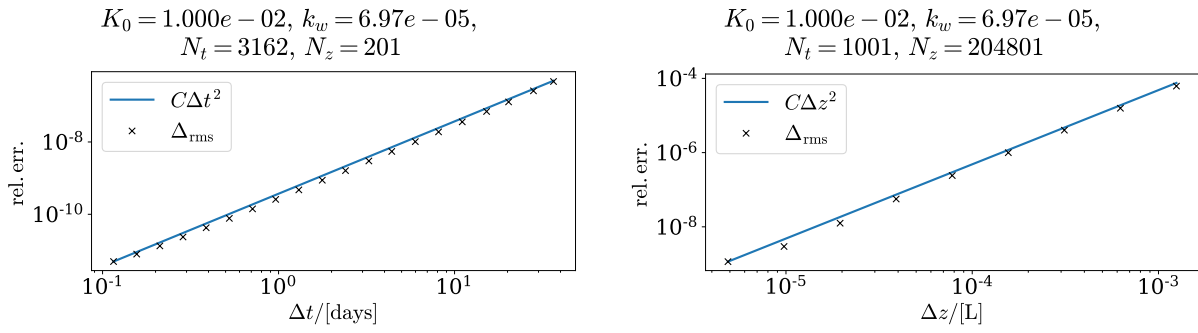


Figure 10: Error, as a function of step length in both time and space.

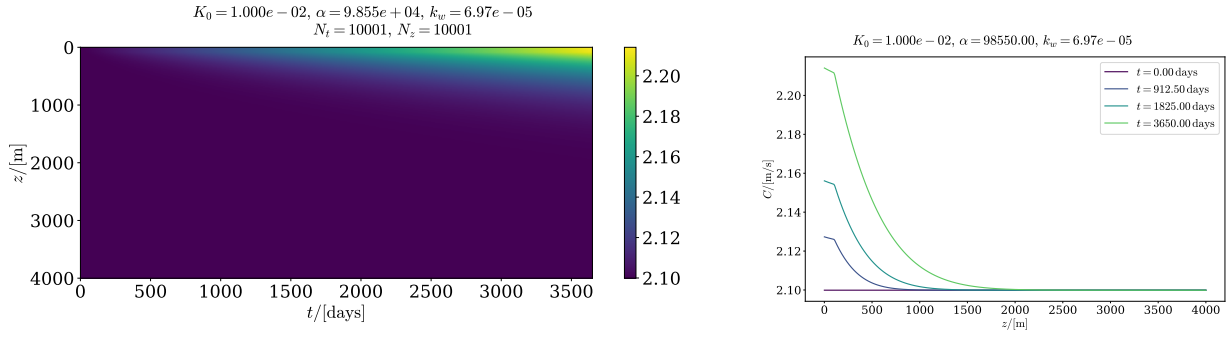


Figure 11: The concentration of CO₂ in as a function of time, over 10 years.

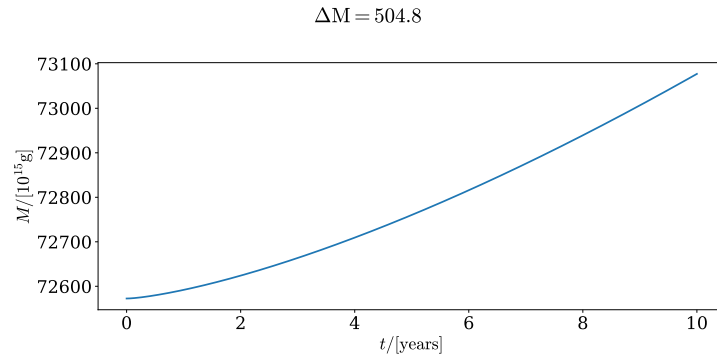


Figure 12: The concentration of CO₂ in as a function of time, over 10 years.

$50.48 \cdot 10^{15} \text{ g/year}$.

References

- [1] NTNU, Institutt for Fysikk. Exercise 3, tfy4235 computational physics, 2021.