Moraine crest or slope: an analysis of the effects of boulder position on cosmogenic exposure age

**Matt D. Tomkins1,2, Jason M. Dortch3, Philip D. Hughes1,2, Jonny J. Huck1, Raimon Pallàs4, Ángel Rodés5, James L. Allard1,2, Andrew G. Stimson1, Didier Bourlès6, Vincent Rinterknecht7,8, Vincent Jomelli8, Laura Rodríguez-Rodríguez 9, Ramon Copons10, Iestyn D. Barr11,2, Chris M. Darvill1,2, Thomas Bishop1**

1Department of Geography, University of Manchester, Manchester, M13 9PL, UK

2Cryosphere Research at Manchester, Manchester, UK

3Kentucky Geological Survey, University of Kentucky, Lexington, USA

4Departament de Dinàmica de la Terra i de l’Oceà, Universitat de Barcelona, 08028 Barcelona, Spain

5Scottish Universities Environmental Research Centre, Rankine Avenue, East Kilbride G75 0QF, UK

6Aix-Marseille Univ., CNRS, IRD, INRA, Coll France, UM 34 CEREGE, Technopôle de l’Environnement Arbois-Méditerranée, BP80, 13545 Aix-en-Provence, France

7Department of Earth and Environmental Sciences, University of St Andrews, Fife KY16 9AL, UK

8Université Paris 1 Panthéon-Sorbonne, CNRS Laboratoire de Géographie Physique, 92195 Meudon, France

9Dpto. Geología, Universidad de Oviedo, Arias de Velasco s/n, 33005 Oviedo, Spain

10Snow and Mountain Research Centre of Andorra (CENMA), Andorran Research Institute (IEA), Sant Julià de Lòria, Andorra

11School of Science and the Environment, Manchester Metropolitan University, Manchester, UK

**Keywords**

Cosmogenic

Moraine

Exhumation

Stabilisation

Schmidt hammer

**Abstract**

Terrestrial cosmogenic nuclide dating (TCN) of ice-marginal moraines can provide unique insights into Quaternary glacial history. However, geomorphic processes can profoundly influence the distribution of TCN ages. Models of moraine evolution predict the greatest ground lowering at moraine crests, but these areas are preferentially sampled due to perceived stability. In this study, we assess the relative utility of moraine crest and moraine slope sampling for a range of moraine types (terminal, lateral, latero-frontal) and moraine ages (Holocene to Last Glacial Maximum) using new and published 10Be and 36Cl TCN ages (*n* = 19) and Schmidt hammer sampling (SH; *n* = 635 moraine boulders, ~19,050 SH *R*-values) of ice-marginal moraines in the Pyrenees mountains, France/Spain. These data show that

These data show that exhumation is widespread for large, matrix-rich and steep-sided moraines deposited at the Last Glacial Maximum, but sampling on the crest is not effective at isolating exhumed boulders. For these moraines, the distribution of exhumed boulders is effectively random while the probability of selecting an exhumed boulder is comparable for moraine crests, ice-proximal and -distal slopes. Within this context, we recommend abandoning the crest-only strategy to widen the population of boulders to select from. However, there is a lack of quantitative evidence linking boulder characteristics to well-clustered TCN datasets; a situation which sets the context for boulder selection “superstition”. Based on our compilation of boulder characteristics, we find that these factors have little explanatory power for the studied moraines, a result which raises questions about the efficacy of geomorphic insight. Instead, it is increasingly clear that landform stability should be prioritised over spatial and boulder criteria, with boulder-rich, matrix-poor moraines more likely to stabilise rapidly after deglaciation. For non-ideal landforms, which are often priority targets for TCN dating, we recommend preliminary SH sampling to identify exhumed boulders and to prioritise individual boulders for analysis.

**Introduction**

Ice-marginal moraines are classic features of glaciated mountain ranges (Buckland, 1840; Penck, 1905) and are prominent terrestrial records of glacial history (Hallet and Putkonen, 1994). By constraining the timing of moraine deposition, it is possible to reconstruct the growth and decay of glaciers and ice sheets through the Quaternary and the palaeoclimatic drivers of glacial cycles (Hays et al., 1976; Broecker and Denton, 1990). Recent developments in terrestrial cosmogenic nuclide (TCN) dating (Nishiizumi et al., 1989; Phillips et al., 1990) have transformed our understanding of Quaternary glaciations by permitting direct analysis of the fragmentary glacial stratigraphic record (Gibbons et al., 1984; Zreda and Phillips, 1995). Despite this progress, TCN dating can be complicated by geomorphic processes which result in pre- or post-depositional exposure of rock surfaces and which account for apparent TCN ages that pre- or post-date the assumed age of the landform (Hallet and Putkonen, 1994; Applegate et al., 2010). While prior exposure is possible (i.e. nuclide inheritance), post-depositional erosion, exhumation and shielding have been shown to profoundly influence TCN age distributions (Shanahan and Zreda, 2000; Briner et al., 2005; Zech et al., 2005; Hein et al., 2011; Heyman et al., 2011; Dortch et al., 2013; Murari et al., 2014).

To avoid geomorphic outliers, researchers typically select samples based on the depositional context and characteristics of individual surfaces (Akçar et al., 2011; Rinterknecht et al., 2014; Palacios et al., 2019). While these criteria have good qualitative reasoning, many have not been tested quantitatively. Previous studies have advocated sampling:

* boulders on moraine crests (Akçar et al., 2011) or on flat, stable surfaces (Gosse et al., 1995),
* the tallest boulders, to minimise the likelihood of post-depositional shielding (Heyman et al., 2016),
* the largest boulders (Akçar et al., 2011) or boulders embedded in the moraine matrix (Ivy-Ochs et al., 2007), to minimise the likelihood of post-depositional instability,
* well-rounded boulders which preserve evidence of glacial transport (Darvill et al., 2015a; Hughes et al., 2016), to minimise the likelihood of pre-depositional exposure,
* hard, “fresh” looking boulders or boulders which elicit a ping instead of a thud when struck with a hammer (*see* Dortch et al., 2010).

Of these, boulder location is critical, as moraine crest boulders are prioritised, while moraine slope boulders are rejected, irrespective of their individual characteristics. It should be noted that this list is not exhaustive and additional criteria may be applied in other settings (e.g. wildfire spalling; Kendrick et al., 2016).

However, there is a paucity of quantitative research correlating boulder characteristics with tightly clustered TCN datasets (e.g. χ2 ≤ 1), and with unclear results generated though ad hoc analysis (Dortch et al., 2010). Although a statistically significant correlation was found between increasing boulder height and dataset clustering (Heyman et al., 2016), the overall effect was minor, as a dominant fraction (>50%) of tall boulder groups were sufficiently scattered to fail a reduced chi square test (χ2 ≤ 2). More fundamentally, the influence of landform characteristics is often overlooked in ad hoc analysis (e.g. moraine sedimentology; Zreda et al., 1994; Putkonen and O’Neal, 2006). Within this context, further work is required to test existing criteria for TCN sample selection and to develop quantitative methods which minimise and account for geomorphic processes (Dortch et al., 2020; Ward et al., 2020). These developments have the potential to significantly improve the robustness of TCN datasets and the chronological utility of the moraine record (Applegate et al., 2012; Dortch et al., 2013).

In turn, this paper focuses on a fundamental component of TCN sample selection; the effect of moraine crest sampling on boulder exposure age. Early numerical models of moraine evolution predicted the greatest ground-lowering at moraine crests (Hallet and Putkonen, 1994; Putkonen and Swanson, 2003) with a period of maximum instability as glaciers retreat and as oversteepened ice-proximal slopes erode and stabilise (Porter and Swanson, 2008). However, moraines continue to degrade through time as a function of moraine height (Putkonen and Swanson, 2003; Putkonen et al., 2008; Schaller et al., 2009), as slope diffusion removes fine-grained material from moraine crests and deposits material at the base of moraine slopes (Applegate et al., 2010). Over time, these processes drive exhumation of boulders which have been shielded from cosmogenic exposure. In turn, the age distribution of moraine crest boulders may primarily reflect an initial stabilisation phase (~1 ka; Briner et al., 2005; Dortch et al., 2010), modified by the ongoing process of moraine degradation, rather than the timing of initial moraine deposition (Putkonen and O’Neil, 2006). Despite this limitation, moraine crests are preferentially sampled for TCN dating (Hallet and Putkonen, 1994). In contrast, slope diffusion models and lichenometric methods predict relative stability on moraine slopes (Hallet and Putkonen, 1994; Putkonen and O’Neil, 2006), but these are rarely sampled for TCN, in part due to the perceived risk that boulders may rotate, shift or roll throughout the lifetime of the moraine. This dichotomy between model predictions and sampling procedures raises a fundamental and currently unanswered question: are moraine crests more informative than moraine slopes?

To address this uncertainty, this paper utilises new and published 10Be and 36Cl TCN ages (*n* = 19) and Schmidt hammer sampling (SH; *n* = 645 moraine boulders, ~19,050 SH *R*-values) of ice-marginal moraines in the Pyrenees mountains, France/Spain, to assess the relative utility of moraine crest and moraine slope sampling for a range of moraine types and moraine ages.

**Methods**

**Moraine selection**

Five moraines of varying age and geomorphology were selected in the Pyrenees (Fig. 1); a mountain range which was extensively glaciated during Quaternary glacial stages (*see* Fig. 1F; Calvet et al., 2011; Delmas et al., 2015; Oliva et al., 2019) but is now on the verge of total deglaciation (Marti et al., 2015). Selected moraines encompass the primary deglaciation phases of the Pyrenees since the global Last Glacial Maximum (gLGM) and feature large populations of granitic moraine boulders sourced from the Hercynian Axial Zone (*see* Fig. 1E; 2A; Crest et al., 2017). While this focused approach does not comprise *all* moraine types or depositional settings (Barr and Lovell, 2014), these sites do encompass a range of moraine types commonly found in cirque and valley landsystems and which are often priority targets for TCN dating (i.e. ≤ gLGM). In contrast, for moraines with long post-depositional histories, alternative techniques are often more appropriate (e.g. Hughes et al., 2007; Lewis et al., 2009; Darvill et al., 2015b), although TCN dating has proved invaluable at sites where long-term moraine denudation is minimal (e.g. Owen et al., 2006; Pallas et al., 2010; Morgan et al., 2011; Dortch et al., 2013; Dietsch et al., 2015; Balter et al., 2020).

Selected moraines include both left and right latero-frontal moraines in the Arànser catchment, Cerdanya (Fig. 3A, 2B). These moraines are large, matrix-rich, steep-sided (~30 - 40°), heavily forested (~100% cover), and record the maximum ice extent (MIE) of the Arànser glacier during the Würmian glacial stage (11.7 - 110 ka; Calvet et al., 2011). The right latero-frontal moraine has previously been dated using 36Cl (*n* = 2; Palacios et al., 2015). Here, we present the results from a further 10 10Be dated boulders (Table 1). On the north side of the Pyrenees, lateral moraines were selected in the Gave de Pau catchment, Hautes-Pyrénées (Fig. 3B, 2A; Jalut et al., 1992; Calvet et al., 2011). At least two proximal (~60 m) but distinct lateral moraine ridges have been identified (Soum d’Ech moraines), with the outer moraine previously dated using 10Be (*n* = 4; Rodés, 2008). As at Arànser, these moraines likely correspond to the Würmian MIE but their distinctive morphologies (multiple nested ridges *vs*. a single large moraine) likely reflects a topographic control on moraine deposition (open topography *vs*. confined valley; Palacios et al., 2015).

On the south side of the Pyrenees, and in the Val de Molières catchment of the Noguera Rigaborçana, sampled sites include the boulder-rich, matrix-poor Outer Pleta Naua terminal moraine (Fig. 4B, 2C), previously assigned to the Younger Dryas chronozone based on 10Be (*n* = 3; Pallàs et al., 2006), and the Tallada cirque moraine (Fig. 4A, 2C), which consists of a single sharp-crested, arcuate terminal moraine with two minor ice-proximal ridges. Although undated, the Tallada moraine is assumed to be late-Holocene in age based on its elevation (~2400 m), topographic setting (small enclosed cirque; ~0.16 km2), aspect (NNE) and the presence of a permanent snowfield (~0.03 km2; Pallàs et al., 2006).

**Sampling approach**

To investigate the depositional and post-depositional histories of these moraines, glacial boulders were selected to cover the entire moraine surface, including the moraine crest (C), the inner ice-proximal slope (IS) and the outer ice-distal slope (OS). In turn, boulder selection was primarily motivated by spatial location and the construction of a dense matrix of sampling points, rather than individual boulder characteristics (Fig. 1). The number of selected boulders varied as a function of moraine size and both boulder density and distribution (*n* = 60 - 275; Table 1) with typical patch sizes (mean area per sampled boulder) of ~14 m2 to ~46 m2. Each boulder was sampled using the Schmidt hammer (SH) to assess the relative degree of weathering following the sampling approach of Tomkins et al. (2018a). All boulders were of sufficient size (Sumner and Nel, 2002) and sampled areas were free of surface discontinuities (Williams and Robinson, 1983) and lichen (Matthews and Owen, 2008). SH R-values were recorded perpendicular to the tested surface to reduce the risk of frictional sliding of the plunger tip (Viles et al., 2011), with single impacts separated by at least a plunger width (Aydin and Basu, 2005). 30 R-values were recorded for each boulder by a single operator (MT) and no outliers were removed following Niedzielski et al. (2009). Schmidt hammer functioning was assessed regularly using the Proceq test anvil (Aydin and Basu, 2005), with instrument calibration performed following the guidelines of Dortch et al. (2016) and Tomkins et al. (2018a). In total, 645 moraine boulders were sampled and ~19,050 SH R-values were generated. To compliment these data, a wide range of boulder characteristics were recorded including dimensions, morphology (angularity, sphericity), surface characteristics (vegetation, fracturing) and depositional context (slope angle, matrix). All data are provided in the Supplementary Information (S1-S2).

These data were analysed based on the assumption that the degree of weathering, as represented by SH R, is inversely correlated with the exposure age of the rock surface (Goudie, 2006). However, this assumption is only valid in the absence of lithological variation between tested rock surfaces (McCarroll, 1989; 1991; Winkler, 2005; Tomkins et al., 2016). Crucially, however, *intra*-landform variability in rock type is absent, as all sampled boulders share a common source area (*see* Fig. 2). Moreover, *inter*-landform variability in rock type is likely minimal, as all sampled boulders were coarse- to medium-grained granites and granodiorites sourced from the Hercynian Axial Zone (Crest et al., 2017). As granitic lithologies have provided particularly effective for calibrated-relative age dating (*see* Engel et al., 2007; 2011; Tomkins et al., 2016; 2018a, 2018b; Wilson et al., 2019; Zasadni et al., 2020), SH R is used here as a proxy for exposure age based on a 10Be-SH calibration dataset developed by Tomkins et al. (2018b). This dataset comprises 52 10Be TCN ages obtained from granite and granodiorite glacial boulders and glacially-sculpted bedrock from across the central and eastern Pyrenees (Pallàs et al., 2006; 2010; Delmas et al., 2008; Crest et al., 2017) and their corresponding SH R values (Fig. 5A; Tomkins et al., 2018b). To further develop this dataset, two additional 10Be dated surfaces from the Val de Molières (Fig. 2C; MUL01 and MUL03; Pallàs et al., 2006) were sampled using the Schmidt hammer (*see* Supplementary Information; Table S1).

However, surface weathering is not only a function of exposure age and lithology, as the rate and style of weathering may be modified by climate (*see* Small et al., 1997; Riebe et al., 2001; 2004; Kuhlemann et al., 2008; 2009; Portenga and Bierman, 2011; Marrero et al., 2018). Crucially, however, *intra*-landform variability in climate is absent as all boulders on an individual landform share a common precipitation and temperature regime. While climate variability at the *inter*-landform scale is evident, for example between the Atlantic and Mediterranean-influenced climates of the northern and southern Pyrenees respectively, the strong correlation between exposure ages and R values for the calibration dataset (R2 = 0.96, *p* < 0.01; Tomkins et al., 2018b), which comprises samples obtained from a range of elevations, catchments and topographic settings (*see* Fig. 5A), indicates that regional climate variability modifies rather than obscures the signal of exposure.

**Calculating calibrated boulder exposure ages**

To utilise this existing correlation, 10Be ages analysed by Tomkins et al. (2018b), in addition to the two new samples described above, were normalised using the CRONUS Earth Web Calculator (Version 2.0; Marrero et al., 2016, available at: <http://cronus.cosmogenicnuclides.rocks/2.0/>), the latest production rates (Borchers et al., 2016), the time-dependent Lm scaling scheme (Lal, 1991; Stone, 2000) and assuming 0 mm ka-1 erosion. To ensure consistency, all 10Be and 36Cl TCN ages discussed in this paper have been recalibrated using these input parameters (full sample details used for exposure age calculation are provided in the Supplementary Information and are available on GitHub: <https://github.com/matt-tomkins/moraine-paper-2020>). In turn, a TCN-SH calibration curve was constructed using logarithmic orthogonal distance regression (ODR, Boggs and Rogers, 1990a); an approach which improves on ordinary least squares regression by minimising orthogonal residuals to account for measurement uncertainties in both the independent and dependent variables. This unweightedapproach returns prediction estimates (1σ) of ± 1.6 - 1.8 ka (calculated using the ODR covariance matrix; Boggs and Rogers, 1990b), with a distribution that corresponds to the empirical rule (~70% of calibration data within 1σ interval, ~96% within 2σ, 100% within 3σ).

However, to produce wider (i.e. more tolerant) prediction intervals which explicitly incorporate measurement errors (*see* Viles et al., 2011; Jull et al., 2015), we utilise Monte Carlo simulated datasets (*n* iterations = 104) in which input variables are randomised based on their associated uncertainties (TCN ± external age uncertainty; SH ± Standard Error of the Mean). This approach is preferable to a weighted ODR which requires unnecessary assumptions regarding weighting constants and is biased by TCN age-uncertainty collinearity (*as* age ↗, uncertainty ↗; Ivy-Ochs et al., 2007; Dortch et al., 2020). Our analytical procedure, which returns wider prediction estimates (1σ) of ± 2.0 - 2.2 ka, is described fully in the Supplementary Information (S4) and has been implemented on SHED-Earth (<http://shed.earth>), an online calculator developed to enable wider and more consistent application of our approach (Tomkins et al., 2018a).

Based on this calibration curve, mean R-values from the 635 sampled boulders were converted into *calibrated exposure ages* through interpolation. While uncertainty estimates for individual calibrated boulder exposure ages are larger than typical uncertainties associated with individual TCN exposure ages, landform age estimates can be of comparable precision to established techniques when derived from large SH datasets (e.g. *n* boulders ≥ 30; Tomkins et al., 2018b; 2018c) and when appropriate statistical approaches for outlier identification and error propagation are employed (Applegate et al., 2012; Dortch et al., 2013; 2020). In addition, and to avoid circularity, it is important to note that all of the 54 10Be calibration surfaces utilised here are independent of the studied moraines.

**Calculating landform ages**

To determine the timing of moraine deposition at each site, we analysed calibrated boulder exposure ages using the Probabilistic Cosmogenic Age Analysis Tool (P-CAAT Version 1.0; Dortch et al. 2020). This method utilises non-linear curve fitting and a Monte Carlo style approach to isolate component Gaussian distributions to account for positive (prior exposure) and negative skew (incomplete exposure) of age datasets (*see* Dortch et al. 2020 for full details). For each landform, we selected the highest probability component Gaussian distribution to represent the age of the landform (Dortch et al., 2020). The results of this analysis are presented in Fig. X and Table X. To assess the validity of these landform ages, we compared these data to previously published 10Be and 36Cl ages (*n* = 9; Pallàs et al., 2006; Rodés, 2008; Palacios et al., 2015) and present 10 new 10Be ages from the Arànser catchment (Table X; Supplementary Information S4).

Based on landform age analysis (Table 2), individual boulders were sorted into “good” and “bad” groups, which include interpolated boulder ages which are within or outside the 2σ (95%) age boundaries of the landform age respectively. Selection of a broad 2σ threshold is appropriate given the uncertainties associated with SH sampling (Aydin and Basu, 2005; Viles et al., 2011), in addition to the systematic (Jull et al., 2015; Borchers et al., 2016) and geologic uncertainties (Hallet and Putkonen, 1994) inherited from TCN dating. In addition, the use of logistic analysis is also appropriate because the magnitude of the deviation from the landform age is typically not of interest. For example, inherited boulders which return ages 50% and 300% greater than the landform age are numerically distinct, but are both “bad” if the objective was to determine the timing of landform deposition. In turn, “good” boulders are interpreted to reflect the timing of moraine deposition or initial stabilisation, while those classed as “bad” are likely compromised by pre- or post-depositional exposure.

**Spatial analysis**

Finally, the spatial distribution of “good” and “bad” boulders was analysed using global and local Moran’s *I* spatial autocorrelation (Moran, 1950) based on Queens contiguity of Voronoi cells (polygons that share either an *edge* or a *vertex*; *see* Fig. X). At the global level, Moran’s *I* is used to assess whether the overall clustering of the data is significantly different from a random distribution. For datasets that are non-random (*p* < 0.05), local Moran’s *I* is used to identify the location of statistically significant boulder clusters. Current sampling approaches are based on the qualitatively-sound but often quantitatively-untested assumptions that (i) the distribution of “good” boulders is non-random and that (ii) “good” clusters are more likely on moraine crests. These assumptions can be explicitly tested for the studied moraines using global and local Moran’s *I* respectively.

**Uncertainty propagation**

The above analyses provide information on the relative occurrence and spatial clustering of “good” and “bad” boulders for a range of moraine types and ages. However, each individual calibrated boulder exposure age has an associated uncertainty which must be propagated fully. To achieve this, the above analyses have been automated using a Monte Carlo style approach as follows:

1. For each landform, calibrated exposure ages were simulated using a truncated normal distribution (*n* = 104, *minimum* = 0, *maximum* = ∞; Mersmann et al. 2018) using calibrated exposure ages and their 1σ prediction bounds (± 2.0 - 2.2 ka) as input parameters. These simulations follow the empirical rule, where ~68% of simulated ages fall within the 1σ prediction bounds, ~95% within the 2σ prediction bounds and ~99.7% within the 3σ prediction bounds.
2. Using these simulated data, we utilised an automated version of P-CAAT (Dortch et al., 2020) to calculate a landform age for each simulated dataset, using the same model parameters for consistency.
3. Using these new landform ages, simulated ages were sorted into “good” and “bad” groups based on the 2σ (95%) age boundaries of the landform age. The relative proportion (%) of these groups was recorded.
4. Finally, the spatial distribution of “good” and “bad” ages for each simulated dataset was assessed using global and local Moran’s *I*, as described previously. For each model run, and when global Moran’s *I* revealed a statistically significant non-random distribution (*p* < 0.05), local Moran’s *I* was used to identify “good” and “bad” clusters. The locations of these clusters was recorded with cluster occurrence presented as a proportion (%) of the total number of model runs (*n* = 104).

In turn, these analyses shed light, in a statistically robust manner, on the relative utility of moraine crest or moraine slope sampling for the studied moraines.

**Results**

**Calibrated boulder exposure ages**

There is a very strong correlation between recalibrated 10Be ages and SH R-values, based on a large calibration dataset (*see* Fig. 5; *n* = 54; Tomkins et al., 2018b). Of the 19 36Cl and 10Be samples from the studied moraines, 15 were located and re-sampled with the SH (*see* Supplementary Tables 1-2). Of these, the vast majority closely match the existing calibration dataset (*n* = 13, *see* Fig. 5). These observations indicate that when lithological variation is minimised, and when a consistent sampling approach is adopted, the relative degree of rock surface weathering can be used as a proxy for surface exposure age.

The exceptions to this are samples ECH-03 (17.2 ± 3.5 ka) and ECH-04 (16.8 ± 3.3 ka) from the Soum d’Ech moraines (Rodés, 2008) which are significantly more weathered (~38 R) than their corresponding TCN ages would predict (~47 R). This difference is attributed to sub-surface weathering in the near-surface soil zone prior to boulder exhumation at ~17 ka, potentially as a function of the morphological properties of the rock surfaces (Chartres and Walker, 1988) or variation in soil moisture content (Dixon et al., 2006). However, the scale of this influence is unlikely to be universal given the close correspondence between sample ECH-01 (19.7 ± 3.6 ka) and the existing calibration dataset (*see* Fig. 5). While sub-surface weathering of boulders under thin soil cover (~25 cm) can occur (Allen, 2002; Darmody et al., 2005; 2008), boulders are often protected from weathering by sediment burial, as evidenced by the emergence of unweathered boulders from glacial tills (Birkeland, 1999) and alluvium (Ehlmann et al. 2008). In turn, while the influence of sub-surface weathering cannot be excluded as an additional contributory factor, subsequent analysis of the Soum d’Ech moraines is based on an assumed correspondence between weathering (SH R) and surface exposure.

**Landform ages**

Landform ages calculated using P-CAAT, in addition to associated model fits (Dortch et al. 2020), are reported in Table 3. Based on this approach, latero-frontal moraines in the Arànser catchment were deposited at 23.3 ± 1.1 ka (left) and 22.3 ± 0.9 ka (right). As these estimates are consistent within measurement uncertainties, and given the comparable stratigraphic positions of these deposits, we consider moraine deposition to be contemporaneous. No independent dating evidence is available for the left lateral moraine, but 12 TCN ages are now available for the right lateral moraine (36Cl, *n* = 2, Palacios et al., 2015; 10Be, *n* = 10). The distribution of these data is skewed (0.67) and non-normal (Shapiro-Wilk test, *W* = 0.90, *p* = 0.14) but the oldest sample (SAL-10; 21.6 ± 1.8 ka) is consistent with both landform ages within measurement uncertainty. The selection of the oldest moraine age accords with standard procedures for interpreting moraine ages as minimum limiting ages (Putkonen and Swanson, 2003; Briner et al., 2005).

In the Gave de Pau catchment, calibrated boulder exposure ages from the proximal Soum d’Ech lateral moraine ridges return landform ages of 25.9 ± 1.0 ka (outer, *n* = 61) and 26.2 ± 0.8 ka (inner, *n* = 39). While these moraines are stratigraphically distinct, they cannot be statistically distinguished. In turn, it is possible that moraine deposition occurred within the resolution of our sampling approach, or that differences in moraine age have been masked by moraine stabilisation, degradation or sub-surface boulder weathering. As the temporal distribution of calibrated boulder ages is near identical (Table 3; Fig. 6), we assign these deposits a combined landform age of 26.1 ± 1.3 ka (), with uncertainties propagated through summation in quadrature, and with subsequent analyses performed on the combined dataset for computational ease.

While it must be noted that this estimate is significantly older than the corresponding 10Be ages (*n* = 3; Rodés, 2008), limiting 14C ages obtained from a proximal palaeolake sediment sequence at Lac de Lourdes (~2 km) appear consistent with an older age for moraine deposition. While the oldest radiometric 14C ages from this over-deepened glacial basin are now considered suspect due to contamination from mineral carbon *(see* Reille and Lowe, 1993; Reille and Andrieu, 1995; Pallàs et al., 2006), a younger AMS 14C age from glaciolacustrine clays suggests initial ice-free conditions by 22.1 ± 0.4 14C ka BP, while an AMS 14C age from overlying gyttja supports the culmination of glaciolacustrine sedimentation and deglaciation of the lower Gave de Pau by 16.8 ± 0.3 14C ka BP (Reille and Andrieu, 1995). Based on these data, the younger ages from Soum d’Ech can be considered suspect (ECH-03, -04). While the oldest sample (ECH-01; 19.7 ± 3.6 ka) is closer to the assumed age of deglaciation, it appears likely that this exposure age is representative of final moraine stabilisation, rather than initial deposition.

In the Val de Molières catchment of the Noguera Rigaborçana, recalibrated 10Be ages on the Outer Pleta Naua moraine range from 12.6 ± 1.5 ka to 13.2 ± 1.6 ka (*n* = 3; Pallàs et al., 2006) and overlap within measurement uncertainties with the SH R-derived landform age (*n* = 60; 12.4 ± 0.3 ka). Finally, the Tallada cirque moraine returned a landform age of 3.4 ± 0.7 ka. While this estimate cannot be independently verified, the limited weathering of the moraine boulders (SHR ≥ 60), in combination with the topographic setting of the Tallada cirque, appears consistent with a late-Holocene origin. This estimate should probably be regarded as a maximum age for cirque deglaciation at this site as glacier re-growth or expansion during the Little Ice Age is possible (Hughes, 2018). Minor moraines occur inbound of the large arcuate terminal (*see* Fig. 4A), with evidence of incision of these deposits on their ice-proximal slopes, potentially by glacial erosion or rock/snow avalanche activity. However, the absence of unweathered glacially-modified boulders (e.g. SH R ≥ 70) on the outer moraine suggests that any subsequent glaciers were minimally erosive.

**Temporal analysis**

While estimated landform ages are generally consistent with independent TCN ages and match wider Pyrenean glacier chronologies (Calvet et al., 2011; Delmas, 2015; Oliva et al., 2018), the age distribution of calibrated boulder ages varies significantly between the sampled moraines (*see* Fig. 6). The Arànser and Soum d’Ech moraines exhibit strong negative skew (Table 3), in line with exhumation models (Applegate et al., 2012), while Tallada is normally distributed with a slight positive skew (Shapiro-Wilk test, *W* = 0.98, *p* = 0.56); a trend which may reflect prior exposure or reworking of glacial material (Applegate et al., 2010). In contrast, boulder ages on the Outer Pleta Naua moraine are normally distributed (*W* = 0.96, *p* = 0.07) and tightly clustered (IQR = 0.55 ka).

In light of these trends, the proportion of “good” and “bad” boulders varies between the sampled moraines. Based on the Monte Carlo style approach described in Section X, the proportion of “good” boulders is highest on the Outer Pleta Naua moraine (98 ± 4%) and lowest on the Arànser left (56 ± 4%) and Arànser right moraines (49 ± 4%). For the older moraines (~gLGM), most “bad” boulders are younger than the assumed age of deglaciation, while the Tallada moraine contains a small but significant component of boulders which are older than the assumed age of deglaciation (12 ± 4%).

**Spatial analysis**

Any trends

and exhibit no significant spatial clustering (Global Moran’s *I* = -0.023, *p* = 0.152). These data appear consistent with rapid post-depositional stabilisation. Statistically significant spatial clustering is absent at the Tallada moraine (*I* = 0.111, *p* = 0.055) but is clearly identified at both Arànser (*I* = 0.101, *p* = 0.004) and the Gave de Pau moraines (*I* = 0.112, *p* = 0.027) which may reflect a combination of post-depositional processes which drive moraine evolution.

However, the spatial distribution of boulder ages does not conform to either a crest or slope stability model (Fig. 4; Hallet and Putkonen, 1994), with no consistent trends between or within individual moraines. The distribution of ages is comparable between moraine groups (Table 3; Fig. 5) while the distribution of statistically significant clusters at the Arànser and Gave de Pau moraines is effectively random, as “good” and “bad” clusters occur on moraine crests, ice-proximal and -distal slopes. There are minor exceptions to this rule, where statistically significant boulder clusters have plausible geomorphological explanations. At Arànser, a cluster of high *R* exhumed boulders occurs near the front of the moraine, which likely reflects fluvial incision of the terminus and degradation of the lateral flanks (e.g. Duran valley; Palacios et al., 2015b). An additional exhumed cluster occurs where a minor stream has cross-cut and incised the moraine crest (Fig. 2C, 4A); this is likely a major meltwater source from the other moraines coming out of the other cirque!

**Boulder characteristics**

~~Based on MANOVA, there is a significant difference in boulder characteristics between “good” and “bad” groups at the population level (~~*~~p~~* ~~< 0.01). Logistic regression reveals four significant predictors at~~ *~~p~~* ~~= 0.05. These include the characteristics of the underlying matrix (~~*~~p~~* ~~= 0.046; 0 = matrix-supported, 0.5 = some combination of matrix/clast-support, 1 = clast-supported), as sampling a clast-supported boulder increases the probability of selecting a “good” boulder by a factor of 1.70 (95% CI: 1.01 - 2.90). Boulder angularity is also significant (~~*~~p~~* ~~< 0.01; Powers, 1953), as sampling a rounded over an angular boulder influences the probability by a factor of 0.30 (95% CI: 0.16 - 0.54). The remaining predictors relate to the depositional environment, as sampling a boulder on the ice-proximal slope rather than the moraine crest influences the probability by a factor of 0.50 (~~*~~p~~* ~~= 0.021; 95% CI: 0.28 - 0.90). Finally, the height of the boulder relative to the moraine crest is also significant (~~*~~p~~* ~~< 0.01) and influences the probability by a factor of 0.31(95% CI: 0.13 - 0.71). In summary, matrix-supported, rounded boulders, deposited near crests but on ice-proximal slopes have a higher probability of returning a “bad” age at the population level.~~

~~It is important to note, however, that these factors are likely inherited from landform characteristics, as well-clustered boulders on the Outer Pleta Naua moraine are predominantly clast-supported (56/60) and angular (33/60). In turn, this landform, and the characteristics of its boulders, has a disproportionate effect on population level analysis. When these analyses are repeated at the landform level, MANOVA reveals no significant differences between groups at~~ *~~p~~* ~~= 0.05 for the Gave de Pau (~~*~~p~~* ~~= 0.17), Outer Pleta Naua (~~*~~p~~* ~~= 0.98) and Tallada moraines (~~*~~p~~* ~~= 0.23). Significant differences are recorded at Aranser (~~*~~p~~* ~~= 0.018) but the identified predictors are inconsistent with population level results. Significant predictors (~~*~~p~~* ~~= 0.05) include the presence of surface vegetation (~~*~~p~~* ~~= 0.008; Factor = 7.99, 95% CI: 1.78 - 39.20), the degree of fracturing (~~*~~p~~* ~~= 0.004; Factor = 4.35, 95% CI: 1.64 - 11.99), the underlying slope angle (~~*~~p~~* ~~= 0.002; Factor = 0.94, 95% CI: 0.91 - 0.98) and the height of the boulder relative to the moraine crest (~~*~~p~~* ~~= 0.019; Factor = 0.02, 95% CI: 0.001 - 0.49). For this moraine, fractured boulders with high surface vegetation cover are more likely to return “good” ages, while increases in slope angle and relative height have a negative effect. This dichotomy between population and landform level results highlights the challenge of untangling the respective influences of boulder and landform characteristics. In turn, the most striking result is that no predictor has consistent explanatory power across moraine groups and at the population level. This analysis points to a decoupling of boulder characteristics and exposure ages and raises questions about the efficacy of geomorphic insight (Akçar et al., 2011).~~

**Discussion**

Accurate dating of ice-marginal moraines is critical to our understanding of Quaternary glacial history and the climatic drivers of glacial cycles (Hallet and Putkonen, 1994). Within this context, efforts to minimise geomorphic bias of moraine TCN datasets may significantly improve the utility of the moraine record. While slope diffusion models predict significant boulder exhumation at moraine crests and relative stability on moraine slopes (Hallet and Putkonen, 1994; Putkonen and Swanson, 2003; Applegate et al., 2010), crests are preferentially sampled for CRE dating, while moraine slopes are generally avoided due to increased risks of boulder instability.

In this study, we have critically assessed these two competing ideas and found that spatial criteria (i.e. boulder position) cannot explain the distribution of boulder ages and, in the absence of detailed geomorphological assessment (e.g. Palacios et al., 2019), should not be used to guide CRE sampling strategy. We found no consistent trends between or within individual moraines while the likelihood of selecting an exhumed boulder is comparable for moraine crests, ice-proximal and -distal slopes (Table 3). In general, however, exhumed boulders are far more likely for steep-sided, matrix-rich moraines, as these are particularly susceptible to slope diffusion, while fluvial incision can explain both lateral degradation near termini and cross-cutting by streams. Importantly, moraine sedimentology appears to place a key control on post-depositional stability (Zreda et al., 1994; Putkonen and O’Neal, 2006), as age distributions for matrix-rich moraines are strongly negatively skewed, implying prolonged and continuous boulder exhumation after deposition. The most significant exhumation signal is evident at Aranser (*IQR* = 7.8 ka) and this may be partially explained by forest growth and boulder toppling (Ivy-Ochs et al., 2007); a process which may account for a similar pattern of exhumation for maxima moraines in the neighbouring Têt catchment (MIE Unit; Samples A1-A3 and B1-B3; Delmas et al., 2008). In contrast, the boulder-rich, matrix-poor Outer Pleta Naua moraine stabilised rapidly after glacial retreat, in line with similar well-dated moraines in the SE Pyrenees (Pallas et al., 2010) and the Alps (Ivy-Ochs et al., 2007). In the absence of a supporting sediment matrix, boulder-rich moraines stabilise quickly and appear less susceptible to subsequent erosion. Finally, for young cirque moraines, reworking of older glacial deposits appears more significant than exhumation (Table 3; Tallada), in line with previous studies (Heyman et al., 2011). In these environments, the age of the oldest boulder may overestimate the “true” age of the moraine (Putkonen and Swanson, 2003; Briner et al., 2005).

**Boulder characteristics**

If spatial criteria cannot be used to isolate exhumed boulders, then the depositional context and characteristics of individual boulders may have stronger explanatory power. Much attention has been devoted to choosing the “best” boulders based on geomorphic expertise (Akçar et al., 2011; Rinterknecht et al., 2014). Previous studies have advocated sampling (i) the tallest boulders, to minimise the likelihood of post-depostional shielding (Heyman et al., 2016), (ii) the largest boulders, to minimise post-depositional instability (Akçar et al., 2011), (iii) those on flat, stable surfaces (Gosse et al., 1995) or on moraine crests (Akçar et al., 2011), (iv) boulders embedded in the moraine matrix (Ivy-Ochs et al., 2007), (v) hard, “fresh” looking boulders, (vi) boulders which elicit a ping instead of a thud when struck with a hammer (*see* Dortch et al., 2010) and (vii) rounded boulders with evidence of glacial transport, as subglacial erosion may have removed the accumulated in-situ cosmogenic signal (Hughes et al., 2016). In contrast, angular boulders may have originated from rockfalls, and through supraglacial transport to glacier margins or direct deposition onto moraines (Porter and Swanson, 2008), may retain a cosmogenic signal from prior exposure. However, there is a paucity of quantitative research correlating boulder characteristics (e.g. i-vii) to tightly clustered CRE datasets (e.g. χ2 ≤ 1), with unclear results generated though ad hoc analysis (Dortch et al., 2010). Although a statistically significant correlation was found between increasing boulder height and dataset clustering (Heyman et al., 2016), the overall effect is minor, as a dominant fraction (>50%) of tall boulder groups were sufficiently scattered to fail a reduced chi square test (χ2 ≤ 2). More fundamentally, any ad hoc approach is inherently flawed as compiled boulder sets were initially collected according to a set of specific criteria (Tylmann et al., 2018), thus precluding a truly representative compilation. Finally, the influence of landform characteristics is often overlooked in ad hoc analysis, but the degree to which the wider depositional context overrides boulder characteristics is not known. These circumstances set the context for boulder selection “superstition”, which equates some boulder characteristics to reliably clustered datasets in the absence of quantitative evidence (Skinner, 1948).

Based on our compilation of boulder dimensions, surface and depositional characteristics (*see* Data Repository), it is clear that these factors have little explanatory power for the studied moraines. While these data do not negate the importance of careful geomorphological assessment of moraine structure (Palacios et al., 2019), the decoupling of boulder characteristics and interpolated exposure ages raises questions about the prioritisation of boulder characteristics in CRE sample selection (Akçar et al., 2011). For moraines that stabilise rapidly (i.e. Outer Pleta Naua), even seemingly non-ideal boulders can return “good” ages (e.g. OPN17, ~0.5 m height, angular, ice-proximal slope, major fractures). Conversely, on moraines with long exhumation histories (i.e. Arànser), even ideal boulders can be exhumed (e.g. AR28, ~8.9 ka younger than landform age, ~1.1 m height, partially embedded on the crest, 7.6° slope). Within this context, we suggest that landform stability should be prioritised for CRE sample selection, with tightly clustered CRE datasets far more likely for rapidly stabilising matrix-poor, boulder-rich moraines (Ivy-Ochs et al., 2007; Pallàs et al., 2010). However, non-ideal landforms are often priority targets for CRE dating, as these often record the retreat stages of the largest glacier systems (e.g. Gave de Pau) or represent “missing” glacial stages in existing chronologies. In turn, quantitative methods to guide CRE sample selection for non-ideal landforms have clear utility for Quaternary glacial research.

**Future directions**

* No penalty to sampling on moraine crests. If there are sufficient good boulders, no problem (however a major challenge identifying these!)
* However, there appears no penalty to sampling moraine slopes - there may be ones that have rolled, shifted etc… but no evidence that this occurs at a faster rate than exhumation of the crest
* There is a potential advantage to widening sampling - currently we’re limiting sampling to a small population. Although density may be higher at the crest (), this is often a small % of the total overall number.
* Clearest message is that “landform stability” is more significant that boulder location - the pattern of “good” vs. “bad” is consistent at the landform scale. In turn, selection of the landform is more critical than the selection of the boulder.
* In this study, the OPN moraine is the best - matrix poor, boulder rich - likely stabilised rapidly. Clustering of 10Be ages and common weathering characteristics of the boulders.
* The worst is both Aranser moraines (skewed, exhumation, boulder toppling from trees, matrix rich).
* However, and given the scope of this project - these recommendations cannot be applied universaslly - focus on matrix poor might lead to rock avalanche sampling (climatic signal?) or even rock glaciers which have there own issues (Crump, Akcar). Alternatively, sampling matrix rich landform does not guarantee a mixed results bag, for example see results from Licciardi on matrix rich, partially forest, laterals in the Sierras. More work is clearly necessary to identify which landform characteristics can be used more equivocally (e.g. Drumlins vs. moraines), or whether a mixed bag is the reality.
* The second is that boulder characteristics alone appear to not have a great deal of predictive power for these moraines. A wide range of metrics, shown to have very poor correlation with the weathering of the boulders (R), and by association their age.
* Finally, SH testing may be the way forward (Tylmann), needs exploring further but importantly it’s not a panacea - it will not gaurentee success - doesn’t work on all rock types (Tomkins, 2016), is sensitive to surface characteristics which are unaffected by TCN, and may be sensitive to other processes which don’t affect the exposure age (e.g. sub-surface weathering). However, evidence suggests that when these limitations are accounted for (e.g. large sampling, consistent approach), there is correlation with TCN.
* Then, pre-sampling with the SH may be effective. For moraines with higher likelihoods of exhumation, pre-sampling may be effective. Sample a statistically robust number of boulders (30-50, ~1 day in the field), and analayse the skew to select a population for dating. We hope that this will imporove clustering.

Our approach, in light of strong regional evidence for a link between SH *R* and exposure ages for granitic surfaces (Engel et al., 2007; 2011; Tomkins et al., 2018a; 2018b), indicates that preliminary SH sampling may be an effective technique for identifying exhumed boulders. Prior to CRE sampling, we recommend collection of statistically viable population of *R* values from a large sample of moraine boulders to refine CRE sampling strategies. Sample size should reflect the relative stability of the studied landform, with large populations required for moraines susceptible to prolonged post-depositional erosion. To assess how this requirement varies between the studied moraines, we randomly sampled each dataset 1000 times (without replacement), selecting *n* samples per model run and used an automated version of the Gaussian decomposition model of Dortch et al. (2013) to determine the associated landform age. We increased *n* for each model run until ≥ 95% of the 1000 simulations returned component Gaussians with a peak age which fell within the 2σ landform age boundaries (95% CI). Based on this approach, the number of required samples scales with the complexity of the underlying age distribution, with a small number required for tightly clustered and normally-distributed moraines (Outer Pleta Naua, *n* = 7) while more are required for moraines compromised by exhumation (Aranser, *n* = 9; Lourdes, *n* = 14). The greatest number of samples are required at Tallada (*n* = 19) which reflects the influence of both pre- and post-depositional exposure. Importantly, as the full age distribution of the studied moraine is unknown before sampling, we recommend sampling ≥ 20 boulders per moraine to account for exhumation or pre-exposure of boulders on non-ideal landforms. However, this value should be regarded as a minimum, as compilation of a larger dataset will widen the population of boulders for CRE selection. In the absence of confounding factors, we argue that SH *R* is an accurate proxy for surface exposure age and should be used routinely to identify exhumed boulders (boulder *R* > modal *R*), to assess the age distribution of the studied moraine (i.e. dataset skewness) and to pre-select boulders for CRE dating (modal *R*).

So what?

**Conclusions**

Intensive SH sampling of ice-marginal moraines in the Pyrenees, in combination with 10Be analysis, provides unique insight into the effects of boulder position on exposure age. We show that large, matrix-rich and steep-sided moraines are particularly susceptible to post-depositional erosion. However, restricting sampling to moraine crests is not effective at isolating exhumed boulders as their distribution is effectively random, and only isolated clusters have clear geomorphological explanations. As such, we recommend abandoning the crest-only strategy to widen the population of boulders to select from. However, we also show that boulder characteristics have limited explanatory power for the studied moraines. In light of this complexity, our data indicate that landform stability should be prioritised over boulder and spatial characteristics. Moraine sedimentology exerts a key control on post-depositional stability, with boulder-rich, matrix-poor moraines stabilising rapidly after deglaciation. In turn, CRE sample size should reflect the relative stability and age of the studied landform, with a large number of samples required in many settings to account for exhumation. Finally, our approach highlights the value of preliminary SH sampling to isolate exhumed boulders and to prioritise boulders for CRE dating, especially for non-ideal landforms. Wider application of this method may improve the chronological robustness of CRE datasets and the value of the moraine record in our understanding of Quaternary landscape evolution.

**Acknowledgements**

MT was funded by a University of Manchester President’s Doctoral Scholar Award. The fieldwork for this paper was supported by a British Society for Geomorphology Postgraduate Research grant awarded to MT. 10Be TCN analysis was supported by X. MT, JD, PH and RP designed the project, with refinements from IB and CD. MT, JA and AS conducted Schmidt hammer sampling. 10Be sampling and analysis was undertaken by RP, AR, DB, VR, VJ, LRR and RC. MT and JH conducted spatial analysis. MT, JH and TB improved error propagation. All authors contributed to the manuscript. We thank Dr. Geoff Evatt and Dr. Andrew Smedley at the University of Manchester for their statistical insights.

|  |  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- | --- |
| **Table 1.** Summary statistics for the sampled moraines | | | | | | | |
| Moraine | Type | # samples | # ISa | # Ca | # OSa | Area (km2)b | Patch sizec |
| Tallada | Terminal | 70 | 16 | 29 | 25 | 0.015 | ~15 m2 |
| Outer Pleta Naua | Terminal | 60 | 20 | 20 | 20 | 0.011 | ~14 m2 |
| Arànser (Left) | Latero-frontal | 275 | 199 | 51 | 25 | 0.421 | ~39 m2 |
| Arànser (Right) | Latero-frontal | 130 | 57 | 23 | 40 | 0.281 | ~46 m2 |
| Soum d’Ech | Lateral | 100 | 37 | 50 | 13 | 0.111 | ~33 m2 |
| a Inner ice-proximal slope (IS), moraine crest (C) and outer ice-distal slope (OS), b Sampling area, c mean area per sampled boulder | | | | | | | |

|  |  |  |  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- | --- | --- | --- |
| **Table 2.** Summary data for terrestrial cosmogenic exposure ages from the sampled morainesa | | | | | | | | | |
| Moraine | Name | Isotope | Latitude (°) | Longitude (°) | Elevation (m) | Age (ka) | Internal ± (ka) | External ± (ka) | SH R± SEMb |
| Outer Pleta Nauac | OPN01 | 10Be | 42.6365 | 0.7399 | 2217 | 13.2 | 1.3 | 1.6 | - |
| OPN02 | 10Be | 42.6365 | 0.7406 | 2197 | 13.0 | 1.7 | 2.0 | 51.68 ± 0.5 |
| OPN03 | 10Be | 42.6365 | 0.7409 | 2195 | 12.6 | 1.2 | 1.5 | - |
| Arànser (Left)d | SAL-01 | 10Be | 42.4283 | 1.6300 | 2000 | 17.6 | 0.6 | 1.5 | 47.57 ± 0.83 |
| SAL-02 | 10Be | 42.4273 | 1.6321 | 1983 | 19.2 | 0.6 | 1.5 | 45.07 ± 0.84 |
| SAL-03 | 10Be | 42.4270 | 1.6326 | 1975 | 21.1 | 0.6 | 1.7 | 44.07 ± 0.82 |
| SAL-04 | 10Be | 42.4254 | 1.6358 | 1933 | 18.0 | 0.6 | 1.5 | 47.57 ± 0.84 |
| SAL-05 | 10Be | 42.4240 | 1.6389 | 1912 | 17.0 | 0.9 | 1.6 | 48.9 ± 0.77 |
| SAL-06 | 10Be | 42.4237 | 1.6395 | 1908 | 19.2 | 0.6 | 1.5 | 44.53 ± 0.74 |
| SAL-07 | 10Be | 42.4229 | 1.6415 | 1896 | 16.7 | 0.5 | 1.4 | 47.43 ± 0.96 |
| SAL-08 | 10Be | 42.4223 | 1.6447 | 1863 | 17.1 | 0.6 | 1.4 | 47.7 ± 0.9 |
| SAL-09 | 10Be | 42.4215 | 1.6481 | 1820 | 20.7 | 0.9 | 1.7 | 44.77 ± 0.8 |
| SAL-10 | 10Be | 42.4213 | 1.6489 | 1808 | 22.4 | 0.7 | 1.8 | 43.03 ± 0.95 |
| PIR-11-13 | 36Cl | 42.4213 | 1.6495 | 1809 | 18.2 | 1.6 | 2.1 | - |
| PIR-11-14 | 36Cl | 42.4209 | 1.6499 | 1805 | 17.3 | 1.7 | 2.2 | 47.6 ± 0.83 |
| Soum d'Eche | ECH01 | 10Be | 43.0863 | -0.0870 | 776 | 19.7 | 3.2 | 3.6 | 42.43 ± 0.98 |
| ECH02 | 10Be | 43.0858 | -0.0880 | 778 | 59.0 | 43.2f | 43.0 | - |
| ECH03 | 10Be | 43.0862 | -0.0873 | 779 | 17.2 | 3.3 | 3.5 | 38.86 ± 1.11 |
| ECH04 | 10Be | 43.0865 | -0.0867 | 781 | 16.8 | 3.0 | 3.3 | 38.77 ± 1.05 |
| a Full sample information used for exposure age calculation is provided in the Supplementary Information or is available on GitHub: <https://github.com/matt-tomkins/moraine-paper-2020>, b Mean of 30 SH R-values ± the Standard Error of the Mean, c OPN samples from Pallàs et al. (2006), d PIR samples from Palacios et al. (2015), e ECH samples from Rodés (2008), f Measurement error, *see* Rodés (2008). | | | | | | | | | |

|  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- |
| **Table 3.** Age statistics for the sampled moraines | | | | | | |
| Moraine | Group | Bandwidtha | Model fitb | Age (ka)c | IQRd | Skew |
| Tallada | - | 0.3731 | 0.9985 | 3.24 ± 0.68e | 1.2 ka | 0.34 |
| Outer Pleta Naua | - | 0.1289 | 0.9972 | 12.40 ± 0.32e | 0.6 ka | -0.24 |
| Arànser | Left | 0.7003 | 0.9978 | 23.29 ± 1.12f | 7.9 ka | -1.02 |
| Right | 0.6796 | 0.9991 | 22.30 ± 0.91f | 6.9 ka | -1.13 |
| Soum d’Ech | Outer | 0.5211 | 0.9971 | 25.94 ± 1.00f | 3.5 ka | -1.49 |
| Inner | 0.4620 | 0.9978 | 26.23 ± 0.81f | 3.5 ka | -1.05 |
| Combined | - | - | 26.05 ± 1.29f,g | 3.6 ka | -1.49 |
| a Numeric bandwidth used for kernel density estimation, b All model *p* values < 0.01, c Reporteduncertainty (±) is the 1σ bounds (68%) of the highest probability component Gaussian, d Interquartile range, e Calculated using the “STD/IQR” bandwidth estimator (Silverman, 1986; Dortch et al., 2020), f Calculated using the “MAD” bandwidth estimator (*see* Dortch et al., 2020), g Weighted mean of Soum d’Ech outer and inner landform ages, with uncertainties propagated through summation in quadrature. | | | | | | |

|  |  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- | --- |
| **Table 4.** Spatial statistics for the sampled morainesa | | | | | | | |
|  | Global Morans *I* | Class statistics (%) | | | “Good” boulder (%)b | | |
| Moraine | # significant datasets | Young | Good | Old | IS | C | OS |
| Tallada | 3% | 4b | 84 | 11c | 85 ± 1 | 79 ± | 90 ± |
| Outer Pleta Naua | < 1% | 0 | 98 | 2 | 100 ± | 95 ± | 100 ± |
| Arànser (Left) | 89% | 43d | 57 | 0 | 54 ± | 59 ± | 76 ± |
| Arànser (Right) | 92% |  |  |  |  |  |  |
| Soum d’Ech | 97% | 46d | 53 | 1 | 57 ± | 50 ± | 54 ± |
| a Derived from simulated datasets. b rockfall or avalanching, c reworking, d exhumation | | | | | | | |

Figure 1. Site photographs of the (A) Tallada, (B) Outer Pleta Naua, (C) Arànser and (D) Soum d’Ech moraines (denoted by red arrows). (E-F) Topographic maps of the Pyrenees (ASTER GDEM V3, UTM projection), showing the locations of the studied catchments and selected moraines and the distribution of Axial Zone granites within those catchments (Nomenclature following Porquet et al., 2017; Cauterets-Panticosa, Néouvielle, Maladeta, Mont-Louis-Andorra). The latter was derived from a 1:400,000 geological map produced by the IGME (*Spain*) and the BRGM (*France*). Also shown are the locations of major summits (Aneto, Carlit, Estats, Plana de Lles, Monte Perdido, Posets, Vignemale) and the maximum ice extent (MIE) during the global Last Glacial Maximum (gLGM; Calvet et al., 2011).

**Figure 2**. Moraine locations and likely boulder flow pathways (yellow arrows) for the (A) Soum d’Ech, (B) Arànser and (C) Val de Molières moraines (Outer Pleta Naua, Tallada). Relevant glacier limits and ice flow pathways are shown following Pallàs et al. (2006) and Calvet et al. (2011). The distribution of Axial Zone granites is shown in (A), but is excluded from plots (B) and (C) for clarity. Boulders on the Soum d’Ech moraines were likely sourced from the Cauterets-Panticosa granites surrounding Vignemale (3298 m), a major accumulation area for the Gave de Pau glacier, but a contribution from the Néouvielle granites to the east cannot be excluded (*see* Fig. 1E; Porquet et al., 2017). In (C), the location and recalibrated ages of samples MUL01 and MUL04 are shown (Pallàs et al., 2006). These samples correspond to a stabilisation or readvance of the Molières glacier following the gLGM, and comprise two of the 54 TCN-SH calibration surfaces (Tomkins et al., 2018b).

Figure 3. Geomorphological maps for the (A) Arànser and (B) Soum d’Ech moraines. These moraines likely correspond to the maximum ice extent (MIE) during the Würmian glacial stage (11.7 - 110 ka; Calvet et al., 2011). Locations and sample names for TCN dated boulders are shown (white circles; Rodés, 2008; Palacios et al., 2015). In (A), the locations of the proximal Fornell (F) and Setut (S) moraines are highlighted. These moraines are stratigraphically distinct from the sampled Arànser moraines but are currently undated. The Arànser left moraine can be tracked further up valley (*see* Fig. 2B) but sampling was focused on the illustrated moraine area (light purple shading) in which the moraine margins are easily delineated (≤ 2 km from glacier terminus).

**Figure 4**. Geomorphological maps for the (A) Tallada and (B) Outer Pleta Naua moraines in the Val de Molières catchment of the Noguera Rigaborçana. Locations and sample names for TCN dated boulders are shown (white circles; Pallàs et al., 2006).

**Figure 5**. (A) Location of exposure age calibration sites (blue points) in the Bassies (B, *n* = 6), Carlit (C, *n* = 3), Noguera Rigaborçana (N, *n* = 4), Maladeta (Ma, *n* = 9), Malniu (Mn, *n* = 21), Molières (Mo, *n* = 2), Orri (O, *n* = 3) and Querol catchments (Q, *n* = 6). Also shown are the locations of sampled moraines (orange points; *see* Fig. 1F) and the maximum ice extent (MIE) during the global Last Glacial Maximum (gLGM; Calvet et al., 2011). (B) Monte Carlo-derived orthogonal distance regression (ODR) between 54 10Be exposure ages (blue points ± external age uncertainty; Pallàs et al., 2006; 2010; Delmas et al., 2008; Crest et al., 2017) and their corresponding SH R-values (mean of 30 R-values ± Standard Error of the Mean; Tomkins et al., 2018b). 1σ (blue dashed lines) and 2σ prediction limits (grey dashed lines) were calculated using the ODR covariance matrix (Boggs and Rogers, 1990b). New TCN samples (10Be, 36Cl) from the studied moraines (*n* = 15) are shown as orange points. Inherited outliers from the original calibration dataset (*n* = 2; Tomkins et al., 2018b) are not shown for clarity. (C) Example of 10Be dated boulder from the Arànser right moraine (SAL-10). Sample information for all exposure ages (*n* = 73) is provided in the Supplementary Information and on GitHub: <https://github.com/matt-tomkins/moraine-paper-2020>

**Figure 6.** Gaussian decomposition of calibrated boulder exposure ages for the Tallada (A), Outer Pleta Naua (B), Arànser (C-D) and Soum d’Ech moraines (E-F). Following P-CAAT guidelines (Dortch et al., 2013; 2020), we selected the highest probability component Gaussian to represent the age of the landform as all are ≤ gLGM (red shaded Gaussian). The summed probability density estimate (PDE) and lower probability component Gaussians are denoted by black and grey distributions respectively.

**Figure 7.** Results of Monte Carlo simulated (*n* = 104) local Moran’s *I* spatial autocorrelation for the Tallada (A-B) and Outer Pleta Naua moraines (C-D). Results are based on Queen’s contiguity of Voronoi cells with the binary “Good/Bad” grouping as the predictor variable. Results are plotted as a proportion (%) of the total number of model runs. The age distribution of “good” (green shading) and “bad” boulders (grey shading) is shown for each moraine in plots B and D.

**Figure 8.** Results of Monte Carlo simulated (*n* = 104) local Moran’s *I* spatial autocorrelation for the Arànser (A-B) and Soum d’Ech moraines (C-D). Results are based on Queen’s contiguity of Voronoi cells with the binary “Good/Bad” grouping as the predictor variable. Results are plotted as a proportion (%) of the total number of model runs. The age distribution of “good” (green shading) and “bad” boulders (grey shading) is shown for each moraine in plots B and D.

**Figure 9.** The likelihood of sampling a “good” boulder (%; within 2σ of the landform age) for each of the studied landforms, subset by boulder position (inner ice-proximal slope, moraine crest, outer ice-distal slope). These data are based on Monte Carlo simulated datasets (*n* = 104), where bar heights represent the mean proportion (%) of “good” boulders and uncertainty bars represent the standard deviation.

**References**

Akçar, N., Ivy-Ochs, S., Kubik, P.W., Schlüchter, C., 2011. Post-depositional impacts on ‘Findlinge’ (erratic boulders) and their implications for surface-exposure dating. Swiss J Geosci 104, 445–453. <https://doi.org/10.1007/s00015-011-0088-7>

Allen, C.E., 2002. The influence of schistocity on soil weathering on large boulder tops, Kärkevagge, Sweden. CATENA, The interpretation and significance of weathering mantels 49, 157–169. <https://doi.org/10.1016/S0341-8162(02)00022-X>

Applegate, P.J., Urban, N.M., Keller, K., Lowell, T.V., Laabs, B.J.C., Kelly, M.A., Alley, R.B., 2012. Improved moraine age interpretations through explicit matching of geomorphic process models to cosmogenic nuclide measurements from single landforms. Quaternary Research 77, 293–304. <https://doi.org/10.1016/j.yqres.2011.12.002>

Applegate, P.J., Urban, N.M., Laabs, B.J.C., Keller, K., Alley, R.B., 2010. Modeling the statistical distributions of cosmogenic exposure dates from moraines. Geoscientific Model Development 3, 293–307. <https://doi.org/10.5194/gmd-3-293-2010>

Aydin, A., Basu, A., 2005. The Schmidt hammer in rock material characterization. Engineering Geology 81, 1–14. <https://doi.org/10.1016/j.enggeo.2005.06.006>

Balco, G., 2006. Converting Al and Be isotope ratio measurements to nuclide concentrations in quartz. Cosmogenic Nuclide Lab, University of Washington.

Balco, G., Stone, J.O., Lifton, N.A., Dunai, T.J., 2008. A complete and easily accessible means of calculating surface exposure ages or erosion rates from 10Be and 26Al measurements. Quaternary Geochronology, Prospects for the New Frontiers of earth and Environmental Sciences 3, 174–195. <https://doi.org/10.1016/j.quageo.2007.12.001>

Balter, A., Bromley, G., Balco, G., Thomas, H., Jackson, M.S., 2020. A 14.5 million-year record of East Antarctic Ice Sheet fluctuations from the central Transantarctic Mountains, constrained with cosmogenic 3He, 10Be, 21Ne, and 26Al. The Cryosphere Discussions 1–41. <https://doi.org/10.5194/tc-2020-57>

Barr, I.D., Lovell, H., 2014. A review of topographic controls on moraine distribution. Geomorphology 226, 44–64. <https://doi.org/10.1016/j.geomorph.2014.07.030>

Birkeland, P., 1999. Soils and Geomorphology, 3 edition. ed. Oxford University Press, New York.

Boggs, Paul T, Rogers, J.E., 1990. The Computation and Use of the Asymptotic Covariance Matrix for Measurement Error Models (No. Internal Report 89-4102). National Institute of Standards and Technology, Gaithersburg, MD, Applied and Computational Mathematics Division.

Boggs, Paul T., Rogers, J.E., 1990. Orthogonal distance regression, in: “Statistical Analysis of Measurement Error Models and Applications: Proceedings of the AMS-IMS-SIAM Joint Summer Research Conference Held June 10-16, 1989,.” Presented at the Contemporary Mathematics, p. 186. <https://doi.org/10.6028/nist.ir.89-4197>

Borchers, B., Marrero, S., Balco, G., Caffee, M., Goehring, B., Lifton, N., Nishiizumi, K., Phillips, F., Schaefer, J., Stone, J., 2016. Geological calibration of spallation production rates in the CRONUS-Earth project. Quaternary Geochronology 31, 188–198. <https://doi.org/10.1016/j.quageo.2015.01.009>

Briner, J.P., Kaufman, D.S., Manley, W.F., Finkel, R.C., Caffee, M.W., 2005. Cosmogenic exposure dating of late Pleistocene moraine stabilization in Alaska. GSA Bulletin 117, 1108–1120. <https://doi.org/10.1130/B25649.1>

Broecker, W.S., Denton, G.H., 1990. The role of ocean-atmosphere reorganizations in glacial cycles. Quaternary Science Reviews 9, 305–341. <https://doi.org/10.1016/0277-3791(90)90026-7>

Brown, E.T., Brook, E.J., Raisbeck, G.M., Yiou, F., Kurz, M.D., 1992. Effective attenuation lengths of cosmic rays producing 10Be and 26Al in quartz: Implications for exposure age dating. Geophysical Research Letters 19, 369–372. <https://doi.org/10.1029/92GL00266>

Brown, E.T., Edmond, J.M., Raisbeck, G.M., Yiou, F., Kurz, M.D., Brook, E.J., 1991. Examination of surface exposure ages of Antarctic moraines using *in situ* produced 10Be and 26Al. Geochimica et Cosmochimica Acta 55, 2269–2283. <https://doi.org/10.1016/0016-7037(91)90103-C>

Buckland, W., 1840. On the evidences of glaciers in Scotland and the north of England. Proceedings of the Geological Society of London 3, 332–337.

Calvet, M., Delmas, M., Gunnell, Y., Braucher, R., Bourlès, D., 2011. Chapter 11 - Recent Advances in Research on Quaternary Glaciations in the Pyrenees, in: Ehlers, J., Gibbard, P.L., Hughes, P.D. (Eds.), Developments in Quaternary Sciences, Quaternary Glaciations - Extent and Chronology. Elsevier, pp. 127–139. <https://doi.org/10.1016/B978-0-444-53447-7.00011-8>

Cerling, T.E., Craig, H., 1994. Geomorphology and in-situ cosmogenic isotopes. Annual Review of Earth and Planetary Sciences 22, 273–317. <https://doi.org/10.1146/annurev.ea.22.050194.001421>

Chartres, C.J., Walker, P.H., 1988. The effect of Aeolian accessions on soil development on granitic-rocks in south eastern Australia. III. Micromorphological and geochemical evidence of weathering and soil development. Soil Res. 26, 33–53. <https://doi.org/10.1071/sr9880033>

Crest, Y., Delmas, M., Braucher, R., Gunnell, Y., Calvet, M., 2017. Cirques have growth spurts during deglacial and interglacial periods: Evidence from 10Be and 26Al nuclide inventories in the central and eastern Pyrenees. Geomorphology 278, 60–77. <https://doi.org/10.1016/j.geomorph.2016.10.035>

Darmody, R.G., Thorn, C.E., Allen, C.E., 2005. Chemical weathering and boulder mantles, Kärkevagge, Swedish Lapland. Geomorphology, Weathering and landscape evolution 67, 159–170. <https://doi.org/10.1016/j.geomorph.2004.07.011>

Darmody, R.G., Thorn, C.E., Dixon, J.C., 2008. Differential rock weathering in the ‘valley of the boulders’’, kärkevagge, swedish lapland.’ Geografiska Annaler: Series A, Physical Geography 90, 201–209. <https://doi.org/10.1111/j.1468-0459.2008.339.x>

Darvill, C.M., Bentley, M.J., Stokes, C.R., 2015a. Geomorphology and weathering characteristics of erratic boulder trains on Tierra del Fuego, southernmost South America: Implications for dating of glacial deposits. Geomorphology 228, 382–397. <https://doi.org/10.1016/j.geomorph.2014.09.017>

Darvill, C.M., Bentley, M.J., Stokes, C.R., Hein, A.S., Rodés, Á., 2015b. Extensive MIS 3 glaciation in southernmost Patagonia revealed by cosmogenic nuclide dating of outwash sediments. Earth and Planetary Science Letters 429, 157–169. <https://doi.org/10.1016/j.epsl.2015.07.030>

Delmas, M., Gunnell, Y., Braucher, R., Calvet, M., Bourlès, D., 2008. Exposure age chronology of the last glaciation in the eastern Pyrenees. Quaternary Research 69, 231–241. <https://doi.org/10.1016/j.yqres.2007.11.004>

Dixon, J.C., Campbell, S.W., Thorn, C.E., Darmody, R.G., 2006. Incipient weathering rind development on introduced machine-polished granite discs in an Arctic alpine environment, northern Scandinavia. Earth Surface Processes and Landforms 31, 111–121. <https://doi.org/10.1002/esp.1241>

Dortch, J., Saha, S., Tomkins, M.D., Murari, M.K., Schoenbohm, L.M., Curl, D., 2019. Probability-based interpretation of terrestrial cosmogenic radionuclide ages: P-CAAT, a tool for the ages. Presented at the AGU Fall Meeting 2019, AGU. <https://agu.confex.com/agu/fm19/meetingapp.cgi/Paper/502207>

Dortch, J.M., Hughes, P.D., Tomkins, M.D., 2016. Schmidt hammer exposure dating (SHED): Calibration boulder of Tomkins et al. (2016). Quaternary Geochronology 35, 67–68. <https://doi.org/10.1016/j.quageo.2016.06.001>

Dortch, J.M., Owen, L.A., Caffee, M.W., 2013. Timing and climatic drivers for glaciation across semi-arid western Himalayan–Tibetan orogen. Quaternary Science Reviews 78, 188–208. <https://doi.org/10.1016/j.quascirev.2013.07.025>

Dortch, J.M., Owen, L.A., Caffee, M.W., Li, D., Lowell, T.V., 2010. Beryllium-10 surface exposure dating of glacial successions in the Central Alaska Range. Journal of Quaternary Science 25, 1259–1269. <https://doi.org/10.1002/jqs.1406>

Ehlmann, B.L., Viles, H.A., Bourke, M.C., 2008. Quantitative morphologic analysis of boulder shape and surface texture to infer environmental history: A case study of rock breakdown at the Ephrata Fan, Channeled Scabland, Washington. Journal of Geophysical Research: Earth Surface 113. <https://doi.org/10.1029/2007JF000872>

Engel, Z., 2007. Measurement and age assignment of intact rock strength in the Krkonoše Mountains, Czech Republic. Zeitschrift für Geomorphologie, Supplementary Issues 69–80. <https://doi.org/10.1127/0372-8854/2007/0051S-0069>

Engel, Z., Traczyk, A., Braucher, R., Woronko, B., Křížek, M., 2011. Use of 10Be exposure ages and Schmidt hammer data for correlation of moraines in the Krkonoše Mountains, Poland/Czech Republic. Zeitschrift für Geomorphologie 175–196. <https://doi.org/10.1127/0372-8854/2011/0055-0036>

Fabel, D., Ballantyne, C.K., Xu, S., 2012. Trimlines, blockfields, mountain-top erratics and the vertical dimensions of the last British–Irish Ice Sheet in NW Scotland. Quaternary Science Reviews 55, 91–102. <https://doi.org/10.1016/j.quascirev.2012.09.002>

Gibbons, A.B., Megeath, J.D., Pierce, K.L., 1984. Probability of moraine survival in a succession of glacial advances. Geology 12, 327–330. [https://doi.org/10.1130/0091-7613(1984)12<327:POMSIA>2.0.CO;2](https://doi.org/10.1130/0091-7613(1984)12%3c327:POMSIA%3e2.0.CO;2)

Goudie, A.S., 2016. The Schmidt Hammer in geomorphological research: Progress in Physical Geography. <https://doi.org/10.1177/0309133306071954>

Hallet, B., Putkonen, J., 1994. Surface Dating of Dynamic Landforms: Young Boulders on Aging Moraines. Science 265, 937–940. <https://doi.org/10.1126/science.265.5174.937>

Hays, J.D., Imbrie, J., Shackleton, N.J., 1976. Variations in the Earth’s Orbit: Pacemaker of the Ice Ages. Science 194, 1121–1132. <https://doi.org/10.1126/science.194.4270.1121>

Hein, A.S., Fogwill, C.J., Sugden, D.E., Xu, S., 2014. Geological scatter of cosmogenic-nuclide exposure ages in the Shackleton Range, Antarctica: Implications for glacial history. Quaternary Geochronology, Tracking the pace of Quaternary landscape change with cosmogenic nuclides 19, 52–66. <https://doi.org/10.1016/j.quageo.2013.03.008>

Heyman, J., Applegate, P.J., Blomdin, R., Gribenski, N., Harbor, J.M., Stroeven, A.P., 2016. Boulder height – exposure age relationships from a global glacial 10Be compilation. Quaternary Geochronology 34, 1–11. <https://doi.org/10.1016/j.quageo.2016.03.002>

Heyman, J., Stroeven, A.P., Harbor, J.M., Caffee, M.W., 2011. Too young or too old: Evaluating cosmogenic exposure dating based on an analysis of compiled boulder exposure ages. Earth and Planetary Science Letters 302, 71–80. <https://doi.org/10.1016/j.epsl.2010.11.040>

Hughes, P.D., 2018. Little Ice Age glaciers and climate in the Mediterranean mountains: a new analysis. Cuadernos de Investigación Geográfica 44, 15–45. <https://doi.org/10.18172/cig.3362>

Hughes, P.D., Glasser, N.F., Fink, D., 2016. Rapid thinning of the Welsh Ice Cap at 20–19ka based on 10Be ages. Quaternary Research 85, 107–117. <https://doi.org/10.1016/j.yqres.2015.11.003>

Hughes, P.D., Woodward, J.C., Gibbard, P.L., 2007. Middle Pleistocene cold stage climates in the Mediterranean: New evidence from the glacial record. Earth and Planetary Science Letters 253, 50–56. <https://doi.org/10.1016/j.epsl.2006.10.019>

Ivy-Ochs, S., Kerschner, H., Schlüchter, C., 2007. Cosmogenic nuclides and the dating of Lateglacial and Early Holocene glacier variations: The Alpine perspective. Quaternary International, From the Swiss Alps to the Crimean Mountains - Alpine Quaternary stratigraphy in a European context 164–165, 53–63. <https://doi.org/10.1016/j.quaint.2006.12.008>

Jalut, G., Marti, J.M., Fontugne, M., Delibrias, G., Vilaplana, J.M., Julia, R., 1992. Glacial to interglacial vegetation changes in the northern and southern Pyrénées: Deglaciation, vegetation cover and chronology. Quaternary Science Reviews 11, 449–480. <https://doi.org/10.1016/0277-3791(92)90027-6>

Jull, A.J.T., Scott, E.M., Bierman, P., 2015. The CRONUS-Earth inter-comparison for cosmogenic isotope analysis. Quaternary Geochronology, The CRONUS-EARTH Volume: Part I 26, 3–10. <https://doi.org/10.1016/j.quageo.2013.09.003>

Kendrick, K.J., Partin, C.A., Graham, R.C., 2016. Granitic Boulder Erosion Caused by Chaparral Wildfire: Implications for Cosmogenic Radionuclide Dating of Bedrock Surfaces. The Journal of Geology 124, 529–539. <https://doi.org/10.1086/686273>

Kohl, C.P., Nishiizumi, K., 1992. Chemical isolation of quartz for measurement of *in-situ* -produced cosmogenic nuclides. Geochimica et Cosmochimica Acta 56, 3583–3587. <https://doi.org/10.1016/0016-7037(92)90401-4>

Kuhlemann, J., Krumrei, I., Danišík, M., Borg, K. van der, 2009. Weathering of granite and granitic regolith in Corsica: short-term 10Be versus long-term thermochronological constraints. Geological Society, London, Special Publications 324, 217–235. <https://doi.org/10.1144/SP324.16>

Kuhlemann, J., van der Borg, K., Bons, P.D., Danišík, M., Frisch, W., 2008. Erosion rates on subalpine paleosurfaces in the western Mediterranean by in-situ 10Be concentrations in granites: implications for surface processes and long-term landscape evolution in Corsica (France). Int J Earth Sci (Geol Rundsch) 97, 549–564. <https://doi.org/10.1007/s00531-007-0169-z>

Lal, D., 1991. Cosmic ray labeling of erosion surfaces: in situ nuclide production rates and erosion models. Earth and Planetary Science Letters 104, 424–439. <https://doi.org/10.1016/0012-821X(91)90220-C>

Lewis, C.J., McDonald, E.V., Sancho, C., Peña, J.L., Rhodes, E.J., 2009. Climatic implications of correlated Upper Pleistocene glacial and fluvial deposits on the Cinca and Gállego Rivers (NE Spain) based on OSL dating and soil stratigraphy. Global and Planetary Change 67, 141–152. <https://doi.org/10.1016/j.gloplacha.2009.01.001>

Marrero, S.M., Hein, A.S., Naylor, M., Attal, M., Shanks, R., Winter, K., Woodward, J., Dunning, S., Westoby, M., Sugden, D., 2018. Controls on subaerial erosion rates in Antarctica. Earth and Planetary Science Letters 501, 56–66. <https://doi.org/10.1016/j.epsl.2018.08.018>

Marrero, S.M., Phillips, F.M., Borchers, B., Lifton, N., Aumer, R., Balco, G., 2016. Cosmogenic nuclide systematics and the CRONUScalc program. Quaternary Geochronology 31, 160–187. <https://doi.org/10.1016/j.quageo.2015.09.005>

Matthews, J.A., Owen, G., 2008. Endolithic lichens, rapid biological weathering and schmidt hammer r‐values on recently exposed rock surfaces: storbreen glacier foreland, jotunheimen, norway. Geografiska Annaler: Series A, Physical Geography 90, 287–297. <https://doi.org/10.1111/j.1468-0459.2008.00346.x>

McCarroll, D., 1991. The Schmidt Hammer, weathering and rock surface roughness. Earth Surface Processes and Landforms 16, 477–480. <https://doi.org/10.1002/esp.3290160510>

McCarroll, D., 1989. Potential and Limitations of the Schmidt Hammer for Relative-Age Dating: Field Tests on Neoglacial Moraines, Jotunheimen, Southern Norway. Arctic and Alpine Research 21, 268–275. <https://doi.org/10.2307/1551565>

Mersmann, O., Trautmann, H., Steuer, D., Bornkamp, B., 2018. truncnorm: Truncated Normal Distribution.

Moran, P.A.P., 1950. Notes on Continuous Stochastic Phenomena. Biometrika 37, 17–23. <https://doi.org/10.2307/2332142>

Morgan, D.J., Putkonen, J., Balco, G., Stone, J., 2011. Degradation of glacial deposits quantified with cosmogenic nuclides, Quartermain Mountains, Antarctica. Earth Surface Processes and Landforms 36, 217–228. <https://doi.org/10.1002/esp.2039>

Murari, M.K., Owen, L.A., Dortch, J.M., Caffee, M.W., Dietsch, C., Fuchs, M., Haneberg, W.C., Sharma, M.C., Townsend-Small, A., 2014. Timing and climatic drivers for glaciation across monsoon-influenced regions of the Himalayan–Tibetan orogen. Quaternary Science Reviews 88, 159–182. <https://doi.org/10.1016/j.quascirev.2014.01.013>

Niedzielski, T., Migoń, P., Placek, A., 2009. A minimum sample size required from Schmidt hammer measurements. Earth Surface Processes and Landforms 34, 1713–1725. <https://doi.org/10.1002/esp.1851>

Nishiizumi, K., Imamura, M., Caffee, M.W., Southon, J.R., Finkel, R.C., McAninch, J., 2007. Absolute calibration of 10Be AMS standards. Nuclear Instruments and Methods in Physics Research Section B: Beam Interactions with Materials and Atoms 258, 403–413. <https://doi.org/10.1016/j.nimb.2007.01.297>

Nishiizumi, K., Winterer, E.L., Kohl, C.P., Klein, J., Middleton, R., Lal, D., Arnold, J.R., 1989. Cosmic ray production rates of 10Be and 26Al in quartz from glacially polished rocks. Journal of Geophysical Research: Solid Earth 94, 17907–17915. <https://doi.org/10.1029/JB094iB12p17907>

Oliva, M., Palacios, D., Fernández-Fernández, J.M., Rodríguez-Rodríguez, L., García-Ruiz, J.M., Andrés, N., Carrasco, R.M., Pedraza, J., Pérez-Alberti, A., Valcárcel, M., Hughes, P.D., 2019. Late Quaternary glacial phases in the Iberian Peninsula. Earth-Science Reviews 192, 564–600. <https://doi.org/10.1016/j.earscirev.2019.03.015>

Owen, L.A., Caffee, M.W., Bovard, K.R., Finkel, R.C., Sharma, M.C., 2006. Terrestrial cosmogenic nuclide surface exposure dating of the oldest glacial successions in the Himalayan orogen: Ladakh Range, northern India. GSA Bulletin 118, 383–392. <https://doi.org/10.1130/B25750.1>

Palacios, D., Gómez-Ortiz, A., Alcalá-Reygosa, J., Andrés, N., Oliva, M., Tanarro, L.M., Salvador-Franch, F., Schimmelpfennig, I., Fernández-Fernández, J.M., Léanni, L., 2019. The challenging application of cosmogenic dating methods in residual glacial landforms: The case of Sierra Nevada (Spain). Geomorphology 325, 103–118. <https://doi.org/10.1016/j.geomorph.2018.10.006>

Palacios, D., Gómez-Ortiz, A., Andrés, N., Vázquez-Selem, L., Salvador-Franch, F., Oliva, M., 2015. Maximum extent of Late Pleistocene glaciers and last deglaciation of La Cerdanya mountains, Southeastern Pyrenees. Geomorphology 231, 116–129. <https://doi.org/10.1016/j.geomorph.2014.10.037>

Pallàs, R., Rodés, Á., Braucher, R., Bourlès, D., Delmas, M., Calvet, M., Gunnell, Y., 2010. Small, isolated glacial catchments as priority targets for cosmogenic surface exposure dating of Pleistocene climate fluctuations, southeastern Pyrenees. Geology 38, 891–894. <https://doi.org/10.1130/G31164.1>

Pallàs, R., Rodés, Á., Braucher, R., Carcaillet, J., Ortuño, M., Bordonau, J., Bourlès, D., Vilaplana, J.M., Masana, E., Santanach, P., 2006. Late Pleistocene and Holocene glaciation in the Pyrenees: a critical review and new evidence from 10Be exposure ages, south-central Pyrenees. Quaternary Science Reviews 25, 2937–2963. <https://doi.org/10.1016/j.quascirev.2006.04.004>

Penck, A., 1905. Glacial Features in the Surface of the Alps. The Journal of Geology 13, 1–19.

Phillips, F.M., Zreda, M.G., Smith, S.S., Elmore, D., Kubik, P.W., Sharma, P., 1990. Cosmogenic Chlorine-36 Chronology for Glacial Deposits at Bloody Canyon, Eastern Sierra Nevada. Science 248, 1529–1532. <https://doi.org/10.1126/science.248.4962.1529>

Porquet, M., Pueyo, E.L., Román-Berdiel, T., Olivier, P., Longares, L.A., Cuevas, J., Ramajo, J., order, the G. working group by alphabetical, Antolín, B., Aranguren, A., Auréjac, J.B., Bouchez, J.-L., Casas, A.M., Denèle, Y., Gleizes, G., Hilario, A., Izquierdo-Llavall, E., Leblanc, D., Oliva-Urcia, B., Santana, V., Tubía, J.M., Vegas, N., 2017. Anisotropy of magnetic susceptibility of the Pyrenean granites. Journal of Maps 13, 438–448. <https://doi.org/10.1080/17445647.2017.1302364>

Portenga, E.W., Bierman, P.R., 2011. Understanding Earth’s eroding surface with 10Be. GSAT 21, 4–10. <https://doi.org/10.1130/G111A.1>

Porter, S.C., Swanson, T.W., 2008. 36Cl dating of the classic Pleistocene glacial record in the northeastern Cascade Range, Washington. Am J Sci 308, 130–166. <https://doi.org/10.2475/02.2008.02>

Putkonen, J., Connolly, J., Orloff, T., 2008. Landscape evolution degrades the geologic signature of past glaciations. Geomorphology, Glacial Landscape Evolution - Implications for Glacial Processes, Patterns and Reconstructions 97, 208–217. <https://doi.org/10.1016/j.geomorph.2007.02.043>

Putkonen, J., O’Neal, M., 2006. Degradation of unconsolidated Quaternary landforms in the western North America. Geomorphology, Quaternary landscape change and modern process in western North America 75, 408–419. <https://doi.org/10.1016/j.geomorph.2005.07.024>

Putkonen, J., Swanson, T., 2003. Accuracy of cosmogenic ages for moraines. Quaternary Research 59, 255–261. <https://doi.org/10.1016/S0033-5894(03)00006-1>

Putnam, A.E., Bromley, G.R.M., Rademaker, K., Schaefer, J.M., 2019. In situ 10Be production-rate calibration from a 14C-dated late-glacial moraine belt in Rannoch Moor, central Scottish Highlands. Quaternary Geochronology 50, 109–125. <https://doi.org/10.1016/j.quageo.2018.11.006>

Raisbeck, G.M., Yiou, F., Bourlès, D., Brown, E., Deboffle, D., Jouhanneau, P., Lestringuez, J., Zhou, Z.Q., 1994. The AMS facility at Gif-sur-Yvette: progress, perturbations and projects. Nuclear Instruments and Methods in Physics Research Section B: Beam Interactions with Materials and Atoms 92, 43–46. <https://doi.org/10.1016/0168-583X(94)95972-2>

Reille, M., Andrieu, V., 1995. The late Pleistocene and Holocene in the Lourdes Basin, Western Pyrénées, France: new pollen analytical and chronological data. Veget Hist Archaebot 4, 1–21. <https://doi.org/10.1007/BF00198611>

Riebe, C.S., Kirchner, J.W., Finkel, R.C., 2004. Erosional and climatic effects on long-term chemical weathering rates in granitic landscapes spanning diverse climate regimes. Earth and Planetary Science Letters 224, 547–562. <https://doi.org/10.1016/j.epsl.2004.05.019>

Riebe, C.S., Kirchner, J.W., Granger, D.E., Finkel, R.C., 2001. Minimal climatic control on erosion rates in the Sierra Nevada, California. Geology 29, 447–450. [https://doi.org/10.1130/0091-7613(2001)029<0447:MCCOER>2.0.CO;2](https://doi.org/10.1130/0091-7613(2001)029%3c0447:MCCOER%3e2.0.CO;2)

Rinterknecht, V., Börner, A., Bourlès, D., Braucher, R., 2014. Cosmogenic 10Be dating of ice sheet marginal belts in Mecklenburg-Vorpommern, Western Pomerania (northeast Germany). Quaternary Geochronology, Tracking the pace of Quaternary landscape change with cosmogenic nuclides 19, 42–51. <https://doi.org/10.1016/j.quageo.2013.05.003>

Rodés, Á., 2008. La última deglaciación en los pirineos: de superficies de exposición mediante 10be, y modelado numérico de paleoglaciares (http://purl.org/dc/dcmitype/Text). Universitat de Barcelona.

Schaller, M., Ehlers, T.A., Blum, J.D., Kallenberg, M.A., 2009. Quantifying glacial moraine age, denudation, and soil mixing with cosmogenic nuclide depth profiles. Journal of Geophysical Research: Earth Surface 114. <https://doi.org/10.1029/2007JF000921>

Shanahan, T.M., Zreda, M., 2000. Chronology of Quaternary glaciations in East Africa. Earth and Planetary Science Letters 177, 23–42. <https://doi.org/10.1016/S0012-821X(00)00029-7>

Silverman, B.W., 1986. Density Estimation for Statistics and Data Analysis. CRC Press.

Small, D., Fabel, D., 2015. A Lateglacial 10Be production rate from glacial lake shorelines in Scotland. Journal of Quaternary Science 30, 509–513. <https://doi.org/10.1002/jqs.2804>

Small, E.E., Anderson, R.S., Repka, J.L., Finkel, R., 1997. Erosion rates of alpine bedrock summit surfaces deduced from in situ 10Be and 26Al. Earth and Planetary Science Letters 150, 413–425. <https://doi.org/10.1016/S0012-821X(97)00092-7>

Stone, J.O., 2000. Air pressure and cosmogenic isotope production. Journal of Geophysical Research: Solid Earth 105, 23753–23759. <https://doi.org/10.1029/2000JB900181>

Sumner, P., Nel, W., 2002. The effect of rock moisture on Schmidt hammer rebound: tests on rock samples from Marion Island and South Africa. Earth Surface Processes and Landforms 27, 1137–1142. <https://doi.org/10.1002/esp.402>

Tomkins, M.D., Dortch, J.M., Hughes, P.D., 2016. Schmidt Hammer exposure dating (SHED): Establishment and implications for the retreat of the last British Ice Sheet. Quaternary Geochronology 33, 46–60. <https://doi.org/10.1016/j.quageo.2016.02.002>

Tomkins, M.D., Dortch, J.M., Hughes, P.D., Huck, J.J., Stimson, A.G., Delmas, M., Calvet, M., Pallàs, R., 2018a. Rapid age assessment of glacial landforms in the Pyrenees using Schmidt hammer exposure dating (SHED). Quaternary Research 90, 26–37. <https://doi.org/10.1017/qua.2018.12>

Tomkins, M.D., Huck, J.J., Dortch, J.M., Hughes, P.D., Kirkbride, M.P., Barr, I.D., 2018b. Schmidt Hammer exposure dating (SHED): Calibration procedures, new exposure age data and an online calculator. Quaternary Geochronology 44, 55–62. <https://doi.org/10.1016/j.quageo.2017.12.003>

Viles, H., Goudie, A., Grab, S., Lalley, J., 2011. The use of the Schmidt Hammer and Equotip for rock hardness assessment in geomorphology and heritage science: a comparative analysis. Earth Surface Processes and Landforms 36, 320–333. <https://doi.org/10.1002/esp.2040>

Wadell, H., 1935. Volume, Shape, and Roundness of Quartz Particles. The Journal of Geology 43, 250–280. <https://doi.org/10.1086/624298>

Ward, D., Licciardi, J.M., Goehring, B.M., 2019. Three-isotope cosmogenic dating reveals a complex deglaciation history in the western Teton Range. Presented at the AGU Fall Meeting 2019, AGU.

Williams, R.B.G., Robinson, D.A., 1983. The effect of surface texture on the determination of the surface hardness of rock using the schmidt hammer. Earth Surface Processes and Landforms 8, 289–292. <https://doi.org/10.1002/esp.3290080311>

Wilson, P., Dunlop, P., Millar, C., Wilson, F.A., 2019. Age determination of glacially-transported boulders in Ireland and Scotland using Schmidt-hammer exposure-age dating (SHD) and terrestrial cosmogenic nuclide (TCN) exposure-age dating. Quaternary Research 1–13. <https://doi.org/10.1017/qua.2019.12>

Winkler, S., 2005. The Schmidt hammer as a relative‐age dating technique: Potential and limitations of its application on Holocene moraines in Mt Cook National Park, Southern Alps, New Zealand. New Zealand Journal of Geology and Geophysics 48, 105–116. <https://doi.org/10.1080/00288306.2005.9515102>

Zasadni, J., Kłapyta, P., Broś, E., Ivy-Ochs, S., Świąder, A., Christl, M., Balážovičová, L., 2020. Latest Pleistocene glacier advances and post-Younger Dryas rock glacier stabilization in the Mt. Kriváň group, High Tatra Mountains, Slovakia. Geomorphology 358, 107093. <https://doi.org/10.1016/j.geomorph.2020.107093>

Zech, R., Glaser, B., Sosin, P., Kubik, P.W., Zech, W., 2005. Evidence for long-lasting landform surface instability on hummocky moraines in the Pamir Mountains (Tajikistan) from 10Be surface exposure dating. Earth and Planetary Science Letters 237, 453–461. <https://doi.org/10.1016/j.epsl.2005.06.031>

Zreda, M.G., Phillips, F.M., 1995. Insights into alpine moraine development from cosmogenic 36Cl buildup dating. Geomorphology, Glacial Geomorphology: Process and Form Development 14, 149–156. <https://doi.org/10.1016/0169-555X(95)00055-9>

Zreda, M.G., Phillips, F.M., Elmore, D., 1994. Cosmogenic 36Cl accumulation in unstable landforms: 2. Simulations and measurements on eroding moraines. Water Resources Research 30, 3127–3136. <https://doi.org/10.1029/94WR00760>