Problem 1: Temperature of a satellite shadowed by the Earth

- A black satellite is orbiting at an altitude H=2000km above the ground.
- 1) Compute the solid angle subtended by the Earth.
- 2) If the Earth irradiates as a black body at T=255 K, compute the Earth radiation flux impinging onto the satellite (assumed as a sphere of radius $R_{sat}=1$ m).
- 3) Compute the radiative equilibrium temperature of the satellite when it is in the shadow of the Earth.
- 4) How does your answer change when the satellite is also illuminated by the Sun?

Computation of the solid angle: [2p]

Solid angle subtended by a spherical cap

$$d\Omega = d\phi \sin\theta \, d\theta$$

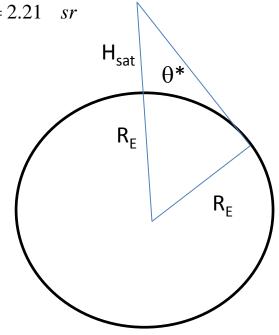
1st method

$$\Omega_{spherical cap} = \int d\Omega = \iint_{\varphi,\theta} d\phi \sin(\theta) d\theta = 2\pi \left[1 - \cos(\theta^*)\right] = 2.21 sr$$

$$\sin(\theta^*) = \frac{R_E}{H_{sat} + R_E} = \frac{6400}{8400} = 0.762$$

$$\cos(\theta^*) = \sqrt{1 - \sin^2(\theta^*)} = 0.648$$

$$\theta^* = \arcsin(0.762) = 0.8662 = 49.6^\circ$$



Note that

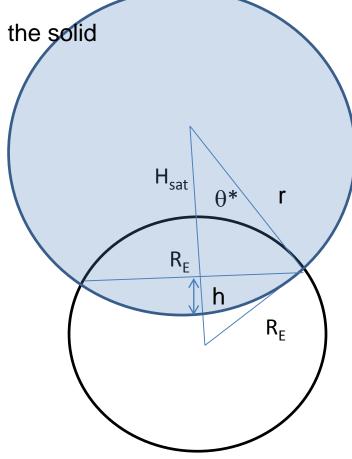
$$\Omega_{\it spherical cap} o 2\pi \ \it for \ H_{\it sat} o 0$$

Computation of the solid angle from area of spherical cap

$$\Omega = \frac{A_{\perp}}{r^2}$$

The solid angle subtended by the Earth is the same as the solid angle subtended by the

Yellow spherical cap



Computation of the solid angle from area of spherical cap

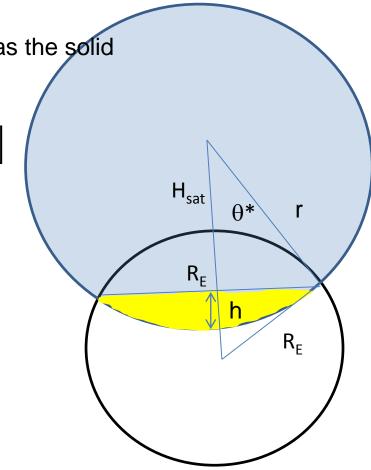
$$\Omega = \frac{A_{\perp}}{r^2}$$

The solid angle subtended by the Earth is the same as the solid angle subtended by the

Yellow spherical cap

$$A_{\perp} = 2 \pi r h = 2 \pi r r \left[1 - \cos(\theta^*) \right] = 2 \pi r^2 \left[1 - \cos(\theta^*) \right]$$

$$\Omega = \frac{2\pi r^2 \left[1 - \cos(\theta^*)\right]}{r^2} = 2\pi \left[1 - \cos(\theta^*)\right]$$



If the Earth irradiates as a black body at T=255 K, compute the Earth radiation flux impinging onto the satellite (assumed as a sphere of radius R_{sat}=1m)

$$F_{\lambda} = \int_{4\pi} d\Omega \cos\theta \ I_{\lambda} \quad \left[W \, m^{-2} \, m^{-1} \right]$$

$$\sin(\theta^*) = 0.762$$

$$I_{\lambda} = B_{\lambda}(T_{Earth})$$

$$F_{\lambda} = \int_{0}^{2\pi} d\phi \int_{0}^{\theta^{*}} I_{\lambda} \cos\theta \sin\theta d\theta = 2\pi I_{\lambda} \sin^{2}\theta^{*}/2 = \pi I_{\lambda} \sin^{2}\theta^{*}$$
 expect to get less than that!!

$$F = \int F_{\lambda} d\lambda = \sigma T_{Earth}^{4} \sin^{2}(\theta^{*}) = 240 \times 0.58 = 139 W m^{2}$$

$$\sigma = 5.67 \times 10^{-8} W m^{-2} K^{-4}$$

 W/m^2

Note that for a collimated source, looking straight into it (sun): $F=I\Delta\Omega$

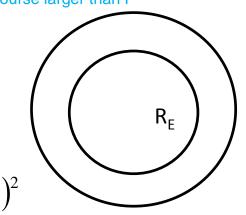
$$F_{coll\ source} = I\ \Delta\Omega = \frac{240}{\pi} \times 2.21 = 168.8\ W\ /\ m^2$$
 which is of course larger than F

2nd method

Just imposing conservation of energy:

$$P_{\lambda}[R_E] = P_{\lambda}[H_{sat} + R_E]$$

$$F_{\lambda}[R_{E}]4\pi R_{E}^{2} = \pi I_{\lambda} 4\pi R_{E}^{2} = F_{\lambda}[H_{sat} + R_{E}]4\pi (H_{sat} + R_{E})^{2}$$



Radiative equilibrium [3p]

$$F_{\lambda} = \pi I_{\lambda} \sin^2(\theta^*)$$

$$P_{abs}^{Earth} = \pi R_{sat}^{2} \int F_{\lambda} d\lambda = \pi^{2} R_{sat}^{2} \int I_{\lambda} d\lambda \sin^{2}(\theta^{*}) = \pi R_{sat}^{2} \sigma T_{E}^{4} \sin^{2}(\theta^{*}) = 437.2 W$$

$$F = 139.2 \text{ W/m}^{2}$$

$$239.7 \text{ W/m}^{2}$$

$$F = \frac{139.2 \text{ W/m}^{2}}{\pi}$$
Stefan's Constant 5.670 x 10⁻⁸ W / m² / K⁴

Stefan's Constant $5.670 \times 10^{-8} \text{ W} / \text{m}^2 / \text{K}^4$

$$P_{em} = \sigma T_{sat}^4 4\pi R_{sat}^2$$

$$P_{abs}^{Earth} = P_{em} \implies T_{sat} = T_E \left[\frac{\sin^2(\theta^*)}{4} \right]^{0.25} = 255 \sqrt{\frac{0.762}{2}} = 157.4 K$$

Contribution from the Sun [3p]

$$P_{abs}^{Sun} = \pi R_{sat}^2 \int F_{\lambda} d\lambda = \pi^2 R_{sat}^2 \underbrace{\int I_{\lambda} d\lambda}_{Sat} \sin^2(\theta_{Sun}) = \pi R_{sat}^2 \underbrace{\sigma_{SB} T_S^4 \left(\frac{R_{Sun}}{d_{Sun}}\right)^2}_{C_0 = 1380 \, \text{W} / m^2} = 4331 \, \text{W}$$

$$\underbrace{\tau_{Sun}^2 = 5800 \, \text{K}}_{T_{Sun}} = 5800 \, \text{K}$$
Sun Diameter = 1,390,600 km

Black body assumption

$$P_{abs}^{Earth} + P_{ass}^{Sun} = P_{em}^{sat}$$

Sun-Earth distance=1.5 x 10⁸ km

Sun Diameter = 1,390,600 km

$$\pi R_{sat}^2 C_0 + \pi R_{sat}^2 \sigma T_E^4 \sin^2(\theta^*) = \sigma T_{sat}^4 4\pi R_{sat}^2$$

$$T_{sat} = \left[\frac{C_0}{4\sigma} + \frac{T_E^4 \sin^2(\theta^*)}{4} \right]^{1/4} = \left[60.7 \times 10^8 + 6.1 \times 10^8 \right]^{1/4} = 286 \ K$$

Strong temperature gradient are therefore expected to be experienced by satellite -> proper design needed

Problem 2

A ground-based radiometer operating at $\lambda = 0.45 \mu m$ is used to measure the solar intensity $I_{\lambda}(0)$. For a solar zenith angle of $\theta = 30^{\circ}$, $I_{\lambda}(0) =$ $1.74 \times 10^7 \ Wm^{-2} \mu m^{-1} sr^{-1}$, and for $\theta =$ 60° , $I_{\lambda}(0) = 1.14 \times 10^{7} Wm^{-2} \mu m^{-1} sr^{-1}$ is measured. From this information, determine the top-of-the-atmosphere solar intensity S_{λ} and the atmospheric optical thickness τ_{λ} .



TOA

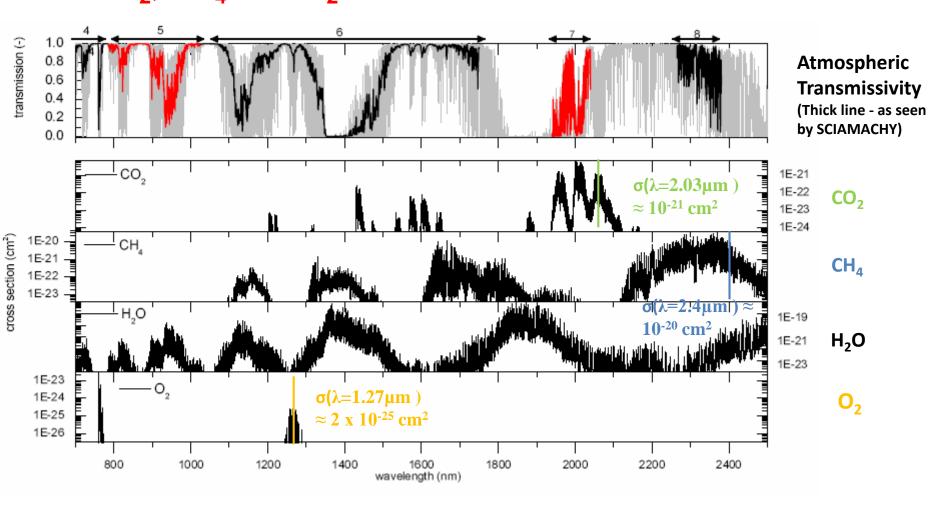
Beer Lambert's law

$$I_{ground}(\theta) = I_{TOA}e^{-\frac{\tau_{nadir}}{\cos(\theta)}} = \begin{cases} I_{TOA}e^{-\frac{2\tau_{nadir}}{\sqrt{3}}} & \theta = 30^{\circ} \\ I_{TOA}e^{-2\tau_{nadir}} & \theta = 60^{\circ} \end{cases} \longrightarrow \frac{I_{ground}(\theta = 30^{\circ})}{I_{ground}(\theta = 60^{\circ})} = e^{\left(2 - \frac{2}{\sqrt{3}}\right)\tau_{nadir}}$$

$$\tau_{nadir} = \ln \left[\frac{I_{ground} \left(\theta = 30^{0} \right)}{I_{ground} \left(\theta = 60^{0} \right)} \right] / \left(2 - \frac{2}{\sqrt{3}} \right) = \ln \frac{1.74}{1.14} / 0.85 = 0.5$$

$$I_{TOA} = I_{ground} (\theta = 60^{\circ}) e^{2\tau_{nadir}} = I_{ground} (\theta = 60^{\circ}) e = 3.1 \times 10^{7} \text{ W}/(m^{2} \text{ } \mu m \text{ } sr)$$

Estimate the transmissivity of well mixed gases CO_2 , CH_4 and O_2 in the near-infrared



Estimate the transmissivity of well mixed gases CO_2 , CH_4 and O_2 in the near-infrared

Transmissivity \mathcal{T} is given by (ignoring scattering)

$$T(\lambda) = \exp{-\{\sigma(\lambda) c L\}}$$

Or more accurate:
$$T(\lambda) = \exp{-\int_{L} \sigma(\lambda, p(L), T(L)) \times c(L) dL}$$

Which can be approximated by :
$$T(\lambda) = \exp{-\left\{\sigma(\lambda, p, T) \times \chi \times VCD_{air}\right\}}$$

With vertical column density of air (molecules per area):

Well Mixed Gases:

- CO₂: χ = 380 ppm and $\sigma(\lambda$ = 2.03 µm) = 1E-21 cm² -> \mathcal{T} = exp{-8.1} = 3E-3
- CH₄: χ = 1800 ppb and $\sigma(\lambda$ = 2.4 µm) = 1E-20 cm² -> \mathcal{T} = exp{-0.39} = 0.68
- O₂: $\chi = 0.2$ and $\sigma(\lambda = 1.27 \,\mu\text{m}) = 2\text{E}-25 \,\text{cm}^2$ -> $\mathcal{T} = \exp\{-0.86\} = 0.42$

Compare the atmospheric transmissivity for Rayleigh scattering in UV with larger wavelength (600nm, 900 nm)

Rayleigh Cross Section

λ, nm	σ, cm ²
300	6.00 E-26
400	1.90 E-26
600	3.80 E-27
1000	4.90 E-28
10,000	4.85 E-32

Compare the atmospheric transmissivity for Rayleigh scattering in UV with larger wavelength (600nm, 900 nm)

Rayleigh Cross Section

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Transmissivity (ignoring absorption or Mie scattering)

$$\mathcal{T}(\lambda) = I_t(\lambda) / I_0(\lambda)$$

$$= \exp{-\tau} = \exp{-k_s(\lambda) L}$$

$$= \exp{-\sigma c L} = \exp{-\sigma VCD}$$

with Vertical Column Density of air (VCD) = 2.14 E25 molec/cm2

Transmissivity:
$$\mathcal{T}(\lambda=300\text{nm}) = \exp\{-6\text{E}-26 \text{ x } 2.14\text{E}25\} = \exp\{-1.3\} = 0.27$$

$$\mathcal{T}(\lambda=600\text{nm}) = \exp\{-0.08\} = 0.92$$

$$\mathcal{T}(\lambda=1000\text{nm}) = \exp\{-0.01\} = 0.99$$

Only 25% of direct light reaches surfaces in UV but 99% at 1000nm.