## Tri-Layer Setup

# Maxime Delorme, Sacha Brun, Adam Finley ${\rm May}\ 2024$

#### 1 Domain

The domain is a cartesian slab. Vertical direction is indicated by the z axis. The top of the domain is located at  $z_{min}$  while the bottom of the domain is located at  $z_{max}$ . Increasing z means going further down in the domain. Hence gravity will be positively signed. The vertical decomposition of the domain consists in three zones:

- 1. a stable zone extending in  $z \in [z_0 = z_{min}, z_1]$  of size  $\Delta z_0 = z_1 z_{min}$ ,
- 2. a convective zone in  $z \in [z_1, z_2[$  of size  $\Delta z_1 = z_2 z_1,$
- 3. a final stable zone spanning  $z \in [z_2, z_{max}]$  of size  $\Delta z_2 = z_{max} z_2$

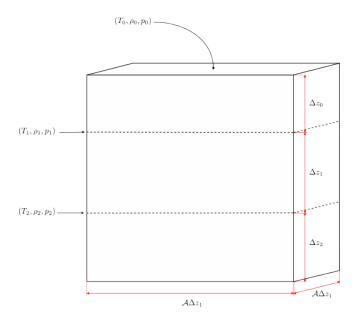


Figure 1: Description of the domain and the associated initial quantities

In the following each quantity subscripted by a number refers to the value of the quantity at the top of the corresponding layer. For instance,  $\rho_1$  indicates the value of the density at the top of the convective zone. The localization of the quantities is described on Figure 1.

Along the horizontal axis (x, and y), the domain is periodic and spanning a certain surface which is proportional to the aspect ratio  $\mathcal{A}$  between the horizontal span and the vertical span of the convective layer. Hence, if we set  $x_{min} = y_{min} = 0$ , we have:

$$x_{max} = y_{max} = \mathcal{A}\Delta z_1 \tag{1}$$

#### Free parameters:

- The size of each zone  $\Delta z_i$
- The aspect ratio of the box A

## 2 Hydrodynamic State

Each zone of the domain is initialized as a polytropic model. For a position in a given zone  $i: z \in [z_i, z_{i+1}]$  we have the following initial conditions on the temperature T, the density  $\rho$  and the pressure P:

$$T(z) = T_i + \theta_i(z - z_i) \tag{2a}$$

$$\rho(z) = \rho_i \left(\frac{T(z)}{T_i}\right)^{m_i} \tag{2b}$$

$$P(z) = P_i \left(\frac{T(z)}{T_i}\right)^{m_i + 1} \tag{2c}$$

Where  $\theta_i$  is the temperature gradient in the zone and  $m_i$  the polytropic index associated with the layer. Finally,  $T_i$ ,  $\rho_i$  and  $P_i$  are respectively the temperature, density and pressure at the top of the layer.

Given the value of the temperature at the top of the domain  $T_1$  and the density at the top of the domain  $\rho_1$  we can calculate the values of the  $T_i$  and  $\rho_i$  for i = 1, 2:

$$T_1 = T_0 + \theta_0 \Delta z_0$$
  $\rho_1 = \rho_0 \left(\frac{T_1}{T_0}\right)^{m_1}$  (3)

Pressure is recovered using the ideal gas equation of state with a gas coefficient  $\mathcal{R} = 1$ , having thus  $P = \rho T$ .

Gravity is kept constant in the domain. It is directed strictly towards positive z and is defined as

$$q = \theta_i(m_i + 1) \ \forall i = 0, 1, 2 \tag{5}$$

This puts a strong constraint on the values of each couple  $(\theta_i, m_i)$  since as soon as g is determined by one of these couple, the values of  $\theta_i$  will be automatically inferred from the values of the polytropic indices  $m_i$ . We decide to give the values of  $m_i$  to the user, and any of the  $\theta_i$ . As soon as one value of  $\theta$  is picked, all the other are automatically fixed.

#### Free parameters:

- The state at the top of the box  $\rho_0$ ,  $T_0$
- The polytropic indices  $m_i$
- One of the temperature gradients  $\theta_i$

#### 3 Diffusive effects

To this hydrodynamic setup should be added thermal diffusion and viscosity. The user provides the thermal dissipation coefficient  $C_k$ . This coefficient, along-side the specific heat ratio  $\gamma$  allows the calculation of the thermal conductivity K:

$$K = \frac{\gamma}{\gamma - 1} C_k \tag{6}$$

This thermal conductivity is then "specialized" for each zone depending on their own polytropic index:  $K_0$ ,  $K_1$ ,  $K_2$ . By default, we consider that  $K_1 = K$ . By following the litterature, we have the relations:

$$\frac{K_0}{K_1} = \frac{m_0 + 1}{m_1 + 1} \qquad \qquad \frac{K_2}{K_1} = \frac{m_2 + 1}{m_1 + 1}$$

Hence we get:

$$K_0 = K_1 \left( \frac{m_0 + 1}{m_1 + 1} \right) \tag{7a}$$

$$K_2 = K_1 \left( \frac{m_2 + 1}{m_1 + 1} \right) \tag{7b}$$

The kinematic viscosity coefficient  $\mu$  is constant throughout the box and defined from the Prandtl number  $\sigma$ . In the normalized system, we have the following relation :

$$\mu = \sigma C_k \tag{8}$$

Free parameters:

- $\bullet$   $C_k$  the thermal dissipation coefficient
- $\sigma$  the Prandtl number

## 4 Free parameters and calculations

If we sum the free parameters to define in the run we have the following set :

- The size of each zone  $\Delta z_0$ ,  $\Delta z_1$ ,  $\Delta z_2$
- The aspect ratio of the box A
- The state at the top of the box  $\rho_0$ ,  $T_0$
- The polytropic indices  $m_i$
- One of the temperature gradients  $\theta_i$
- $C_k$  the thermal dissipation coefficient
- $\bullet$   $\sigma$  the Prandtl number

Most of these parameters are left to the user. However defining  $T_0$  and  $rho_0$  for a given run might prove difficult. Since these points correspond to an arbitrary "top" of a solar atmosphere.  $\rho_0$  is the density scaling parameter, and thus should be set depending on the physics/units of the run. Conveniently it can be left to 1, since it does not particularly affect the simulation.

 $T_0$  is more difficult. One way to define it is to use the density ratio between the top and the bottom of the convection zone as a proxy. This ratio is noted  $\chi_{\rho} = \rho_2/\rho_1$ . We can find an explicit expression for  $\chi_{\rho}$  by injecting (3) and (4) and simplifying. We get the following expression:

$$T_0 = \frac{\theta_1 \Delta z_1}{\chi_\rho^{1/m_1}} - \theta_0 \Delta z_0 \tag{9}$$