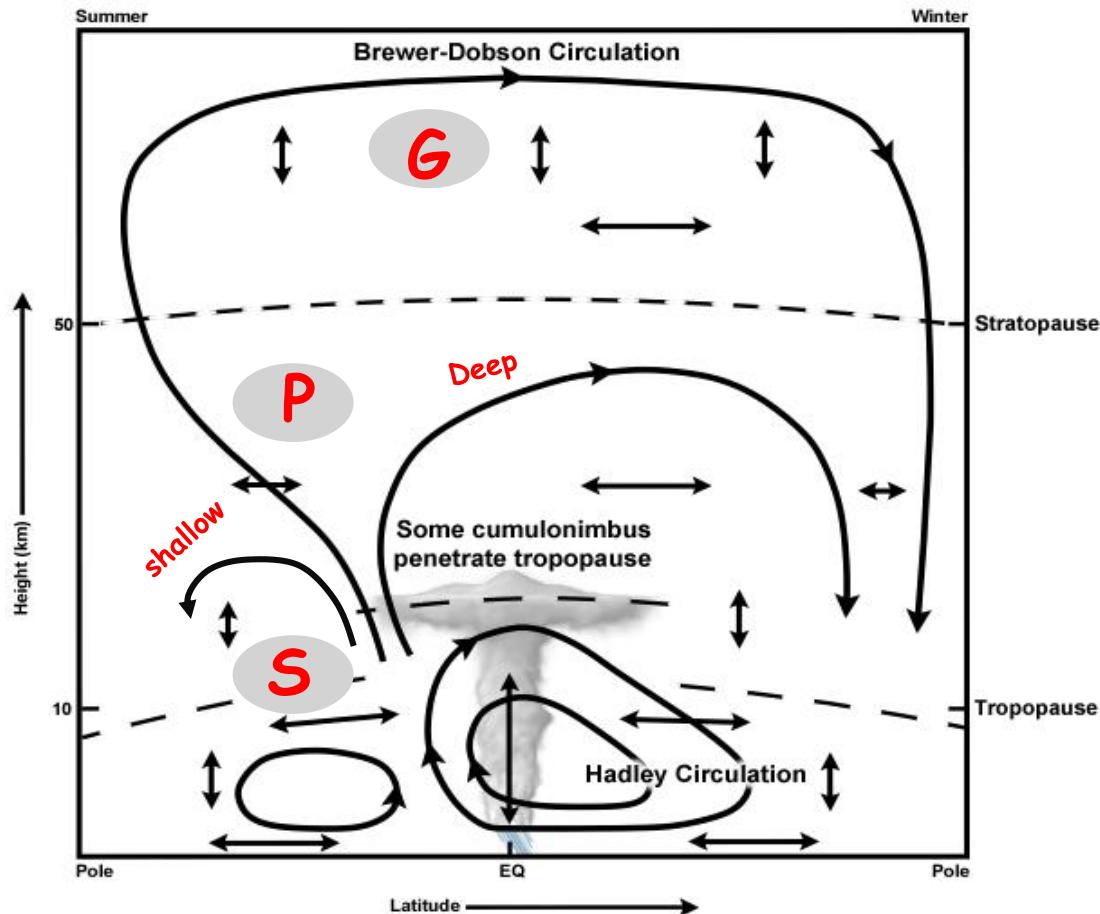


NATURAL CLIMATE VARIABILITY AND STATISTICAL ANALYSIS APPROACH

MOHAMADOU DIALLO

May 29th 2020 | Atmospheric modeling lecture | University of Wuppertal, Germany

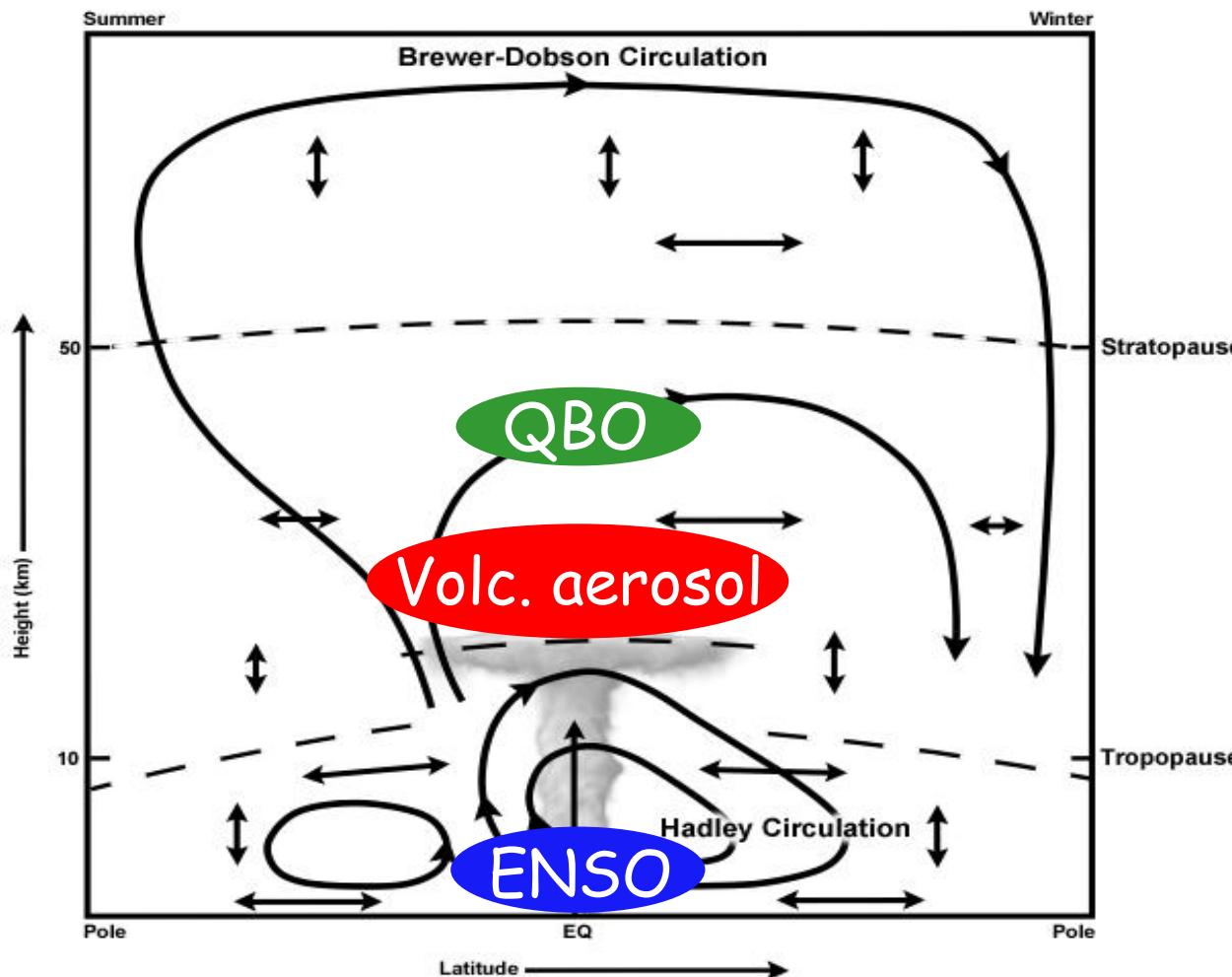
What is the Brewer-Dobson Circulation (BDC)?



[WMO, 1985]

- ❖ **BDC** is the transport of air mass from the inner tropics to the higher latitudes via **transition**, **shallow** and **deep** branches [Brewer, 1949].
- ❖ **BDC** is driven by the **waves breaking** [Plumb, 2002]:
 - ✓ Gravity waves (G)
 - ✓ Planetary-scale waves (P)
 - ✓ Synoptic-scale waves (S)

Modulations of the BDC by natural variability

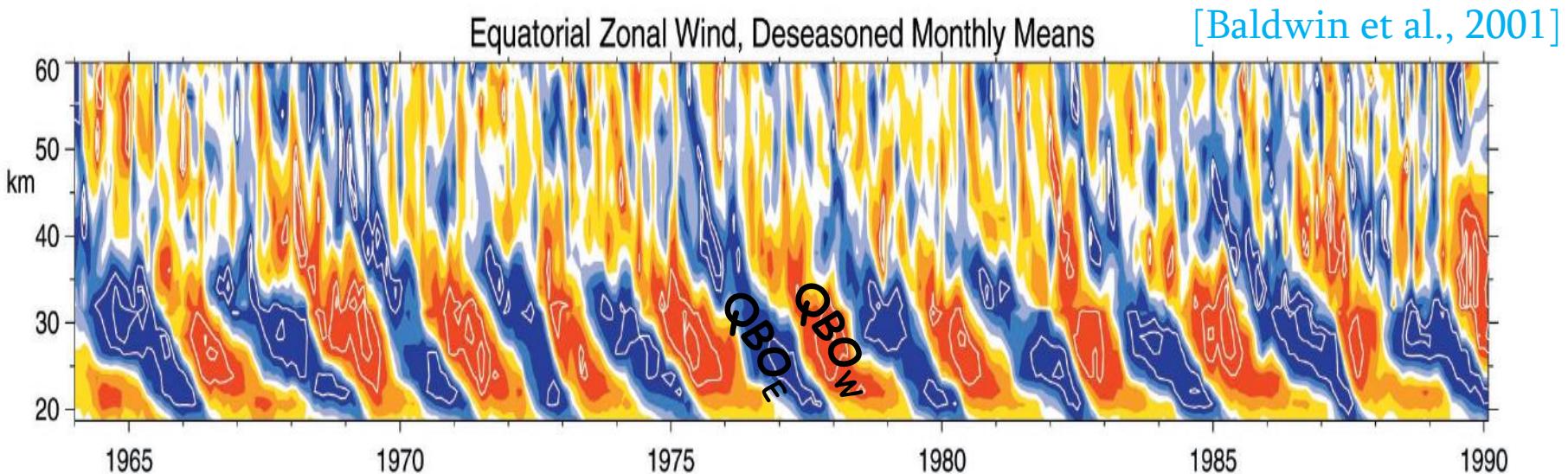


[WMO, 1985]

BDC is modulated by naturally varying signals:

- ❖ **Quasi-Biennial Oscillation phases (QBO)** [Plumb & Bell, QJRMS 1982].
- ❖ **El Niño-Southern Oscillation (ENSO)** [Randel et al., GRL 2009].
- ❖ **Volcanic aerosols** [Robock, GRL 2010].

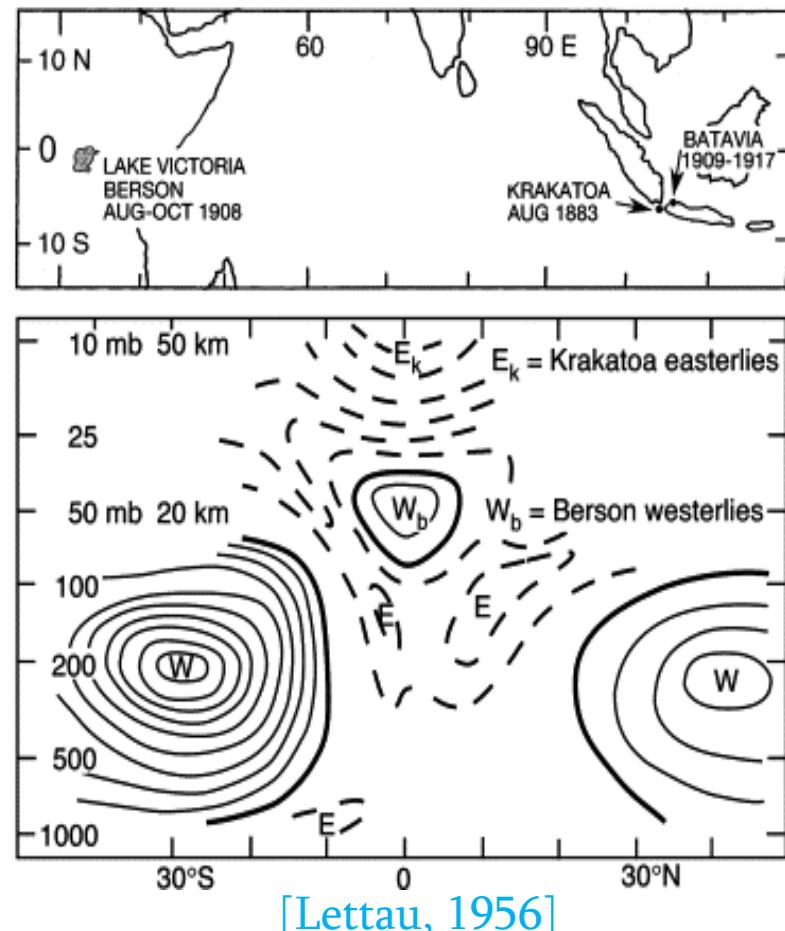
The Quasi-Biennial Oscillation (QBO)



- ✓ The QBO is the mean 28 months oscillation of zonal mean winds.
- ✓ The QBO propagates downward into the lower stratosphere at 1 km/month.
- ✓ Transition between QBO_W and QBO_E regimes is often delayed between 30 and 50 hPa.
- ✓ Both the QBO period and amplitude exhibit considerable variability.

Discovery of the QBO

- ✓ Global spread of Krakatao volcanic plume in Aug 1883 led to a discovery of **Krakatau-easterlies** at 10hPa.
- ✓ In 1910, the discover of **Berson-westerly**, leading to a conundrum of the **Krakatau-easterlies**.
- ✓ In 1924, Van Bemmelen confirmed the existing **Krakatau-easterlies**.
- ✓ Alternating **QBO_{W/E}** was fully described in **1960s** by Edmon and Reed.
- ✓ **Westerly momentum** induced by equatorially trapped **Kelvin waves** and **easterly momentum** by **Rossby-gravity** [Holton & Lindzen, 1972; Plumb, 1977].

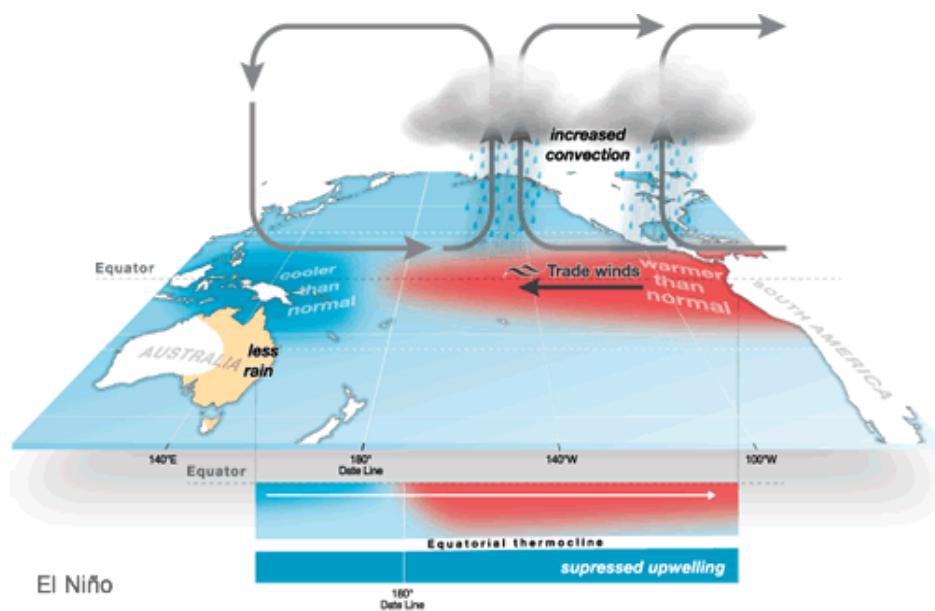


The El Niño Southern Oscillation (ENSO)

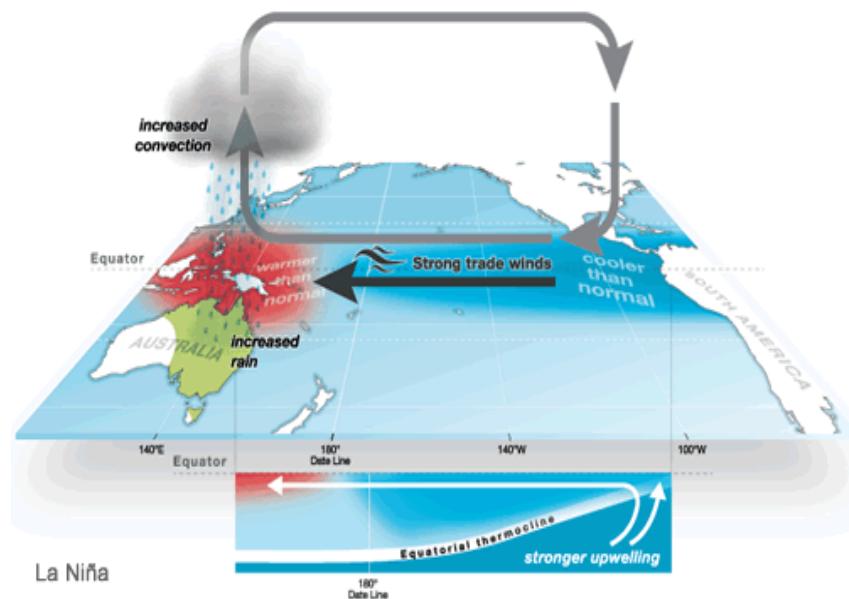
- ✓ ENSO is a coupled ocean-atmosphere phenomenon caused by recurring redistributions of heat and atmospheric momentum in the tropical Pacific [McPhaden, 2002].

The El Niño Southern Oscillation (ENSO)

- ✓ ENSO is a coupled ocean-atmosphere phenomenon caused by recurring redistributions of heat and atmospheric momentum in the tropical Pacific [McPhaden, 2002].
- ✓ The extremes of ENSO, termed **El Niño** and **La Niña**, encompass a wide range of climate conditions [Philander, 1990].

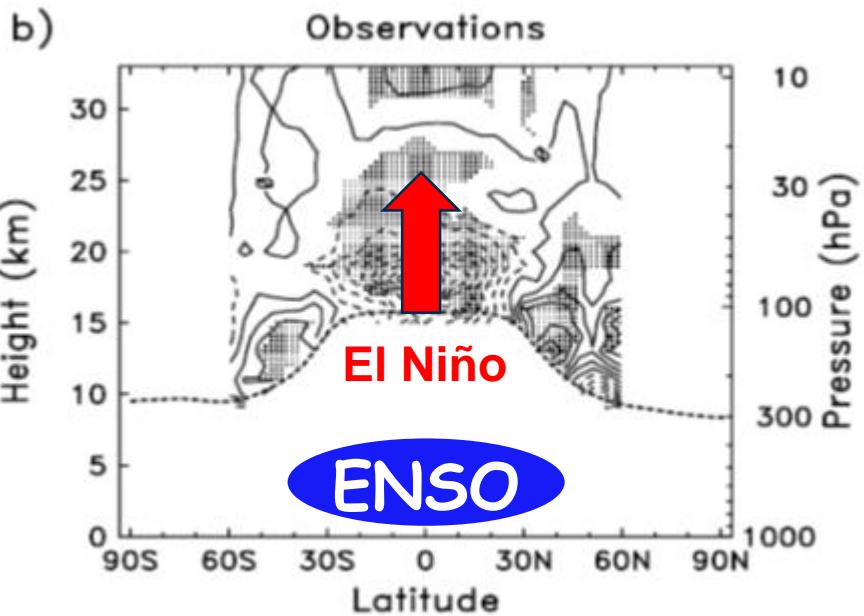


El Niño (anomalously warm SSTs)



La Niña (anomalously cool SSTs)

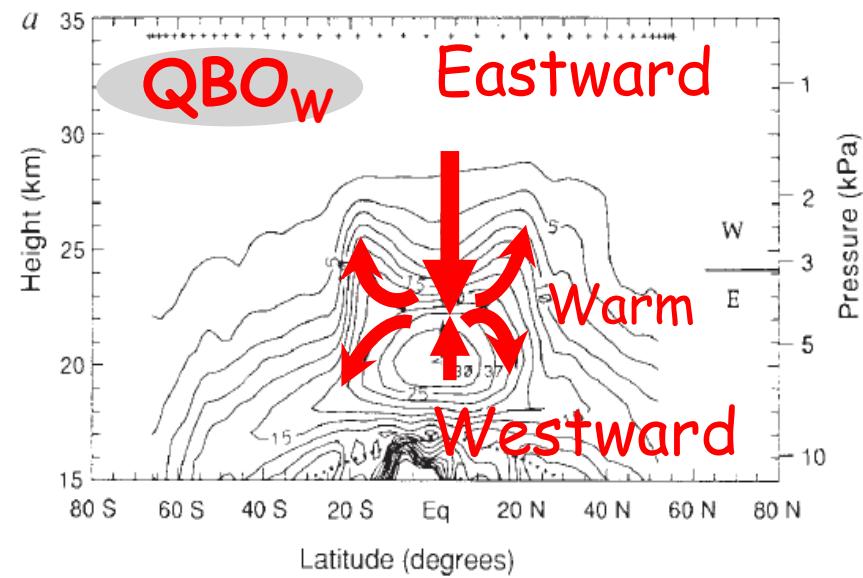
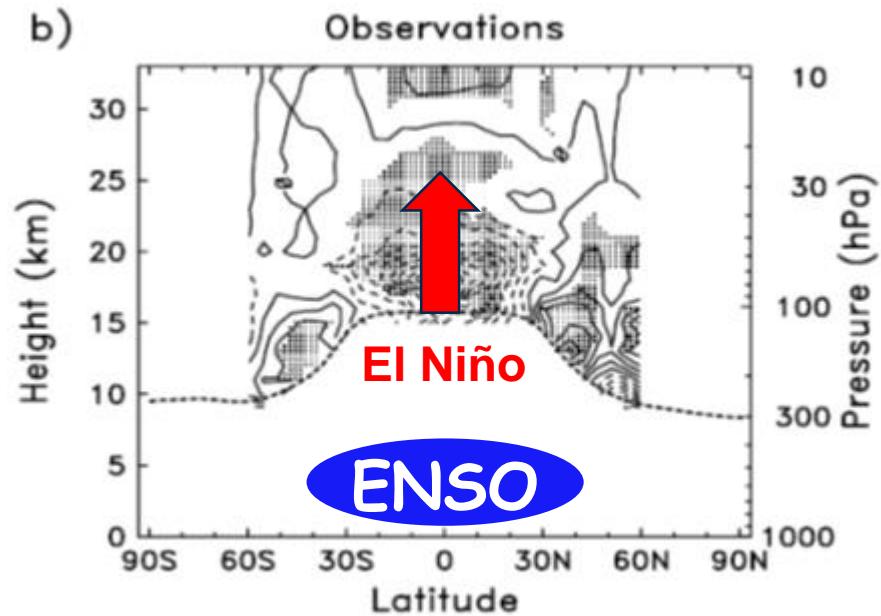
Well-known QBO & ENSO modulations of BDC



- ❖ El Niño enhances the tropical upwelling [Randel et al., 2009; Calvo et al. 2010; Konopka et al., 2016; Diallo et al., 2018].

Well-known QBO & ENSO modulations of BDC

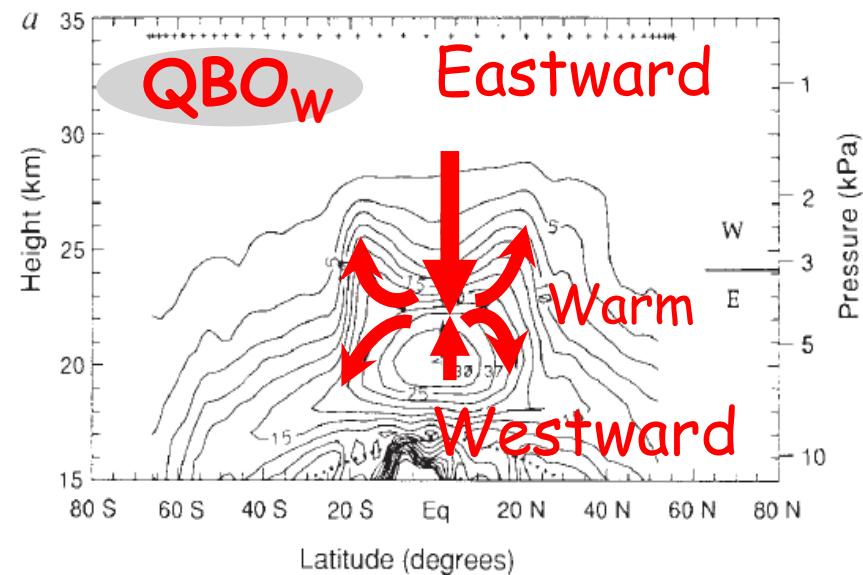
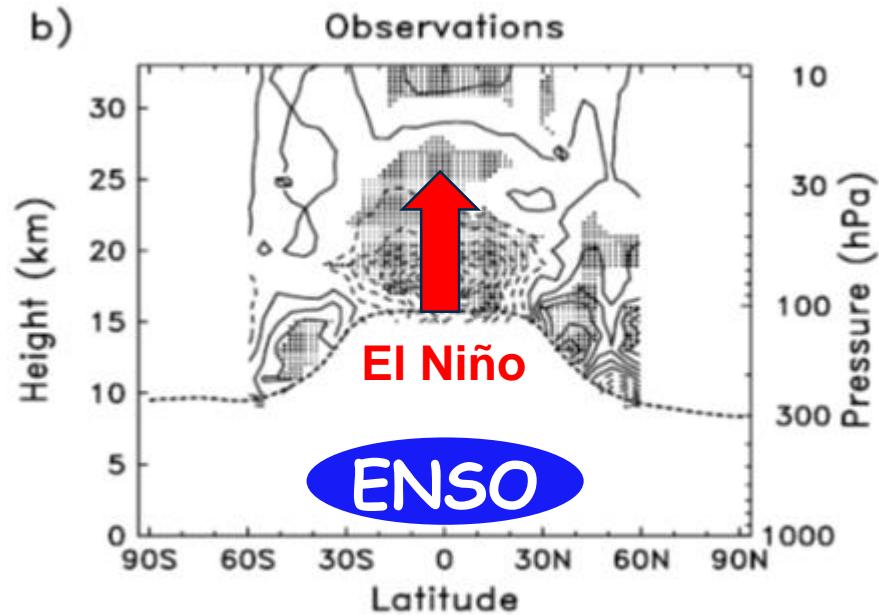
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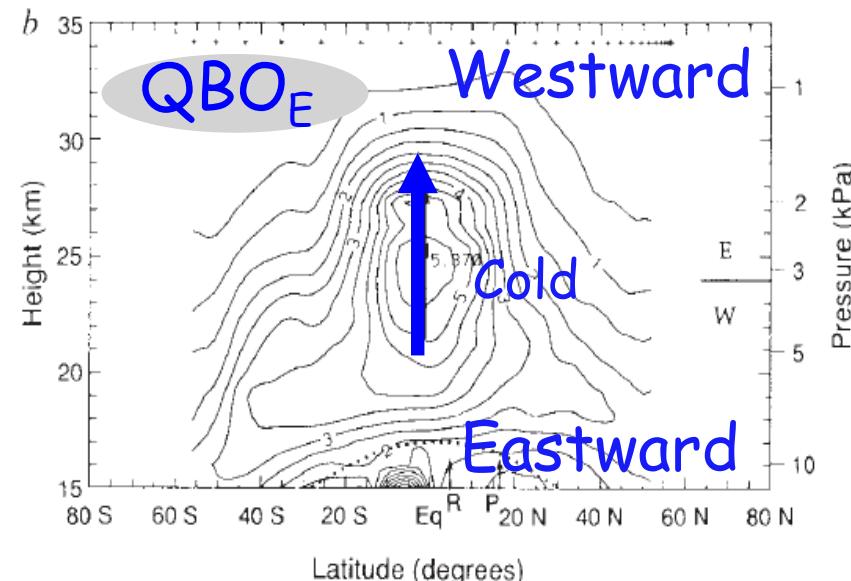
- ❖ El Niño enhances the tropical upwelling [Randel et al., 2009; Calvo et al. 2010; Konopka et al., 2016; Diallo et al., 2018].
- ❖ QBO_W reduces tropical upwelling but enhances horizontal transport.
- ❖ QBO_E enhances tropical upwelling and anomalous cold tropopause [Plumb & Bell, 1982; Trepte et al., 1992]

Well-known QBO & ENSO modulations of BDC

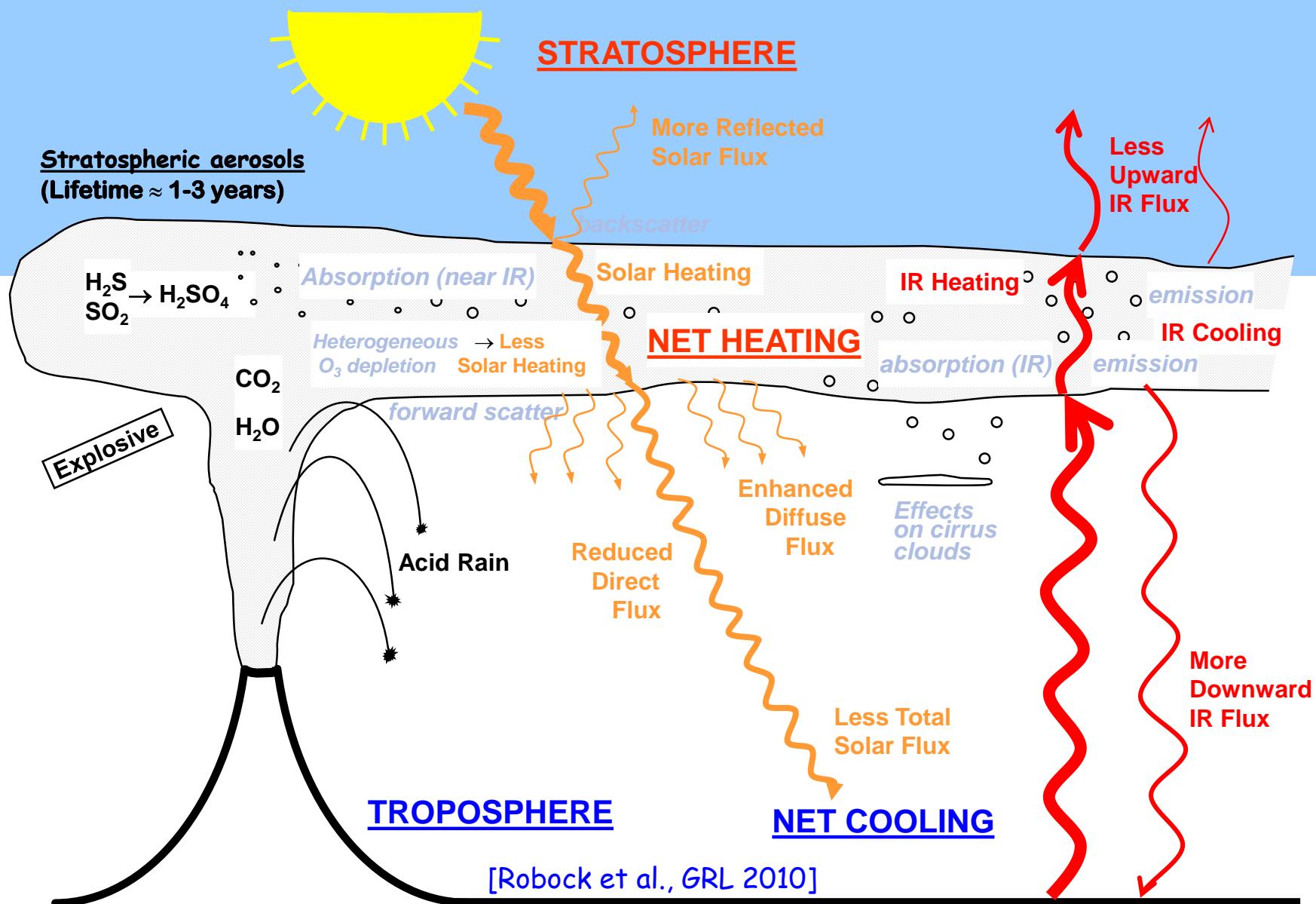
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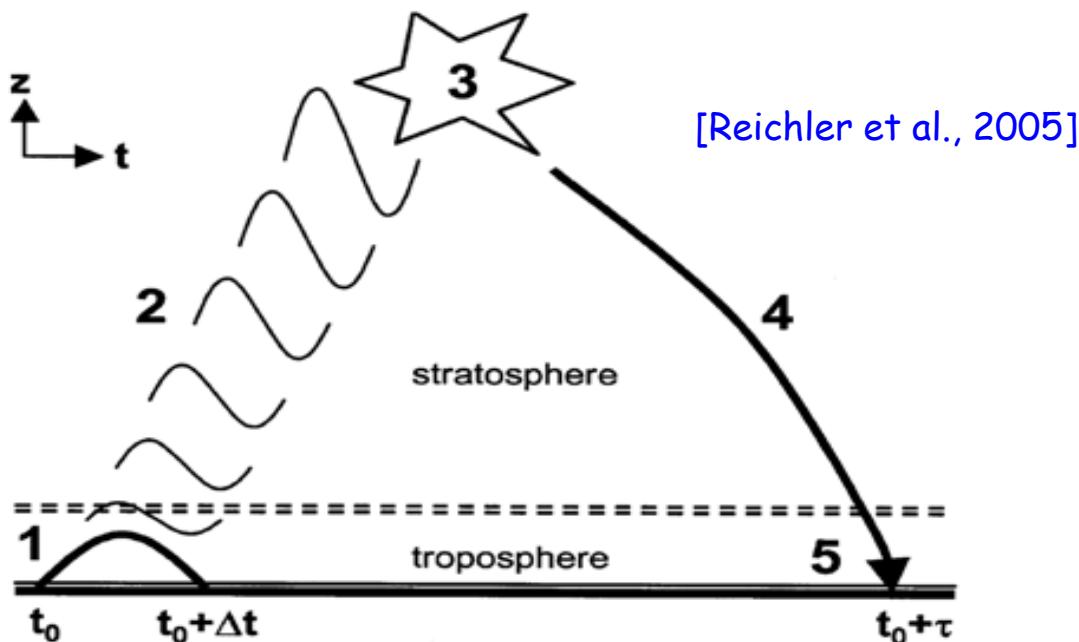
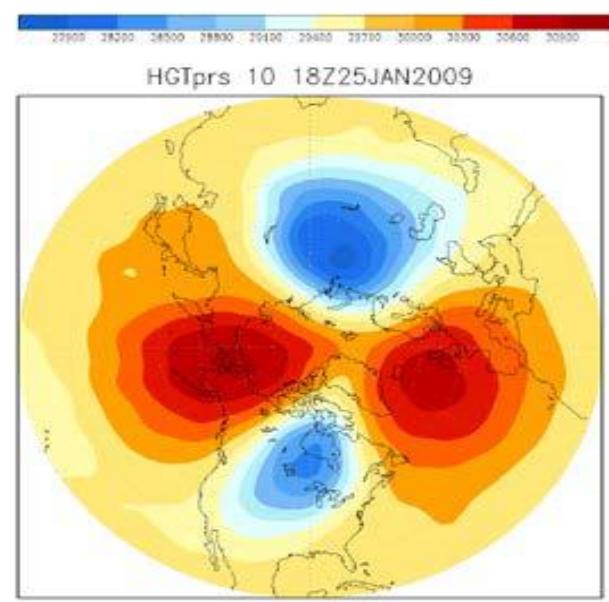
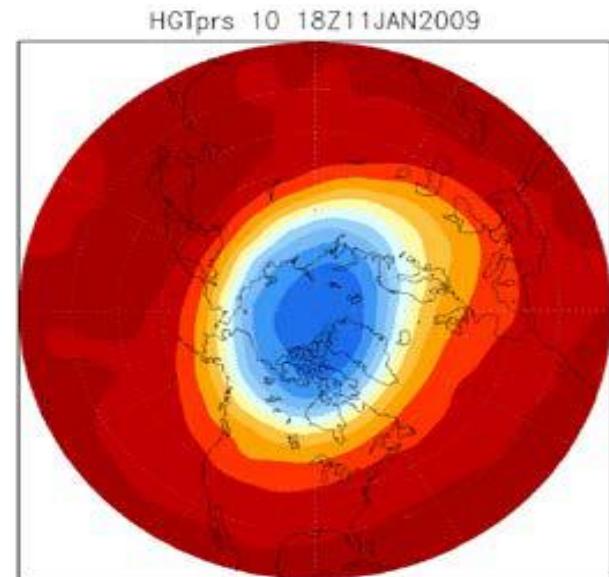


Radiative forcing of the Junge Layer and BDC



Stratos. Sudden Warmings & its tropospheric impact

- ❖ Wintertime stratosphere can undergo large and rapid changes known as **Sudden Warmings**.
- ❖ These are characterized by dramatic changes in **high-latitude wind and temperature**.
- ❖ In the **troposphere**, they have been associated with a change in path of **North-Atlantic weather systems**.

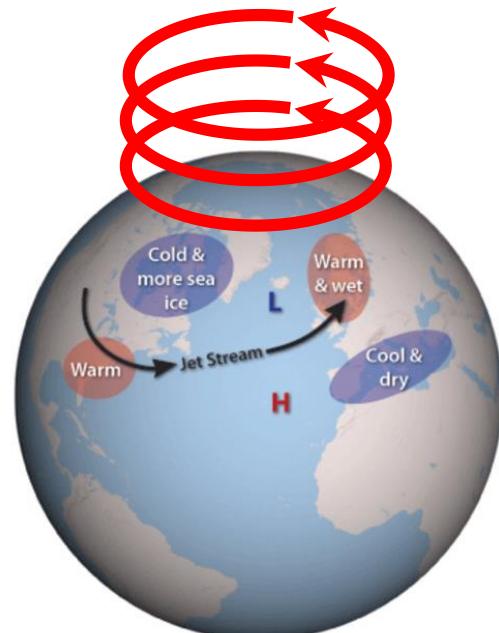


The North Atlantic Oscillation (NAO)

- ❖ Variation of the pressure gradient between Iceland and Azores.
- ❖ Leading mode of large-scale atmospheric circulation patterns over Europe.
- ❖ Control the strength/direction of westerly winds & storm track locations.
- ❖ Phases linked to weather regimes: **Cyclonic (NAO+)** or **Blocking (NAO-)**.

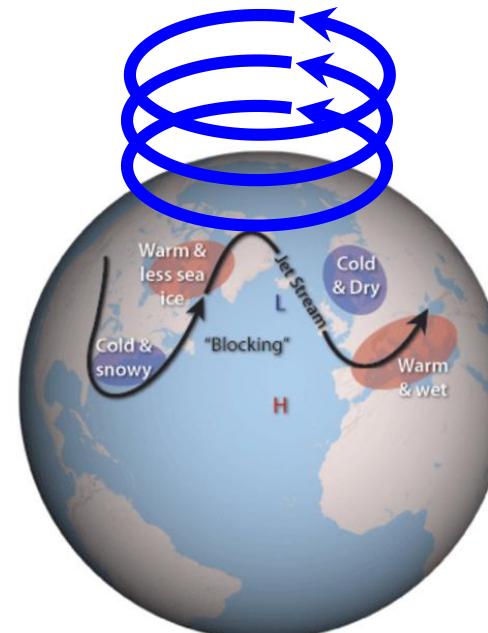
Positive NAO phase

Cyclonic regime



Negative NAO phase

Blocking regime

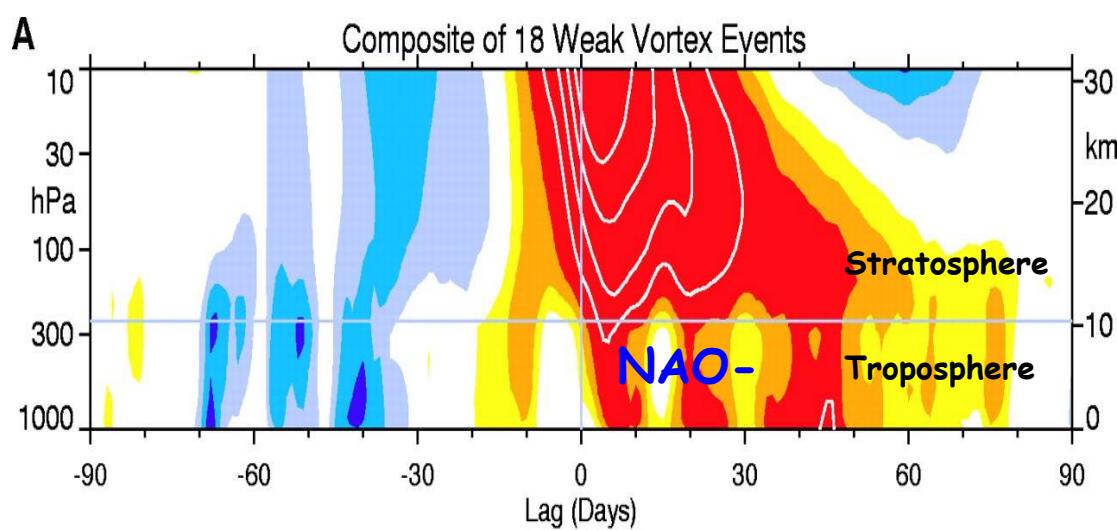


[Cassou, Nat. 2008]
[Grams et al. Nclim 2017]

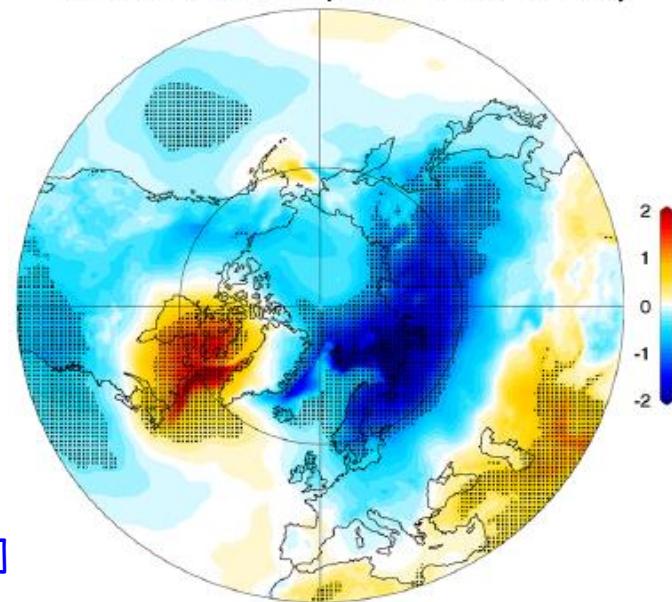
Weather extremes related to Stratospheric Variability

- ❖ Stratospheric variability modulates the **European weather**
- ❖ Winter weather extremes (low temperatures, snow, etc.) are much more common during **weak vortex events**.
- ❖ Atlantic blocking occurs almost exclusively during **weak vortex events**.
- ❖ Strong winds and ocean wave events are much more common during **strong vortex events**.
- ❖ Composites of the **60 days sfc T anomalies (K)** after historical SSWs in the JRA-55 reanalyses.

[Baldwin & Dunkerton, Sci. 2001]



(b) Surface temperature anomaly



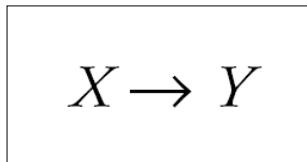
[Butler et al., 2017]

4. Statistical models

“Essentially, all models are wrong, but some are useful.”

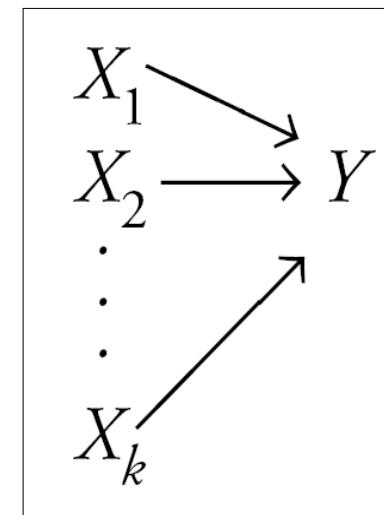
General idea

Simple regression considers the relation between a single explanatory variable and response variable



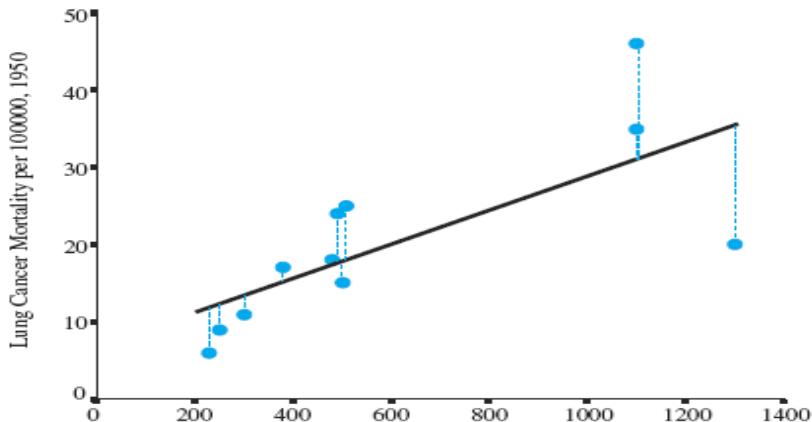
Multiple regression simultaneously considers the influence of multiple explanatory variables on a response variable Y

The intent is to look at the independent effect of each variable X_i , while “adjusting out” the influence of potential confounders $X_j (j \neq i)$



General idea

- A simple regression model (one independent variable) fits a regression *line* in 2-D space



- A multiple regression model with two explanatory variables fits a regression plane in 3-D space

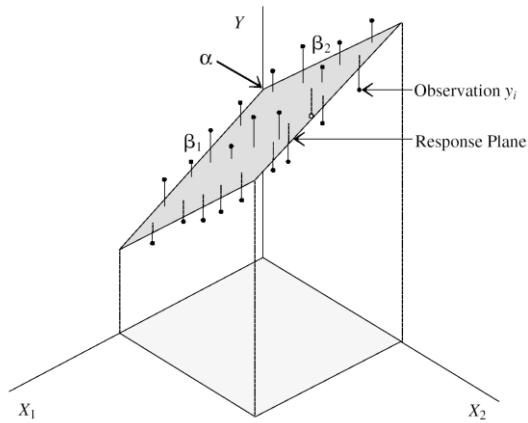
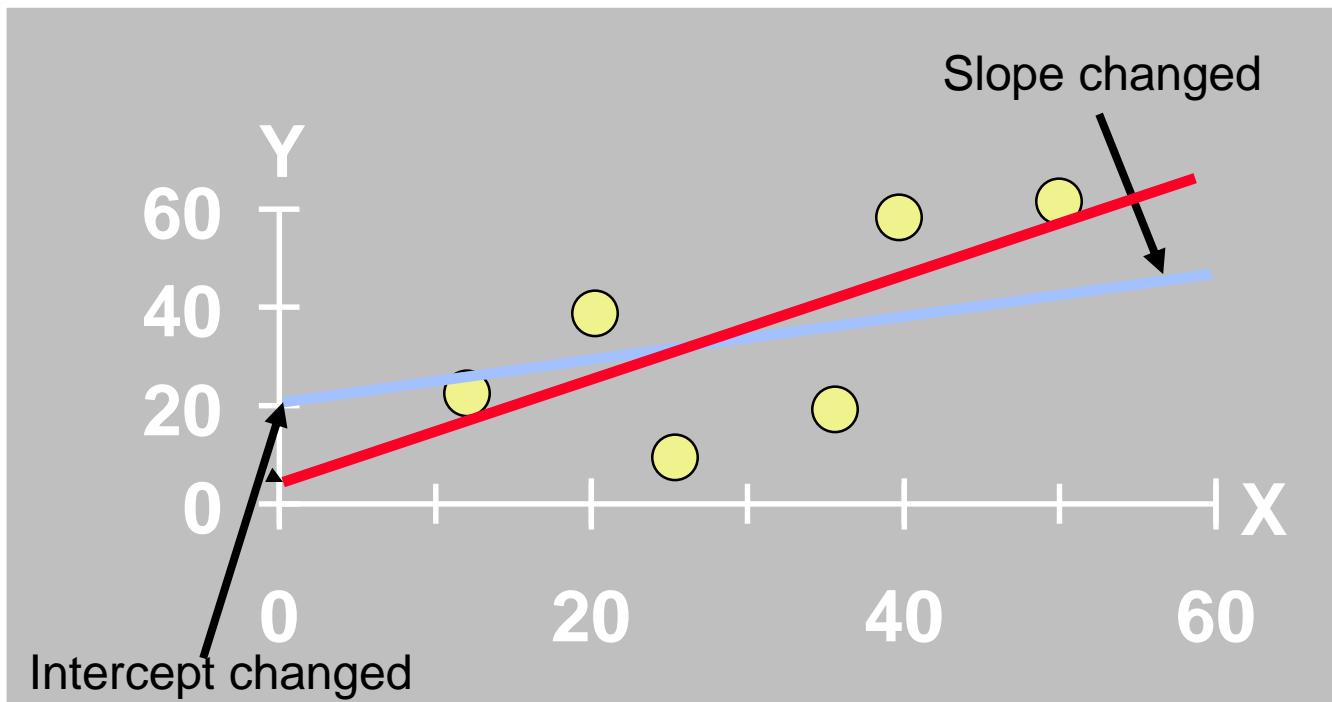


FIGURE 15.1 Three-dimensional response plane.

Thinking Challenge

How would you draw a line through the points?
How do you determine which line ‘fits best’?



Simple linear regression model

- For n given observations, a linear regression model can be written as

$$Y_i = \beta_0 + \beta_1 X_i + \varepsilon_i$$

Diagram illustrating the components of the simple linear regression model:

- Dependent (Response) Variable** points to Y_i .
- Independent (Explanatory) Variable** points to X_i .
- Observation Y-Intercept** points to β_0 .
- Observation Slope** points to β_1 .
- Random Error** points to ε_i .

Simple linear regression model

- The previous formulas can transform into matrix notation

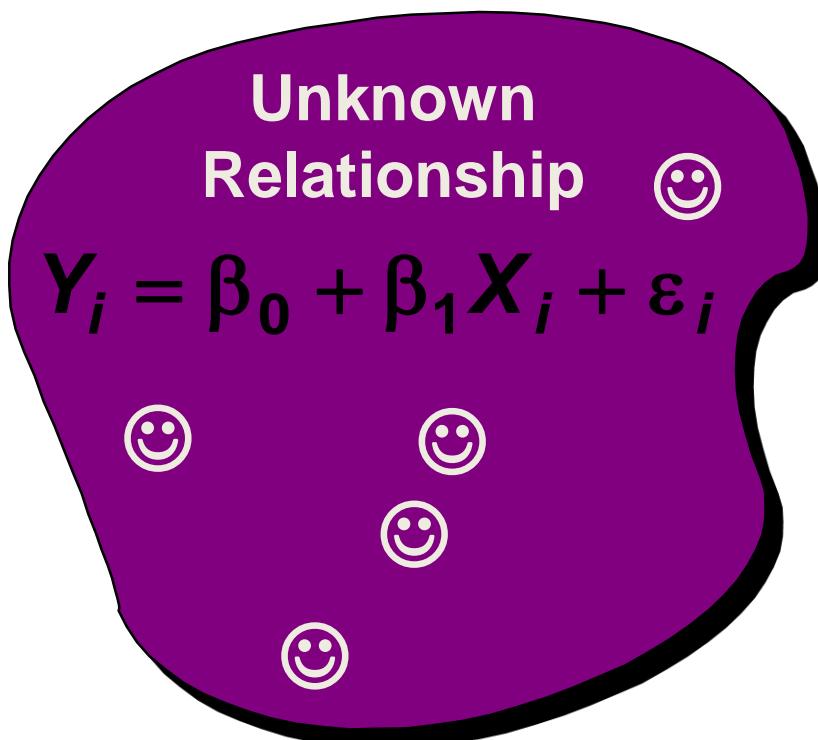
$$Y = X\beta + \varepsilon$$

$$Y = \begin{bmatrix} Y_1 \\ Y_2 \\ \vdots \\ \vdots \\ Y_n \end{bmatrix}; \quad X = \begin{bmatrix} 1 & X_1 \\ 1 & X_2 \\ \vdots & \vdots \\ \vdots & \vdots \\ 1 & X_n \end{bmatrix}; \quad \beta = \begin{bmatrix} \beta_0 \\ \beta_1 \end{bmatrix} \quad \text{et} \quad \varepsilon = \begin{bmatrix} \varepsilon_1 \\ \varepsilon_2 \\ \vdots \\ \vdots \\ \varepsilon_n \end{bmatrix}$$

Observation & sample regression model

- To estimate the parameters, we randomly sample the observations as following:

Observations



$$\beta_0 \quad \beta_1 \quad \sigma^2$$

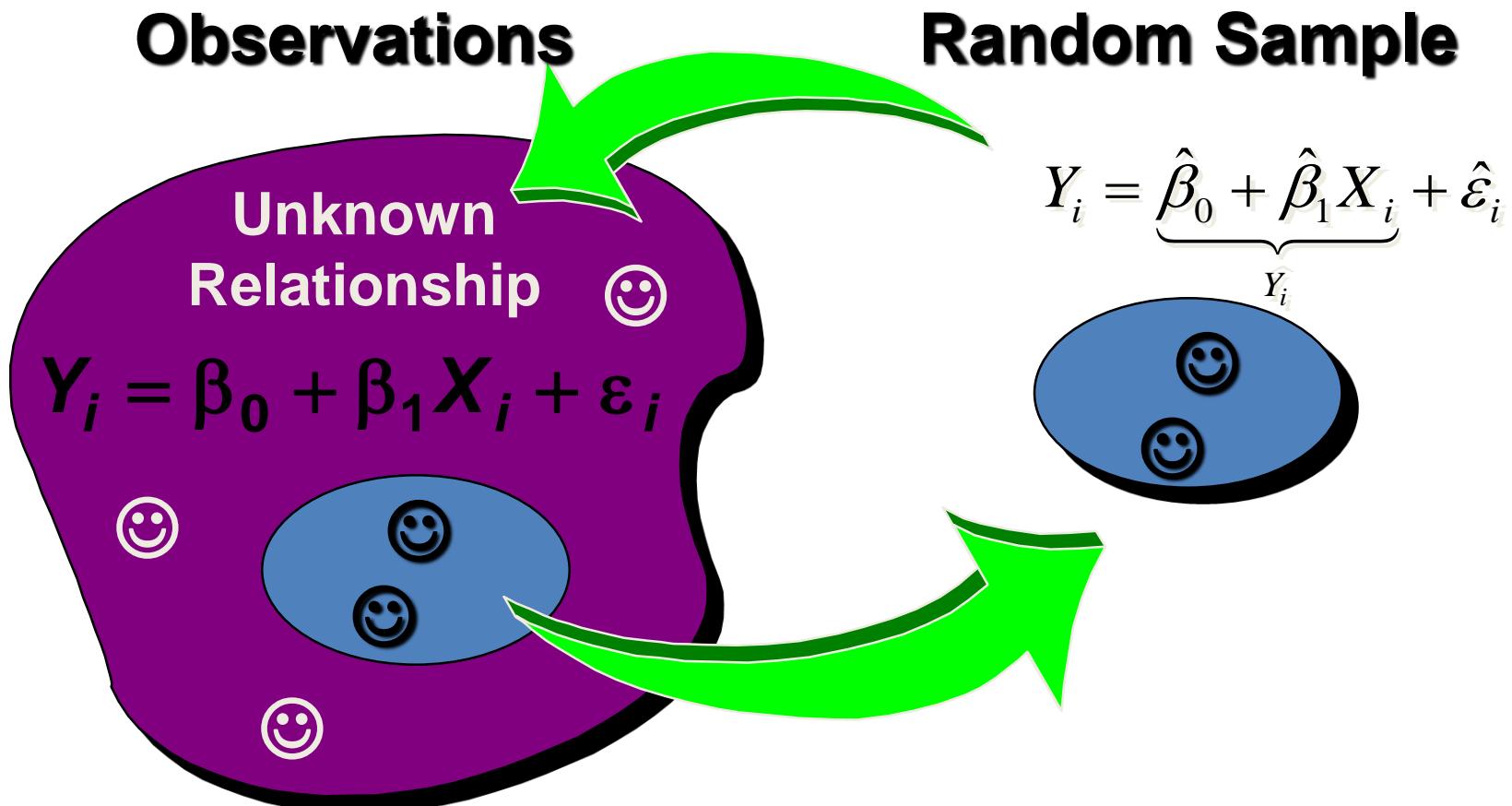
$$\hat{\beta}_0 = E(\beta_0)$$

$$\hat{\beta}_1 = E(\beta_1)$$

$$\sigma^2 = \text{Var}(\varepsilon_i)$$

Observation & sample regression model

- To estimate the parameters, we randomly sample the observations as following:



Least squares method

- **Least squares method** (LS) is used to minimize the sum of the squared errors (SSE).

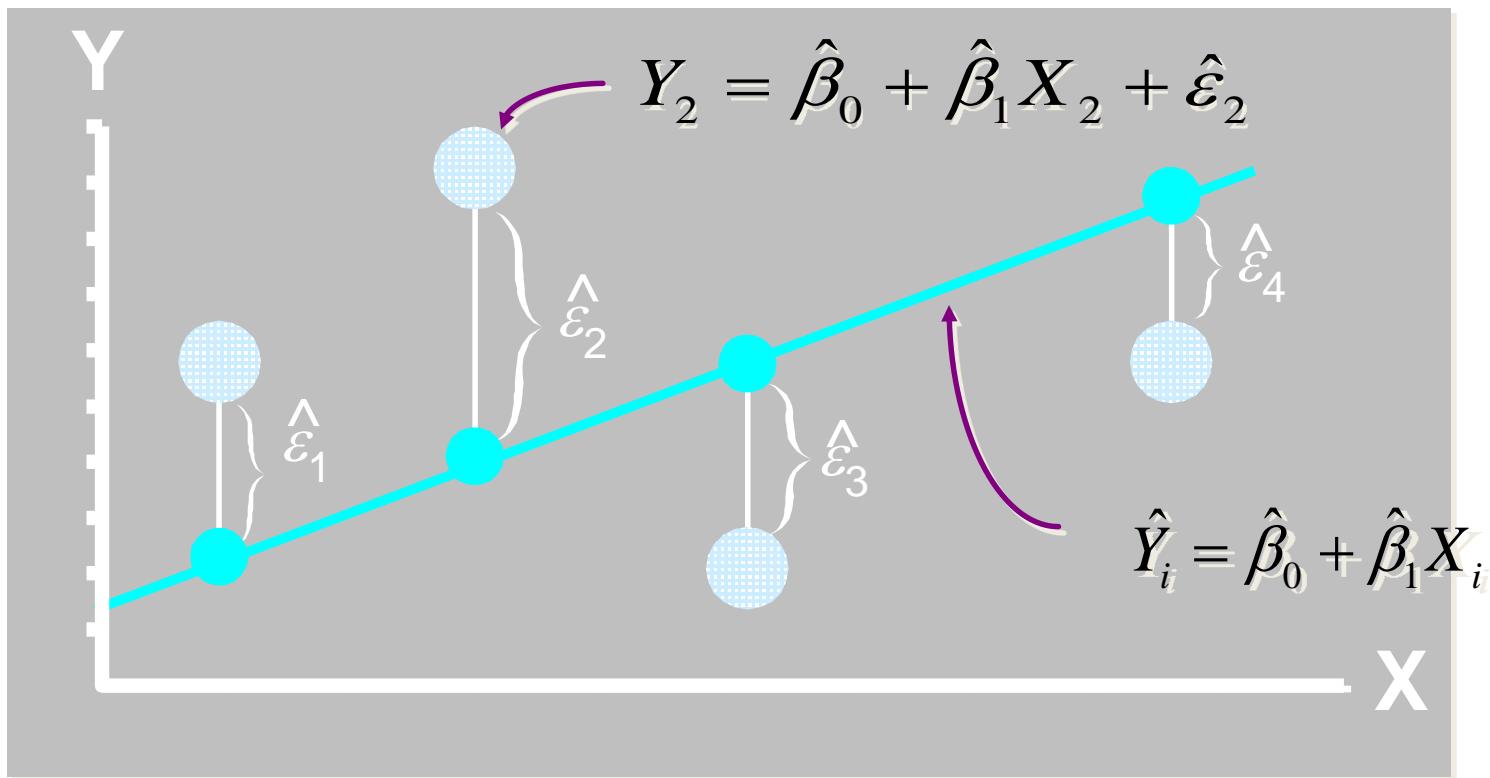
$$\sum_{i=1}^n (Y_i - \hat{Y}_i)^2 = \sum_{i=1}^n \hat{\epsilon}_i^2$$

- “**Best Fit**” means difference between actual **Y** values and predicted **Yhat** values are a **minimum**. *But* positive differences off-set negative. So square errors!

Least squares method

- Graphical representation of LS method.

LS minimizes $\sum_{i=1}^n \hat{\varepsilon}_i^2 = \hat{\varepsilon}_1^2 + \hat{\varepsilon}_2^2 + \hat{\varepsilon}_3^2 + \hat{\varepsilon}_4^2$



Estimating Parameters

- Parameters are estimated using LS method, which minimizes the SSE.

$$(\beta_0, \beta_1) = \operatorname{Arg} \min_{(\beta_0, \beta_1) \in \mathbb{R}^2} \left[\sum_{i=1}^n \varepsilon_i^2 \right]$$

$$= \operatorname{Arg} \min_{(\beta_0, \beta_1) \in \mathbb{R}^2} \sum_{i=1}^n (Y_i - \beta_0 - \beta_1 X_i)^2$$

$$= \min F(\beta_0, \beta_1)$$

$$\begin{cases} \frac{\partial F(\beta_0, \beta_1)}{\partial \beta_0} \Bigg|_{\beta_0=\hat{\beta}_0, \beta_1=\hat{\beta}_1} = 0, \\ \frac{\partial F(\beta_0, \beta_1)}{\partial \beta_1} \Bigg|_{\beta_0=\hat{\beta}_0, \beta_1=\hat{\beta}_1} = 0, \end{cases}$$

- Solve the equation system

Estimating Parameters

- Prediction equation

$$Y_i = \hat{\beta}_0 + \hat{\beta}_1 X_i$$

- Sample slope

$$\hat{\beta}_1 = \frac{SS_{xy}}{SS_{xx}} = \frac{\sum (X_i - \bar{X})(Y_i - \bar{Y})}{\sum (x_i - \bar{X})^2} = \frac{Cov(X, Y)}{Var(X)}$$

- Sample Y - intercept

$$\hat{\beta}_0 = \bar{Y} - \hat{\beta}_1 \bar{X},$$

Coeficcident of determination

- It is a measure of the **overall quality** of the regression.
- Specifically, it is the percentage of **total variation** exhibited in the Y_i data that is accounted for by the sample regression line.

- The sample mean of Y : $\bar{Y} = \frac{1}{n} \sum_{i=1}^n Y_i$
- Total variation in Y : $= \sum_{i=1}^n (Y_i - \bar{Y})^2$
- Residual (unaccounted) variation in Y : $= \sum_{i=1}^n (Y_i - \hat{Y}_i)^2 = \sum_{i=1}^n \hat{\varepsilon}_i^2$

Coeficcident of determination

$$R^2 = \frac{\text{variation accounted for by x variables}}{\text{total variation}}$$
$$= 1 - \frac{\text{variation not accounted for by x variables}}{\text{total variation}}$$

$$= 1 - \frac{\sum_{i=1}^n (Y_i - \hat{Y}_i)^2}{\sum_{i=1}^n (Y_i - \bar{Y}_i)^2}$$

Accuracy of the coefficient estimates

- The **standard error** of an estimator reflects how it varies under repeated sampling. We have

$$SE(\hat{\beta}_1) = \frac{\sigma}{\sqrt{\sum_{i=1}^n (X_i - \bar{X})^2}}; \quad SE(\hat{\beta}_0) = \sigma \sqrt{\frac{1}{n} + \frac{\bar{X}^2}{\sqrt{\sum_{i=1}^n (X_i - \bar{X})^2}}}$$

how close together x values are

$$\sigma^2 = \frac{1}{n-2} \overbrace{\sum_{i=1}^n (Y_i - \hat{Y}_i)^2}^{\text{how large the errors are}}$$

- These standard errors can be used to compute **confidence intervals**. A 95% confidence interval is defined as a range of values such that with 95% probability, the range will contain the true unknown value of the parameter. It has the form

$$\hat{\beta}_1 \pm 2 \cdot SE(\hat{\beta}_1)$$

Mutiple linear regression model

- The previous formulas can transform into matrix (linear algebra) notation

$$\mathbf{Y} = \mathbf{X}\boldsymbol{\beta} + \boldsymbol{\varepsilon}$$

$$Y = \begin{bmatrix} Y_1 \\ Y_2 \\ \vdots \\ Y_n \end{bmatrix}; \quad X = \begin{bmatrix} 1 & X_{11} & \cdots & \cdots & \cdots & X_{1p} \\ 1 & X_{22} & \cdots & \cdots & \cdots & X_{2p} \\ \vdots & \vdots & \ddots & \ddots & \ddots & \vdots \\ \vdots & \vdots & \ddots & \ddots & \ddots & \vdots \\ \vdots & \vdots & \ddots & \ddots & \ddots & \vdots \\ 1 & X_{n1} & \cdots & \cdots & \cdots & X_{np} \end{bmatrix};$$

response vector

$$\boldsymbol{\beta} = \begin{bmatrix} \beta_0 \\ \beta_1 \\ \beta_2 \\ \vdots \\ \vdots \\ \beta_p \end{bmatrix}; \quad \boldsymbol{\varepsilon} = \begin{bmatrix} \varepsilon_1 \\ \varepsilon_2 \\ \vdots \\ \vdots \\ \vdots \\ \varepsilon_n \end{bmatrix}$$

design matrix

slope vector

error vector

Mutiple linear regression model

- What mutiple linear regresion means graphycally?

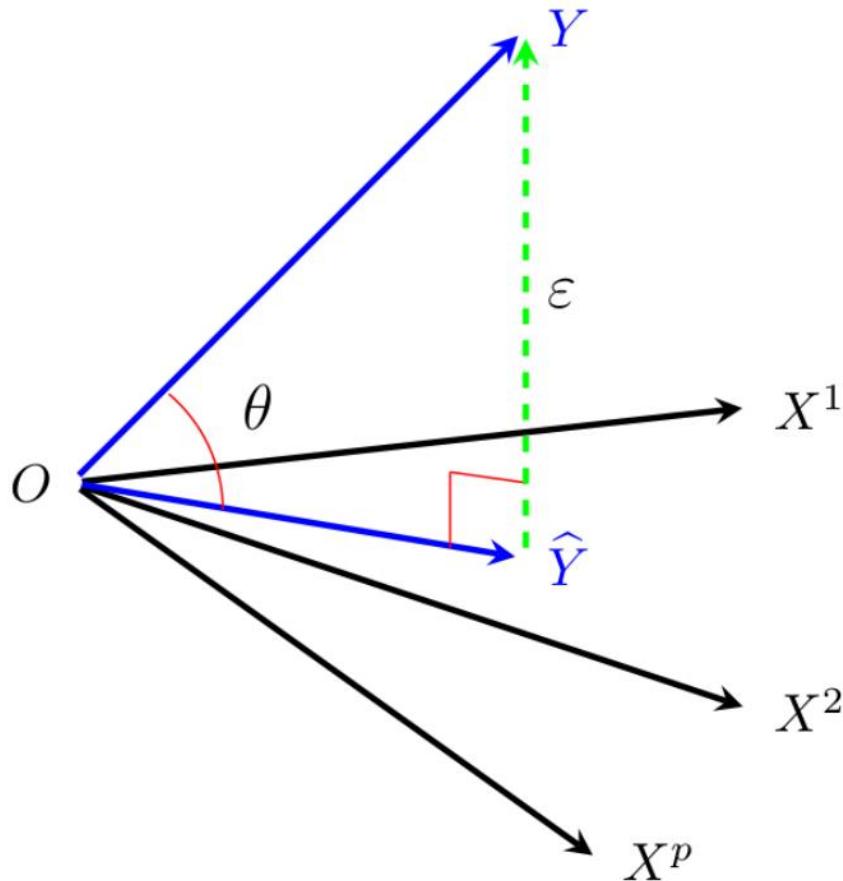


FIGURE 1 – Géométriquement, la régression est la projection \hat{Y} de Y sur l'espace vectoriel $\text{Vect}\{\mathbf{1}, X^1, \dots, X^p\}$; de plus $R^2 = \cos^2(\theta)$.

Least squares method

- In order to estimate β , we take a least squares approach that is analogous to what we did in the simple linear regression case. That is, we want to minimize

$$\min_{\beta_{i=1,\dots,n} \in \mathbb{R}^2} \left[\sum_{i=1}^n \left(Y_i - \beta_0 - \beta_1 X_{i,1} - \dots - \beta_p X_{i,p} \right)^2 \right]$$

over all possible values of the intercept and slopes. It is a fact that this is minimized by setting

$$\beta = (X^T X)^{-1} X^T Y$$

Estimating Parameters

- The **fitted values** are

$$Y = X \beta = X (X^T X)^{-1} X^T Y = HY$$

- The **residuals** are

$$\varepsilon = Y - Y = \left(I - X (X^T X)^{-1} X^T \right) Y = (I - H) Y$$

- The **error standard deviation** is estimated as

$$\sigma^2 = \frac{1}{n-1-p} \overbrace{\sum_{i=1}^n \varepsilon_i^2}^{\text{how large the errors are}}$$

Application to mean age of air

- To disentangle the impact of the natural variability (aerosols, QBO, ENSO) on BD-circulation, the monthly zonal mean age of air from CLaMS and TRACZILLA simulations, as function of latitude and altitude is analysed using a multiple linear regression model. This model yields:

$$\text{AoA}(t, \phi, z) = a(\phi, z) \cdot t + C(t, \phi, z) \quad \text{Trend + Seasonal cycle}$$
$$+ b_1(\phi, z) \cdot \text{qbo}(t - \tau_{\text{qbo}}) \quad \text{Quasi-biennial Oscillation}$$
$$+ b_2(\phi, z) \cdot \text{enso}(t - \tau_{\text{enso}}) \quad \text{El Nino Southern Oscillation}$$
$$+ b_3(\phi, z) \cdot \text{aod}(t - \tau_{\text{aod}}) \quad \text{Volcanic aerosols}$$
$$+ \varepsilon(t, \phi, z) \quad \text{Residual}$$

Application to mean age of air

- In this case $p=5$ and $n=12 \text{ months} \times N_{\text{year}}$.

$$Y = \beta X + \varepsilon$$

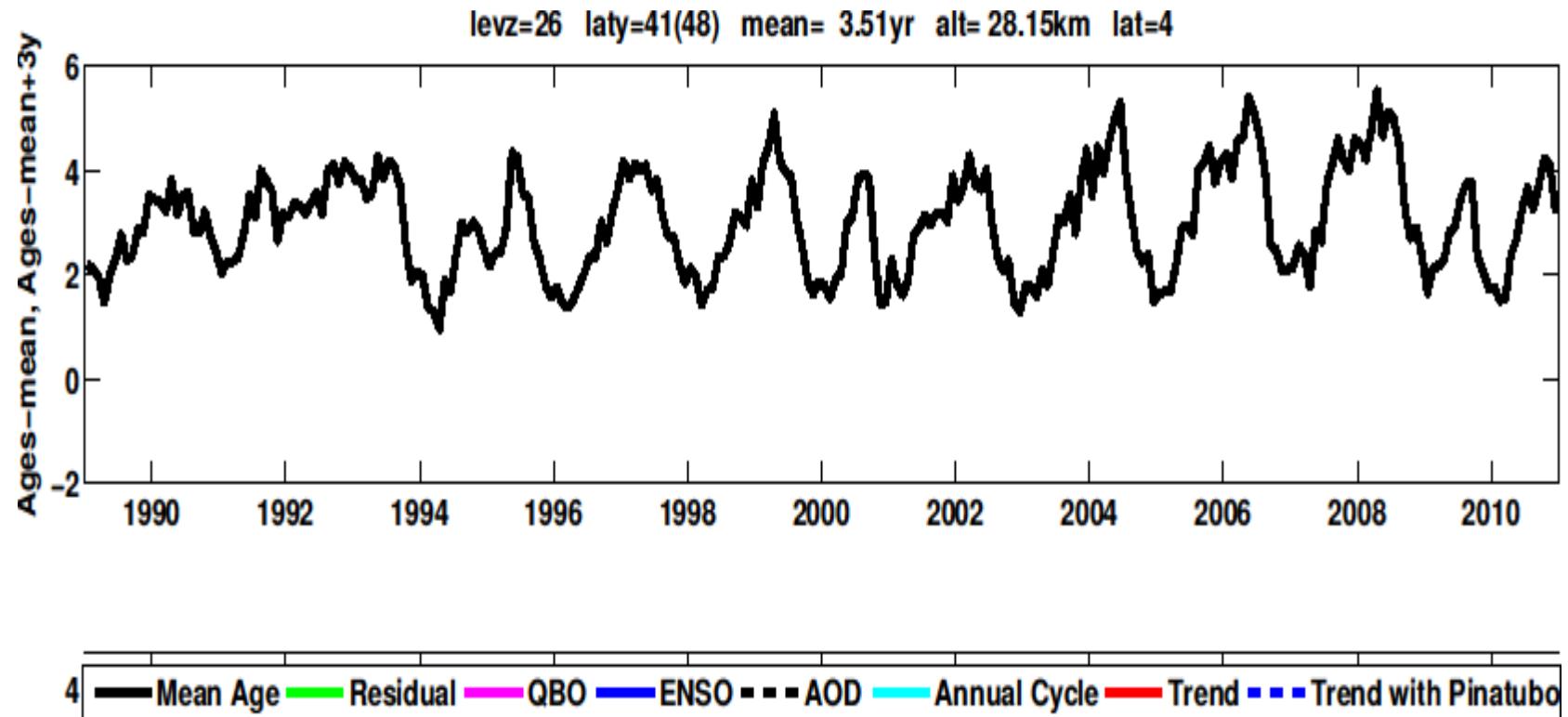
$$\min_{\beta_{i=1,\dots,n} \in \mathbb{R}^2} \left[\sum_{i=1}^n (age_i - d_j - a.t_i - b_1.qbo_i - b_2.enso_i - b_3.ado_i)^2 \right] = \min_{\beta_{i=1,\dots,n} \in \mathbb{R}^2} \sum_{i=1}^n \hat{\varepsilon}_i^2$$

$$j = 1 + [i-1, 12]$$

- Instead of **15** coefficients to estimate we can reduce the system to **4** and use a trick to calculate the **11** coeffs.

Application to mean age of air

- Time serie of the mean age of air at 28 km & 4N



Application to mean age of air

- Time serie of the mean age of air and different components of the best fit derived from the regression model.

