



ANTASed – An Updated Sediment Model for Antarctica

A. Baranov^{1*}, A. Morelli² and A. Chuvaev³

¹Schmidt Institute of Physics of the Earth, Russian Academy of Sciences, Moscow, Russia, ²Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Bologna, Bologna, Italy, ³Institute of Earthquake Prediction Theory and Mathematical Geophysics, Russian Academy of Sciences, Moscow, Russia

We compile existing seismic, gravity, radar and magnetic data, together with the subglacial bedrock relief from the BEDMACHINE project, to build the most detailed sediment model for Antarctica. We interpolate these data according to a tectonic map of Antarctica using a statistical kriging method. Our results reveal significant sediment accumulation in Antarctica with several types of sedimentary basins: parts of the Beacon Supergroup and more recent rifting basins. The basement relief closely resembles major geological and tectonic structures. The thickness of sediments has significant variations around the continent, and depends on the degree of crustal extension. West Antarctica has wide sedimentary basins: the Ross basin (thickness 2–6 km), the Filchner-Ronne basin (2–12 km) with continuations into East Antarctica, the Bentley Subglacial Trench and the Byrd basin (2–4 km). The deepest Filchner-Ronne basin has a complex structure with multi-layered sediments. East Antarctica is characterized by vast sedimentary basins such as the Pensacola-Pole (1–2 km), Coats Land (1–3 km), Dronning Maud Land (1–2 km), Vostok (2–7 km), Aurora (1–3 km), Astrolabe (2–4 km), Adventure (2–4 km), and Wilkes (1–4 km) basins, along with narrow deep rifts filled by sediments: Jutulstraumen (1–2 km), Lambert (2–5 km), Scott, Denman, Vanderford and Totten (2–4 km) rifts. The average thickness of sediments for the whole continent is about 0.77 km. The new model, ANTASed, represents a significant improvement over CRUST 1.0 for Antarctica, and reveals new sedimentary basins. Differences between ANTASed and CRUST 1.0 reach +12/–3 km.

Keywords: sediments, Antarctica, Gondwana, Denman, Vanderford, Astrolabe, Adventure, Scott

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*Correspondence:

A. Baranov
baranov@ifz.ru

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INTRODUCTION

Sediments represent the most heterogeneous layer of the Earth's interior and their characteristics can mask the mantle signature in geophysical measurements. In addition, information on age and density of sediments is also important for geodynamical and gravity modeling. Most of the large sedimentary basins formed due to crustal extension (continental rifting), a process initiated by mantle convection (e.g. Hartley and Allen, 1994). Geophysical data provide important clues on basin structure and origin. Seismic, gravity, radar, magnetic, and subglacial relief data can be used to construct sedimentary thickness maps. Sediment properties, such as their density, can also be estimated using geophysical data, from which their load can be inferred as a parameter for determining isostatic balance. Considerable efforts to investigate the sediments have been undertaken in the past, but some parts of the world are still poorly studied, such as Antarctica. Here, the ice sheet is partly underlain by sedimentary basins, and knowledge of the subglacial sedimentary layer holds particular interest

because of its role in glacier dynamics (e.g., Aster and Winberry, 2017) as sediments reduce basal friction and may thus favor basal sliding (Anandakrishnan et al., 1998). Also, low-lying subglacial sedimentary basins may be below sea level, and thus have important implications on ice sheet stability over global change (Studinger et al., 2001; Joughin and Alley, 2011). Global sediment models such as CRUST1.0 (Laske et al., 2013) and CRUST2.0 (Bassin et al., 2000) are poorly constrained in Antarctica. CRUST2.0 uses crustal types on a 2×2 Cartesian grid to specify sediment properties, including sediment thickness, while model CRUST1.0 uses the sediment model of Laske and Masters (1997) based on maps available at the time, including the 1985 AAPG Tectonic Map of the World assembled for Exxon (Laske et al., 2013). These global models obviously could not take into account the later maps of subglacial relief (Fretwell et al., 2013; Morlighem et al., 2020), magnetic and gravity data (e.g., Aitken, et al., 2014; Frederick et al., 2016; Scheinert et al., 2016; Golynsky and Golynsky 2019). A preliminary regional map of the total sediment thickness has been built by Baranov et al. (2018a). They found large basins of sediments, with maximum thickness under the Filchner-Ronne Ice Shelf (14 km). However, not all data had been taken into account at the time. In addition, that model has low detail in the interior regions. It is therefore important to update the sediment model for Antarctica using newly available geophysical data. Seismic tomographic images indicate cold, high seismic velocity, subcontinental mantle under East Antarctica (Danesi and Morelli, 2001; Becker, 2012; An et al., 2015) and a descending mantle flow (Chuvaev et al., 2020). On the contrary, under West Antarctica, according to tomographic models, we see presence of hot low-velocity upper mantle, and upwelling mantle flow (e.g., Morelli and Danesi, 2004; Lucas et al., 2020). Processes of extension and magmatism continue there at present time (Faccenna et al., 2008). West Antarctica is now a large zone of crustal extension, and can therefore be inferred as a vast area of sediment accumulation (Jordan et al., 2020).

Perhaps the most important tectonic feature of West Antarctica is the West Antarctic Rift System — a 500–600 km wide region of stretched continental crust that borders the Transantarctic Mountains (Davey et al., 2006; LeMasurier, 2008; Jordan et al., 2010; Granot et al., 2013a; Tinto et al., 2019; Jordan et al., 2020; Lucas et al., 2020). This rift system differs from other active rift systems in one important respect: it has very low (subglacial) relief (1,000–2,500 m below sea level). Rifting processes began during and before the Gondwana breakup, and in some places they continue till now (Granot et al., 2013b). Broad-mode Cretaceous rifting linked to the Gondwana breakup has been followed by narrow-mode Cenozoic rifting (Brink et al., 1993; Davey et al., 2006; Huerta and Harry, 2007). These processes have been accompanied by sediment accumulation (Hole and LeMasurier, 1994). The new BEDMACHINE detailed subglacial topography map (Morlighem et al., 2020) reveals some new features of the subglacial structure of Antarctica and creates the prerequisites for improving a sediment thickness map. In this study, we compile seismic, bedrock and other geophysical data to construct a new sediment model of Antarctica. The framework of our map is given by the BEDMACHINE subglacial relief (Morlighem et al.,

2020), and the gravity map from Scheinert et al. (2016). The data and the compilation technique are described in *Data Compilation and Methods* and **Table 1**. Results are described in *Results*. Discussion and concluding remarks are given in *Discussion and Conclusion*.

DATA COMPILATION AND METHODS

Seismic data for Antarctica are geographically rather sparse, especially in inland regions. However, we also resort to new gravity, radar, bedrock and magnetic data to infer information on shallow crustal structure in regions without seismic coverage. These data mainly provide indirect values for sediment thickness. **However magnetic data allow a fairly accurate determination of sediment thickness i.e. depth to magnetic basement** e.g. (Aitken et al., 2014). The data sources are presented in **Table 1**. They result in an uneven data distribution on the whole continent (**Figure 1**). Having collect suitable data, we analyse them, infer additional contours of sediment thickness according to the bedrock relief, and interpolate the data array onto a geographical grid, created with an azimuthal equidistant projection centered at the South Pole, with grid step 0.1° , using standard kriging (linear variogram; SURFER, Golden Software package). We limit the grid by the border of the Antarctic continent. This technique has been described in detail and applied to construct Moho models for Antarctica (Baranov and Morelli, 2013; Baranov et al., 2018a) and the crustal models for Central and Southern Asia (Baranov, 2010; Baranov et al., 2018b). We obtain seismic velocity information for sediments from Deep Seismic Sounding (DSS), wide-angle reflection, and refraction profiles (**Table 1**). We estimate uncertainties of such a model, using standard deviations (SURFER, Golden Software package). Although kriging standard deviations do not necessarily give an accurate picture of the interpolation prediction error (e.g., Cressie, 1990), we provide such a map in **Figure 2**. It can be interpreted as a proxy for the relative geographical reliability of our model. Nevertheless, we believe that the sediment map for Antarctica has significantly been improved by integrating many more information sources than in earlier studies.

RESULTS

It has been suggested that some of the present sedimentary basins in Antarctica began to form during and before Gondwana breakup from Permian age (Powell et al., 1988; Boger and Wilson, 2003) during the formation of the Beacon Supergroup (Collinson et al., 1994). The Beacon Supergroup experienced the intrusions of magmatic sills during the Gondwana breakup, after which sedimentation in the depressions continued and ceased in Eocene-Miocene due to ice covering the continent (Whitehead et al., 2006). Our results, based on geophysical data summarized in *Data Compilation and Methods*, are presented next. The resulting sediment thickness for Antarctica is shown in **Figure 3A**.

TABLE 1 | Data sources used in this study.

Region	Total sediment thickness, km	Upper sediments, km	Middle sediments, km	Lower sediments, km	Upper sediments, Vp, km/s	Middle sediments, Vp, km/s	Lower sediments, Vp, km/s	Source	Additionally
Coats Land	1–3	1–2	0–1	0	4.3–4.4?	?	—	Bamber et al. (2006), Shepherd et al. (2006), Scheinert et al. (2016) (Radio-echo sounding, gravity and magnetic data)	Beacon Supergroup? new rifted sediments, latitudinal strike
Dronning Maud Land	1–2	1–2	-	—	4.7	-	—	Hungeling and Tyssen (1991), Scheinert et al. (2016) (seismic and gravity data)	Local area near P2 profile and for Jutulstraumen Rift
Lambert Rift	2–5	2	0–3	—	3.4–3.5	3.5–3.7	—	Kolmakov et al., (1975) (DSS profiles); Mishra et al. (1999), Scheinert et al. (2016) (gravity and magnetic data)	From Devon to Cenozoic sediments
Scott and Denman rifts	2–4	2	0–2	—	?	?	—	Golynsky and Golynsky (2012), Scheinert et al. (2016) (gravity data)	From Devon to Cenozoic sediments
Vostok Basin	2–7	2	2–5	—	3.4–3.5	3.6–3.7	—	Isanina et al. (2009) (seismic profile with converted waves); Golynsky and Golynsky (2012) (radar, gravity and magnetic data)	From Devon to Cenozoic sediments
Aurora Basin	1–3	1–2	0–1	—	?	?	?	Aitken et al. (2014) (radar, gravity and magnetic data)	?
Adventure Trench	2–4	2	0–2	—	?	?	?	Ferraccioli et al. (2001), Studinger et al. (2004) (gravity and magnetic data)	Rifted sediments?
Astrolabe Subglacial Basin	2–4	2	0–2	—	?	?	?	Aitken et al. (2014) (radar, gravity and magnetic data)	Rifted sediments?
Vanderford and Totten rifts	2–4	2	0–2	—	?	—	—	Aitken, et al. (2014), Golynsky and Golynsky (2019) (radar, gravity and magnetic data)	latitudinal strike
Wilkes Subglacial Basin	1–4	1–2	0–2	—	3.4	3.6–3.7?	—	Ferraccioli et al. (2001), Frederick et al. (2016) (radar, gravity and magnetic data)	Beacon Supergroup, Cenozoic marine sediments?
Ross Ice Shelf	2–6	2	0–4	—	1.9–2.4	4.2	—	Bentley (1973) (BB', DSS profile); Rooney et al. (1987) (two refraction profiles); Brink et al. (1993) (reflection, refraction and gravity profile); Beaudoin et al. (1992) (refraction profile); Munson and Bentley (1992) (reflection, refraction and gravity profile)	Part of West Antarctic Rift System, Beacon Supergroup, Mesozoic-Cenozoic sediments
Byrd and Bentley Subglacial basins	2–4	2	0–2	—	4.1	4.2–4.3	—	Bentley (1973) (BB', CC' DSS profiles)	Part of West Antarctic Rift System, Beacon Supergroup, Mesozoic-Cenozoic sediments
Filchner-Ronne Basin	2–12	2	0–5	0–7	2.6–3.0	3.3–4.0	4.6–4.9	Huebscher et al. (1996) (2 refraction profiles); Leitchenkov and Kudryavtzev (1997) (DSS profile); Golynsky and Aleshkova (1997) (magnetic data)	Sediments of terrigenous, marine and glacial marine origin
Pensacola-Pole Subglacial Basin	1–2	1–2	—	—	4.3–4.4?	—	—	Bentley (1973) (AA' DSS profile); Paxman et al. (2019) (gravity, radar and magnetic data)	Beacon Supergroup?

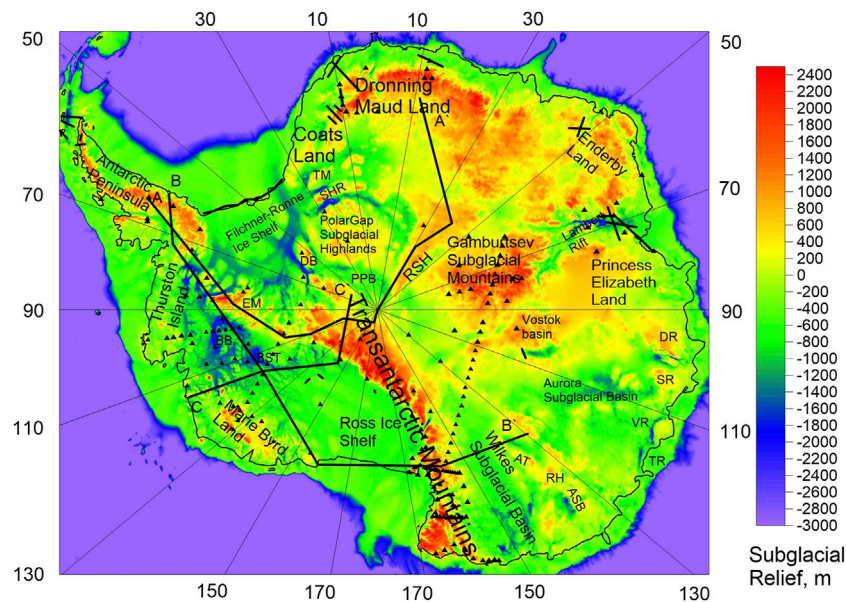


FIGURE 1 | Subglacial relief in Antarctica according to BEDMACHINE (Morlighem et al., 2020). Notation used: black triangles represent seismic stations; black lines represent seismic profiles. Used abbreviation: PPB, Pensacola-Pole Basin; EM, Ellsworth Mountains; BB, Byrd Basin; BST, Bentley Subglacial Trench; SHR, Shackleton Range; TM, Theron Mountains; SR, Scott rift; DR, Denman rift; VR, Vanderford rift; TR, Totten rift; AT, Adventure Trench; ASB, Astrolabe Subglacial Basin; RH, Resolution Highlands; DB, Dufek Block; RSH, Recovery Subglacial Highlands.

East Antarctica

East Antarctica has vast areas with low bedrock elevation and sedimentary basins. These basins can be divided into three blocks: the continuation of the Filchner-Ronne basin into East Antarctica, the sedimentary basins of the Indo-Antarctica and Australo-Antarctica blocks of East Antarctica. At the deep subglacial valleys between PolarGap Subglacial Highlands, Shackleton Range and Theron Mountains there are indications of a 2–3 km thick sedimentary layer, in continuation to the South Pole (Bamber et al., 2006; Shepherd et al., 2006; Scheinert et al., 2016). These sediments are believed to be part of the Paleozoic Beacon Supergroup (Collinson et al., 1994). A Permian sedimentary rock layer of about 1–2 km is found for the Western Dronning Maud Land, whereas on the P1, Novo, SAE33, SAE34 and Kottas seismic profiles sediments were not found (Kogan, 1971; Hungeling and Tyssen 1991; Kudryavtzev et al., 1991; Bayer et al., 2009). Part of the AA' profile from South Pole to the central part of Dronning Maud Land reveals no sediments (Bentley, 1973). No sediments are found at Enderby Land (Kanao et al., 2011). According to subglacial relief (Morlighem et al., 2020) and gravity data (Scheinert et al., 2016), we suggest 1–2 km of sediments for the narrow Jutulstraumen rift.

Lambert Rift is a deep long subglacial valley with low relief and sediments about 2–5 km thick (Kolmakov et al., 1975; Mishra et al., 1999). The process of rifting started at Permian age and continued in the Mesozoic (Boger and Wilson, 2003; Lisker et al., 2003). Further, in addition to rifting, the area apparently also submerged. Terrestrial forested sediments have been found near

the coast in Prydz Bay, with the first evidence for glaciation of Late Eocene age (Whitehead et al., 2006).

Scott and Denman are other long narrow rifts (Morlighem et al., 2020) with a possible continuation in Vostok subglacial Basin (Golynsky and Golynsky, 2012). It is suggested that these rifts are the continuation of the rift between India and Australia that developed during Gondwana break-up (Harrowfield et al., 2005). Lambert, Scott and Denman rifts have continuation in India (Holdgate et al., 2005). According to subglacial relief (Morlighem et al., 2020) and gravity data (Scheinert et al., 2016), we hence suggest 2–4 km of sediments for Scott and Denman trenches.

The seismic profile for the Vostok Basin near the coastline of the Vostok Lake provides sediment thickness of about 2–7 km (Isanina et al., 2009). A more recent study also reveals hard and soft sediments here (Filina et al., 2008).

The Aurora Subglacial Basin is characterized by low subglacial topography and deep trenches. It is a branching system of narrow basins in central East Antarctica trending northwest towards the Aurora coast. The inland system of deep trenches can be tracked over a distance of 1,000 km towards the Transantarctic Mountains under ice in the Aurora basin (Tabacco et al., 2006; Golynsky and Golynsky, 2019; Morlighem et al., 2020). Radar, gravity and magnetic data suggest 1–3 km of sediments for these trenches (Aitken et al., 2014; Scheinert et al., 2016).

Data from the ICECAP/IceBridge project outline the Vanderford and Totten deep subglacial trenches. These trenches are about 50 km wide and 400 km long, with very low subglacial relief (Aitken et al., 2014). It is also suggested that the Vanderford and Totten trenches are another rift that,

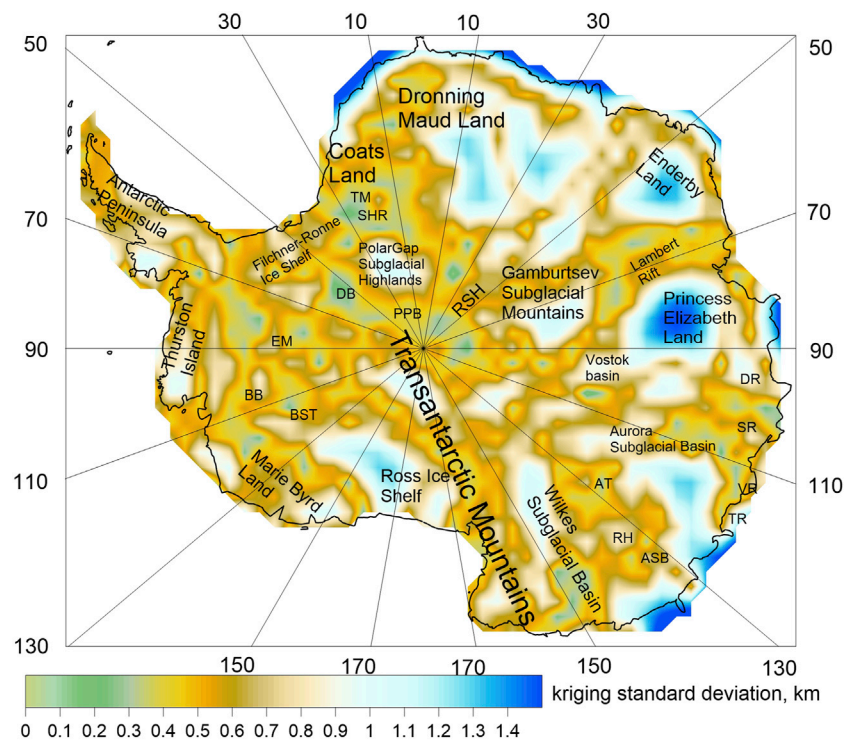


FIGURE 2 | Map of kriging standard deviation. It is used the same abbreviations as in **Figure 1**.

unlike the other East Antarctic rifts, has a latitudinal strike similar to Coats Land rifts. This rift has probably formed during Late Paleozoic–Early Mesozoic, such as Lambert Rift (Boger and Wilson, 2003; Lisker et al., 2003). Gravity and magnetic data suggest that the thickness of sedimentary rocks within the rift can exceed 4 km (Golynsky and Golynsky, 2019).

The Adventure Trench at 132°E is located in the intra-continental realm inside the Australo-Antarctic block. Gravity and magnetic data suggest 2–4 km of sediments here (Ferraccioli et al., 2001; Studinger et al., 2004).

The Astrolabe Subglacial Basin is another isolated deep depression between Adventure Trench and Resolution Highlands. According to magnetic data and low subglacial relief we suggest sediment thickness of 2–4 km here too (Aitken et al., 2014).

The Wilkes Basin is a broad area between Transantarctic Mountains and Aurora Basin with negative subglacial relief (Morlighem et al., 2020) and deep trenches due to possible processes of continental rifting (Ferraccioli et al., 2001). This basin lies sub-parallel to the Transantarctic Mountains and the West Antarctic Rift System. Radar, gravity and magnetic data predict 1–4 km thickness of sediments for the Wilkes basin, interpreted mainly as Beacon Supergroup (Ferraccioli et al., 2001; Ferraccioli et al., 2009; Frederick et al., 2016). Thickness of sediments for the Central basin is about 1–2 km, whereas for the Western Basin it is about 2–4 km.

The Pensacola–Pole Basin (PPB) is located in proximity to the Transantarctic Mountains near the South Pole. New radar data

reveal an elongated subglacial basin with an average subglacial relief of –500 m (Morlighem et al., 2020), lying between Dufek Block, Polargap Subglacial Highlands, Recovery Subglacial Highlands, Pensacola Mountains and South Pole. Seismic refraction (Bentley, 1973), gravity, radar and magnetic studies (Wannamaker et al., 2004; Paxman et al., 2019) indicate that this sedimentary basin has thickness of about 1–2 km.

West Antarctica

In West Antarctica, the ice sheet is grounded below sea level in most parts, so that the basal sediment layer is even more relevant for its evolution than in East Antarctica. The West Antarctic Rift System is the most dominant and relevant tectonic feature connected to sedimentary basins. These basins have low subglacial relief, and their sediments are due to rifting processes. At the Ross Ice Shelf, the sediment thickness reaches 6 km according to seismic data (Bentley, 1973; Rooney et al., 1987; Beaudoin et al., 1992; Munson and Bentley, 1992; Brink et al., 1993; Shen et al., 2018). The Byrd Subglacial Basin and Bentley Subglacial Trench have very low subglacial topography (up to –2,500 m; Morlighem et al., 2020). The sedimentary layer here is about 2–4 km thick (Bentley, 1973). The area below the Filchner-Ronne Ice Shelf is another large sedimentary basin. The sedimentary properties change along the seismic profiles near the coastline from Antarctic Peninsula to Coats Land (Huebscher et al., 1996; Leitchenkov and Kudryavtzev, 1997; Jokat and Herter, 2016). At the center of the basin, the thickness of sediments reaches more than 12 km

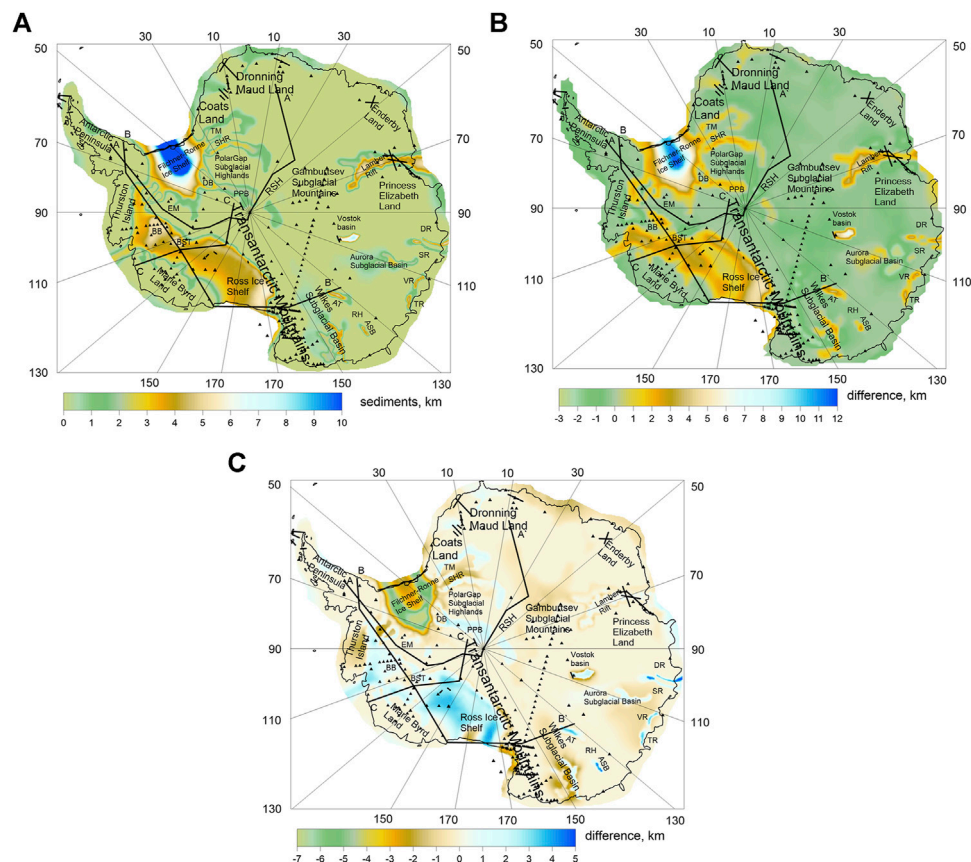


FIGURE 3 | (A) Map of a sediment thickness for Antarctica (ANTASed). It is used the same abbreviations as in **Figure 1**. **(B)** Difference between ANTASed model and the CRUST 1.0 model. **(C)** Difference between ANTASed and preliminary sediment model from Baranov et al. (2018a).

(Huebscher et al., 1996; Leitchenkov and Kudryavtzev, 1997). According to magnetic and seismic Rayleigh wave studies the sedimentary basin continues to the south until Ellsworth Mountains and Dufek block with a continuation into Coats Land of East Antarctica and between Ellsworth Mountains and Antarctic Peninsula (Golynsky and Aleshkova, 1997; Shen et al., 2018).

For the CRUST 1.0 model, the average sediment thickness in the Antarctic continent is about 0.58 km with maxima up to 5 km, while our results (**Figure 3A**) indicate that the **average sediment thickness is about 0.77 km with maxima up to 12 km** under the Filchner-Ronne Ice Shelf. Our model reveals significant differences with respect to both CRUST1.0 (see **Figure 3B**) and the previous sediment model by Baranov et al., 2018a (**Figure 3C**). In CRUST1.0, mainly based on the 1985 AAPG Tectonic Map of the World assembled for Exxon (Laske et al., 2013), several sedimentary basins in East Antarctica are not represented at all—Adventure, Astrolabe and Vostok subglacial basins, rifts of East Antarctica except Lambert Rift, basins of Coats Land. At the same time, in our ANTASed model, sedimentary basins from CRUST1.0 are shown in more detail (Lambert Rift, Aurora, Wilkes, Pensacola-Pole, Filchner-Ronne, Byrd, Bentley and Ross basins). Difference between ANTASed and CRUST1.0 may amount up to 10 km, with average 0.19 km

(standard deviation 1.53 km), that we mainly attribute to the improved resolution of our model (see **Figure 3B**). There are significant differences in regions such as under the Filchner-Ronne Ice Shelf (4–12 km), Lambert, Denman, Scott, Vanderford and Totten Rifts (up to 4 km), under the Ross Ice Shelf (up to 4 km), Wilkes, Bentley and Byrd basins (2–3 km); Pensacola-Pole basin (1–2 km), basins of Coats Land (2–3 km).

For the Filchner Ronne Ice Shelf area, the Exxon map used to build CRUST1.0 (<https://www.worldcat.org/title/tectonic-map-of-the-world/oclc/32224099>) delineates two separate sedimentary basins, the Ronne and the Filchner basins, while according to more recent seismic data we have a single, continuous, basin (Huebscher et al., 1996; Leitchenkov and Kudryavtzev, 1997). At the time of Exxon map creation these seismic profiles did not exist and it was suggested that Berkner Island is a natural boundary between these sedimentary basins. This explains the large difference between the CRUST1.0 sediment model (Laske et al., 2013) and ANTASed, up to 10 km. In the preliminary sediment model by Baranov et al. (2018a) part of the sedimentary basins of East Antarctica is also missing (Astrolabe basin, rifts of East Antarctica except Lambert Rift, Pensacola-Pole and basins of Coats Land).

In our new model other sedimentary basins are mapped with more details (Lambert Rift, Aurora, Wilkes, Filchner-Ronne,

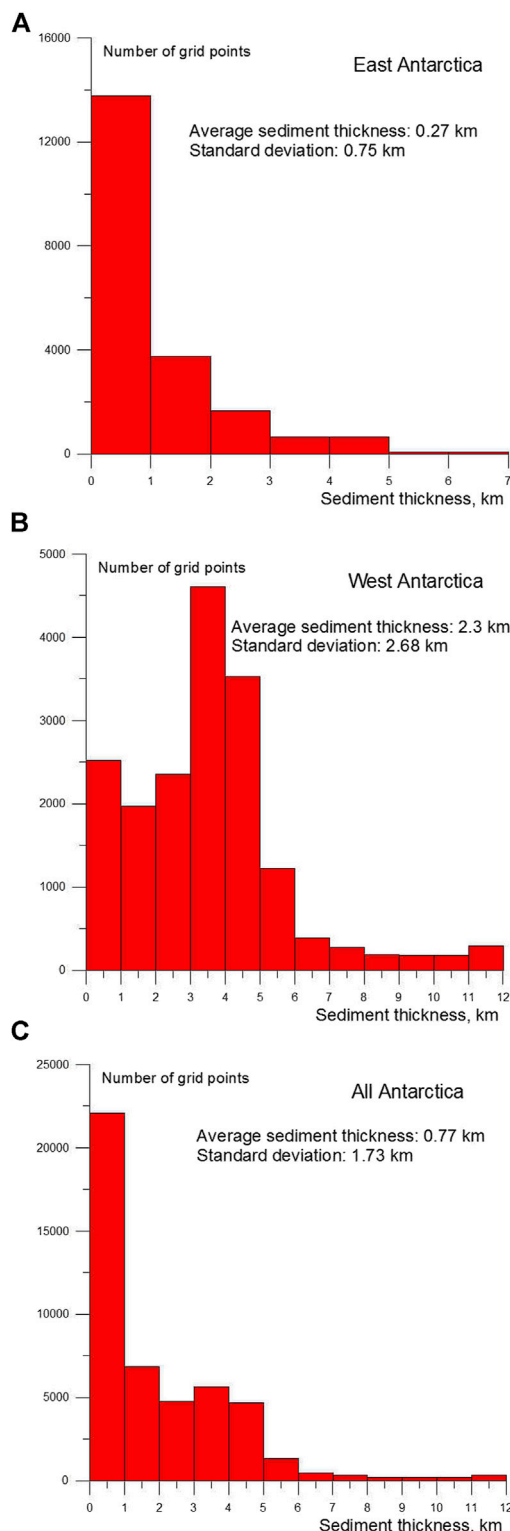


FIGURE 4 | Histograms of the total sediment thickness for (A) East Antarctica, (B) West Antarctica and the whole continent (C).

Byrd, Bentley and Ross basins). Although the entire Wilkes Basin had previously been considered a wide sedimentary basin, newer data show that sediments are mainly located in narrow and deep grabens (Western and Central Basins) of the Wilkes lowland. A similar situation is also true for the Aurora basin. Adventure basin is deeper than in the previous model by Baranov et al. (2018a). Deep and narrow rifts of East Antarctica have properly been recognized in the BEDMACHINE model (Morlinghem et al., 2020). Additional data reveals sediments for these rifts. The boundaries and sediment thickness for the Filcher-Ronne Ice basin are identified by new geophysical data (Jokat and Herter, 2016).

The difference between ANTASed and the preliminary sediment model by Baranov et al. (2018a) is $-7/+5$ km with average -0.06 km (standard deviation 1.1 km) mainly due to the improved resolution of our model (see Figure 3C). There are significant differences in regions such as under the Filchner-Ronne Ice Shelf (-2 to -7 km), Denman, Scott, Vanderford and Totten rifts (up to 4 km), under the Ross Ice Shelf (-2 to $+2$ km), Wilkes, Bentley and Byrd basins (-2 to $+2$ km), Pensacola-Pole basin (1–2 km), Vostok basin (-2 to $+2$ km), Adventure basin (2–3 km); basins of Coats Land (-2 to $+2$ km).

Figure 4 shows the frequency distribution of sediment thickness for East Antarctica, West Antarctica, and for the whole continent. The average sediment thickness of West Antarctica is about 2.3 km (standard deviation 2.68 km) due to large sedimentary basins. The average sediment thickness of East Antarctica is about 0.27 km (standard deviation 0.75 km).

DISCUSSION AND CONCLUSION

We have collected and analyzed different geophysical data: subglacial bedrock relief, gravity, radar, magnetic, and seismic data to provide a regional sediment model of the Antarctic continent. The new sediment thickness grid, ANTASed, is then built, analyzed and compared with previous models. The bulk of the sediments in Antarctica belongs to the Beacon Supergroup, in addition to them, there are also younger Mesozoic and Cenozoic sediments.

Our results, based on the combined processing of geophysical data (see Figure 1), reveal significant differences between the sediment structure of East and West Antarctica. The dimensions of sedimentary basins are determined by the degree of crustal extension - rather moderate for East Antarctica (up to 7 km for Vostok basin) and significantly higher for West Antarctica (up to 12 km for the region below the Ronne Ice Shelf).

The main sedimentary basins are concentrated in broad areas of extension in the West Antarctic Rift System with low subglacial relief (Ross, Filchner-Ronne, Bentley and Byrd basins and their extensions) and in the lowlands of the

Australo-Antarctica block of East Antarctica.

According to our results, West Antarctica contains wide sedimentary basins: the Ross basin (thickness 2–6 km); the Filchner-Ronne basin with its continuations to the Antarctic Peninsula; South Pole and Coats Land (2–12 km), Bentley and Byrd basins (2–5 km). Maximum sediment thickness is reached under the Filchner-Ronne Ice Shelf (up to 12 km) with a complex multi-layered structure.

For East Antarctica there are both narrow sediment-filled rifts: Jutulstraumen (thickness 1–2 km), Lambert (2–5 km), Scott, Denman, Vanderford and Totten (2–4 km); and wider areas such as the Vostok (2–7 km), Aurora (1–3 km), Astrolabe (2–4 km), Adventure (2–4 km) and Wilkes (1–4 km) basins, corresponding to the Australo-Antarctica block of East Antarctica.

The average sediment thickness for Antarctica is about 0.77 km, for West Antarctica 2.3 km, for East Antarctica 0.27 km. Compared to previous models CRUST1.0 (Laske et al., 2013) and ANTACrust (Baranov et al., 2018a), we recognize the new Pensacola-Pole, Coats Land, Denman, Scott, Adventure, Astrolabe, Vanderford and Totten sedimentary basins in East Antarctica. The resolution of the sediment map is also improved. Differences between ANTASed and previous models reach +12/–3 km for CRUST 1.0 (Laske et al., 2013) and +5/–7 km for ANTACrust (Baranov et al., 2018a) respectively. By integrating a wide range of specific geophysical and geological information, ANTASed provides significant improvements over global models for sediments in Antarctica.

The ANTASed model also complements the new three-layer crustal model for Antarctica (Baranov et al., 2021) and sediment

model for surrounding ocean (Straume et al., 2019). We provide our sediment map for Antarctica and hope that it will be useful to scientists (<https://www.itpz-ran.ru/en/activity/current-projects/antased-a-new-sediment-model-for-antarctica/>). Nevertheless, we recognize that new data are still needed to add information particularly in the central parts of Antarctica.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/Supplementary Material, further inquiries can be directed to the corresponding author.

AUTHOR CONTRIBUTIONS

AB designed and performed calculations of the maps, provided a main framework of the research. AM provided a general leadership and corrected the manuscript. AC provided a BEDMACHINE data and corrected the manuscript.

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