



CANAIMOC Workshop



CArbon **N**orth **A**tlantic Irrigation by the **M**eridional **O**verturning **C**irculation

Carbon cycle in the ocean Methods to estimate anthropogenic carbon

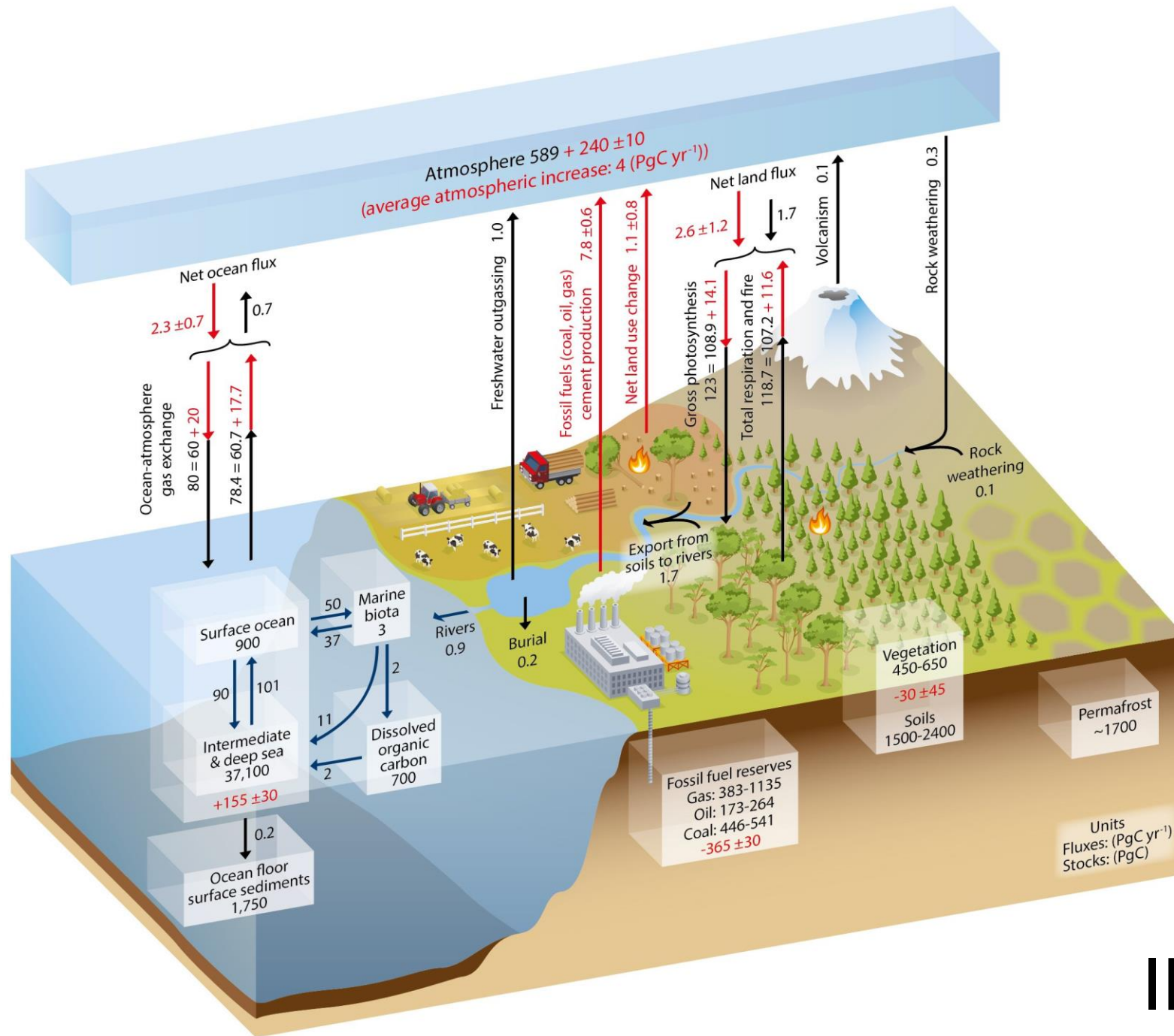
Marcos Fontela, 2021



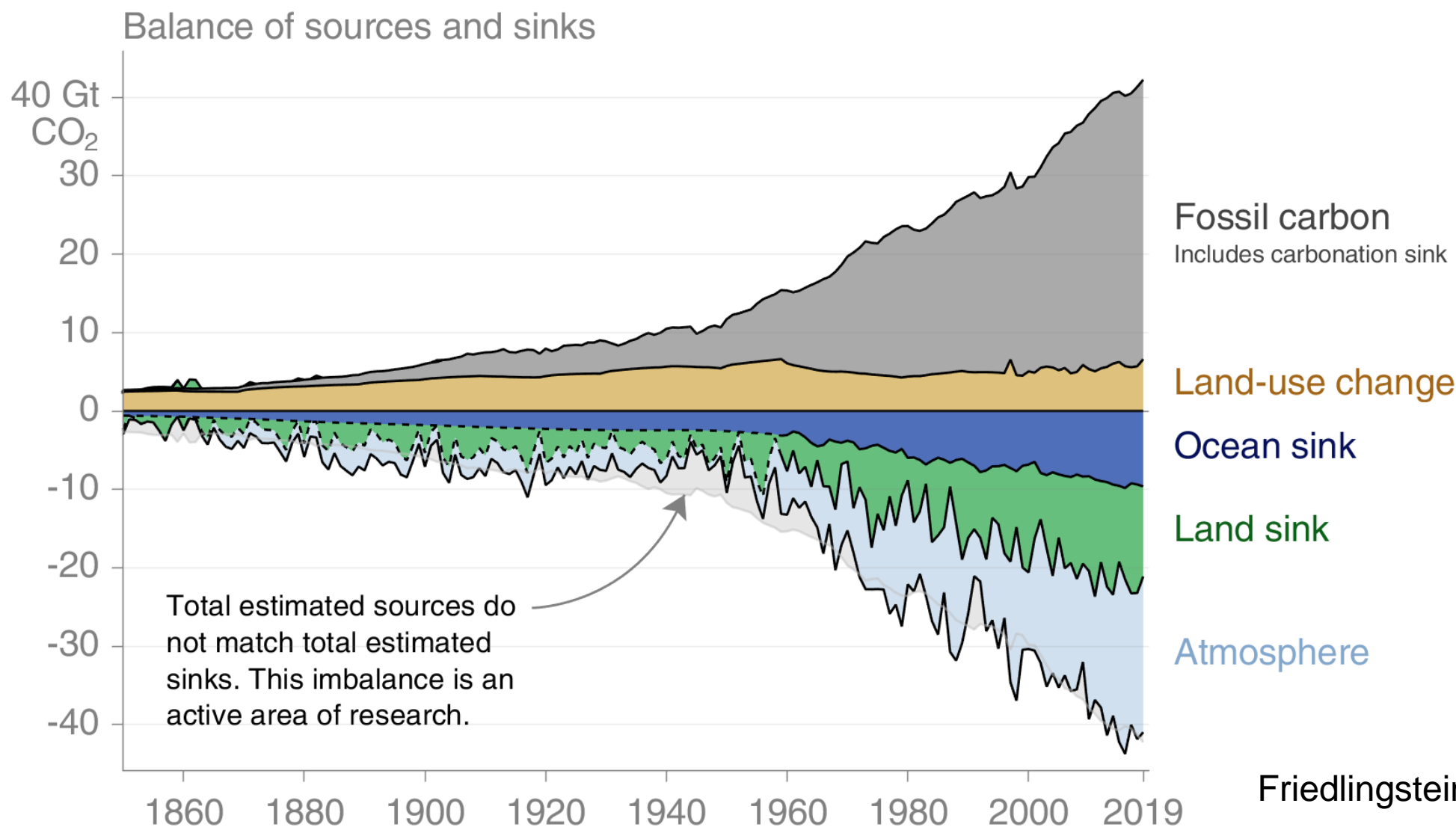
*Materials in:

<https://github.com/mfontela/CANAIMOC>

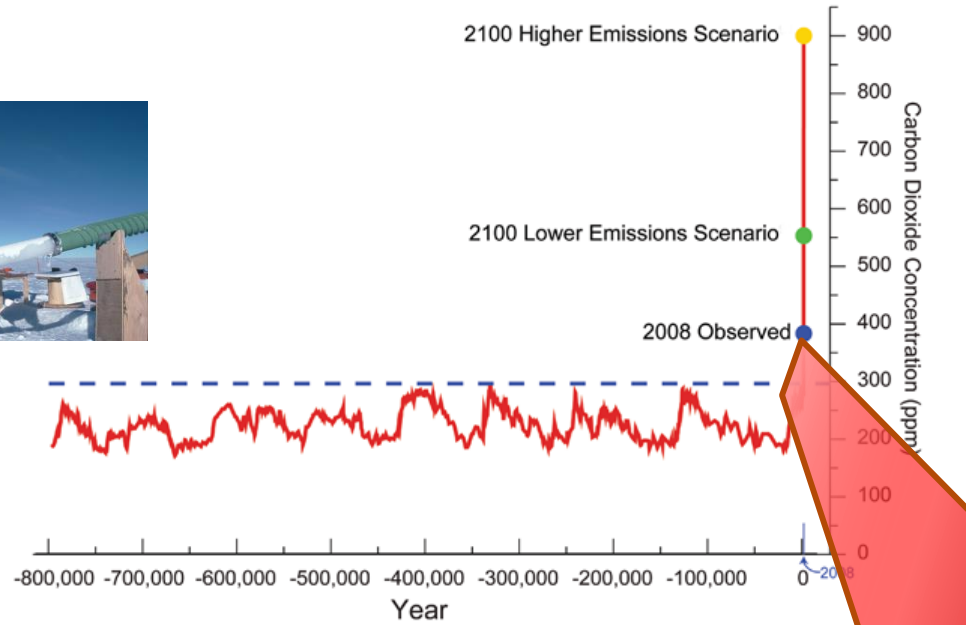
- Carbon cycle
 - Intro
 - Methods for Ocean carbon parameters
 - ❖ Study cases: seacarb
- Anthropogenic carbon
 - Intro
 - Methods: back-calculation vs transient tracers
- ϕC_T^o :
 - Study cases



IPCC AR5

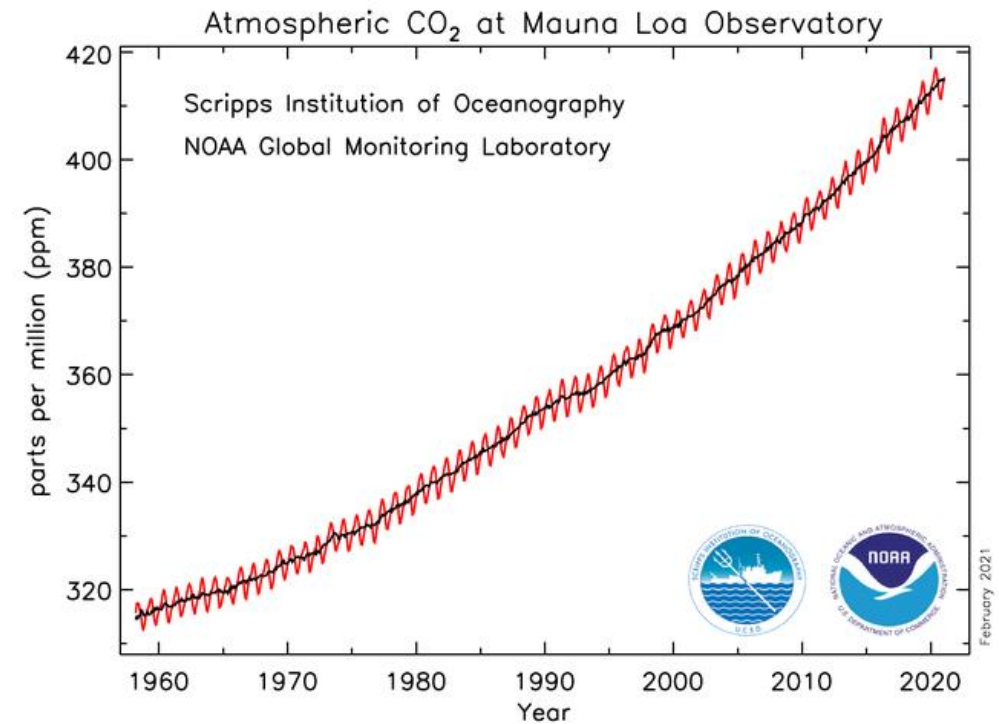


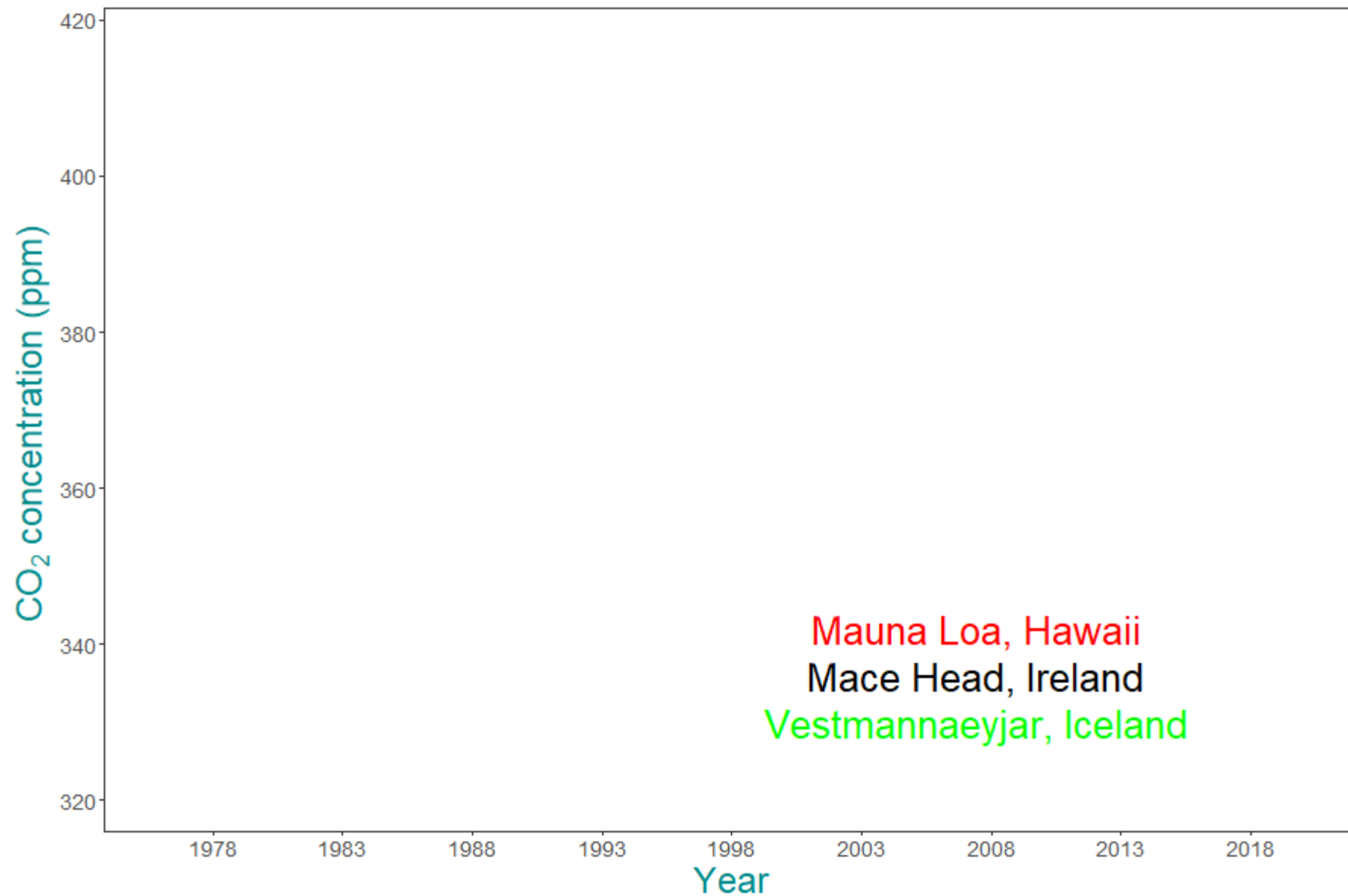
Friedlingstein et al. 2019 ESSD



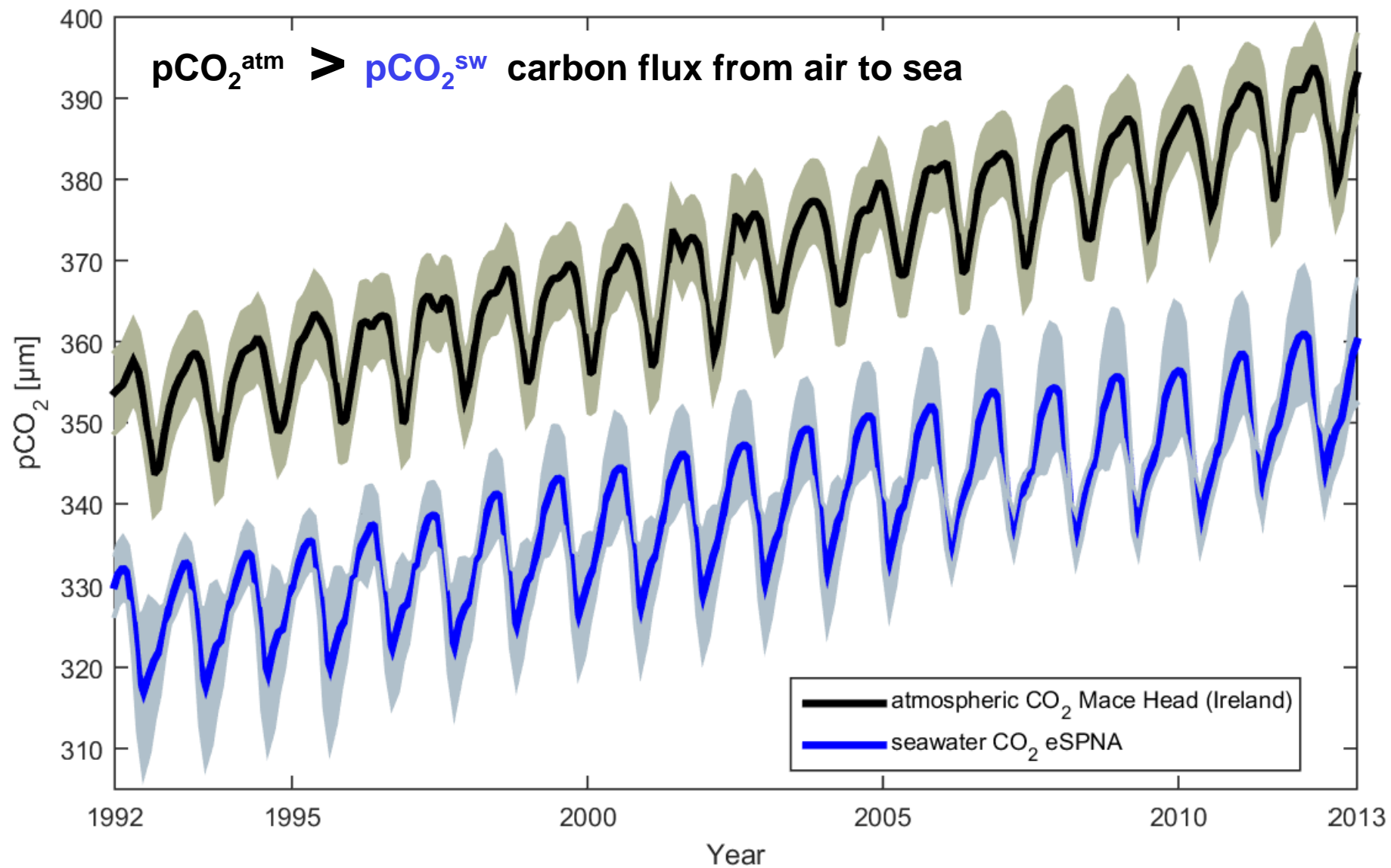
Luthi et al. 2008 Nature

Keeling curve



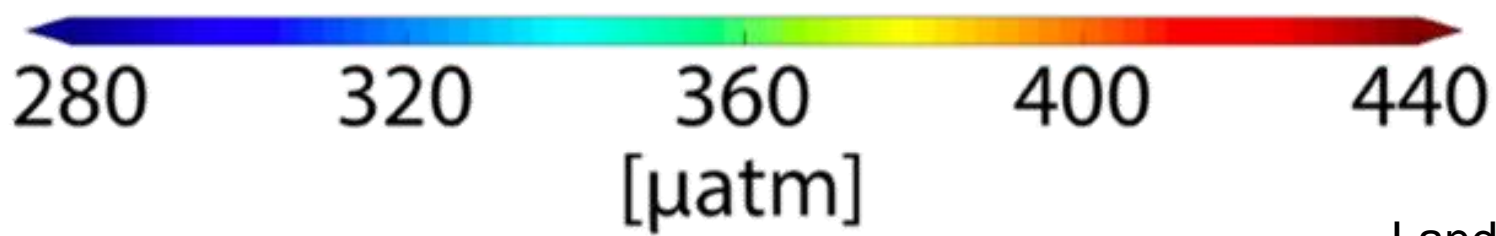
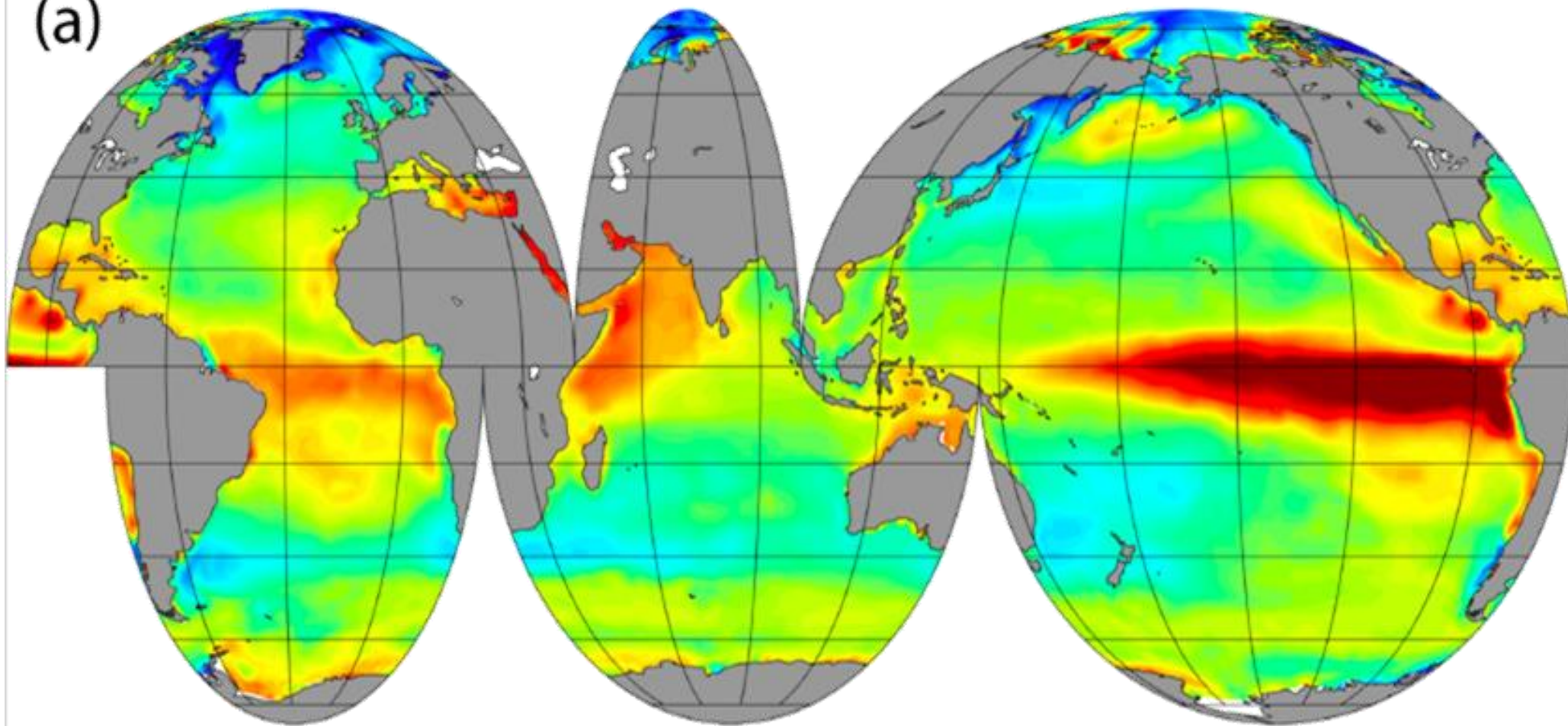


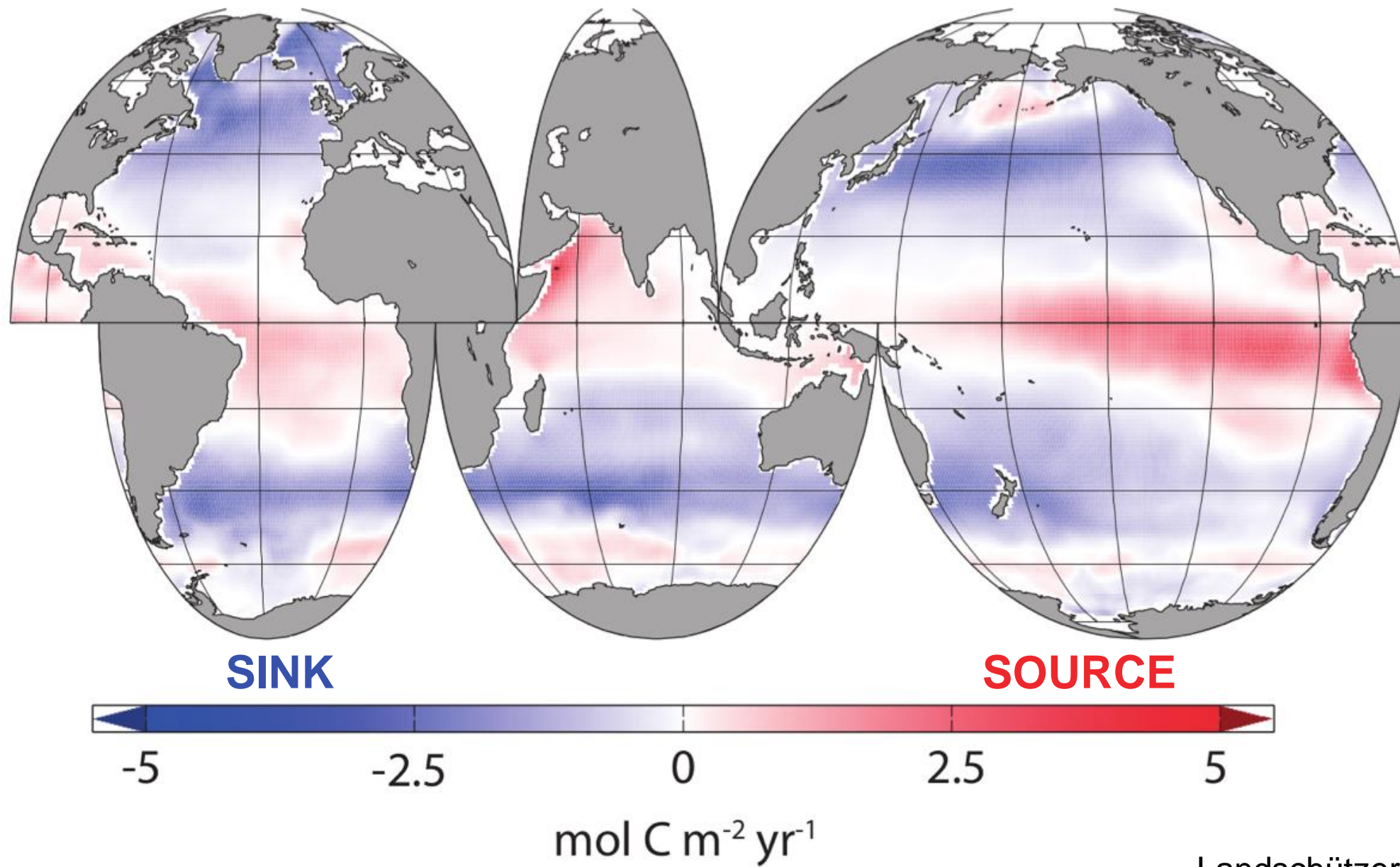
<https://www.esrl.noaa.gov/gmd/dv/data/>



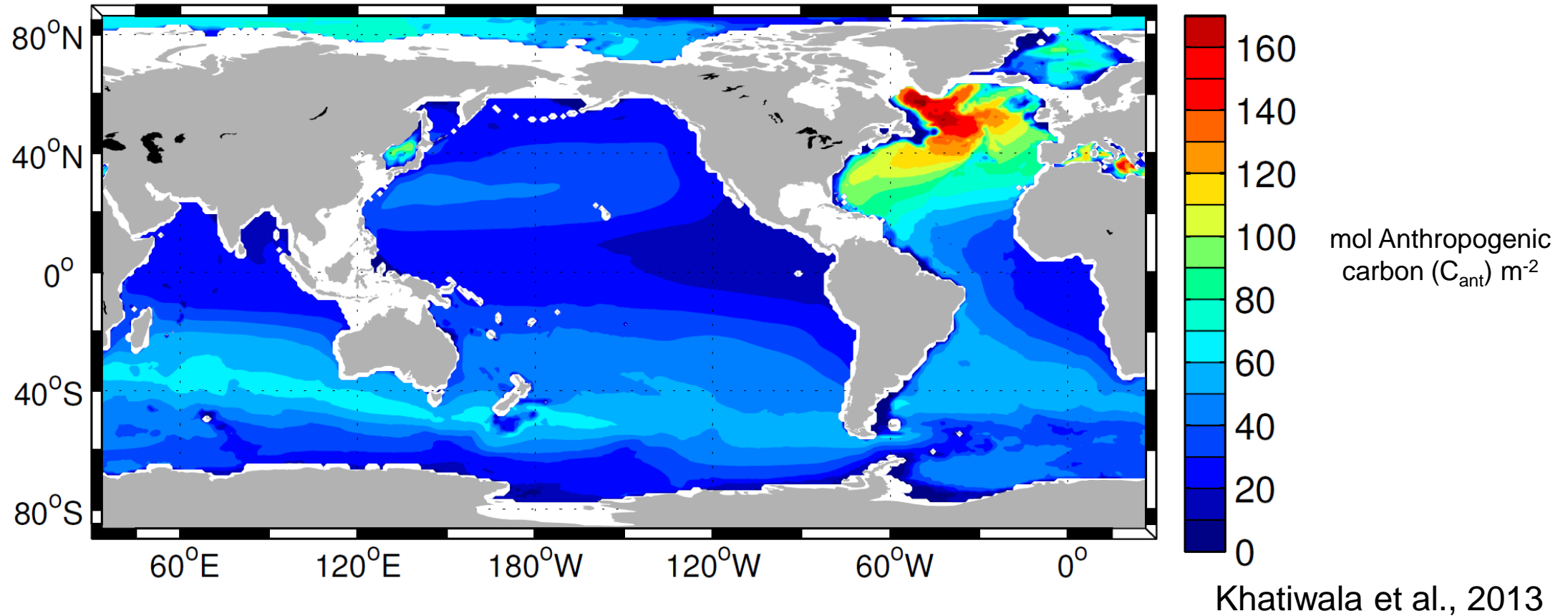
Seawater pCO₂ data from Rödenbeck et al. (2015) BG

(a)





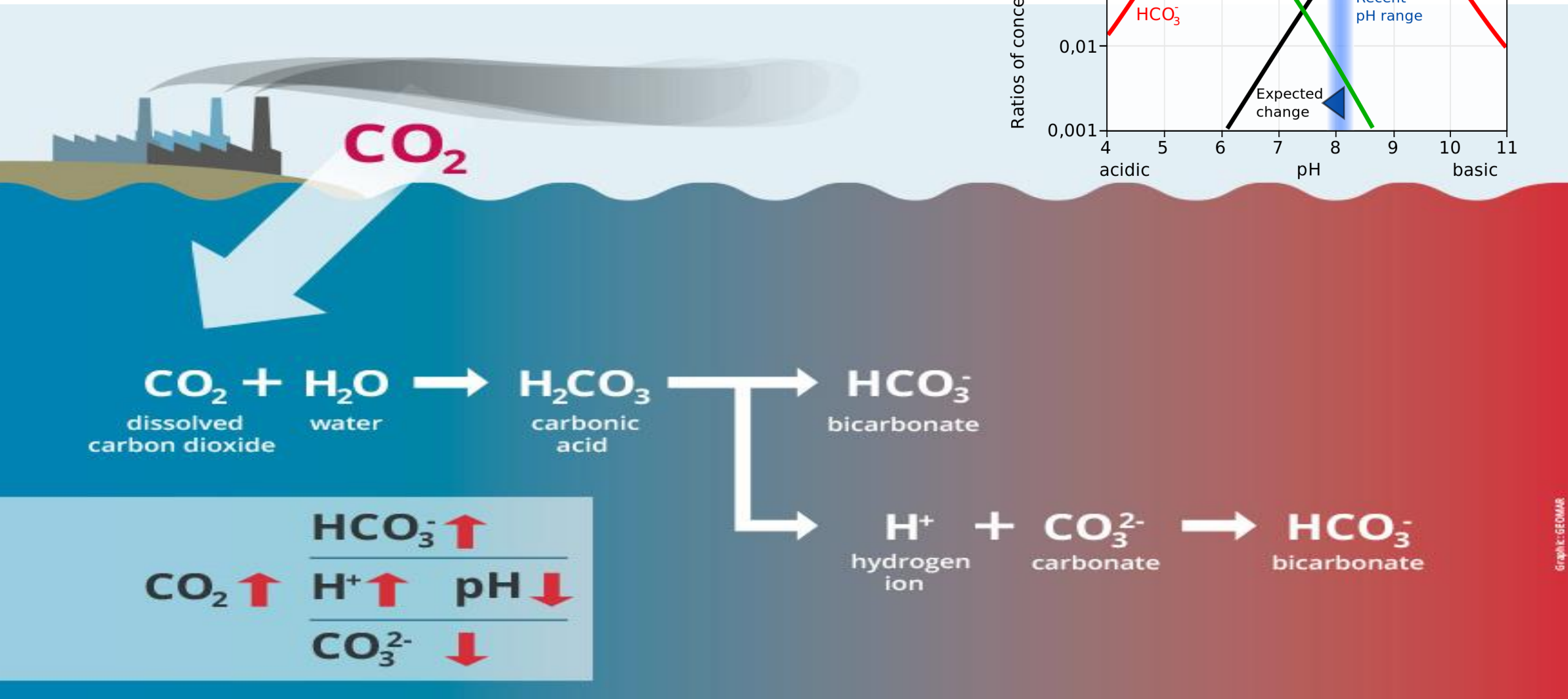
HOMOGENEOUS atmosphere versus...



HETEROGENEOUS ocean!

OCEAN ACIDIFICATION:

The other CO₂ problem, the evil twin of Climate Change...



- Carbon cycle

- Methods for Ocean carbon parameters:

DIC

ALK

pH

pCO₂

“If you know two of these parameters you can compute all the others with a given T^a , salinity and pressure”

- Carbon cycle

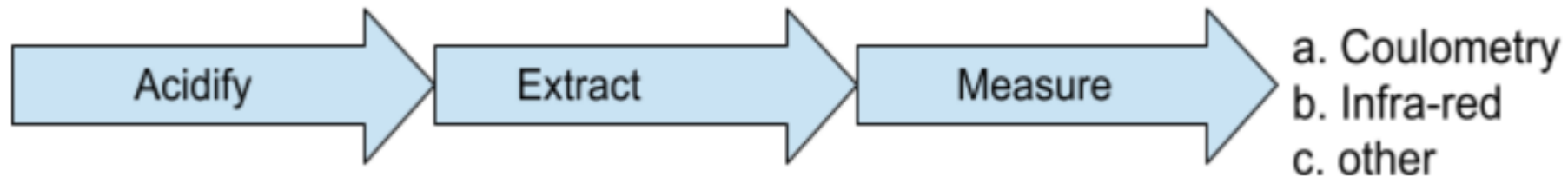
- Methods for Ocean carbon parameters:

DIC

Dissolved inorganic carbon (DIC or CT) is **the sum of the concentrations** of all inorganic carbon species.

$$\text{DIC} = [\text{CO}_2] + [\text{HCO}_3^-] + [\text{CO}_3^{=}]$$

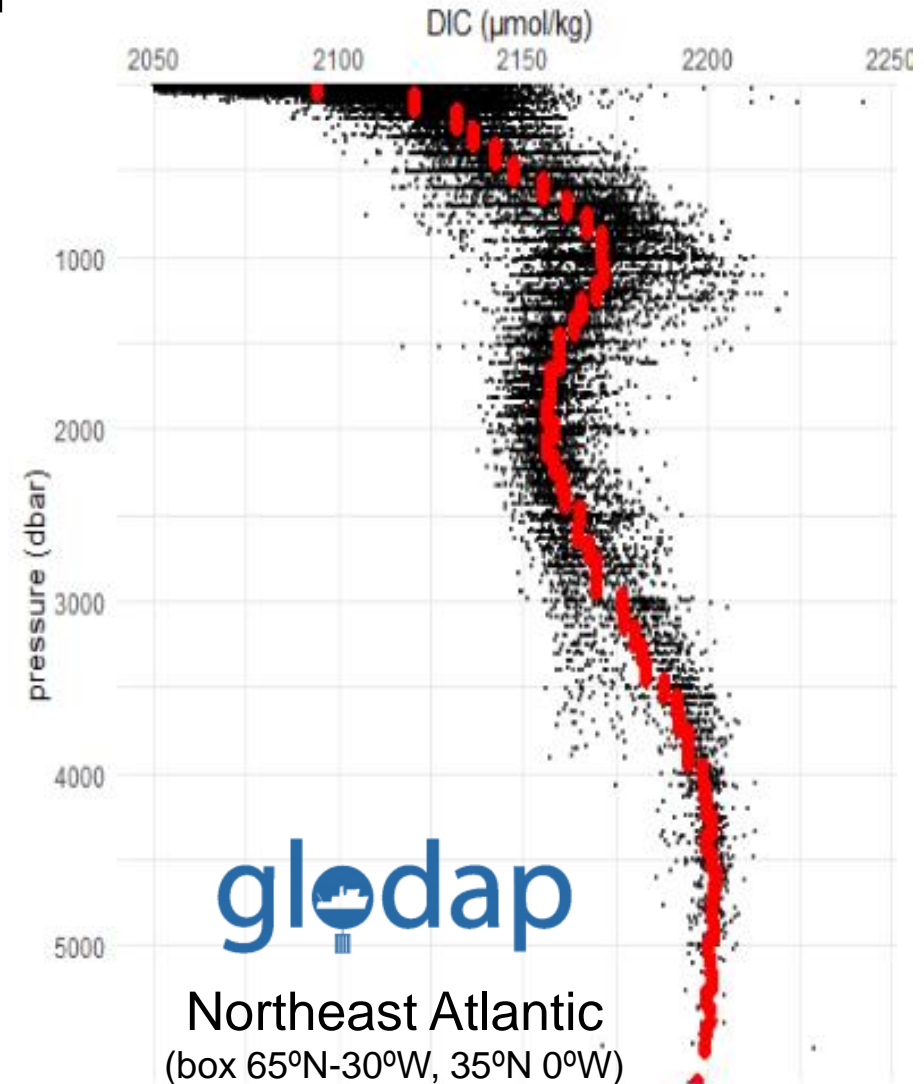
(1:90:9)



- Carbon cycle

- Methods for Ocean carbon parameters:

DIC

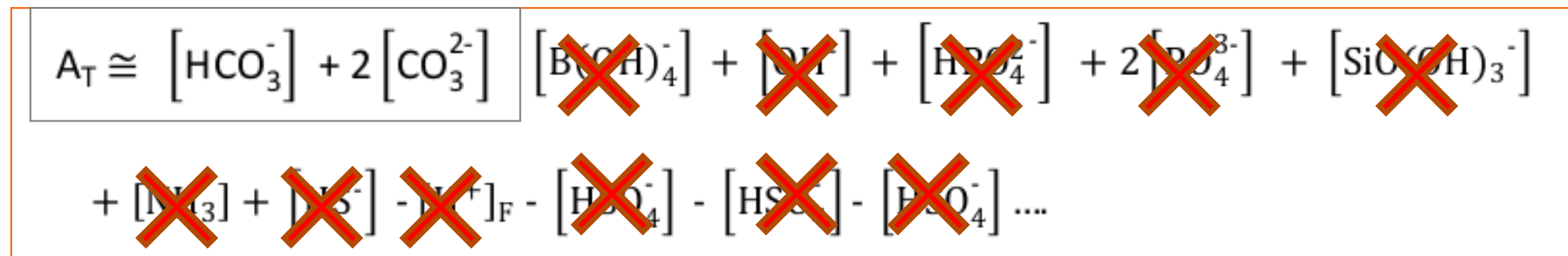


- Carbon cycle

- Methods for Ocean carbon parameters:

ALK

The excess of proton acceptors over proton donors with respect to a certain zero level of protons (Dickson, 1981):

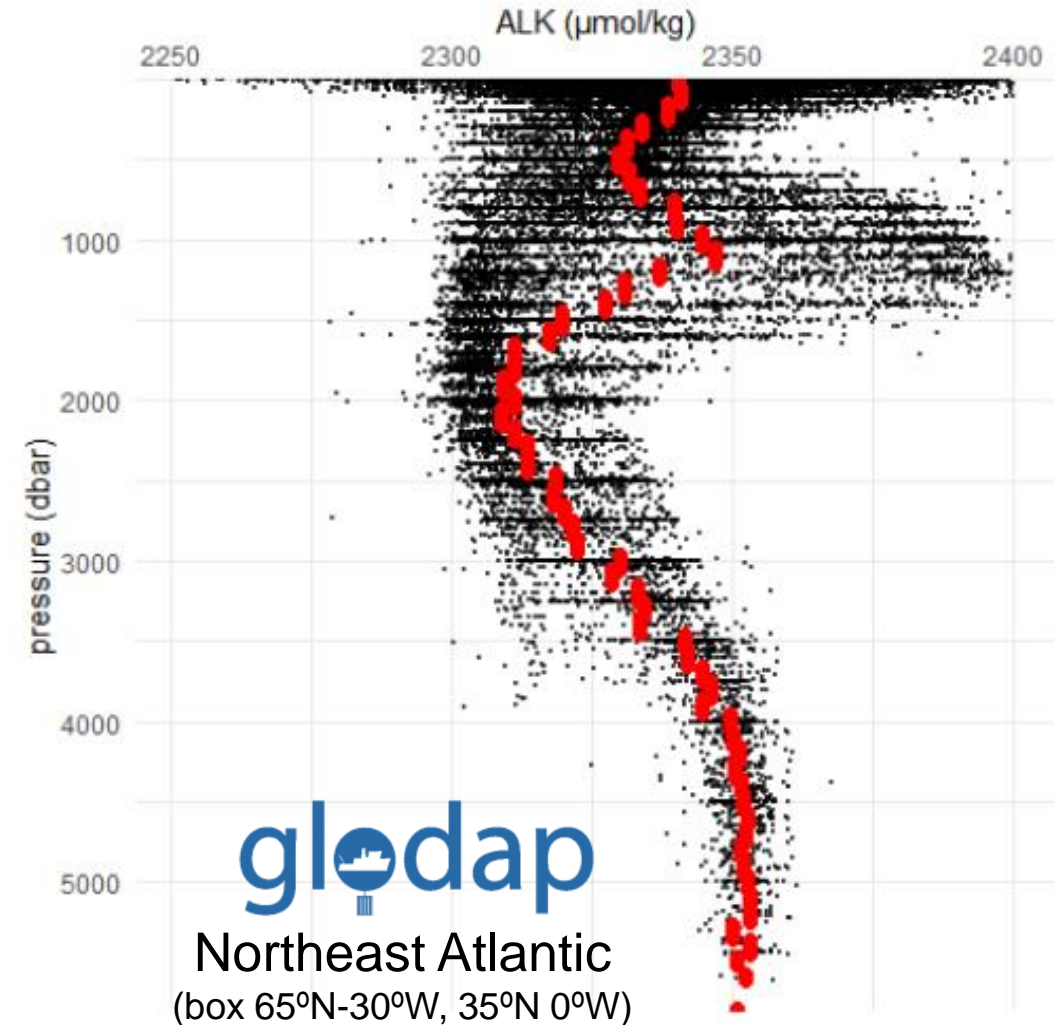


Measurement techniques:

1. Acidimetric titration

- Carbon cycle
 - Methods for Ocean carbon parameters:

ALK



- Carbon cycle

- Methods for Ocean carbon parameters:

pH

defined as the negative of the base 10
logarithm of the hydrogen ion ($[H^+]$)
concentration

$$pH = -\log_{10}[H^+]$$

Measurement techniques:

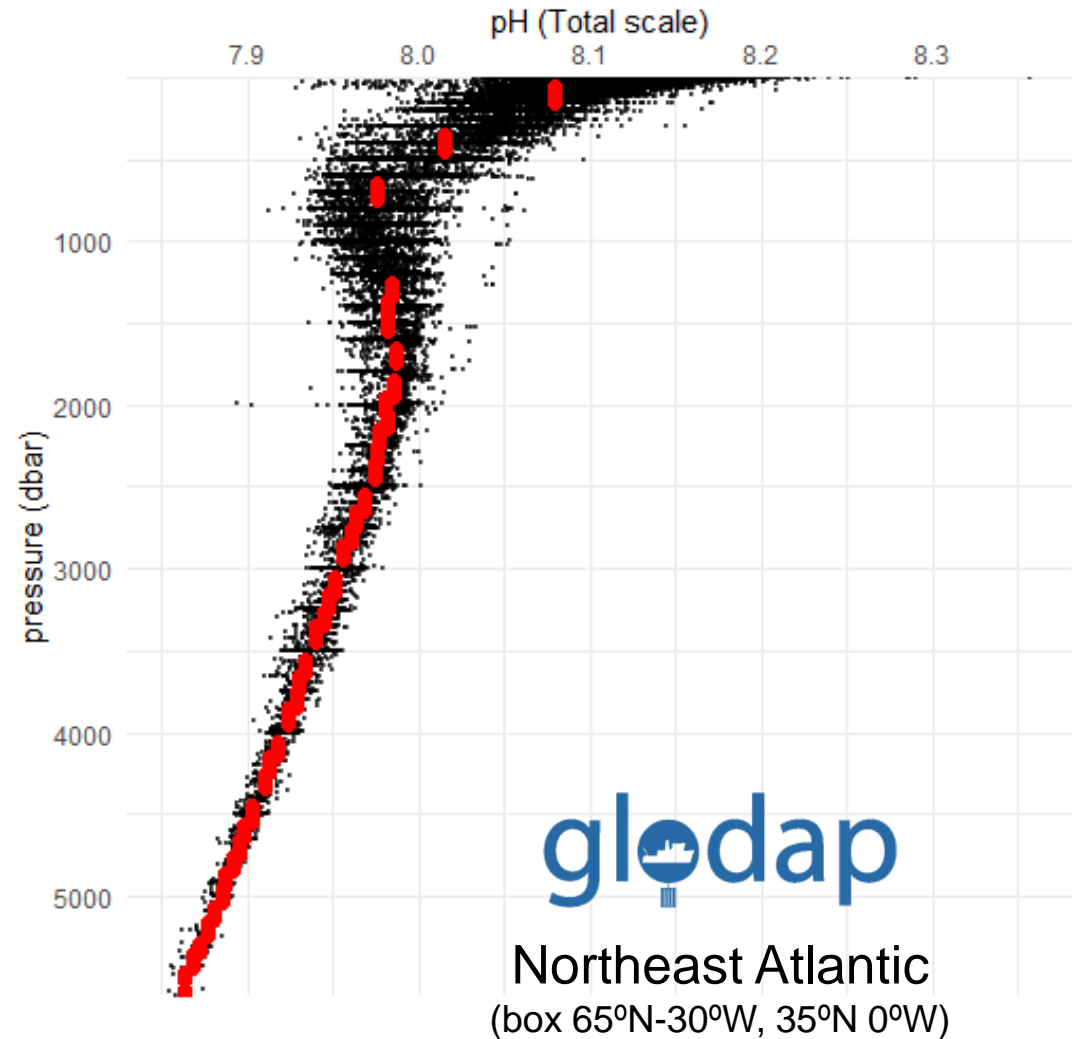
1. Potentiometric technique using a glass/reference electrode cell
2. Spectrophotometric with indicator dye *m*-cresol pruple



- Carbon cycle

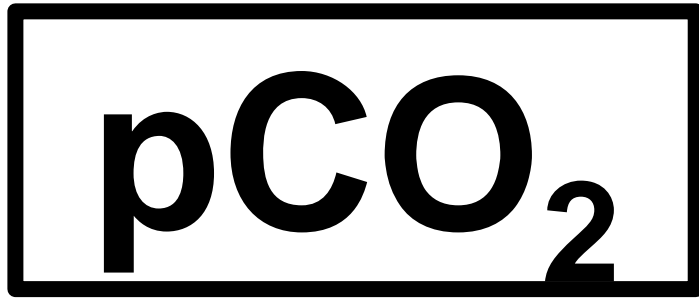
- Methods for Ocean carbon parameters:

pH



- Carbon cycle

- Methods for Ocean carbon parameters:



The product of the mole fraction of CO_2 ($x\text{CO}_2$) in the equilibrated gas phase and the total pressure of equilibration (p):

$$p(\text{CO}_2) = x(\text{CO}_2) \cdot p$$

A temperature-dependent property

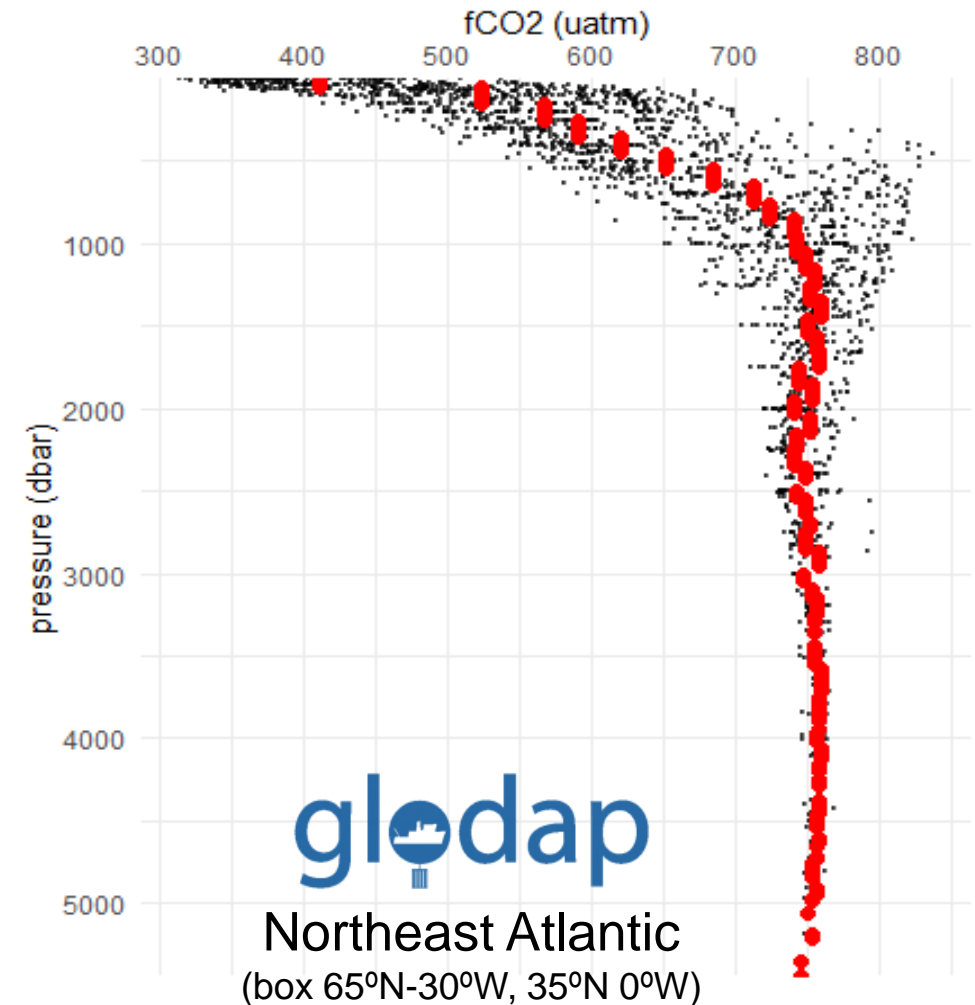
Measurement techniques:

1. from discrete samples in air equilibrium
 2. From continuous stream in air equilibrium
- ➡ Non-dispersive infrared absorption detector

- Carbon cycle

- Methods for Ocean carbon parameters:

pCO₂



- Carbon cycle
 - Methods for Ocean carbon parameters
 - ❖ Study cases: *seacarb*

Table 1. Carbonate system software packages.

Package	Language	Version	Reference
CO2SYS ^a	QBasic	1.05	Lewis and Wallace (1998)
CO2SYS ^b	Excel	24	Pelletier et al. (2007)
CO2SYS ^a	Excel	2.1	Pierrot et al. (2006)
CO2SYS ^a	MATLAB	1.1	van Heuven et al. (2011)
CO2calc ^c	Visual Basic	1.3.0	Robbins et al. (2010)
csys ^d	MATLAB	04–2014	Zeebe and Wolf-Gladrow (2001)
ODV ^e	C++	4.5.0	Schlitzer (2002)
mocsy ^f	Fortran 95	2.0	Orr and Epitalon (2015)
seacarb ^g	R	3.0.6	Gattuso et al. (2015)
swco2 ^h	Excel	2	Hunter (2007); Mosley et al. (2010)
swco2 ^h	Visual Basic	2	Hunter (2007)

^a <http://cdiac.ornl.gov/oceans/co2rprt.html>

^b <http://www.ecy.wa.gov/programs/eap/models.html>

^c <http://pubs.usgs.gov/of/2010/1280/>

^d <http://www.soest.hawaii.edu>

^e <http://odv.awi.de/>

^f <http://ocmip5.ipsl.jussieu.fr/mocsy>

^g <http://cran.r-project.org/package=seacarb>

^h http://neon.otago.ac.nz/research/mfc/people/keith_hunter/software/swco2/

PyCO2SYS: marine carbonate system calculations in Python



Humphreys, Matthew P.; Gregor, Luke; Pierrot, Denis; van Heuven, Steven M. A. C.; Lewis, Ernie R.; Wallace, Douglas W. R.

PyCO2SYS is a Python toolbox for solving the marine carbonate system and calculating related seawater properties. Its core is a Python implementation of CO2SYS for MATLAB. Documentation is available at PyCO2SYS.rtfd.io.

<https://github.com/mvdm7/PyCO2SYS>

<https://github.com/mvdm7/PyCO2SYS-examples>

Orr, J. C., Epitalon, J. M. & Gattuso, J. P. Comparison of ten packages that compute ocean carbonate chemistry. *Biogeosciences* **12**, (2015). ([link](#))

seacarb

Gattuso, J.-P., Epitalon, J.-M., Lavigne, H., Orr, J.,
2020. seacarb: Seawater Carbonate Chemistry.



```
> library(seacarb)
```

KEY FUNCTION: *carb(...)*

```
> carb(flag, var1, var2, S=35, T=25, Patm=1, P=0,  
Pt=0, Sit=0, k1k2="x", kf="x", ks="d", pHscale="T", b="u74",  
gas="potential", warn="y", eos="eos80", long, lat)
```

> carb(**flag**, var1, var2,
S=Salinity, T=Temperature,
Patm=1, P=Pressure **in bar**,
Pt=Total phosphate mol/kg, Sit=Total Silicate mol/kg,
k1k2="x", kf="x", ks="d", pHscale="T", b="u74", gas="potential",
warn="y", eos="eos80", long, lat)



flag = 1 pH and CO₂ given
flag = 2 CO₂ and HCO₃ given
flag = 3 CO₂ and CO₃ given
flag = 4 CO₂ and ALK given
flag = 5 CO₂ and DIC given
flag = 6 pH and HCO₃ given
flag = 7 pH and CO₃ given
flag = 8 pH and ALK given

flag = 9 pH and DIC given
flag = 10 HCO₃ and CO₃ given
flag = 11 HCO₃ and ALK given
flag = 12 HCO₃ and DIC given
flag = 13 CO₃ and ALK given
flag = 14 CO₃ and DIC given
flag = 15 ALK and DIC given

flag = 21 pCO₂ and pH given
flag = 22 pCO₂ and HCO₃ given
flag = 23 pCO₂ and CO₃ given
flag = 24 pCO₂ and ALK given
flag = 25 pCO₂ and DIC given

> carb(flag, var1, var2,
S=Salinity, T=Temperature,
Patm=1, P=Pressure **in bar**,
Pt=Total phosphate mol/kg, Sit=Total Silicate mol/kg,
k1k2="x", kf="x", ks="d", pHscale="T", b="u74", gas="potential",
warn="y", eos="eos80", long=1.e20, lat=1.e20)



flag = 8	pH	and	ALK	given
flag = 9	pH	and	DIC	given
flag = 15	ALK	and	DIC	given

var1

var2

NOTE: the order is important!

> carb(flag, var1, var2,
S=Salinity, T=Temperature,
Patm=1, P=Pressure in bar,
Pt=Total phosphate mol/kg, Sit=Total Silicate mol/kg,
k1k2="x", kf="x", ks="d", pHscale="T", b="u74", gas="potential",
warn="y", eos="eos80", long=1.e20, lat=1.e20)



flag = 8 pH and ALK given
flag = 9 pH and DIC given
flag = 15 ALK and DIC given

var1
var2

Units: mol/kg always expect pH (no units) and pCO₂ (μatm)

NOTE UNITS! is important!

Summary (for a intro level)



These are your carbon data

```
> carb(flag, var1, var2,  
      S=Salinity, T=Temperature,  
      Patm=1, P=Pressure in bar,  
      Pt=Total phosphate mol/kg, Sit=Total Silicate mol/kg,  
      k1k2="x", kf="x", ks="d", pHscale="T", b="u74", gas="potential",  
      warn="y", eos="eos80", long, lat)
```

Almost sure you also
have this data

You can let Nutrient
info in blank

...and forget about this

(unless you like to go into detail, of course!)

These are your carbon data

> carb(**flag, var1, var2,**
S=Salinity, T=Temperature, Patm=1, **P=Pressure in bar**)

*HTML “CANAIMOC.html” done with the Rmarkdown
file “CANAIMOC.Rmd” in:

<https://github.com/mfontela/CANAIMOC>

Also available online here:

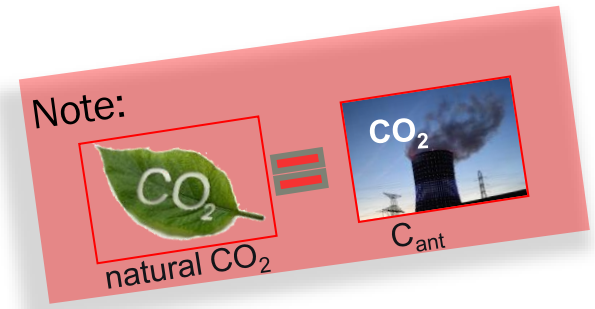
<https://mfontela.github.io/web/CANAIMOC.html>

- Carbon cycle
 - Intro
 - Methods for Ocean carbon parameters
 - ❖ Study cases: seacarb, CO2SYS
- Anthropogenic carbon (C_{ant})
 - Intro
 - Methods: carbon-based vs transient tracers-based
 - Biogeochemical back-calculation ϕC_T^0 method:
 - ❖ Study cases

- Anthropogenic carbon

- Intro

$$\text{DIC} = \text{DIC}_{\text{nat}} + \text{C}_{\text{ant}}$$



- C_{ant} is **not a directly measurable quantity**. It has to be estimated using indirect means.
- The anthropogenic signal in the ocean is only **a few percent (3-4%)** of DIC
- Carbon in the ocean participates in rather complex **in situ biogeochemistry processes**.
- C_{ant} distribution in the ocean is highly **heterogeneous**.

Anthropogenic carbon (C_{ant})

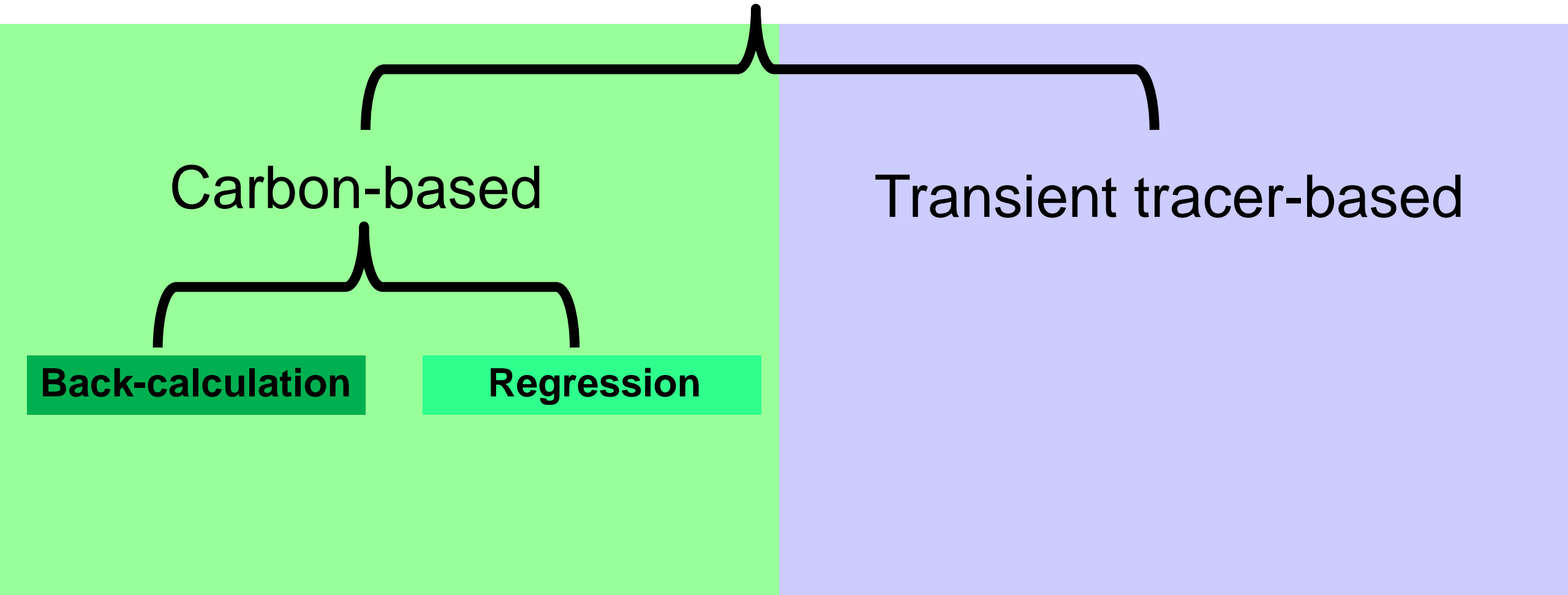
Methods

Carbon-based

Transient tracer-based

Back-calculation

Regression

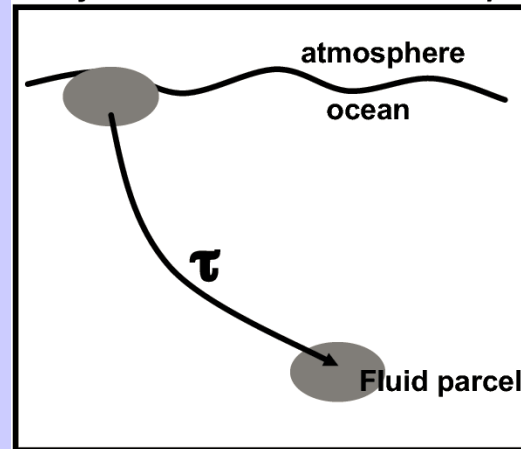


Transient tracer-based

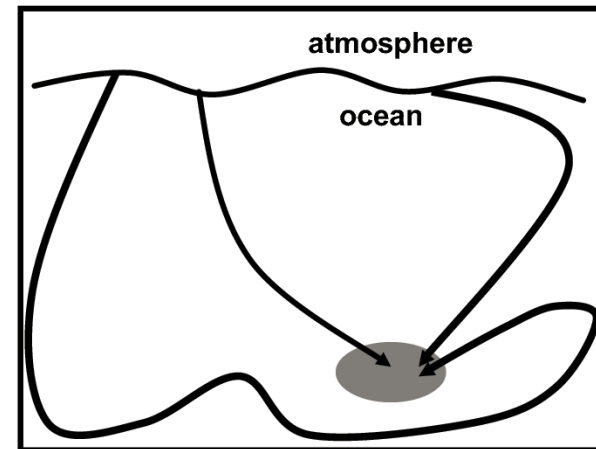
- TTD: Transient tracer distribution (Waugh et al., 2006)
- GF: Green function (Khatiwala et al., 2009)
- OCIM: Ocean Circulation Inverse Model (DeVries, 2014)

(a bit in-between methods: also an inversion method but with $\Delta^{14}\text{C}$ as constraint)

Purely advective view of ocean transport



and the more realistic view ...



CAVEAT: mean air-sea equilibrium times of transient tracers different to CO_2

Methods

```
graph TD; Methods --> Carbon-based; Carbon-based --> Back-calculation; Carbon-based --> Regression;
```

Carbon-based

Back-calculation

Regression

- **multiple linear regressions** on a number of biogeochemical variables
- Assumption: **no temporal trends in the independent variables** and the relationship between dependent and independent variables stays the same.
- Weakness: you need **at least two observations** at approximately the same location for different periods.

Methods

```
graph TD; Methods --- Carbon-based; Carbon-based --- Back-calculation; Carbon-based --- Regression;
```

Carbon-based

Back-calculation

Regression

- MLR (Wallace, 1995)
- eMLR (Friis et al., 2005)
- eMLR(C*) (Clement & Gruber, 2018)

Regression

➤ eMLR(C^*) (Clement & Gruber, 2018)

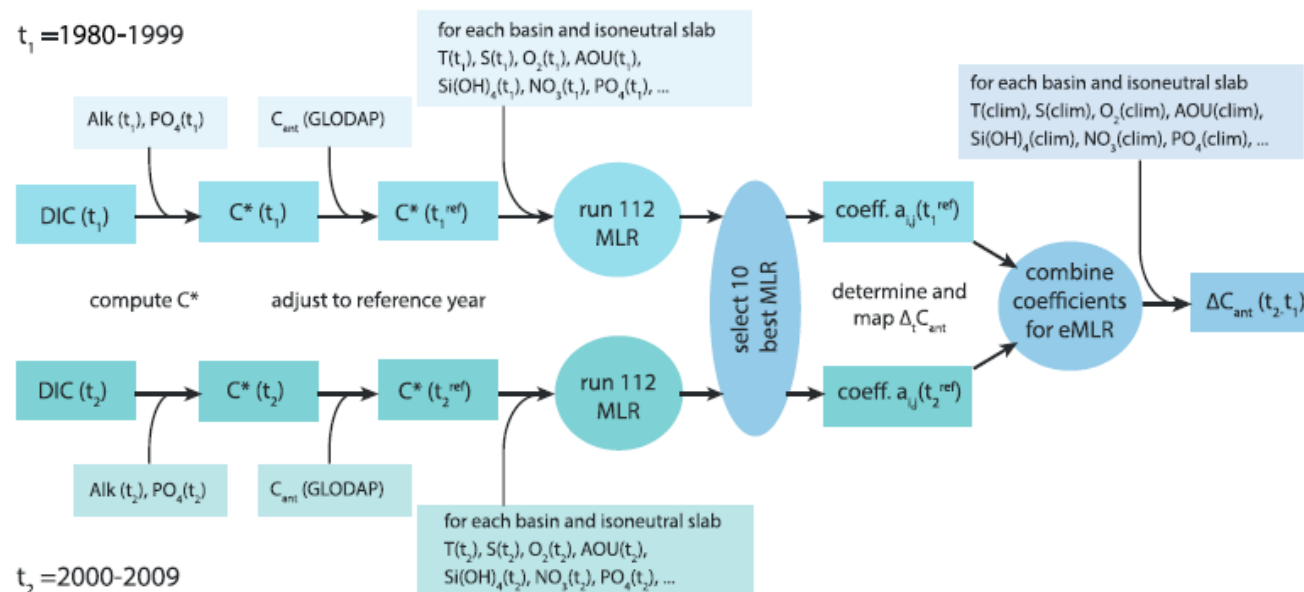
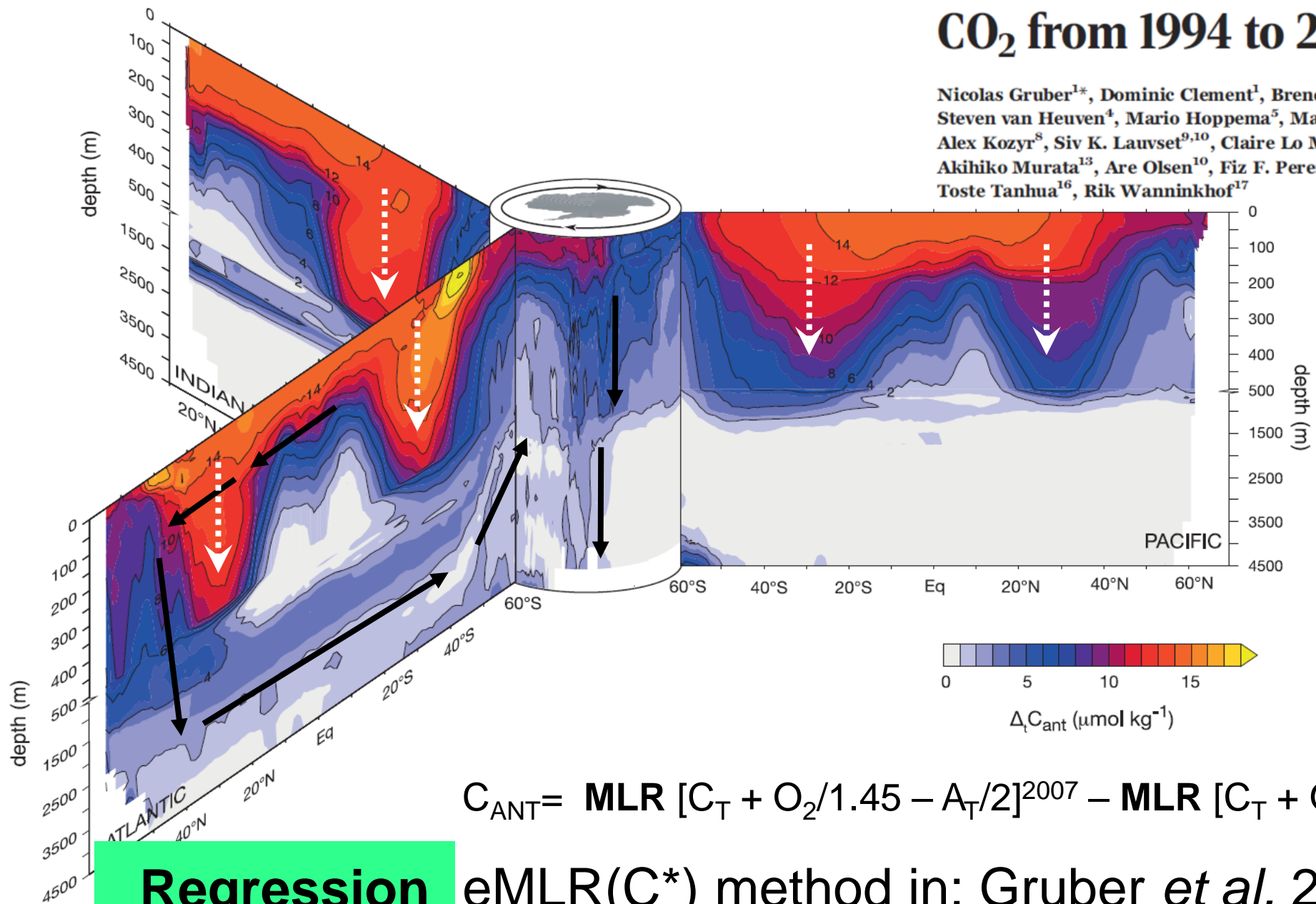


Figure 2. Flowchart of the eMLR(C^*) method to determine the decadal increase in the oceanic content of anthropogenic CO_2 , that is, $\Delta C_{\text{ant}}(t_2 - t_1)$. Also shown are the required input data in each of the four major steps. In the first step, C^* is computed from the measured DIC, Alk, and PO_4 . In the second step, C^* is adjusted to a common reference year for each of the two periods, that is, t_i^{ref} . In the third step, separate multiple linear regressions are run for each time period, using all possible combinations. The best 10 regression models are selected for each isoneutral surface based on their fit to the data and combined to form the eMLR. In the fourth and final step, the eMLR models are combined with global gridded climatological distributions of the predictors to map $\Delta C_{\text{ant}}(t_2 - t_1)$ globally below 150 m.

The oceanic sink for anthropogenic CO₂ from 1994 to 2007

Nicolas Gruber^{1*}, Dominic Clement¹, Brendan R. Carter^{2,3}, Richard A. Feely², Steven van Heuven⁴, Mario Hoppema⁵, Masao Ishii⁶, Robert M. Key⁷, Alex Kozyr⁸, Siv K. Lauvset^{9,10}, Claire Lo Monaco¹¹, Jeremy T. Mathis¹², Akihiko Murata¹³, Are Olsen¹⁰, Fiz F. Perez¹⁴, Christopher L. Sabine¹⁵, Toste Tanhua¹⁶, Rik Wanninkhof¹⁷



$$C_{\text{ANT}} = \text{MLR} [C_T + O_2/1.45 - A_T/2]^{2007} - \text{MLR} [C_T + O_2/1.45 - A_T/2]^{1994}$$

eMLR(C*) method in: Gruber *et al.* 2019 Science

Methods

Carbon-based

Back-calculation

Regression

$$C_{\text{ant}} = \text{DIC} - \Delta C_{\text{bio}} - C_{\text{phys}}$$

- ✓ Modify a measured DIC concentration for the changes that occurred since being no longer in contact with the ocean surface
- ✓ Estimation of the preformed DIC is critical point.

Methods

```
graph TD; Methods --> Carbon-based; Methods --> Regression; Carbon-based --> Back-calculation; Carbon-based --> Regression;
```

Carbon-based

Back-calculation

- Brewer (1978), Chen and Millero (1979)
- ΔC^* (Gruber et al. 1996)
- TrOCA (Touratier & Goyet, 2004).
- ϕCT^o (Vázquez-Rodriguez et al., 2009)

Regression

➤ Brewer (1978)

Back-calculation

VOL. 5, NO. 12

GEOPHYSICAL RESEARCH LETTERS

DECEMBER 1978

DIRECT OBSERVATION OF THE OCEANIC CO₂ INCREASE

Peter G. Brewer

Woods Hole Oceanographic Institution
Woods Hole, Massachusetts 02543

➤ Chen and Millero (1979)

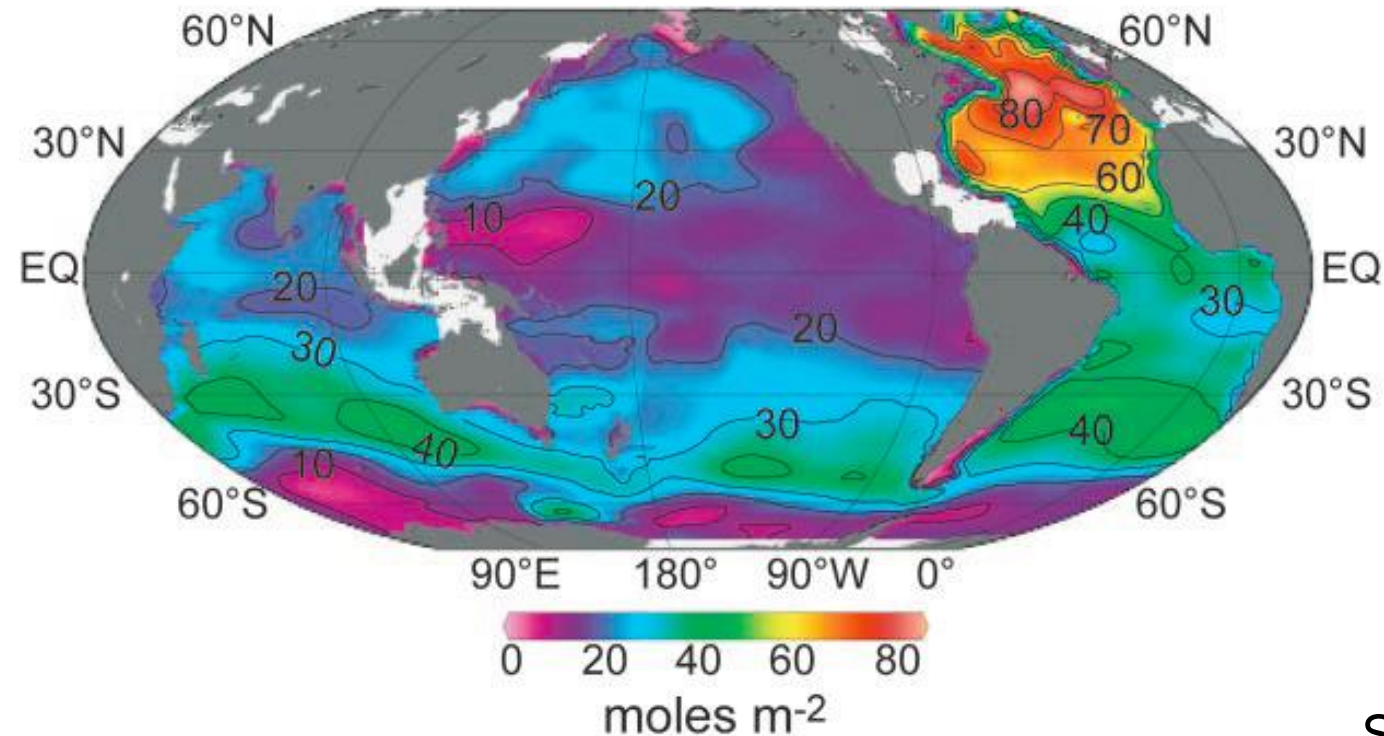
Nature Vol. 277 18 January 1979

Gradual increase of oceanic CO₂

➤ ΔC^* (Gruber et al. 1996)

Back-calculation

$$C^* = DIC - r_{C:O_2}O_2 - \frac{1}{2} (Alk + r_{N:O_2}O_2)$$



Sabine et al., 2004

➤ TrOCA (Touratier & Goyet, 2004)

Back-calculation

Tracer combining **O**xygen, inorganic **C**arbon and total **A**lkalinity

The concentration of $C_{\text{ant}}^{\text{TrOCA}}$ is then estimated using eq. 11:

$$C_{\text{ant}}^{\text{TrOCA}} = \frac{O_2 + 1.279 \left[C_T - \frac{1}{2} A_T \right] - e^{\left(7.511 - (1.087 \times 10^{-2}) \theta - \frac{7.81 \times 10^5}{A_T^2} \right)}}{1.279}.$$

(11)

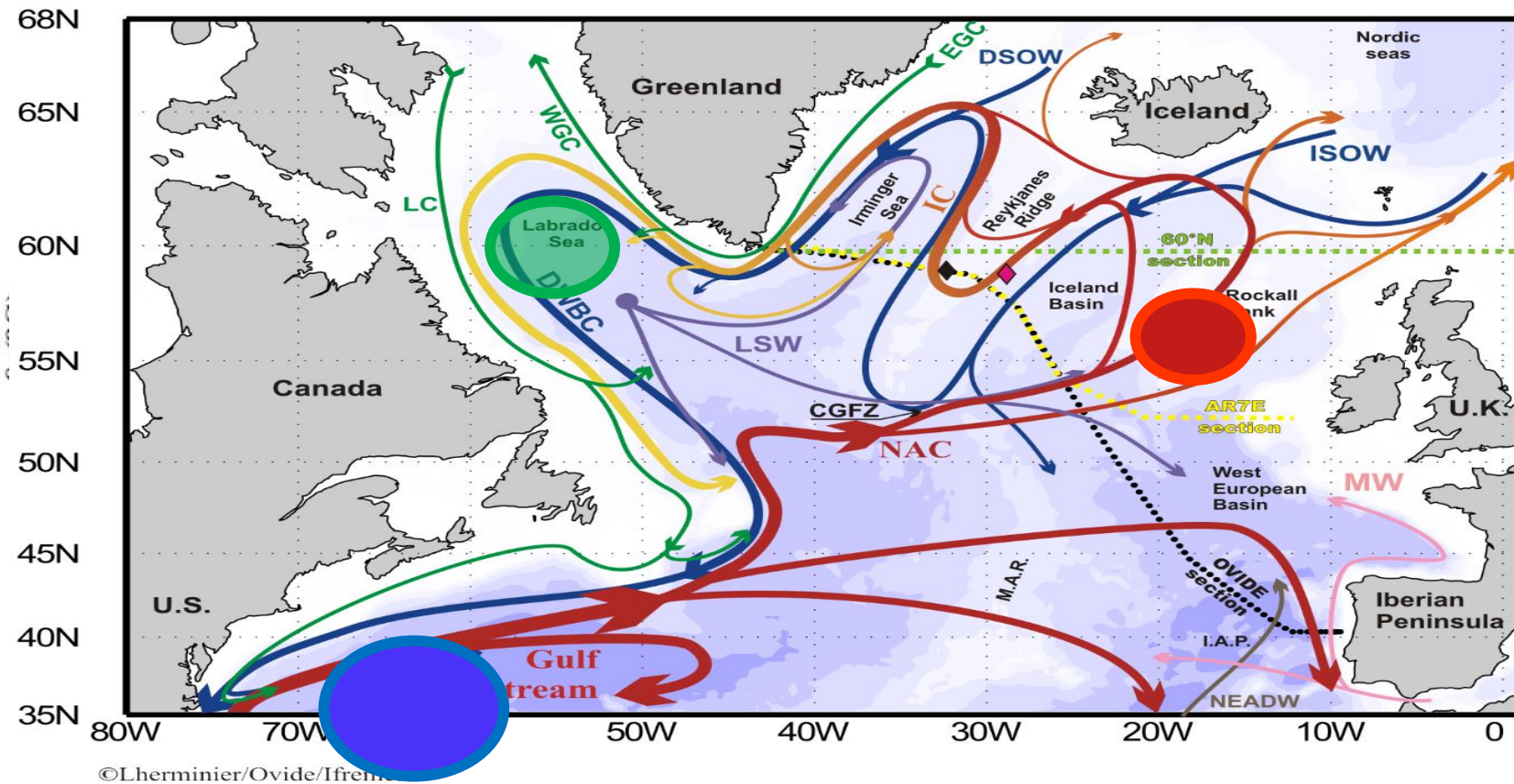
➤ ϕCT° (Vázquez-Rodriguez et al., 2009)

Back-calculation

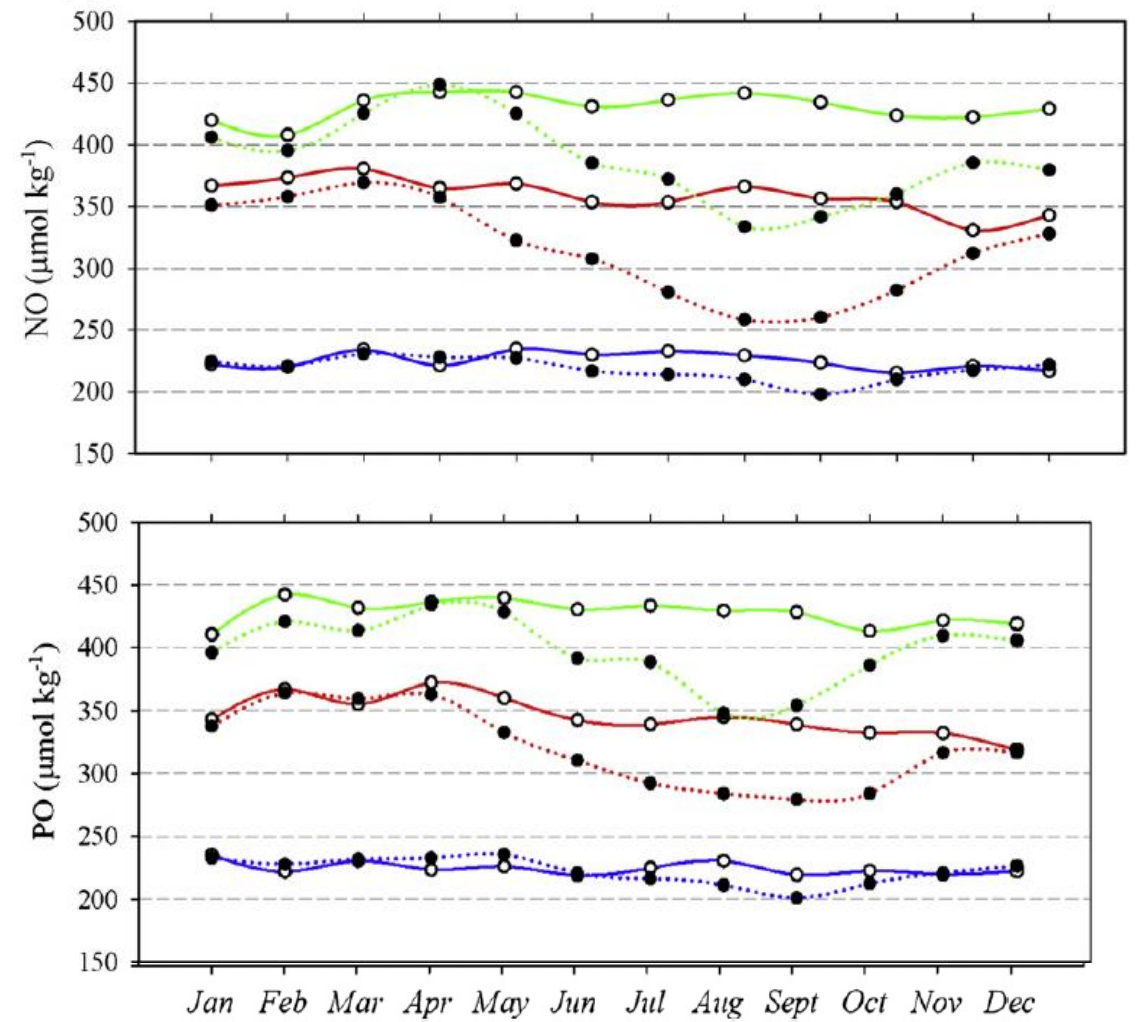
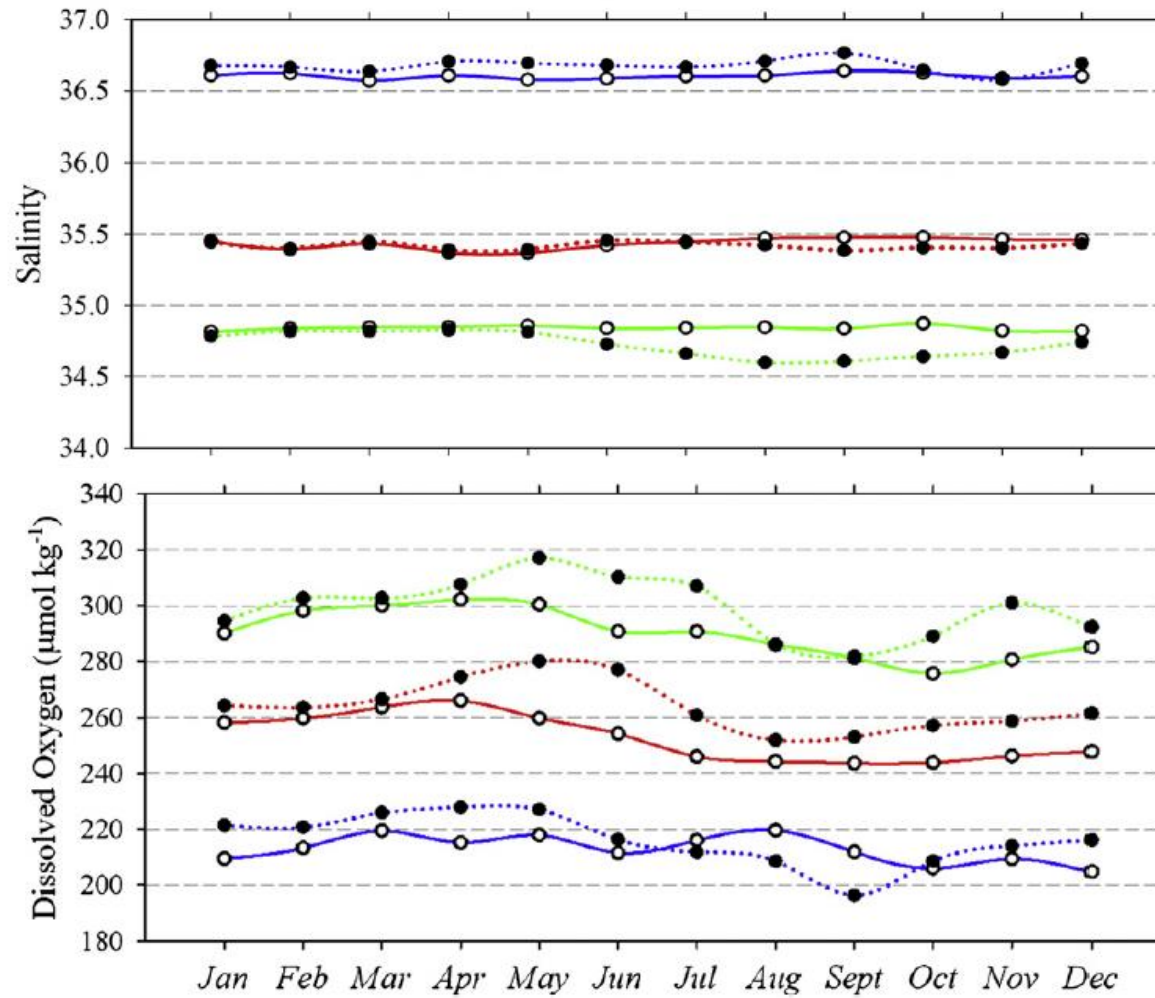
- ✓ C_{ant} is the difference between the water mass DIC at their formation time and the DIC that it would have in preindustrial times.
- ✓ Takes into account changes in carbon concentration due **organic matter remineralization and CaCO_3 dissolution**.
- ✓ Only need **hydrographical** data: carbon, oxygen and inorganic nutrients.
- ✓ Overall uncertainty: **$\pm 5.2 \mu\text{mol}\cdot\text{kg}^{-1}$**
- ✓ Reference layer: **subsurface** layer.

Subsurface layer

— Sargasso Sea — Rockall Plateau — Labrador Sea
— Subsurface layer (100-200 m) Surface layer (0-20 m)



Subsurface layer



- Proper characterization of preformed conditions.
- No needs for CFC's data.
- No assumptions about C_{ant} saturation in the water mass formation time.
- Includes corrections for warming and acidification.
- Does not need a “zero” C_{ant} reference
- Is based on biogeochemical and oceanographic knowledge.

- Not global, only for the Atlantic Ocean.
- Depends in part on a simplified OMP.
- ΔC_{dis} computation is based on multiple linear regressions (MLR) that are potentially improved.
- Assume constant and homogeneous stoichiometric ratios.

More info in:

- Vázquez-Rodríguez, M., F. Touratier, C. Lo Monaco, D. W. Waugh, X. a. Padin, R. G. J. Bellerby, C. Goyet, N. Metzl, a. F. Ríos, and F. F. Pérez (2008), Anthropogenic carbon distributions in the Atlantic Ocean: data-based estimates from the Arctic to the Antarctic, *Biogeosciences Discuss.*, 5(2), 1421–1443, doi:10.5194/bgd-5-1421-2008.
- Pérez, F. F., M. Vázquez-Rodríguez, H. Mercier, A. Velo, P. Lherminier, and A. F. Ríos (2010), Trends of anthropogenic CO₂ storage in North Atlantic water masses, *Biogeosciences*, 7(5), 1789–1807, doi:10.5194/bg-7-1789-2010.
- Pérez, F. F., H. Mercier, M. Vázquez-Rodríguez, P. Lherminier, A. Velo, P. C. Pardo, G. Rosón, and A. F. Ríos (2013), Atlantic Ocean CO₂ uptake reduced by weakening of the meridional overturning circulation, *Nat. Geosci.*, 6(2), 146–152, doi:10.1038/ngeo1680.
- Vázquez-Rodríguez, M., X. a. Padin, a. F. Ríos, R. G. J. Bellerby, and F. F. Pérez (2009), An upgraded carbon-based method to estimate the anthropogenic fraction of dissolved CO₂ in the Atlantic Ocean, *Biogeosciences Discuss.*, 6(2), 4527–4571, doi:10.5194/bgd-6-4527-2009.
- Fajar, N. M., P. C. Pardo, L. Carracedo, M. Vázquez-Rodríguez, A. F. Ríos, and F. F. Pérez (2012), Trends of anthropogenic CO₂ along 20°W in the Iberian Basin | Tendencias del CO₂ antropogénico a lo largo de 20°W en la Cuenca Ibérica, *Ciencias Mar.*, 38(1 B), 287–306.
- Khatiwala, S. et al. Global ocean storage of anthropogenic carbon. *Biogeosciences* **10**, 2169–2191 (2013).
- Gruber, N. et al. The oceanic sink for anthropogenic CO₂ from 1994 to 2007. *Science* (80-.). **363**, 1193–1199 (2019).
- Orr, J. C. et al. Anthropogenic ocean acidification over the twenty-first century and its impact on calcifying organisms. *Nature* **437**, 681–686 (2005).
- Clement, D. & Gruber, N. The eMLR(C*) Method to Determine Decadal Changes in the Global Ocean Storage of Anthropogenic CO₂. *Global Biogeochem. Cycles* (2018). doi:10.1002/2017GB005819

& a thousand more!!!

& the PhD thesis of Marcos Vázquez-Rodríguez...



THANK YOU!
MERCI!
OBRIGADO! 😊



CANAIMOC Workshop



CArbon NOrth Atlantic Irrigation by the Meridional Overturning Circulation

For this work M. Fontela was funded by Portuguese national funds from FCT - Foundation for Science and Technology through project UID/Multi/04326/2019 and CEECINST/00114/2018.

mmfontela@ualg.pt

@MarcosFontela

www.marcosfontela.com