

## Spherical shell coordinates

```
clear
set_demo_defaults;
R_mars = 3389508;      % [m] Mars' mean radius
grav = 3.711;         % [m/s^2] grav. acceleration on Mars
```

We have seen that moving to cylindrical coordinates removed the ambiguity in the interpretation of the polar recharge. To properly incorporate precipitation we need to go to a geometry with a meaningful surface area. In 1D linear coordinates the surface area is an arbitrary function of the undetermined width. In cylindrical coordinates we have a proper surface area, but given that the southern highlands aquifer stretches halfway through the northern hemisphere the we have a huge error in the actual surface area compared to a sphere, see figure. The same would be true for any estimate of the actual groundwater volume.

```
theta_bnd = acos(1/3);
theta_bnd_deg = rad2deg(theta_bnd)
```

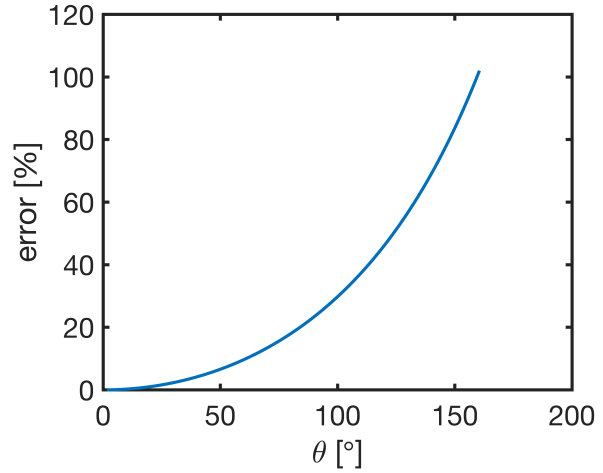
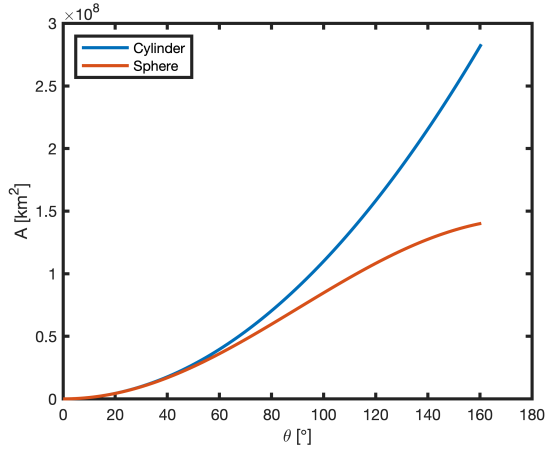
```
theta_bnd_deg = 70.5288
```

```
theta_vec = linspace(0,90+theta_bnd_deg,100);
l = R_mars*deg2rad(theta_vec); % [m] distance to dichotomy bnd

A_cyl = pi*l.^2;
A_cap = 2*pi*R_mars^2*(1-cos(deg2rad(theta_vec)));

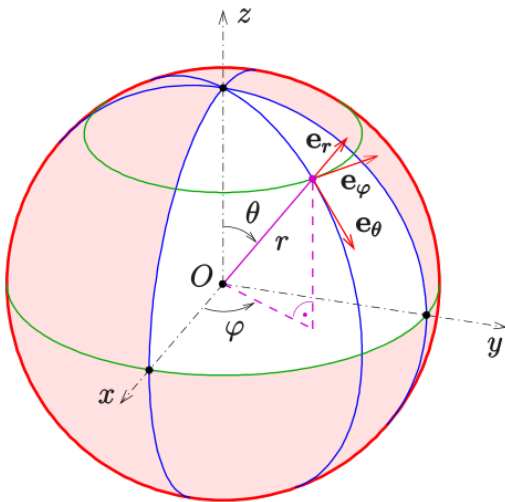
figure('position',[10 10 900 600])
subplot 121
plot(theta_vec,A_cyl/1e6,theta_vec,A_cap/1e6)
xlabel '\theta [\circ]', ylabel 'A [km^2]', pbaspect([1 .8 1])
legend('Cylinder','Sphere','location','northwest')

subplot 122
plot(theta_vec,(A_cyl-A_cap)./A_cap*100)
xlabel '\theta [\circ]', ylabel 'error [%]', pbaspect([1 .8 1])
```



## Spherical coordinates

This motivates us to discretize the discrete operators in spherical coordinates. The definition of standard variables in spherical coordinates is shown in the figure below.



Here  $r$  is the radial coordinate,  $\theta$  is the co-latitude and  $\varphi$  is the circumferential coordinate. The associated definition of the gradient and the divergence

- $\nabla h = \frac{\partial h}{\partial r} \hat{\mathbf{r}} + \frac{1}{r} \frac{\partial h}{\partial \theta} \hat{\boldsymbol{\theta}} + \frac{1}{r \sin \theta} \frac{\partial h}{\partial \varphi} \hat{\boldsymbol{\phi}}$
- $\nabla \cdot \mathbf{q} = \frac{1}{r^2} \frac{\partial(r^2 q_r)}{\partial r} + \frac{1}{r \sin(\theta)} \frac{\partial}{\partial \theta} (\sin(\theta) q_\theta) + \frac{1}{r \sin(\theta)} \frac{\partial q_\varphi}{\partial \varphi}$

where  $\mathbf{q} = [q_r, q_\theta, q_\varphi]$ .

The southern highlands aquifer is in a spherical shell, so that  $r = R_{\text{Mars}}$  is fixed. To obtain a one-dimensional model we assume no change in the circumferential direction,  $\varphi$ , so that  $\partial/\partial\varphi = 0$ . Therefore, the remaining independent variable is the co-latitude,  $\theta$ , and one-dimensional operators in spherical shell geometry,  $x = \theta$ , are

- $\nabla h = \frac{1}{R_{\text{Mars}}} \frac{dh}{dx}$
- $\nabla \cdot \mathbf{q} = \frac{1}{R_{\text{Mars}} \sin(x)} \frac{d}{dx} (\sin(x) q)$

In spherical shell coordinates both the divergence and the gradient change. Again we have to amend the function `build_ops.m`.

## Discrete operators

The discrete divergence and gradient matrix in spherical shell geometry can therefore be obtained as follows:

```
Grid.xmin = 0.1; Grid.xmax = 1; Grid.Nx = 35;
Grid = build_grid(Grid);
[D,G,I] = build_ops(Grid);

% Modification for spherical shell
Grid.R_shell = R_mars;
Rf = spdiags(sin(Grid.xf),0,Grid.Nx+1,Grid.Nx+1);
Rcinv = spdiags(1./(Grid.R_shell*sin(Grid.xc)),0,Grid.Nx,Grid.Nx);
D = Rcinv*D*Rf;
G = G/Grid.R_shell;
L = -D*G;
```

Similar to the cylindrical coordinates we evaluate terms outside the divergence at cell centers and terms inside the divergence at cell faces.

## Spherical shell aquifer with precipitation

### Dimensional

The equations for the steady confined aquifer with precipitation on a spherical shell are given by

$$\frac{1}{R_{\text{Mars}} \sin(\theta)} \frac{d}{d\theta} \left( \sin(\theta) b K \frac{1}{R_{\text{Mars}}} \frac{dh}{dx} \right) = q_p \text{ on } \theta \in [0, \theta_b]$$

with the boundary conditions

$$\left. \frac{dh}{d\theta} \right|_0 = 0 \text{ on } h(\theta_b) = h_o$$

The parameter values are as before

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```
yr2s = 60^2*24*365.25; % second per year
rho = 1e3; % [kg/m^3] density of water
grav = 3.711; % [m/s^2] grav. acceleration on Mars
k = 1e-11; % [m^2] permeability (Hanna & Phillips 2005)
mu = 1e-3; % [Pa s] water viscosity
ho = -500; % [m] sea level
b = 5e3; % [m] aquifer thickness
theta_bnd = pi-acos(1/3); % [rad] angle dichotomy boundary from south pole
% derived values
K = k*rho*grav/mu; % [m/s] hydraulic conductivity
```

## Dimensionless

The angle,  $\theta$ , in radians is defined as ratio of the arc length,  $s$ , to the radius,  $R$ , of the circle,  $\theta = s/R$ , and hence dimensionless and of order one. We define the characteristic scale  $h_c = q_p R_{\text{Mars}}^2 / (bK)$  and the associated dimensionless head  $h' = (h - h_o) / h_c$ . This induces the scale  $q_c = q_p R_{\text{Mars}} / b$  for the flux. The dimensionless equations are

$$-\frac{1}{\sin \theta} \frac{d}{d\theta} \left( \sin \theta \frac{dh'}{d\theta} \right) = 1 \text{ on } \theta \in [0, \theta_b]$$

with the boundary conditions

$$\left. \frac{dh'}{d\theta} \right|_0 = 0 \text{ and } h'(\theta_b) = 0.$$

The dimensionless flux is simply  $q' = -\frac{dh'}{d\theta}$ . The only dimensionless parameter is therefore the angle of the dichotomy boundary  $\theta_b$ .

## Analytic solutions

The dimensionless analytic solution is given by

$$h' = \log \left( \frac{\cos \theta + 1}{\cos \theta_b + 1} \right) \text{ and } q' = \frac{1 - \cos \theta}{\sin \theta} = \csc \theta - \cot \theta.$$

```
hD_ana = @(theta,theta_bnd) log((cos(theta)+1)/(cos(theta_bnd)+1));
qD_ana = @(theta) csc(theta) - cot(theta);
```

The solution is show in the figure below for increasing co-lattitudes of the dichotomy boundary

```
h_scale_plot = 0.3;
theta_bnd_vec = [theta_bnd:-theta_bnd/5:theta_bnd/5];

figure('position',[10 10 900 600])
for i = 1:length(theta_bnd_vec)
    theta = linspace(0,theta_bnd_vec(i),1e2);
    subplot(2,2,1)
    plot(theta,hD_ana(theta,theta_bnd_vec(i))); hold on
    subplot(2,2,3)
    plot(theta,qD_ana(theta)); hold on
    subplot(2,2,[2 4])
    x_h = (1+hD_ana(theta,theta_bnd_vec(i))*h_scale_plot).*sin(theta);
    z_h = (1+hD_ana(theta,theta_bnd_vec(i))*h_scale_plot).*cos(theta);
    plot(x_h,-z_h,'-','linewidth',1.5), hold on
end
subplot(2,2,1)
ylabel 'h'', xlabel '\theta'
subplot(2,2,3)
ylabel 'q'', xlabel '\theta'
subplot(2,2,[2 4])
theta_sphere = linspace(0,pi,5e2);
x_base = sin(theta_sphere);
z_base = cos(theta_sphere);
plot(x_base,z_base,'k','linewidth',1.5)
axis equal
xlim([0 2]), ylim([-2 2])
```

